



1 **Modeling Seismic Hazard and Landslide Potentials in Northwestern Yunnan,**
2 **China: Exploring Complex Fault Systems with multi-segment rupturing in a**
3 **Block Rotational Tectonic Zone**

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5 Jia Cheng^{1*}, Chong Xu², Xiwei Xu¹, Shimin Zhang², Pengyu Zhu²

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7 1. School of Earth Science and Resources, China University of Geosciences (Beijing),

8 Beijing, 100083, China

9 2. National Institute of Natural Hazards, Ministry of Emergency Management of China,

10 Beijing, 100085, China

11 *Corresponding author:

12 Jia Cheng (jiacheng@cugb.edu.cn, jiacheng@gmail.com)

13 Address: China University of Geosciences, No. 29, Xueyuan Road, Haidian, Beijing, 100083,

14 China

15 Phone: +86-10-13466515670

16 Fax: +86-10-82322264

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18 **Abstract**

19 The Northwestern Yunnan Region (NWYR), located on the southeastern edge of the

20 Tibetan Plateau, is characterized by a combination of low-crustal flow and gravitational

21 collapse, giving rise to a complex network of active faults. This presents significant

22 seismic hazards, particularly due to the potential for multi-segment ruptures and

23 resulting landslides, as demonstrated by the historical 1515 *M*7.8 Yongshen Earthquake.

24 This article presented a novel seismic hazard modeling study for the NWYR,



25 integrating fault slip parameters and assessing multi-segment rupturing risks. Among
26 the four potential multi-segment rupture combination models examined, Model 1,
27 characterized by multi-segment rupture combinations on single faults, particularly
28 fracturing the Zhongdian fault, emerges as the most suitable for the NWYR, supported
29 by the alignment of modeled seismicity rates with fault slip rates. Our analysis
30 demonstrated that peak ground-motion acceleration values, calculated with a 475-year
31 return period from modeled seismicity rates, exhibited a strong correlation with fault
32 distribution, averagely higher than the China Seismic Ground Motion Parameters
33 Zonation Map. Furthermore, we conducted simulations to forecast landslide occurrence
34 probabilities across our peak ground-motion acceleration distribution map. Our
35 findings underscored that the observed combinations of multi-segment ruptures and
36 their associated behaviors were in alignment with the small block rotation triggered by
37 the gravitational collapse of the Tibetan Plateau. This result highlighted the intricate
38 interplay between multi-segment rupturing hazards and regional geological dynamics.

39 **Key Words:**

40 Northwestern Yunnan Region; multi-segment rupture; probability seismic hazard
41 analysis; landslide probabilities

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44 **1. Introduction**

45 The collision of the Eurasia Plateau and the Indian plate makes the Tibetan Plateau world
46 highest altitude of 4000+ m averagely. The eastern extrusion of the crust in the Tibetan



47 Plateau, associated with the wedged Eastern Himalayan syntaxis, initiates a clockwise
48 rotation of crustal deformation in the southeastern margin of the Tibetan Plateau (Figure
49 1) (Zhang et al., 2004; Gan et al., 2007; Wang and Shen, 2020). The Northwestern
50 Yunnan Region (NWYR), in the west part of the southeastern margin of the Tibetan
51 Plateau, borders the Tibetan Plateau, with the Lijiang-Xiaojiang fault serving as a
52 boundary fault that separates the Tibetan Plateau, boasting an average altitude of over
53 3000 meters, from the Yunnan Region, which maintains an average altitude of over
54 2000 meters (Yu et al., 2022; Zhang et al., 2022) (see Figure 1). Unlike thrust faults in
55 the plateau boundary, such as the Longmenshan fault ruptured by the 2008 $M_w7.9$
56 Wenchuan earthquake, the Holocene slip type of the Lijiang-Xiaojinhe fault is sinistral,
57 with a strike-slip rate of approximately 3 mm/yr, as determined by geological (Xu et al.,
58 2003; Shen et al., 2005; Ding et al., 2018; Gao et al., 2019) and geodetic data (Gan et
59 al., 2007; Cheng et al., 2012). The peculiar slip behavior of the Lijiang-Xiaojinhe fault
60 has garnered considerable attention in studies pertaining to crustal structure, fault
61 activities, and earthquake hazards (Xu et al., 2003; Cheng et al., 2012; Zhao et al., 2013;
62 Bao et al., 2015; Zhang et al., 2020; Huang et al., 2022; Zhang et al., 2022; Dai et al.,
63 2023). Zhang et al. (2020), a shear-wave velocity model was employed to reveal that
64 three faults - the Longmenshan fault, the Lijiang-Xiaojinhe fault, and the Chenghai fault
65 - delineate a low-velocity belt. This investigation unveiled the presence of low-crustal
66 flow beneath the Northwestern Yunnan Region (NWYR). Similarly, Zhang et al. (2022)
67 utilized magnetotelluric (MT) observations in the southern vicinity of the Lijiang-
68 Xiaojinhe fault, corroborating these findings and emphasizing the NWYR as a pathway



69 for ductile low-crustal flow. Analogously, a GPS study by Cheng et al. (2012) yielded
70 comparable results. Upon eliminating the rigid rotation component from the regional
71 GPS velocity field, they demonstrated a clockwise rotation propelled by ductile crustal
72 flow, particularly accelerated within the NWYR. They posited that this acceleration in
73 clockwise rotation might also be intensified by the tensional drag originating from the
74 Burma Plate.

75 The intricate network of crustal deformation encompassing the Northwestern Yunnan
76 Region (NWYR) introduces complexity to the slip behavior of faults and the focal
77 mechanisms of recent earthquakes. Within this area, three distinct fault slip behaviors
78 are observed: the NE-trending Lijiang-Xiaojinhe fault displays left-lateral strike-slip,
79 NW-trending faults exhibit right-lateral strike-slip, and North-South trending faults
80 demonstrate normal slip (see Figure 2). The presence of faults with diverse rupture
81 behaviors contributes to the complexity of earthquake hazards. Historically, these faults
82 have been associated with significant seismic events and numerous casualties. Notably,
83 three earthquakes with $M7+$ have occurred in the NWYR: the Yongsheng earthquake
84 of 1515 ($\sim M7.5$) on the Chenghai fault, the Midu earthquake of 1652 ($\sim M7$) on the Red
85 River fault, and the Dali earthquake of 1925 ($\sim M7$) on the Diancangshan East fault.
86 Additionally, the 1990 Lijiang earthquake ($M_S7.0/M_W6.6$) occurred on the Yulong East
87 fault, exhibiting dominant normal slip behavior. Historical and paleo-earthquake
88 studies suggest that nearly all of these faults have the potential to generate catastrophic
89 earthquakes (e.g., Ding et al., 2018; Ren et al., 2007; Chang et al., 2014), and induced
90 numerous landslides (Institute of Geology-State Seismological Bureau, and Yunnan



91 Seismological Bureau, 1990; Huang et al. 2021).

92 Fieldwork studies and focal mechanisms of recent earthquakes underscore the
93 complexity of fault slip behaviors in this tectonic environment (Figure 2). Both
94 historical and instrumental earthquakes have affected nearly all faults in the region,
95 emphasizing the seismic risks in NWYR. For instance, the 2013 Deqin earthquake
96 swarm, reaching a maximum magnitude of $M_S5.9/M_W5.7$ on August 31 (Wu et al.,
97 2015), and the 2021 Yangbi earthquake swarm, reaching a maximum magnitude of
98 $M_S5.9/M_W6.1$ on May 21 (Zhou et al., 2022), are noteworthy seismic events (refer to
99 Figure 2). The 2013 $M_W5.7$ Deqin earthquake swarm, characterized by tensional stress,
100 occurred at the intersection of the Zhongdian fault and the southern part of the
101 Jinshajiang fault, illustrating the susceptibility of the regional stress field to disturbance.
102 Conversely, the 2021 $M_W6.1$ Yangbi earthquake swarm occurred at the connection point
103 of the dominant dextral strike-slip faults, namely the Red River fault and the Weiqi-
104 Qiaohou fault, representing a different tectonic environment compared to the 2013
105 $M_W5.7$ Deqin earthquake swarm. This distinct setting suggests that either of these two
106 faults may be at risk of seismic activity during the pre-earthquake period.

107 Due to the high altitude, dense vegetation, and easily weathered conditions, obtaining
108 accurate fault slip rates poses a significant challenge, often resulting in notable errors.
109 Recent studies have provided fresh insights into slip rates and fault behaviors, offering
110 the potential to enhance the precision of seismic hazard models. For instance,
111 determining the dextral slip rate of the Zhongdian fault has proven particularly difficult
112 due to high error margins. Recent research has evaluated the Holocene dextral slip rate



113 to be $\sim 1.5 \pm 0.2$ mm/yr based on displacements of water-ice remains (Wu et al., 2019),
114 and $\sim 2.1 \pm 0.2$ mm/yr based on displacements of river terraces (Chang et al., 2014), both
115 utilizing Optically Stimulated Luminescence (OSL) dating. These values notably differ
116 from the right-lateral slip rate of 4~6 mm/yr estimated by Shen et al. (2001) based on
117 gully displacements from the last glacial period, but are more aligned with the rate
118 derived from GPS velocity data (Cheng et al., 2012). Incorporating these updated fault
119 slip rates into regional seismic hazard models holds the promise of enhancing their
120 accuracy. Therefore, integrating these new slip rates into the regional seismic hazard
121 model is crucial to ensure the reliability of the results.

122 Given the inherent challenges of fieldwork studies on fault activities, only a limited
123 number of investigations have been conducted regarding seismic hazard analysis in the
124 Northwest Yunnan region. Among these studies, Zhou et al. (2004) conducted a micro-
125 zonation of seismic hazards in the NWYR. They examined regional fault activities
126 through field surveys and estimated the potential maximum magnitude of these faults.
127 Their approach involved outlining polygons around the source faults to divide them
128 into different potential seismic sources and calculating historical seismicity rates within
129 these polygons. This methodology is widely employed in seismic hazard modeling in
130 China, particularly in the national seismic hazard map of the China Seismic Ground
131 Motion Parameters Zonation Map (CSGMPZM) (Gao et al., 2015). CSGMPZM also
132 utilized this methodology to assess potential maximum magnitudes and compute
133 seismicity rates. However, their studies often did not integrate fault geometry models,
134 especially fault segmentation models. Consequently, the fault geometry, including



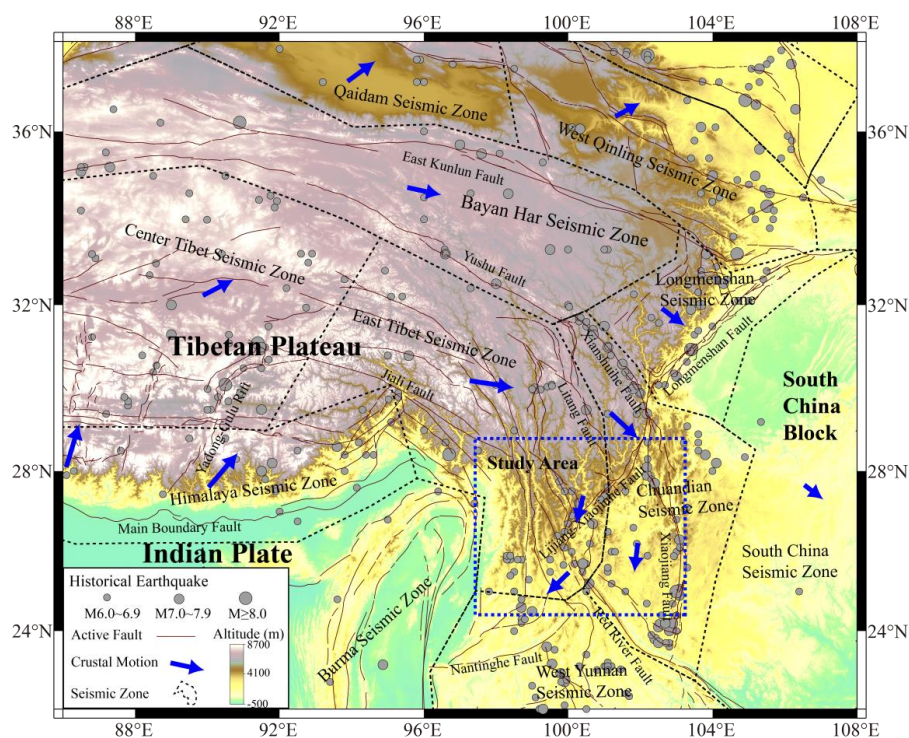
135 rupture length and area, may not be accurately linked to the magnitude of large
136 earthquakes. Furthermore, it is crucial to recognize the potential occurrence of multi-
137 segment rupturing, which has not been documented in historical records. Similarly,
138 seismicity rates were typically derived solely from historical earthquakes and were not
139 synchronized with fault slip rates. Relying solely on historical earthquakes for
140 seismicity rate calculations may lead to either overestimation or underestimation of
141 seismic hazards.

142 In this article, we developed a regional seismic hazard model for the Northwestern
143 Yunnan Region (NWYR), accounting for fault slip behaviors, the potential occurrence
144 of large earthquakes, and the likelihood of multi-segment ruptures. We initially
145 developed fault segmentation models for the primary active faults in the Northwestern
146 Yunnan Region (NWYR), drawing on recent geological research on fault segmentation
147 and geological fault slip rates. Subsequently, we employed the SHERIFS code (Chartier
148 et al., 2017; 2019) to simulate seismicity rates across possible multi-segment
149 combination models. We identified the multi-segment combination model that best
150 aligns with the majority of fault slip rates, considering fault segmentation and historical
151 seismicity rates. Ultimately, we calculated the Peak Ground Acceleration (PGA) with a
152 10% probability of exceedance within 50 years using the seismicity rates from the
153 selected fault segmentation models. The exploration of multi-segment rupture
154 combinations, along with the resultant modeled seismicity rates and PGA values, offers
155 valuable insights into the seismic hazard present in the NWYR. Leveraging the modeled
156 PGA values, we employed a machine learning model to compute the probability



157 distribution of landslides induced by potential seismic hazards. This increased precision
158 and reliability will be invaluable for guiding disaster preparedness initiatives, land-use
159 planning, and infrastructure resilience strategies in the area.

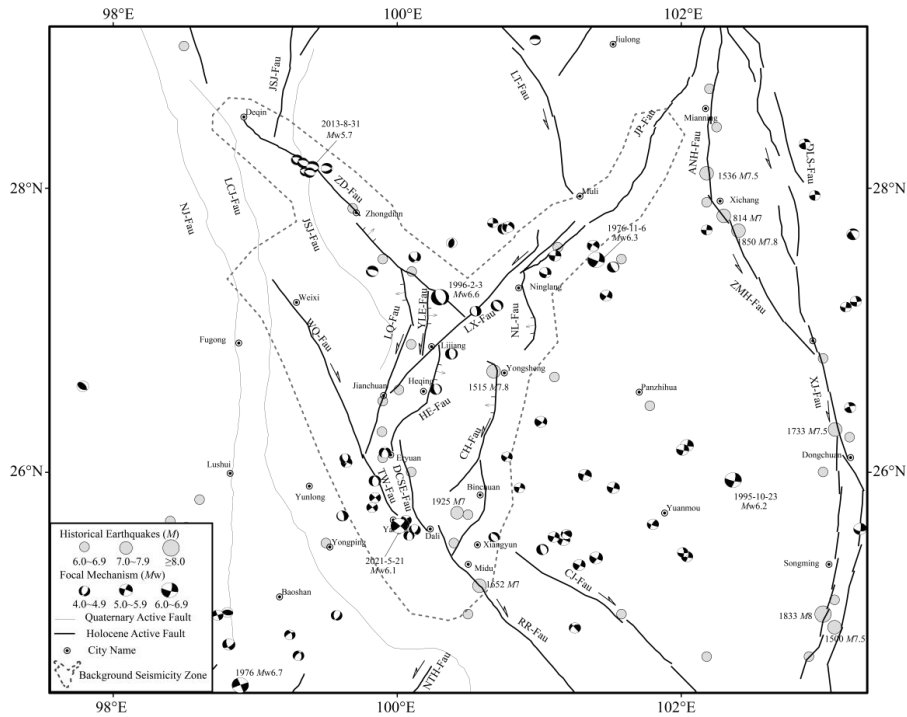
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161

162 Figure 1. Tectonic environment of the Eastern Tibetan Plateau and the location of the
163 Northwestern Yunnan Region (NWYR). The dashed rectangle delineates the study area,
164 while dashed polygons depict the seismic zones delineated by Rong et al. (2020).

165



166

167 Figure 2. Regional active faults and historical earthquake activities in the NWYR. The
 168 focal mechanisms of recent earthquakes (1976–2023) are sourced from the global
 169 centroid moment tensor (GCMT) catalog. Earthquakes with $M6+$ are sourced from the
 170 moment magnitude (M_w) catalog of Cheng et al. (2017). The dashed line represents the
 171 division for background seismicity calculation, which extends 20 km from the faults.
 172 JP-Fau: Jinping fault; LT-Fau: Litang fault; ANH-Fau: Anninghe fault; ZMH-Fau:
 173 Zemuhe fault; XJ-Fau: Xiaojiang fault; CJ-Fau: Chuxiong-Jianshui fault; RR-Fau: Red
 174 River fault; NTH-Fau: Nantinghe fault; DYJ-Fau: Dayingjiang fault; NJ-Fau: Nuijiang
 175 fault; LCJ-Fau: Lancangjiang fault; JSJ-Fau: Jinshajiang fault; ZD-Fau: Zhongdian
 176 fault; LX-Fau: Lijiang-Xiaojinhe fault; WQ-Fau: Weixi-Qiaohou fault; YLE-Fau:
 177 Yulong East fault; LQ-Fau: Longpan-Qiaohou fault; HE-Fau: Heqing-Eryuan fault;
 178 CH-Fau: Chenghai Fault; DCSE-Fau: Diancangshan East Fault; TW-Fau: Tongdian-
 179 Weishan Fault.

180



181 **2. Fault slip rates, Segmentation, and Multi-segment Rupture Combinations**

182 **Models**

183 **2.1 Fault slip rates, Segmentation**

184 In the NWYR, the Lijiang-Xiaojinhe fault, characterized by its left-lateral strike-slip
185 rate, and the northern segment of the Red River fault, which displays significant right-
186 lateral strike-slip movement, play pivotal roles in crustal deformation. Moreover, as a
187 result of the southward extrusion of the Tibetan Plateau, NE-trending faults such as the
188 Lijiang-Xiaojinhe fault also manifest a left-lateral strike-slip component. These
189 observations underscore the complex interplay of fault dynamics in the NWYR, as
190 elucidated by previous studies (Gan et al., 2007; Cheng et al., 2012; Wang and Shen,
191 2020).

192 To counterbalance the southwestward crustal extrusion (Wang et al., 1998; Cheng et al.,
193 2012), several other faults in the region, such as the Chenghai fault, the Ninglang fault,
194 the Heqing-Eryuan fault, the Yulong East fault, and the Longpan-Qiaohou fault (also
195 known as the Jianchuan fault), also exhibit a component of normal slip rate as well
196 (Institute of Geology-State Seismological Bureau, and Yunnan Seismological Bureau,
197 1990; Han et al., 2004). In contrast, the Zhongdian fault and the northern part of the
198 Red River fault, including the Weixi-Qiaohou fault and the Diancangshan East fault,
199 exhibit right-lateral strike-slip movement (Zhou et al., 2004; Han et al., 2005). Recent
200 focal mechanisms of intermediate earthquakes indicate a complex regional stress field,
201 featuring both strike-slip and normal faulting regimes (Figure 2). Table S1 provides an
202 overview of the observed fault slip rates in the NWYR.



203 The Lijiang-Xiaojinhe fault serves as a boundary fault delineating the Tibetan Plateau
204 from the Central Yunnan block (Xu et al., 2003; Cheng et al., 2012). We divided the
205 Lijiang-Xiaojinhe fault into 10 segments (the F1~F10 segments in Figure 3) based on
206 fault geometry and its intersection with other faults. For the F1 segment, known as the
207 Jinpingshan fault, recent fault mapping reveals a Holocene left-lateral slip rate ranging
208 from 1.3~2.7 mm/yr derived from gully displacement across the segment, while the
209 vertical slip rate is approximately 0.2 mm/yr (Mr. Rui Ding, 2024, private
210 communication).

211 Regarding the F5~F10 segments, Gao et al. (2019) demonstrated that the Hongxing-
212 Jianshanying segment (F6 segment in Figure 3) exhibited a Holocene left-lateral slip
213 rate of 3.32 ± 0.22 mm/yr with a normal slip rate component of 0.35 ± 0.02 mm/yr,
214 whereas the Runan-Nanxi segment (F10 segment in Figure 3) had a Holocene left-
215 lateral slip rate of 2.37 ± 0.20 mm/yr. Accordingly, we applied the slip rate of the F6
216 segment for the F4~F7 segments and the slip rate of the F10 segment for the F8~F10
217 segments. Notably, we considered the strike-slip motion of the F5~F10 segments to
218 originate from two sources: the strike-slip Jinpingshan fault and the strike-slip of the
219 Litang fault, aligning with the observed clockwise rotation of regional crustal
220 deformation around the Litang fault and the Lijiang-Xiaojinhe fault. Consequently, we
221 inferred the left-lateral strike-slip rate of the F4 segment to be ~ 2.1 mm/yr, consistent
222 with the southern section of the Litang fault (Zhou et al., 2007). However, the F2 and
223 F3 segments, which link the F1 and F4 segments, lack recorded fault slip rates from
224 fieldwork studies. In this regard, we assigned a conservative estimate, employing half



225 the value of the strike-slip rate of the F1 segment for both the F2 and F3 segments,
226 approximately 1.2 mm/yr.

227 For the Longpan-Qiaohou fault (comprising the F11~F14 segments), we delineated it
228 into four distinct segments based on the fault mapping data provided by Wu et al. (2023).
229 The sinistral slip rate of the Longpan-Qiaohou fault was estimated at ~2.2 mm/yr over
230 the past 3500 years, with a normal slip rate of 0.23 mm/yr (Institute of Geology-State
231 Seismological Bureau, and Yunnan Seismological Bureau, 1990).

232 As for the Yulong East fault, we segmented the fault into two segments, namely the F15
233 and F16 segments, utilizing fault mapping data and Quaternary sedimentary
234 distribution. The slip rate of the Yulong East fault was assessed by Han et al. (2005),
235 who determined that the Quaternary left-lateral and normal slip rates are 0.84 mm/yr
236 and 0.70 mm/yr, respectively, derived from the displacement observed in a gully
237 crossing the fault.

238 Regarding the Zhongdian fault, we partitioned it into six segments, designated as the
239 F17~F22 segments, based on fault mapping data (Wu et al., 2023). The Holocene
240 dextral slip rate of the Zhongdian fault is estimated to be approximately 1.7-2.0 mm/yr,
241 with a minor normal slip rate of 0.6-0.7 mm/yr based on terrace displacement across
242 the fault (Chang et al., 2014).

243 For the Heqing-Eryuan fault, we segmented it into two sections, labeled as the F23 and
244 F24 segments. The Quaternary dextral slip rate and normal slip rate of the Heqing-
245 Eryuan fault were reported to be around 2 mm/yr and 0.7~1.0 mm/yr, respectively, as



246 documented by the Institute of Geology-State Seismological Bureau, and Yunnan
247 Seismological Bureau (1990). Additionally, recent research by Sun et al. (2017) yielded
248 similar fault slip rate results, indicating a left-lateral slip rate of 1.80 mm/yr and a
249 vertical slip rate of 0.28 mm/yr since the Pleistocene.

250 The Ninglang fault is primarily characterized as a left-lateral strike-slip fault, although
251 it exhibits a minor normal slip component of less than 0.1 mm/yr at the basin margin.

252 The strike-slip rate of the Ninglang fault, as determined from fault mapping work
253 conducted by Dr. Panxing Yang from Institute of Earthquake Forecasting, China
254 Earthquake Administration (private communication), was estimated to be less than 1
255 mm/yr. For our analysis, we opted to utilize a median value of 0.5 ± 0.4 mm/yr for the
256 strike-slip rate of the Ninglang fault. Based on the distribution of Quaternary sediments,
257 we divided the Ninglang fault into two distinct segments, designated as the F25 and
258 F26 segments.

259 For the Chenghai fault, the sinistral slip rate has been estimated to range from 2.5 to
260 3.0 mm/yr, determined from the erosion rate of the Jinshajiang River crossing the fault.

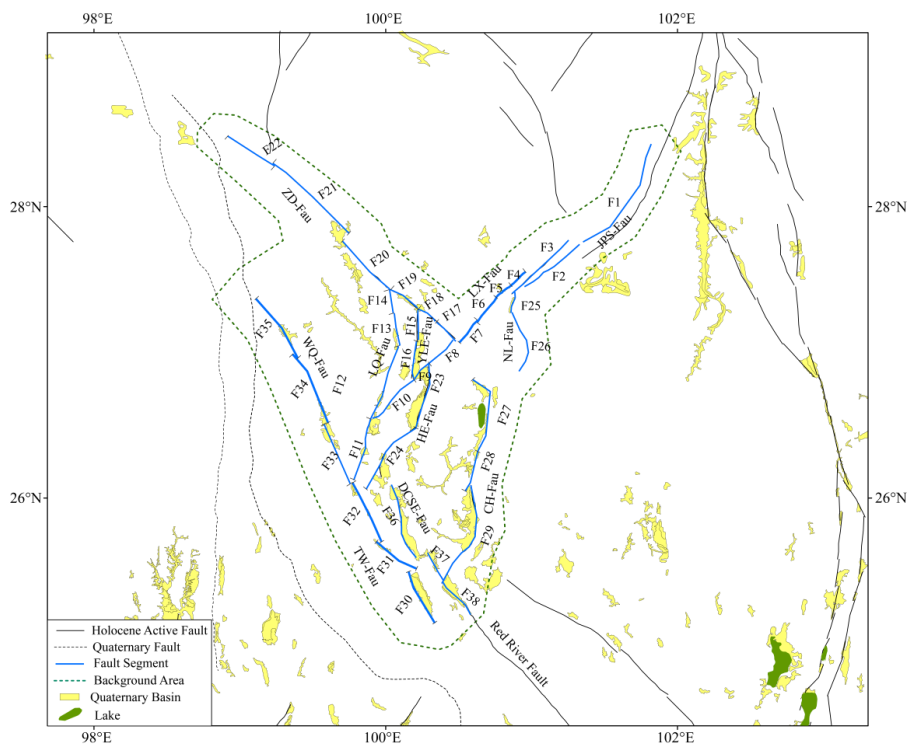
261 Additionally, the normal slip rate is reported to be between 0.7 and 1.0 mm/yr, assessed
262 from the lift rate of the fault scarps (Institute of Geology-State Seismological Bureau,
263 and Yunnan Seismological Bureau, 1990), which is consistent with the findings of Tang
264 et al. (2017). We divided the Chenghai fault into three segments, i.e., the Chenghai
265 segment (the F27 segment), the Qina segment (the F28 segment), and the Bingchuan
266 segment (the F29 segment), based on the sedimentary distribution (Huang et al., 2018;
267 Yu et al., 2005).



268 The southern end of the Longpan-Qiaohou fault separates the Weixi-Qiaohou fault from
269 the Tongdian-Weishan fault. We segmented these two faults into six segments each
270 based on fault mapping data and Quaternary sedimentary distribution. Concerning the
271 slip rate of the Tongdian-Weishan fault, the dextral slip rate in the Late Pleistocene is
272 estimated to be $\sim 1.8\text{-}2.4$ mm/yr, with a normal slip rate of $0.17\text{-}0.35$ mm/yr, calculated
273 from the displacement of fault scarps (Chang et al., 2016). In contrast, for the Weixi-
274 Qiaohou fault, the dextral slip rate is ~ 1.25 mm/yr, while the normal slip rate is ~ 0.91
275 mm/yr since the Late Pleistocene (Ren et al., 2007). Comparing these rates to the dextral
276 slip rate from the middle section of the Red River fault, which is reported to be $1.1 \pm$
277 0.4 mm/yr (Shi et al., 2018), it is evident that the dextral slip rates decrease from the
278 northwest to the southeast across the Red River fault system, encompassing the Weixi-
279 Qiaohou fault, the Tongdian-Weishan fault, and the Red River fault.

280 The Diancangshan East Fault is the seismogenic fault of the 1925 $M7$ Dali earthquake.
281 We deduced that the Diancangshan East fault is a dominant normal slip fault as the
282 boundary fault of the Dali basin and the Erhai Lake. The normal slip rate of this fault is
283 $1\text{-}2$ mm/yr (Guo et al., 1984; Zhou et al., 2004).

284 Additionally, we incorporated the F37 and F38 segments of the northern part of the Red
285 River fault into our segmentation model. The right-lateral strike-slip rate of these two
286 segments is ~ 1.1 mm/yr. Figure 3 illustrated the segmentation model of the faults in the
287 NWYR.



288

289 Figure 3. Fault segmentation model for the Northwestern Yunnan Region (NWYR).

290 In which, the Quaternary Basin distribution were from Deng et al. (2003); the fault

291 data are from Wu et al. (2023).

292

293 2.2 Multi-segment Rupture Combinations Models

294 Based on the segmentation model, fault rupture behaviors, and the intersections
295 among fault segments, we devised four multi-segment rupture combination models for
296 the fault segments in the NWYR (Figure 4). The $M_w6.6$ Lijiang earthquake on February
297 3, 1996, represents a significant normal rupture event that occurred on the Yulong East
298 fault. This earthquake stands out as the most substantial seismic event in the NWYR
299 since the 1970s, underscoring the normal slip behavior of the Yulong East fault. This



300 observation suggests potential implications for the rupture behavior of the Zhongdian
301 fault.

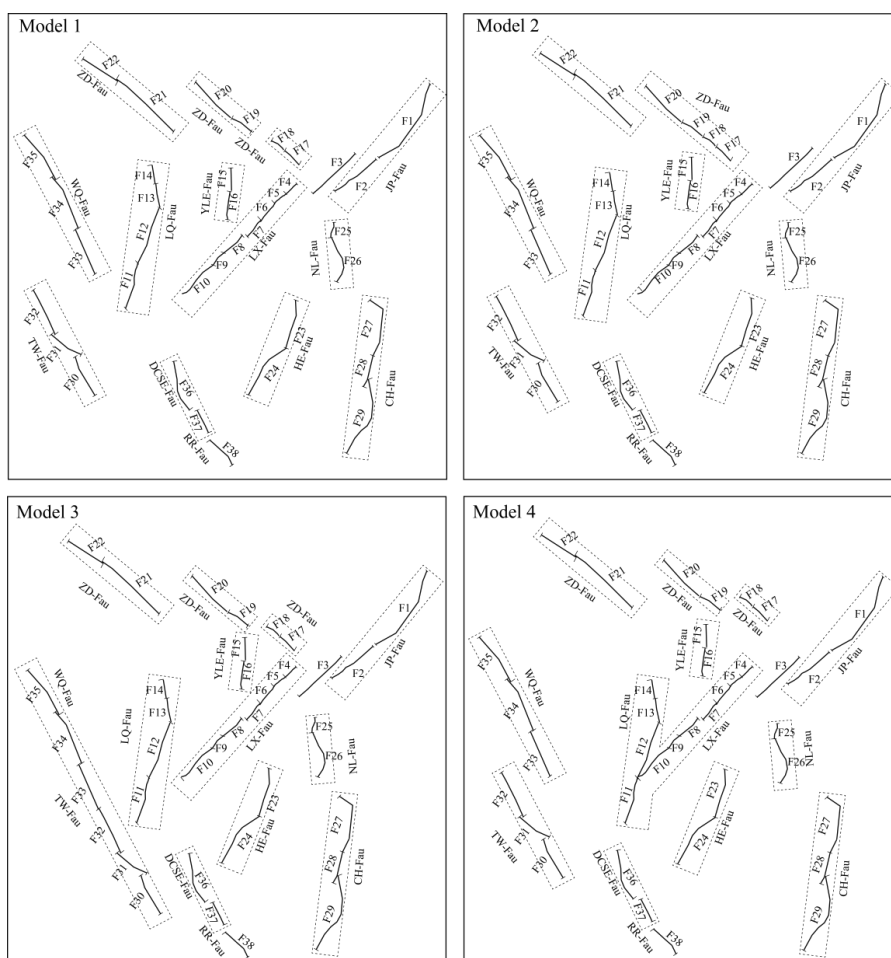
302 In Model 1, we exclusively examined the multi-segment rupture combinations
303 within the same faults. Specifically, for the Zhongdian fault, we integrated the multi-
304 segment rupturing of the F17 and F18 segments, as well as the multi-segment rupturing
305 of the F19 and F20 segments. This approach considered the normal slip behavior of the
306 Yulong East fault (F15 and F16 segments) and its potential impact on Quaternary
307 sedimentary distribution between the F18 and F19 segments of the Zhongdian fault.
308 Subsequently, in Model 2, we evaluated the plausibility of multi-segment rupturing
309 occurring across the F17~F20 segments.

310 In the NWYR, the prevailing fault behavior is sinistral slip along the northeast-
311 trending faults, a trend consistent with the observed clockwise rotation in regional
312 crustal deformation (Cheng et al., 2012) and the presence of ductile low-crust flow
313 (Zhang et al., 2022). The sinistral slip observed along the Longpan-Qiaohou fault may
314 hinder the dextral slip occurring along the Weixi-Qiaohou fault, which extends from the
315 Tongdian-Weishan fault, contributing to the decrease in dextral slip rates observed from
316 the Weixi-Qiaohou fault to the Tongdian-Weishan fault. In Model 3, we integrated the
317 multi-segment rupture combination of the Weixi-Qiaohou fault (the F33~F35 segments)
318 and the Tongdian-Weishan fault (the F30~F32 segments).

319 In 2023, two earthquakes of $M_w7.8$ and $M_w7.5$ successively ruptured the East
320 Anatolia fault region in Turkey (Xu et al., 2023; Petersen et al., 2023). The rupture of
321 the first earthquake, with $M_w7.8$, initiated on the splay Narli fault and propagated
322 bilaterally along the main East Anatolia fault (Liu et al., 2023). Consequently, we took
323 into account the possibility of rupture propagation from one fault to another in our
324 rupture combinations. Using Model 4, we investigated whether the rupture on the



325 Lijiang-Xiaojinhe fault could propagate to the Longpan-Qiaohou fault. This
326 consideration was prompted by similarities in the rupture behavior between the F11
327 segment of the Longpan-Qiaohou fault and the F10 segment of the Lijiang-Xiaojinhe
328 fault, along with a minor difference in strike (Table S1).



329
330 Figure 4 Possible Rupture Combination Models for the Fault Segments in NWYR.
331 Dashed rectangular show the rupture combinations for each Model.
332

333 3. Multi-segment rupture hazard Modeling



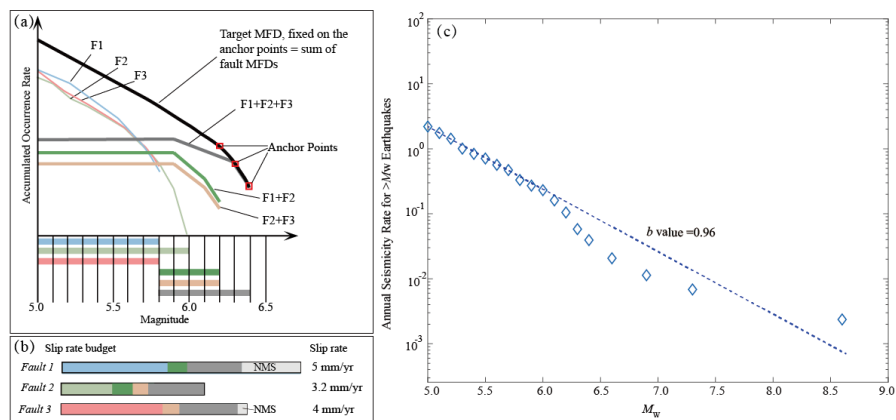
334 3.1 Methodology

335 In the earthquake hazard modeling, the seismicity rate of the earthquakes should
336 reflect both the fault slip rate and the magnitude-frequency distribution (MFD), e.g.,
337 the Gutenberg-Richter (G-R) Relationship (Gutenberg and Richter, 1944) and the
338 Characteristic earthquake model (Schwartz and Coppersmith, 1984). Youngs and
339 Coppersmith (1995) balanced the fault slip rates and the magnitude-frequency
340 relationship in the seismicity rate on the faults. They employed the composite
341 characteristic earthquake model (Y-C) or truncated G-R model to convert the fault slip
342 rate into the seismicity rate on the fault. These converted MFD were widely used in
343 seismic hazard analysis (e.g., Avital et al., 2018; Chartier et al., 2019; Rong et al., 2020).
344 This approach allows for a more comprehensive assessment of earthquake hazards by
345 integrating both fault slip rates and the frequency of seismic events.

346 For assessing the possibilities and probabilities of multi-segment rupturing, it is
347 essential to represent the seismicity rate of such combinations in the magnitude-
348 frequency relationship for each segment. To achieve this, Chartier et al. (2017; 2019)
349 devised a Python-based code known as SHERIFS. This code employed an iterative
350 process, enabling the balancing of occurrence rates for multi-segment rupturing events
351 alongside intermediate and small earthquakes on each fault segment (Figure 5a).
352 Leveraging historical seismicity data, they utilized the slip rate of each fault segment to
353 convert it into the target MFD, such as the G-R, or the Y-C distribution. This method
354 offered a robust framework for assessing seismic hazard, integrating both single and
355 multi-segment rupture scenarios effectively. Determining the maximum magnitudes of
356 individual fault segments and their combinations could rely on fault length, following
357 rupture scaling laws proposed by researchers like Wells and Coppersmith (1994) and
358 Cheng et al. (2020).



359 In the final step, they iterated the seismicity rates across magnitude bins associated
360 with multi-segment rupturing, spanning from large magnitudes down to small
361 magnitudes, according to the target MFD for each fault segment. However, in many
362 cases, the fault slip rate or calculated seismicity rates couldn't fully account for the
363 entire seismic activity. The remaining portion of the fault slip rate for each segment was
364 attributed to non-main-shock slip (NMS), including processes like post-seismic slip and
365 silent creep. A non-main-shock (NMS) ratio of $\leq 30\%$ – 40% was typically considered
366 indicative of a model misfit, potentially due to creeping and specific conditions such as
367 boundary fault segments or creeping segments (as depicted in Figure 5b). Here, we
368 adopted a similar approach in simulating seismic hazard modeling for the regional fault
369 system in the NWYR.



370 **Figure 5** a. Scheme of the occurrence rate iterative process on the fault segments
371 constrained by the magnitude-Frequency relationship and fault slip rates; b. Slip
372 budget of the fault slip rate and its consumption on the earthquakes (Modified from
373 Chartier et al., 2017); c. Calculated b value for the East Tibet Seismic Zone where
374 the NWYR located in Figure 1.

376

377 Given the fractured structure of the crust in the NWYR, as documented by Cao et



378 al. (2023), the seismicity distribution in this area was notably complex and differed
379 significantly from that observed directly on the fault lines. Therefore, in our analysis of
380 seismicity rates for the whole seismicity rates on the regional faults, we opted to utilize
381 the Gutenberg-Richter (G-R) relation (Gutenberg and Richter, 1944) as the Magnitude-
382 Frequency relationship, rather than the Youngs-Coppersmith (Y-C) relation (Youngs
383 and Coppersmith, 1985).

384 For estimating the magnitudes based on rupture length, we applied the relationship
385 proposed by Cheng et al. (2020) to determine the maximum magnitude for each
386 individual fault segment as well as their multi-segment combinations. Additionally, we
387 accounted for a portion of earthquakes with $M < 6.5$ as off-fault seismicity. Specifically,
388 we assigned probabilities of 95%, 90%, 85%, 80%, and 80% for magnitude bins ranging
389 from 6.0 to 6.4, 5.5 to 5.9, 5.0 to 5.4, 4.5 to 4.9, and 4.0 to 4.4, respectively, based on
390 prior studies (Chartier et al., 2019; Cheng et al., 2021).

391 We conducted a calculation of the b-value for the East Tibet Seismic Zone, which
392 encompasses nearly all of the NWYR, as illustrated in Figure 1. The earthquake catalog
393 utilized for this analysis was sourced from Cheng et al. (2017), covering the time period
394 from 780 BC to 2015 AD. Additionally, we incorporated earthquakes from the Global
395 CMT catalog spanning the period from 2016 to 2023 into the dataset. The regressed b-
396 value was approximately 0.96, with completeness times for magnitudes $M_w 4.5$, $M_w 5.0$,
397 $M_w 5.3$, $M_w 5.7$, $M_w 6.1$, and $M_w 6.4$ identified as 1985, 1966, 1928, 1916, 1916, and
398 1900, respectively. It's worth noting that the calculated b-value is slightly higher than
399 the value of 0.86 reported in Rong et al. (2020), likely due to the inclusion of new
400 earthquakes occurring after 2015. Figure 5c provides a visualization of the Gutenberg-
401 Richter relationship in the East Tibet Seismic Zone, in which the b value is 0.96.

402



403 3.2 Comparison of Modeled Results

404 We depicted the NMS ratios and modeled seismicity rates in Figure 6. The right
405 panels showcased the NMS ratios of the segments in Model 1~4. Model 1 exhibited the
406 most balanced results between the modeled seismicity rates and historical ones. In
407 Figure 6a, all segments in Model 1 demonstrated NMS ratios smaller than 30%.
408 Chartier et al. (2019) suggested that NMS ratios below 30%-40% serve as a benchmark
409 to assess the validity of multi-segment combination models, indicating effective
410 consumption of the slip rate of each segment into seismicity rates for each fault segment.
411 The left panels in Figure 6b further underscored the harmony between the modeled and
412 observed seismicity rates. Here, the observed historical seismicity rates closely aligned
413 with the calculated ones, particularly for <M7 earthquakes.

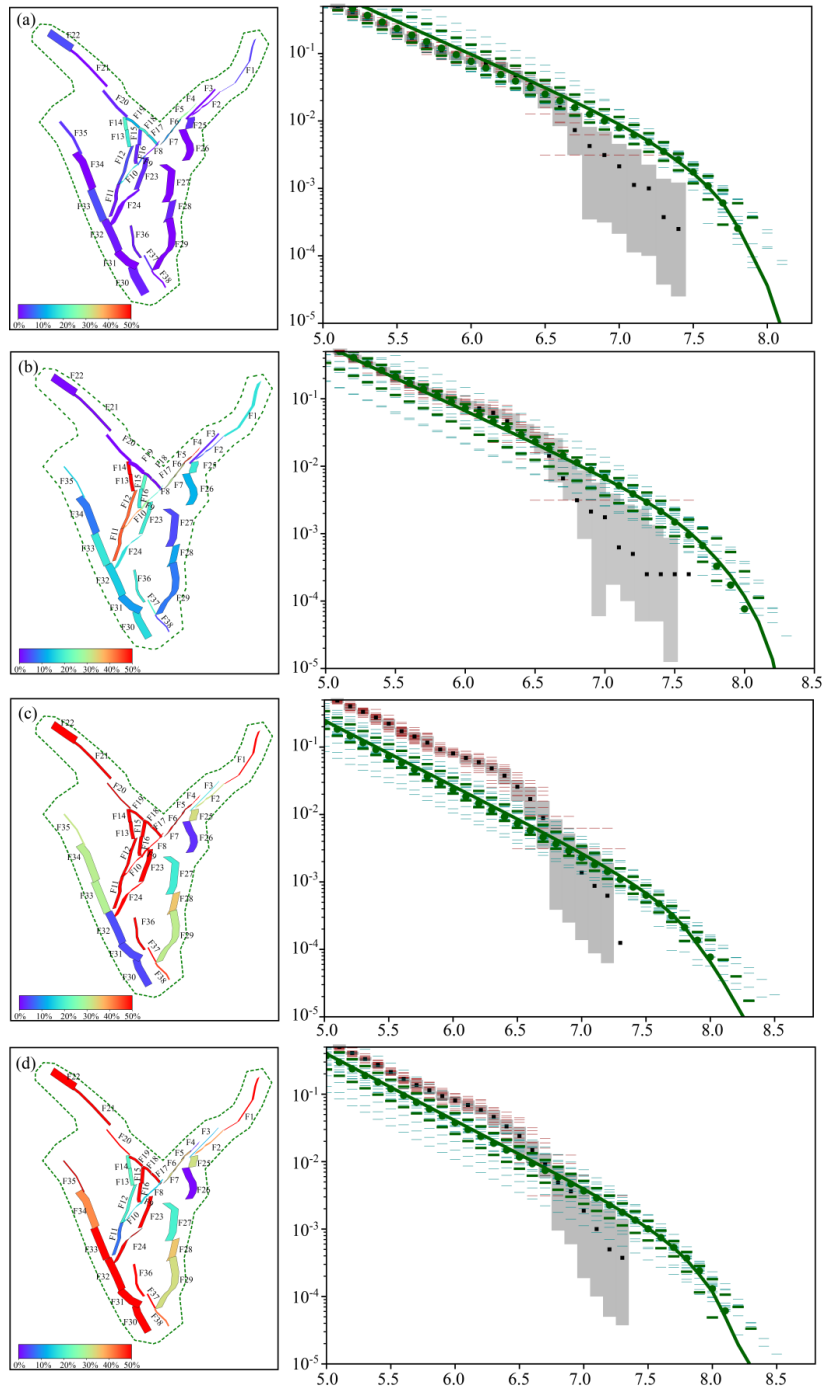
414 Compared to Model 1, Model 2 combined segments F17~F20 as a single unit,
415 instead of considering the F17~F18 segments and the F19~F20 segments separately.
416 The left panel of Figure 6b indicates that the NMS ratios for segments of F11~F14 and
417 F4~F5 are all greater than 40%, while the F6~F7 multi-segment combination has an
418 NMS ratio ranging from 30% to 40%, showing that the combination of segments
419 F17~F20 has an impact on the seismicity rates of these faults. From the right panel in
420 Figure 6b, the historical seismicity rates for each fault segment were similar to those in
421 Model 1. However, the calculated seismicity rates for each segment in Model 2 became
422 smaller than those in Model 1, except for a slightly higher rate in the magnitude range
423 of 6.0~6.5. This result indicated that the fault slip rates are not being adequately
424 accounted for in Model 2, unlike in Model 1 (Figure 6a).

425 In Model 3, the rupture combination comprised segments of F30~F35, rather than
426 considering them separately as F30~F32 and F33~F35. Most segments exhibited high
427 NMS ratios in the left panel of Figure 6c. The calculated seismicity rates were generally



428 smaller than the historical ones in the right panel of Figure 6c. Similarly, Model 4 was
429 utilized to investigate whether the great earthquakes of the Y-shaped rupture, combining
430 segments F4~F10 with F11~F14, could occur. The NMS ratios for each segment and
431 the calculated seismicity rates were comparable to those observed in Model 3 (Figure
432 6d).

433 In addition, we also presented the results using the rupture scaling relationship
434 proposed by Wells and Coppersmith (1995) in Figure S1. Model 1 exhibited the most
435 consistent outcomes, with the maximum NMS ratio observed on F14 at 39.3%. The
436 NMS ratios for all other segments were below 30%. For the calculated seismicity rates
437 obtained from Model 2 to Model 4 using the rupture scaling relationship of Wells and
438 Coppersmith (1995), we observed similar patterns. All three models showed segments
439 with NMS ratios exceeding 40%. Furthermore, we found that these models utilizing the
440 rupture scaling of Wells and Coppersmith (1995) consistently yielded higher NMS
441 ratios on average compared to those obtained from the rupture scaling of Cheng et al.
442 (2020).



443
444 Figure 6 Calculated NMS ratios and comparison results for different models using the



445 G-R relation. (a) Modeled Non-Mainshock Slip (NMS) Ratio; (b) Comparisons
446 between the historical Seismicity rates for different models. Dashed green lines are the
447 MFD of each model, and the solid green line is the mean MFD, green patches represent
448 the uncertainty (16-84 percentiles). The dotted black line is the rate from the catalog;
449 the dashed red lines are sampled rates of the catalog exploring the uncertainties on the
450 magnitudes of earthquakes, and gray rectangular show the one-sigma uncertainty on the
451 earthquake rates in statistical analysis.

452
453 Based on the comparison among different rupture combination models, Model 1
454 demonstrated the most consistent results among the multi-segment rupture
455 combinations, fault segment slip rates, and the Magnitude-Frequency relationship.
456 Therefore, we utilized the seismicity rates from Model 1 to calculate the Peak Ground
457 Acceleration (PGA) values for the NWYR.

458

459 **3.3 Comparison with the results of current national seismic hazard map**

460 We utilized the OpenQuake Engine v3.10 (Pagani et al., 2014) to calculate the
461 Peak Ground Acceleration (PGA) values for the NWYR. In this computation, we
462 employed a logic tree model comprising the Abrahamson et al. (2014); Chiou and
463 Youngs (2014); Campbell and Bozorgnia (2014); and Boore et al. (2014) branches.
464 Each branch was assigned an equal weight of 0.25, following the selection criteria
465 established by Dangkua et al. (2018) for mainland China. These Ground Motion
466 Prediction Equations (GMPEs) are tailored for earthquakes characterized by moment
467 magnitude (M_w) and the distance to the rupture plane (R_{rup}) or its surface projection
468 (R_{JB}).

469 Figure 7a illustrated the distribution of Peak Ground Acceleration (PGA) for the



470 site condition of firm to hard rock ($V_{s30}=760$ m/s, or NEHRP B) resulting from the
471 seismicity model in Model 1, corresponding to a 10% probability of exceedance in 50
472 years, which is equivalent to a return period of 475 years. The analysis revealed
473 concentrations of high values exceeding 0.40 g near fault sources, particularly in areas
474 with multiple fault sources. These areas include the F2~F5 segments of the Lijiang-
475 Xiaojinhe fault, the vicinity of the Yulong East Fault (YLE-Fau), the southern part of
476 the Zhongdian fault (ZD-Fau), and the northern extent of the Heqing-Eryuan fault (HE-
477 Fau).

478 The area of the F2~F5 segments includes three parallel faults, with the sum of the
479 strike-slip of ~ 3 mm/yr, makes the PGA values relatively higher. The maximum
480 magnitude of the combinations of F17~F18 segments and the F15~F16 segments are
481 both approximately $M_w 6.6$. These areas exhibit a prevalence of moderate earthquakes
482 with short recurrence intervals and high Peak Ground Acceleration (PGA) distributions
483 over a 475-year period. The modeled seismicity rates of the F23 segment and the F24
484 segment both complied with the G-R relationship, containing enough intermediate
485 earthquakes, induced the high PGA values around. Along the Chenghai fault, high PGA
486 values are also observed around the F27~F28 segments with strike-slip rates of 3.0
487 mm/yr but are lower around the F29 segments with a strike-slip rate of 2.5 mm/yr. For
488 the Red River fault and its extensions, including the Tongdian-Weishan fault and the
489 Weixi-Qiaohou fault, high PGA values were concentrated around the F37~F38
490 segments and the intersection points of the F11, F32, and F32 segments.

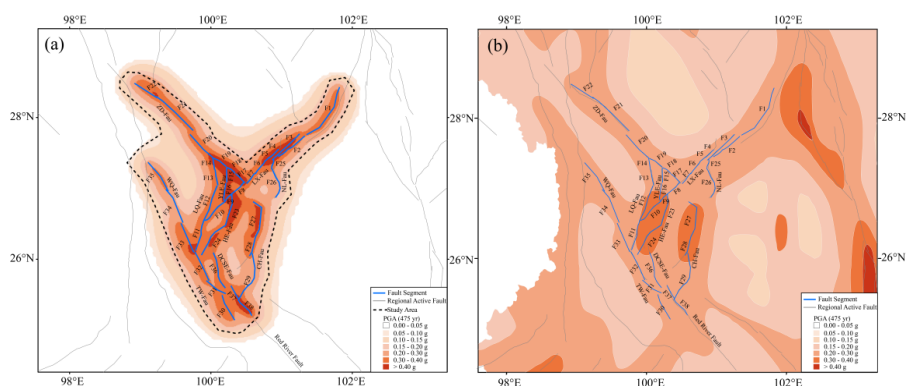
491 Comparison with the national seismic hazard map of the China Seismic Ground
492 Motion Parameters Zonation Map (CSGMPZM) (Gao et al., 2015) (Figure 7b) for the
493 site condition of dense soil and soft rock ($V_{s30} = 500$ m/s, or NHERP C) (Chen et al.,
494 2021), our Peak Ground Acceleration (PGA) values are consistently much higher and



495 more intricate. The V_{s30} of 500 m/s is equivalent to the Type II in the classification
496 table of CSGMPZM, while the value of 760 m/s belongs to Type I1. Table 1 was the
497 adjustment factors used by CSGMPZM for site amplification (Gao et al., 2015). Even
498 if we applied these site amplification adjustment factors to convert our PGA values
499 from type I1 to type II, the PGA values would not change obviously as the adjustment
500 are the near to 1 for PGA values of 0.30~0.40 g, and 1 for PGA values of ≥ 0.40 g. In
501 figure 7b, the CSGMPZM indicates two high PGA values ranging from 0.30 to 0.40 g
502 in the NWYR, specifically around the F23~F24 segments and the F27~F28 segments,
503 respectively. PGA values in other areas surrounding the fault segments in our model
504 range from 0.20 to 0.30 g. In the development of the CSGMPZM, the region in the
505 around China was divided into 29 large seismic source zones to calculate the parameters
506 of the Magnitude-Frequency Distribution (MFD). Additionally, over 1,000 potential
507 fault sources across China were incorporated into the model. Historical seismicity rates
508 on the MFD were employed to predict future seismic activity. This methodology led to
509 lower anticipated seismicity rates in regions with limited historical earthquake records.
510 The identification of potential fault sources in the CSGMPZM relied on expert opinions
511 gleaned from research on historical surface rupture, fault segmentation, and the
512 distribution of past earthquakes. These data sources were subsequently utilized to
513 allocate predicted seismicity rates based on the MFD. Furthermore, the utilization of
514 different Ground Motion Prediction Equations (GMPEs) in the CSGMPZM compared
515 to our results could also contribute to variations in PGA values. The CSGMPZM
516 utilized GMPEs from Yu et al. (2013) based on surface magnitude (M_s) and epicentral
517 distance (R_{epi}). Their GMPEs result in higher PGA values for distances less than 80
518 km but lower values for distances ≥ 80 km (Cheng et al., 2021). Consequently, the
519 seismicity rates derived from the fault slip rates and the multi-segment rupture



520 combinations were key factors that rendered our modeled PGA values higher than those
 521 from CSGMPZM.



522
 523 Figure 7 Comparison of the Modeled PGA distribution of 10% in the next 50 years,
 524 (a) the PGA results in this article (b) the PGA results in the CSGMPZM.
 525

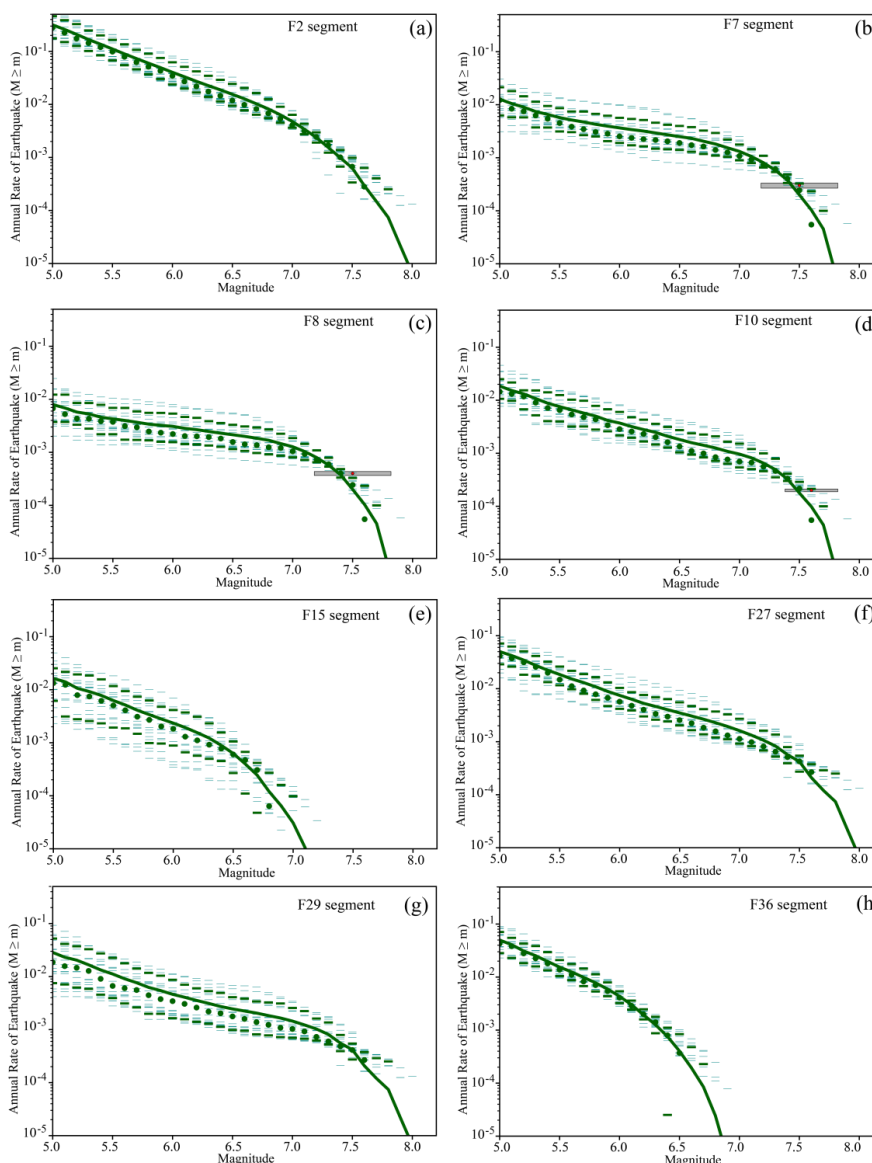
526
 527 Table 1 Adjustment factors for PGA values of different Site condition via Type II

PGA values for type II	Site condition type				
	I ₀	I ₁	II	III	IV
≤ 0.05 g	0.72	0.80	1.00	1.30	1.25
0.10 g	0.74	0.82	1.00	1.25	1.20
0.15 g	0.75	0.83	1.00	1.15	1.10
0.20 g	0.76	0.85	1.00	1.00	1.00
0.30 g	0.85	0.95	1.00	1.00	0.95
≥ 0.40 g	0.90	1.00	1.00	1.00	0.90

528
 529 In Figure 8, we further illustrated seismicity rates for several typical fault segments
 530 to elucidate the reasons behind the observed high PGA values. In Figure 8a, the
 531 seismicity rates of the F2 segment exhibit a typical G-R relationship, leading to a high
 532 PGA distribution in the surrounding area. We compared the seismicity rates on the
 533 F7~F8 segments, and the F10 segment with the recurrence intervals from paleo-
 534 earthquake studies. In Figures 8b~8c, the red bars illustrate that our modeled seismicity



535 rates align with the recurrence interval of approximately 3000 years for a magnitude
536 7.5 earthquake, as determined by Ding et al. (2018). We have also observed that
537 segments F7 to F8 of the Lijiang-Xiaojinhe fault tend to conform to the characteristic
538 earthquake model based on their seismicity rate distribution. Segment F10, with a
539 length of approximately 44 km, experienced rupture during the 1751 M6.8 earthquake.
540 Tang et al. (2014) identified three paleo-earthquake events with a recurrence interval of
541 around 5300 years for earthquakes of $M_{6.5+}$ earthquakes on segment F11. They
542 suggested that the two paleo-events before 1751 AD were considerably stronger than
543 the one in 1751 AD, implying multi-segment rupturing involving combinations of
544 segments F11+F12, F11~F13, and F11~F14 resulting in magnitudes of $M_w 7.4\sim 7.6$
545 earthquakes. Additionally, we illustrate the seismicity rates on segments F15, F27, F29,
546 and F36 in Figures 8e~8f, which closely resemble the G-R distribution, leading to high
547 PGA distributions in their vicinity. These results demonstrate that the occurrence rate
548 of intermediate earthquakes influences the high PGA distributions.
549



550
551 Figure 8 Modeled Seismicity rates for different magnitude on the fault segments. The
552 solid line is the mean MFD, and small patches represent the uncertainty (16-84
553 percentiles). The dotted line is the rate from the catalog with uncertainties. The red
554 circle is the occurrence rate of the repeated large historical earthquake rate, and the gray
555 box is the associated uncertainty.



556

557 **3.4 Landslide Probabilities**

558 Utilizing the modeled PGA values for rock site conditions presented in Figure 7a
559 as a foundation, we enhanced our analysis by incorporating the site amplification effect
560 derived from Chen et al.'s (2021) comprehensive site condition map. Their research,
561 leveraging geological unit data, culminated in a detailed site condition map covering
562 mainland China. Leveraging this invaluable resource, we integrated their site condition
563 map along with the amplification factors for each geological type compared to type II
564 (referenced in Table 1) to refine the PGA value distribution map (Figure 9a). Our
565 methodology involved multiplying the PGA values for specific site conditions by the
566 ratio of type II PGA values to those of the specific type. This approach effectively
567 magnified PGA values across different site conditions, enriching the granularity of our
568 analysis. Figure 9a illustrates the resultant PGA distribution map, now encompassing
569 site amplifications specifically tailored for the NWYR region. Notably, our findings
570 reveal minimal alterations in the PGA distribution, particularly in proximity to fault
571 lines, where PGA values remain consistent or exceed 0.4 g (as detailed in Table 1).

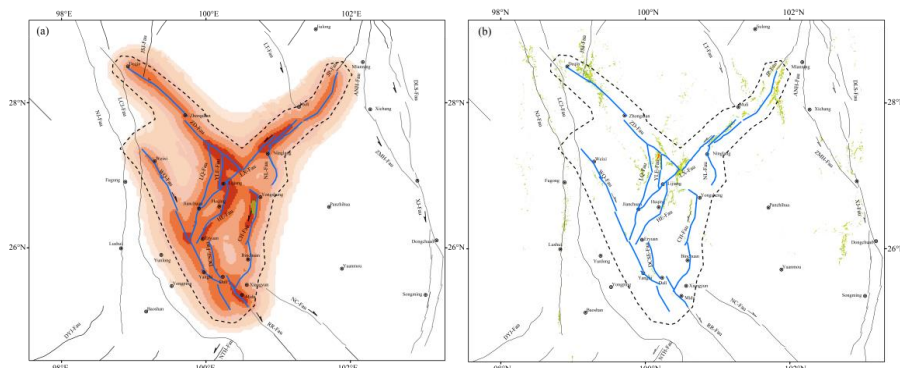
572 Using simulated ground motion data from potential earthquake scenarios, we
573 conducted a thorough assessment of landslide susceptibility in the affected regions. Our
574 analysis employed a machine learning framework, following the methodology outlined
575 by Xu et al. (2019), to develop a predictive model for earthquake-induced landslides.
576 This model was trained utilizing data from nine earthquakes, ranging from the 1999
577 $M_w7.7$ Chichi earthquake to the 2017 $M_w6.5$ Jiuzhaigou earthquake, all of which
578 occurred within or near China. The training dataset comprised samples of earthquake-
579 induced landslides alongside 13 relevant factors. These factors encompassed diverse
580 parameters such as elevation, slope angle, slope aspect, land cover, proximity to faults,



581 geological characteristics, average annual rainfall, and PGA. Leveraging this rich
582 dataset, we constructed a robust predictive model capable of discerning landslide
583 probabilities.

584 Figure 9b illustrates the resultant landslide probability map for the NWYR region.
585 Notably, areas exhibiting high PGA distribution correspond closely to regions with
586 elevated landslide probabilities. For instance, notable areas include the northern end of
587 the Zhongdian fault, the Jinpingshan fault, the Yulong East fault, the northern end of
588 the Heqing-Eryuan fault, the northern part of the Chenghai fault, and the eastern section
589 of the Lijiang-Xiaojinhe fault (the F2~F4 segments). Of particular significance are
590 regions surrounding the Yulong East fault and the convergence zone of the Lijiang-
591 Xiaojinhe fault and the Zhongdian fault. These areas exhibit pronounced differences in
592 altitude, ample rainfall, and elevated PGA values, making them particularly susceptible
593 to landslide occurrences.

594 By integrating multiple geospatial factors and leveraging advanced machine
595 learning techniques, our analysis provides valuable insights into landslide susceptibility
596 in earthquake-prone regions, aiding in effective risk management and mitigation
597 strategies.



598
599 **Figure 9. (a) PGA distribution Map considering different site amplifications; (b) the**
600 **probabilities of landslide occurrence impacted by the PGA values.**
601



602 **4 Discussion and Conclusion**

603 In seismic hazard analysis, understanding fault slip behaviors, including slip rates
604 and fault geometries, is pivotal for accurately modeling future seismicity rates.
605 Concurrently, historical earthquake occurrence rates provide a foundation for
606 estimating these future rates. Of significant note, attention must also be directed towards
607 earthquakes involving multi-segment ruptures, which may not be documented in
608 historical records. In this article, we unveil a new seismic hazard model for the NWYR,
609 where the boundary of the Tibetan Plateau intersects with local low-crustal flow.

610 **4.1 Multi-segment Rupturing Hazards in NWYR**

611 The complex fault system results in earthquake occurring almost all the faults with
612 various rupture behaviors in the NWYR, while the catalog of historical and paleo-
613 earthquake data only recorded a small part of the rupturing events. The NWYR serves
614 as the boundary region between China and Myanmar. This area is predominantly
615 inhabited by ethnic minorities in China, resulting in limited written documentation of
616 its history, particularly regarding earthquake disasters. However, some significant
617 earthquakes have been documented, particularly those that have had a seismic impact
618 on major cities like Dali, e.g., the 1515 $M7.8$ earthquake in Yongsheng ruptured two
619 continuous segments of the Chenghai fault (Institute of Geology-State Seismological
620 Bureau and Yunnan Seismological Bureau, 1990). The historical earthquake catalog
621 used in our seismic hazard modeling often struggles to include all these combinations
622 of ruptured scenarios that occurred in the past. What we have done is to search for
623 possible rupture combinations and calculate their seismicity rates to include in our
624 model. These rupture combinations might be constrained by various factors, such as the
625 geometry of fault segments, the width of the step-over between each pair of segments,
626 and the maturity of the fault steps (Cunningham and Mann, 2007; Biasi and Wesnousky,



627 2017). For strike-slip faults, a width of 5 km is often used to assess the reasonableness
628 of the rupture combinations (e.g., Biasi and Wesnousky, 2017). However, in the NWYR,
629 where faults are located in the conduit of ductile low-crust flow, all step-overs have
630 widths of less than 5 km except the one of ~ 7 km between F20 and F21 segments.
631 Hence, we advocate that the intersection relationship between faults is the primary
632 determinant of whether multi-segment rupture events occur among fault segments in
633 this region.

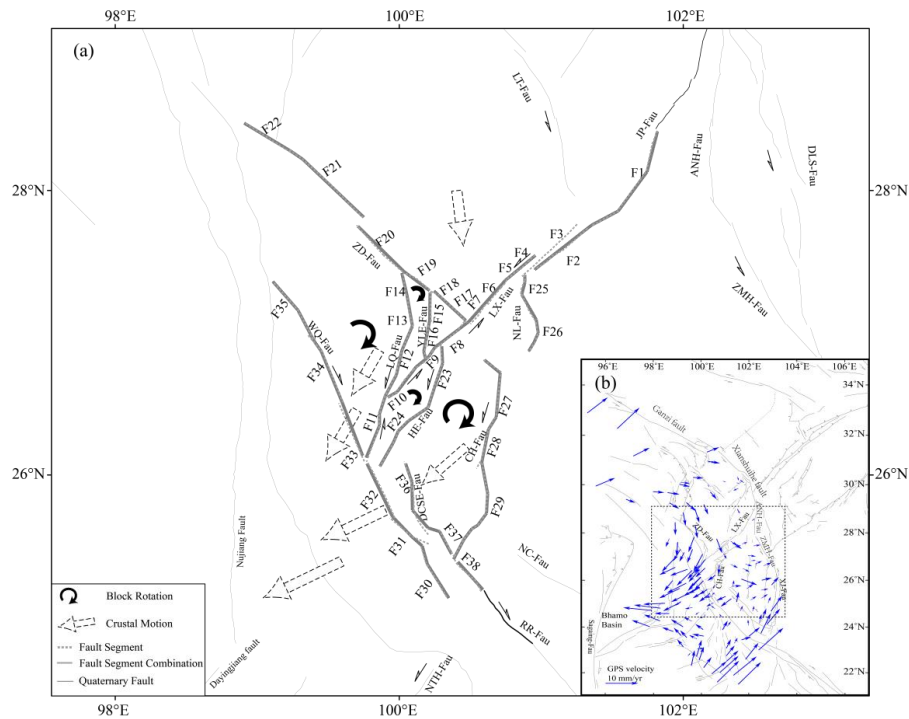
634 **4.2 Implication of the small-block rotation in NWYR**

635 The Holocene strike-slip motion of the Lijiang-Xiaojinhe fault behaves the
636 dominant role in this region, and intersects the Heqing-Eryuan fault and the Yulong East
637 fault. Model 1 also confirmed the capability of the entire rupture of the F4~F10
638 segments of the Lijiang-Xiaojinhe fault. The Chenghai fault and the Zhongdian fault
639 also are separated by the Lijiang-Xiaojinhe fault, which differs the view of Wang et
640 al. (1998) that the Dali fault (including the Longpan-Qiaohou fault and the Chenghai
641 fault) is the primary fault in this region. The Longpan-Qiaohou fault obstructs the
642 westward continuation of the Lijiang-Xiaojinhe fault, and simultaneously, the F11
643 segment also resists rupturing in conjunction with the Lijiang-Xiaojinhe fault (Model 4
644 in Figure 4). In contrast, the Weixi-Qiaohou fault (WQ-F) and the Tongdian-Weishan
645 fault (TW-F) are part of distinct small-blocks and therefore cannot rupture
646 simultaneously, as depicted in Figure 10a. This indicates that the northern end of the
647 Red River fault is intercepted by the Longpan-Qiaohou fault. The Zhongdian fault (ZD-
648 F) was separated to rupture in our model (Model 1 in Figure 4), especially for the
649 F17~F18 segments combination and the F19~F20 segments combination. Here, we
650 propose that the normal- and strike-slip of the Yulong East fault poses a greater
651 destructive potential to the Zhongdian fault compared to the strike-slip of the Longpan-



652 Qiaohou fault.

653 Hence, our configurations of multi-segment ruptures portrayed in Model 1 of
654 Figure 4 correspond to the rotational patterns noted in the small block delineated in the
655 NWYR by Wang et al. (1998). We illustrated this clockwise rotation of the small
656 blocks in the NWYR in Figure 10a. This clockwise rotation was further supported by
657 GPS observations to the west of the Xianshuihe fault, the Anninghe fault, and the
658 Xiaojiang fault, after eliminating the entire movement (Figure 10b) (Cheng et al., 2012).
659 In Figure 10b, the area where the Nujiang fault intersects with the Dayingjiang fault
660 experiences the strongest extensional forces. Rangin et al. (2013) and Lindsey et al.
661 (2023) proposed that the dynamic source of this extensional tectonic environment was
662 the side effect of the gravitational collapse of the Tibetan Plateau with the westwards
663 of upper crust faster than the lower crust (Rangin et al., 2013; Lindsey et al., 2023).
664 This extensional force exerts obviously on the faults in our model, making the rotation
665 of small blocks, and the normal slip of the regional faults, e.g., the Diancangshan fault
666 and the Chenghai fault.



667

668 Figure 10 a. Kinematic Model of the faults in the NWYR; b. Regional GPS
669 motion after removing the whole movement (Cheng et al., 2012).

670

671 In conclusion, our study has provided valuable insights into the seismic hazard
672 present in the NWYR. By developing fault segmentation models based on recent
673 geological research and utilizing advanced simulation techniques, we have enhanced
674 our understanding of fault activities and seismicity rates across the region. Through
675 careful analysis and consideration of fault segmentation, fault slip rates, and historical
676 seismicity, we have identified multi-segment models that best represent the observed
677 data. Our calculations of PGA with a 10% probability of exceedance within 50 years
678 offer crucial information for assessing seismic risk in the NWYR. The PGA values,
679 associated with obvious latitude difference, abundant precipitation, are prone to
680 occurrence of landslides.



681 Furthermore, our investigation into multi-segment rupture combinations has
682 illuminated potential scenarios for seismic events in the region. Through the integration
683 of these findings, we have generated a more comprehensive assessment of seismic
684 hazards and landslide probabilities. These are intertwined with the regional small block
685 rotation induced by the low-crustal flow and gravitational collapse along the
686 southeastern frontier of the Tibetan Plateau.

687 Moving forward, continued research and monitoring efforts are essential for
688 refining our understanding of seismic hazards in the region. Further investigations into
689 fault behaviors, fault interactions, and the potential for multi-segment ruptures will be
690 crucial for enhancing the accuracy and reliability of seismic and induced geological
691 hazard assessments. By remaining vigilant and proactive in our approach to seismic risk
692 management, we can better protect communities and infrastructure in the face of future
693 seismic events in the NWYR and beyond.

694

695 *Code availability*

696 In this study, we have used the code related to Chartier et al. (2019,
697 <https://doi.org/10.1785/02201803320>), which can be downloaded from the webpage
698 (<https://doi.org/10.1785/02201803320>, last accessed in May, 2024).

699

700 *Data availability*

701 The focal mechanism data are from Global CMT catalog (www.globalcmt.org, last
702 accessed in May, 2024) Table S1 in the supplementary material for this paper includes
703 the fault segments, historical and paleo-earthquakes and their associated slip parameters.

704



705 *Author contributions.*

706 Jia Cheng was responsible for methodology, software, and writing the original draft.

707 Chong Xu worked for the landslide occurrence probabilities calculation. Xiwei Xu

708 and Shimin Zhang contributed to design the fault rupture combination models.

709 Pengyu Zhu contributed to seismic hazard modeling.

710

711 *Competing interests.*

712 The authors declare that they have no known competing financial interests or personal

713 relationships that could have appeared to influence the work reported in this paper.

714

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726

727

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