



Prediction of landslide induced debris' severity using machine learning algorithms: a case of South Korea

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29 Abstract

Rainfall-induced landslides frequently occur in the mountainous region of Korean peninsula. 30 The resulting landslide induced debris cause extreme property damages, huge financial losses 31 32 and human deaths. To mitigate their effect different landslide susceptibility mapping are 33 frequently used. However, these methods do identify regions with potential landslides but they do not quantify their severity. In this paper multi-category ordered machine models, namely, 34 35 proportional odd logistic regression (POLR), random forest (RF), support vector machine (SVM), and extreme gradient boosting (EGB) methods, are proposed to fill the specified gap. 36 37 Moreover, the exploratory data analysis on landslide induced derris's dataset has been 38 conducted on to examine patterns and relationship between landslide-induced debris 39 severity(size), causal factors(rainfall) and influencing factors. Findings revealed that cumulative three days' rainfall and slope length were most responsible for the severity of 40 landslide-originated debris severity and slopes between 20° to 40° was identified as most 41 vulnerable region. Furthermore, the predictive accuracy statistics were compared to assess the 42 suitable model for debris severity for Korean case. The RF and EGB ranked higher with an 43 overall accuracy of 90.07% and 86.09% and kappa of 0.72 and 0.61 on the validation set, 44 respectively. The findings of this research may be useful in the identification of high risk zones 45 46 for extreme rainfall-induced debris to elaborate mitigation and resilience policies, post-disaster 47 rehabilitation planning and land use management.

48

Keywords: Rainfall-induced debris Severity; Proportional odd logistic regression; Random
forest; Support vector machine; Extreme gradient boosting; Machine learning, South Korea





51 1. Introduction

52 Rainfall-induced debris is a natural phenomenon that occurs when the slope fails due to the saturation of soil after the rainfall exceeds a certain threshold (Au 1998; Takara et al. 2010; 53 Peruccacci et al. 2017; Segoni et al. 2018; Crawford et al. 2019; Coppola et al. 2022). 54 55 Rainwater penetrates the soil through cracks or pores (Zeng et al. 2022) which destabilizes the 56 slope and induces landslides (Franzluebbers 2002). Furthermore, the volume of landslide-57 induced debris depends on the geological condition of the terrain, rainfall intensity and duration (Chang et al. 2011). Extreme rainfall is the triggering factor for landslides which is one of the 58 most damaging natural disasters with the expensive cost of repair and indemnification 59 (Kockelman 1986; Gariano and Guzzetti 2016). In addition, heavy windstorms, typhoons, and 60 extensive rainfall have destroyed many properties and taken many human lives yearly (Liu et 61 al. 2018). Furthermore, landslides have caused enormous environmental degradation, 62 infrastructure damage, casualties, and loss of life, which disturb the socio-economic aspect of 63 the community (Li et al. 2012; Sarkar and Dorji 2019; Zhao et al. 2019; Taylor et al. 2020; 64 Lacroix et al. 2020; Winter 2020; Negi et al. 2020; Ju et al. 2020; Van et al. 2021). Most rainfall-65 induced landslides were found to be shallow (de Jesús et al. 2019; Liu et al. 2021; Chang et al. 66 2021) however, some were very extreme and resulted in severe human and financial damage 67 (Turner, 2018; Meena et al., 2021). Klose et al. (2016) found that from 1980 to 2013, landslides 68 took thousands of lives and an annual average of about \$20 bilion of economic losses, which 69 was 17% of the total (\$121 billion) annual mean of global disaster-induced losses. 70

The Korean peninsula is characterized by mountainous, which makes it more prone to rainfall-induced landslides (Lee et al. 2013). Lee et al. (2012) found that the triggering factor for landslides was short-duration heavy rainfall. Park et al. (2013) reported that the annual property damage caused by rainfall-induced landslides in South Korea averaged between US\$500M to US\$1000M and approximately 36 human deaths per year from 1997 to 2010.





76 Therefore, to mitigate the effect of landslides in South Korea, different studies have been 77 carried out on landslide assessment (Kadavi et al. 2019; Lee and Winter 2019; Sameen et al. 78 2020; Panahi et al. 2020; Hakim et al. 2022). Lee et al. (2020) applied the Naïve and Bayesian 79 Networks model for landslide susceptibility mapping in Umyeon Mountain. Lee et al. (2012) 80 used physical slope and probabilistic model, i.e., decision trees and logistic regression for landslide susceptibility mapping in Gangwon-do. Lee et al. (2013) developed the binary 81 logistic regression model for predicting the occurrence of landslides. Woo et al. (2014) 82 constructed a landslide hazard map using binary logistic regression. Park and Kim (2019) 83 84 compared boosted trees and random forest model's performance in landslide susceptibility mapping for Umyeon Mountain; the same methods were previously applied at Pyeong-Chang 85 by Kim et al. (2018). It was observed that the objectives of previous studies were to predict 86 87 landslide susceptibility; they did not specify how severe the occurring landslides would be. 88 Further, most of studies were performed on a small scale and only predicted the occurrence, not the size. Therefore, in the present study, we analyzed landslide-induced debris severity 89 90 based on the causative variables and influencing factors. This study is novel in expressing the relationship of debris' severity, causative and influencing factors. It is an extention on landslide 91 92 mapping which quantifies the magnitude of landslide-induced debris. The quantification of debris severity may be useful in land management by highlighting regions prone to higher 93 94 rainfall-induced debris to know whether economic activities that may be carried out on the given region may not be vulnerable to extreme landslide hazards. The severity of debris is 95 measured in unit of volume (m³) and classified as shallow(below 500m³) small(500-2000m³), 96 medium (2000-5000m³) and critical, i.e., above 5000m³. Words severity of debris, debris 97 98 volume or size of debris express the same quantity in different ways and are used 99 interchangeably in this manuscript.





100 2. Study region

101	Korean peninsula is located in the northern hemisphere, between China and Japan in Northeast
102	Asia. Its climate has continental and oceanic features with wide temperature differences. The
103	yearly mean temperature ranges from 10°C to 16°C, that is, from -6°C to 7°C in winter and
104	23°C to 27°C in summer. In South Korea, the rainy season range from June to September, with
105	1000mm to 1800mm of precipitation in the southern part and 1100 to 1400mm in the central
106	region (https://web.kma.go.kr/).
107	The altitude ranges from 0 to 1911 meters, with mount Halla (in Jeju Island) being the
108	highest peak in South Korea. The Gangwon Province is the most mountainous region of about
109	64% of all tallest mountains in Korea, that is, 23 of 36 mountains. The surface geology of the

Korean peninsula is mostly composed of igneous, sedimentary, and metamorphic rock (Chough et al., 2000). The arable soil depth varies between 1 to 2m (Lee and Winter, 2019). Due to the high intensity rainfall and weak geological formation in the mountainous region causes high frequency of landslides. Figure 1(a) illustrates the distribution of landslides by the volume of landslides, while subplots (b & c) exhibited the relationship of slope length, altitude, and rainfall with the landslide size.





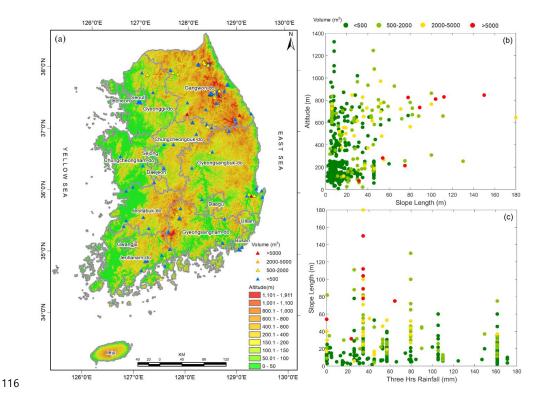


Figure 1. (a) Landslide location in South Korea, (b) Distribution of landslide volume per
slope length and altitude, and (c) Distribution of landslide volume per slope length
and rainfall (Data source: elevation data acquired from NGII, 2018).

120 3. Methodology

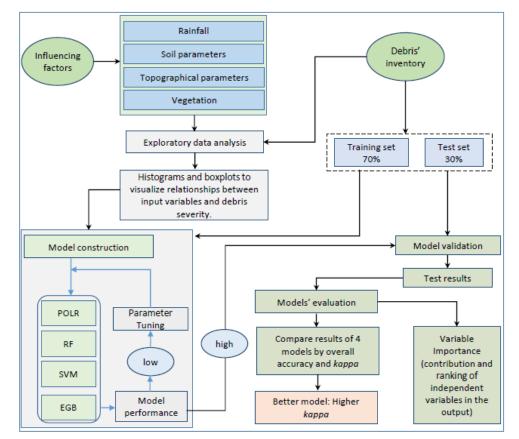
121 3.1. Problem formulation

Predictive models that deal with multi-variate random variables were investigated to predict the severity of rainfall-induced landslides. Among those predictive models, proportional odd logistic regression and other machine learning-based algorithms such as extreme gradient boosting, random forest, and support vector machine are widely used to deal with classification problems (Marjanović et al. 2011; Lee et al. 2012; Woo et al. 2014; Wang et al. 2022a). The main steps for the modeling process were to analyze the interaction of variables





- 128 that influence the severity of landslides with results in a higher size of debris. For the purposes,
- 129 four machine learning models i.e., POLR, RF, SVM, and EGB, were used to assess the most
- 130 suitable predictive model for landslide-induced debris in south Korea (;Su et al.,2022). The
- 131 comparison was made using predictive accuracy and the value of kappa. Figure 2 shows the
- 132 steps followed in the construction of the model.



133

134 Figure 2. Modelling workflow process for the prediction of landslide induced debris' severity

135 using machine learning algorithms.

136 3.2. Data description

137 The dataset contains 455 debris inventory collected from field surveys with the help of





portable GPS, a laser ranger and a clinometer. Variable definitions and descriptions for each 138 139 feature in the dataset are presented in Table 1. The rainfall data were collected from Korea 140 meteorological Administration stations scattered around the country nearest each landslide-141 induced debris site. It was revealed that the duration and quantity of rainfall directly affect 142 landslides (Berti et al. 2012; Kim et al. 2014; Sarkar et al. 2019; Liu et al. 2020; Ngo et al. 2021). Rainfall data were grouped into twelve variables: cumulative rainfall, continuous hourly 143 rainfall, three hours, six hours, nine hours, twelve hours, one day, three days, seven days, two 144 weeks, three weeks, and four weeks' rainfall. Different measures of rainfall were captured due 145 146 to the time-dependent cumulative effect of rainfall on the slope stability, and prolonged rainfall has a more damaging effect in mountainous regions (Baum and Godt 2010; Hidalgo et al. 2017; 147 148 Meena et al. 2022). The conditioning factors, i.e., soil type, topsoil depth, altitude, slope, slope-149 length, slope aspect, and vegetation (leafage, size of tree, age of trees, and fire history), were 150 collected. The soil type was classified as sandy loam, lithosols, silt loam, and clay. The soil depth was grouped into below 20 centimeters, between 20-50, and 50-100 centimeters. The 151 152 rainfall infiltrating the topsoil causes saturation, which initiates the landslide and then results in debris flow (Baum and Godt 2010; Vahedifard et al., 2017; Zhu and Zhang 2019). The anti-153 154 erosive drainage presence and status were categorized into: very good, good, and bad. Dranage channels reduces the concentration of water in soil and effect on water flow, saturation, soil 155 moisture, and valley landslides(Shahabi and Hashim 2015). The vegetation-covered and 156 necked lands have different affect of landslide (Lee et al. 2013; Ozioko and Igwe 2020; Huang 157 158 et al. 2021). The foliage information was classified into pines, broad-leaved, and mixture, while, 159 the size of trees was classified as large, small, and medium. The age of trees was grouped into 160 seven classes viz. 1-5 5-15, 15-25, 25-35, 35-45 and >45 years. The forest fire history was also 161 considered as a influencing factor due to its erosive nature(Huang et al. 2020). Geographical features were found to contribute to the severity of landslides at different levels; steep slopes 162





were found to fail as the intensity of the rainfall increased (Brand et al. 1984; Au 1998; Charles
and Shi 1998; Nandi and Shakoor 2008; Pham et al. 2018; Chen et al. 2020). It was also noted
that plane areas beneath steep slopes face the damage caused by debris flowing from the top of
mountains (Raja et al. 2017, Wang et al. 2022b).
The output variable (volume) of landslides has been classified into four categories, i.e.,
below 500m³, between 500-2000m³, 2000-5000m³, and above 5000m³(Fig. 4c). There are

different types of landslides; which are classified based on the cause or shape after occurrence (Causes 2001). Landslides may result from liquefaction, earthquakes, intense surface water flow due to precipitation, underground water, ice melting, human activities, tectonic movements etc.(Alexander 1992; Causes, 2001; van der Beek, 2021;McColl 2022). In this paper, we only considered landslide-induced debris originating from rainfall. Table 1 summarizes the characteristics of debris: its types, frequency and size.

Volume (m ³) Destroyed area	<500	500-2000	2000-5000	>5000	Total
valley erosion	1	1	1		3
falling rocks	1				1
mixed/ complexes	3	2	1		6
slope	1	1		1	3
scour	1				1
curved wedges	4	1			5
a circular arc	205	45	14	2	266
Plane	120	35	10	5	170
Total	336	85	26	8	455

176 Table 1: Landslide-induced debris types and frequency

177 *3.3 Exploratory data analysis*

The relationship between the influencing factors and the debris size were analyzed. We consider categorical variables, also known as qualitative set of information, that is divided into groups. It describes data which are non-numerical and serve qualitative purposes, such as



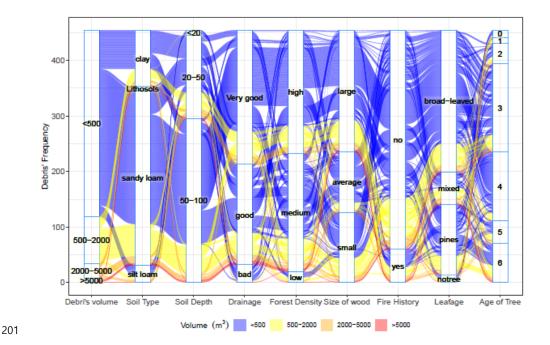


representing a certain value or distinctive character of objects (nominal), when there is no order 181 182 in levels and ordinal when there is an order in the dataset (Bilder and Loughin 2014). In addition, continuous features were presented in form of histograms and boxplots to analyze the 183 dispersion and influence on the size of debris. Figure 3 illustrates that most of the occurred 184 debris was shallow (below 500m³, 73.85%), followed by 500-2000m³ (18.68%), and the 185 extreme debris (above $5000m^3$) was the least frequent (1.76%). Debris with a size above 186 5000m³ was mainly associated with sandy loam soil of depth above 20cm and a non-perfected 187 drainage system, where the forest density was medium and in a place that experienced wildfire. 188 189 The region with pines leafage experienced shallower debris compared to other types of leafage. The region with older trees, above 45 years of age, experienced more severe debris than 190 younger forests. Šilhán and Stoffel (2015) highlighted that the area with timber age of above 191 192 45years were more sensitivity to landslide occurrences. The size of debris for wildfireexperienced regions was observable; compared to the frequency of cases with no wildfire 193 history, the severity was quite higher. The wildfire influence on the rainfall-induced debris is 194 195 due to the reduced infiltration of water into the soil, which increases the erodibility of soil (Ranger et al. 2020; Tiwari et al. 2020). The inadequate drainage system resulted in severe 196 debris(Popescu 2002), where the system was very good, no severe debris occurred. The 197 resulting huge number of shallow debris for a very good drainage system is due to the fact 198 199 those systems are usually created in the most vulnerable regions and the shallow debris is an indication of improvement in landslide mitigation. 200

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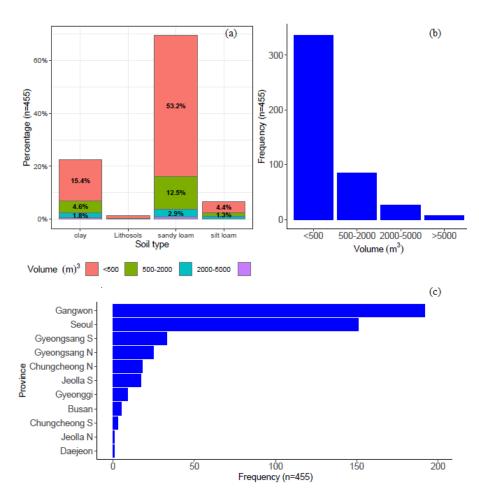


202 Figure 3. Interlinkage between the size of debris and categorical influencing factors.

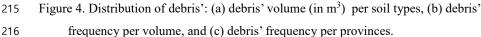
203 Figure 4a illustrates that sandy soil, which totalizes about 70% of all debris, was the most vulnerable soil, followed by clay (22%), silt loam (7%) and lithosols (1%). Similarly, 204 sandy soil not only ranked higher in terms of frequency but also in terms of size. The high 205 sensitivity of sandy soil to rainfall-induced debris may be due to its high coefficient of 206 permeability which facilitates fast saturation of topsoil during the rainfall period that induces 207 shallow debris (Lee et al. 2013). Overall, about 73.85% of all debris was shallow, that is, below 208 500m³, 18.68 % (500-2000m³) was small, 5.71% (2000-5000m³) was medium, and only 1.76% 209 were critical, i.e., above 5000m³, as depicted in Fig. 4b. Figure 4c, illustrates that about 74% 210 211 of landslide-induced debris occurred in Gangwon and Seoul; this made the two provinces more vulnerable than other regions. South and North Gyeongsang provinces ranked third and fourth 212 with 7% (25 cases) and 5% (18 cases) of all debris inventory, respectively. 213











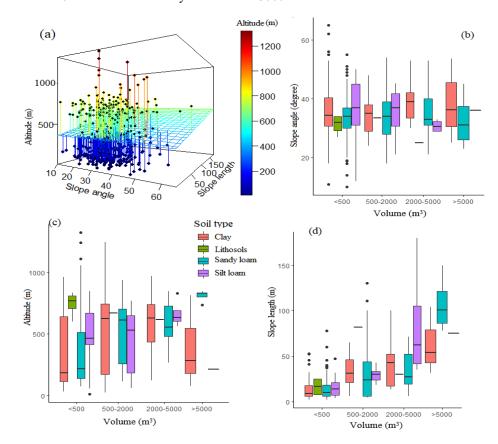
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To examine the relationship between continuous explanatory variables and their effect on the size of debris, 3D scatter plots and boxplots were presented (Fig. 5). It was observed on the 3D scatterplot that most debris occurred at slope between 20° and 40° (Fig. 5a), this was also confirmed by the boxplot (Fig. 5b), where few exceptions were observed for shallow debris for clay and sandy soil where debris occurred at small or at very large slope angles as outliers. Shallow debris was independent of the slope angle as depicted in Fig.5c; outliers were





scattered on all slope angles, and as the size of debris increased, the occurrence converged at slope between 20° and 40°. There was a pseudo-decreasing trend between altitude and size of debris (Fig. 5c), the highest altitudes (above 600m) were associated with shallow debris and critical debris occurred at altitude between 200 and 900m. On the other hand, a quasi-increasing relationship between size of debris and slope length was observed in all type of soil (Fig. 5d). Debris below 500m³ was associated with slope length below 80m, the highest quartile was about 140m associated with clay soil of above 5000m³ of debris size.



231

Figure 5. Distribution of debris and continuous predictors: (a) 3D Scatter plot between altitude, topographic slope and slope length, (b) boxplot of volume of debris per slope angle and soil types, (c) boxplot of volume of debris per altitude and soil types, and (d) boxplot of volume of debris per slope length and soil types.

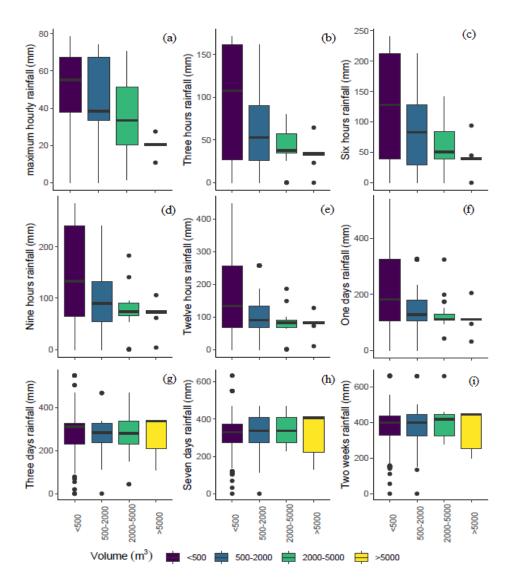




236	The analysis of the relationship between the size of debris and rainfall over different
237	time intervals has shown that short time rainfall was associated with shallow debris(fig.6 a-f),
238	as the cumulative time interval of rainfall increased, the debris size increased and stabilized at
239	three days cumulative rainfall(Fig. 6g). we observed that the increase in severity of debris was
240	associated with lower precipitation as reflected in the yellow boxplot for debris above
241	5000m3(Fig. 6. g-f), the precipitation of occurrence was below the corresponding median
242	rainfall. Therefore the lower precipitation on prolonged time greater or equal to three days was
243	responsible for severe debris. On the other hand, short-duration heavy rains were responsible
244	for shallow debris(Polemio and Petrucci 2000). The relationship between shallow debris and
245	heavy rain was reflected in the boxplot representing debris below 500m ³ (Fig. 6 a-f), where the
246	interquartile range of precipitation was large.







247

Figure 6. Distribution of severity of debris per antecedent rainfall groupped in different timeintervals

250

The precipitation at the time of the incident exhibited an inverse relationship between the size of debris and rainfall intensity (Fig. 7 a-f). From three days' cumulative of antecedent rainfall, the relationship became almost constant in all time-based cumulative rainfall (Fig. 7 g-h). The rainfall of lower intensity falling over a prolonged period was observed to trigger the

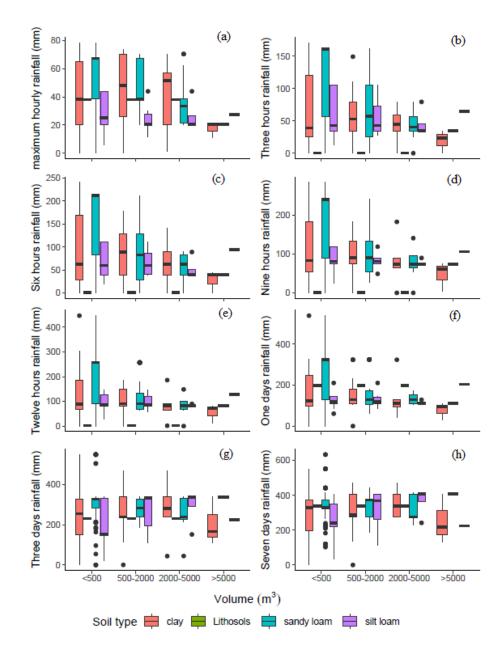




255	large size of rainfall-induced debris (Tang et al. 2017; Rivas et al. 2022). The threshold for the
256	triggering factor was the rainfall duration of atleast 3days (Chinkulkijniwat et al., 2020;
257	Rahardjo et al., 2020). In terms of soiltype, clay exhibited shallow debris at lowest
258	precipitation followed by sandy soil. The median of the occurrence of shallow debris was the
259	lowest for silt loam compared with other soil types. From three days cumulative rainfall(Fig.
260	7g), the median of occurrence of all size of debris stabilized around 300mm of precipitation
261	and clay was more likely to produce severe debris.









263 Figure 7. Relationship between rainfall and the size of debris.

264





265 3.3. Method Description

266 The analysis, construction, and evaluation of models were done in the following chronological order as depicted in Fig. 2. First, the data set was curated: formatting the 267 variables to match their types, that is, numerical variables into a numerical format, and 268 269 categorical variables into factor and ordered factors according to their natural characteristics 270 (Table 1). Second, the dataset was split into training and test set. Third, the four machine learning algorithms (Kainthura and Sharma 2022; Su et al., 2022), were applied to the training 271 and test set on all variables consecutively. Finally, confusion matrices for each method were 272 generated to compare the performance based on the overall accuracy and kappa. The variable 273 ranking plot was generated to identify the cause of differences in the predictive power of the 274 275 four methods.

All analyses were done using the following packages in R software: Caret (Max 2022) for the creation of confusion matrix, dplyr (Wickham et al. 2022) for data manipulation and formatting, MASS for running proportional odd logistic model, Random Forest for running the random forest algorithm, Xgboost for running extreme gradient boost, SVM for supporting vector machine, Ggplot2 (Wickham 2016), alluvial (Brunson 2020) for plotting, and Matrix for creation of sparse matrix which is used in training extreme gradient boosting.

Variable importance is a systematic approach for identifying the contribution of input variable in the prediction of the outcome variable in the predicative model. For graphical representation of variable importance (Biecek and Burzykowski 2021), the plot was made using DALEX libraries (Biecek 2018) and the e1071 package (Meyer et al. 2021) in R (Team 2021).

286 This plot ranks variables according to its influence on the predictive power of the model.

The selected methods for modeling were chosen based on low parsimony and are frequently applied to ordered outcome problems. The predictive performance for each model was evaluated using confusion matrix information accuracy and *kappa* (Caelen 2017). The





- kappa statistic *k* measures the agreement between the observed and predicted values to quantify the ability of the model to classify the output variable into their appropriate classes or categories. Let both the observed Y variable and predicted Z variable have g categories or levels. Let f_{ij} be the frequencies of observations in the *i*th categorical output variable Y and the *j*th category of the predicted values, then the frequency table known as the confusion matrix can be arranged as follows: Z=1 Z=2 ... Z=g

296

297 The observed (actual values) ratio of agreement between Y and Z is expressed as:

$$p_0 = \frac{1}{n} \sum_{i=1}^{g} f_{ii}$$
 (1)

299

298

300 and the expected agreement by chance is defined as:

301
$$p_e = \frac{1}{n^2} \sum_{i=1}^g f_{i+} f_{+i}$$
(2)

302

303 where f_{i+} is the total for the ith row f_{i+} is the total for the ith column. The value of kappa is 304 the estimate of the population coefficient calculated using the following formula:

305
$$k = \frac{\Pr[Y=Z] - \Pr[y=z|Y \text{ and } Z \text{ are independent}]}{1 - \Pr[y=z|Y \text{ and } Z \text{ are independent}]}$$
(3)

306





307 The confusion matrix is useful for analyses, control tunes of different classifiers, and 308 identification of a combination of classes with its recognition values or rates (Susmaga 2004). 309 The confusion matrix is a $(g \times g)$ dimension table (matrix) which matches predicted values 310 from the model vs. actual values from the dataset, where g stands for levels of the outcome 311 variable, entries on the diagonal represents the correct classification, and non-diagonal elements represent misclassification. The accuracy is defined as the ratio of correctly classified 312 entries and the sum of correctly classified and misclassification. The accuracy of the model is 313 compared to the value of the No information rate (NIR). The NIR is the baseline for assessing 314 315 performance, not 0.5. For the model to be useful (better than random guess), the lower bound 316 for a 95% confidence interval (CI) of prediction must be greater than NIR (Garson 2021). The 317 next paragraphs describe each method in detail.

The proportional odds logistic model (POLR) (McCullagh 1980) is one of the usefull methods designed to handle ordered or ranked outcome variables when the outcome categories (levels) are more than two. This model is constructed based on cumulative probability distribution (Brant 1990), $y_j = \Pr(y \le j)$ and is expressed in the form:

logit
$$[y_j/(1-y_j)] = \theta_j - \beta^t X$$
 (4)

where y is a set of N and independent observation taking values j= 1, 2, ..., k, X is a vector of independent variables, $\theta_1 < \theta_2 < \cdots < \theta_{k-1}$ and β^t are unknown parameters. To use ordinal regression, assumptions must be satisfied.

To use the proportional odds logistic regression, the proportional odd assumption or the parallel regression conditions must be satisfied. The first states that no independent variable has a disproportionate effect on any level of the dependent variable (McNulty 2021). If this condition is not satisfied, other methods such as adjacent logit models may be used (Agresti, 2010; Harrell, 2015). To test the parallel regression, the Brant-Wald test is used and this test





compares the general ordinal logistic regression (with no assumption of proportional odd) with 331 332 POLR. It is the test of significance in the difference between the two models, which produces 333 chi-square statistics. If the p-value of each of the different coefficients of variables in the model 334 is greater than 0.05, then the parallel regression assumption holds (Brant 1990) or the 335 assumption is violated otherwise. Binary logistic regression is frequentry used for landslide susceptiblity mapping(Yesilnacar and Topal 2005; Yilmaz 2009; Lin et al. 2017; Lombardo and 336 Mai 2018; Sun et al. 2021), the ordered logistic regression is an extension for binary logistic 337 applied to solve problems with multilevel ordered outcomes(Brant 1990). 338

339 The second used method was the Random Forest which is a classification and regression 340 methods. The RF algorithm (Breiman 2001) is a combination of tree predictors and every tree is made based on values of random vectors, which are sampled independently using the same 341 342 distribution to create all trees of the forest. In this study, the RF classification method (Biau 343 and Scornet 2016; Lechner and Okasa 2020;) is appropriate due to its capability to handle multiple outcome-related problems (Diaz-Uriarte and de Andrés 2005). It was applied in 344 345 different regions for landslide susceptibility mapping(He et al. 2021; Sun et al. 2021; Huang et al. 2022). 346

The third used method is SVM, which was widely used for mapping the likelihood of landslides (Lee et al., 2017). Among the multiple class prediction methods, the SVM method performed better for protein fold recognition (Ding and Dubchak 2001;Huang and Zhao 2018), landslide hazard (Hong et al. 2015), landslide spatial prediction (Pham et al. 2018), etc. SVM performed not only in multiclass problems but also better in ordered multilevel problems and it worked better than traditional regression methods (Li et al. 2012). More details about the SVM algorithm is described by Noble (2006).

The last but not least to choose was the EGB. Extreme gradient boosting is also a machine learning algorithm known for its high-speed performance and efficient prediction accuracy





- 356 (Chang et al. 2018). The algorithm is based on Friedman's work (Friedman 2001) and
- implemented in R under the package xgboost (Chen et al. 2015). This method is known for its
- high performance on larger sample sizes (Georganos et al. 2018).

359 4. Analysis & discussion

To identify the suitable model for the prediction of the size of debris, previously discussed

361 methods were explored, and results were summarized and compared in this section.

362 4.1 Debris prediction using POLR

The significance of the observed relationship between the size of the debris and the associated explanatory variables was measured using proportional odd logistic regression. Variable selection was made using backward selection (Andersen 2010). The above coefficients can be interpreted as follows: taking into consideration of p-values, and for each case supposing that all other coefficients were held constant, the decrease in three hours' rainfall is associated with a 94% lower odd the size of debris, an increase is six hours' rainfall was associated with 4% increase in higher odd of the size of the debris (Table 2).

370 Table 2. POLR model coefficients.

Variables	Coefficient	P value	Odds ratio
Three hours rain	-0.051	< 0.01	0.949
Six hours rain	0.039	< 0.01	1.040
One day rain	-0.009	0.01	0.990
Slope length	0.039	< 0.01	1.040
altitude	0.001	0.02	1.001
Drainage: Good	1.246	0.13	3.477
Drainage: Very good	1.818	0.04	6.164
Intercepts:			
<500 500-2000	2.313	< 0.01	10.105
500-2000 2000-5000	4.455	< 0.01	86.102
2000-5000 >5000	6.428	< 0.01	619.397

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- The goodness of fit and model diagnostic result shows that McFadden pseudo-R² was 0.23 where the values between 0.2 and 0.4 indicated an excellent fit (Louviere et al. 2000). To test proportional odd assumptions, the Brant-Wald test (Brant, 1990) was used, and the result was summarized in Table 3. It shows that all probabilities for all variables including the omnibus are greater than the 0.05 threshold, except for slope length, which proves that the parallel regression assumption is satisfied.
- 379 Table 3. Parallel regression assumption test.

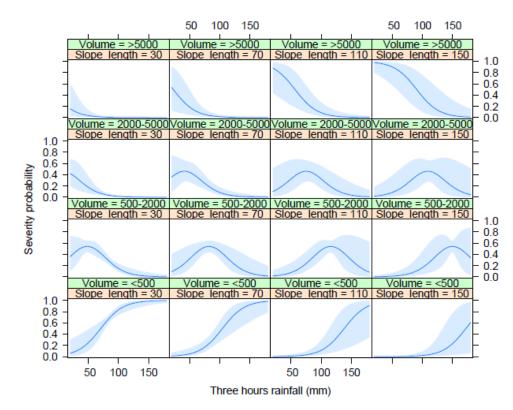
Test for	X ²	df	Probability
Omnibus	12.6	14	0.56
Three hours rain	1.76	2	0.41
Six hours rain	1.18	2	0.56
One day rain	1.18	2	0.55
Slope length	7.1	2	0.03
altitude	1.31	2	0.52
Drainage	0.19	4	1

380

The effect plot was used to demonstrate the change, in the likelihood of occurrence of 381 landslides of a given volume, associated with the change in selected predictors. Figure 7 depicts 382 the variation of probability of landslides of a given volume per three days' rainfall and slope 383 384 length. The likelihood of occurrence of debris below 500m³ was directly proportional to an increase in the rainfall for the slope length below 30m. There was a decrease in the occurrence 385 386 of debris larger than 500m³ as the rainfall increased. This decrease is an indication that shorter slope length was associated with a shallow debris (First column of Fig. 7). Moving from the 387 first to the 4th column, the probability of shallow debris (below 500m³), the probability shifted 388 from 0.9 to 0.6 as slope length increased from 30m to 170m, and the long tail of probability 389 plots for debris of size above 500m³ disappeared as the slope-length and rainfall increased. This 390 shifted up the probability curve for critical debris from 0.2 to 0.9 for slope length below 30m 391 392 to 0.9 for slope length above 110m, respectively.







393

394 Figure 8. Effect of three hours of rainfall and slope length on size of debris.

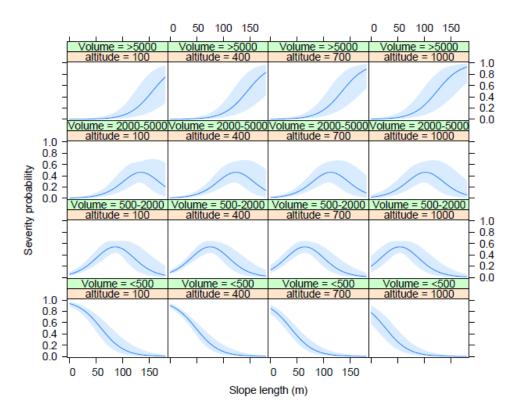
395

Figure 9 illustrates that the maximum size of debris occurred at slope lengths above 50m 396 and with rainfall between 80mm to 110mm. Looking at the two upper light hand corners of Fig. 397 9, the probability of occurrence of debris above 5000m³ peaked at slope length above 100m 398 and fade as the rain value increased, this fading associated with rain does not mean that there 399 is an inverse relationship with rainfall but it associated with the rarity of heavy rain in the 400 401 dataset. The probability of occurrence of shallow debris decreased as both altitude and slope 402 length increased. On the other hand, the chance of occurrence of debris of size between 500m³ 403 and 2000m³ increased with slope length and attained the maximum length between 70 and 80m 404 with the maximum probability of occurrence of 0.6, and the last one decreased for slope length





- 405 above 100m. For the medeium sized debris, the maximum probability of occurrence was 0.7
- 406 and associated with slope lengths between 120m and 150m. The probability of occurrence of
- 407 critical debris increased exponentially with slope length and altitude.



408

409 Figure 9. Effect of altitude and slope length on size of debris.

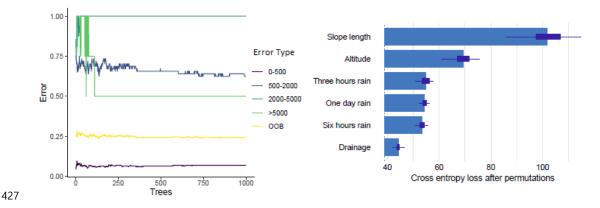
The intuitive explanation of the continuous decrease with shallow debris is associated with the cumulative character of the model, which means that, what looks like a decrease is not a real decrease, it is a shift from a lower level to a higher level associated with an increase in variables into consideration. For altitude below 300m, shifting from a size below 500m³ to 500-2000m³ happened at 20m of slope length, while shifting from 500-2000m³ to 2000-5000m³ and above occurred at slope length between 60 to 80m.





416 4.2 Debris prediction using RF

The random forest model was run on the training set of 304 observations. The number of 417 grown trees was 1000, the variable tried at each split was 4, and the out-of-bag error estimates 418 419 (OOB) error rate was 24.42%. Figure 10a depicts the predictions of a class below 500m³ that 420 had the least errors. Since the model is for classification, to calculate the prediction error, the 421 model was run on the test set. The variable importance graph illustrates the slope length was the most contributing variable to the model accuracy (Fig. 10b). The mean decrease Gini (cross-422 423 entropy loss after permutation) value is the measure of the contribution of each variable to the 424 homogeneity of nodes and leaves in the random forest (Martinez-Taboada and Redondo 2020). 425 The higher the value, the more important the variable is, which exhibited that slope length is more influencing factor for the size of debris. 426



428 Figure 10. RF model: (a) training error rate, and (b) variable importance.

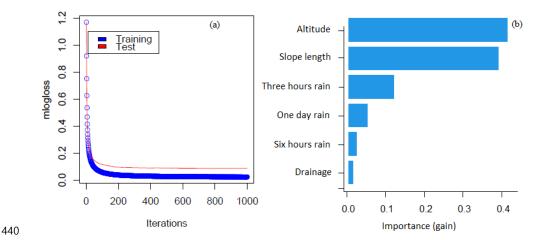
429 4.3 Debris prediction using EGB

To run the EGB, the train and test set were transformed into a sparse matrix, as it is run on numerical data matrices (Chen et al. 2022). The optimum model was obtained at the 881 number of iterations and the learning rate of 0.3. Since the task was a multi-classification, the multi: softmax objective was used (Chen et al. 2022). The evaluation metric was mlogloss





(Multi-class log loss) (Kabani and El-Sakka 2016). While training the model, the minimum mlogloss were 0.025 and 0.088 for the training and test sets respectively (Fig. 11a). The difference between training and the test error is due to the larger variance associated with the fewer number of observations in the test set. The extreme gradient boosting associated with higher importance to slope length, altitude, slope and three hours of rainfall, respectively (Fig. 11b).



441 Figure 11. EGB model: (a) error rate and (b) variable importance.

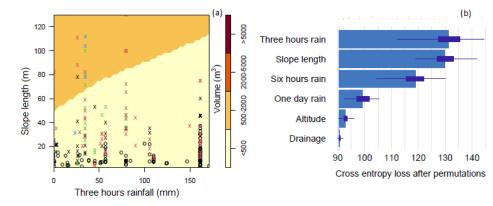
442 4.4 Debris prediction using SVM

The support vector classification model with linear kernel was applied, and the number of support vectors was 142. The outcome of SVM model shown a higher performance rate on the test than the training set. Figure 11a depicts a two-dimensional projection of train data using slope length and three hours of rainfall showing different shading and support vectors. One of the weaknesses of the SVM, it classified predictions into two lower classes as indicated in Fig. 12a, it couldn't distinguish the moderates from extreme debris. The SVM assigned higher importance to slope length and three hours of rainfall as the second high ranking variable as





450 shown in Fig. 11b.



451

452 Figure 12. SVM model: (a) SVM classification plot, and (b) variable importance.

453 *4.4 Discussion and model suitability assessment for landslide-induced debris severity*

454 prediction

This study analyzed the relationship between the severity of landslide-induced debris. The 455 exploratory data analysis revealed that 93% of occurred debris was below 2000m³. The 456 Gangwon province and Seoul were more vulnerable regions in terms of the frequency of 457 incidents. Despite a higher frequency of debris in Seoul, their size was small compared to 458 Gangwon province with higher number of cases and more large sized-debris. To analyse the 459 significance of the relationship between the size of debris with different models ranked the 460 461 slope length as the most influential variable for the size of the debris. To visualize the 462 association of the slope lengh with the size of debris, the scatter plot (Fig.13) revealed that the pattern of increasing trend of slope length and size of debris was more remarkable across all 463 provinces, shallow debris were associatted with slope length blow 50m. The debris cases were 464 465 clustered between 20° and 40°, as depicted in Fig.13, and critical debris tended to be clustered 466 around 30 degrees; the association of slope angle and size of debris was not statistically significant. For the soil-debris size relationship, sandy soil and clay were associated with a 467

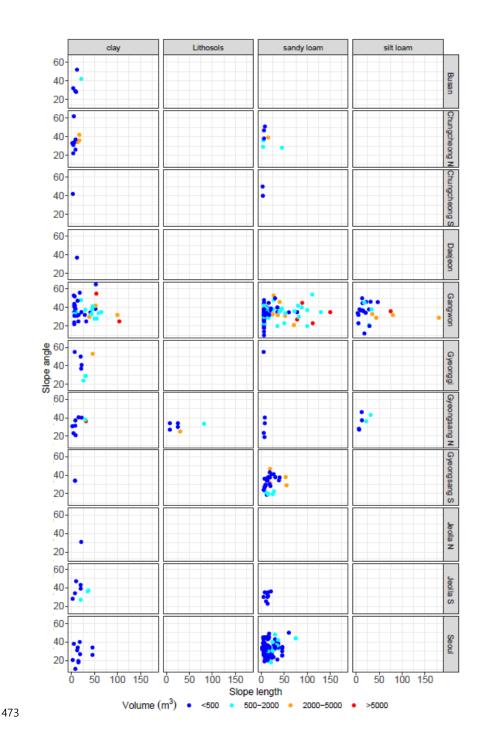




- 468 higher frequency of debris; they exhibited shallow debris in all provinces except in Gangwon
- 469 where severe debris occurred. It was observed that even though the silt loam soil was not highly
- 470 frequent, it was vulnerable to severe debris as the slope length increased. Gangwon province
- 471 was the region where the increasing relationship between the size of debris was observed, other
- 472 provinces were not prone to severe debris.







474 Figure 13. Province-wise scatter plot of debris' size per slope length, slope angle and soil





475 types.

476

477 To assess the suitability of utilized models, it was observed that the ordinal regression prediction power was low due to its inherent weakness; it does not perform well on the 478 479 imbalanced dataset (Agresti 2010). That is, the outcome variable has some more prevalent 480 levels than others. It also estimates the coefficient using maximum likelihood techniques, 481 which require a large sample size, the used dataset was highly imbal 482 anced. Figure 7 revealed that almost 80% of all occurred debris was shallow, as result, the model coefficients became unstable with larger prediction intervals. Despite the weakness, the 483 ordinal model has the effect display (Fox and Weisberg 2019), which clearly shows the effect 484 of each variable to each extent and the associated probability. The predictive performance for 485 each of the four discussed methods was summarized in Table 4 to facilitate their comparison. 486 The random forest model performed well in all cases on the training set and validation as well. 487 488 This model associated the influencing factor with higher importance and lower importance to 489 rain-associated variables. RF prediction accuracy was very high on the training set (0.93), and 0.90 on the test set. The prediction at 95% of confidence interval width ranged from 0.84 to 490 491 0.94 on training and test sets, respectively. The NIR shifted from 72.28% to 74.17%, and this small increase is due to a small sample of the validation set. 492

The model accuracy for POLR was quite moderate based on the *kappa* value of 0.38. The performance accuracy was better, the no information rate, NIR=0.7228<0.7314 lower boundaries of CI on the training set. This last condition was not satisfied for the test set, that is NIR=0.74 was higher than 0.69, which was the lower CI for the prediction interval on the test set. Based on the p-value for the prediction on test 0.26 > 0.05, the performance was not reliable. This is confirmed by the overall performance metric *kappa* =0.30, which was quite moderate.

499





500

501	The EGB model ranked the second best model after the RF, it satisfied all prediction
502	conditions the overall performance ($kappa = 0.6121$) was slightly below the random forest
503	0.7273. The SVM model result was satisfactory for the training set. The lower bound of very
504	close to the information rate and the overall performance was moderate (kappa=0.32), the
505	weakness of the model was its incapacity to distinguish the moderate debris from the extreme
506	ones; as result, it predicted all debris into two lower categories in Fig. 8a. The NIR fell into the
507	prediction interval on the validation set, and the p-value was 0.11>0.05, which is an indication
508	of moderate prediction accuracy.

509 Table 4. Model accuracy statistics for the four methods.

Model accuracy statistics						
Method	Data	Accuracy	95% CI	NIR	P-Value	kappa
	Train	0.93	(0.89, 0.95)	0.72	< 0.001	0.82
RF	Test	0.90	(0.84, 0.94)	0.74	< 0.001	0.72
	Train	0.78	(0.73, 0.82)	0.72	0.011	0.38
POLR	Test	0.76	(0.69, 0.83)	0.74	0.26	0.3
	Train	0.77	(0.72, 0.82)	0.72	0.015	0.32
SVM	Test	0.78	(0.71, 0.85)	0.74	0.11	0.3
	Train	0.86	(0.81, 0.89)	0.72	< 0.001	0.63
EGB	Test	0.86	(0.79, 0.91)	0.74	0	0.6

510

The landslide-induced debris prediction is an extension of landslide susceptibility mapping and 511 may be useful in the quantification and prediction of debris resulting from a rainfall-induced 512 513 landslide. This quantification can facilitate risk management (Ho and Ko 2009), in the identification of regions prone to severe debris and the making of policies for mitigation 514 (Carmela and Mario Parise 2022). For example the decision of planting more vegetation that 515 fits the conditions of the region to strengthen the soil or deciding an appropriate activity to be 516 517 done in a given region to improve stability ,safety and efficient land use (Mayer et al. 2008) . Furthermore, some activities in regions prone to severe debris may be prohibited for the safety 518





and well-being of the public(Frattini et al. 2010; DeGraff and Romesburg 2020; Di Napoli et 519 520 al. 2020). In addition, the model may serve the disaster manager to create appropriate funds 521 for post-disaster recovery. For example, if a region is expected to have shallow debris, the manager may establish a small fund for paying minor labour to repair the damaged environment. 522 523 In extreme cases, a big fund may be created to pay for machinery and construction of preventive walls, plantation, and cost of machinery to remove debris in the affected region, to rehabilitate 524 the impacted economic activities in the neighbourhood (Kachi et al. 2016). Due to the lack of 525 financial data associated with the inventories the cost of post-disaster recovery was not 526 527 estimated, more studies in the future may be carried out to fill this gap. The approach in this paper is valid for the studied area based on the user input data; more research in the future may 528 529 be conducted to know whether the findings in this paper are general for regions with different 530 characteristics or settings.

531 6. Conclusions

532 The study analyzed the relationship between the size of rainfall-induced debris and causal factors, i.e., time-based cumulative rainfall and influencing factors: soil types, vegetation, and 533 geomorphology features. The exploratory data analysis revealed that the Gangwon province is 534 535 prone to more frequent and more severe landslide-induced debris. Soil-related information revealed that the landslide-induced debris was more frequent in sandy soil and more severe, 536 but its influence was not statistically significant in the predictive model. The region with non-537 538 perfected drainage systems also experienced severe debris. The regions with old timber that 539 experienced fire had a higher debris likelihood. To examine the significance and to identify the 540 suitable model for landslide-induced debris severity, four predictive modeling techniques i.e., 541 POLR, RF, EGB and SVM, were applied to examine the causal and influencing factors of the 542 severity of rainfall-induced debris in South Korea. The performance metrics, accuracy and





543	kappa were applied to compare the predictive power of each of the four methods. The findings
544	of this research revealed that three hours' rainfall, one-day rainfall, and slope length were the
545	most influencing factor and altitude took the second place. This finding was consistent with the
546	results of Lee et al. (2012), stating that short-duration rainfalls were responsible for landslides
547	and are the cause of their severity. The comparative analysis has shown that random forest had
548	better predictive power with an accuracy of 90% and $kappa = 0.72$, and extreme gradient
549	boosting followed with an accuracy of 86% (<i>kappa</i> = 0.6). The last two methods SVM with an
550	accuracy of 78% (4% above NIR), and POLR performance was moderate at 76%, which is
551	only 2% above 74% performance decision basis (No information rate NIR), but we did not
552	have enough information to confirm their use as a basis for creation early warning system for
553	rainfall-induced extreme debris. This is because POLR does not perform well on limited and
554	imbalanced data (Rahman et al. 2021), which is the root cause of a wide range of prediction
555	intervals. Thus, RF and EGB may be used as a suitable models for rainfall-induced debris
556	prediction. The creation of a nationwide landslide database would solve the shortage of reliable
557	data and allow the usage of more alternative methods, which will result in more improved
558	models. The findings of this research may be used for the elaboration of rainfall-induced debris
559	mitigation policies such as post-disaster rehabilitation planning and land use management.

- 560 List of abbreviations
- DALEX: moDel Agnostic Language for Exploration and explanation.
- MASS: Modern Applied Statistics with S.
- SVM: Support Vector Machine.
- POLR: Proportional Odd Logistic Regression.
- EGB: Extreme Gradient Boosting.
- RF: Random Forest.





- NIR: No information Rate.
- Acc: Accuracy.
- CI: confidence interval.
- OOB: Out Of Bag error estimates.
- CFM: Confusion Matrix.
- Caret: Classification and Regression Training

573 Acknowledgments:

- 574 This research was supported by Basic Science Research Program through the National
- 575 Research Foundation of Korea (NRF) funded by the Ministry of Education
- 576 (2021R1A6A1A03044326), and the grant (NRF-2021R1C1C2003316) funded by the National
- 577 Research Foundation of Korea (NRF).

578 Data availability:

579 The datasets used and/or analyzed during the current study are available from the 580 corresponding author on reasonable request.

581 **Declaration of Competing Interest:**

- 582 The authors declare that there are no conflicts of interest.
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