

Seismic Signal ~~Classification~~ Characterization of Snow Avalanches using Distributed Acoustic Sensing in Grasdalen, Western Norway

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Abstract. We show the ~~usage~~use of Distributed Acoustic Sensing for analyzing seismic signals from snow ~~avalanche events~~avalanches. For three winter seasons, we continuously recorded seismic data using Distributed Acoustic Sensing (DAS) on a ~~section of a~~ standard telecommunication fiber along a mountain road in Grasdalen, western Norway. Multiple snow ~~avalanche events~~avalanches were registered, alongside various other signals, such as road traffic and detonations from remote avalanche triggering. We describe the signal characteristics of natural and manually triggered ~~avalanche events~~avalanches and present a comparison with other observed signals in both time and frequency domain. Our frequency analysis shows that avalanche signals are most visible between ~~20-50~~20 - 50 Hz.

For larger avalanches, we observe weak low-frequency precursor signals, which correspond to the ~~avalanches approach~~avalanches approaching towards the fiber. The more prominent high-amplitude signals appear to be produced by the snow masses impacting the stopping cones or the steep terrain near the road. ~~In one natural avalanche event, we interpret distinct spike signals as likely corresponding to the stopping snow mass, based on similar findings from previous studies using geophones. For a smaller, naturally released avalanche, we observe impulsive signals, interpreted as the snow mass stopping. These results agree with earlier studies that used geophone measurements.~~

Automatic detection, tested with simple STA/LTA thresholds in the ~~20-50~~20 - 50 Hz range, presents challenges due to false positives from road traffic. Further refinement and testing are required to improve detection reliability in this complex environment. Our study represents an initial exploration into the application of Distributed Acoustic Sensing for snow avalanche detection, showcasing its potential as an effective monitoring tool for long road networks in mountainous regions.

1 Introduction

Snow avalanches and landslides pose a considerable risk ~~for to~~ road safety and infrastructure in Norway (Bjordal and Larsen, 2009). According to Lunde and Njå (2019), this problem will become even more prevalent in the following years with ~~the~~-ongoing climate change. Avalanche monitoring systems can help improve road safety by enabling immediate and automatic road closures when necessary. By gathering real-time observational data about avalanche activity, these systems can reduce the risk of vehicles entering ~~avalanche-prone areas. Additionally~~avalanche-prone areas. In addition, assessing whether vehicles were present near the danger zone during an avalanche helps determine ~~if~~-whether a rescue mission is required.~~Hence, gathering observational data about avalanche activity is an important factor, optimally leading to a functioning monitoring and warning system.~~

Avalanche detection and monitoring have previously been done using either optical remote sensing or acoustic/seismic methods. Optical systems such as doppler-radar (Schreiber et al., 2001) or webcams (Fox et al., 2024) require direct line of sight and are limited in their spatial coverage. In the case of webcams the detection capability can be strongly decreased by bad weather conditions such as heavy snowfall. The second type of detection systems makes use of acoustic or seismic waves generated by ~~avalanche events~~avalanches (Suriñach et al., 2001). They offer long detection distances up to a few ~~kilometres~~kilometers (Marchetti et al., 2019) and immunity to bad weather conditions. Additionally, no direct line of sight is necessary, ~~making it a reliable tool for automatic detection.~~

Seismic signals of snow avalanches were first recorded by St. Lawrence and Williams (1976). Using observations from Valle de la Sionne in Switzerland, Suriñach et al. (2001) showed that seismic sensors can be used for studying avalanche characteristics. Analyzing seismic data from the same test site, Biescas et al. (2003) demonstrated that running frequency spectra show a gradual increase in higher frequency signals as the ~~avalanche approaches the geophone~~avalanches approach the geophone placed in a bunker at the valley floor, resulting in a triangular shaped spectrum. This was observed for large avalanches that reached the sensor. Smaller avalanches which stopped 1000 m or further away from the sensor did not exhibit this triangular shape. Pérez-Guillén et al. (2016) studied frequency signatures of both dry and wet avalanches using geophones. For powder snow avalanches, the majority of seismic energy was found below 10 Hz, produced by the turbulent flow part at the front of the avalanche. For wet snow avalanches however, most of the seismic energy was generated by the dense flow regime, resulting in signals with a frequency content mainly above 10 Hz. Recording and ~~localizing~~locating multiple avalanche signals can be achieved by utilizing a geophone array in combination with beam-forming algorithms (Lacroix et al., 2012) or statistical models (Heck et al., 2018).

Distributed Acoustic Sensing (DAS) is a technology that utilizes optical fibers for measuring dynamic strain or strain rate along the cable's axial direction. This is achieved by sending laser pulses into the fiber and recording the backscattering, subsequently extracting phase differences between different back-scattered light bits. The phase difference measurement can either be performed between the backscattering response from adjacent regions of fiber for the same incident pulse or between successive pulses but for the same region of fiber. Phase differences are then converted to strainrate or strain (Waagaard et al., 2021). DAS has been successfully tested and proven useful in a wide range of seismic applications, from oil and gas exploration (Liu et al.,

2024) over monitoring of wells (Horst et al., 2015), urban activity (Lindsey et al., 2020), tides (Buisman et al., 2023), whales (Bouffaut et al., 2022) to microseismic events (Walter et al., 2020). It offers both high spatial and temporal resolution, with sensing distances in the meter range (Waagaard et al., 2021) and a frequency resolution between 0.01 Hz (Taweesintananon et al., 2023) and 50 ~~kHz~~-kHz (Parker et al., 2014), depending on the interrogator. Since existing telecommunication fibers can be utilized there is a great potential for monitoring applications without the immediate need of new installations. The key advantage over local installations like radar, geophones, and infrasound systems is that with a single interrogator, more than 100 km of fiber can be monitored continuously (Waagaard et al., 2021), potentially enabling the observation of multiple snow avalanche sites using just one setup.

First detection tests of triggered snow avalanches with DAS were conducted by Prokop et al. (2013). Using a fiber following an avalanche path, Fichtner et al. (2021) were able to observe snow avalanche signals and differentiate between flow regimes. Paitz et al. (2023) employed an unsupervised clustering algorithm (Bayesian Gaussian Mixture Model) on a set of features extracted from DAS avalanche recordings from the same test site. Amongst other findings, they could recognize seismic far field signatures before the avalanche hit the fiber. This gives hope for a realistic monitoring system where the sensor position does not necessarily lie in the avalanche path but is instead situated some distance away from it, such as a telecommunication fiber in a valley. ~~As mentioned, previous experiments were conducted on test sites with fibers following the avalanche paths. With readily available telecommunication fibers along exposed roads there is a large potential of increasing avalanche monitoring systems without the need for further installations.~~ Within this paper, we present results from a seismic avalanche test monitoring system in western Norway based on Distributed Acoustic Sensing (DAS), utilizing a standard telecommunication fiber along a mountain road as the receiver.

2 Monitoring Site Description

During the winter seasons of 2021/2022, 2022/2023 and 2023/2024 we recorded DAS data in the mountain valley of Grasdalen in the region of Innlandet, Norway. The measurement period was approximately 4 months for each season. The DAS survey was conducted on a 2 km section of a single mode standard telecommunication fiber buried next to the road. ~~Overview of the Grasdalen Survey Setup. Note that the different sections of the fiber are color-coded based on the surroundings for better visual discriminability. The same color coding is used in alle subsequent figures. The altitude difference between neighbouring contour lines is 100 m. Aerial footage by Norwegian Mapping Authority (2020)~~

The environment ~~of surrounding~~ the road changes throughout the array: ~~In the beginning and the end~~ At both ends, the road is going into tunnels, ~~whereas in the middle~~. In the middle section it is housed in a concrete gallery that shields it from rockfall and snow avalanches.

On the part southwest of the bridge, the road runs next to 10 m high conical breaking mounds protecting it from avalanches (labeled as cones) ~~protecting it from avalanches~~. On both sides of the gallery, steel cattle grids are embedded into the road. Due to the changing surroundings we expect to observe different noise levels throughout the array. The ~~different cable environments are setup is~~ displayed in Figure 1 ~~as colored lines. The mountainside is simultaneously monitored with a commercial, infrasound-based avalanche detection system located 100 m east of the bridge. Additionally, a webcam located near the bridge was used for verification of avalanche activity and extent of avalanche runout zones. and in Figure A1 (bird's eye perspective).~~

The DAS array is monitoring ~~two avalanche prone mountainsides: the east and the southeast face of~~ the avalanche prone east face of the mountain called Sætreskarsfjellet. ~~The first one~~ It stretches from 1300 m.a.s.l. down to the valley floor at 875 m.a.s.l. Between 950 m.a.s.l. and 1250 m.a.s.l. the slope is mostly between 30 and 35 degrees steep, with rocky bands in between that are up to 45 degrees steep (NVE, 2024). Avalanches can be manually triggered by setting off explosives. The detonation point is located at the top of the mountainside at 1300 m.a.s.l. (see Figure 1) ~~A small section of the southeast face of Saetreskarsfjellet is also facing the open road south of the avalanche gallery. The altitude difference in this part is up to 175 m (between 1000 m.a.s.l. at the top and 835 m.a.s.l. next to the road) with a steepness of up to 45 degrees. The southeast mountain face above the southern tunnel is considerably higher and oftentimes steeper than 50 degrees based on (NVE, 2024).~~

The mountainside is simultaneously monitored with a commercial, infrasound-based avalanche detection system located 100 m east of the bridge. We do not have direct access to the infrasound data and can only use the event entries that are available on Varsom Regobs. Additionally, a webcam located near the bridge was used for verification of avalanche activity and extent of avalanche runout zones.

Based on previous avalanche reports from Grasdalen (~~Varsom Regobs~~) on Varsom Regobs (norwegian web portal for registering and sharing natural hazard warnings), the altitude difference and ~~steepness~~ slope, we expect avalanches of up to size 3 (size level based on European Avalanche Warning Services (3/1/2024)) ~~on the east as well as the southeast face~~. This means avalanches of a length over several 100 m with a total volume of 10000 m³, capable of crossing the road and destroying cars. ~~On the small stretch of open road below the southeast face avalanches of size of 2 can be expected. 3D view of the~~

105 ~~Grasdalen valley and the Survey setup (fiber sections are marked with same colors as in Figure 1). The inset image shows the east face of Saetreskarsfjellet as seen from the bridge. Avalanche prone areas are highlighted in orange, the detonation point is marked with a red star. The parking lot at start of the northeastern tunnel is marked with a 'P'. Aerial footage by Norwegian Mapping Authority (2020); inset image by Google (2024).~~

~~An ASN OptoDAS interrogator was used for data acquisition. The~~ We recorded strain rate with an ANS OptoDAS interrogator
110 with 2 m channel spacing, variable gauge lengths between 3.1 m and 5.1 m and sampling rates of 250 and 500 Hz for different seasons. The gauge length in our system is defined as the length of fiber over which phase measurements of the backscattered optical signal are spatially averaged (Waagaard et al., 2021). The acquisition parameters utilized for each season can be found in Table 1. ~~For the last (2023/24) season, the sampling rate was reduced to 250 Hz. This adjustment was made because higher frequency data were not deemed necessary, since the dominant avalanche-related energy is found below 100 Hz (see below~~
115 ~~in the next section). Additionally this led to more manageable file sizes during data handling. The interrogator is located at a~~ service point inside the southern tunnel (see Figure 1).

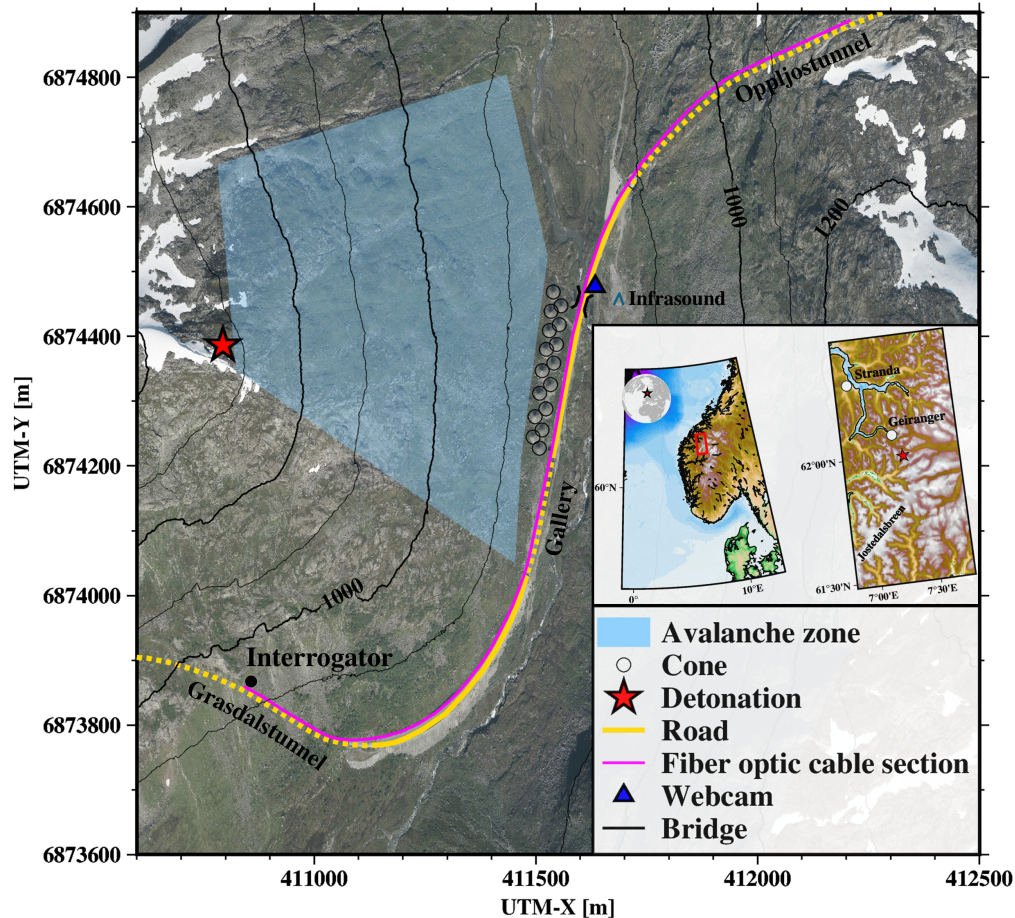


Figure 1. Acquisition parameters Overview of the Grasdalen survey setup. The tunnels sections of the road (and fiber) are displayed as dashed lines. The altitude difference between neighbouring contour lines is 100 m. Aerial footage by Norwegian Mapping Authority (2020)

3 Signal Classification

3.1 Types of Signals

During the surveying period ~~the following types of signals were observed along the DAS array : Road traffic, detonations from avalanche triggering-,~~ the DAS array recorded distinct seismic signals originating from road traffic, avalanche-triggering explosions, and snow avalanches (both natural and artificially triggered ~~snow avalanches, and background noise-~~), characterized by differences in spatial patterns, amplitude distributions, and frequency content. Avalanche signals are particularly dominant in the 20 - 50 Hz range and localized amplification of traffic signals is observed near the bridge and cattle grids.

This section provides a comparison between these signal types, followed by an in-depth analysis of selected ~~avalanche events-~~ avalanches. The avalanches were identified in the DAS data using the avalanche catalog from Varsom Regobs (Varsom Regobs). This catalog is based on observations from webcams, ~~an-~~ the infrasound system in the valley and reports from human observers, including personnel from the state road authority responsible for avalanche risk assessments. ~~An overview about the four most important recorded signal types in terms of number of appearance and importance-~~ An overview of four representative signals is displayed in Figure 2. Figure 2a illustrates the signal from an artificially triggered avalanche on ~~February-10, 2022-~~ The February 2022 09:01 (UTC). The detonation occurs approximately 40 s earlier and is displayed separately (Figure 2c).

The avalanche impacted the cones and ~~subsequently-~~ covered the road ~~section-~~ between the bridge and the gallery with a half-meter thick layer of snow (RegObs Avalanche Observation, 10.02.2022). The signal is divided into two distinct areas: the ~~initial smaller slide, first~~ starting at 10 seconds between 800 and 880 m (marked with 1 in Figure 2a), and the ~~main slide second~~ (marked with 2) between 650 and 775 m.

Figure 2b shows the signal of a naturally released snow avalanche from ~~February-2, 2024-~~ February 2024 17:21 (UTC). Three distinct features are observed: The first event between 7.5 s and 9 s (marked with 1) shows the highest amplitudes. Since those are observed west of the bridge near the cones, we interpret this as the slide hitting one or multiple of the cones. The second sharp event (2) occurs at the end of the open road section close to the tunnel. Since there is a steep incline from the riverbed up to a parking lot next to the tunnel entrance, this signal ~~likely corresponds to the snow mass could be generated by the~~ avalanche impacting the ramp of the parking lot (marked with 'P' in Figure A1). The recurring activity in the area marked with 3 is interpreted as multiple smaller slides going off after the first one. Additionally, four short events (encircled in blue) with relatively high amplitudes, each lasting approximately 0.5 s, were detected in this area. ~~Those could be connected to parts of the avalanche coming to a stop-~~

Figure 2c shows signals generated by cars passing the valley from northeast to southwest on ~~February-10, 2022-~~ February 2022 11:07 (UTC). In total, 15 individual traffic sources can be identified. By comparing these signals with webcam-confirmed traffic events, we conclude that the stronger amplitudes (marked with a 'T') likely correspond to trucks, while the other signals are attributed to passenger cars. For reference, a confirmed truck signal is added in the supplementary figures (A2). The strength of the individual traffic signals fluctuates across the array sections, with the gallery area appearing to be less sensitive. ~~This variation is likely due to differences in the cable environment-~~

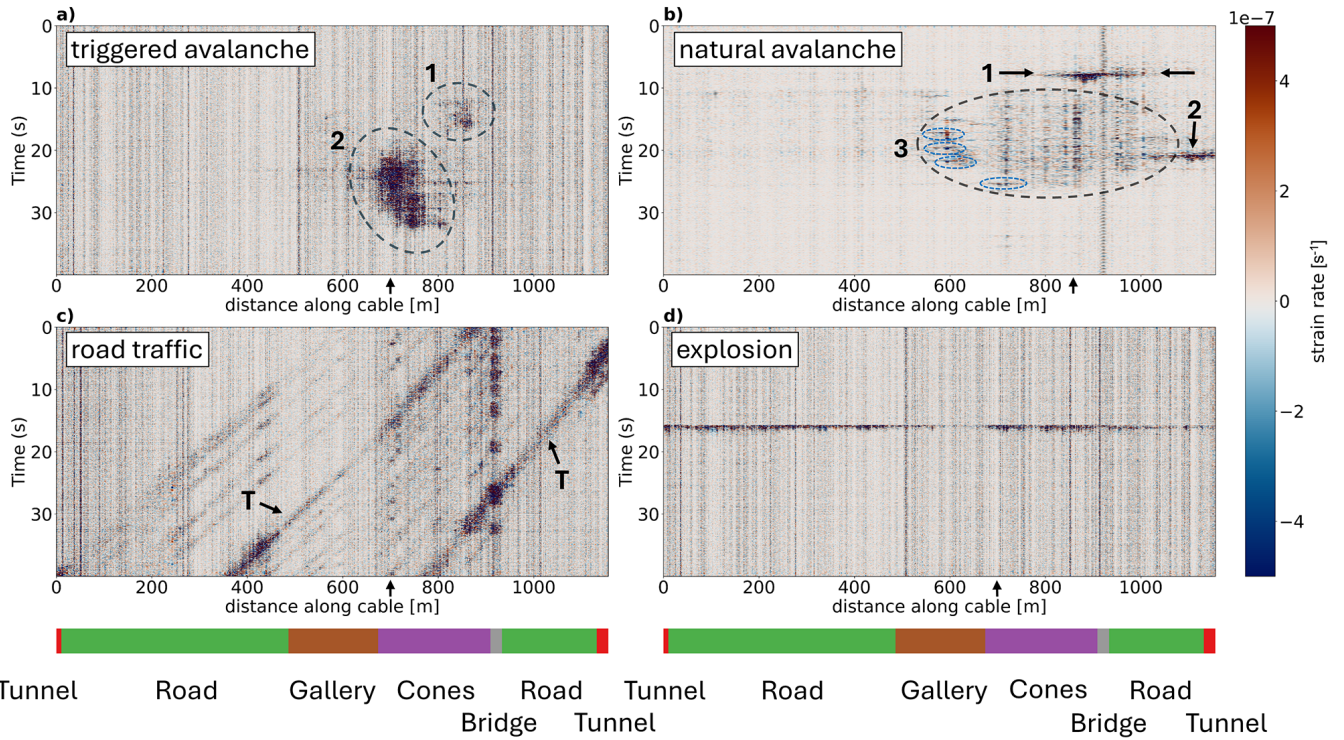


Figure 2. Strain-rate plots of four different signal types: a) Triggered snow avalanche from 10 /02/February 2022 09:01 (UTC) resulting in two signals that are marked with (1) and (2) b) natural snow avalanche from 02/02/2 February 2024 17:21 (UTC) with three distinct features (marked with 1,2 and 3). Within the third event group, four noticeable spike events are encircled in blue. c) traffic signals from 10 /02/February 2022 involving a truck signal-11:07 (UTC) including two strong signals visible across all channels, associated with heavier vehicles like trucks (marked with a 'T'). d) explosion signal from 10 /02/February 2022 09:00 (UTC) which triggered the avalanche shown in a). A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. Arrows mark the single channels that were used for frequency analysis (Figure 3).

The bridge section, in particular, exhibits noisy signals, which are further amplified when traffic passes over it. Similarly, the cattle grids at 10 m, 430 m, 705 m, and 1130 m show increased signal noise under traffic conditions, as shown in Appendix A2.

In Figure 2d, the explosion signal that triggered the avalanche shown in Figure 2a is displayed. It is visible throughout the whole array, with smaller amplitudes in the gallery section. The maximum duration is approximately 1-second. The first arrival appears as a slightly curved line with a maximum arrival time difference of 0.08 s between first and latest arrival for the whole array (not visible in the Figure due to the long time scale of 40 s). This is likely due to the varying distance between source and receivers as a simple direct wave modeling example shows ?? a similar behaviour. one second.

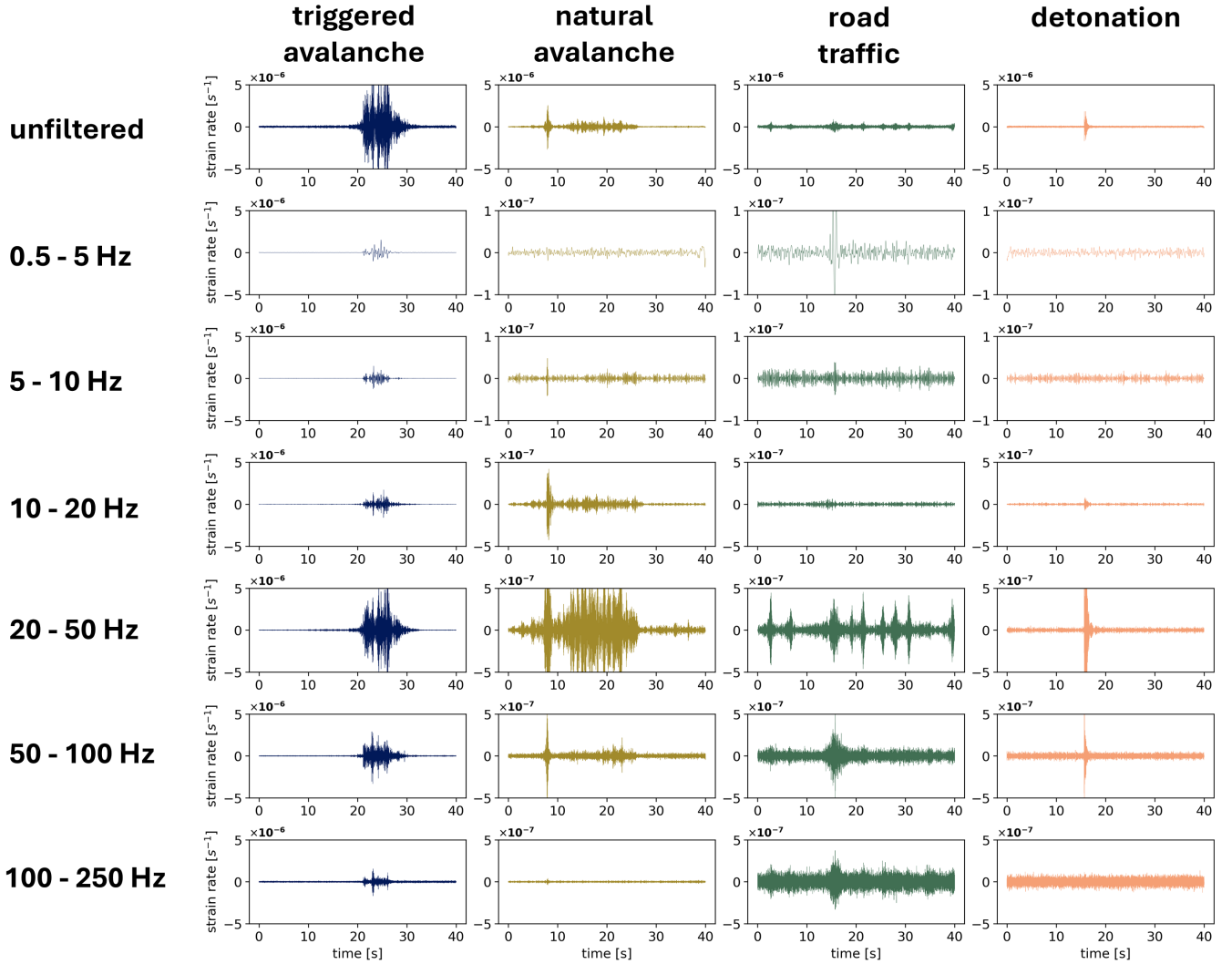
To analyze the frequency content and distribution of the events, the seismic data were filtered into different We filtered the seismic data into several frequency bands and compared using data then compared the signals from selected channels. For the artificial avalanche events, a channel was selected in the most relevant section of the array adjacent For the artificially triggered avalanches, we chose the channel from the array section next to the cones , where the avalanche signal was strongest. This same channel was used to compare (where the amplitudes were highest). We used the same channel to analyze traffic and explosion data for consistency. Since For the natural avalanche occurred , the highest amplitudes were observed closer to the bridge; therefore, a different channel was chosen where the signal for that event was more prominent. located near the bridge was chosen for analysis.

In Figures 3, the four different events from Figure 2 are compared, showing one channel each. Figure 3a represents an artificially triggered avalanche, Figure 3b a natural avalanche, Figure 3c road traffic, and Figure 3d an explosion used for avalanche triggering. Note that the y-axis scaling varies in Figure 3 to improve visibility and accommodate small values. This process was repeated for a second round of the same type of events, as shown in Figure ??.

For both the triggered and natural avalanches, most of the energy is found between 20 and 50 Hz. In Figure 3 the The triggered avalanche signal is visible across all frequency bands, with the strongest highest amplitudes between 20 and 50 Hz. The This frequency band was selected for later detection studies. We suspect that the broad frequency response can likely be attributed to the size and extent of the event, as it impacted the cones and covered the road, including the area directly over the cable. In contrast, the second triggered avalanche in Figure ?? is mainly visible in the 20-50 Hz band and does not show large amplitudes at lower frequencies. The signals of the two natural avalanches differ significantly from each other. The avalanche from February stems from the impulsive impact of the snow mass hitting the cones.

The natural avalanche from 2 , February 2024 , exhibits much larger amplitudes and shows a strong spike (1) 17:21 (UTC) shows a spike at the beginning, likely due to the avalanche hitting the cones. The second one from May 13, 2022, did not hit the cones and is only visible in the 20-50 Hz band as a slight swell between 15 and 25 s and a short spike event at 31 s. The signals from

Signals from road traffic are more pronounced in Figure ?? . In particular also prominently present in the 20 - 50 Hz range (Figure 3c). Notably, the signal at 50 seconds is strong and 16 s is visible across all frequency bands, stronger than the other ear signals. Individual car peaks are easily identifiable, especially in the 20-50 20 - 50 Hz band. The detonation signal is visible in the 20-100 Hz range in both Figures 3 and ??. Since the snow avalanche signal are strongest in the 20-50 Hz band and especially weak events like the natural snow avalanche displayed in Figure ?? are best detectable in that range , this frequency band was selected for later detection studies 20 - 100 Hz range.



Frequency comparison of single channel (that are indicated with arrows in Figure 2) data for different event types in Winter:
a) Triggered snow avalanche from-

Figure 3. Frequency comparison of single channel (that are indicated with arrows in Figure 2) data for different event types in Winter: a) Triggered snow avalanche from 10 February 2022 09:01 (UTC) b) natural snow avalanche from 2 February 2024 17:21 (UTC) c) car signals from 10 February 2022 11:07 (UTC) d) explosion signal from 10 February 2022 09:00 (UTC). Note that the y-axis scaling varies.

A comparison of the observed event types in the F-X domain is shown in Figure 4. The vertical stripe patterns likely result from variable coupling between the cable and the ground along the array. Similarities between triggered avalanches are evident in both their spatial extent and frequency content. Also, the triggered avalanche from 10 /02/2022 b) February 2022 and the natural snow-avalanche from 02 /02/February 2024 e) car signals from 10/02/ exhibit comparable spectral activity in the range between 850 m and 900 m. Traffic noise is particularly pronounced near the bridge (marked with dotted lines) and the cattle

195 grids (indicated by arrowheads). These locations exhibit high spectral intensities between approximately 20 and 220 Hz. The increased visibility of noise from the truck on 16 April 2022 ~~d) explosion signal from 10/02/2022. Note that y-axis ranges differ from subplot to subplot~~ is likely due to seasonal change: by this later date, the road was dry and the cattle grids were no longer filled with snow, causing increased noise generation when vehicles are crossing the exposed steel rods.

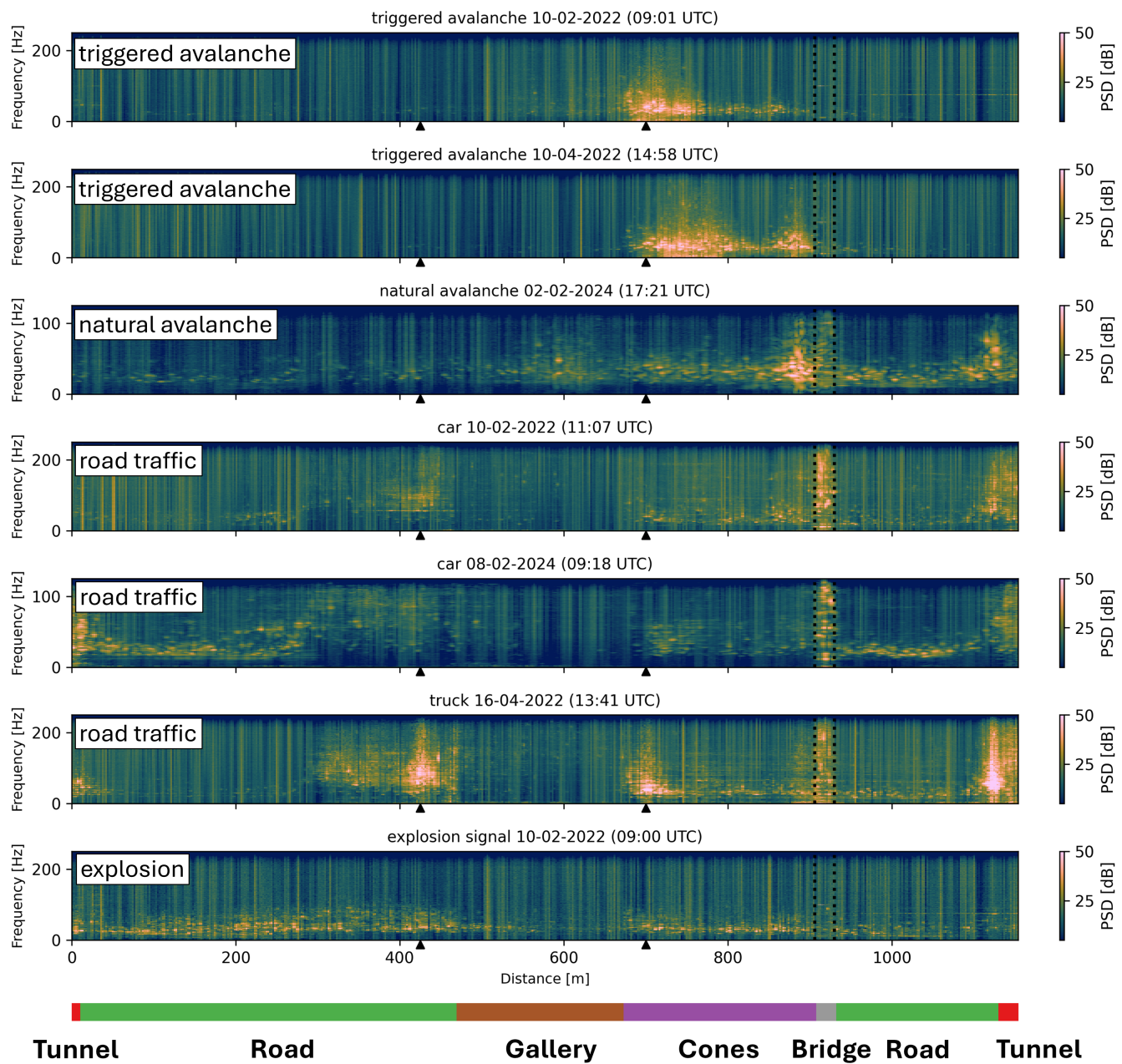


Figure 4. Event comparison in the F-X domain. The arrow heads mark the positions of the cattle grids and the dotted lines the location of the bridge. The maximum observable frequency for better visibility the 2024 dataset is lower (125 Hz) since the dataset of 2024 was recorded with 250 Hz instead of 500 Hz.

Frequency comparison of single channel data for different event types for spring season: a) Triggered snow avalanche from 10/04/2022 b) natural snow avalanche from 13/05/2022 c) car signals from 11/04/2022 d) explosion signal from 10/04/2022. Note that the y-axis range for the natural avalanche is reduced for better visibility.

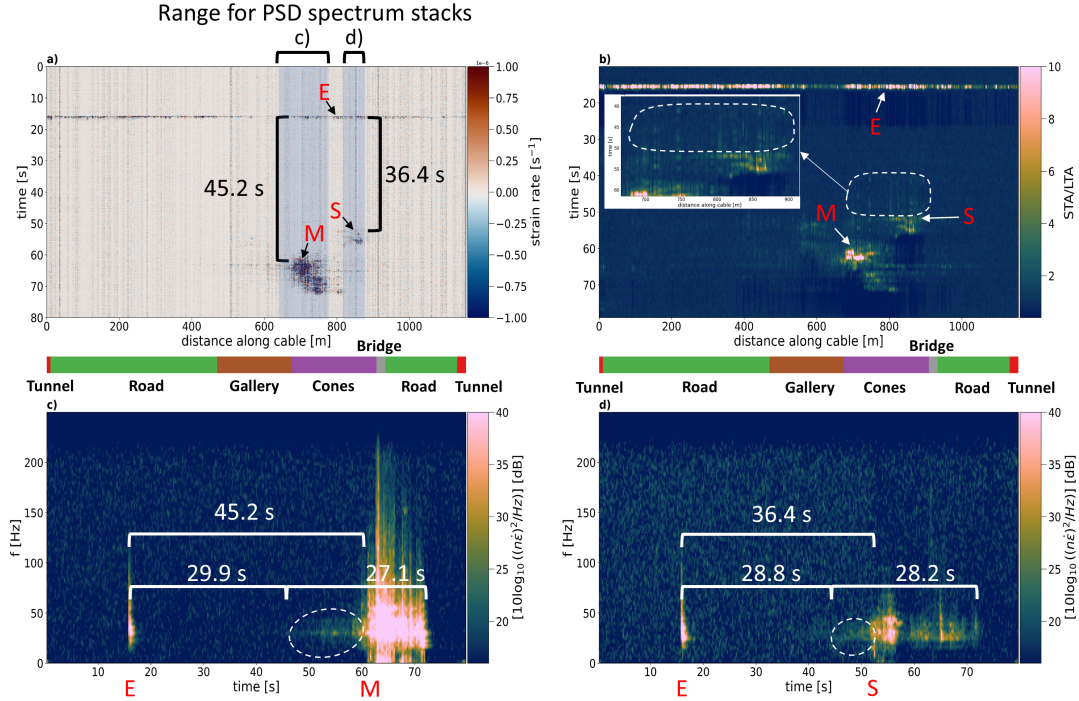


Figure 5. Temporal overview of the triggered avalanche from [10.02.2022 10 February 2022 09:01 \(UTC\)](#). a) Strain rate plot including explosion (E) and two avalanche signals (S for the small event and M for the main event). For reference, the time difference between the explosion and the two main signals is added (45.2 s for M, 36.4 s for S). Blue-shaded areas highlight the channels that are used to compute the stacked **PSD-spectra-spectrograms** shown in c) and d). A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. b) STA/LTA in [20-50-20-50](#) Hz band (short window: 1 s, long window: 10 s). Explosion (E) and two avalanche signals (S and M) are highlighted as in a). Precursor activity before the strong signal onsets is encircled in white. (c) and (d) Stacked **PSD-spectra-spectrograms** for the two channel ranges indicated in (a). The time intervals between the explosion and the start of precursor activity, as well as between the explosion and the strong signal onsets (onsets are labeled with M and S, respectively), are shown with white brackets, along with the corresponding time differences. Precursor activity occurring before the strong signal onsets is encircled in white.

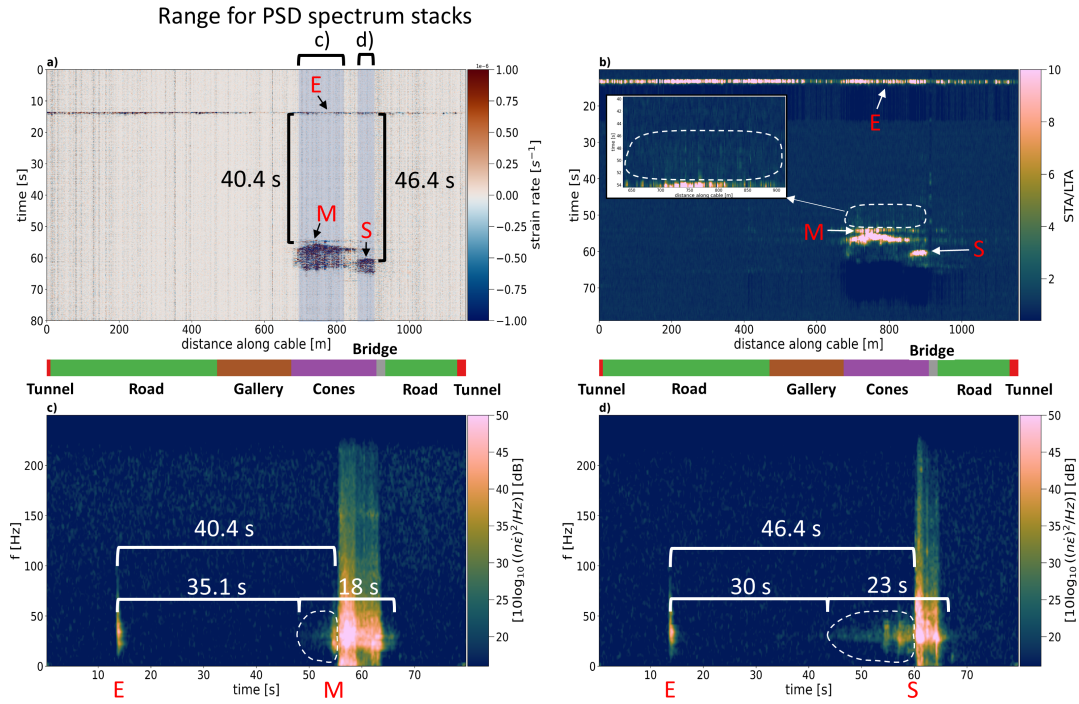


Figure 6. Temporal overview of the triggered avalanche from ~~10.04.2022~~ 10 April 2022 14:59 (UTC). a) Strain rate plot including explosion (E) and two avalanche signals (S for the small event and M for the main event). For reference, the time difference between the explosion and the two main signals is added (40.4 s for M, 46.4 s for S). Blue-shaded areas highlight the channels that are used to compute the stacked **PSD** spectrum plots shown in c) and d). A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. b) STA/LTA in ~~20-50-20-50~~ 20-50-20-50 Hz band (short window: 1 s, long window: 10 s). Explosion (E) and two avalanche signals (S and M) are highlighted as in a). Precursor activity before the strong signal onsets is encircled in white. (c) and (d) Stacked **PSD spectra-spectrograms** for the two channel ranges indicated in (a). The time intervals between the explosion and the start of precursor activity, as well as between the explosion and the strong signal onsets (onsets are labeled with M and S, respectively), are shown with white brackets, along with the corresponding time differences. Precursor activity occurring before the strong signal onsets is encircled in white.

In this section, we present a more detailed analysis of the two largest triggered avalanches (both reached the road), focusing on their time evolution and signal characteristics. ~~The analysis involves the computation of the Short-Time Average over Long-Time Average (STA/LTA) and Power Spectral Density (PSD). The STA/LTA was calculated using a frequency band of~~ The data is bandpass-filtered (20 - 50 Hz where the snow avalanche signal is the most prominent. We used a short window length-), after which the STA/LTA is computed using window lengths of 1 second ~~and a long-window length of (STA) and 10~~ length-), after which the STA/LTA is computed using window lengths of 1 second ~~and a long-window length of (STA) and 10~~ length-), after which the STA/LTA is computed using window lengths of 1 second and a long-window length of (STA) and 10 seconds (LTA). In this case, the window lengths for the STA and LTA were chosen to effectively capture the onset of avalanche signals, including short, ~~"spiky"-impulsive~~ "spiky"-impulsive events, such as those encircled in blue in Figure 2b. ~~The power spectral density (PSD)-~~

The spectra were calculated using a windowed Fast Fourier transform with Hanning windows of 0.5 s duration and 50% overlap. The PSD is-
 The spectrogram is calculated for each channel individually, then stacked within the regions of avalanche activity and normalized by the number of stacked traces for comparison purposes. Figure 5 illustrates the time evolution of the avalanche triggered on February-10, 2022. February 2022 09:01 (UTC).
 The strain rate plot (Figure 5a) shows the initial explosion at 16 seconds, followed by a small event occurring 36 seconds later, and the main event appearing 45 seconds after the explosion. While the strain-rate plot does not indicate significant activity before these events, minor disturbances are observed in the STA/LTA plot between 40 and 50 seconds (circled in white). In the stacked PSD-spectra-spectrograms (Figures 5c and 5d, a weak signal between 20 Hz and 35 Hz emerges at 45.9 seconds for the first stacking window and at 44.8 seconds in the second. The onset of strong signals (M and S in Figure 5a) corresponds to a sudden-increase-in-power-and-a-broadening of frequency content in the PSD-spectra-stacked spectrograms. For both stacking windows the activity stops at 73 seconds, yielding total event durations of 27.1 seconds and 28.2 seconds, respectively. This aligns well with the infrasound measurement of the avalanche, reporting a time of 28 seconds (RegObs Avalanche Observation, 10.02.2022). Given the size and impact of this avalanche, the infrasound measurement is considered reliable, consistent with previous findings on the effectiveness of infrasound in detecting large avalanches (Mayer et al., 2020).
 The second analyzed event is a triggered avalanche from April-10, April 2022 14:59 (UTC), shown in Figure 6. The strain rate plot in Figure 6a reveals two distinct events: the main event (M) and a secondary, smaller event (S). Both events appear within a similar range of channels as those observed on February-10, 2022. February 2022 09:01 (UTC). This consistency suggests that the same two avalanche paths were active on both days. Notably, the time interval between the explosion and the onset of the strong signal is longer this time, measuring 40.4 seconds for the main event and 46.4 seconds for the smaller one. Due to the lack of additional data regarding the event, the reason for this extended duration remains unknown.
 Although a clear precursor signal is not visible in Figure 6a, there is a slight indication of activity in the STA/LTA plot in Figure 6b. In the stacked PSD-spectra-spectrograms, slight activity starts at 48.7 s for the main event (Figure 6c) and 43.6 s for the small event (Figure 6d).
 At the onset of the strong signals in the strain rate plot, we observe a corresponding-increase-in-power-in-the-PSD-spectra, along-with-a-broadening of the frequency content. This likely corresponds to seismic noise-waves stemming from the snow mass hitting the cones. For both events, the activity in the PSD-spectrum-spectrograms resumes at 66.6 s. This leads to total event durations of 18 seconds and 23 seconds, respectively, which are significantly longer than the 6-second duration reported from the infrasound data.
 The signal evolution for-of the natural avalanche from-February-recorded on 2, February 2024, at 17:21 (UTC) is displayed in the-appendix-A3-Appendix A3. The event duration is 25 s, with a 6 s precursor signal is visible in the stacked spectrogram preceding the first impulsive event (marked with '1'). Two distinct high-amplitude impulses are observed, possibly corresponding to major snow impacts, followed by lower-amplitude signals that may indicate subsequent smaller slides.

4 Discussion and Outlook

We recorded seismic signals from snow avalanches using Distributed Acoustic Sensing (DAS) along a mountain road. The avalanche signals varied significantly, depending primarily on the size of the avalanche and whether it reached the cones or the road. The highest amplitude signals, present over the full recordable frequency range between 0 and 250 Hz, were observed on fiber stretches adjacent to the cones and the steep terrain incline leading up to the road. We believe that this could result from the snow mass colliding with the mentioned features, producing high amplitude surface waves. We also detected pre-impact signals using the Power Spectral Density Spectrum spectrogram. Those lie in the frequency range from 20 - 50 Hz, likely caused by the approaching avalanche (Figures 5 and 6).

For the natural avalanche on February 2, February 2024 17:21 (UTC), several short signals were observed at different parts of the cable as described in Figure 2 (encircled in blue). Since three of those events are detected in the section of the avalanche gallery with no obvious obstacles present, we associate them with several small parts of the slide coming to an abrupt stop. This would fit to the findings of Suriñach et al. (2000) who reported short duration signals at the end of the avalanche signal trains, commonly referred to as the stopping phase. They recorded artificially triggered snow avalanches using 3 - C geophones 1 - 3 km away from the avalanche path. Similar signals were detected and analyzed in more detail by Suriñach et al. (2020), using geophones placed along the avalanche path. In one case of their recorded events, a wet snow avalanche stopped before reaching the geophone, presenting a scenario similar to the one observed in this study. The sudden stop of the snow mass produced a spike in the seismic signal over the full frequency range of 1 - 40 Hz. Based on these findings, we are confident that think the same explanation is also could also be valid for the aforementioned signals observed by our DAS system.

We attempted to apply frequency-wavenumber (fk) filters to remove any noise, including traffic, approaching the fiber cable at relatively low apparent velocities. However, as seen on e.g. Figure 2c and better on Figure A2, the traffic signal is incoherent from trace to trace. Averaging over several traces though helps to make the signal coherent, however, that also applies to the avalanche signal and subsequent fk-filters will also remove a significant part of the avalanche signal while only reducing the coherent part of the traffic signal, still leaving much traffic noise in the data. Hence, fk-filters applied to this case study does not solve the problem, but more advanced transformation of the noise should further be tried out in order to detect smaller slides where the thresholds for triggering would need to be lowered without too many false detections.

In the two triggered events, two spatially distinct signals suggest the presence of two slightly different paths for the avalanches. Although this is not verified yet, we aim to investigate this further using avalanche simulations. Although we different avalanche paths approximately 80 m apart. Although we detect small precursor signals for two large avalanches, they are weak and difficult to identify, making real-time detection challenging. Additional instrumentation, such as extending the fiber up the mountainside, could improve this and also be used for studying avalanche dynamics in more detail, as shown by Paitz et al. (2023).

Automated real-time avalanche detection remains challenging due to the variability of avalanche signals as well as the presence of traffic noise. As an initial attempt, we tested STA/LTA threshold triggering in the 20-50-20 - 50 Hz frequency band with threshold values between 5 and 10, focusing on the open road section next to the cones. While both triggered avalanches are

picked up by this method, smaller, less characteristic events such as the side ~~event~~ events displayed in Figure A3a (labeled with 3) are not detected. As mentioned, road traffic noise along the fiber path adds to the problem, since especially large trucks exhibit strong signals present over the full recorded frequency range. This effect is amplified at the cattle grids and the bridge where the noise spreads over multiple channels as displayed in Figure A2a. ~~Consequently, this~~ This leads to numerous
280 false positives ~~along the vehicle signal track~~, which could not be filtered out as of now. ~~Detecting smaller slides would require lowering the threshold further, making the system even more susceptible to false positives from traffic noise. In conclusion, traffic noise needs to be identified and excluded through alternative methods to improve detection accuracy. One potential solution is a two-step approach. First, a conventional STA/LTA threshold trigger would be applied continuously, and in the case of an event detection, a data snippet surrounding the event would be analyzed again. The data snippet (e.g., a 10-second~~
285 ~~recording) could be treated as an image and analyzed using a Convolutional Neural Network (CNN) to distinguish traffic signals from all other events. A CNN approach could be promising given that we have a large amount of training data available for traffic.~~

5 Conclusion

~~We have made valuable observations of snow avalanches using Distributed Acoustic Sensing (DAS) along a mountain road. Detecting snow avalanches is critical, and DAS presents a promising method for using existing infrastructure to address this challenge. Our results demonstrate the successful detection of both triggered and natural avalanches. We have characterized the observed signals, including those from road traffic and explosions, noting that avalanche signals were most prominent in the 20-50 Hz~~

In a recent study by Kang et al. (2024) it was demonstrated that utilizing the phase information of Cross-Spectral Density
295 Matrices can be used for detection of rock-slope failures, leveraging the presence of event-specific coherence patterns. This was achieved using a semi-supervised classification approach. This method could also be promising for snow avalanche detection. However, an implementation on this dataset might be challenging due to the limited number of confirmed avalanche events available as training data. Additionally, while the authors used frequencies within the 2 - 8 Hz range, our results show that in this dataset avalanche signals are most prominent at higher frequencies (20 - 50 Hz). It is uncertain whether the method
300 would maintain its effectiveness in this frequency band. ~~Weak precursor signals were identified for larger events, though these are insufficient for reliable automatic detection. We believe that the large amplitudes are generated when the snow mass collides with the snow-stopping cones next to the road and the slope leading up to the road next to the northeastern tunnel. For natural avalanches, we detected short, spike-like signals across several channels, which we attribute to the snow mass coming to a stop. However, road traffic, particularly large trucks, generates high-amplitude signals across the entire frequency~~
305 ~~spectrum, complicating automatic detection efforts. Further work is necessary to enhance detection capabilities, with the goal of ultimately achieving reliable automatic detection.~~

One other potential solution shown by Truong et al. could involve using edge detection techniques to identify the typical diagonal lines that correspond to vehicles followed by applying the Hough transform (Duda and Hart, 1972) to extract these

features. Given the large amount of available training data for traffic signals in this study, this approach could be further
310 enhanced by integrating a Convolutional Neural Network (CNN) to improve line detection accuracy, as demonstrated by
Xie et al. (2025). The process would start with applying a conventional STA/LTA threshold trigger to a segment of data. If
an event is detected, the resulting trigger matrix (a binary image containing only the trigger values) would be analyzed for
diagonal (slanted) features, for example, using the Sobel operator (Kanopoulos et al., 1988). The Sobel-detected edges would
then be processed with the Hough transform for line extraction. If no lines are detected, the event could be flagged as a potential
315 avalanche.

| | <u>2022</u> | <u>2023</u> | <u>2024</u> |
|----------------------------|-------------|-------------|-------------|
| <u>channel spacing [m]</u> | <u>2</u> | <u>2</u> | <u>2</u> |
| <u>sampling rate [Hz]</u> | <u>500</u> | <u>500</u> | <u>250</u> |
| <u>Gauge length [m]</u> | <u>3.1</u> | <u>5.1</u> | <u>5.1</u> |

Table 1. Acquisition parameters

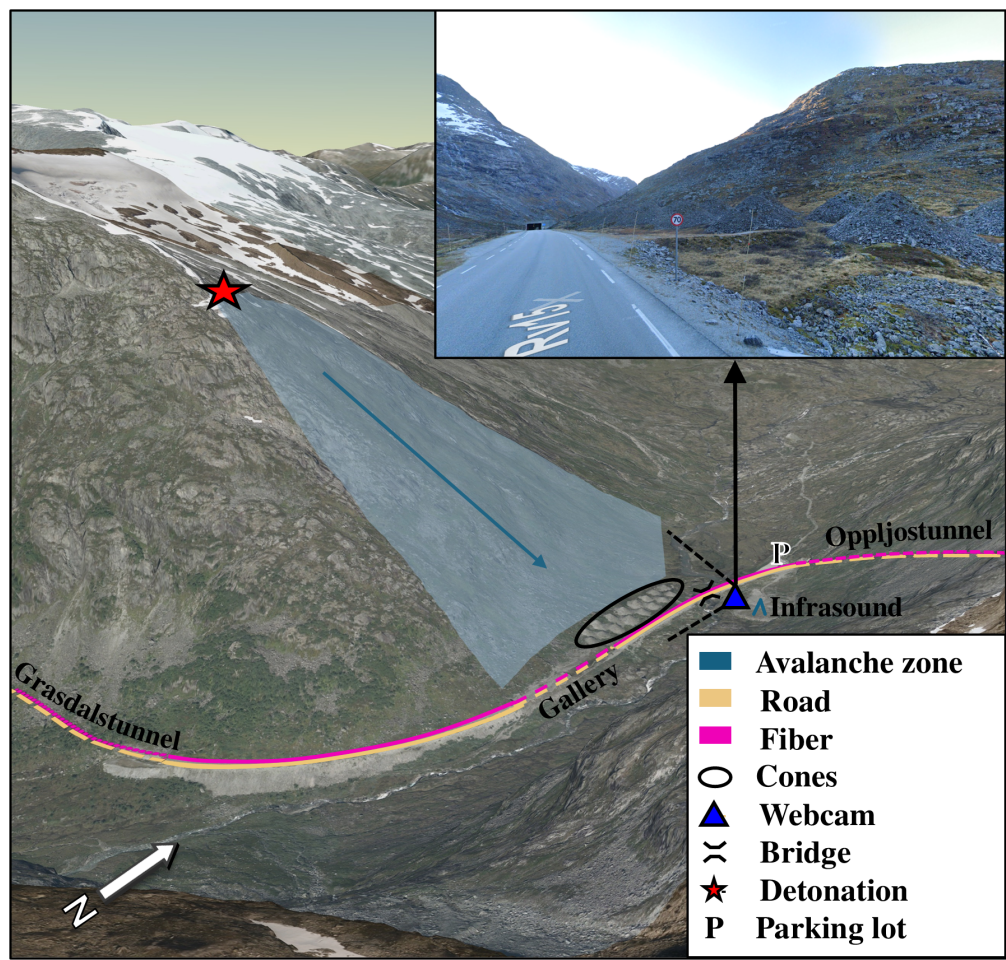


Figure A1. 3D view of the Grasdalen valley and the survey setup. The inset image shows the east face of Saetreskarsfjellet, the cones and the road as seen from the bridge. Aerial footage by Norwegian Mapping Authority (2020), inset image by Google (2024).

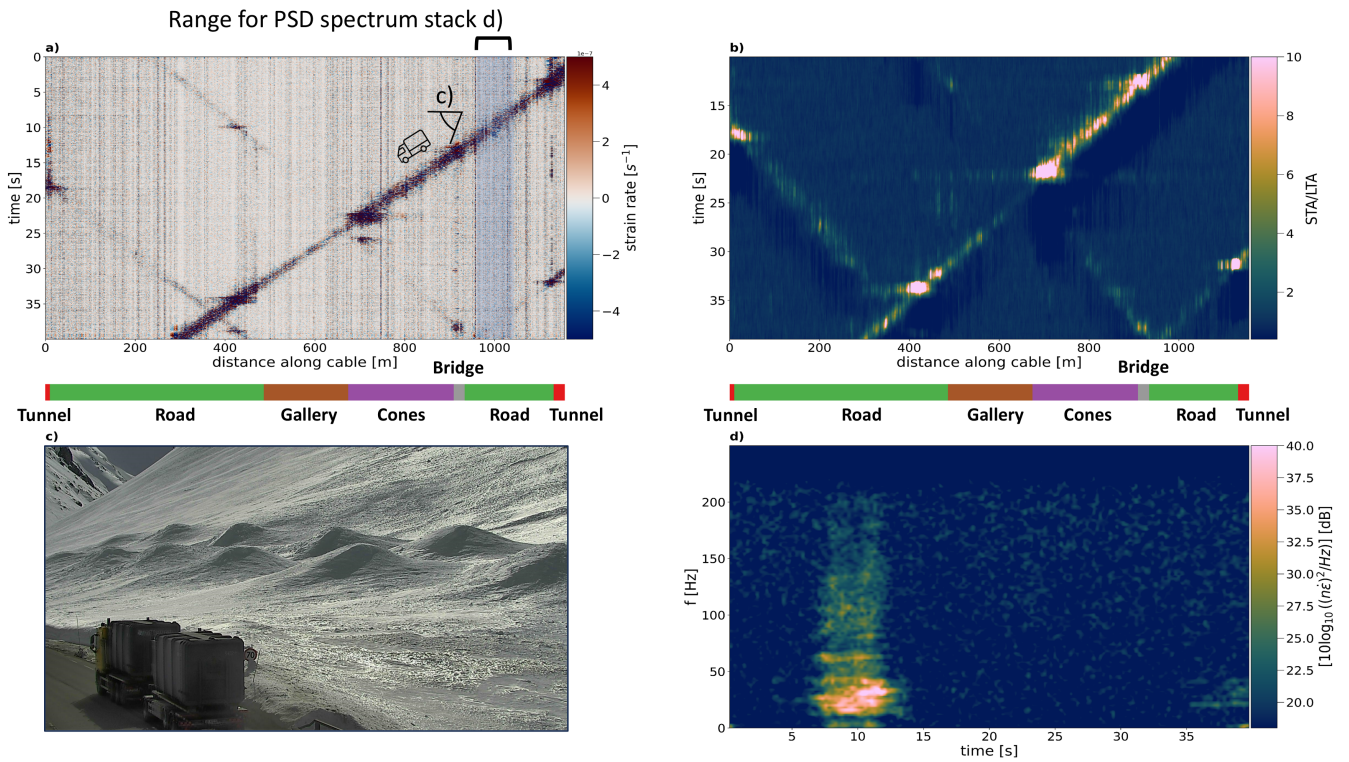


Figure A2. Road traffic signals from ~~16.04.2022~~ 16 April 2022 13:41 (UTC). The strongest signal belongs to a two-axle truck passing the array from northeast to southwest. a) Strain rate plot with signals of ~~6-4~~ individual vehicles. The strongest signal train is caused by a truck, passing the road from northeast to southwest. This was verified by the webcam footage shown in c). Webcam and truck position from the webcam image are marked. Note the noisy signals created by the vehicles passing over the kettle grids at 10 m, 430 m, 705 m and 1130 m as well as the bridge. The blue-shaded area highlights the channels that are used for the stacked PSD spectrum plot shown in d). A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. b) STA/LTA in ~~20-50-20~~ 20-50-20 Hz band (short window: 1 s, long window: 10 s). c) Webcam Image showing the truck next to the cones. Webcam and truck position are ~~highlighted sketched~~ highlighted sketched in a). d) Stacked PSD Spectra for the two range of channels.

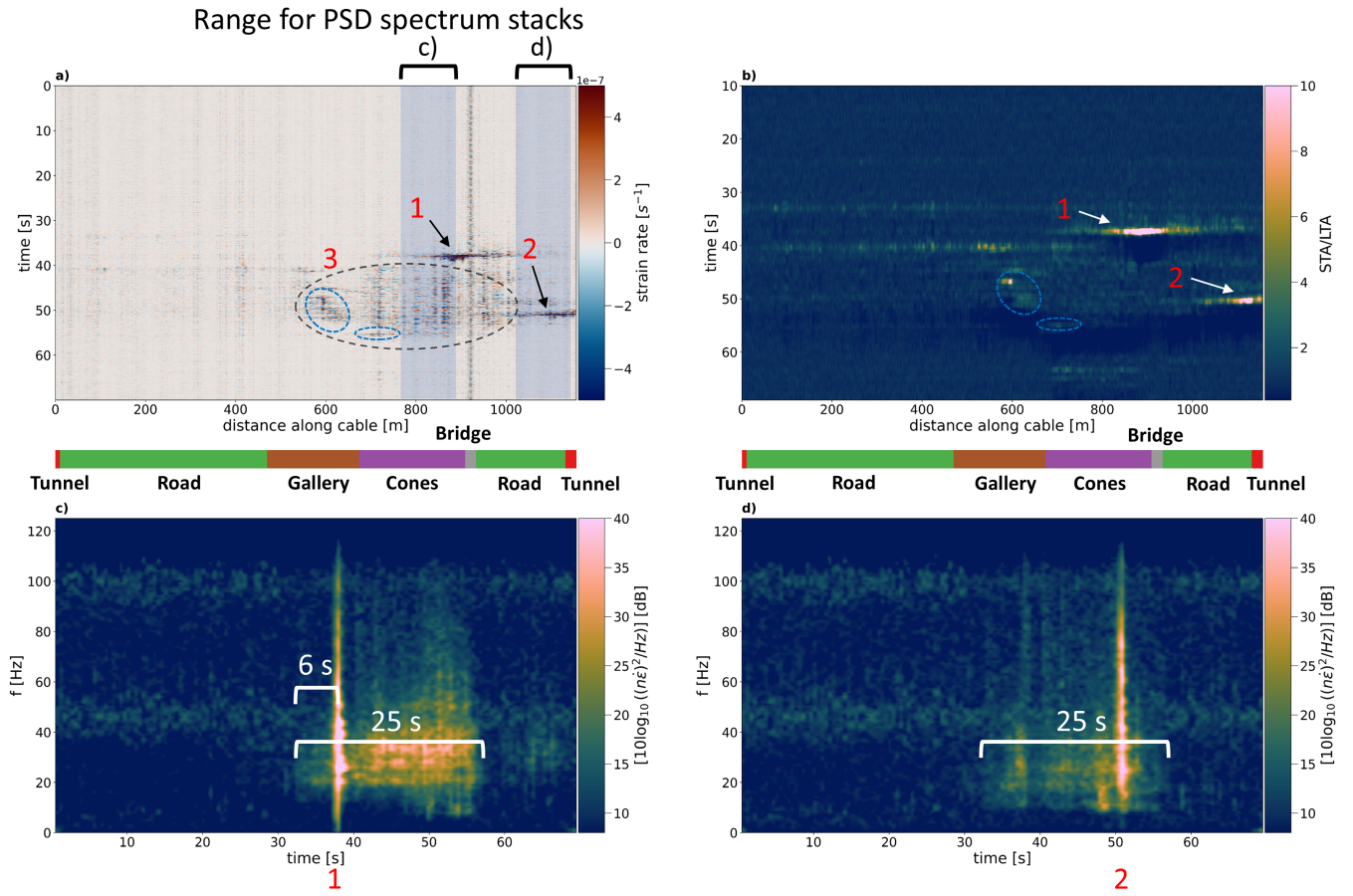


Figure A3. Temporal overview of the natural avalanche from [02.02.2024](#) [February 2024 17:21 \(UTC\)](#). a) Strain rate plot including two strong avalanche signals (labeled 1 and 2) and several less strong avalanche signals (labeled with 3). Four noticeable short events within this range are encircled in blue. Blue-shaded areas highlight the channels that are used to compute the stacked PSD spectrum plots shown in c) and d). A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. b) STA/LTA in [20-50-20 - 50](#) Hz band (short window: 1 s, long window: 10 s). The two strong avalanche signals (1 and 2) are highlighted as in a). The short events mentioned in a) are encircled in blue. c) and d) Stacked PSD spectra for the two channel ranges indicated in (a). The total event time is indicated with white brackets. Precursor activity occurring before the strong signal onsets is encircled in white.

~~First Arrival of Explosion signal from 10.02.2022—comparison of recorded and modeled signal. Recorded strain rate plot is displayed in grayscale. A highpass filter (third order Butterworth) with a cutoff frequency of 0.5 Hz was used to suppress low frequency channel noise. The modelled first arrival (direct travel path, constant velocity of $2500 \frac{m}{s}$, topography neglected) is overlaid.~~

320 *Author contributions.* AW and VO initiated the research; AW gathered the data; FK analysed the data, plotted graphs and maps and wrote the manuscript; AW and CB reviewed graphs and assisted with data analysis; VO and ML reviewed and edited the manuscript.

Data availability. At the moment, the DAS data of this study not publicly available. Access can be granted on individual request.

Competing interests. The authors declare that they have no conflict of interest.

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