

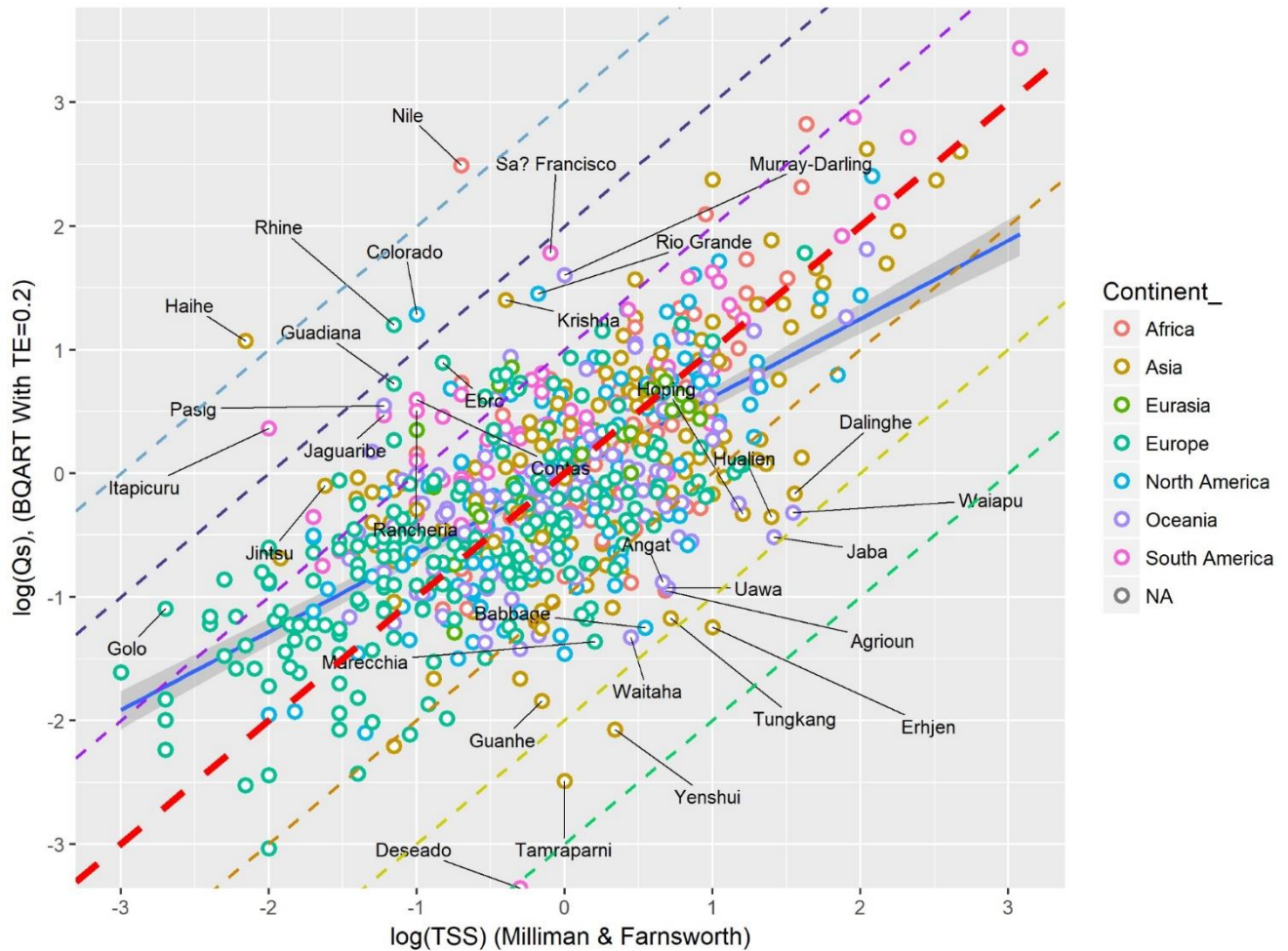
1 **Supplementary Material**

2 **The contribution of diminishing river sand loads to beach** 3 **erosion worldwide**

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6 **S1 River solid discharge models**

7 River sediment discharge has been computed on a global scale using the BQART formula with a unique
8 retention coefficient ($Te = 0.2$, Syvitski and Milliman, 2007). For the calculation, we used data from the
9 HydroBasins database. The basins are associated with the nearest coastal cell along the coastline. Thus, each
10 point on the coast receives the sum of the watersheds whose outlet is close (some may not receive any sediment
11 discharge). To verify the soundness of this calculation, we compared it to Milliman and Farnsworth's (2011)
12 database (Figure S1). There are some problems in linking the two databases (HydroBasins and Milliman and
13 Farnsworth), mostly due to the fact that the gauging stations in Milliman and Farnsworth's database are far
14 upstream from the mouth. Sometimes even one basin in HydroBasins corresponds to two basins in Milliman and
15 Farnsworth (2011). That said, Figure S1 shows that the BQART calculation gives information of the right order of
16 magnitude for the vast majority of rivers. The basins that depart significantly from the 1:1 line in Figure S1 are
17 probably those that are greatly influenced by anthropogenic effects.



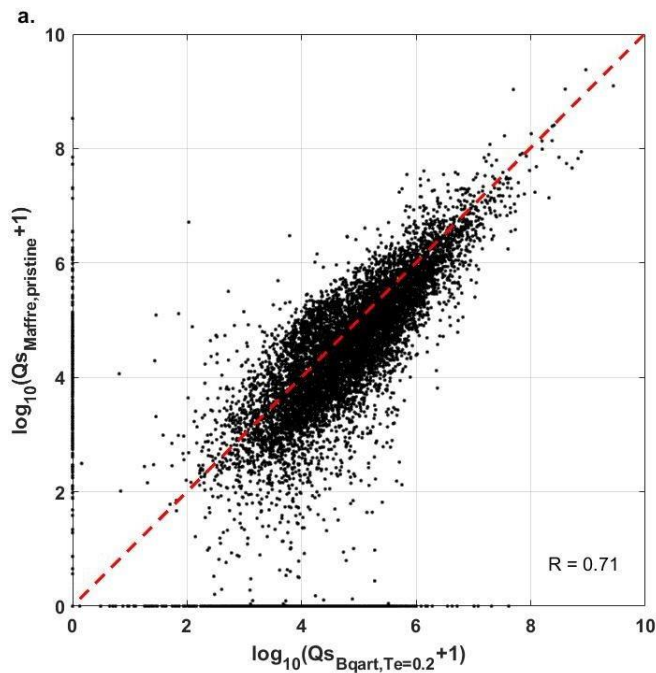
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Figure S1. Observed vs. modeled solid river discharge. Scatter of observed vs. modeled solid river discharge (logarithmic scale). Observations (X-axis) are from Milliman & Farnsworth (2011), modeled rates (Y-axis) are computed using BQART. Red dashed line is the 1:1 line and the blue solid line is the linear regression over the displayed dataset. The labeled basins are those that depart significantly from the 1:1 line.

19 **S2 Comparison of two river input models**

20 In order to better evaluate the impact of dams on the sediment flux in different catchments, we used a model
 21 where the sediment flux can be calculated on every pixel and either summed or partially removed by dams in
 22 order to calculate the sediment outflux to the ocean River. This model was proposed by Maffre et al. (2018) on
 23 the basis of a 3.75 longitude by 1.9 latitude resolution.

24 Figure S2 is a scatter plot to compare the two methods without the effect of dams. Obviously, the BQART outputs
 25 correlate well with Maffre’s model as they were both calibrated against observed river discharges.



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Figure S2. Comparison of the solid river discharges computed using BQART method with the solid river discharges computed using Maffre et al. method, considering a world with no dams and $Te = 0.2$

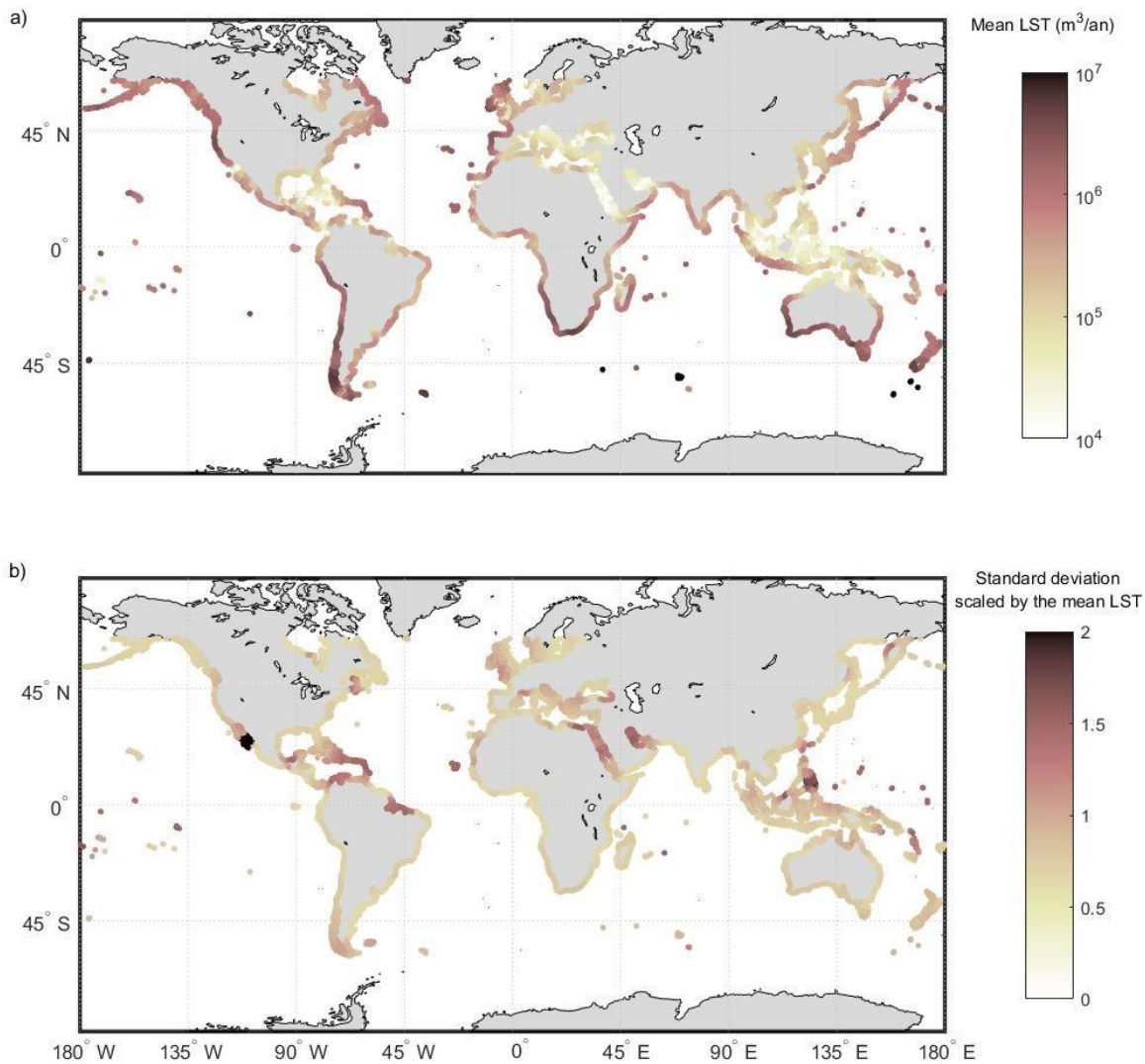
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S3 Longshore sediment transport formula

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In order to evaluate the sensitivity to the choice of the longshore sediment transport (LST), we initially considered four classical formulas: Kamphuis (1991), Kaczmarek et al. (2001), CERC (1984) and Bayram et al., (2007). Common input to the formulas are the nearshore bathymetry (surf zone slope) and a time series of wave height, period, and direction. The average (for each cell) obtained with the four methods is mapped on Figure S3.a. The standard deviation between the four methods, scaled by the local LST average is mapped on Figure S3.b. The scaled standard deviation is systematically below 1 at high latitudes, meaning that the four LST evaluations agree within one order of magnitude. This relative scaled standard deviation increases in closed environments (Figure S3.b) where low mean LST rates are experienced (Figure S3.a). These bulk LST formula have a known tendency to overestimate, particularly in low energy wave environments (Tran et al., 2021) and therefore have to be used cautiously.

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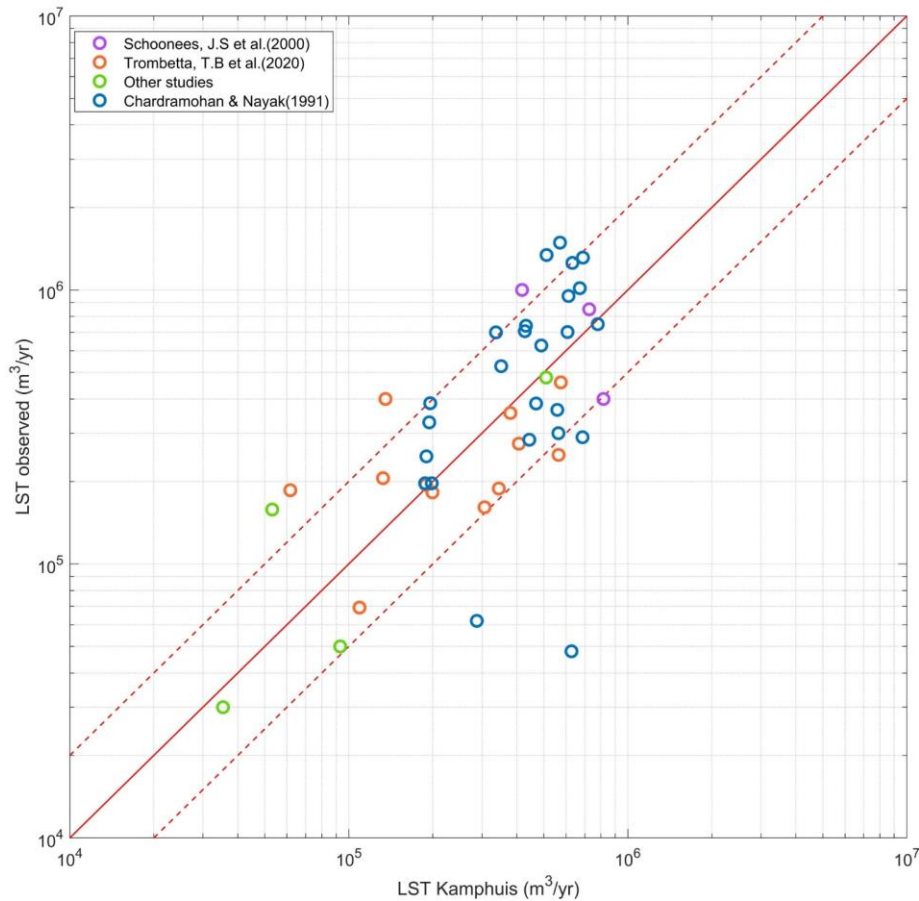


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Figure S3. LST computed fluxes and deviation depending on the model used. (a) average LST computed with four different formulas (Kamphuis, Kaczmarek, CERC and Bayram). (b) standard deviation of the results of the different models, scaled by the average LST at each point (average value mapped in panel a).

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41 In an exercise to quantitatively validate our necessarily coarse global sediment model and the use of Kamphuis
 42 formula, we compared our computed longshore sediment transport (LST) at worldwide monitored sites (Figure
 43 S4). Overall, Figure S4 shows that our modeled longshore sediment transport rates are of the same order of
 44 magnitude compared to observations. However, we note LST rates computed with Kamphuis formula sometimes
 45 fall out of the range of observations (e.g. the 2 blue dots far below the 1:2 line), especially for observed LST rates
 46 below $10^5 \text{ m}^3/\text{yr}$.



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Figure S4. Observed vs. modeled longshore sediment transport. Scatter of observed vs. modeled longshore sediment transport on logarithmic scales. Observations are from : Schoonees et al., 2000; Trombetta et al., 2020; Chardramohan & Nayak, 1992; Aagaard et al., 2004; Roger and Ravens, 2008; Appendini et al., 2012; Cipriani and Stone, 2001. Modeled rate are computed using Kamphuis (1991) formula. Red solid line is the 1:1 line and dashed red lines are 2:1 and 1:2 lines.

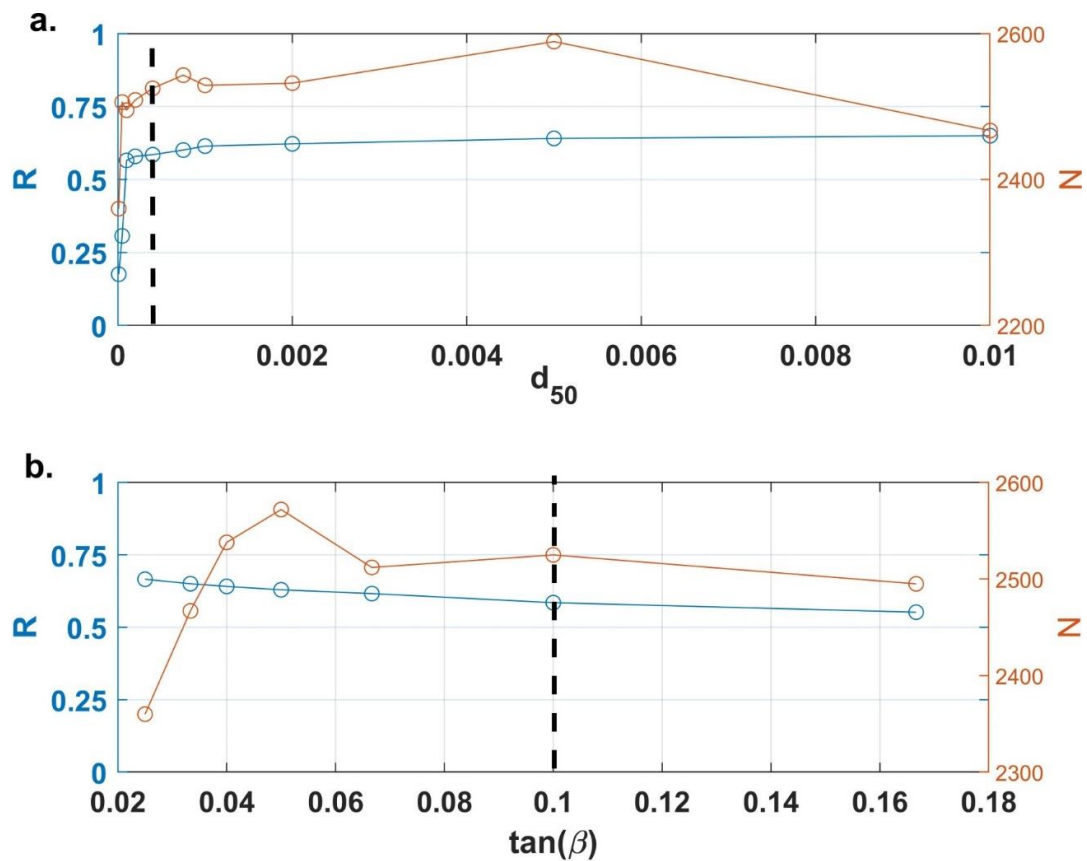
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S4 Sensitivity of model to sediment grain size and coastal slope

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Figure S5 shows the sensitivity of the model to variations in the input parameters d_{50} (median grain size) and $\tan(\theta)$ (beach slope). Model output are generally stable even for large changes in d_{50} and $\tan(\theta)$.

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Figure S5: Sensitivity of the model to two important parameters: sediment grain size and beach slope. Correlation coefficient between modeled sediment budgets and observed erosion trends, R (in blue) and number N of cells (in brown) as a function of: a. the median diameter of sand grains d_{50} ; b. the beach slope $\tan\beta$. Vertical dashed lines show the values used in this study.

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