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2   # Assessing Landslide Damming susceptibility in Central Asia

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12   **Abstract**

13   Central Asia regions are characterized by active tectonics, high mountain chains with extreme topography with

14   glaciers and strong seasonal rainfall events. These key predisposing factors make large landslides a serious natural

15   threat in the area, causing several casualties every year. The mountain crests are divided by wide lenticular or

16   narrow, linear intermountain tectonic depressions, which are incised by many of the most important Central Asia

17   rivers and are also subject to major seasonal river flood hazard. This multi-hazard combination is a source of

18   potential damming scenarios which can bring cascading effects with devastating consequences for the surrounding

19   settlements and population. Different hazards can only be managed with a multi-hazard approach coherent within

20   the different countries, as suggested by the requirements of the Sendai Framework for Disaster Risk Reduction.

21   This work was carried out within the framework of the SFRARR Project ("*Strengthening Financial Resilience

22   and Accelerating Risk Reduction in Central Asia*") as a part of a multi-hazard approach with the aim of providing

23   a damming susceptibility analysis at a regional scale for Central Asia. To achieve this, a semi-automated GIS-

24   based mapping method, centred on a bivariate correlation of morphometric parameters defined by a morphological

25   index, originally designed to assess the damming susceptibility at basin/regional scale, was modified to be adopted

26   nationwide and applied to spatially assess the obstruction of the river network in Central Asia for mapped and

27   newly formed landslides. The proposed methodology represents an improvement of the previously designed,

28   requiring a smaller amount of data, bringing new information on the damming hazard management and risk

29   reduction for the Central Asia regions.

30   **1   Introduction**

31   The mountainous areas of the Djungaria, Tien Shan, Pamir and Kopetdag in Central Asia territories are

32   characterized by complex and active tectonic and are the sources of most of Central Asia rivers. A rugged

33   topography along with complex geological structure and high seismicity are ideal setting for large slope failures.

34   In general, when landslides completely obstruct a river channel, they generate a landslide dam whose consequences

can be a serious hazard forming upstream backwater and causing catastrophic downstream flooding, changes in the riverbed course, embankments instability triggering other landslides with a cascading effect (Swanson et al. 1986; Costa and Schuster 1988; Casagli and Ermini 1999). The effects of impounded water and anomalous flood waves, resulting from a dam breach, have significant economic and social impacts in upstream and downstream areas with economic and human losses (King et al. 1989; Dai et al. 2005; Chen and Chang 2016). Rebuilding costs can be extensive, as they are direct (e.g., infrastructure and buildings reconstruction, safety measures) and indirect (e.g., loss in real estate value and damage caused to industrial and agricultural production), harder to estimate.

Most of landslide dams have a short life as about 40% of them collapse within 24 hours after formation and about 80% within one year (Tacconi Stefanelli et al, 2015; Fan et al., 2020). Given the limited available time, a complete and reliable analysis of the risks, requiring in-depth study of the phenomenon, is not achievable during the event and only rapid assessments for the dam stability are suitable. When the people to evacuate are too many or the related risk is too high, engineering measures for the hazard reduction are attempted: among these are for example modification of slope geometry, drainage, retaining structures and internal slope reinforcement (Popescu and Sasahara 2009; Schuster and Evans 2011). Therefore, part of the effects from landslide damming can be avoided or at least reduced thanks to mitigation and prevention measures (e.g., slopes stabilization or re-profiling) if the most critical areas with the highest damming probability are known. Consequently, planning and prevention tools, such as risk and susceptibility mapping, are essential to reduce the costs of natural hazard and improve the efficiency of environmental management.

Reactivation of ancient landslides triggered during different climatic and environmental conditions may often generate new mass movements (Casagli and Ermini 1999; Canuti et al. 2004; Dikau and Schrott 1999; Borgatti and Soldati 2010; Crozier 2010). Landslides generated in the past are now dormant, with strength parameters of the sliding surface close to the residual ones, and difficult to recognize because vegetation, erosion and superficial alteration hide their morphology. Natural causes, such as earthquakes, river undercutting, rainfall, and snowmelt, or even anthropic activity can reactivate these ancient phenomena. Therefore, all dormant landslides capable to reach a river along their pathway can potentially dam it and should be investigated. New landslides, instead, may develop wherever are present suitable conditions along the slopes. The spatial occurrence probability is commonly assessed by landslide susceptibility analysis, highly dependent on landslide volume (Catani et al., 2016), which is difficult to accurately predict.

Landslides in Central Asia are quite common and a considerable number of them have huge dimension, often induced by strong earthquakes but also by floods, heavy rainfall and snowmelt (Behling et al., 2014; Golovko, 2015; Havenith et al., 2015a, 2015b, 2006b; Kalmetieva et al., 2009; Rosi et al., 2023; Saponaro et al., 2014; Strom and Abdrakhmatov, 2017, 2018). Concerning landslide dam events, in Central Asia regions several mass movements of considerable size produced the obstruction of a river section, of which more than 100 still are existing with a lake (Strom, 2010). Although many of these could be considered stable (Strom, 2010), the occurrence of devastating outburst floods in the last century show that their potential hazard should never be overlooked also considering the seismicity of the region. In the Rushan and Murgab districts of Gorno-Badakhshan Autonomous Oblast (Pamirs, Tajikistan) along the Murghab river, the Usoi dam is one of the most famous of the many cases in the regions. Its impounded lake, called Lake Sarez, is 60 km long with 500 m of depth and has a

73 stored volume of about 17 km³, representing the world deepest landslide-dammed lake (Costa and Schuster, 1991;

74 Fan et al., 2020). It was originated on February 18th, 1911, when a MW 7.2 earthquake triggered a giant wedge-

75 failure of about 2.2 km³ of rock (mainly quartzite, schist, shale and dolomite) and debris that blocked the Murgab

76 River and a tributary valley, forming the 560 m high, 5 km long and 4 km wide Usoi dam, impounding Lake Sarez,

77 also creating the smaller Lake Shadau (Strom, 2010).

78 Landslide dam evolution, according to some studies (Swanson et al. 1986; Ermini and Casagli, 2003; Dal Sasso et

79 al. 2014; Tacconi Stefanelli et al. 2016), can be estimated through geomorphological indexes which require

80 parameters characterizing the landslide (or the dam) and the river (or the lake). Geomorphological indexes are a

81 powerful classification tool, but their prediction power depend mainly on long studies, a large amount of data and

82 measurement efforts given their empirical nature. Many of these indexes need parameters not always available and

83 easy to acquire, such as grain size distribution (Liao et al. 2022) or landslide velocity (Swanson et al., 1986).

84 In this work, we propose a simple semi-automatic GIS-based mapping methodology to verify the damming

85 susceptibility of river networks at national scale from existing and neo-formed landslides trough a

86 geomorphological index. This activity research was carried out in the framework of the SFRARR Project

87 ("*Strengthening Financial Resilience and Accelerating Risk Reduction in Central Asia*") as a part of a multi-hazard

88 approach (Bazzurro et al., 2023).

89 The proposed mapping methodology represents an innovation in terms of application simplicity, availability of

90 data and of extension of the analysed area, bringing new information on the damming hazard in the Central Asia

91 regions where the landslide susceptibility is quite high (Rosi et al., 2023) and a set of input data required for the

92 methodology application were available.

93 **2   Study area**

94 Central Asia is a region of vast diversity encompassing high mountain chains, deserts, and steppes (Figure 1). The

95 southern and eastern parts of the region are predominantly occupied by the mountainous areas of Djungaria, Tien

96 Shan, Pamir, Kopetdag, and a small part of Western Altaj, with peaks exceeding 7,000 m above sea level (a.s.l)

97 (Strom, 2010). These intraplate mountain systems, developed in the Cenozoic as a result of the India-Asian

98 collision, is located between the Tarim Basin and the Kazakh Shield (Molnar and Tapponier 1975, Abdrakhmatov

99 et al., 1996; 2003; Zubovich et al., 2010, Ullah et al., 2015). This study focusses the attention on the territories of

100 Central Asia that includes Turkmenistan, Kazakhstan, Kyrgyz Republic, Uzbekistan, and Tajikistan, covering a

101 surface of more than 4·10⁶ km². Mountain building began in the Oligocene (Chedia 1980) or later (Abdrakhmatov

102 et al. 1996), resulting in a complex system of basement folds interrupted by several thrusts and reverse faults with

103 lateral offset of important amounts (Delvaux et al. 2001).

104

Figure 1. Geographical and geomorphological framework of the study area. Lake's polygons from Esri,

Garmin International, Inc.; topographic base from NASA's SRTM project (Far and Kobrick, 2000).

107   The mountain belts contain several regional fault zones, and others cross the mountain systems with a NW-SE axis

108   (Trifonov et al. 1992). Paleozoic crystalline rocks form, for the most part, the mountain ridges which correspond

109   to a neotectonic anticline and are separated by tectonic depressions, with lenticular or linear shapes. These

110   intermountain depressions host the primary river valleys and are filled by Neogene and Quaternary deposits,

111   principally sandstone, siltstone interbedded by gypsum, and conglomerates (Strom and Abdrakhmatov, 2017).

112   Lithologies from Mesozoic and Paleogene are characteristic of the areas at the foot of mountain ranges. This main

113   deeply incised river network, fed by glaciers, snowmelt water and rain, is linked by narrow deep gorges up to 1-2

114   km deep (Strom and Abdrakhmatov, 2018) and is the origin of most of the rivers in Central Asia.

115   The retreat and shrinkage of glaciers in Central Asia regions induced by the global warming produce a seasonal

116   variation in river discharge and consequently an increase of its induced hazards such as glacial lake outburst floods

117   (GLOFs) resulting in countless losses of human life and destroyed infrastructure (Petrov et al. 2017; Wang et al.,

118   2013). The high seismicity, frequent floods and a complex geological and topographic structure contribute to

119   predispose the region to frequent landslides which can potentially obstruct the narrow valleys of the mountain

120   ranges and in turn be the cause of chain risks (CAC DRMI, 2009; Havenit et al, 2017).

3   Materials and Methods

The Morphological Obstruction Index (MOI) (Tacconi Stefanelli et al., 2016) is a bivariate index able to evaluate the potential hazard posed by landslide dams that requires only simple morphometrical parameters which are easily extracted from common Digital Elevation Models. The MOI has been used in several studies, such as in Italy and Peru (Tacconi Stefanelli et al., 2016; 2018), to assess landslide damming susceptibility showing better results than others popular indexes (Swanson et al., 1986). It is a useful tool for identifying high-risk areas and for prioritizing mitigation efforts in landslide-prone regions.

The MOI is calculated by dividing the volume of the landslide, V_l (m³), by the width of the river valley, W_v (m), at the dam location.

$$MOI = log\left(V_l/W_v\right)$$
(1)

The MOI is based on the principle that the higher the ratio of the landslide volume to the river width, the greater the potential for dam formation. It is important to point out that river width, W_v, shall be defined as the width of the river valley which can potentially be obstructed creating a dammed lake and not of just the channel where the river flows, as is often misinterpreted, although in narrow mountain valleys these often coincide.

Landslide dams analyzed by the index can be grouped within three evolutionary classes: formed, not formed and of uncertain evolution. The limits of these domains are depicted by two lines, the lower red "Non-formation line" and the upper blue "Formation line" (Figure 2).
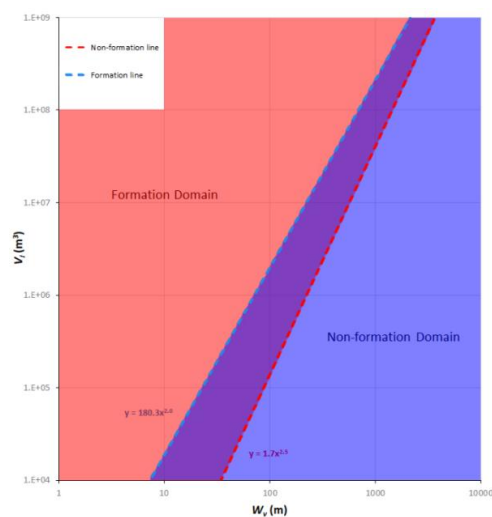
Figure 2. Schematic plot of the non-Formation line and Formation line.

The equation of the former is expressed as follows:

$$V_l' = 1.7 \cdot W_v^{2.5}$$
(2)

142   Where V_l' is the "Non-formation volume" and is the minimum landslide volume able to potentially block a river

143   with a given width W_v. Smaller volumes cannot completely dam the river. The latter expression draws the upper

144   limit for not formed dams and is expressed as follows:

145    $$V_l'' = 180.3 \cdot W_v^2 \tag{3}$$

146   Where V_l'' is the "Formation volume", is the minimum landslide volume able to dam the river valley, with a

147   confidence of 99%, and the inferior boundary of the Formation domain (which includes only formed dams).

148   As originally proposed by Tacconi Stefanelli et al. (2020), these two equations, Eq.(2) and (3), can be used to

149   apply a simple semi-automatic methodology in order to verify at basin scale the damming susceptibility from

150   existing and neo-formed landslides. The following semi-automated procedure, inspired by the one of Tacconi

151   Stefanelli et al. (2020) of which this represents an improvement, is applied on a national scale and can be

152   reproduced entirely in a GIS (Geographic Information System) environment.

153   Within an even medium-long time interval the valley width in each river stretch does not change significantly and

154   can be considered an immutable factor in the MOI equation (Eq.(1)). Starting from this assumption, along with

155   Eq. (2) and Eq. (3), if the average river width W_v of each river stretch can be assessed, the two threshold landslide

156   volumes V_l' (Non-formation volume) and V_l'' (Formation volume) can be estimated for each river stretch.

157   Landslides that cause river obstruction are in many cases reactivations of ancient movements that are still in a

158   condition of partial instability and that have not reached a potential equilibrium reaching the valley floor.

159   Therefore, with a landslide inventory it is possible to assess, with some assumptions and simplifications, which

160   among the mapped landslides are able to dam their river section. Each landslide that is not already laying in the

161   valley floor with a volume bigger than V' and V" are identified as potentially prone to block the river in the future

162   in that point. Then, a "Map of Damming Susceptibility" for reactivation of existing landslides can be generated.

163   The likelihood prediction for new landslides, with volume bigger than V_l' and V_l'', is a much more difficult task

164   as the volume is a complex value to be estimated (Catani et al., 2016). The exceeding probability of landslide

165   volume used by Tacconi Stefanelli et al. (2020) was reached thanks to the knowledge of the alpha exponent of the

166   statistical frequency distribution of the landslide volumes in the whole study area. To achieve this, a database of

167   landslides with a very high number of events (tens or even hundreds of thousands) should be available (Catani et

168   al., 2016), which in our study area unfortunately is not. To have an assessment of the damming susceptibility for

169   neo-formed landslides the two volume threshold values, evaluated for all the river networks, can be used as well.

170   After estimating the river width of every river stretches, the V_l' and V_l'' values of each of them can be computed

171   through the corresponding two equations. In this way there will be two reference values to be able to assess whether

172   the volume of a new landslide can potentially obstruct an affected river stretch.

173   The input data needed for the procedure are a Digital Elevation Model, a vector layer of the river network, and an

174   updated landslide inventory. The data quality and resolution such as the landslides inventory completeness, the

175   river network reliability and the DEM's pixel size heavily affect the quality of the result (Tacconi Stefanelli et al.,

176   2020). Thus, it was decided to use the DEM with the higher resolution freely available from the NASA's SRTM

177   project (Far and Kobrick, 2000) with a 1 arc-second, or about 30 meters of resolution. The river network came

178   from Coccia et al., (2023). The latter input data is a database of 8910 landslides, that is a compilation of several

179   different inventories collected through decades of field surveys, studies and remote sensing analysis in the study

180   area, shown in Figure 3.

181   Hereafter the detail of each inventory:

182   •   The "Rockslides and Rock Avalanches of Central Asia" (Strom and Abdrakhmatov, 2018): an inventory

183   including more than 1000 of very big (>=1 Mm^3) rockslides and rock avalanches, covering central Asian countries

184   (excluding Turkmenistan and Altai) and also Chinese Tien Shan and Pamir, and Afghan Badakhshan. Collected

185   in decades of field survey and analysis of aerial/satellite imaging, it includes also information on morphometric

186   parameters (runout, area), dammed lakes, head-scarps, and quantitative characteristics (such as area, volume) for

187   about 600 cases.

188   •   The "Tien Shan landslide inventory" (Havenith et al., 2015a): is the biggest database in the study area.

189   Assembled through field work, remote sensing and geophysical data interpretation, it includes the elements of the

190   previous inventory alongside other smaller landslides in soft sediments (Havenith et al. 2006a; Schlögel et al.,

191   2011) for a total of 3,462 landslides polygons, including information on landslide length and area.

192   •   The "Multi-temporal landslide inventory for a study area in Southern Kyrgyz Republic derived from

193   RapidEye satellite time series data (2009 – 2013)" (Behling et al., 2014; 2016; 2020), includes 1,582 landslide

194   polygons mapped from multi-sensor optical satellite time series data (from 1986 to 2013) over an area of 2,500

195   km^2 in the Fergana valley rim in southern Kyrgyz Republic, and include information on landslide activity (area

196   and year of trigger).

197   •   The "Tajikistan landslide database" produced by the Institute of Water Problems, Hydropower,

198   Engineering and Ecology of Tajikistan (IWPHE), with 2,710 landslide polygons and 114 landslide-prone areas,

199   including information on landslides length and area.

200   •   The Institute of Seismology of the Academy of Science of Uzbekistan (ISASUZ) provided an inventory

201   which covers the Tashkent province composed by a point inventory (Nyazov R.A. 2020) and a polygon inventory

202   (345 landslide) digitized from the maps in Juliev et al., 2017.
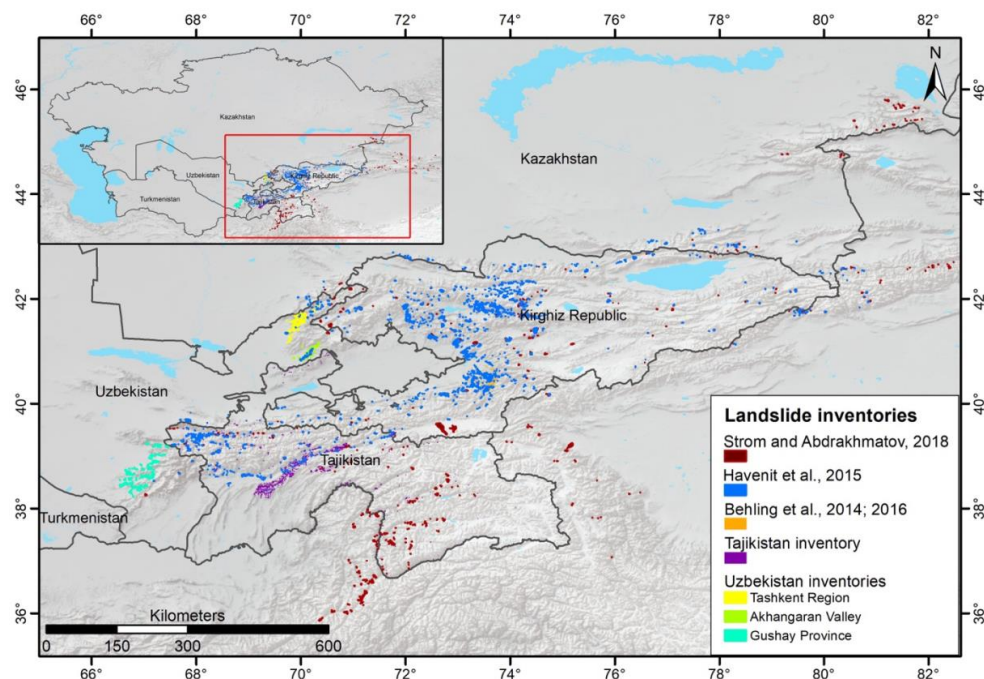
Figure 3. Map of the landslide inventories in the study area. Lake's polygons from Esri, Garmin
International, Inc.; basemap from Esri, USGS, NOAA.

The methodology adopted to obtain the maps of damming susceptibility, derived from Tacconi Stefanelli et al.
(2020), is summarized in the following main steps displayed in Figure 4. According to the literature (Swanson et
al., 1986; Fan et al. 2014, 2020, 2021; Tacconi Stefanelli et al., 2015, 2018), river obstructions occur in most of
the time within hilly or mountainous areas and specially along steep slopes. Therefore, considering the extension
of the study area, in order to reduce the time of elaboration and improve the visualization of the results, in step I
of Figure 4 a series of unnecessary data were removed from the calculations during some preliminary operations.
For this reason, river that flaw in flat areas (with less than 4° slopes) were not considered in the elaborations, since
their damming probability is certainly very low, and the potential impounded lake should have a negligible volume.
Additionally, to have maps easier to manage and display, the river network was split in 5 km long river stretches
consecutive to each other.

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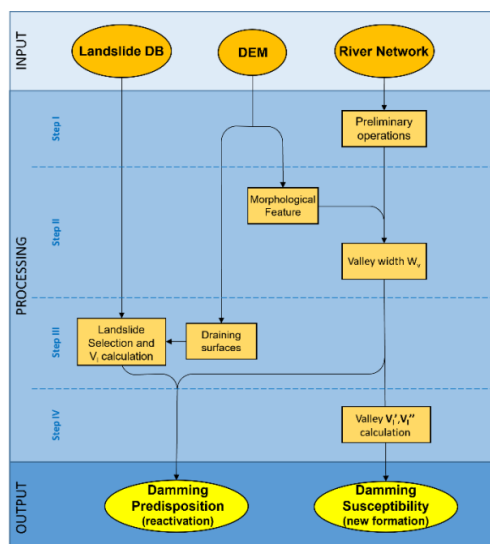
Figure 4. Flow chart of the main steps of the mapping methodology.

In applied geomorphology and natural science studies the analysis and characterization of the landscape has evolved during the last decades with the increasing accessibility of remote sensing data and the development of different algorithms able to automatically extract morphological features and landform information even at broad scales (Drăguţ and Dornik 2016; Maxwell and Shobe 2022; Righini and Surian 2018; Wang et al. 2010).

As already mentioned, the clear definition of the width of a river can be subjective and its measurement difficult to repeat especially if performed by different operators. In step II of Figure 4, an objective automatic method to extract morphometrical parameters have been chosen also for this reason. Wood (2009) implemented the "*LandSerf*" software (already incorporated in SAGA GIS or QGIS software), designed to automatically classify landforms from DEMs. Similarly for pattern detection and texture analysis within image processing, the software extracts land-surface parameters (e.g., slope, aspect, and curvature) from DEMs through a multi-scale approach. During these processing, the algorithm performs a classification of the landscape, grouping the landforms with homogeneous morphometric characteristics (pits, channels, peaks, ridges, passes, and planes) as shown in Figure 5. Thanks to this algorithm of morphological forms analysis proposed by Wood (2009), the polygons representing the morphological unit of the river valley can be automatically defined objectively even in a large area and extracted.

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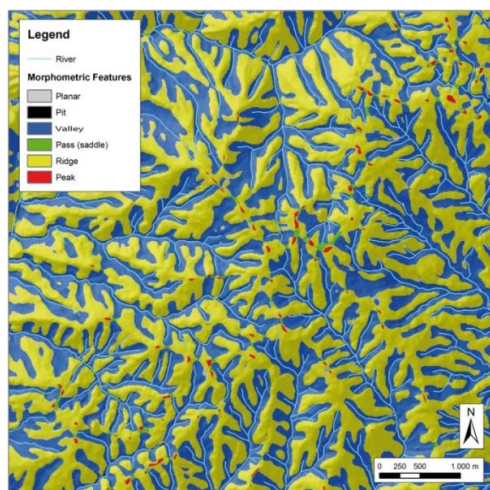


**Figure 5. Classification of the landscape into morphological classes according to Wood (2009) (modified
from Tacconi Stefanelli et al., 2020).**

236   The effectiveness of distinguishing different morphological landforms of this automatic tool is greater in
237   mountainous regions characterized by significant differences in elevation, compared to flat areas where
238   distinctions between landforms are less evident. The accuracy of the output is directly correlated with the resolution
239   of the DEM, which should ideally be about a few meters. Coarser resolutions result in landslide volumes with a
240   corresponding level of uncertainty.

241   The following phase is to provide to each river stretch a value of a valley width, W_v. A series of 1 km long lines
242   (hereafter "transects") are generated, perpendicular to the stretches of the river network, outdistanced by 500
243   meters apart from each other. The created river valley polygons are used to "cut" the transects and then to measure
244   the distance between the two river valley borders through the length of the cut transects.

245   Next, the valley widths (W_v) for each segment of the river are determined by assigning them an average value
246   based on N perpendicular transects, excluding the extreme values (maximum and minimum, respectively W_{max}
247   and W_{min}), as in the fallowing equation:

248    $W_v = \left(\sum_{i=1}^{n} W_i - W_{min} - W_{max} \right) \frac{1}{n-2}$ (4)

249   By utilizing an updated database of landslide polygons, in the step III of Figure 4 it is possible to determine if a
250   reactivated landslide is big enough to cause a complete river blockage thanks to the comparison with the boundary
251   volumes of V_l' (below which a landslide cannot completely block the river) and V_l'' (above which the river valley
252   is dammed for sure). A reactivated landslide should follow a downhill path akin to the flow of surface water.
253   Within each slope, the drainage directions can be easily determined along the river network using a GIS software.
254   Each mass movement can then be linked to the corresponding river stretch it would reach if reactivated based on
255   their corresponding draining surfaces.

	$V_1 > V_1'$ (V_1'')	$V_1 < V_1'$ (V_1'') $< V_1 * 1.2$	$V_1 < V_1'$ (V_1'')
Classification Value	2	1	0

286

287

Figure 7. Predisposition matrix used for the assignment of the damming predisposition intensity to the
mapped landslides (after Tacconi Stefanelli et al., 2020).

290 Even if the proposed method is objective, it is certainly not free from uncertainties and errors. The 20% increase

291 applied to mapped landslide volumes to reduce underestimation errors can in turn produce false positives for

292 overestimation errors. While a false positive is preferable to a false negative (according to a principle of prudence),

293 too many high-risk false positive cases "spread" an unreal risk throughout the area instead of concentrating it in

294 sites of real risk. Therefore, it can be assumed that the landslide bodies which have previously reached the valley

295 floor have already generated most of their effect on the river network (Strom, 2010) or have had no effect, spending

296 their potential risk component. These landslides, also with a volume higher than V_1' and V_1'' and therefore classified

297 with Very High dam predisposition, even if reactivated probably will not produce any further effect in the future.

298 For these reasons, it was decided to downgrade the classification of those landslides that intersect the river network

299 by reducing its position of the classification of damming predisposition by one class.

300 Using the W_v value for each river stretches estimated during the step III of Figure 4, in the last step (IV) the two

301 boundary landslide volumes, namely "Non-formation volume" and "Formation volume" (V_1' and V_1''), can be

302 estimated by applying the equations of the "Non-formation" (Eq. (2)) and "Formation" lines (Eq. (3)). These two

303 values can be used both to classify the damming susceptibility of the river network (for new landslides) and of the

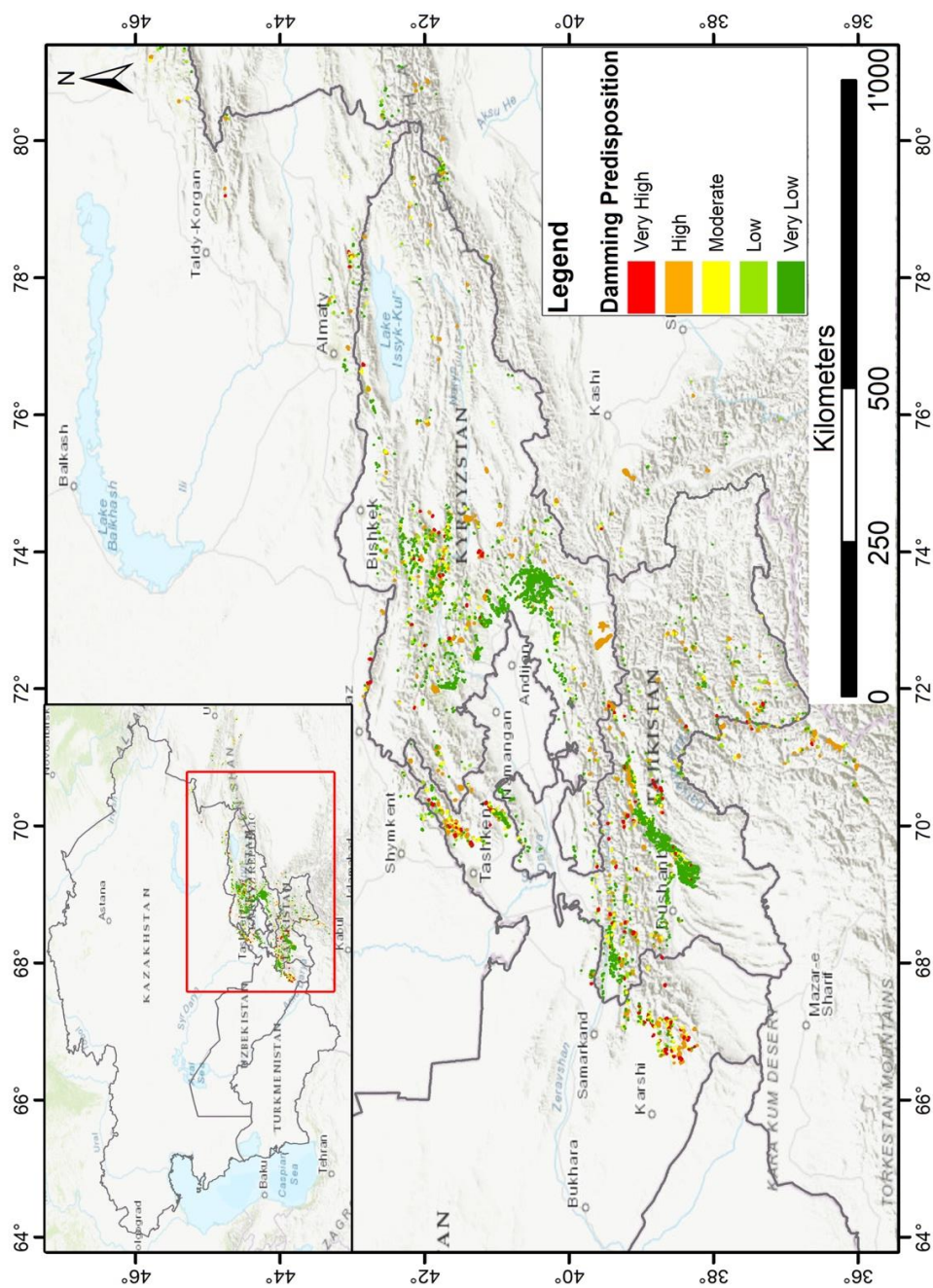
304 landslides inventory (for their reactivation). For the first case, the computed volume values V_1' and V_1'' embody

305 the required volumes of a new landslide to have a potential or certain (with 99% of confidence) obstruction for

306 each river stretches.

4   Results

The mapping methodology was applied to all the studied territories of the Central Asia region in order to analyze and evaluate the results. Two smaller basins, the upper Pskem river and the Fergana valley, were selected to verify the reliability at a catchment scale of the results obtained from a methodology applied on a national scale. The assessment of damming predisposition on the available landslide inventory on the Central Asia regions is shown in the map of Figure 8, while a closer detail is reported in Figure 10 showing the Kyrgyz Republic territory. The number of landslides (644 cases) classified with Very High damming predisposition from the whole inventory before the class reduction due to the river intersection was unjustifiably and unreasonably large posing excessive concern and risk perception. After the change, this number decreased by 75% up to 166 cases, a high number but more reasonable concerning such a large area. In the class distribution of the damming predisposition shown in Figure 9 the most frequent class is the Very Low, with 81% of the whole database, followed by the Low and High classes both with 6% and the remaining percentage divided among Moderate (5%) and Very High (2%).

This distribution is quite coherent with the landslide volumes frequency distribution since it is reasonable to associate landslides with very low volume (83%, shown in Figure 6) with those classified with very low susceptibility (81%, Figure 9). The landslides classified with the higher values of susceptibility (Moderate, High, and Very High with a total of 13%) instead do not only include landslides with higher volumes (more than 100 million m^3 representing 4% of the total), meaning that also even smaller landslides can potentially block narrow river stretches in these regions.
Figure 8. Map of Central Asia Landslide Damming Susceptibility. Basemap source: Esri, HERE, Garmin

Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance

328   Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User

329   Community.

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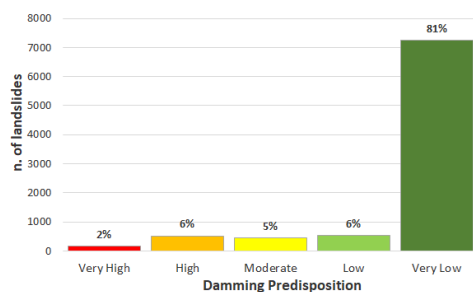
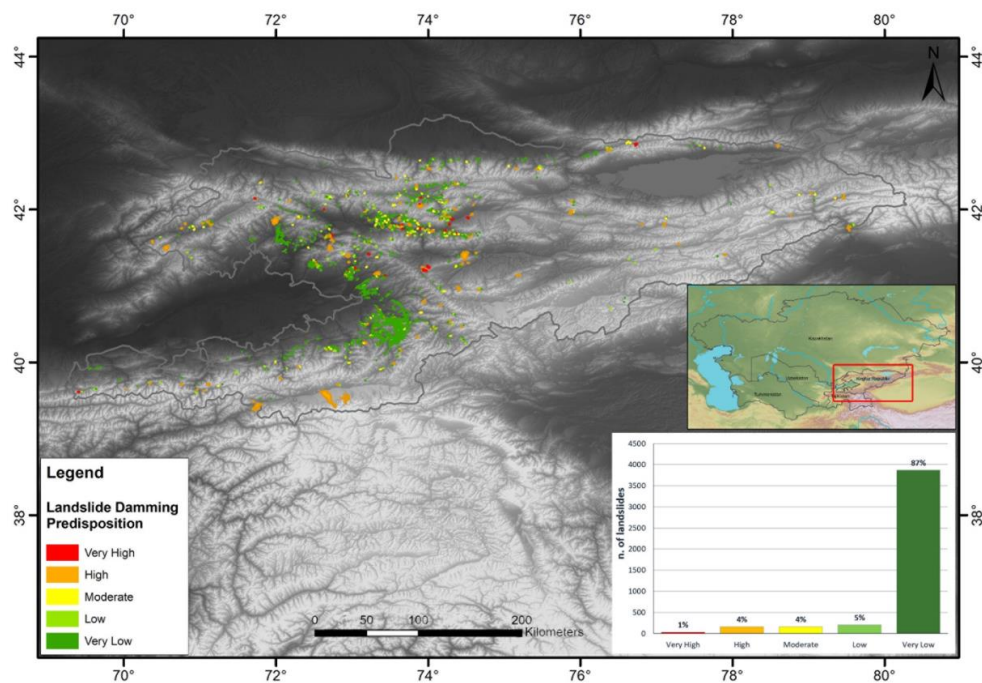

332   **Figure 9. Classes distribution of the damming predisposition for landslides reactivation.**

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334   **Figure 10. Map of Damming Predisposition by landslides reactivation in Kyrgyz Republic territory.**

335   Topographic base from NASA's SRTM project (Far and Kobrick, 2000).

336   Concerning the damming susceptibility caused by new landslides along all the river network in the study area, two

337   different maps of the river networks have been produced using the Non-formation and Formation volumes values.

338   Although counterintuitive at first glance, these maps provide complementary information. The former provides

339   the volumes of landslides that surely create an obstruction, while the latter the volumes below which it definitely

does not form. According to the preliminary steps of the described methodology, in the river stretches running in flat areas (slope degree less than 4° representing the 88.4% of the entire river network) the analysis has been not applied, due to the extreme unlikelihood that a complete obstruction will occur in such areas. The magnitude of the damming susceptibility of the river networks has been classified in five classes and shown in Figure 11 and Figure 14. The five volumes intervals describing damming susceptibility were decided according to general value distribution of landslides volumes and an expert judgement. Since small landslides are more frequent than large ones, as reported in Figure 6, the lower is the landslide volume required to realize an obstruction, the higher is the magnitude. In the map of damming susceptibility related to the "Non formation", reported in Figure 11, the central classes, Moderate and Low are the most frequent with 4.4% and 5.8% respectively, as reported in Figure 12. This means that in most of the river stretches in the study area the minimum landslide volume able to potentially dam the riverbed is between the limit values of the two classes, from 2,5 to 25 million m^3. The following most frequent class is the Very Low with 0.8% and only a very small portion of the river stretches are classified as High and Very High with just 0.4% and 0.2% with a required landslide volume less than 2.5 million m^3. An example of close-up on the Tajikistan territory is reported in Figure 13.

Regarding the map of damming susceptibility related to Formation values, the map in Figure 14 shows slightly different results. The most frequent classes are the two lower ones, Low and Very Low with 4.4% and 6% respectively, as described in Figure 15. Only just the 0.3% and 0.4% fall in the classes Very High and High damming susceptibility. A close-up on the Kyrgyz Republic is reported in Figure 16.

The results of the classification for the river networks of each state are shown in Figure 17 to Figure 21. The landslides of Tajikistan, Kyrgyz Republic, Uzbekistan and Kazakhstan regions have been classified according to damming predisposition (Figure 17-a., Figure 18-a., Figure 19-a. and Figure 20-a). In the Turkmenistan territory, it was not possible to assess any damming predisposition by landslides reactivation since the absence of any available landslide inventory. The results of Uzbekistan and Kazakhstan regions (Figure 19-a. and Figure 20-a.) are a bit different from Kyrgyz Republic and Tajikistan regions due to the different availability of landslide inventories and a different reliefs orographic structure and valleys morphology the formers national territories. As already mentioned, for a clearer comprehension of the damming susceptibility classification of the river network at the national level, the river stretches flowing in lowlands have not been considered in the analysis. Concerning the Damming Susceptibility of Non-Formation (Figure 17-b., Figure 18-b., Figure 19-b., Figure 20-b. and Figure 21-a.), the most frequent are Low and Moderate classes, followed by Very Low class. Fortunately, only very few river stretches have been classified as Very High and High. For the Damming Susceptibility of Formation (Figure 17-c., Figure 18-c., Figure 19-c., Figure 20-c. and Figure 21-b.) most of the rivers fall into Very Low and Low classes, followed by Moderate class. Also in this case, only very few river stretches have been classified as Very High and High. The results of the Tajikistan territory are quite similar to the Kyrgyz Republic and Uzbekistan with which it shares a similar orographic distribution and morphology of the territory. Turkmenistan and Kazakhstan show a slightly different distribution with higher percentage on Moderate class in the Damming Susceptibility of Non-Formation and Low class in the damming susceptibility of Formation.

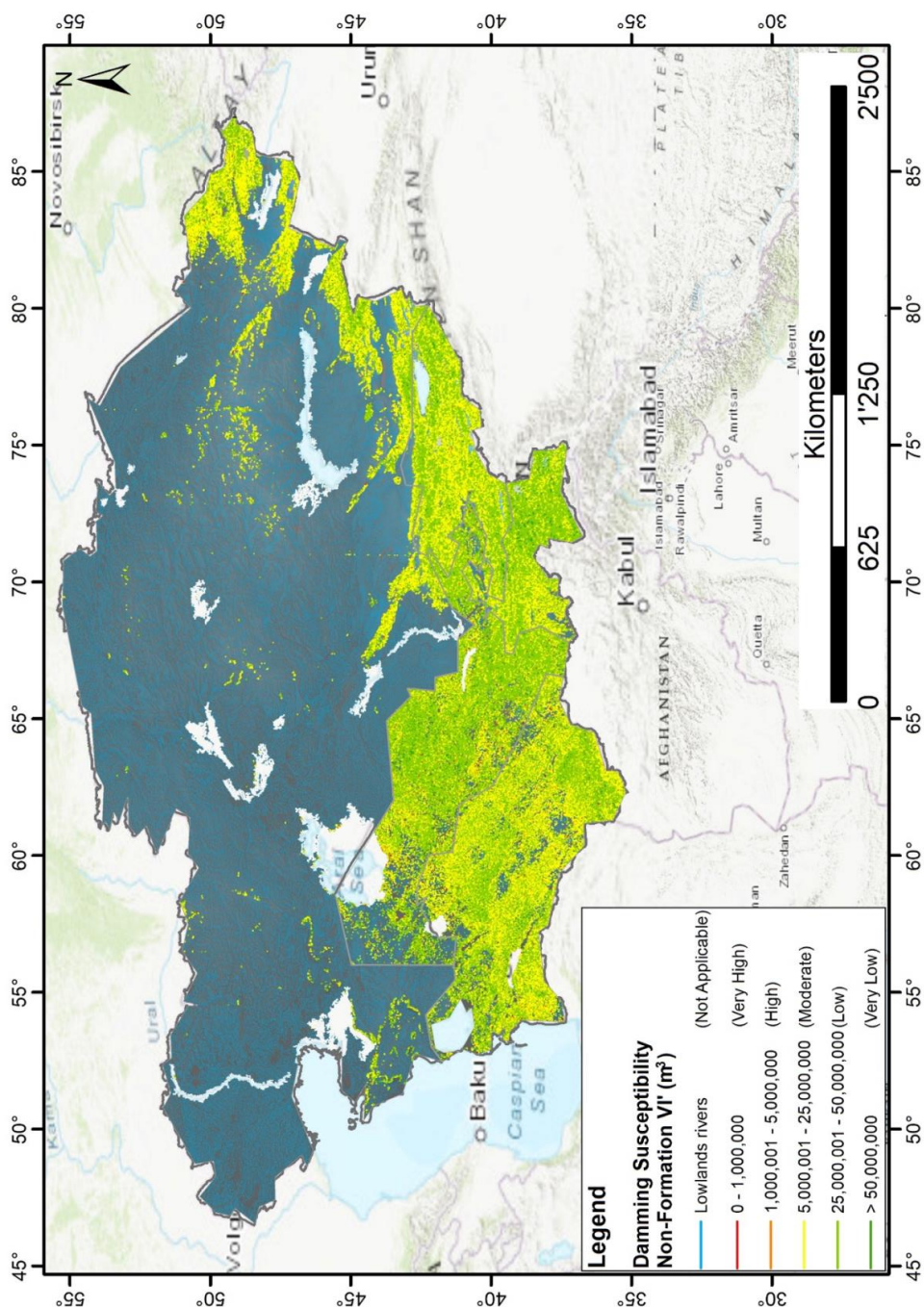
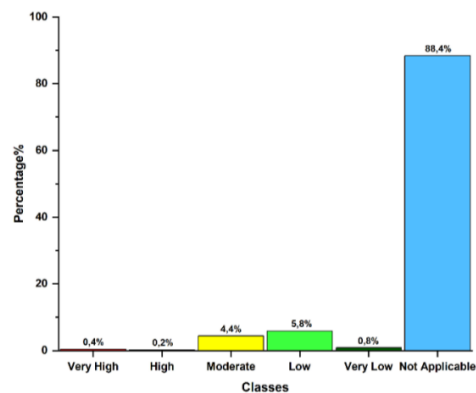
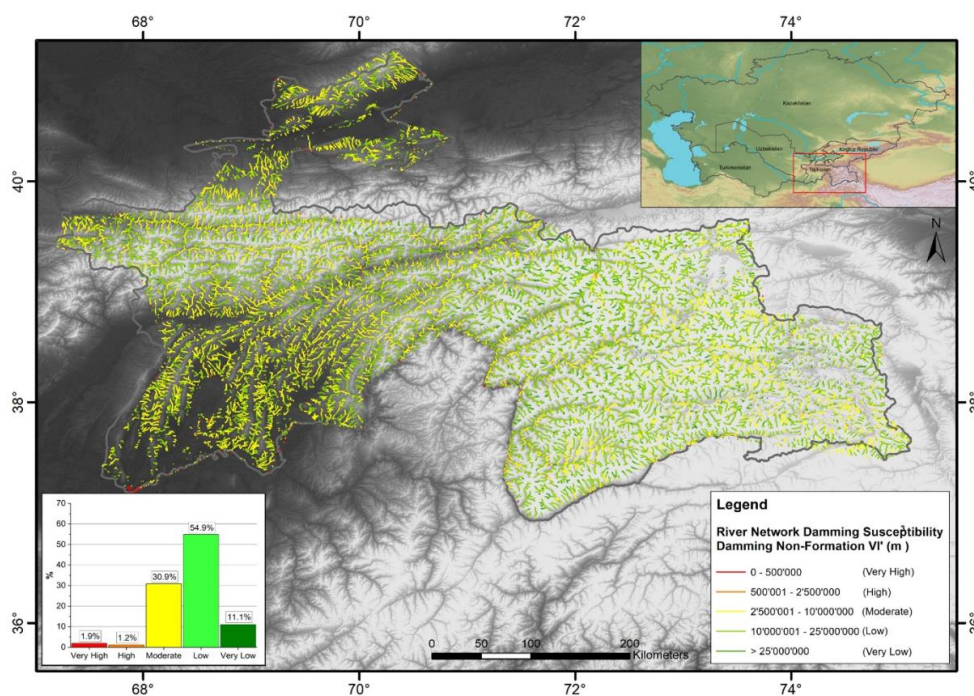



Figure 11. Damming susceptibility map of non-formation of river stretches by new landslides in the region. River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap,

379 increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri

380 Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

381

382 **Figure 12. Distribution of the damming susceptibility in the study area by new landslides related to Non**

383 **formation boundary values.**

384

385 **Figure 13. Damming Susceptibility Map of non-formation of river stretches by new landslides in**

386 **Tajikistan.** River network database from Coccia et al., (in prep.). Topographic base from NASA's SRTM

387 project (Far and Kobrick, 2000).

388

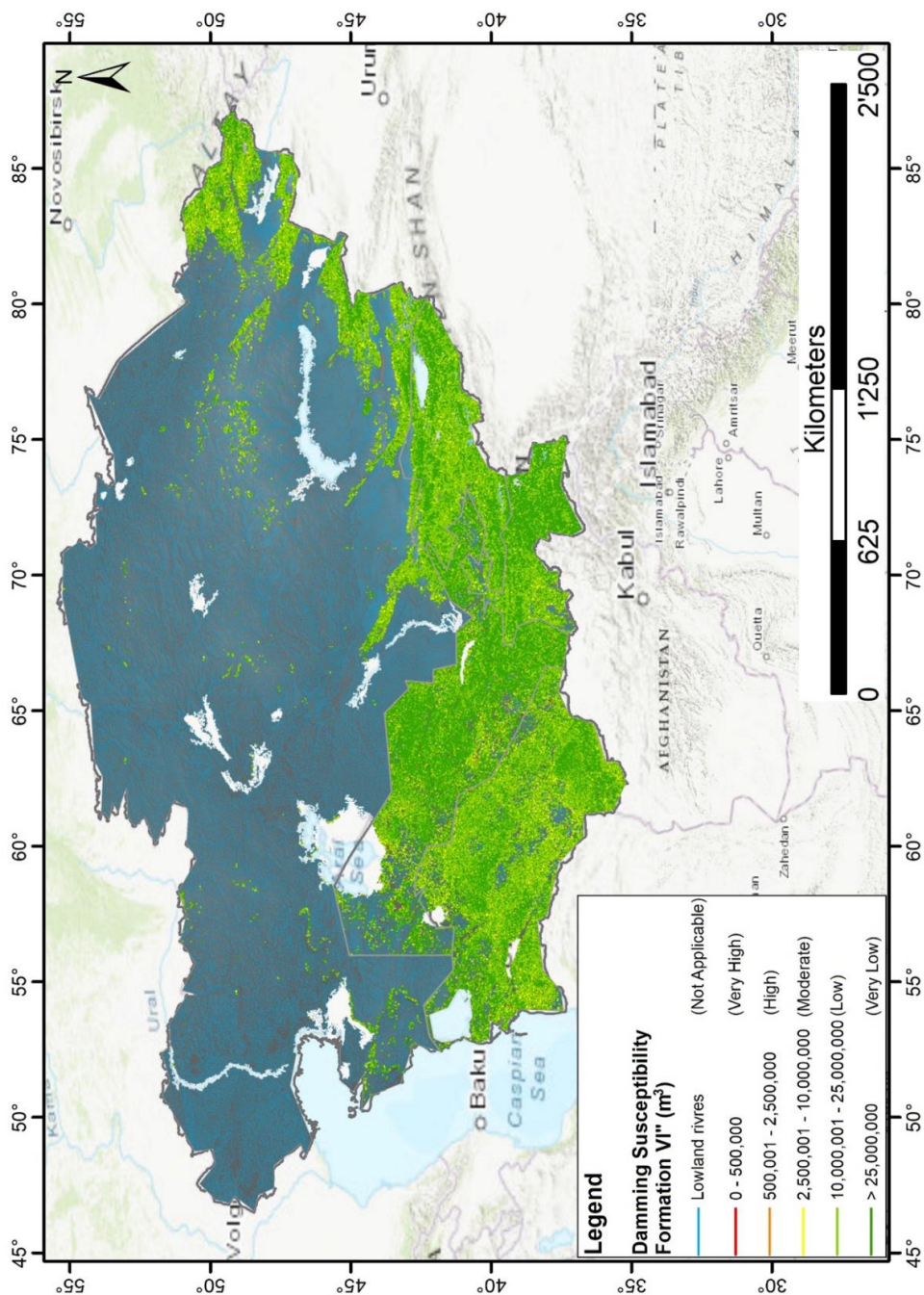



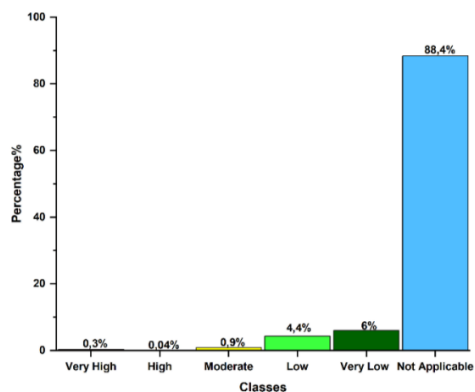
Figure 14. Damming Susceptibility Map of Formation of river stretches by new landslides in the region.

River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap, increment

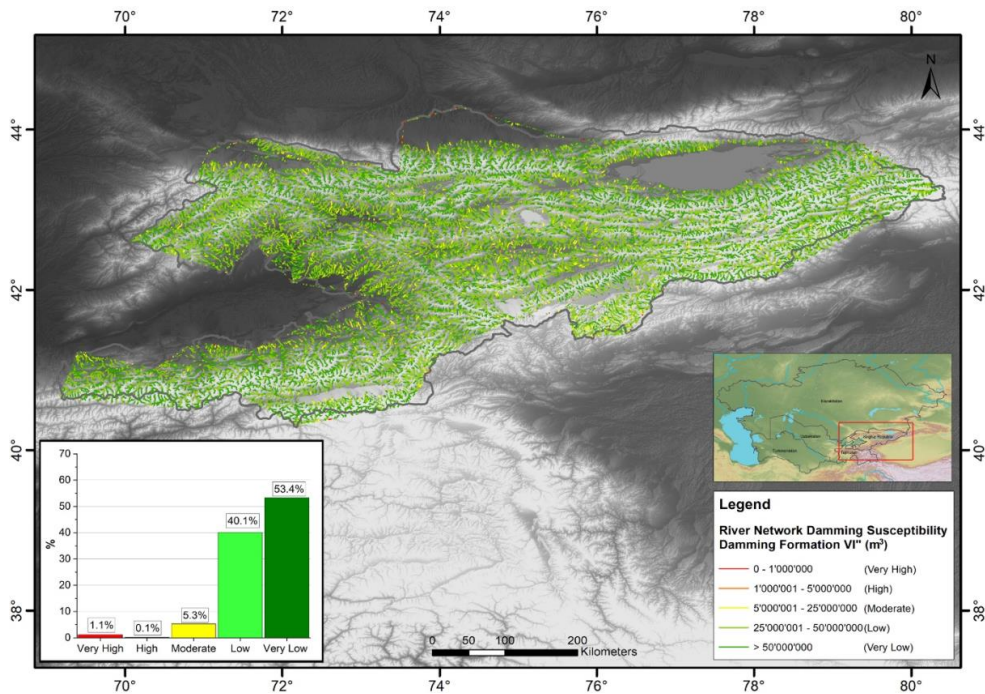
389

390

391 P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan,

392 METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

393

394 **Figure 15. Distribution of the Damming Susceptibility in the study area by new landslides related to**

395 **Formation boundary values.**

396

397 **Figure 16. Damming Susceptibility Map of formation of river stretches by new landslides in the Kyrgyz**

398 **Republic territory.** River network database from Coccia et al., (in prep.). Topographic base from NASA's

399 SRTM project (Far and Kobrick, 2000).

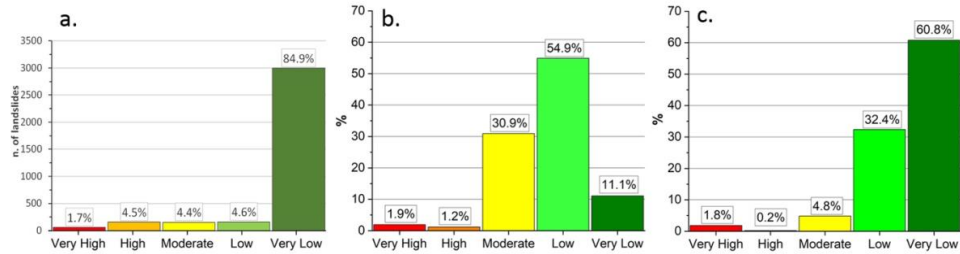



Figure 17. Classes distribution in Tajikistan of the Damming Predisposition for landslides reactivation (a.), Damming Susceptibility of Non-Formation (b.) and of Formation (c.) for new landslides.

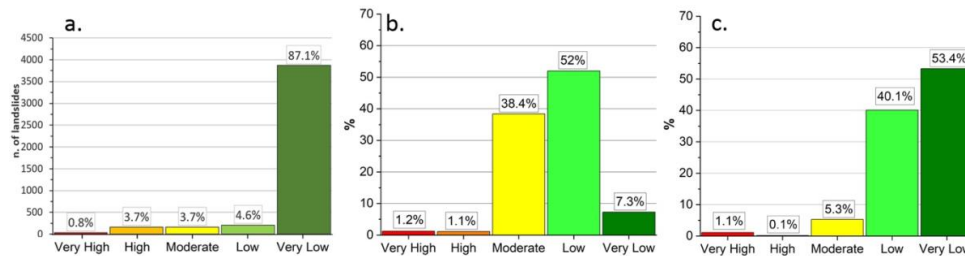



Figure 18. Classes distribution in the Kyrgyz Republic of the Damming Predisposition for landslides reactivation (a.), Damming Susceptibility of Non-Formation (b.) and of Formation (c.) for new landslides.

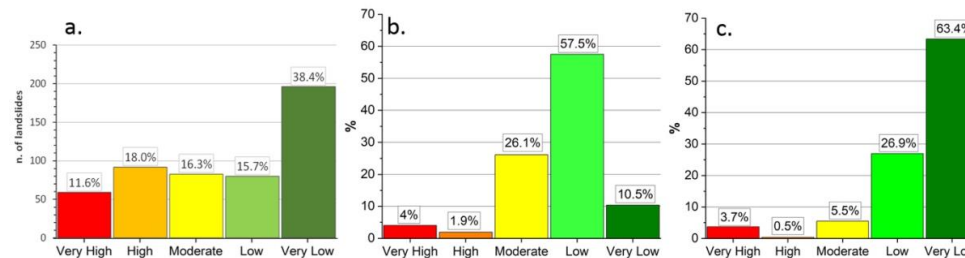



Figure 19. Classes distribution in Uzbekistan of the Damming Predisposition for landslides reactivation (a.), Damming Susceptibility of Non-Formation (b.) and of Formation (c.) for new landslides.

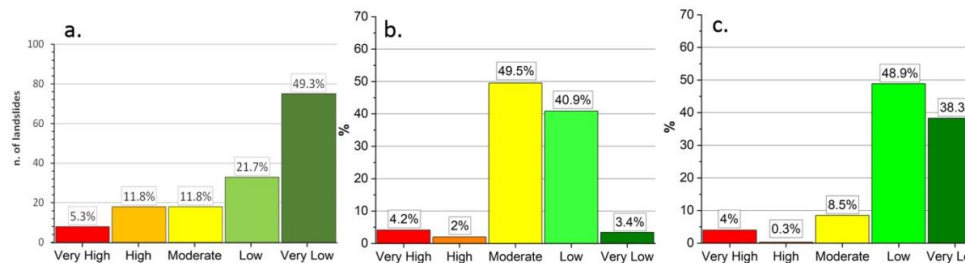



Figure 20. Classes distribution in Kazakhstan of the Damming Predisposition for landslides reactivation (a.), Damming Susceptibility of Non-Formation (b.) and of Formation (c.) for new landslides.

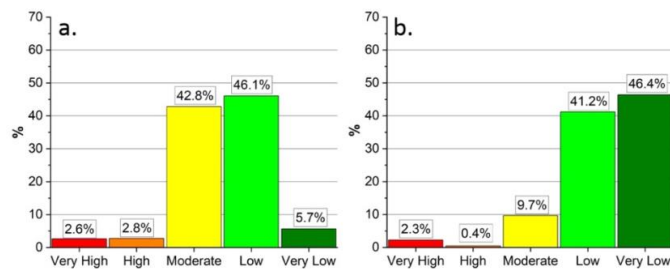



Figure 21. Classes distribution in Turkmenistan of the Damming Susceptibility of Non-Formation (a.) and of Formation (b.) for new landslides.

4.1 Upper Pskem river valley (Uzbekistan)

The Pskem river, locate in the Tashkent region of Uzbekistan, is a right-hand tributary of the Chirchik River that is the feeder of the Syr Darya river basin (in the Western Tien-Shan). The river originates from the confluence of the Maidantal and Oygaing rivers and is one of the main tributaries of the Charvak Lake (Semakova et al., 2016). This artificial lake is central for the local economy for its functions as reserve for fishing and water, as well as a source of hydroelectric energy and because of that various villages arise around it and downstream. The formation of a natural obstruction and an upstream impoundment in the Pskem basin could be a serious threat due to the possible instability of the earth dam and for the possible catastrophic cascade effects that its collapse could have downstream on the artificial basins and their 168 meters high earthfill dam.

With a careful observation of the map of Damming Predisposition by landslides reactivation in the lower Pskem basin in an area of 443 km^2 (Figure 22), some of the 53 mapped landslides should be subjected to further study. Among all, most landslides were classified with a Very Low and Low predisposition value, respectively 21 and 11 cases (39.6% and 20.8%), and only 4 landslides with a Very High value (7.5%), 10 with High (18.9%) and 7 with Moderate (13.2%). Landslides named A, B, C, D and E in Figure 22, if reactivated will potentially cause an obstruction of the main river section of the Pskem, being classified the first three and the latest two respectively with High and Very High damming predisposition. As shown in Table 2, the volumes of all these landslides are way bigger than the boundary volume of Non-Formation and Formation from Figure 23 and Figure 24. It is important to notice that the landslides A, B and C are laid down in the valley floor, meaning that in the past they had probably already dammed the river in that point, and the classification of their damming predisposition have been reduced by one, from Very High to High. Due to the considerable volumes of the landslides in the basin and the presence of landslides that have probably already blocked the river in the past, this relatively small area is certainly worthy of attention.

Table 2. Landslides volumes and damming parameters W$_v$, V$_l$', V$_l$'' of the landslides in Figure 20 computed using the described method.

Landslide	V_1 - Landslide volume (m³)	W_v – River Width (m)	V_1' - Volume of Non-formation (m³)	V_1'' - Volume of Formation (m³)
A	200.000.000	300	2.600.000	16.200.000
B	12.000.000	235	1.500.000	10.000.000
C	34.000.000	318	3.000.000	18.200.000
D	73.000.000	513	10.100.000	47.400.000
E	61.000.000	575	13.500.000	60.000.000

439

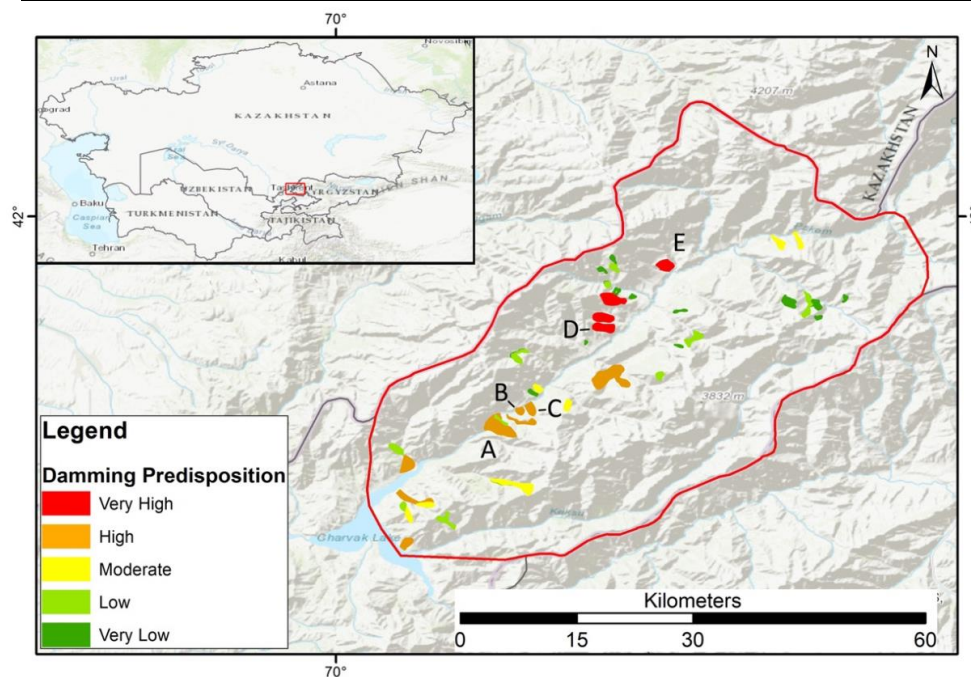



Figure 22. Map of Damming Predisposition by landslides reactivation in the lower Pskem basin. Basemap source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

The obstruction of the Pskem river by one of these landslides would cause an upstream impoundment with a surface from 2 to 10 km² or more, depending on the dam position and height. The dam collapse could release a catastrophic flooding wave with destructive effects in the downstream areas. In the worst scenario, even the earthfill dam located few kilometers downstream could be seriously damaged with unpredictable effects. Since the reliability of this mapping method is strictly correlated to the quality of the input data, when the used DEM has a coarse resolution, in similar cases of possible risk to people's life it is always advisable to do a second "manual

450 check" even using some free satellite imaging services like Google Earth. In fact, when the DEM resolution is too
451 rough, the GIS tool used in this methodology to evaluate the extension of the riverbed morphologic unit can
452 produce inconsistent and incorrect results, causing improper damming susceptibility evaluations. The results of
453 the measurements on Google Earth orthophotos in Table 3 show that the difference between the river width values
454 calculated with the mapping method (W_v) and those measured on Google Earth (W_{vGE}) can in some cases be
455 substantial modifying the calculated boundary volumes V' and V'', although in this case they do not modify
456 drastically the final classification of the five landslides.

457 The river network of the upper Pskem valley have been also classified producing the maps of Damming
458 Susceptibility of Non-formation and Formation (Figure 23 and Figure 24 respectively). Concerning the Damming
459 Susceptibility Map of Non-formation (Figure 23), the most frequent are Low and Moderate classes with 65.1%
460 and 22.6% respectively, followed by Very Low class with 11.1%. Only just 1.3% have been classified as High and
461 0.0% as Very High. For the Damming Susceptibility Map of Formation (Figure 24) most of the rivers fall into
462 Very Low and Low classes with 69.8% and 27.7%, followed by Moderate class with 2.1%. Only 0.4% have been
463 classified as High and 0.0% as Very High.

464 **Table 3. Damming parameters W_{vGE}, $V_l'_{GE}$, $V_l''_{GE}$ of the landslides in Figure 22 computed with Google**
465 **Earth observation.**

Landslide	W_{vGE} – River Width (m)	$V_l'_{GE}$ - Volume of non-formation (m³)	$V_l''_{GE}$ - Volume of Formation (m³)
A	415	6.000.000	31.000.000
B	310	2.800.000	17.300.000
C	260	1.800.000	12.100.000
D	530	11.000.000	50.000.000
E	450	7.300.000	36.500.000

466

467 The general damming susceptibility of the valley is low but a singular river stretch, marked by a black circle in
468 Figure 23 and Figure 24, classified with High susceptibility in both maps should be carefully evaluated. This river
469 part is clearly noticeable in the middle of the area along the main river path, a bit upstream from the landslides
470 named B and C. The high classification values mean that geographically in that point the valley width undergoes
471 a shrinkage and for this reason even a relatively small landslide generated from the surrounding slopes can create
472 an obstruction, therefore it would be worthy of a more detailed investigation.

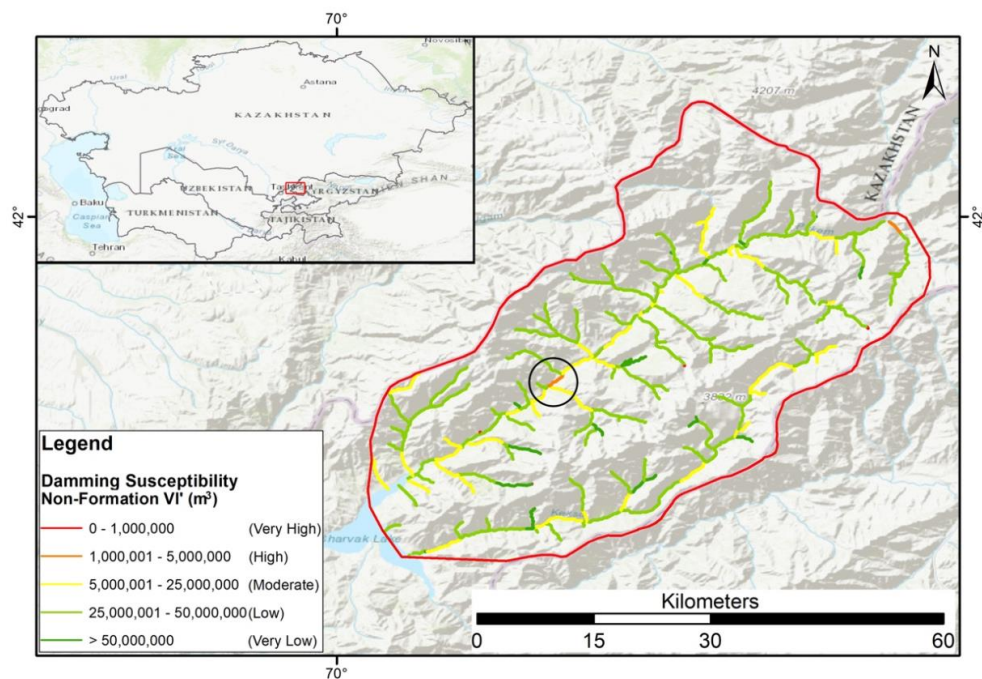



Figure 23. Damming Susceptibility Map of Non-formation of river stretches by new landslides in the lower Pskem basin. The black circle highlights a river stretch with unusually high values. River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

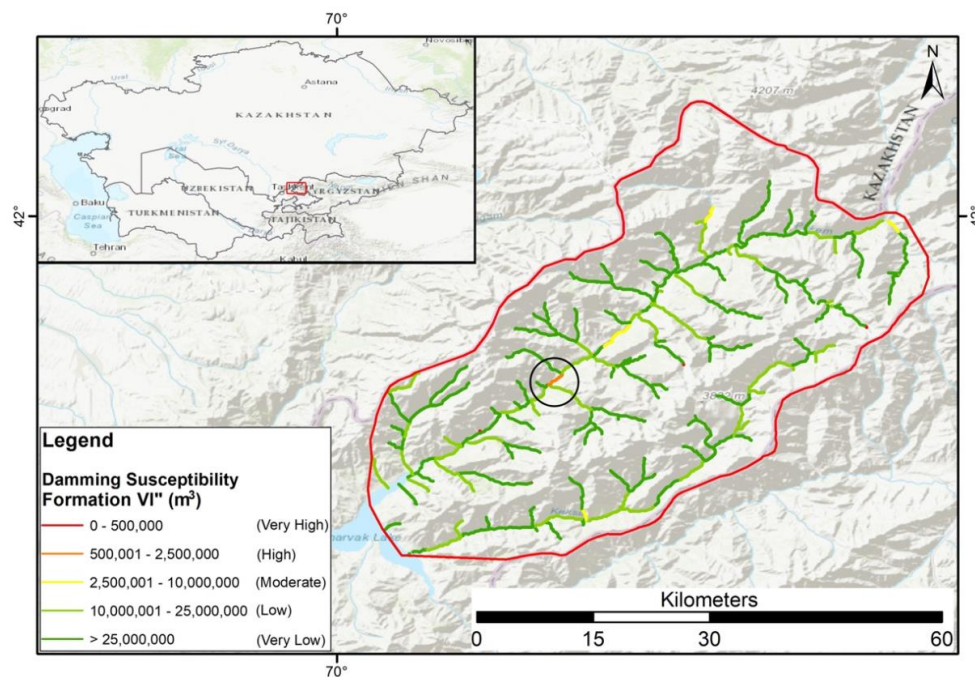



Figure 24. Damming Susceptibility Map of Formation of river stretches by new landslides in the lower Pskem basin. The black circle highlights a river stretch with unusually high values. River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

4.2 The Fergana valley mountainous rim (Tajikistan-Kyrgyz Republic-Uzbekistan)

The Fergana valley is one of the largest intermountain depressions in Central Asia located between Uzbekistan, Kyrgyz Republic, and Tajikistan. It hosts two main rivers, the Naryn and the Kara Darya, which join together to form the Syr Darya. In this populated area landslide activity is recurrent, causing every year damage to infrastructure and loss of human life, and triggered by complex interactions between multiple factors such as tectonic, geological, morphological and meteorological (Danneels et al., 2008; Schlöegel et al., 2011; Piroton et al., 2020). The mapping methodology have been applied also to the Fergana valley and a total of 3370 landslides, coming from various data sources, have been classified as shown in Figure 25. Comparably to the classification result of the entire inventory (Figure 8) most of the cases (94%) have a Very Low damming predisposition, followed by Low and Moderate (with 2.5% and 1.8% respectively) as reported in Table 4. Just very few landslides fall into High and Very High classes (with 1.4% and 0.3% respectively). For the classification of the river network of the Fergana valley, the maps of Damming Susceptibility of Non-formation and Formation have been produced (Figure 26 and Figure 27 respectively). As a method with a multi-scale approach, in such large areas, this damming

499 susceptibility method is suitable to provide territorial planning suggestions rather than indications on single

500 interventions at local scale. The overall damming predisposition of the Fergana valley is quite low, considering

501 the presence of 3370 mapped landslides in total, even if there are few landslides (10) classified with Very High

502 damming predisposition which should be studied with more attention through localized analysis of damming

503 susceptibility to ensure that downstream areas are not at risk and therefore require a specific monitoring.

504 Table 4 have been reported the distribution of the percentages of the damming susceptibility classes of those river

505 stretches that are not running in flat areas, since these lowland rivers represent 53.6% of the total. Concerning the

506 Damming Susceptibility Map of non-formation of the remaining river stretches (Figure 26), the most frequent are

507 Low and Moderate classes with 53.4% and 36.2% respectively, followed by Very Low class with 7.0%. Only just

508 2.1% and 1.3% have been classified as Very High and High. For the Damming Susceptibility Map of Formation

509 (Figure 27) most of the rivers fall into Very Low and Low classes with 54.5% and 38.1%, followed by Moderate

510 class with 5.2%. Only 1.9% and 0.2% have been classified as Very High and High respectively.

511 **Table 4. Distribution of Damming Susceptibility classes on existing landslides (Figure 25) and on the river**

512 **stretches for non-formation (Figure 26) and Formation of new landslides (Figure 27).**

Damming Susceptibility	Landslides		non-formation	Formation
	n.	%	%	%
Very High	10	0.3%	1.9	1.7
High	48	1.4%	1.2	0.2
Moderate	61	1.8%	7.0	5.3
Low	83	2.5%	53.2	38.8
Very Low	3168	94.0%	6.7	54.0

513

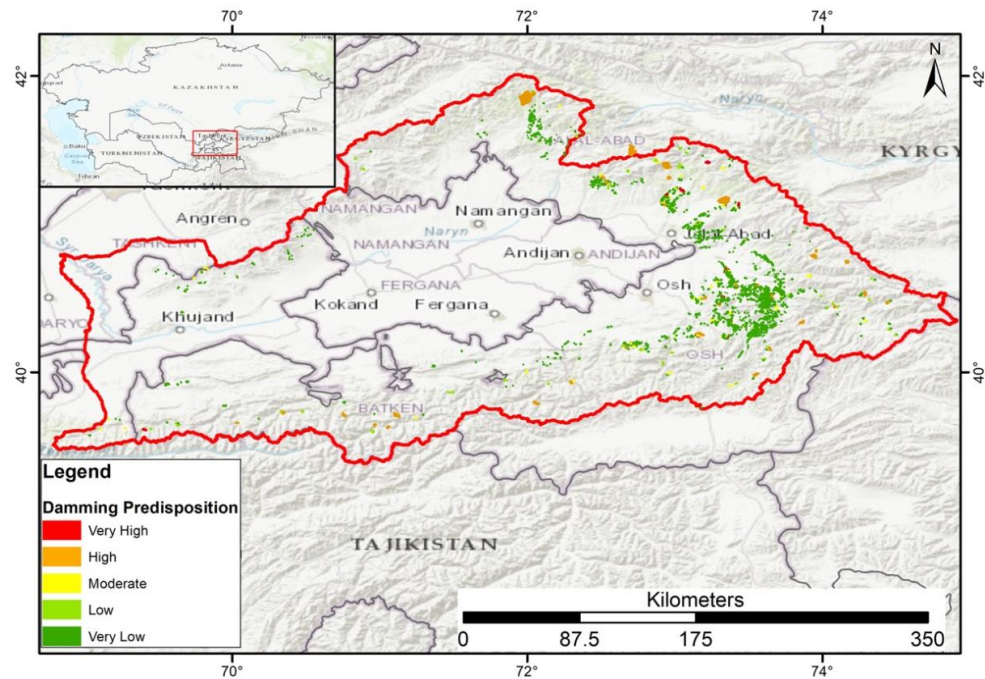



Figure 25. Map of Damming Predisposition by landslides reactivation in the Fergana valley. Basemap

source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN,

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and the GIS User Community.

519
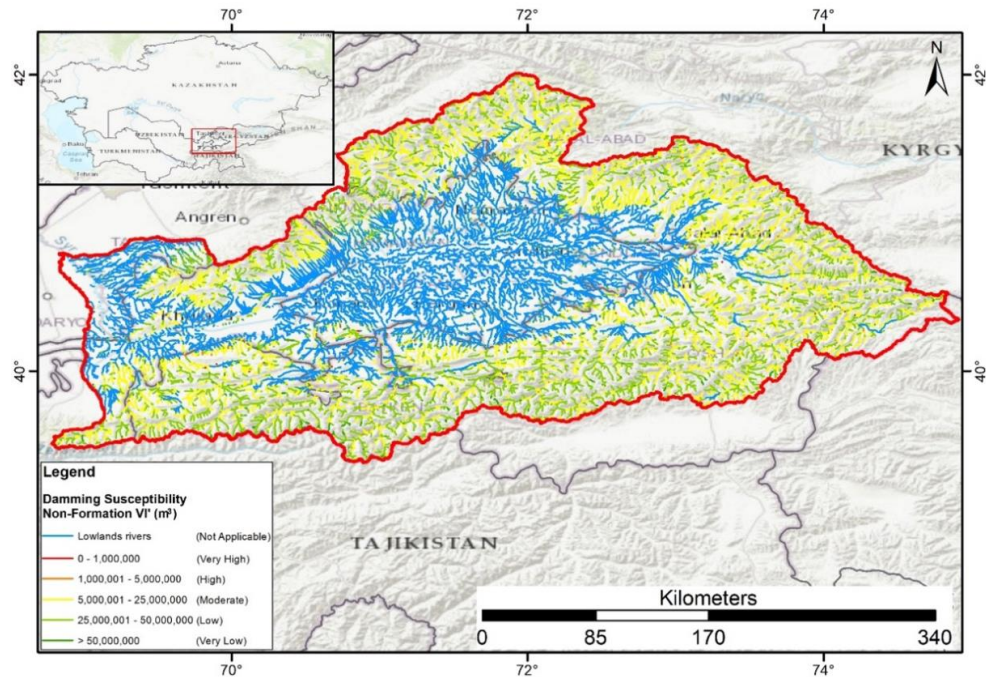
Figure 26. Damming Susceptibility Map of Non-formation of river stretches by new landslides in the Fergana valley. River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

Figure 27. Damming Susceptibility Map of Formation of river stretches by new landslides in the Fergana valley. River network database from Coccia et al., (in prep.). Basemap source: Esri, HERE, Garmin Intermap, increment P Corp, GEBCO, USGS, FAO, NPS, NRCAN, GeoBase, IGN, Kadaster NL, Ordnance Survey, Esri Japan, METI, Esri China (Honk Kong), © OpenStreetMap contributors, and the GIS User Community.

5   Discussion

During the application of the damming mapping methodology, the main issues encountered was the extremely wide study area, the amount of data and the processing time required. The used mapping methodology based on the MOI equations (Eq.(1)), was originally designed to assess the damming susceptibility at basin/regional scale (Tacconi Stefanelli et al., 2016; 2020), where the morphological parameters essential for the correct application of the tool proposed by Wood (2009) must be correctly found to have an accurate river width required in the MOI equations (Eq.(1)). This time-consuming phase has been simplified in this research, according to the wide dimension of the study area, taking into account not the basins but the different states in the Central Asia region. Furthermore, the results quality is directly proportional to the resolution and quality of the input data, which on the other hand is inversely proportional to the processing time. In this regard, a further criticality of this process is the reliability on the landslides volumes assessment method, since a higher quality of landslides data (sliding geometry and depth) allows the application of a more accurate volume calculation and therefore a better final result.

543   Thus, even if the results are not always highly reliable at local scale, requiring many in-depth specific studies in

544   the areas identified with the higher predisposition, they can be undoubted useful in very large countries to adopt

545   risk reduction measures, for planning purposes and for land development management.

546   Considering the size of the area, in Figure 10 the number of landslides classified with Very High damming

547   predisposition (166 cases) is reasonable in absolute value, even if a bit high if compared with the total number of

548   landslides present in the inventory (8910 cases). Without a detailed study it is not possible to say how many of

549   these are false positives or not, however it is important to remember that this type of risk mapping methods gives

550   information on if and where, not when these events may occur.

551   The two maps of damming susceptibility (Figure 11 and Figure 14), while not providing probability values as done

552   by Tacconi Stefanelli et al. (2020), offer information (the volumes of landslides) that can be more easily spent and

553   interpreted even by operators who are not specifically expert, and for this reason have more practical utility.

554   Furthermore, the classification of the river stretches thus produced, not requiring the alpha parameter (linked to

555   the probability of landslide occurrence) as in the original method proposed by Tacconi Stefanelli et al. (2020), it

556   is much easier to obtain and for this reason it can be considered an improvement within a view of wider usability.

557   **6   Conclusions**

558   The price of a river obstruction, in terms of reconstruction and losses on both economic and lives, can be much

559   higher compared with the costs of a proper environmental planning and land-use management. Be able to define

560   the areas with higher risk could considerably lower the costs, allowing to focus the economic resources in effective

561   preventive interventions, planning and monitoring activities.

562   In this work a damming mapping methodology have been proposed and carried out on the Central Asia regions.

563   The used method, originally developed applying the Morphological Obstruction Index at basin scale, have been

564   modified to fit such a large study area and the available data. The improvement of the original method allows a

565   simpler use and the need for less data, more easily available, although the absence of a validation of the results

566   inevitably remains. The main aim of this study was to propose a practical tool to assess where the damming

567   susceptibility from reactivation of mapped landslides and formation of new landslides are higher at national scale.

568   This second result of the mapping damming susceptibility from new landslide can be particularly useful in area of

569   the world where there is a lack of diffuse assessment of landslide activity and incomplete landslide inventories.

570   *Code and data availability*. The landslide dam mapping susceptibility method was implemented by using the cited

571   landslide inventory maps, published by the following authors: Behling et al., 2014, 2016, 2020; Havenith et al.,

572   2015a; Strom and Abdrakhmatov, 2018. The SRTM DEM data are available from https://earthexplorer.usgs.gov/.

573   The river network and other landslide inventories were provided by the SFRAAR project partners: RED (Risk,

574   Engineering + Development – Pavia, Italy), OGS (National Institute of Oceanography and Experimental

575   Geophysics, Seismological Research Center, Trieste, Italy), IWPHE (Institute of Water problems, Hydropower,

576   Engineering and Ecology, Dushanbe, Republic of Tajikistan), ISASUZ (Institute of Seismology of the Academy

577   of Science of Uzbekistan, Tashkent, Uzbekistan), LLP (Institute of Seismology of the Science Committee of the

578   Republic of Kazakhstan, Almaty).

Author contribution. Carlo Tacconi Stefanelli implemented the damming mapping method, William Frodella conceived with Carlo Tacconi Stefanelli the article structure and collected the data, Francesco Caleca supported the method application on part of the study area. Francesco Caleca also performed statistical analysis involving the method results. All the aforementioned Authors contributed to the writing of the article and the figure graphics. Veronica Tofani coordinated the work and reviewed the paper. Zhanar Raimbekova and Ruslan Umuraliev provided environment and geomorphology information and part of the landslide database for Kazakhstan and Kyrgyz Republic.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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