

Review article: Towards Improved Drought Prediction in the Mediterranean Region – Modelling Approaches and Future Directions

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Abstract.

There is a scientific consensus that the Mediterranean ~~region~~ (MEDRRegion (MedR)) is warming and as the temperature continues to rise, ~~extreme events such as~~ droughts and heat waves are becoming more frequent, severe, and widespread. Given the detrimental effects of droughts, it is crucial to accelerate the development of forecasting and early warning systems to minimize their negative impact. This paper ~~examines~~ reviews the current state of ~~knowledge in~~ drought modeling and prediction ~~using~~ applied in the MedR, including statistical, dynamical, and hybrid statistical-dynamical models. ~~By considering the multifaceted nature of droughts, the study encompasses meteorological, agricultural, and suggests some hydrological drought forms and spans a variety of forecast scales, from weekly to annual timelines. Our objective is to pinpoint the knowledge gaps in literature and to propose potential research prospects to further trajectories to improve drought the prediction of droughts in this region. The review finds that while all methods have their each method has its unique strengths and shortcomings limitations, hybrid statistical-dynamical methods can perform models appear to hold the most promising potential for skillful prediction with a long seasonal to annual lead times.~~ However, the application of these methods is still challenging due to the lack of high-quality observational data and the limited computational resources. Finally, the paper concludes by discussing the importance of using a combination of sophisticated methods such as data assimilation techniques, machine learning models, and copula models and integrating data from different sources (e.g., remote sensing data, in-situ measurements, and reanalysis) to improve the accuracy and efficiency of drought forecasting.

Key Words: drought, forecasting, data assimilation, machine learning, Mediterranean, review

1 Introduction

Drought is a recurrent phenomenon in the Mediterranean ~~basin~~ (MEDBRegion (MedR)). Throughout time, adaptation to this kind of climate ~~events~~ event has been an important issue for the development of many countries in the region. Yet, with the disruptive accelerated impact of global warming, already reflected in more regular and intense droughts around the Mediterranean in the last few decades, building resilience to extreme weather conditions remains a true challenge (Satour et al., 2021). ~~For these reasons among others,~~ the region is often described as a ~~Hotspot~~ hotspot for climate change (Tuel and Eltahir, 2020). The Intergovernmental Panel on Climate Change (IPCC) pointed out in the Sixth Assessment Report (AR6) that global warming has been more rapid in the Mediterranean than in the rest of the world (IPCC, 2021). ~~This report projected an increase in the~~

38 frequency and/or severity of agricultural and ecological droughts across the Mediterranean and Western Africa-
39 (IPCC, 2021). A global increase of 2 °C is thought to correspond to a 3 °C increase in the daily maximum
40 temperature in the MEDB MedR (Seneviratne et al., 2016; Vogel et al., 2021). If this increase in temperature
41 continues at the same pace, the Mediterranean region (MEDR) MedR is susceptible to experience fearful
42 desertification by the end of the 21st century, driving an increase in aridity- (Carvalho et al., 2022).

43 This will surely lead to irreversible biodiversity loss and reduced diminish the capacity capability of semi-arid
44 Mediterranean ecosystems to function as an effective carbon sinks sinks in the future (Valentini et al., 2000;
45 Briassoulis, 2017; Zeng et al., 2021). forthcoming. All these These conditions exacerbate water stress that which
46 in turn, enhances in turn the probability of wildfire. (Turco et al., 2017a). A phenomenon already witnessed these
47 two last summers (2021 and 2022) in several Mediterranean countries (Turkey, Greece, Italy, Algeria, and
48 Morocco), displacing thousands, killing hundreds, and causing irreparable damage (Rodrigues et al., 2023; Yilmaz
49 et al., 2023; Eberle and Higuera Roa, 2022).

50 The Mediterranean Sea (MEDS) is the body of water that separates three continents: lying between Africa,
51 Europe, and Asia. serves as a substantial source of moisture and heat, affecting atmospheric circulation and
52 weather patterns (Mariotti et al., 2008). Its narrow connection to the Atlantic Ocean via the 14 km wide Strait of
53 Gibraltar is only 14 km wide. The MEDS is surrounded by and the surrounding varied topography (Fig. 1), with
54 vegetated areas to the north and desert areas to the south and east with narrow vegetated areas around contribute
55 to the east region's complex climate dynamics (Michaelides et al., 2018). The topography of land surrounding the
56 MEDS is varied with the existence of complex mountain ranges with high altitudes (Fig. 1). This is one of the
57 reasons that render the dynamic characteristics of the atmospheric flow complex at various scales, playing a critical
58 role in the regional and local climate (Michaelides et al., 2018).

59 The Mediterranean climate MedR is defined as characterized by a mid-latitude temperate climate with mild rainy
60 winters and hot, dry summers (Lionello et al., 2023). Notably, this area is positioned in a transitional band
61 between the midlatitude and subtropical regions, which makes climate modeling for this region quite challenging
62 (Planton et al., 2012). Precipitation has a marked annual cycle, with The Mediterranean climate exhibits a strong
63 spatial gradient in precipitation, with generally decreasing precipitation values towards the south and hardly any
64 precipitation during the summer. It is also unevenly distributed and characterized by a strong spatial gradient, with
65 values decreasing toward the South (Lionello, 2012). Droughts occurring during the wet season (or during the crop
66 growing season) Such conditions pose challenges in climate modeling and can severely impact lead to severe
67 impacts on water supply, and agricultural production agriculture, especially for countries in regions relying mostly
68 on rain-fed agriculture (Tramblay et al., 2020).

69 Water availability is unevenly distributed among the Mediterranean countries with 72% in temperate countries of
70 the North, against 5% in the South, and 23% in the East (Milano et al., 2013). Accordingly, several countries such
71 as Algeria, Morocco, Egypt, Libya, Malta, and some countries of southern Europe such as Portugal and Spain are
72 experiencing a structural water shortage that is likely to increase with the expected population growth. (Sanchis-
73 Ibor et al., 2020). This situation is further aggravated when multi-annual droughts hit the region. Therefore In this
74 challenging context, drought forecasting at a sufficient that provides seasonal to annual lead time is of primary
75 importance times becomes critically important for the proactive management of agricultural and water resources
76 and agriculture in this difficult context. management.

77 Growing concern about the drought phenomenon in the last decades has spurred the development of improved
78 systems that predict the full cycle of drought (onset, duration, severity, and recovery) via a large number of indices
79 and models. Common approaches to predicting drought can be subdivided into two categories of models: statistical
80 models and dynamical models. Statistical models, also named data-driven models, rely on the estimated
81 correlations between several predictors (large-scale climate variables) and predictands (local climate variables
82 represented by historical observations). ~~While~~ The climatology-based or persistence-based models, like the
83 Ensemble Streamflow Prediction (ESP) system, form an essential tool in this category, leveraging both historical
84 and near real-time data to generate a probabilistic forecast of future drought events (AghaKouchak, 2014a; Turco
85 et al., 2017b; Torres-Vázquez et al., 2023). ~~Meanwhile,~~ dynamical drought prediction relies on the use of Global
86 Climate Models (GCMs) to simulate the dynamical processes that govern hydroclimatic variability. Nevertheless,
87 despite the usefulness of these models in drought prediction and early warning systems, their forecast accuracy
88 remains limited for longer lead times (exceeding one month) (Wood et al., 2015). The post-processing and multi-
89 model ensemble techniques are usually used to improve prediction skills by avoiding systematic bias related to the
90 coarse resolution of GCMs (Han and Singh, 2020). Recently, drought prediction has also been tackled by the hybrid
91 statistical-dynamical models which combine the two approaches mentioned above. These models constitute a
92 promising tool for long lead-time drought forecasting (Ribeiro and Pires, 2016).

93 Despite the efforts made to predict drought phenomena, it remains largely little understood due to its multiple
94 causing mechanisms and contributing factors (Kiem et al., 2016; Hao et al., 2018). The complexity and variability
95 depicted by many physical mechanisms such as Sea Surface Temperature (SST), North Atlantic Oscillation
96 (NAO), El Niño—Southern Oscillation (ENSO), Mediterranean Oscillation (MO), and land-atmosphere feedback
97 are also responsible for the low performance of drought monitoring and forecasting (Ayugi et al., 2022). ~~The~~
98 ~~MEDB is positioned in a transitional band between the midlatitude and the subtropical regions rendering climate~~
99 ~~modeling very challenging (Planton et al., 2012).~~ Understanding the synoptic conditions leading to the drought
100 phenomenon becomes increasingly important given the upward trend in temperature in ~~particular in the~~
101 ~~Mediterranean region~~ MedR. Further investigations to assimilate how large-scale teleconnections affect local
102 weather and climate anomalies, as well as how these ~~latter~~ later feedback into the larger context, are much needed
103 in this context.

104 To address these questions, many numerous review papers tried have sought to bring together consolidate the
105 scientific advances in the field of drought prediction from different regions of the world (e.g., Mishra and Singh,
106 2011; Hao et al., 2018; Fung et al., 2019; Han and Singh, 2020). ~~However, drought is a region-specific phenomenon~~
107 ~~since the meteorological conditions that drive its onset (precipitation deficit, high temperature, soil moisture,~~
108 ~~evapotranspiration [ET]...)~~ ~~depend highly on the considered region. Consequently, solutions developed and~~
109 ~~successfully applied in one region may not necessarily be appropriate to others.~~ While these studies provided a
110 comprehensive overview of drought prediction at a global scale, our paper offers an in-depth analysis of drought
111 prediction methodologies specifically applied to the Mediterranean context. This is achieved through an
112 examination of the applicability, strengths, and limitations of statistical, dynamical, and hybrid statistical-
113 dynamical models, in line with the regional specifics of the MedR. This specificity is vital given that drought, as a
114 phenomenon, is highly region dependent. The unique meteorological conditions of the MedR necessitate dedicated
115 studies, as solutions developed for other regions may not be applicable or effective here.

Tramblay et al., (2020) emphasized the urgent need to develop for drought modeling and forecasting tailored methods designed for the Mediterranean context. This research highlights the complexity and challenges for, particularly as climate change continues to exacerbate drought conditions in this region. Building on this, our work not only emphasizes the complexities of drought assessment in the MEDR under anthropogenic and climate change effects. This paper is intended to fill the knowledge gaps in the Mediterranean drought, reviews the but also conducts a critical review of recent drought forecast methods, and focus on the prospects forecasting methodologies applied specifically to the MedR. In addition to shedding light on the merits and limitations of these methods, our investigation also helps identify underexplored areas that constitute a promising tool to overcome the actual warrant further research. Detecting these gaps is a crucial aspect of our work, as it directs future research towards these relatively unexplored realms of drought prediction weaknesses.

The structure of this paper is as follows: Section 2 highlights the difficulty related to the definition of drought from different perspectives. The causes of drought in MEDR MedR are provided in section 3. Sections 4, 5, and 6 present the recent advances in drought prediction with statistical, dynamical, and hybrid statistical-dynamical models respectively. Section 7 discusses the results found in this review, providing insights into the current state of drought forecasting in the MEDR MedR and highlighting potential areas for improvement. The challenges in drought prediction are reviewed with the prospects in section 8. Finally, the 9th section presents the conclusions of the whole paper.

Figure 1 Topography of the Mediterranean Region: (30°N - 46°N in latitude and 10°W - 40°E in longitude).

2 Drought Definitions, Classification, and Indices

Drought is a compound phenomenon of creeping nature. Establishing an accurate prediction, well describing its starting date and duration is extremely hard. The multidisciplinary and multiscale nature of drought renders the understanding of this phenomenon very challenging: (AghaKouchak et al., 2021). As a matter of fact, literature gives numerous definitions for drought.

In the eighties, Wilhite and Glantz (1985) found more than 150 published definitions of drought that can be categorized into four broad groups: meteorological, agricultural, hydrological, and socioeconomical. This classification based on both physical and socioeconomic factors is still adopted today. As this classification is human-centered, some recent works emphasized the need to consider the ecological drought as well, which creates multiple stresses in natural ecosystems, see for example (Crausbay et al., (2017); Vicente-Serrano et al., (2020); Bradford et al., (2020); and Zhang et al., (2022). Since the aim of this study is to review forecasting drought methods, we will focus only on the first three categories that provide direct methods to quantify drought as a physical phenomenon.

In an attempt to associate a mathematical definition with each drought type, several drought indices have emerged. These indices are typically based upon some hydroclimatic variables or parameters (indicators) such as temperature, precipitation, soil moisture, streamflow, and snowpack to describe three major characteristics of the drought event: severity, duration, and frequency. However, the lack of a universal definition of drought is also apparent in the huge variety of indices (more than 100) that have been developed for drought prediction: (Lloyd-Hughes, 2014). Unfortunately, this plethora of indices creates more confusion than clarity (Lloyd-Hughes, 2014) and makes the choice of the most suitable indices a difficult task.

2.1. Meteorological Drought

The World Meteorological Organization (WMO) characterizes meteorological drought as “a prolonged absence or marked deficiency of precipitation deficit over a continuous period (”. Similarly, the IPCC defines meteorological drought as “a period of abnormally dry spell). This definition is weather in a region-specific because the determination of the over an extended period”. The threshold used to state if distinguish between a period is dry or wet period often depends on the average amount of rainfall in typical for the specific area under study area. Hence, there is. This gives rise to a considerable number variety of meteorological definitions belonging to different, each tailored to the distinct conditions of diverse regions or countries (Isendahl, 2006). Therefore, coming up with Regarding the MedR, creating a single encompassing definition of meteorological drought in the MEDR, that takes into account the is particularly challenging. This complexity of its stems from the diverse climate and conditions across the region, particularly the pronounced variability between the eastern and western meteorological conditions responsible for the that contribute to drought. is complicated.

The Standardized Precipitation Index (SPI) (McKee et al., 1993) and the Standardized Precipitation Evapotranspiration Index (SPEI) (Vicente-Serrano et al., 2010, 2010a) are two of the most prevalent indicators used to describe meteorological drought. They owe their popularity to the recommendation of the World Meteorological Organization (WMO) (Svoboda et al., 2012). The SPI has been extensively used in previous studies for its ease of computation, its probabilistic nature, and its ability to detect drought at multiple time scales (Madadgar and Moradkhani, 2013; Chen et al., 2013; Li et al., 2020; Mesbahzadeh et al., 2020; Das et al., 2020). By fitting a probability distribution to observed precipitation data, the SPI is calculated and subsequently transformed into a standard normal distribution with a mean of 0 and a standard deviation of 1 (Livada and Assimakopoulos, 2007). Consequently, SPI values can be compared across various regions and timeframes (e.g., 1, 3, 6, 12, or 24 months). This multiscale nature of SPI enables it to capture diverse aspects of drought depending on the selected time scale. The shorter time scales (1-3 months) are suitable for monitoring agricultural drought, while longer time scales (6-12 months or more) are better suited for evaluating hydrological drought. However, it should be noted that the SPI considers only precipitation data and neglects the variability of temperature and potential evapotranspiration (PET), ignoring the effect of warming on droughts. Indeed, in relatively wet regions, precipitation deficit can constitute an important indicator for drought onset (Gamelin et al., 2022). Yet, in midlatitude (or extratropic) regions such as the Mediterranean where the climatological precipitation is modest or low, precipitation deficit may not be sufficient to measure extreme droughts. Furthermore, knowing the upward trend in temperature and the influence of high atmospheric evaporative demand (AED) in increasing severity of recent drought events in the MEDR MedR (Tramblay et al., 2020; Mathbout et al., 2021; Bouabdelli et al., 2022), the choice of drought indices needs to prioritize those including these variables in their formulation such as SPEI, or Palmer Drought Severity Index (PDSI) (Palmer, 1965) and Reconnaissance Drought Index (RDI) (Tsakiris and Vangelis, 2005) to mention but a few.

The SPEI was developed by Vicente-Serrano et al. (2010, 2010a) using the climatic water balance concept of climatic water supply and AED. It is based on precipitation and PET and has the advantage of combining the multi-scalar character of the SPI with the ability to include the effects of temperature variability (Vicente-Serrano et al., 2010, 2010a).

193 A global assessment of drought indices conducted by Vicente-Serrano et al. (2012) found that SPEI provided a
194 superior capability in capturing drought impacts, particularly during the crucial summer season. Bouabdelli et al.
195 (2022) used SPI and SPEI indices and Copula theory to study the impact of temperature on agricultural
196 drought characteristics under future climate scenarios over seven vast Algerian plains located in the Mediterranean
197 region. MedR. The results of this study confirmed that the frequency of drought events is much higher using SPI
198 while their duration and severity are more intense using SPEI. Russo et al. (2019) performed drought
199 characterization in MEDR MedR using both the SPEI and the SPI by considering the period 1980–2014. They
200 concluded Their findings indicated that SPEI is better correlated for the exhibits a stronger correlation with drought
201 conditions over a 3-months' month time scale and while SPI shows a better correlation for the 9 months, which
202 reflects the capacity month duration. This result highlights the ability of SPEI to capture earlier the early shifts in
203 the balance between ET evapotranspiration and precipitation more efficiently than SPI (Russo et al., 2019).
204 However,

205 Despite the main weakness utility of this index is its sensitivity to SPEI in drought characterization, it does have a
206 noteworthy limitation. The effectiveness of SPEI significantly relies on the method that estimates used for
207 estimating PET such as the Penman-Monteith equation, the Thornthwaite method, the Hargreaves method, and the
208 Priestley-Taylor method among others. These estimation methods can yield varying results, leading to
209 inconsistencies in SPEI values. In essence, the sensitivity of SPEI to the PET estimation method used could
210 potentially affect the accuracy and reliability of the index in representing drought conditions (Vicente-Serrano et
211 al., 2010 2010b; Stagge et al., 2014).

212 The PDSI has also been widely used to quantify the drought characteristics for a given location and time. It includes
213 precipitation, temperature, and soil moisture data to estimate water supply and demand and to reflect long-term
214 drought. But it has shown some inconsistencies when used at various locations (Wells et al., 2004). A self-
215 calibrating variant of this index (scPDSI) was proposed by Wells et al. (2004) to automatically calibrates the
216 behavior of the index by replacing empirical constants in its computation with dynamically estimated values to
217 account for the variability of precipitation and the climate characteristics between locations (Wells et al., 2004).
218 Ionita and Nagavciuc (2021) evaluated the drought characteristics at the European level over the period 1901–2019
219 using SPI, SPEI, and scPDSI. The results based on SPEI and scPDSI show that the increase in mean air temperature
220 and PET are making central Europe and the Mediterranean region MedR dryer, whereas Northern Europe is getting
221 wetter. While results based on SPI using only precipitation data did not reveal this drought variability. This
222 underscores the findings of Vicente-Serrano et al. (2012), who emphasized the benefits of using more integrative
223 indices like SPEI in understanding and predicting drought variability more effectively.

224 The MedPDSI, which is an update of the PDSI formulation in terms of its soil water balance to consider real
225 evapotranspiration (based on reanalysis data instead of PET) in the MEDB MedR, has allowed an earlier
226 identification of longer and more severe droughts (Paulo et al., 2012). (Paulo et al., (2012) compared SPI, SPEI,
227 PDSI, and MedPDSI in detecting drought characteristics in Portugal for the period 1941 to 2006. They concluded
228 that PDSI and MedPDSI are likely to identify better the supply-demand dynamics and that they may be of great
229 interest for drought warning applications, aiming namely at agriculture (Paulo et al., 2012).

230 2.2. Agricultural Drought

231 Agriculture is very sensitive to climate variation especially extreme weather. Due to its dependency on water
232 availability, this sector is strongly impacted by drought events. ~~In the Mediterranean basin, agriculture is mainly~~
233 ~~rain-fed (wheat, barley, olive, and orange trees...).~~ ~~If~~ ~~In the Mediterranean Basin, agricultural practices span both~~
234 ~~rain-fed and irrigated systems. Rain-fed agriculture is prevalent, particularly for crops such as wheat and barley,~~
235 ~~while crops like olives and citrus fruits, such as oranges, often utilize controlled irrigation systems to supplement~~
236 ~~natural precipitation (Rodrigo-Comino et al., 2021). Regardless of the system employed, if meteorological drought~~
237 ~~lasts for a prolonged period, it can lead to a reduction in soil moisture to such a level that it harmfully affects crop~~
238 ~~production, especially during the active plant growth season. (Wilhite and Glantz, 1985; Mishra and Singh, 2010).~~
239 ~~At this stage the agricultural drought sets in.~~

240 Therefore, in addition to meteorological factors, the agricultural drought definition is also related to the retention
241 capacity of soil in the crop growth season (Kuśmierk-Tomaszewska and Żarski, 2021) which depends on crop
242 types, soil characteristics, and soil management. All these indicators can be employed to develop relevant
243 agricultural drought indices. Among them, we cite Crop Moisture Index (CMI) (Palmer, 1968); Soil Moisture
244 Deficit Index (SMDI); Evapotranspiration Deficit Index (ETDI) (Narasimhan and Srinivasan, 2005); Normalized
245 Soil Moisture index (NSMI) (Dutra et al., 2008) and Empirical Standardized Soil Moisture Index (SSMI) (Carrão
246 et al., 2016).

247 ~~All~~ ~~The formulation of these indices~~ ~~include~~ ~~integrates~~ ~~soil moisture data~~ ~~,~~ ~~leveraging a variety of assessment~~
248 ~~techniques, each with unique advantages. These include~~ ~~in their formulation,~~ ~~situ soil moisture probes, cosmic-~~
249 ~~ray neutron probes, and physically driven models such as the ISBA land surface model (Tramblay et al., 2019).~~
250 ~~Each of these techniques has distinct advantages and is suitable for different application contexts (Miralles et al.,~~
251 ~~2010; Martens et al., 2017). However, when faced with the scarcity of observed soil moisture data~~ ~~are still limited.~~
252 ~~Currently, the only way,~~ ~~remote sensing comes to~~ ~~obtain~~ ~~the forefront. It furnishes extensive and frequent~~
253 ~~measurements of soil moisture characteristics~~ ~~is through remotely sensed,~~ ~~effectively supplementing areas where~~
254 ~~observed data~~ ~~(Gruber and Peng, 2022). Those have already some known~~ ~~falls short. Yet, it is crucial to be aware~~
255 ~~of the limitations~~ ~~such as the~~ ~~of these tools. Despite its indispensable role, remote sensing is constrained by factors~~
256 ~~such as coarse time~~ ~~temporal and~~ ~~space~~ ~~spatial resolution, low depth of~~ ~~limited penetration depth, and incompatible~~
257 ~~governing hydrologic principles (Mohanty et al., 2017; Gruber and Peng, 2022).~~ As an alternative, hydrological
258 models have been commonly used to simulate and calibrate this variable in the context of agricultural drought
259 forecasts (Hao et al., 2018). Mimeau et al., (2021) used a modeling framework to estimate soil moisture sensitivity
260 to changes in precipitation and temperature at 10 plots located in southern France. They concluded that the current
261 climate change scenarios may induce longer periods of depleted soil moisture content, corresponding to
262 agricultural drought conditions.

263 In general, when soil moisture in the root zone reaches a critical level, farmers resort to irrigation to save crops-
264 ~~(Kang et al., 2000).~~ However, ~~if~~ nowadays agriculture consumes approximately 85% of global fresh water for
265 irrigation (D'Odorico et al., 2019; Tatlhego et al., 2022), ~~this figure tends~~ ~~which is expected to~~ increase in the years
266 to come by growing population, increasing food consumption, and rising temperatures that accelerate PET and
267 ~~promotes~~ ~~promote~~ hydrological stress.

268 2.3. Hydrological Drought

269 Unlike agricultural drought which is mainly affected by the depletion of soil moisture after a dry period, a lack of
270 precipitation impacts many components of the hydrological system in a river basin or watershed (streams,
271 reservoirs, and lakes). These define water availability that can be used for commercial navigation, generation of
272 hydroelectric power, irrigation of farmlands, industry, and domestic activities for several months after the
273 deficiency in precipitation. Consequently, hydrological drought lags behind the occurrence of meteorological and
274 agricultural droughts. This lag time is a characteristic of the watershed, which is defined based on many physical
275 drivers such as evapotranspiration capacity, soil properties, vegetation types, snow accumulation/melt, local water
276 management such as dams' construction and control, water supply operation rules, and irrigation strategy (Van
277 Loon and Laaha, 2015).

278 A hydrological drought is generally proclaimed when the water levels in streamflow, reservoirs, lakes, aquifers,
279 and other water storage systems fall below a specific threshold. Therefore, the hydrological drought prediction
280 necessitates the analysis of climate variables such as precipitation and temperature and initial catchment conditions
281 (e.g., snow cover, and soil moisture) (Hao et al., 2018).

282 In the Mediterranean ~~basin~~Basin, a common tendency for water levels to drop in shallow lakes and aquifers has
283 motivated many researchers to study the hydrological drought in this region: Greece (Myronidis et al., 2012);
284 Turkey (Akyuz et al., 2012); Tunisia (Hamdi et al., 2016); Lebanon (Al Sayah et al., 2021); Italy (Di Nunno et al.,
285 2021); Portugal (Mendes et al., 2022); Algeria (Bouabdelli et al., 2022); Syria (Mohammed et al., 2022). The most
286 common hydrological drought indices include Palmer Hydrologic Drought Index (PHDI) (Palmer, 1965), the
287 Streamflow drought index (SDI) (Nalbantis, 2008), and Standardized Runoff Index (SRI) (Shukla and Wood,
288 2008).

289 As part of the effort made by Palmer in the sixties, the PHDI has been developed by using the same two-layer soil
290 model as the PDSI, but it applies a stricter criterion for determining the ends of drought to account for long-term
291 drought events that reduce surface and groundwater supply. (Vasiliades and Loukas, (2009) tested the Palmer
292 indices in a Mediterranean basin (in Greece) they concluded that these indices were successful in the identification
293 of drought severity of historical events, but they were unable to identify drought duration.

294 The SRI is an index that uses the same computational principles as SPI but uses monthly mean streamflow rather
295 than precipitation only to account for the hydrologic process that determines seasonal lags in the influence of
296 climate on streamflow (Shukla and Wood, 2008). Shukla and Wood (2008) compared the SRI and the SPI results
297 during drought events in a snowmelt region. They concluded that the SRI can be used as a complement to the SPI
298 for depicting hydrologic aspects of drought.

299 The SDI is also a simple index that uses the cumulative monthly streamflow volumes for a given hydrological year
300 to predict wet and dry periods and identify the severity of a hydrological drought (Nalbantis, 2008). **Bouabdelli et**
301 **al., (2022) compared (2020) conducted a comparison study of the SPI and the SDI and, focusing on their**
302 **characteristics in across three watersheds in the karst area of northwestern Algeria. They found a good**
303 **agreement Their analysis revealed a substantial similarity between meteorological drought events (as represented**
304 **by SPI-12) and hydrological drought events expressed (as indicated by SPI-12 and SDI-6, respectively, which**
305 **reflects the sensitivity). This correlation emphasizes the sensitive and responsive nature of the response of a basin**

306 towards these basins to dry conditions, further illustrated by the swift transition from meteorological to
307 hydrological drought events in the studied basins (Bouabdelli et al., 2020).

308 The application of hydrological drought indices seems appears to be very useful. But valuable. However, the main
309 problem challenge in applying these indices is lies in the need requirement for a long-time-term series of climatic
310 data. According to the WMO, up to 30 years of continuous rainfall data according to the WMO suggestion may
311 be necessary for accurate drought index calculations (WMO, 1994). This condition is not always fulfilled which
312 makes the rainfall-runoff transformation a difficult task (De Luca et al., 2022). Modern hydrological models can
313 offer a valuable counterpart to existing climate-based drought indices by simulating hydrologic variables such as
314 land surface runoff (Shukla and Wood, 2008).

315 3 Overview of the physical mechanisms causing drought in the Mediterranean region

316 It is difficult to determine the physical mechanisms causing droughts in the Mediterranean basin since the region
317 covers a complex landscape with high topographic and climatic heterogeneity, strong land-sea contrasts, and high
318 anthropic pressure (De Luca et al., 2022).

319 Assuming that any type Considering the various forms of drought starts first by being meteorological, an accurate
320 droughts, characterized by a deficit in precipitation, are commonly recognized as marking the onset of drought
321 prediction conditions. This initial stage is automatically intrinsically linked to precipitation predictability, which
322 depends on is driven by large-scale atmospheric motions (such as Walker circulations and Rossby wave), forced
323 by SST anomaly waves, influenced by factors like SST anomalies, radiative forcing changes (both natural and
324 anthropogenic changes in radiative forcing), and land surface interactions (Hao et al., 2018; Wood et al., 2015).
325 However, because of due to the inherently chaotic nature of the atmospheric circulation, this predictability became
326 unreliable, particularly for meteorological droughts, tends to diminish beyond a one-month lead time. It is crucial
327 to note that the reliability of these predictions can differ when considering other drought types (such as agricultural
328 or hydrological droughts) or altering the forecast scale, with seasonal forecasts often displaying more reliability
329 months in advance, while daily forecasts may face limitations from around two weeks.

330 The discovery of teleconnections between SST anomalies and hydroclimatic phenomena constitutes a major
331 advance in drought forecasting and early warning (Wood et al., 2015). Indeed, Notably, it is widely established
332 within the scientific community gathers that some certain ocean-atmospheric teleconnections, such as ENSO, can
333 have a strong correlation with profoundly influence the onset of drought onset conditions in many various regions
334 of the world. Based on this correlation worldwide, particularly in the tropics (Ropelewski and Halpert, 1987;
335 Shabbar and Skinner, 2004; Hoell et al., 2014; Vicente-Serrano et al., 2017). For instance, during the peak phase
336 of El Niño or La Niña in the tropical Pacific, a skillful seasonal drought prediction at corresponding change in
337 precipitation patterns can be observed several months later in North American winter climate (Livezey and Smith,
338 1999; Hoerling and Kumar, 2003). This delayed impact provides a crucial window for predicting potential drought
339 conditions with a long lead time (>1 month) became possible. However exceeding one month (Johnson and Xie,
340 2010). Moreover, this lagged correlation allows for proactive drought management strategies, with the ability to
341 anticipate and prepare for drought conditions based on forecasted ENSO conditions. Nevertheless, drought
342 predictability is seasonally and spatially variable. In general, Typically, the accuracy of seasonal drought prediction

343 skill is high over ~~superior in the tropics, while it is still challenging over~~ in the extra-tropics (Tureo ~~Doblas-Reyes et~~
344 ~~al., 2017~~ 2013).

345 In the Mediterranean region MedR, the response of climate to ENSO is complex. It varies over time and depends
346 on the maturity of the ENSO state, and the co-occurrence with NAO (Kim and Raible, 2021; Brönnimann et al.,
347 2007; Mariotti et al., 2002). Although many authors have found a non-negligible correlation between ENSO and
348 precipitation anomalies in the MEDR MedR, it remains insignificant compared to the tropics (Mariotti et al., 2002).

349 ~~In contrast, many studies rather~~ the NAO is commonly identified ~~the NAO as an a~~ prominent factor influencing
350 factor in Mediterranean climate variability during the winter season (Ulbrich and Christoph, 1999; Vicente-Serrano
351 et al., 2011; Kahya, 2011; Santos et al., 2014; Cook et al., 2016). It is important to note, however, that while
352 acknowledging the profound impact of the NAO on the climate dynamics of the MedR, its predictability, especially
353 on seasonal scales, continues to be a considerable challenge in the field of climate science (Czaja and Frankignoul,
354 1999; Saunders and Qian, 2002; Scaife et al., 2014; Dunstone et al., 2016). ~~The~~

355 During the positive phase of the NAO ~~is related to,~~ below-average precipitation rates are observed over large parts
356 of the northern and western MEDR MedR. While in the negative phase of NAO, the climate is wetter and warmer
357 (Lionello, 2012). Kim and Raible, (2021) analyzed the dynamics of multi-year droughts over the western and
358 central Mediterranean for the period of 850–2099. ~~The~~ This analysis shows that droughts occur more frequently
359 during the positive NAO phase and La Niña-like conditions. This study also confirmed that suggests Mediterranean
360 droughts are from 850-1849 CE were mainly driven by the internal variability of the climate system ~~rather than,~~
361 including elements like barotropic high-pressure systems, positive NAO phases, and La Niña-like conditions.
362 Conversely, external forcing such as volcanic eruptions were found to be associated with wetter Mediterranean
363 conditions. In the period 1850-2099 CE, however, anthropogenic influences amplified land-atmosphere feedback,
364 leading to persistent dry conditions in the Mediterranean (Kim and Raible, 2021).

365 Paz et al., (2003) analyzed monthly mean Sea Level Pressure anomalies (SLP) from the 1958–1997 record over
366 the Mediterranean Basin. They identified a significant anomalous SLP oscillation between North Africa (NA) and
367 West Asia (WA) and concluded that the regional trend of the NAWA index could explain increased drought
368 processes in the eastern Mediterranean after the late '70s, in relation to northern hemispheric circulation.

369 The climate heterogeneity in the Mediterranean area may also be explained by the regional Mediterranean
370 Oscillation (MO) characterized by the opposite precipitation patterns between the eastern and western regions
371 (Düneloh and Jacobeit, 2003). More recently Redolat et al., (2019) proposed a new version of MO that uses areas
372 instead of observatories or isolated points. The new index which is referred to as the Upper-Level Mediterranean
373 Oscillation index (ULMOi) is based on the differences in geopotential height at 500 hPa to improve the
374 predictability of seasonal anomalies in the Mediterranean climate (Redolat et al., 2019). According to this study,
375 ULMOi has reported higher confidence than the MO index for rainfall predictability (Redolat et al., 2019). Other
376 teleconnections influencing the climate of MEDR MedR can be found in the reviews done by (Paz et al., (2003)
377 and (Lionello, (2012). Recent works have also shed light on the impact of Madden Julian Oscillation (MJO) on
378 water availability in the region, especially during heavy rainy episodes, see for example (Chaqdid et al., 2023)

379 ~~At the regional scale, land surface interactions from various surface conditions (e.g., soil moisture, snow cover,~~
380 ~~vegetation cover, etc.) can play a prominent role in exacerbating the drought but could also contribute to their~~

381 predictability on sub-seasonal time scales (Dirmeyer et al., 2021). Therefore, drought forecasting skill also depends
382 on the accuracy in representing these land-atmosphere processes.

383 In addition, the Mediterranean is a hotspot region that comprises a nearly enclosed sea (source of moisture and
384 heat) surrounded by highly urbanized littoral which results in complex interactions between ocean-atmosphere-
385 land processes that have a high impact on the climate and hydrological cycle, including extremes weather events
386 that frequently affect the region (Ducroeq et al., 2018).

387 The Mediterranean basin's climate is also shaped by the complex interaction of ocean-atmosphere-land processes,
388 which can significantly influence the region's hydrological cycle and contribute to droughts (Lionello et al., 2012;
389 Ducroeq et al., 2018). The nearly enclosed Mediterranean Sea serves as a substantial source of moisture and heat,
390 affecting atmospheric circulation and weather patterns (Mariotti et al., 2008). Coastal areas in the Mediterranean
391 basin experience land-sea breeze circulation due to temperature differences between land and sea surfaces
392 (Drobinski et al., 2018). This daily circulation pattern can impact the distribution of precipitation, potentially
393 leading to prolonged dry spells, especially during transitional seasons (Ducroeq et al., 2018).

394 The region's complex topography, featuring mountain ranges and valleys, gives rise to orographic effects that
395 impact precipitation patterns (Ricard et al., 2012). Orographic lifting forces moist air to rise over mountains
396 (Chaqdid et al., 2023), causing drier conditions on leeward slopes (Drobinski et al., 2016). This dynamic results in
397 localized climate conditions and can intensify drought events. In addition, the highly urbanized littoral in the
398 Mediterranean basin is subject to the urban heat island (UHI) effect, where urban areas exhibit significantly higher
399 temperatures than their rural counterparts (Santamouris, 2014). This phenomenon alters local atmospheric
400 circulation, intensifies heat waves, and exacerbates drought conditions, particularly in densely populated areas
401 (Giannakopoulos et al., 2009).

402 Land-use/cover changes driven by human activities, such as deforestation, urbanization, and agricultural
403 expansion, further influence the regional climate and hydrological cycle (Lambin et al., 2003), affecting surface
404 albedo, evapotranspiration rates, and soil moisture, ultimately altering the intensity and frequency of drought
405 events (Duveiller et al., 2018).

406 In conclusion, several complex factors that influence the predictability of drought are not yet fully understood,
407 especially those related to climate change. Therefore, more research on the physical mechanisms causing drought
408 in the MEDR_{MedR} is needed to improve the predictability of drought forecasts.

409 Expanding our grasp of the physical factors causing drought in MEDR_{MedR}, we will now delve into drought
410 forecasting models. By leveraging insights from these mechanisms, scientists have developed numerous
411 approaches and techniques including statistical, dynamical, and hybrid statistical-dynamical models to boost the
412 accuracy and trustworthiness of drought predictions.

413 **4 Statistical Drought Prediction Methods**

414 Once the major sources of predictability are identified, the task of the statistical models is to uncover the spatial
415 and/or temporal relationship between a set of these potential predictors and the predictand. When a large number
416 of predictors are identified within the same region, dimension reduction techniques like Principal Component
417 Analysis (PCA) or Linear Discriminant Analysis (LDA) can improve model accuracy and efficiency by reducing

418 the number of dimensions while preserving essential information. On the other hand, feature selection methods
419 such as decision trees or Random Forests can help eliminate irrelevant predictors. These approaches can prevent
420 overfitting, leading to enhanced model performance and interpretability (Hao et al., 2018; Ribeiro and Pires, 2016).
421 The next sections will present the frequently used data-driven models and how they were employed to predict
422 different types of droughts at different spatiotemporal resolutions in the ~~MEDR~~ **MedR**.

423 4.1. Time Series models

424 During the last few decades, several methods have been developed to analyze the stochastic characteristics of
425 hydrologic time series (Morid et al., 2007; Rafiei-Sardooi et al., 2018; Band et al., 2022; Zarei and Mahmoudi,
426 2020). Moving average (MA), Autoregressive (AR), and Autoregressive Integrated Moving Average (ARIMA)
427 are all linear models that analyze past observations of the same variable to predict its future values. Normality and
428 stationarity of observations are two of the basic assumptions of these time-series models. Therefore, if some trends
429 or seasonality are detected in observations, they should be removed before the modeling to avoid any drift in the
430 concepts to be captured.

431 ARIMA ~~and Seasonal ARIMA (SARIMA) are~~ is the most frequently used time-series ~~models-model~~ (Zhang et al.,
432 2003). The popularity of ~~these modelsthis model~~ this model is related to ~~their~~ its ability to search systematically for an
433 adequate model at each step of the model building (identification, parameter approximation, and diagnostic check).
434 This method is based on the concept that nonstationary data could be made stationary by “differencing” the series
435 (Box et al., 2015). The approach involved considering a value Y at time point t and adding/subtracting based on
436 the Y values at previous time points and adding/subtracting error terms from previous time points. The formula
437 can be written as:

$$Y_t = c + \varphi_1 Y_{t-1} + \dots + \varphi_p Y_{t-p} + \theta_1 e_{t-1} + \dots + \theta_q e_{t-q} + e_t, \quad (1)$$

438 where:

439 Y_t is the value of the variable at time t; c is a constant term; p and q are the orders of AR and MA models,
440 respectively; φ_i and θ_i are model parameters; $e_{t-1} \dots e_t$ are the error terms.

441 The AR component captures the impact of past values on the current value, the I component handles any non-
442 stationarity in the data (i.e., changes in the mean or variance over time) by “differencing” the time series, and the
443 MA component captures the impact of random shocks or errors in the data.

444 The ARIMA model is generally expressed with the three terms p, d, and q. The order of differencing in the I
445 component is denoted by the value of (d) in the ARIMA(p,d,q) notation. It represents the number of times that the
446 data must be “differenced” to produce a stationary signal. The lag order (p) represents the number of prior
447 observations having a strong correlation with the current observation. While (q) is the size of the moving window
448 and is identified by determining the number of lag errors that have a significant impact on the current observation.

449 The SARIMA is a more specific version of ARIMA that includes a seasonal component, which takes into account
450 the repeating patterns that occur at regular intervals (e.g., daily, weekly, monthly) in the data. This makes it more
451 appropriate for forecasting seasonal time series data.

(Bouznad et al., (2021) used conducted a comparative analysis of ARIMA and SARIMA models using precipitation, temperature, and evapotranspiration data to assess seasonal drought conditions in the Algerian highlands by analyzing precipitation, temperature, and ET data from 1985. These models were compared based on their ability to 2014, then by computing the aridity index, the SPI, replicate and the Normalized Difference Vegetation Index (NDVI). They identified forecast the data series accurately. The SARIMA model emerged as the best model better choice as it returned exhibited significant p-values for all the studied variables under study. This implies that the model was statistically significant in predicting the variables and thus outperformed the ARIMA model in this specific context. In the same country (Achite et al., (2022) investigated the meteorological and hydrological drought in the Wadi Ouahrane basin Basin using ARIMA and SARIMA models applied to SPI and SRI indices. A validation based on R² revealed high equality accuracy for SPI and SRI of 0.9796 and 0.5197 respectively, at 1-month lag. Additional examples of the use of the time-series model in drought forecasting in MEDR MedR can be found in Table 1.

Although time series models have shown good predictability of drought characteristics, these methods present certain limitations as they are based solely on the persistence of some drought indicators (trend, seasonality) without worrying about their interactions.

Table 1 Main studies using the Time series model to forecast drought in the MEDR MedR.

4.2. Regression analysis

Regression models are commonly applied in drought forecasting due to their straightforwardness, interpretability, and proficiency in revealing potential connections between hydroclimatic variables. These models use various predictors (independent variables), including precipitation, temperature, and other relevant climate indices, to approximate drought indices or related target variables (dependent variables).

Simple and multivariate linear regression (MLR) models have been broadly applied for projecting extreme hydrological phenomena such as droughts (Sharma et al., 2018). These models shed light on the linear connections between various predictors and predictands, offering a valuable method to understand the primary factors of drought conditions and their interactions (Mishra et al., 2011).

An MLR model that predicts drought from ~~multiple~~ multiple drought predictors X_1, X_2, \dots, X_n can be formulated as:

$$Y = \beta_0 + \beta_1 X_1 + \beta_2 X_2 + \dots + \beta_n X_n + \varepsilon \quad (2)$$

Where:

β_0 is the y-intercept or the constant term,

$\beta_{i(i=1,2,\dots,n)}$ are the regression coefficient for each independent variable $X_{i(i=1,2,\dots,n)}$,

ε is the model's error term.

On the other hand, when drought forecasts have a binary or dichotomous nature, such as drought vs. no drought, logistic regression models can be particularly useful. In these cases, the dependent variable (drought) is expressed

485 as a probability or likelihood of occurrence. The main goal of logistic regression is to estimate the relationship
486 between a set of predictors and the probability of the binary outcome (Rahali et al., 2021; Hosmer et al., 2013).
487 Some of the applications of regression analysis for drought forecasting in the MEDR MedR are discussed below
488 and summarized in (Table 2).

489 **Table 2 Main studies using regression analysis to forecast drought in the MEDR MedR.**

490 Sousa et al., (2011) analyzed the spatiotemporal evolution of drought conditions ~~in~~ across the MEDR MedR during
491 the 20th century using monthly precipitations, NAO, and SST as independent variables and scPDSI as a dependent
492 variable ~~of~~. ~~Their study successfully developed a calibrated robust stepwise regression model. A six-month lead~~
493 ~~prediction capable of predicting summer drought conditions six months in advance with a high correlation of 0.79~~
494 ~~between simulated and observed scPDSI time series, thus demonstrating its utility in forecasting future drought~~
495 ~~conditions in the region.~~ Tigkas and Tsakiris, (2015) used the MLR model with variables that include the minimum
496 temperature and RDI as the main independent variable for the assessment of drought effects on wheat yield in two
497 rural areas of Greece. The results of this analysis showed a high correlation between RDI and the wheat yield
498 during the winter months which proves that satisfactory prediction of the drought impacts on wheat ~~yield~~ yields 2
499 to 3 months before the harvest can be achieved using the MLR model. Martínez-Fernández et al., (2016)
500 ~~investigated the agricultural drought conducted a study in the REMEDHUS (Soil Moisture Measurement Stations~~
501 ~~Network) area (Spain) by computing in Spain, aiming to monitor agricultural drought on a weekly time scale and~~
502 ~~provide early warning to farmers for adapting irrigation strategies. They computed a specific agricultural drought~~
503 ~~index (SWDI),) using data from the SMOS satellite. Several~~ Within this study, various computation approaches
504 ~~have been~~ were analyzed, and the ones that yielded the most promising results were those based directly based on
505 soil attributes or parameters extracted from pedo-transfer function (PTF) ~~(~~). These approaches utilized a multiple
506 regression analysis, using the with soil water parameters as dependent variables and many incorporated other
507 relevant soil characteristics such as independent variables (texture, bulk density, and porosity) ~~), showed the best~~
508 results.

509 Although regression models have been valuable in drought forecasting, they exhibit certain limitations such as the
510 linearity assumption, limited interactions between variables, sensitivity to overfitting and multicollinearity (Rafiei-
511 Sardooi et al., 2018). Consequently, their ability to accurately represent complex real-world phenomena is often
512 insufficient (Zhang, 2003). To address these shortcomings, more advanced models capable of capturing non-linear
513 relationships and interactions are required, ultimately improving the forecasting of complex hydroclimatic events
514 such as droughts.

515 4.3. Machine Learning and Hybrid Models

516 One of the big challenges in drought prediction is the random and nonlinear nature of the hydroclimatic variables.
517 (Agana and Homaifar, 2017). Over the last two decades, intelligent techniques such as Artificial Neural networks
518 (ANN), Support Vector Machines (SVMs SVM), and Fuzzy Logic (FL) have proven to be very promising tools
519 for modeling nonlinear and dynamic time series (Mokhtarzad et al., 2017; Dikshit et al., 2022; Prodhan et al.,
520 2022). Therefore, these These algorithms have received great attention thus garnered significant interest in the
521 field realms of drought modeling and forecasting and modeling (Prodhan et al., 2022). In the context of modeling,

522 they are used to develop mathematical representations of complex drought systems, capturing the interplay of
523 various atmospheric, hydrological, and land surface processes that lead to these phenomena. In forecasting, the
524 models derived from these algorithms are employed to anticipate future drought conditions, assisting in risk
525 assessment and mitigation strategies. Table 3 highlights the key studies utilizing intelligent that utilize machine
526 learning models to predict for drought prediction in the Mediterranean region MedR.

527 Prodhon et al., (2022) stated in their review of machine learning methods for drought hazard monitoring and
528 forecasting on a global scale that the ANN was the most popular model in peer-reviewed literature, and they
529 suggested that higher use of the ANN model is anticipated because it has non-linear properties that make it more
530 robust for identifying all possible interactions between predictors.

531 ANN is a mathematical model inspired by biological brain neural networks. It consists of an interconnected group
532 of nodes (artificial neurons) and processes information using a connectionist computation (Fig. 2). In the case of
533 drought forecasting, ANN architecture is usually made of three layers: an input layer which consists of the drought
534 predictors, hidden layer(s) which comprises a function that applies weights to the input variables and passes them
535 using a non-linear activation function, and an output layer that consists of the drought target variable or drought
536 index (Han and Singh, 2020).

537 **Figure 2 Drought forecasting based on a simple ANN architecture.**

538 For the proper functioning of a neural network, the optimization of network weights (known as the learning or
539 training process) is an essential step (Dikshit et al., 2022). Back-propagation, Feed Forward, Gradient Descent,
540 Stochastic Gradient Descent, Adam and Levenberg–Marquardt are among the common training algorithms
541 (Bergou et al., 2020). The role of these algorithms is to minimize the difference between predicted and observed
542 values by adjusting the network weights and biases of the model.

543 Di Nunno et al., (2021) used a non-linear AutoRegressive with eXogenous inputs (NARX) neural network (a
544 particular type of recurrent dynamic ANNs) to predict spring flows in the Umbria region (Italy). The results of this
545 study show a good performance of the NARX model in predicting spring discharges for both short (1 month:
546 $R^2 = 0.901290 - 0.984298$, $RAE = 0.093309 - 0.255725$) and long-term lag time (12 months: $R^2 = 0.900590 -$
547 0.983898 , $RAE = 0.096309 - 0.240924$). Achour et al., (2020) also confirmed the performance of the ANN model
548 with multi-layer perceptron networks architecture and Levenberg–Marquardt calibration algorithm in predicting
549 drought in seven plains located in northwestern Algeria with 2 months lead time ($R^2=0.81$, $RMSE<0.41$ and MAE
550 <0.23).

551 SVM is also a robust supervised learning model that investigates data for classification and regression analysis. It
552 designates the best separating line to classify the data with more safety margins. Besides, the good performance in
553 solving linear problems, SVMs could also transfer a non-linear classification to a linear one using the kernel
554 function and be able to solve high-dimensional problems (El Aissaoui et al., 2021).

555 In the context of drought studies, SVM is particularly beneficial due to its ability to handle many inputs, use a
556 small dataset for training, and its resistance to overfitting compared to ANN (Hao et al., 2018). These features
557 make SVM less sensitive to data sample size, enhancing the robustness of the drought model. On the forecasting
558 aspect, SVM uses employs a kernel function to map predictors in a high-dimensional hidden space and
559 subsequently transforming the predictand to the output space (Hao et al., 2018). It can use a small data set for

560 training and can handle many inputs. Therefore, SVM is less sensitive to data sample size and less prone to
561 overfitting than ANN.

562 El Aissaoui et al., 2021). This process allows the SVM model to generate effective and accurate forecasts about
563 potential future drought events, given the input variables.

564 El Aissaoui et al. (2021) used the Support Vector Regression (SVR) model with three kernel functions (linear,
565 sigmoid, polynomial, and radial basis function [RBF]) for the prediction of drought in the region of Upper
566 Moulouya (Morocco) through the SPI and SPEI indices. ~~They have demonstrated a good performance of the
567 prediction model and~~ Their research underscores the SVR model's effectiveness, particularly with the RBF was
568 designated as the best model kernel function, in predicting the weather forecasting drought index with indices SPI
569 $R = 0.92$ for the SPI) and SPEI ($R = 0.89$ for the SPEI.) Mohammed et al., (2022) evaluated the applicability of
570 4 Machine Learning algorithms namely bagging (BG), random subspace (RSS), random tree (RT), and random
571 forest (RF) in predicting agricultural and hydrological drought events in the eastern Mediterranean region MedR
572 based on SPI. The results of this study revealed that hydrological drought (SPI-12, -24) was more severe over the
573 study area and BG was the best model in the validation stage with RMSE ≈ 0.62 – 0.83 and $r \approx 0.58$ – 0.79 .

574 To further improve the prediction accuracy of AI models, preprocessing of data using wavelet decomposition
575 (WD), PCA, or empirical mode decomposition (EMD) is recommended. These techniques known as hybrid models
576 have gained attention due to their potential to improve prediction accuracy and better capture complex
577 relationships in the data (Yoo et al., 2015; Liu et al., 2020). The preprocessing techniques are used to extract and
578 represent the essential features and patterns within the data and statistical methods, such as ANN, SVM, or RF,
579 model the relationship between the input variables and the target drought index. El Ibrahim and Baali, (2018)
580 explored the prediction of short-term (SPI-3) and long-term (SPI-12) drought conditions using 6 models: SVR,
581 ANN-MLP, Adaptive Neuro-Fuzzy Inference Systems (ANFIS), WA-SVR, WA-MLP, and WAANFIS in the
582 Saïss Plain (Morocco). They argued that ANN models were more efficient than SVR models and that the use of
583 wavelet analysis has enhanced the prediction skill of ANN models which is probably due to their capacity in
584 detecting local discontinuities and non-stationary characteristics of the data.

585 **Table 3 Main studies using Artificial Intelligence Models to forecast drought in the MEDR MedR.**

586 (Özger et al., (2020) evaluated the effect of using EMD and WD for decomposing time series data on drought
587 prediction using the self-calibrated Palmer Drought Severity Index (sc-PDSI) and machine learning models ANN
588 and SVM. They found that the accuracy of standalone machine learning models in mid-term midterm sc-PDSI
589 predictions was unsatisfactory, but it significantly improved when EMD and WD techniques were introduced,
590 particularly for hybrid wavelet models.

591 In summary, machine learning and hybrid models, which combine preprocessing techniques with statistical
592 methods, have demonstrated their efficiency in drought forecasting, as they can effectively handle intricate,
593 nonlinear relationships and adjust to a diverse range of input data characteristics. However, the applicability of
594 these models may be challenging when input variables exhibit strong dependence on each other. This dependency
595 can lead to several issues such as multicollinearity, overfitting, and diminishing returns (Maloney et al., 2012). To
596 address these limitations and improve drought forecasting performance, it is essential to consider joint probability
597 models (Madadgar et al., 2014; Hao et al., 2018).

598 4.4. Joint Probability Models

599 The probabilistic analysis of drought events plays a significant role in the planning and management of water
600 ~~resources~~ **resource** systems, particularly in arid or semi-arid Mediterranean regions known for low annual and
601 seasonal precipitation. Drought return periods, which estimate the frequency of drought events, can provide
602 valuable information for responsible water management during drought conditions. The univariate frequency
603 analysis is a common method for analyzing drought events. As mentioned above, drought is usually characterized
604 by its severity, duration, and frequency which can be extracted using the theory of runs introduced by Yevjevich
605 (1967). These characteristics present a dependence structure that can be ignored by the univariate approach,
606 resulting in an under/overestimation of drought risks. As such, several joint probability theories have been recently
607 incorporated into drought risk analysis including two or more variables. One of the most important joint probability
608 models that have garnered increasing attention in the hydrologic community over the last decade is the copula
609 model (Jehanzaib et al., 2021; Pontes Filho et al., 2020; Das et al., 2020; Zellou and Rahali, 2019; Mortuza et al.,
610 2019; Ozga-Zielinski et al., 2016; Xu et al., 2015; **AghaKouchak, 2014b**; Madadgar and Moradkhani, 2013; Chen
611 et al., 2013).

612 There are numerous copula families and classes, such as elliptic, Archimedean (Clayton, Frank, Gumbel, Joe),
613 extreme value, and Bayesian to cite but a few. The choice of the most suitable copula family depends on the
614 specific modeling goals and the structure of the data being modeled: **(Genest and Favre, 2007 ; Joe, 2014)**.

615 A brief overview of the bivariate copula theory is given here to initiate readers about their concept and application.
616 However, for additional details on the theory and concepts of the copula, readers may refer to the monographs by
617 Joe (1997) and Nelsen (2007). ~~For the construction~~ **Furthermore, comprehensive methodological understanding of**
618 **constructing high-dimensional copulas, such as pair-copula-construction** **Pair Copula Construction (PCC) and**
619 **nested** **Nested Archimedean construction** **Construction (NAC), readers may refer to can be garnered from the works**
620 **of Aas and Berg (2009) and Savu and Trede (2010).**

621 Let F be a 2-dimensional distribution function, with univariate margins F_1 and F_2 for random variables U and V ,
622 respectively. According to Sklar's theorem (Sklar, 1959), there exists a copula C such that:

$$F(U, V) = C(F_1(U), F_2(V)) \quad U, V \in R \quad (3)$$

623 with C unique when $F_1(U)$ and $F_2(V)$ are continuous marginal distributions, so that

624 $C: [0,1]^2 \rightarrow [0,1]$ that satisfies the boundary conditions $C(u, 0) = C(0, v) = 0$

625 and $C(u, 1) = C(1, u) = u$ (Uniform margins) for any $u \in [0,1]$ and the so-called 2-increasing property
626 (Papaioannou et al., 2016).

627 The main advantage of the copula over the traditional multivariate distributions is its ability to model the nonlinear
628 dependence structure between variables independently from the choice of their marginal distributions (Salvadori
629 and De Michele, 2004). This concept simplifies the joint probability analysis and its application in high dimensions
630 (with a large number of variables or predictors) becomes possible.

631 Serinaldi et al. (2009) constructed a four-dimensional joint distribution using the copula approach and SPI to model
632 the stochastic structure of drought variables in Sicily (Italy). Drought return periods were next computed as mean

633 interarrival time, taking into account two drought characteristics at a time by means of the corresponding bivariate
634 marginals of the fitted four-dimensional distribution. Bouabdelli et al. (2020) investigated the joint probability and
635 joint return period of drought severity and duration using copula theory to assess the hydrological drought risk in
636 the reference period and its probability of occurrence in the future under two climate change scenarios in three
637 basins located in northern Algeria. Bonaccorso et al. (2015) evaluated the conditional probability of future SPI
638 classes under the hypothesis of multivariate normal distribution of NAO and SPI series in Sicily (Italy). The results
639 of this study indicated that transition probabilities toward equal or worse drought conditions increase as NAO
640 tends toward extremely positive values. Table 4 displays additional examples of the application of the Joint
641 Probability Models to forecast drought in the **MEDR** **MedR**.

642 **Table 4 Main studies using Joint Probability Models to forecast drought in the **MEDR** **MedR**.**

643 All the above-mentioned studies confirm that copulas can accurately capture the joint distribution and dependence
644 structure between multiple drought predictors without making strong assumptions about their marginal
645 distributions. By combining the strengths of machine learning models with the flexibility of copulas, researchers
646 can develop more accurate and reliable hybrid methods that better represent the intricacies of hydrological
647 processes and climatic variables, even in the presence of strong dependence among the input variables (Jiang et
648 al., 2023; Li et al., 2022; Wu et al., 2022; Zhu et al., 2020).

649 **4.5. Ensemble Streamflow Prediction**

650 The ESP method, a commonly used technique in hydrological forecasting, was primarily intended for medium to
651 long-term streamflow prediction (Day, 1985). However, its utility extends to the prediction of hydrological
652 droughts, characterized by low streamflows (KyungHwan and DegHyo, 2015 ; Sutanto et al., 2020 ; Troin et al.,
653 2021).

654 ESP operates on the principle of employing historical data to generate an ensemble of possible future climate
655 conditions (Turco et al., 2017b). The process begins by determining the current state of the system, considering
656 parameters such as current streamflow, soil moisture levels, and reservoir levels which serves as the initial
657 conditions for the forecast (Wood et al., 2016). The generation of the ensemble involves choosing a historical
658 record at each time (day, week or month) of forecast that will provide the meteorological inputs (Day, 1985). By
659 repeating this process for every time in the historical record, an ensemble of forecasts is produced, each member
660 representing a potential future scenario. The hydrological model is run for each ensemble member, using the
661 chosen meteorological inputs and initial conditions to generate a range of potential future states of the system
662 (Harrigan et al., 2018). The ensemble of forecasts is then analyzed to derive probabilistic predictions.

663 As new data becomes available, forecasts can be updated by re-initializing the system's state and generating a new
664 ensemble of forecasts. A significant advantage of this method is that it enables the uncertainty prediction by
665 producing a variety of potential future streamflow forecast scenarios which can increase the confidence of this
666 approach, specifically for its operational use in water management (Troin et al., 2021).

667 However, the limitations of the ESP method must be noted. For instance, it presupposes that future behavior will
668 mirror past behavior, a concept that may not hold under changing climatic conditions (Wood et al., 2016).

669 Furthermore, the method's performance is heavily reliant on the quality and duration of the historical
670 meteorological records used in the ensemble generation process (Turco et al., 2017b).

671 ESP is frequently employed as a benchmark for comparison with more sophisticated forecasting methods, such as
672 dynamical climate models or hybrid statistical-dynamical models (AghaKouchak, 2014a; Turco et al., 2017b;
673 Torres-Vázquez et al., 2023). Although these more complex methods can outperform ESP in some instances, the
674 computationally efficient ESP method often exhibits comparable performance, particularly when forecasting a few
675 months ahead (Turco et al., 2017b; Torres-Vázquez et al., 2023).

676 4.5.4.6. Markov Chain Models

677 Unlike some regions of the world, subjected to well-known phenomena like ENSO (e.g., tropical regions), the
678 governing factors of drought are not clearly identified in the MEDR. Consequently, drought prediction becomes a
679 challenging task, particularly on seasonal and longer time scales. The stochastic analysis of drought episodes may
680 then be a promising alternative to handle this issue. Markov chains are effective tools to understand the stochastic
681 characteristics of drought events and their temporal dependency. These models ~~are based on the assumption~~ assume
682 that future states depend only on the current state.

683 Mathematically, Markov chain is a stochastic process X , such as at any time t , X_{t+1} is conditionally independent
684 from $X_0, X_1, X_2, \dots, X_{t-1}$, given X_t ; the probability that X_{t+1} takes a particular value j depends on the past only
685 through its most recent value X_t :

$$P\{X_{t+1} = j | X_0, X_1, \dots, X_t\} = P\{X_{t+1} = j | X_t = i\} \forall i, j \in S, t \in T \quad (4)$$

686 A Markov chain is characterized by a set of states, S , and by the transition probability, P_{ij} , between states. The
687 transition probability P_{ij} is the probability that the Markov chain is at the next time point in state j , given that it is
688 at the present time point in state i .

689 The drought prediction using this concept can be expressed as the transition from wet or normal state to dry state
690 (or the other way around) or the transition from one drought severity state to another (e.g., no drought, mild
691 drought, moderate drought, extreme drought). Habibi et al. (2018) studied meteorological drought in North
692 Algeria's Chélif-Zahrez basin, employing both localized and spatially-distributed probabilities for temporal
693 transitions using Markov Chains, and recurrence probabilities using an optimal time series model, the APARCH
694 approach. Paulo and Pereira (2007) used Markov chains, incorporating homogeneous and non-homogeneous
695 formulations, to predict drought transitions up to three months ahead, based on the SPI derived from 67 years of
696 data in Southern Portugal. The non-homogeneous Markov model outperformed its counterpart by considering the
697 initial month and seasonal rainfall variations. Table 5 lists additional studies that apply Markov chain models for
698 MEDR MedR drought forecasting.

699 **Table 5 Main studies using Markov Chains Model to forecast drought in the MEDR MedR.**

700 These studies generally support the effectiveness of Markov chain models in providing valuable drought insights.
701 However, it is essential to consider the challenges associated with applying Markov chains within the
702 MEDR MedR, as the region's complex topography, considerable interannual climate fluctuations, limited data
703 availability, and the non-stationarity resulting from climate change can adversely affect the models' core

704 assumptions and constrain their long-term forecasting accuracy. Addressing these challenges calls for the adoption
705 of more sophisticated techniques that encompass both stochastic and physically-based approaches, ultimately
706 enhancing the accuracy and reliability of drought predictions in this region (Paulo and Pereira, 2007).

707 5 Dynamical Drought Prediction Methods

708 ~~Future~~ Dynamical drought projections and near-real-time prediction methods are ~~challenging since several relevant~~
709 ~~variables and complex processes contribute to the occurrence and severity of this phenomenon (Balting~~ generally
710 based on the use of seasonal climate forecasts derived from comprehensive GCMs. The European Centre for
711 Medium-Range Weather Forecasts (ECMWF)'s System 4 (SYS4), the Hadley Centre's Global Environmental
712 Model (HadGEM), the Community Earth System Model (CESM), and the National Centers for Environmental
713 Prediction (NCEP)'s Climate Forecast System (CFS) are some widely recognized examples. Designed to emulate
714 ~~et al., 2021).~~ The dynamical drought prediction is frequently based on GCMs. These models can represent the
715 physical processes in ~~across~~ the atmosphere, ocean, and land surface ~~and project future climate changes under~~
716 ~~different scenarios to provide estimates of climate variables.~~ these GCMs can produce near-term forecasts for
717 various climatic factors ~~such as precipitation, temperature, surface pressure, and winds~~ on a global scale.
718 ~~However, GCMs have generally quite~~ these models typically provide a global overview and possess a relatively
719 coarse resolution ~~relative to the scale of exposure units in most impact assessments with a horizontal resolution~~
720 ~~varying between~~ which spans from 150 and km to 300 km, km horizontally, encompassing 10 to 20 vertical
721 atmospheric layers in the atmosphere, and up to 30 oceanic layers in ~~This level of detail may not offer the oceans.~~
722 ~~Therefore~~ specificity necessary for local-scale impact assessments. To counter this, post-processing including ~~steps.~~
723 encompassing downscaling and bias correction ~~is often an essential step before using~~ are crucial when employing
724 GCM forecasts in practice (Tuel et al., 2021; Gumus et al., 2023). The goal of this step is to provide high-
725 resolution climate projections for impact studies on ~~main objective here is to refine the global, coarse-grained GCM~~
726 data into higher-resolution forecasts. These refined forecasts are far more pertinent for predicting seasonal drought
727 events at a regional and local ~~scales.~~ scale within the MedR.

728 The most common approaches to downscale GCM forecasts include statistical models, dynamic or nested models,
729 and hybrid statistical-dynamical models (Wilby et al., 2004). In statistical downscaling, large-scale variables are
730 used as the predictors and desired near-surface climate variables are the predictands (Gutiérrez et al., 2019). The
731 role of statistical models is then to measure the correlations between predictors and predictands. Whereas
732 dynamical downscaling refers to the use of high-resolution regional simulations to dynamically extrapolate the
733 effects of large-scale climate processes to regional or local scales based on a nesting approach between GCMs and
734 Regional Climate Models (RCMs) (Giorgi and Gutowski, 2015). However, it is known that GCMs contain
735 significant systematic biases that may propagate into RCMs through the lateral and lower boundary conditions
736 and thus degrade the dynamically downscaled simulations and lead to large uncertainties (Maraun, 2016). Besides,
737 climate predictions from a single climate model simulation are sensitive to initial oceanic and atmospheric states
738 and can represent only one of the possible pathways the climate system might follow.

5.1. Multi-Model Ensemble

To allow probabilistic estimates of climate variables with uncertainties in quantification, it is necessary to carry out an ensemble of simulations with different initial conditions from each model and to combine various models as ensemble members. The frequently used Multi-Model Ensemble (MME) and bias correction methods include quantile mapping (Wood et al., 2002) and Bayesian Model Averaging (Krishnamurti et al., 1999; Seifi et al., 2022). These methods proceed by adjusting the modeled mean, variance, and/or higher moments of the distribution of climate variables, to match the observations. However, such MME simulations can be very computationally demanding. Therefore, some international dynamical downscaling intercomparison projects were carried out such as the Coordinated Regional Downscaling Experiment (CORDEX, Wilby et al., 1998) and its Mediterranean initiative MEdCORDEX (Ruti et al., 2016) to provide present and future climate simulations with a high spatial resolution (~~~12km~~; 12 km). In a study conducted by Turco et al. (2017b), the accuracy and reliability of ECMWF's System 4 (SYS4) in forecasting drought conditions, characterized by a six-month SPEI6, across Europe from 1981 to 2010 was evaluated. They found that the SYS4 model effectively projected the spatial patterns of SPEI6 and various drought conditions (ranging from extreme to normal) with a reasonable degree of precision up to a lead time of 2 months. In the same geographical context, Ceglar et al. (2017) demonstrate the power of dynamical models in the agricultural sector by investigating the relationship between large-scale atmospheric circulation and crop yields in Europe. Their research highlights the significant potential of such models in developing effective seasonal crop yield forecasting, and consequently, in advancing dynamic adaptation strategies to climate variability and change.

~~Baronetti et al. (2022) analyzed the expected characteristics of drought episodes in the near (2021–2050) and far (2071–2100) future compared to the baseline conditions (1971–2000) for northern Italy using EURO-CORDEX and MEdCORDEX GCMs/RCMs pairs at a spatial resolution of 0.11 degrees for the Representative Concentration Pathways (RCPs) (4.5 and 8.5) scenarios. The results indicated that the GCM/RCM pairs performed generally well, while in complex environments such as coastal areas and mountain regions, the simulations were affected by considerable uncertainty. Dubrovský et al. (2014) used an ensemble of 16 GCMs to map future drought and climate variability in the Mediterranean region. Bağcaei et al. (2021) compared the capacity of the latest release Coupled Model Intercomparison Project Phase 6 (CMIP6) model ensembles in representing near-surface temperature and precipitation of Turkey in comparison with its predecessor CMIP5 to better understand the vulnerability degree of the country to climate change. All these studies confirmed the good performance of MME methods in providing probabilistic drought forecasts for 1 to 2 months of lead time, and improving drought onset detectability. However, much effort should be made in selecting the most skilled GCM ensembles in reproducing the large and synoptic scale atmospheric and land-surface conditions associated with drought development in the MEDR. MedR. By prioritizing ensembles that adequately capture the region's distinct climate characteristics, spatial-temporal variability, and land-atmosphere interactions, the MME forecasts can mitigate biases related to key meteorological variables such as temperature or precipitation and significantly improve the precision and reliability of drought predictions (Li et al., 2023; Ahmed et al., 2019).~~

5.2. Coupled hydrological models.

~~On the other hand, GCMs often struggle to accurately represent some drought-relevant variables complex elements of the hydrological cycle, such as soil moisture, streamflow, groundwater level, and PET, which are integral parts~~

of the hydrological cycle, are not necessarily well represented. The inherent complexities of these variables and the broad spatial scale of GCMs make it challenging to fully capture their behavior. This gap can limit the effectiveness of GCMs in the GCMs drought prediction and modelling (Balting et al., 2021). Consequently, to dynamically forecast agricultural and hydrological droughts, the water balance should be correctly simulated by hydrological models forced by climate forecasts (Wanders and Wood, 2016). Among the most used models to forecast hydrological drought, we cite, the Soil and Water Assessment Tool (SWAT) (Arnold et al., 1998), the Variable Infiltration Capacity (VIC) (Liang et al., 1994), and the Community Land Model (CLM) (Oleson et al., 2004). These models can incorporate data on soil moisture, vegetation, snow water equivalent, groundwater level, and other initial hydrologic conditions with climate forecasts to simulate the movement of water through the hydrological cycle, including the processes of precipitation, evaporation, infiltration, and runoff. Crop growth models can also be coupled with hydrological models to make an accurate prediction of agricultural drought and its impact on crop yields- (Narasimhan and Srinivasan, 2005; Abhishek et al., 2021).

Coupled hydroclimatic models can improve drought forecasting by allowing for the consideration of feedback between the hydrological and climatological components of the Earth system. Indeed, drought conditions can affect the availability of water for evapotranspiration, which in turn can affect the amount of moisture in the atmosphere and the likelihood of precipitation. By incorporating this feedback into the model, it is possible to produce more accurate forecasts of drought conditions.

In a recent study, Brouziyne et al. (2020) combined meteorological and hydrological drought indices (SPI and SDI) with a SWAT model forced by bias-corrected CNRM-CM5 data to predict future droughts under two RCPs (4.5 & 8.5) in Bouregreg watershed, Morocco. They confirmed that using multiple drought indices and a comprehensive water budget indicator such as Total Water Yield provided a valid approach to evaluate drought conditions in a Mediterranean context. Marx et al. (2018) analyzed a multi-model ensemble of 45 hydrological simulations based on three RCPs (2.6, 6.0, and 8.5), five GCMs (CMIP5), and three state-of-the-art hydrological models (mHM, Noah-MP, and PCR-GLOBWB) to investigate how hydrological low flows are affected under different levels of future global warming. Based on the analysis of the results, the authors recommended using multiple hydrological models in climate impact studies and to embrace uncertain information on the multi-model ensemble as well as its single members in the adaptation process.

5.3. Long-term drought projection under climate change.

As climate change continues to influence drought events in the MedR, it is vital to integrate long-term climate projections into drought forecasting strategies (Tramblay et al., 2020). In this regard, GCMs are essential for projecting future climate changes under varying scenarios, such as Representative Concentration Pathways (RCPs) or Shared Socioeconomic Pathways (SSPs¹). Coupled with downscaling techniques, these models offer region-specific projections of critical climate variables including precipitation, temperature, surface pressure, and winds. These projections are instrumental in estimating long-term drought events, facilitating a more comprehensive risk assessment for stakeholders and decision makers. Baronetti et al. (2022) analyzed the expected characteristics of drought episodes in the near (2021–2050) and far (2071–2100) future compared to the baseline conditions (1971–

¹ SSPs are the latest climate change scenarios used in CMIP6. They not only incorporate greenhouse gas emissions scenarios like their predecessor, RCPs from CMIP5, but also integrate socioeconomic factors, such as population growth, economic development, and technological progress. Essentially, SSPs provide a more holistic view of possible future climate scenarios by considering both environmental and societal changes.

2000) for northern Italy using EURO-CORDEX and MedCORDEX GCMs/RCMs pairs at a spatial resolution of 0.11 degrees for the RCPs (4.5 and 8.5) scenarios. The results indicated that the GCM/RCM pairs performed generally well, while in complex environments such as coastal areas and mountain regions, the simulations were affected by considerable uncertainty. Dubrovský et al. (2014) used an ensemble of 16 GCMs to map future drought and climate variability in the MedR. Bağçacı et al. (2021) compared the capacity of the latest release Coupled Model Intercomparison Project Phase 6 (CMIP6) model ensembles in representing near-surface temperature and precipitation of Turkey in comparison with its predecessor CMIP5 to better understand the vulnerability degree of the country to climate change. In a study conducted by Cos et al. (2022), the authors compared climate projections from CMIP5 and CMIP6 models to assess the impacts of climate change in the MedR. The findings reveal a robust and significant warming trend across all seasons, with CMIP6 models projecting stronger warming compared to CMIP5. While precipitation changes show greater uncertainties, a robust and significant decline is projected over large parts of the region during summer by the end of the century, particularly under high emission scenarios. (Seker and Gumus, 2022) uses 22 global circulation models from CMIP6 to project future precipitation and temperature changes in the MedR. The MMEs outperform individual GCMs in simulating historical data, and the projections indicate a decrease in precipitation by 15% for SSP2–4.5 and 20% for SSP5–8.5. Table 6 shows the main studies using dynamical models to forecast drought in MedR.

Table 6 Main studies using dynamical models to forecast drought in the MEDR MedR.

In summary, significant recent advancements in recent years have led to improvements in the accuracy and reliability of dynamical seasonal drought forecasting. Key developments include higher resolution with dynamical models encompass increased climate resolution climate models, enhanced process, improved representation through advanced land surface and hydrological models, and the implementation of physical processes, improved initialization methods using data assimilation techniques to better incorporate observed data (Liu et al., 2020). Additionally, the adoption of ensemble forecasting methods has improved the assessment of forecast uncertainty (Zhou et al., 2022), use of multi-model ensembles (Wanders and Wood, 2016; Seker and Gumus, 2022), while the integration of coupled climate models has captured the influence of large scale climate patterns on regional drought conditions (Guion et al., 2022). However, they still have some limitations related to computational coupled modeling approaches (Guion et al., 2022), and the development of sub-seasonal to seasonal predictions (Zhou et al., 2021). These steps have contributed to more accurate and reliable drought predictions. However, even with these improvements, predicting drought months in advance remains a significant challenge due to the inherent complexity, data requirements, and reduced skill at longer lead times and chaos of the climate system.

6 Hybrid Statistical-Dynamic Methods

6.1 While statistical dynamical methods

As mentioned above the major limitations of statistical models are related to models, when appropriately fine-tuned, can effectively predict seasonal drought events, a significant limitation arises from the non-stationary relationship between the predictors and predictands and predictors used to forecast drought. Statistical models do not consider climate changes used in the forecasting process (AghaKouchak et al., 2022). This can limit their ability to accurately predict unprecedented drought events, which fall beyond the scope of their historical training data

852 (Hao et al., 2018). means that they may not be able to adequately forecast drought events that have not occurred
853 in the past. While On the other hand, dynamical models can integrate climate change signals through some Shared
854 Socioeconomic Pathways (SSPs) scenarios and can capture are proficient at capturing the nonlinear interactions
855 in among the atmosphere, land, and ocean, their forecast skill is still limited for a long enhancing their ability to
856 detect the onset of droughts (Turco et al., 2017b; Ceglar et al., 2017). However, despite their advanced capabilities,
857 their forecast proficiency is generally constrained to a few months of lead time due to the inherent uncertainty in
858 predicting future events. (Turco et al., 2017b). To address the shortcomings associated with seasonal forecasting
859 skills, hybrid models employ statistical or machine learning methods to merge a broad variety of forecasts from
860 statistical and dynamical models into a final probabilistic prediction product (Slater et al., 2022). The frequently
861 used merging methods include the regression analysis, BMA, and Bayesian post-processing method (Hao et al.,
862 2018; Strazzo et al., 2019; Han and Singh, 2020; Xu et al., 2018). The BMA method involves the estimation of
863 the posterior probability density function (PDF) of model parameters based on the observed data and using this
864 PDF to weight each individual model forecast (Tian et al., 2023). The hybrid forecast is then generated as the
865 weighted average of the individual forecasts from statistical and dynamical models. The BMA weights estimation
866 with simultaneous model uncertainty quantification can also be used in selecting the best-performing ensemble
867 members to reduce the cost of running large ensembles (Raftery et al., 2005). There is also an opportunity to
868 enhance the probabilistic seasonal forecast skill through Bayesian post-processing methods such as the Calibration,
869 Bridging, and Merging (CBaM) technique (Schepen et al., 2014; Schepen et al., 2016; Strazzo et al., 2019). The
870 calibration step consists in optimizing the dynamical forecasts from multiple GCMs by analyzing their correlation
871 to observed data through a statistical model. In the bridging step, the dynamical forecasts from GCMs are calibrated
872 using some large-scale climate indices (e.g., ENSO, NAO, PDO, AO), and finally, the merging component
873 combines the forecasts of the two previous steps.

874 These hybrid statistical-dynamical models combine the strengths of both modeling approaches and offer several
875 advantages compared to either statistical or dynamical models alone. Thereby, seasonal drought forecasting using
876 hybrid models has recently become an active area of research (Madadgar et al., 2016; Strazzo et al., 2019;
877 AghaKouchak et al., 2022). In On global scale, Yan and Wood (2013) analyzed the MEDR, capability of seasonal
878 forecasting of global drought onset and found that despite climate models increasing drought detection, a
879 significant proportion of onset events are still missed. Their findings underscore the urgent need for implementing
880 reliable, skillful probabilistic forecasting methods to better manage the inherent uncertainties and potentially
881 improve drought predictability. Dutra et al. (2014) confirmed that the uncertainty in long lead time forecasts
882 suggested that drought onset might fundamentally be a stochastic problem. Mo and Lyon (2015) also found that
883 improvements in near-real-time global precipitation observations could yield the most substantial advances in
884 global meteorological drought prediction in the near term. This reinforces the notion that the effectiveness of
885 dynamical models is fundamentally associated to the quality of initial data and the inherent stochastic nature of
886 drought onset.

887 In line with these findings, a unique approach was undertaken by Ribeiro, and Pires, (2016) in the MedR. They
888 proposed a two-step hybrid scheme combining that combines dynamical model forecasts from the UK Met Office
889 (UKMO) operational forecasting system and with past observations as predictors on in a statistical downscaling
890 approach based on MLR models to forecast model for long-range regional drought index SPI forecasting in Portugal

(Table 7). They concluded that hybridization improves drought forecasting skills in comparison to purely dynamical forecasts.

Table 7 Main studies using hybrid statistical-dynamical models to forecast drought in the MEDR.

Moreover, the Leveraging these advantages of hybrid statistical-dynamical models make the prediction of flash droughts has become possible. Indeed, these events can develop rapidly by a quick decline in soil moisture and streamflow that may cause devastating economic and ecological impacts in a short period (from a few days to 1–2 months) (Mo and Lettenmaier, 2015) which makes them, particularly challenging to forecast. By providing a more nuanced understanding of the drought contributing factors, hybrid statistical-dynamical models help to identify potential warning signs of an imminent drought event, improve drought early warning systems, and reduce false alarm rate of drought onset (Xu et al., 2018), thus tackling some of the limitations and challenges highlighted in the earlier studies.

7 Discussion

7.1. Drought types and indices

The indices adopted by the surveyed studies were grouped according to three distinct drought categories: meteorological, agricultural, and hydrological. Figure 3 illustrates the percentage of usage for each index by category. Meteorological droughts were the most common, appearing in 63.00% of the examined studies, followed by agricultural droughts with approximately 22.20%, whereas hydrological droughts were the least prevalent, making up only 14.80%.

The SPI was the primary indicator, used in 70.59% of meteorological drought studies. But it also served as an indicator for hydrological and agricultural droughts, with usage rates of around 25% and 8.33%, respectively.

Despite the apparent versatility of the SPI, its reliance on precipitation data limits its ability to account for other influential factors such as evapotranspiration, soil moisture, land usage, and water management practices. Consequently, an overemphasis on the SPI could potentially constrain our comprehension of drought phenomena in the MedR. To enrich this understanding, it is recommended to incorporate a broader range of indicators and models that include a more diverse set of variables. By fitting a probability distribution to observed precipitation data, the SPI is calculated and subsequently transformed into a standard normal distribution with a mean of 0 and a standard deviation of 1 (Livada and Assimakopoulos, 2007). Consequently, SPI values can be compared across various regions and timeframes (e.g., 1, 3, 6, 12, or 24 months). This multiscale nature of SPI enables it to capture diverse aspects of drought depending on the selected time scale. The shorter time scales (1–3 months) are suitable for monitoring agricultural drought, while longer time scales (6–12 months or more) are better suited for evaluating hydrological drought. It is important to recognize, however, that the SPI does not consider other factors influencing drought, such as evapotranspiration, soil moisture, land use, and water management practices. In regions with high temperatures and evapotranspiration rates like the Mediterranean, the SPI may not offer a comprehensive assessment of drought conditions.

Using multivariate drought indices like such as the SPEI, PDSI, and/or sc-PDSI, or alternatively, a combination of multiple indices, can help account for contribute to a more comprehensive view by including regional feedback

928 mechanisms in the forecast process ~~and better assess the impact~~. This approach also enhances our capacity to
929 evaluate the impacts of global warming on drought severity and intensity in MEDR (the MedR (see Marcos-Garcia
930 et al., 2017; Gouveia et al., 2017)).

931 Figure 3 Pie chart showing the proportion of use of indices in the MEDR surveyed studies in MedR for different drought
932 types.

933 On the other hand, SDI was the most applied index in hydrological drought studies in the MEDR MedR (37.550%).
934 It is calculated by comparing the current streamflow to the long-term average or median streamflow for a specific
935 location and time of year (Nalbantis & Tsakiris, 2009). Despite its usefulness, there are some limits to using SDI
936 in MEDR MedR. Indeed, this region is known for highly variable climates with strong seasonality (wet winters and
937 dry summers) and the presence of transient streams or intermittent rivers that flow only during and after rainfall
938 events, especially in sub-humid and semi-arid areas. Groundwater recharge principally occurs during the wet
939 season, when precipitation infiltrates the soil and replenishes aquifers (Scanlon et al., 2002). In these regions, the
940 SDI may not provide an accurate representation of the hydrological drought as it relies solely on streamflow data.
941 Therefore, the use of SDI should be done in combination with other drought indices that consider variables such
942 as groundwater, soil moisture, runoff, and regional variations in precipitation and streamflow patterns for accurate
943 hydrological drought assessment.

944 One can notice from Fig. 3 that the agricultural drought studies are characterized by more diversity of indices. This
945 diversity can be explained by the varied range of agro-climatic conditions that characterize the MEDR MedR,
946 including a wide range of soil types, topography, and vegetation cover. These diverse conditions can result in
947 varying impacts of drought on agricultural production, which require different drought indices to accurately
948 capture the extent and severity of the drought. In addition, the MEDR MedR is also home to a diverse range of
949 crops, each with different sensitivities to drought (Feres & Soriano, 2007). This diversity of crops can require
950 different indices to assess the impact of drought on each crop.

951 Overall, a suitable index should be able to capture the impacts of drought, detect changes over time, and
952 differentiate between different levels of severity, while also being accurate and easily interpretable by stakeholders.

953 7.2. Drought forecasting accuracy

954 Key obstacles in drought modeling include the absence of a one-size-fits-all model, choosing suitable inputs,
955 determining an index that accurately represents drought tracking in various regions, and the uneven geographical
956 influence that leads to discrepancies in model accuracy (Mishra & Desai, 2005; IPCC, 2012). Consequently,
957 contrasting different methodologies is crucial for developing a reliable prediction model.

958 The accuracy of drought prediction depends on various factors such as the quality and availability of data, spatial
959 and temporal scales, prediction lead time, and model complexity, to cite but a few (Wilhite et al., 2014; Mishra &
960 Singh, 2010). For consistency, this analysis only includes studies that use R^2 as evaluation criteria of the forecast
961 with a lead time of 1 month. Joint probability models were excluded from this analysis since the accuracy
962 evaluation criteria were different. Moreover, the concept of lead time is not addressed in the majority most of the
963 surveyed studies. It is also important to note that this analysis does not include hybrid statistical-dynamical models,
964 as the number of studies applying this approach in the MEDR MedR was quite limited. Consequently, the available

965 research is insufficient to offer a comprehensive understanding of the applicability and effectiveness of these
966 models in the region.

967 **Figure 4 Box and whiskers plot ~~showing to show~~ the performance of drought prediction models denoted by the**
968 **coefficient of determination (R^2) for the surveyed studies in MEDR.**

969 ~~Figure 4 shows a box and whisker plot of drought forecasting model accuracy based on R^2 in the surveyed studies~~
970 ~~in the MEDR. **MedR. The lower box shows the 25th percentile, the upper box shows the 75 percentile and the median**~~
971 ~~**(50th percentile) is represented by the black line inside the box. The whiskers show the extent to the minimum and**~~
972 ~~**maximum values within 1.5 times the interquartile range (IQR) from the box.**~~

973 ~~Figure 4 shows a box and whisker plot of drought forecasting model accuracy based on R^2 in the surveyed studies~~
974 ~~in the **MedR** (see table 1 in Appendix).~~ According to the graph, hybrid models appear to be the most accurate and
975 consistent, with the highest median and shortest box height. Markov chains and AI models also have relatively
976 short box heights, indicating high agreement and accuracy across studies. Meanwhile, dynamical and regression
977 models exhibit moderate to high accuracy (both have median equal to 0.79), but the height of the dynamical model
978 box is shorter than that of the regression models, suggesting greater consistency. Time series models also show
979 moderate to high accuracy, with a median equal to 0.82.

980 Nonetheless, Fig. 4 provides valuable information about the relative performance of different models across
981 multiple studies in the ~~MEDR~~ **MedR**. The consistently high median of hybrid models suggests that they are
982 particularly effective for drought forecasting in the region. Similarly, the consistent performance of the AI and
983 Markov chain models, suggests that these models also show promise. The variability in the performance of the
984 regression, and the time series, as indicated by their taller boxplots, suggests that there may be more variability in
985 the effectiveness of these models across different studies and regions. The results also show that dynamical models
986 can provide valuable insights into drought conditions. However, the high variability in their performance, suggests
987 that there may be room for improvement in the development and implementation of these models in ~~MEDR~~ **MedR**.

988 This analysis concludes that simple statistical models such as Markov chains, regression, and time series can still
989 be useful in some situations and are generally more transparent and easier to interpret. For example, when focusing
990 on a single variable to forecast drought (e.g., precipitation using SPI), simple models like ARIMA can effectively
991 capture the temporal patterns and provide reasonable forecasts. Or, when drought conditions can be effectively
992 represented by discrete states or categories, Markov chains can be employed to model the transition probabilities
993 between these states and forecast future drought conditions (Habibi et al., 2018; Nalbantis and Tsakiris, 2009;
994 Paulo and Pereira, 2007). Also, when working with a limited number of variables and moderate interactions,
995 simple regression models like linear or logistic regression can provide adequate predictions of drought conditions
996 (Sharma et al., 2017). The effectiveness of simple models in these situations depends on the specific context and
997 the data quality and quantity. When more complex relationships or high-dimensional data are involved, it may be
998 necessary to employ more advanced models like dynamical models or combine simple models with techniques
999 like machine learning, copulas, or hybrid approaches to improve forecasting performance. Hybrid statistical-
1000 dynamical models present a promising avenue for enhancing forecast accuracy, particularly for extended lead
1001 times and in situations where intricate processes and interactions are critical (AghaKouchak et al., 2021; Mehran
1002 et al., 2020; Madadgar et al., 2016). The relatively nascent emergence of these hybrid techniques has resulted in a
1003 limited number of studies applying them in the ~~MEDR~~ **MedR**. This can be ascribed to factors such as data

1004 constraints, computational complexity, and model uncertainty. Moreover, proficiency in both statistical and
1005 dynamical modeling is needed, and interdisciplinary cooperation is frequently deficient. **Notwithstanding these**
1006 **challenges, there is an increasing interest not only in refining drought forecasting abilities, with enhancing**
1007 **traditional dynamical models but also in the prospect of wider adoption development and utilization of hybrid**
1008 **models as. As research advances progresses and resources become more accessible, these hybrid models may see**
1009 **wider adoption for their potential to improve predictive accuracy.**

1010 7.3. Spatial and Temporal Scales of Drought

1011 Figure 5 displays the spatial and temporal scales of drought forecasting studies in the **MEDR MedR** with a pie chart
1012 indicating the percentage of use of drought forecasting method: statistical, dynamical, and hybrid statistical models
1013 for each spatiotemporal scale. This figure shows that the number of droughts forecasting studies tends to decrease
1014 as the spatial scale increases and increases as the time scale increases. We can also notice from this figure that the
1015 majority of studies in the **MEDR MedR** focused on the local scales (e.g., city or catchment), particularly at annual
1016 and seasonal time scales. In contrast, very few studies were conducted at the **MEDR MedR** scale, and only a few
1017 studies were conducted at the country scale.

1018 **Figure 5 Spatial and temporal scales of drought forecasting studies in the Mediterranean region with a pie chart**
1019 **indicating the percentage of use of drought forecasting method: statistical, dynamical, and hybrid statistical models for**
1020 **each spatiotemporal scale.**

1021 When considering the spatial scale, drought forecasting becomes more challenging at larger scales due to various
1022 factors. One of the major challenges is the complexity of the interactions between different factors that contribute
1023 to droughts, such as precipitation, temperature, soil moisture, and vegetation cover (Sheffield & Wood, 2011).
1024 These interactions are nonlinear and difficult to capture accurately, especially at larger scales where there are more
1025 variability and heterogeneity (AghaKouchak et al., 2015). For instance, at the country scale, there could be
1026 different microclimates, topography, and land use practices that affect these factors differently (Vicente-Serrano
1027 et al., 2010, 2010a). This heterogeneity tends to increase as the spatial scale increases, making it harder to calibrate
1028 and validate drought forecasting models. On the other hand, the small number of studies that focused on large
1029 geographic areas is probably due to the challenge of data availability and homogeneity, which arises due to
1030 limitations in data collection and standardization, particularly at larger spatial scales (Dai, 2011). This can lead to
1031 incomplete or inconsistent datasets, which in turn can impact the accuracy of drought forecasting models. Remote
1032 sensing technologies can provide a solution to this problem by allowing for the collection of large-scale, high-
1033 resolution data that can improve the accuracy of forecasting models (Gouveia et al., 2017). The role of remote
1034 sensing data in improving drought prediction will be further discussed in sect. 8.2.

1035 When considering the time scale, the number of droughts forecasting studies tends to increase as the scale
1036 increases. Drought research often emphasizes seasonal, annual, or decadal scales due to various factors. The slow-
1037 onset nature of droughts necessitates studying their progression and recovery over extended periods (Mishra &
1038 Singh, 2010). Investigating longer time scales also allows researchers to analyze the impact of large-scale climate
1039 drivers, such as ENSO or NAO, on drought events (Dai, 2011). Moreover, focusing on these time scales enables
1040 a better assessment of drought consequences on water resources, agriculture, and ecosystems, which are more
1041 pronounced over extended periods (Wilhite & Pulwarty, 2017). Additionally, data availability and reliability tend

1042 to be higher for longer time scales, facilitating more robust analyses. Long-term trends and climate change impacts
1043 on droughts can also be better understood at longer time scales (Trenberth et al., 2014).

1044 Notably, only one study focused on the weekly time scale. Drought forecasting at small scales or weekly time
1045 scales offers several advantages, including early warning and improved water management (Pulwarty &
1046 Sivakumar, 2014), quick response to flash droughts (Mo & Lettenmaier, 2015), support for agricultural decision-
1047 making (Hansen et al., 2011), improved accuracy of longer-term forecasts (Yuan et al., 2015), and model
1048 improvement and validation (Wood et al., 2016). However, drought forecasting at such a small scale may be more
1049 challenging due to the chaotic nature of the atmosphere, making it difficult to accurately model complex
1050 interactions between atmospheric conditions, land surface characteristics, and water management practices over
1051 short periods (Lorenz, 1963; Seneviratne et al., 2012).

1052 On the other hand, the most commonly used forecasting methods were statistical and hybrid statistical models,
1053 with only a few studies applying dynamical models and the percentage of studies applying this last approach
1054 increases with an increase in the temporal scale. There could be several reasons for these findings. Dynamical
1055 models require large amounts of high-quality input data, which may not be readily available for the **MEDR** **MedR**
1056 due to limitations in historical data and spatial coverage (Giorgi & Lionello, 2008). Statistical and hybrid statistical
1057 models often have lower data requirements and are generally computationally more efficient than dynamical
1058 models, making them more suitable for regions with limited data availability and computational constraints.
1059 Furthermore, the percentage of studies applying dynamical models increases with an increase in the temporal scale
1060 because these models are better suited for capturing long-term climate variability and the influence of large-scale
1061 climate drivers (Dai, 2011; Sheffield et al., 2012). Statistical and hybrid statistical models, conversely, are more
1062 effective at capturing short-term variability and local-scale processes, which are often more relevant for drought
1063 forecasting in the **Mediterranean region** **MedR** (Mehran et al., 2014). Lastly, data availability at shorter temporal
1064 scales can be a limiting factor for developing and validating dynamical models (Shah et al., 2018).

1065 In summary, while increasing the spatial scale can decrease the accuracy of drought forecasting studies, increasing
1066 the time scale can improve the accuracy by allowing for a more comprehensive understanding of the various factors
1067 that contribute to drought conditions. It is essential to consider both spatial and temporal scales when conducting
1068 drought forecasting studies to ensure the most accurate predictions possible.

1069 **8 Challenges and Future Prospects**

1070 In the earlier discussion, we analyzed drought indices, factors affecting the accuracy of drought forecasts, and the
1071 significance of spatial and temporal scales in drought predictions within the **MEDR** **MedR** context. Building on
1072 this understanding, the following sections will focus on the challenges and prospects within the realm of drought
1073 forecasting, which will help to pinpoint potential avenues for progress and innovation in this area.

1074 **8.1. Data Assimilation**

1075 The lack of in-situ measurement networks and coarse global seasonal forecast skills
1076 has hindered drought forecasting facilities, especially in data-poor regions (Pozzi et al.,
1077 2013; Haile et al. 2020). In this regard, Data Assimilation (DA) provides a powerful approach to enhancing drought
1078 forecasting accuracy by incorporating different observations and climate forecasts into a hydrologic model to

1079 generate more precise initial conditions (Hao et al., 2018; Tang et al., 2016). Therefore, many studies have referred
1080 to this method to better forecast hydroclimatic variables (e.g., Bazrkar and Chu, 2021; Peng, 2021; Xu et al., 2020;
1081 Liu et al., 2019; Steiger et al., 2018; Steiger and Smerdon, 2017). The ensemble Kalman Filter (EnKF) (Evensen,
1082 1994) algorithm is one of the most popular DA techniques applied by the hydrologic community. However, this
1083 assimilation method is subject to some inherent drawbacks especially in nonlinear dynamic systems thus resulting
1084 in suboptimal performance and violation of water balance (Abbaszadeh et al., 2018). Given these limitations,
1085 emphasis should be placed on the development of improved DA algorithms better adapted to hydrologic models,
1086 which allow the modeling of different temporal and spatial scales and the improvement of water balance. This can
1087 be achieved by modifying the standard approaches such as the ensemble Kalman filter or variational algorithms
1088 so that, accurate predictions can be obtained at a reasonable computational cost. These include among others hybrid
1089 EnKF-Var methods (Bannister, 2017; Bergou et al., 2016; Mandel et al., 2016) and AI algorithms for ensemble
1090 post-processing (Grönquist et al., 2021). One recent advance in data assimilation techniques for drought
1091 forecasting is the use of machine learning algorithms to improve the accuracy of predictions. For example,
1092 researchers have used machine learning techniques to develop models that can analyze large amounts of data from
1093 a variety of sources and generate more accurate forecasts of drought conditions (Aghelpour et al., 2020; Rhee and
1094 Im, 2017; Feng et al., 2019). These models can also be updated in real-time as new data becomes available,
1095 allowing for more accurate and up-to-date forecasts. Another advance in data assimilation techniques for drought
1096 forecasting is the use of remote sensing data and reanalysis to improve the accuracy of predictions, which may be
1097 particularly beneficial in areas where ground-based observations are limited (Shahzaman et al., 2021b; Shi et al.,
1098 2011).

1099 8.2. Remote Sensing and Reanalysis

1100 Various challenges in drought modeling in the MEDR ~~MedR~~ are related to data availability. The lack of climatic
1101 and hydrological observations in ungauged catchments, low station density, short data records, data gaps, and
1102 limited data access in some Mediterranean countries. All these challenges can limit the accuracy and reliability of
1103 drought predictions. ~~Finding~~ Although many efforts are being deployed by developing new complete datasets in
1104 the MEDR (Tuel and El Moçayd, 2023), finding alternative data sources and modeling techniques is essential to
1105 tackle these challenges.

1106 Remote sensing data can provide real-time information about the Earth's surface facilitating effective drought
1107 forecasting, monitoring, and early warning (Zhang et al., 2016). Agricultural drought can be assessed by analyzing
1108 changes in vegetation cover over time. Indeed, drought can lead to marked changes in the health and vigor of
1109 vegetation, and these changes can be detected using remote sensing data (Belal et al., 2014). By analyzing changes
1110 in vegetation greenness over time, it is possible to identify areas that are experiencing or are at risk of experiencing
1111 drought stress. Moreover, drought conditions related to vegetation or evapotranspiration can also be monitored
1112 with drought indices from remote sensing products, such as NDVI or Evaporation Stress Index (ESI) (Shahzaman
1113 et al., 2021a). Microwave satellite data can also be used to estimate soil moisture levels during crop growing
1114 season, which can be used to predict and monitor potential agricultural droughts (Le Page and Zribi, 2019; Yuan
1115 et al., 2015).

1116 In addition, satellite observations of precipitation and soil moisture such as IMERG (Huffman et al., 2015),
1117 PERSIANN-CCS (Sadeghi et al., 2021), CHIRPS (Funk et al., 2015), ~~and~~ SMAP (Entekhabi et al., 2010), MSWEP

1118 V2 (Beck et al., 2019), GLEAM v3 (Martens et al., 2017), and DROP (Turco et al., 2020) can be used in
1119 conjunction with the in-situ observations and ground-based radar observations data to fill observational gaps.

1120 Moreover, data from numerical weather forecasting reanalysis such as ERA5-land were used instead or along with
1121 direct observations to forecast drought in many studies (Babre et al., 2020; Junqueira et al., 2022; Parker et al.,
1122 2021). ERA5-land is a state-of-the-art global reanalysis dataset that can provide a consistent view of the evolution
1123 of land variables (e.g., precipitation, temperature) over several decades at an enhanced resolution (~~~10km~~10 km).
1124 This product obtained by assimilating observations through a 4D-VAR data assimilation technique can be used as
1125 ground truth in data-poor regions. For example, ERA5-land can be used to calibrate and validate climate forecasts
1126 and to choose an ensemble of the most skilled GCMs in reproducing the actual observed climate in a specific
1127 region.

1128 Similarly, SAFRAN, a high-resolution meteorological reanalysis, has shown its utility in regions with sparse
1129 observational data. Trambly et al. (2019) used SAFRAN to generate a high-resolution (5 km) gridded daily
1130 precipitation datasets for Tunisia between 1979 and 2015. Their study, which combined data from 960 rain gauges
1131 with the SAFRAN analysis, demonstrated that SAFRAN surpassed other standard interpolation methods like
1132 Inverse Distance, Nearest Neighbors, Ordinary Kriging, or Residual Kriging with altitude. The outcome was a
1133 highly accurate gridded precipitation dataset that could be instrumental for climate studies, model evaluation, and
1134 hydrological modeling to support the planning and management of surface water resources.

1135 Finally, remote sensing data and reanalysis remain valuable tools for drought forecasting and monitoring, as it
1136 provides timely land surface information that can fill the observational gaps, help to identify areas at risk of
1137 potential drought conditions and to monitor the progression of drought over time.

1138 8.3. Uncertainty analysis in drought forecasting

1139 In spite of the large number of studies that have been carried out on the probabilistic characterization of drought,
1140 the quantification of uncertainty of these forecasts is still ignored in major studies. Uncertainty analysis is an
1141 important aspect of probabilistic drought forecast, as it allows users to understand the degree of confidence
1142 associated with the forecasted probabilities (Hao et al., 2016; Dehghani et al., 2014). Therefore, more efforts
1143 should focus on quantifying the uncertainty beyond just an ensemble of model simulations (AghaKouchak et al.,
1144 2022).

1145 Drought forecasting is subject to epistemic and aleatory uncertainties. The first one arises from incomplete
1146 knowledge of drought processes and can be reduced with improved understanding, more data, and good models'
1147 calibration and validation. The second one is related to the inherent variability and randomness in natural systems
1148 and is often difficult to reduce (Pappenberger & Beven, 2006). In addition, uncertainties in drought forecasting
1149 can vary by region, spatial scale, and temporal scale. As we discussed in sect. 7.3, even well calibrated and
1150 validated, the drought forecasting model will not necessarily perform equally well in all periods or locations. By
1151 considering the uncertainty of the drought model as a nonstationary process in space and time, researchers can
1152 gain new insights into the variability of uncertainty and its underlying causes (AghaKouchak et al., 2022). This
1153 perspective can help identify regions or periods where the uncertainties are particularly high, which can guide
1154 further research, data collection, and model development efforts. Additionally, understanding the space-time

1155 variability of uncertainty can inform the development of more robust and reliable forecasting and decision-making
1156 approaches that account for the changing nature of uncertainty.

1157 Various techniques can be employed to quantify drought forecast uncertainty, including ensemble forecasting
1158 (Palmer et al., 2004), Bayesian methods (Vrugt et al., 2008), sensitivity analysis (Saltelli et al., 2008) and
1159 probabilistic forecasting (Gneiting et al., 2005). Probabilistic drought prediction can also involve the use of data
1160 assimilation techniques to integrate different data sources, including remote sensing data, ground-based
1161 observations, and output from meteorological and hydrological models. Lately, hybrid statistical-dynamical
1162 models have shown their potential in reducing uncertainties associated with both statistical and dynamical methods
1163 (Yuan et al., 2015; Madadgar et al., 2016). For example, shortcomings in dynamical model physics or data can be
1164 counterbalanced by the empirical associations in statistical models. While, uncertainties in statistical models
1165 resulting from shifting climate conditions can be tackled by the physically-based dynamical models (Yuan et al.,
1166 2015).

1167 In summary, probabilistic drought prediction with uncertainty analysis can be useful tools for decision-makers,
1168 as they provide a more comprehensive view of the potential impacts of drought and allow for more informed risk
1169 management decisions. However, what is missing in the current drought forecasting models is not just the
1170 uncertainty quantification, but also a lack of awareness of it (AghaKouchak et al., 2022).

1171 **8.4. Drought Information Systems**

1172 A critical component of proactive approaches to drought preparedness is providing timely and
1173 reliable climate information, including seasonal forecasts, that helps decision-makers prepare
1174 management policies (Manatsa et al., 2017). Identifying drought risk timely depends on our
1175 ability to monitor and forecast its physical causing mechanisms at the relevant spatiotemporal
1176 scale. An integrated national drought monitoring and early warning system has been
1177 implemented in many regions and countries such as the United States, New Zealand, South Asia, India, and Europe
1178 (Prabhakar and Rama, 2022) but has not taken place until recently in developing countries (e.g., the Southern and
1179 Eastern Mediterranean countries). This is probably due to the lack of a drought information system, the sparse
1180 observation networks, and the low predictability of seasonal precipitation in these countries. To overcome these
1181 limitations, there is a need for developing a Drought Information System with a complete approach allowing data
1182 collection and preprocessing, accurate probabilistic drought risk prediction using a combination of ensemble
1183 climate seasonal forecasts, ground-based observations, reanalysis, conventional and remote-sensing observations,
1184 ~~Artificial Intelligence, Data Assimilation~~artificial intelligence, data assimilation and hydrological models and
1185 drought information dissemination through a web-based Drought Early Warning System (DEWS).

1186 **9 Conclusions**

1187 This study reviewed the recent statistical, dynamical, and hybrid statistical-dynamical methods used to forecast
1188 droughts and their application on the MEDR MedR. Drought definitions, classification, indices, and causative
1189 physical mechanisms were also presented in the context of the MEDR MedR. The main conclusions of this review
1190 are:

- 1191 1. There are only a few studies on the analysis of physical mechanisms causing droughts in the MEDR MedR.
1192 The review of these studies confirmed that seasonal drought predictability skills are still very limited over
1193 the region due to its relatively poor teleconnection with ENSO compared to the tropical and subtropical
1194 regions. Besides, MEDR MedR is strongly influenced by other climate patterns, such as the NAO, regional
1195 MO, ULMOi, and NAWA which can also affect the region's weather and climate but their relationship to
1196 drought onset is rather weak and could not explain major droughts in the region. Land surface memory can
1197 also contribute to the predictability of seasonal and sub-seasonal droughts. Thereby, an accurate
1198 representation of these land-atmosphere processes is needed to improve drought forecasting skills in mid-
1199 latitude regions such as the Mediterranean.
- 1200 2. Statistical models were largely used to forecast droughts in the MEDR MedR. One of the major limitations
1201 of these models is that they often assume a stationary relationship between the predictors and the
1202 predictands which can lead to potentially inaccurate forecasts. In this regard, AI models such as SVR, SVM,
1203 and ANN have proven good capacity in detecting local discontinuities and non-stationary characteristics of
1204 the data and show satisfactory forecasting skills at less than 6 months lead time. Moreover, sophisticated
1205 statistical models, incorporating a data pre-processing technique such as wavelet analysis, EMD, or PCA
1206 with AI models have proven to be more efficient than using a single model and can extend the lead time of
1207 the drought forecast up to 12 months. The copulas can also provide valuable insights into the complex
1208 relationships between different drought predictors. The use of copulas enables a more in-depth analysis of
1209 the nonlinear dependencies between variables such as temperature, precipitation, and soil moisture, yielding
1210 a more comprehensive understanding of the factors that contribute to drought risk in a specific region. This
1211 leads to a more sophisticated and reliable forecast of drought probability. Thus, copulas are a highly useful
1212 resource in the ongoing effort to understand and manage the consequences of drought.
- 1213 3. Dynamical models can, given their ability to capture the nonlinear interactions in across the
1214 atmosphere, land, and ocean, but their forecast skill is still limited offer considerable potential for a long
1215 lead time due to the more accurate and reliable seasonal drought predictions. However, the inherent
1216 chaotic nature of the atmosphere. In addition, the reliability of the restricts their forecast skill to a few
1217 months in advance. The dynamical models is related to the quality of data used to drive hydrological
1218 models (e.g., initial hydrologic condition and downscaled drought forecasting has seen notable
1219 advancements, such as enhanced climate forecasts) and the quality of the model calibration and validation
1220 which also depends on the quantity and quality of observations used in these resolution, refined
1221 representation of physical processes. On the other hand, climate predictions from a single GCM cannot
1222 represent all the climate pathways. Therefore, more efforts should focus on the probabilistic estimation,
1223 improved initialization methods, the application of multi-model ensembles, and the development of
1224 coupled modeling approaches. These developments have indeed bolstered the accuracy and reliability
1225 of drought predictions. Nevertheless, the implementation of climate variables, which involves
1226 uncertainty quantification on various GCMs as ensemble members, these models in the MedR is
1227 constrained by challenges such as limited data availability, computational complexity, and inherent
1228 model uncertainties.
- 1229 4. Hybrid statistical-dynamical models can be promising tools to potentially enhance the accuracy and
1230 reliability of drought forecasting in the MEDR MedR. By merging a broad variety of forecasts from

1231 statistical and dynamical models into a final probabilistic prediction, hybrid models benefit from the
1232 strengths of both modeling approaches and improve the forecast skill compared to an individual model. But
1233 their applicability remains challenging due to several constraints. Indeed, the hybrid model may require
1234 careful calibration and validation to ensure that they are performing optimally which can be time-
1235 consuming, requiring a large amount of data, specialized expertise, and high computational resources.

1236 5. One of the major challenges in drought forecasting in the ~~MEDR~~MedR is the lack of long-term, high-
1237 quality hydroclimatic observations to convey the nonstationary patterns and the variability of the climate.
1238 In addition, hydrologic model predictions are often poor, due to model initialization, parametrization, and
1239 physical errors. To address these challenges, it is important to improve the availability and quality of data
1240 for drought forecasting in this region. This could involve implementing better monitoring systems and
1241 increasing the number of weather stations in the region. In addition, efforts should be made to improve the
1242 performance of drought forecasting models by using more advanced data assimilation and machine learning
1243 techniques and to incorporate data from other sources such as state-of-art satellite observations and
1244 reanalysis with relatively high spatiotemporal analysis to provide a superior hydrologic and climate states
1245 estimate and consequently a skillful agricultural and hydrological drought forecasting.

1246 6. Drought mapping is the final stage in which drought risk information is disseminated and communicated
1247 to end users. Major studies in the ~~Mediterranean region~~MedR analyze drought risk using some drought
1248 indices without applying a visualization via maps or presenting the risk on a single map showing the overall
1249 risk situation. An informative visualization of results via probabilistic drought risk maps with regard to
1250 cartographic rigor is recommended, whereby color gradations or contouring are used to
1251 effectively illustrate the range of probabilities. Ensuring cartographic rigor, such maps should maintain
1252 spatial accuracy, use appropriate scaling, and include a clearly defined legend to decrypt different
1253 probability levels. Uncertainties related to drought modeling and prediction also need to be perspicuously
1254 defined, discussed and communicated to increase the intelligibility and comprehensibility of decision-
1255 makers, farmers, and other end users.

1256 7. Finally, much effort should be done to improve the communication and dissemination of drought forecasts
1257 which can help in extending their lead time by ensuring that decision-makers and stakeholders have access
1258 to the most up-to-date information.

1259 **Index of Acronyms**

Adaptive neuro-fuzzy inference systems (ANFIS)	Autoregressive integrated moving average (ARIMA)
Akaike's Information Criterion (AIC)	Autoregressive moving average (ARMA)
Anderson-Darling (AD)	Autoregressive moving average time series of order (11) (ARMA)
Artificial neural network of multilayered perceptron (ANN-MLP)	Autoregressive moving average time series of order 1 (MA1)
Asymmetric Power Autoregressive Conditional Heteroskedasticity (APARCH)	Autoregressive moving average time series of order 2 (MA2)
Atmospheric water deficit (AWD)	Autoregressive time series of order 1 (AR1)
Automated Statistical Downscaling (ASD)	Autoregressive time series of order 2 (AR2)
AutoRegressive (AR)	Bagging (BG)
Autoregressive Conditional Heteroskedasticity time series of order 1 (ARCH)	Bagnouls-Gausson aridity index (BGI)
	Bayesian Information Criterion (BIC)

Breaks for Additive Season and Trend (BFAST)
 Coefficient of efficiency (CE)
 Convolutional neural network long short-term memory (CNN-LSTM)
 Co-ordinated regional climate downscaling experiment for the Mediterranean area (MedCORDEX)
 Corrected and unbiased trend-free-pre-whitening (TFPWcu)
 Coupled Model Intercomparison Project (CMIP)
 Cramers-von Mises (CvM)
 Crop moisture index (CMI)
 Drought class transition probabilities (DCTP)
 Empirical Mode Decomposition (EMD)
 Exponential General Autoregressive Conditional Heteroskedasticity time series of order (11) (EGARCH)
 False alarm ratio (FAR)
 Frequency bias (FB)
 Generalized Autoregressive Conditional Heteroskedasticity time series of order (11) (GARCH)
 Geometric Brownian Motion (GMB)
 Geometric Brownian Motion time series model with asymmetric Jumps (GBMAJ)
 Global Historical Climatology Network-Monthly (GHCN)
 Global Land Data Assimilation System (GLDAS)
 Groundwater Resource Index (GRI)
 Growing season minimum and maximum values (gsmm)
 Hadley Centre Coupled Model version 3 (HadCM3)
 Kolmogorov-Smirnov (K-S)
 Land Surface Temperature (LST)
 Maximum likelihood methods (MLIKE)
 Mean absolute error (MAE)
 Mean error (ME)
 Model output statistics (MOS)
 Moderate Resolution Imaging Spectroradiometer (MODIS)
 Modified Fournier Index (MFI)
 Monthly average relative humidity (MARH)
 Monthly mean solar radiation (MMSR)
 Moving average (MA)
 Multiple Linear Regression (MLR)

National Center for Atmospheric Research (NCAR)
 National Centers for Atmospheric Prediction (NCEP)
 National Oceanic and Atmospheric Administration (NOAA)
 NDVI anomaly index (NDVIA)
 Non-linear AutoRegressive with eXogenous inputs (NARX)
 Normalized Difference Vegetation Index (NDVI)
 Normalized Difference Water Index (NDWI)
 North Atlantic Oscillation (NAO)
 Pedotransfer functions (PTF)
 Periodic autoregressive (PAR)
 Periodic autoregressive moving average (PARMA)
 Principal component analysis (PCA)
 Probability of detection (POD)
 Probability of false detection (POFD)
 Proportion of correct predictions (PC)
 Random forest (RF)
 Random subspace (RSS)
 Random tree (RT)
 Reconnaissance Drought Index (RDI)
 Root mean squared error (RMSE)
 Sea Surface Temperature (SST)
 Seasonal-ARIMA (SARIMA)
 Soil and Terrain Database (SOTER)
 Soil Moisture (SM)
 Soil Moisture Agricultural Drought Index (SMADI)
 Soil Moisture and Ocean Salinity (SMOS)
 Soil moisture anomaly index (SMAI)
 Soil Moisture Deficit Index (SMDI)
 Soil moisture percentiles (Wp)
 Soil Water Deficit Index (SWDI)
 Soil Wetness Deficit Index (SWetDI)
 Standardized Water-Level Index (SWI)
 Streamflow drought index (SDI)
 Support vector Regression (SVR)
 Temperature Condition Index (TCI)
 The Second Generation of Canadian Coupled General Circulation Model (CGCM2)
 Vegetation Condition Index (VCI)
 Vegetation Health Index (VHI)
 Wavelet Analysis (WA)
 Wavelet decomposition (WD)

1260

1261 **Competing Interests**

1262 The authors declare that they have no conflict of interest.

1263 **Author contribution**

1264 Each author has made substantial contributions to the creation of this manuscript. BZ was responsible for
 1265 conceptualization, methodology, investigation, analysis, drafting the manuscript, and reviewing and editing. NEM

1266 contributed to the methodology, analysis, ~~writing, reviewing~~review, and editing processes. EHB was involved in
1267 the methodology, analysis, review, and editing stages.

1268 Disclaimer

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Table 41 Main studies using the Time series model to forecast drought in the **MEDR**

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Bouznad et al., 2021)	Precipitation, temperature, and ET	Aridity index, SPI, NDVI	ARIMA, SARIMA	Monthly, annual	Algeria	Meteorological	Baseline 1985–2014 Future 2015–2024
(Achite et al., 2022)	Monthly precipitation	SPI12, SRI12	ARIMA, SARIMA	Annual	Algeria	Meteorological, hydrological	1972–2018
(Al Sayah et al., 2021)	LANDSAT imageries at a 3-year interval, and meteorological indicators	MFI, BGI, VHI, VCI, TCI, NDWI, NDVI	ARIMA/SARIMA	Annual	Lebanon	Meteorological, hydrological and agricultural	1990–2018
(Tatli, 2015)	IPCC observed precipitation	PDSI	Hurst exponent, Mann - Kendall test	Monthly	Turkey	Meteorological	1966–2010
(Pablos et al., 2017)	LST, NDVI Satellite SM data (SMOS BEC L4 and MODIS SR) and In Situ SM Data	SWDI, SMADI, SMDI, SWetDI, AWD CMI	POD; POFD; FAR; FB	Weekly	Spain	Agricultural	2010- 2016
(Hadri et al., 2021)	NDVI ; Rainfall	SPI, SWI	The Mann-Kendall and Sen's slope	Seasonal	Morocco	Meteorological, agricultural	2008-2017
(Ben Abdelmalek and Nouiri, 2020)	Monthly rainfall series in 16 main meteorological stations	SPI, RDI, Annual PET	Mann - Kendall test, Weighted Inverse Distance interpolation	Annual	Tunisia	Meteorological, agricultural	1973–2016
(Karabulut, 2015)	Precipitation	SPI	Cumulative Deviation Curve	Monthly, seasonal, annual	Turkey	Meteorological	1975–2010
(Jiménez-Donaire et al., 2020)	Rainfall, soil moisture, and vegetation (NDVI)	SPI, NDVIA SMAI	Combined Drought Indicator	Monthly, seasonal, annual	Spain	Agricultural	2003–2013
(Ben Mhenni et al., 2021)	SM (SOTER); MedCORDEX daily grided reanalysis of meteorological data; NOAA weekly NDVI	SPI, SPEI, PDSI, and Wp	Lag-correlation analysis	Seasonal, annual	Tunisia	Meteorological, agricultural	1982–2011
(Derdous et al., 2021)	Rainfall	SPI	the Mann–Kendal, Sen's slope estimator, and the Pettitt test;	Monthly, seasonal, annual	Algeria	Meteorological	1936 –2008

(Mendes et al., 2022)	Precipitation, water level in reservoirs	SPI14	BFAST	Seasonal	Portugal	Hydrological	1978-2020
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1993 **Table 22** Main studies using regression analysis to forecast drought in the MEDR MedR

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Sousa et al., 2011)	Monthly rainfall SST, NAO	PDSI, scPDSI	Calibrated Stepwise Regression	Monthly, seasonal, annual	MEDR MedR	Meteorological	1901–2000
(Papadopoulos et al., 2021)	Monthly precipitation	SPI, RDI	Fuzzy linear regression analysis	Monthly, seasonal, annual	Greece	Meteorological	1996–2016
(Martínez-Fernández et al., 2016)	In situ hourly SM, daily rainfall, daily PET, and SMOS data	SWDI	PTF; linear regression	<u>Weekly</u> , Seasonal	Spain	Agricultural	2010–2014
(Tigkas and Tsakiris, 2015)	Monthly rainfall; average monthly mean, max, and min temperature	PET, RDI	Multiple regression models	Monthly, seasonal, annual	Greece	Agricultural	47-50 years

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1995 **Table 33** Main studies using Artificial Intelligence Models to forecast drought in the MEDR MedR

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Mohammed et al., 2022)	Precipitation	SPI	BG, RSS, RT, and RF	Monthly, seasonal, annual	Syria	Agricultural, Hydrological	1946-2005
(Di Nunno et al., 2021)	Precipitation and discharge		NARX neural networks	Seasonal	Italy	Hydrological	1997-2020
(El Aissaoui et al., 2021)	Monthly average precipitation; Monthly min/max air temperature; MARH; MMSR	SPI, SPEI	SVR1: linear; SVR2: Polynomial; SVR3: RBF; SVR4: sigmoid	Monthly	Morocco	Meteorological	1979–2013
(Achour et al., 2020)	Monthly rainfall data	SPI	TFPWcu; ANN	Monthly, seasonal and annual	Algeria	Meteorological	1960–2010

(El Alaoui El Fels et al., 2020)	Monthly rainfall amount	SPI	PCA, Frequency analysis, ANN	Monthly, annual	Morocco	Meteorological	1970–2017
(El Ibrahimy and Baali, 2018)	Observed SPI	Predicted SPI	ANFIS; ANN-MLP; SVR, ANN, WA-ANFIS, WA-SVR, WA-ANN-MLP	Monthly, seasonal, annual	Morocco	Meteorological	1978–2014
(Djebouai and Souag-Gamane, 2016)	Historical monthly rainfall	SPI	ARIMA, SARIMA, WA-ANN	Monthly, seasonal, annual	Algeria	Meteorological	1936–2008
(Myronidis et al., 2012)	Monthly precipitation Monthly in-situ measurements of water lake levels	SPI	ARIMA-ANN	Annual and seasonal	Greece	Meteorological	1973–2008
(Danandeh Mehr et al., 2022)	Rainfall and temperature time series	SPEI-3 and SPEI-6	CNN-LSTM, genetic programming, ANN, LSTM and CNN	Monthly	Turkey	Meteorological	1971–2016
(Başakın et al., 2021)	Monthly sc-PDSI	Predicted sc-PDSI	ANFIS, EMD-ANFIS	Monthly, seasonal,	Turkey	Meteorological	1900–2016
(Özger et al., 2020)	Monthly sc-PDSI	Predicted sc-PDSI	EMD, WD, ANFIS, SVM, WD-ANFIS, EMD-ANFIS, WD-SVM,	Monthly, seasonal	Turkey	Meteorological	1900–2016

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Table 44 Main studies using Joint Probability Models to forecast drought in the **MEDR** **MedR**.

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Bouabdelli et al., 2020)	Monthly precipitation, temperature 9 GCMs of CMIP5	SPI12, SDI6	Copula theory, Hydrological modeling using GR2M	Seasonal, annual	Algeria	Meteorological, Hydrological	Baseline: 1941–2011, Future: 2021–2100
(Bonaccorso et al., 2015)	NAO; areal monthly precipitation series;	SPI	DCTP (SPI, NAO)	Monthly, seasonal	Sicily, Italy	Meteorological	1921–2008
(Serinaldi et al., 2009)	Mean areal precipitation, aggregated at 6 months	SPI; joint return periods of drought	Probabilistic analysis of drought characteristics using Copula	Seasonal	Italy	Meteorological	1921–2003
(Hamdi et al., 2016)	Daily streamflow data,	The joint probabilities and bivariate	Two-dimensional copula model;	Annual	Tunisia	Hydrological	1966–2008

		return periods	the threshold level method				
(Esit and YUCE, 2022)	Monthly precipitation	SPI	Two-dimensional copula model	Seasonal	Turkey	Meteorological	1963–2016
(Tosunoglu and Can, 2016)	Monthly rainfall series	SPI; probabilistic properties of droughts	Two-dimensional copula model	Monthly	Turkey	Meteorological	1966–2006

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Table 55 Main studies using Markov Chains Model to forecast drought in the [MEDR](#) [MedR](#).

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Habibi et al., 2018)	Annual precipitation from 65 meteorological stations	SPI	Markov chain models, DI and 11 time series models (GMB, GBMAJ, APARCH, AR1, AR2, ARCH, ARMA, EGARCH, GARCH, MA1, MA2)	Annual	Algeria	Meteorological	1960–2010
(Paulo and Pereira, 2007)	67-year averages of monthly precipitation	SPI	Non-homogeneous and homogeneous Markovian modeling	Monthly, seasonal, annual	Portugal	Meteorological	1931/32 – 1998/99
(Lazri et al., 2015)	Annual precipitation maps from meteorological satellite data; 219 rain gauges and radar precipitation	SPI	Markov chain model; Transition probability matrix	Annual	Algeria	Meteorological	2005–2010
(Nalbantis and Tsakiris, 2009)	Monthly Precipitation, monthly streamflow	SPI, SDI	Non-stationary Markov chain	Monthly, seasonal, annual	Greece	Hydrological	1970–71 to 1999–2000.
(Akyuz et al., 2012)	Observed annual streamflow	Probabilities and return periods of droughts	First-order Markov chain model, second-order Markov chain model	Annual	Turkey, New work, Sweden	Hydrological	1938–2005
(Cancelliere et al., 2007)	Monthly Precipitation in 43 precipitation stations	SPI	Markov chain model	Seasonal, annual	Sicily, Italy	Meteorological	1921–2003

2000

2001

Table 66 Main studies using dynamical models to forecast drought in the [MEDR](#) [MedR](#).

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
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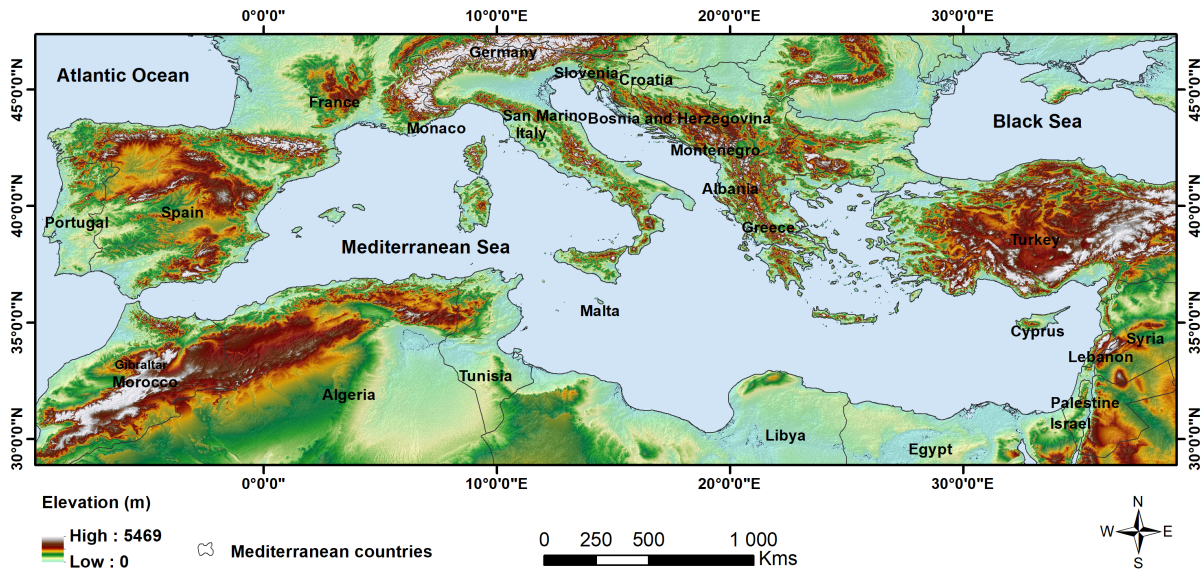
(Elkharrim and Bahi, 2014)	Historical precipitation; HadCM3(monthly precipitation and temperature); Observed GHCN v3; NCEP and NCAR reanalysis	SPI	ASD	Seasonal and annual	Morocco	Meteorological	Baseline 1961-2010 Future 2014-2099
(Marx et al., 2018)	GCMs: GFDL-ESM2M, HadGEM2-ES, IPSL-CM5A-LR, MIROC-ESM-CHEM, NorESM1-M		Hydrological models: mHM, Noah-MP, and PCR-GLOBWB	Annual	Europe	Meteorological and hydrological	Baseline 1971–2000
(Vasiliades and Loukas, 2009)	Observed runoff	PDSI, Weighted PDSI, PHDI and the moisture anomaly Z-index; runoff and soil moisture	monthly UTHBAL conceptual water balance model	Monthly	Greece	Meteorological and hydrological	1960–2002
(Brouziyne et al., 2020)	CNRM-CM5 (RCP4.5, RCP8.5); GLDAS 25 km reanalysis data; Observed daily rainfall and temperature (max and min) series	SPI-12; SDI-12; Monthly runoff, rainfall; Future water yield.	Hydrological model SWAT;	Annual	Morocco	Meteorological, Hydrological	Baseline 1985-2005; Future 2030–2050 and 2080–2100
(Mendicino et al., 2008)	Monthly precipitation, temperature, SPI, NDVI	GRI	A water balance model	Seasonal , annual	Italy	Meteorological and Hydrological	1959–2006
(Dubrovský et al., 2014)	Monthly and daily precipitation and temperature outputs from 16 GCMs simulations (IPCC-AR4)	PDSI, Z-index	Multi-GCM forecast	Seasonal	MEDRMe dR	Meteorological	Baseline 1961–1990; Future 2070–2100
(Ruffault et al., 2014)	Daily precipitation, temperature and global radiation from ARPEGE-Climate model Version 4; Historical observations from SAFRAN dataset	Maps of summer precipitations , number of wet days in summer and drought intensity	Water balance model, quantile mapping/ anomaly method	Annual seasonal	France	Agricultural, Hydrological	Baseline 1961–1990 Future 2071–2100

2003

Table 77 Main studies using hybrid statistical-dynamical models to forecast drought in the **MEDR**

Reference	Inputs	Outputs	Methods	Time scale	Study area	Drought type	Study period
(Ribeiro and Pires, 2016)	UKMO operational forecasting system	SPI3	MLR	Seasonal, annual	Portugal	Meteorological, agricultural, and hydrological	1987–2003

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Figure 1 Topography of the Mediterranean Region: **(30°N - 46°N in latitude and 10°W - 40°E in longitude).**

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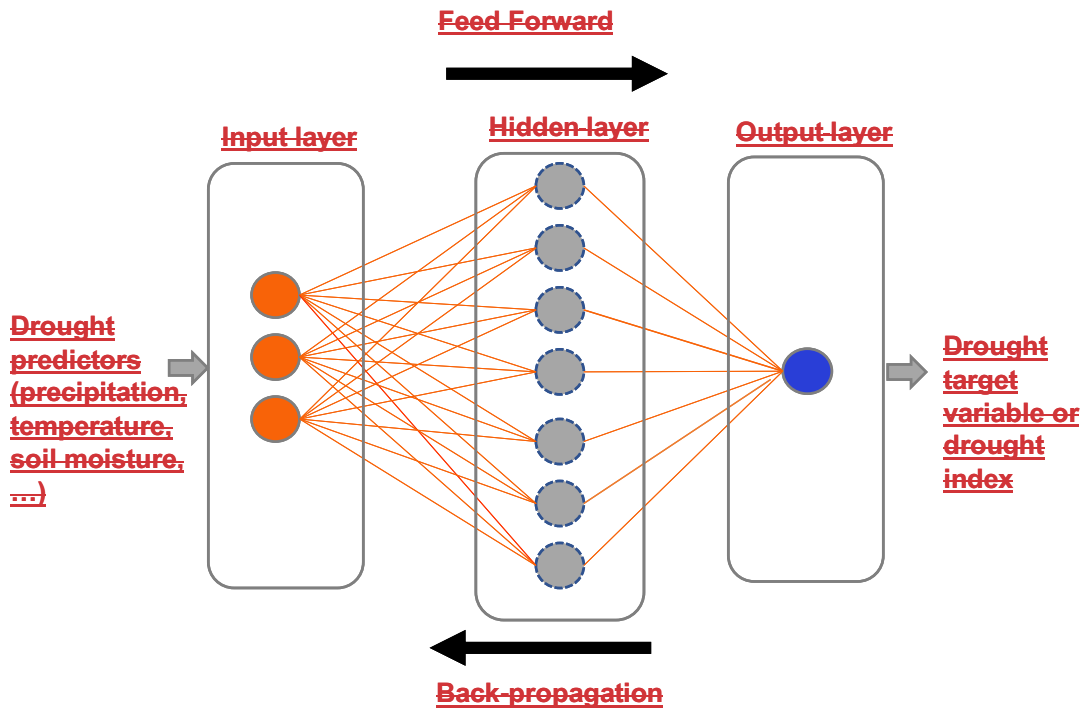
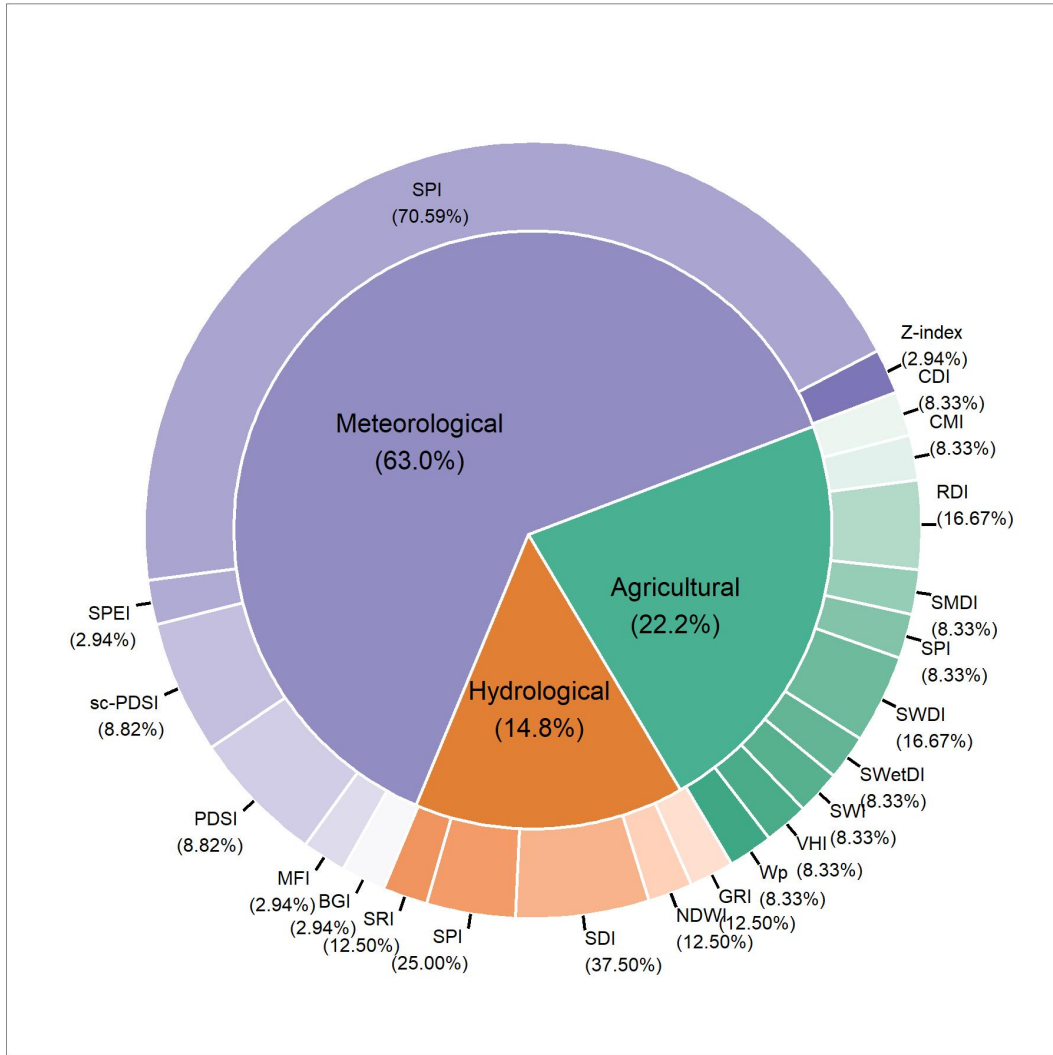


Figure 2 Drought forecasting based on a simple ANN architecture.

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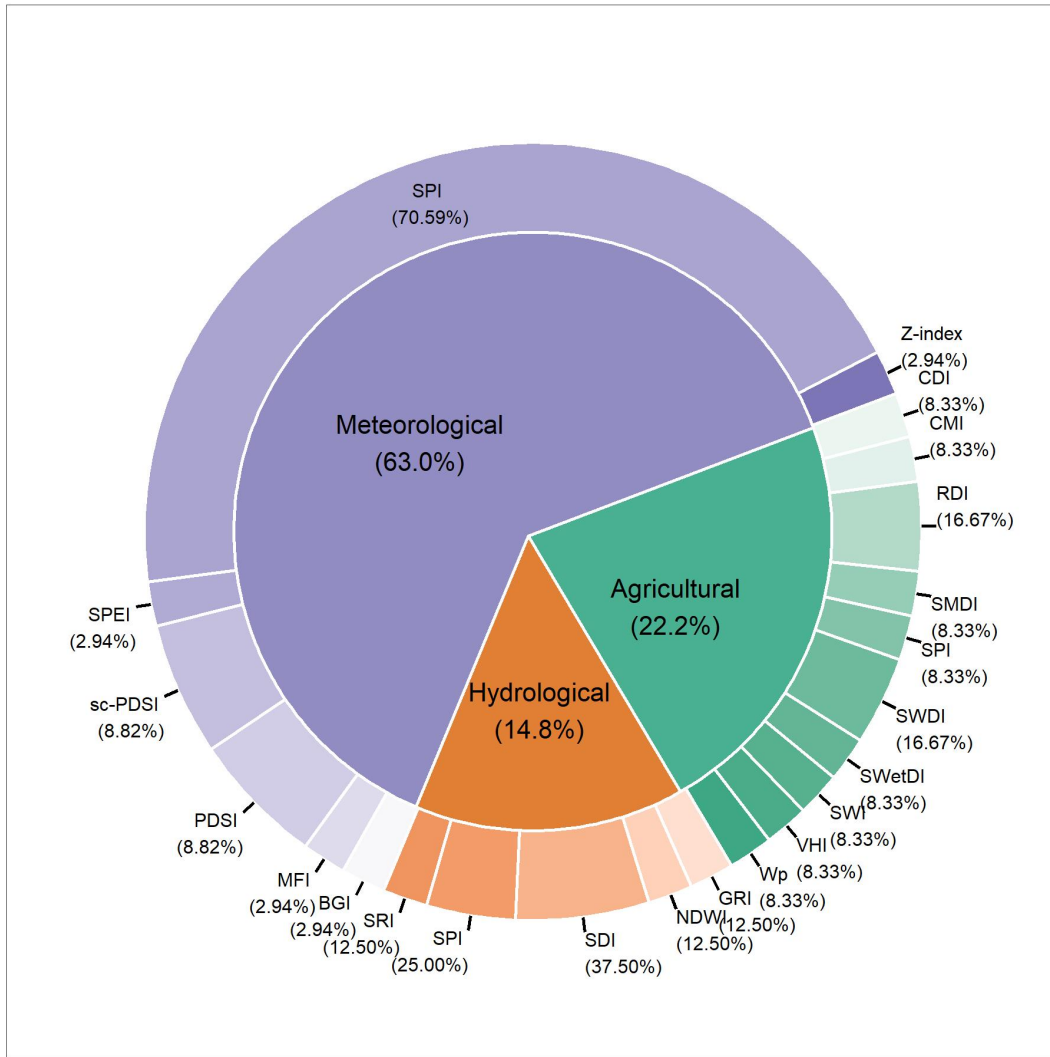
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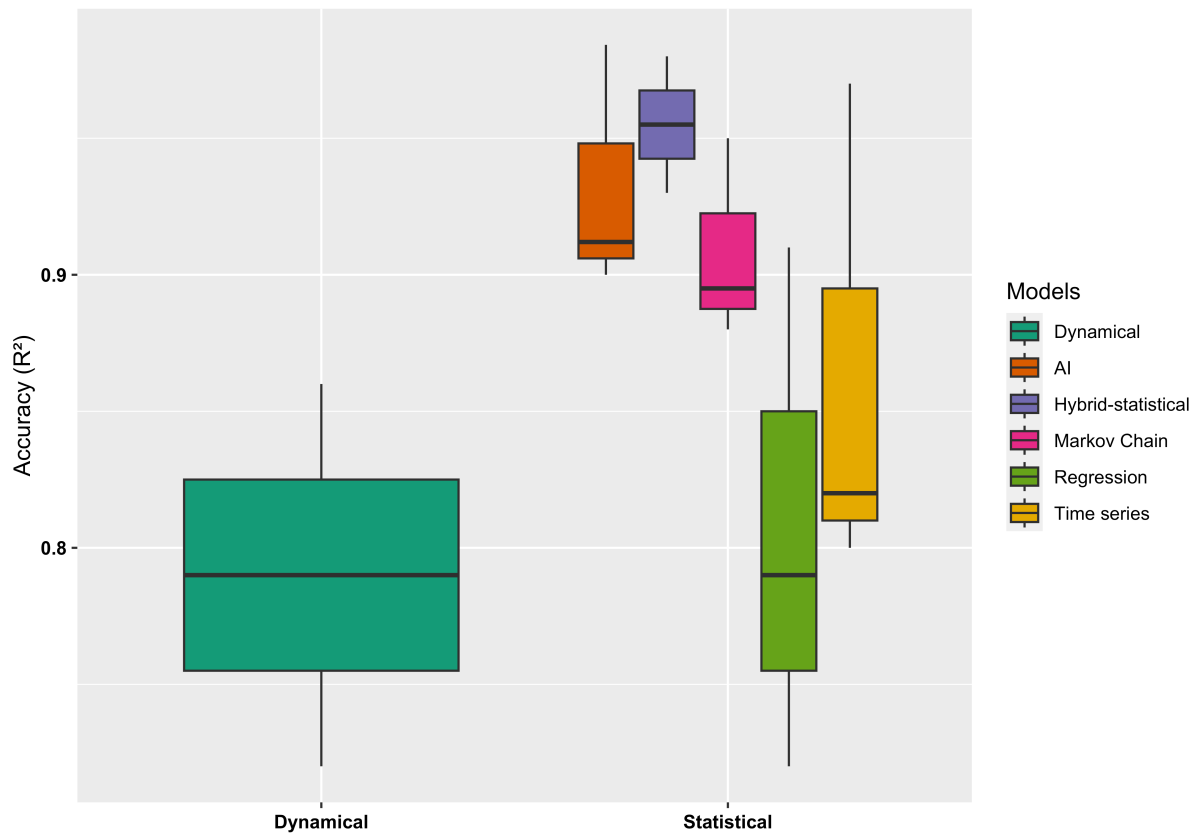
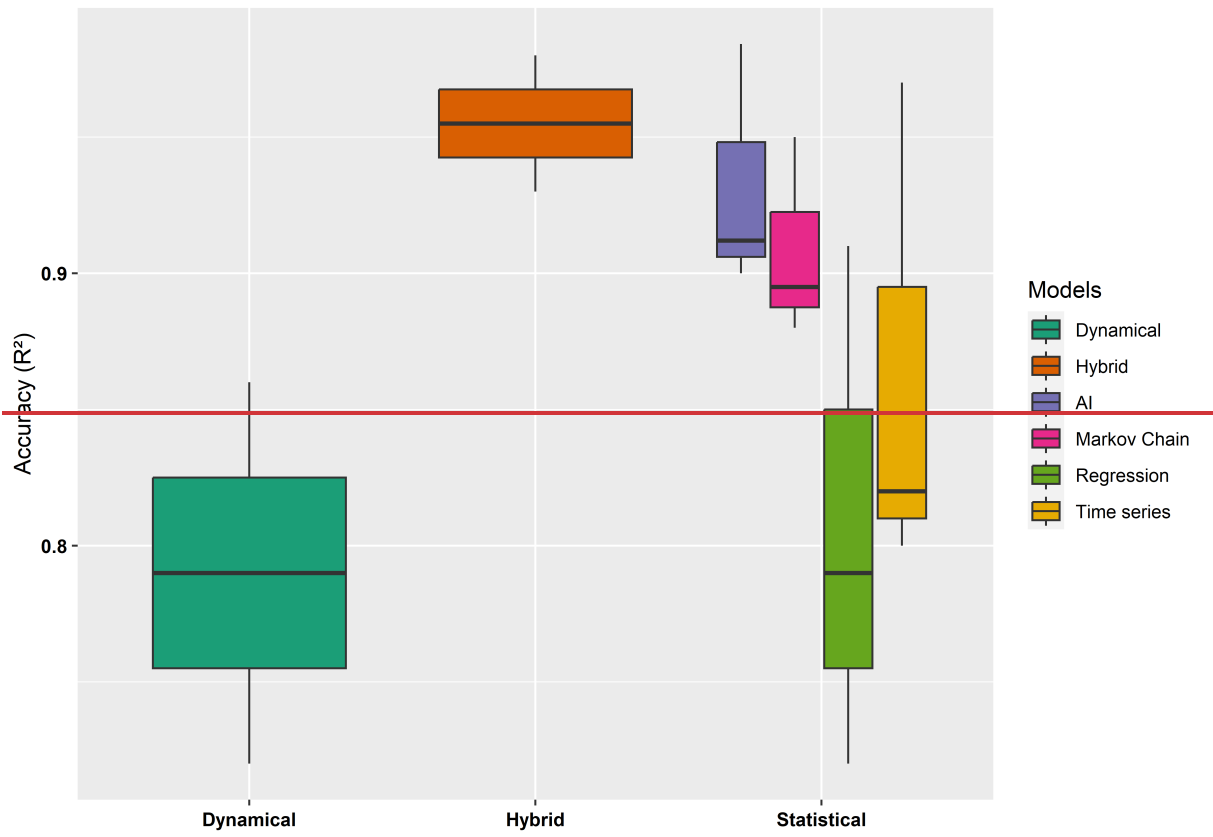
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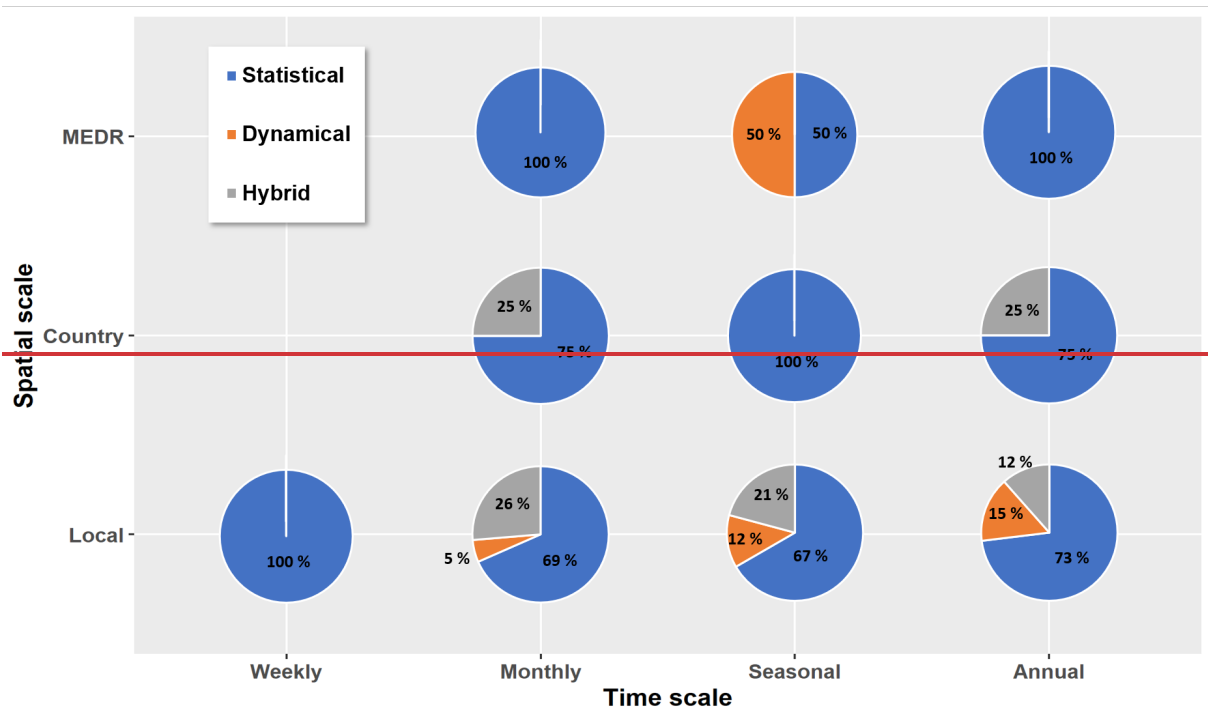


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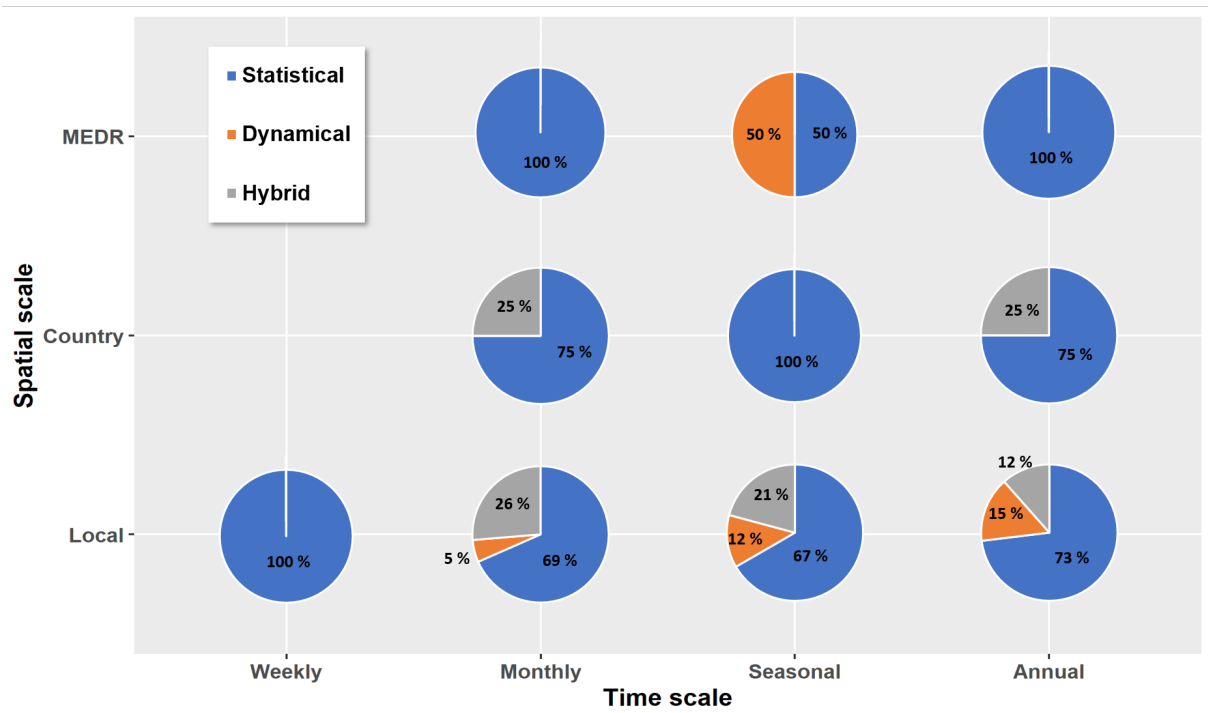
Figure 3 Pie chart showing the proportion of use of indices in the MEDR surveyed studies in MedR (Tables 1-7) for different drought types.



2045 **Figure 4** Box and whiskers plot showing the performance of drought prediction models denoted by the coefficient of
 2046 **determination (R²) for the surveyed studies in MEDR.**



2047 **MedR. The lower box shows the 25th percentile, the upper box shows the 75th percentile and the median (50th percentile)**
 2048 **is represented by the black line inside the box. The whiskers show the extent to the minimum and maximum values**
 2049 **within 1.5 times the interquartile range (IQR) from the box.**
 2050



2051 **Figure 5** Spatial and temporal scales of drought forecasting studies in the **Mediterranean region MedR** with pie chart
 2052 **indicating the percentage of use of drought forecasting method: statistical, dynamical and hybrid—statistical models for**
 2053 **each spatio-temporal scale.**
 2054