Simulating sea level extremes from synthetic low-pressure systems

Jani Särkkä[1], Jani Räihä[1], Mika Rantanen[1], and Matti Kämäräinen[1]

[1] Finnish Meteorological Institute, P.O.Box 503 , FI-00101 Helsinki, Finland

Correspondence: Jani Särkkä (Jani.Sarkka@fmi.fi)

Abstract. In this article we present a method for numerical simulations of extreme sea levels using synthetic low-pressure systems as atmospheric forcing. Our simulations can be considered as estimates of the high sea levels that may be reached when a low-pressure system of high intensity and optimal track passes the studied region. We test the method using sites located along the Baltic Sea coast and simulate synthetic cyclones with various ~~intensities and~~ tracks. To model the effects of the cyclone properties on sea level, we simulate internal Baltic Sea water level variations with a numerical two-dimensional hydrodynamic model, forced by an ensemble of time-dependent wind and air-pressure fields from synthetic cyclones. The storm surges caused by the synthetic cyclones come on top of the mean water level of the Baltic Sea, for which we used a fixed upper estimate of 100 cm. We find high extremes in the northern Bothnian Bay and in the eastern Gulf of Finland, where the sea level extreme due to the synthetic cyclone reach up to 3.5 meters. In the event that the mean water level of the Baltic Sea has a maximal value (1 meter) during the cyclone, highest sea levels of 4.5 meters could thus be reached. We find our method to be suitable for use in further studies of sea level extremes.

1 Introduction

A large part of the population and infrastructure in the Baltic Sea region is concentrated in the vicinity of the coastline. The potential for extreme sea level conditions must be accounted for when considering the safety of citizens and the need for coastal flood protection. A comprehensive investigation of the present-day sea level variability and of the possible scenarios for mean sea level in the future is essential for the determination of coastal flooding risks (Pellikka et al., 2018; Leijala et al., 2018; Pellikka et al., 2023). The impact of atmospheric conditions on sea level extremes is an important research topic within this context.

Sea level is measured by tide gauges situated on the coastline. Tide gauge records are available for many locations on the Baltic Sea coast only for the past 100-200 years, but several historical floods can be reconstructed from other sources, e.g., historical records or studies of coastal sediments (Rutgersson et al., 2022). Nonetheless, it may be assumed that many sea level extremes during the past 1000 years have left no traces in historical records, especially in historically sparsely populated coastal areas such as around the coast of the Bothnian Bay. In the northern Baltic Sea, the coastline has retreated due to land uplift, and the coastal strip affected by flooding in the past is now further away from the shoreline. Therefore the effects of high sea levels have been reduced by the relative decline in the mean sea level with respect to the shoreline.

The Baltic Sea is a shallow semi-enclosed sea, connected to the Atlantic Ocean through the narrow and shallow Danish Straits. The local short-term sea level variations are due to winds, air pressure, currents, tides, and internal oscillations (seiches). Long-term changes are caused by global mean sea level variations, postglacial land uplift and the water flow between the North Sea and the Baltic Sea through the Danish Straits (Leppäranta and Myrberg, 2009; Johansson et al., 2014). The sea level extremes in the Baltic Sea are highest during winter months, most commonly in November, December and January (Wolski and Wiśniewski, 2023). The short-term sea level extremes (on hourly or daily timescales) are usually caused by moving extratropical cyclones which in favorable atmospheric conditions can strengthen to powerful windstorms (Della-Marta et al., 2009). The annual number of extratropical cyclones and windstorms in the Baltic Sea region has a large inter-annual variability and no significant trends were observed during 1980–2019 (Laurila et al., 2021). The effects of short-term sea level fluctuations are largest in the narrow bays of the Baltic Sea. In particular, the narrow Gulf of Finland is vulnerable to coastal flooding due to windstorms, such as Storm Gudrun, which caused damages in 2005 (Tõnisson et al., 2008; Averkiev and Klevannyy, 2010). Major coastal flooding occurred in St. Petersburg in 1777, 1824, and 1924 (Kulikov and Medvedev, 2017), of which the flood in 1824 was the most extreme, with maximum sea level height 4.21 m. Even in cases where in situ sea level measurements are not available, traces of extreme sea level events can be found in sediment studies (Piotrowski et al., 2017; Leszczyńska et al., 2022). More information on sea level extremes in the Baltic Sea can be found in recent reviews considering the region (Weisse et al., 2021; Rutgersson et al., 2022).

~~In~~ The estimation of extreme sea levels is often made using statistical methods for observed sea level data (Arns et al., 2013). Within the observations, the most extreme weather systems, like tropical cyclones, are often missing or represented by single outliers. This complicates the extreme value analysis of sea level data, especially in the regions affected by tropical cyclones (O'Grady et al., 2022). In the Baltic Sea, even extratropical cyclones may cause high sea levels due to the shallow depth and the shape of the coastline. Variations in the track of the cyclone strongly affect the extreme sea levels, especially in the ends of bays. In this study, ~~we simulate~~ our aim is to estimate the lower limits of extreme sea levels ~~exceeding the values found in tide gauge observations. We~~ in the Baltic Sea caused by cyclones. The extreme value estimates based on statistical methods on observed data are too conservative, if the cyclones causing the highest sea levels have not occurred during the measurement period at the site studied. Thus, we have chosen to use numerical simulations to evaluate extremes in this study. To obtain values for return periods for sea level extremes, statistical analyses of simulation data need to be made (as in Särkkä et al. (2017)), this will be left as a topic of future studies.

To study how high sea levels a storm with optimal track could cause at the Baltic Sea coast, we reconstruct idealised cyclones ~~storms~~ that are physically realistic, but which have not occurred during the era of tide gauge observations. We ~~model~~ describe the cyclones by ~~describing~~ a pressure field with a spatially varying time-dependent function that reproduces the typical cyclone characteristics. ~~This~~ The pressure and wind data is used as input to a numerical sea level model to simulate extreme sea levels. Similar method has been used to study extreme sea levels in the Gulf of Finland (Averkiev and Klevannyy, 2010; Kalyuzhnaya et al., 2015; Apukhtin et al., 2017), and in the Bothnian Bay (Gordeeva and Klevannyy, 2020). In a recent study (Andrée et al., 2022), synthetic uniform wind fields were used in extreme sea levels simulations and an empirical-statistical method was used to model the site-specific relations between wind and extreme sea level. In addition, synthetic cyclones have been used to assess
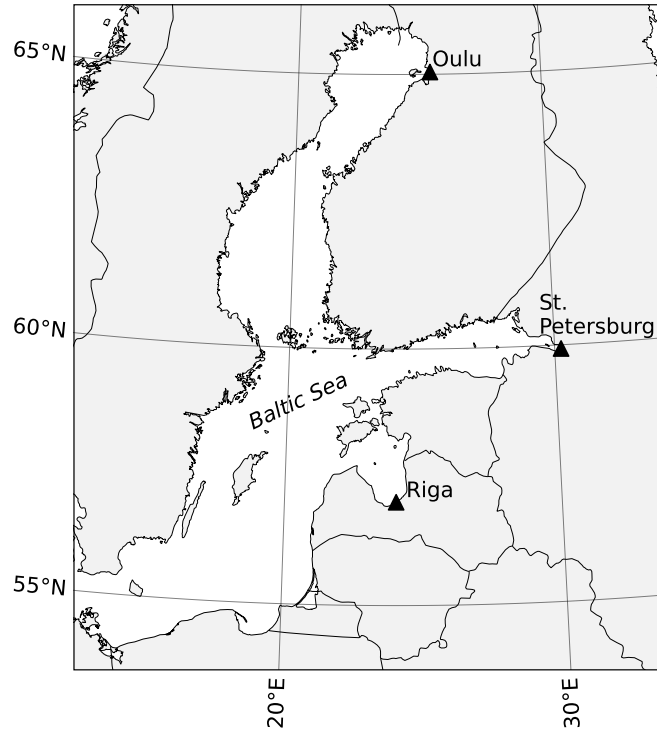
Figure 1. The Baltic Sea and locations (Oulu, St. Petersburg and Riga) where high sea levels were studied.

hazards caused by high sea level and waves in tropical regions ~~(Leijnse et al., 2022)~~(Leijnse et al., 2022; O'Grady et al., 2022)
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~~In this study, we simulate sea levels at three locations on the Baltic coast using synthetic cyclones as an atmospheric input for the sea level model. We choose these locations on the basis of where particularly high sea levels are expected, representing~~
65 ~~different sub-basins of the Baltic Sea: Oulu for the Bothnian Bay, St. Petersburg for the Gulf of Finland, and Riga for the Bay of Riga.~~

2 Data and methods

2.1 Synthetic atmospheric forcing

To model the sea levels associated with our synthetic cyclones, we need information on the atmospheric conditions for the
70 forcing of the sea level model. To this end, we use a method where the mean sea level pressure and surface winds of the cyclone are calculated from the pressure field of a cyclone propagating with a constant velocity. The pressure field has the form of a Gaussian function
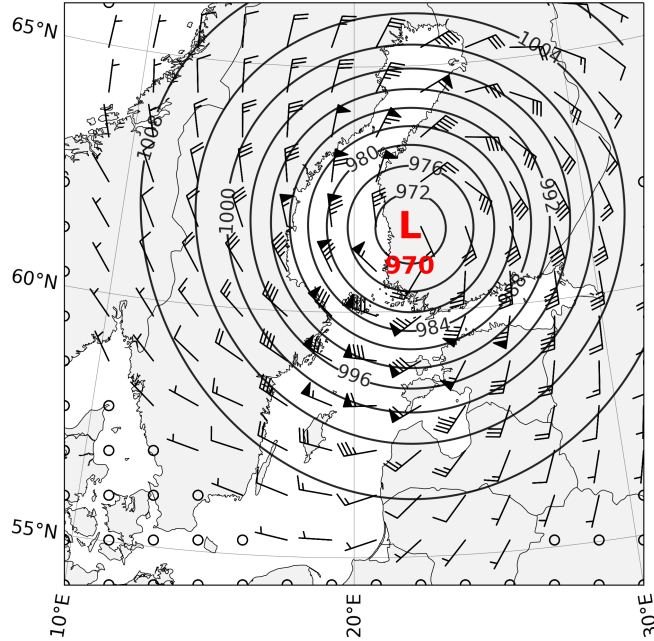
Figure 2. Example of synthetic cyclone and induced winds. Contours represent sea level pressure (hPa) and wind barbs represent 10 m winds (knots).

$$P_{x,y} = P_0 + \Delta P \exp\{-\frac{(x-x_0)^2}{2\sigma^2} - \frac{(y-y_0)^2}{2\sigma^2}\}, \tag{1}$$

where P_0 is the air pressure sufficiently far from the cyclone center, ΔP is the pressure anomaly in the cyclone center, x_0 and y_0 are the time-dependent coordinates of the cyclone center, and σ, the width of the pressure anomaly, is related to the radius R of the cyclone via $\sigma = 0.25R$. 10-metre geostrophic wind components (u_g, v_g) are calculated from the pressure field using the geostrophic balance as

$$u_g = -\frac{1}{f\rho}\frac{\partial P}{\partial y}, v_g = \frac{1}{f\rho}\frac{\partial P}{\partial x}, \tag{2}$$

where u_g (m s^{-1}) is the zonal geostrophic wind component, v_g (m s^{-1}) is the meridional geostrophic wind component, $f = 2\Omega\sin(\phi)$ is the Coriolis parameter (Ω is the rotation rate of the Earth and ϕ is the latitude) and ρ is the density of air (1.3 kg m^{-3}). An example of the pressure and wind field of a synthetic cyclone is shown in Fig. 2.

Actual wind components (u, v) differ from the geostrophic winds (u_g, v_g), especially near the surface where friction acts to slow winds down, leading to winds turning anticlockwise in the Northern Hemisphere, and where curvature effects (the centrifugal force) are important. We study this difference between the theoretical geostrophic and the actual near-surface wind by using the analysis previously applied by Särkkä et al. (2017), with linear regression applied separately for wind

speed and direction in the ERA-Interim reanalysis dataset (Dee et al., 2011) over the period 1979–2012. We obtain spatially varying coefficients for wind speed and direction change by comparing the actual 10-metre winds of the reanalysis dataset with geostrophic winds that we calculate from the mean sea level pressure field. We then modify the geostrophic winds in the cyclone generator with coefficients from Särkkä et al. (2017) to achieve more realistic wind fields. Finally, we save these wind

90 fields and pressure fields in a numerical grid and use them as the input of the sea level model.

2.2   Limitations of the synthetic cyclone

It should be noted that the use of synthetic cyclones is inevitably subject to certain limitations. One of these is the Gaussian shape of the cyclone. As can be seen from Fig. 2, the spatial field of the simulated cyclone is not entirely realistic, as the cyclone is symmetrical and lacks the frontal structures typical of extratropical cyclones. However, the Gaussian shape was

95 originally chosen for two main reasons: (1) it requires only two parameters to describe it, and (2) it describes the general shape of cyclones with sufficient accuracy. It is clear, of course, that not all possible variations in cyclone shape can be described by the Gaussian simplification. However, we would like to emphasise the importance of other factors in determining sea level variations. For example, far more relevant to simulated sea levels than the deviations of the shapes from Gaussian are (1) the depth and extent of the cyclone, and (2) the track and propagation speed of the cyclone.

100 Furthermore, the pressure anomaly of the simulated cyclone is not time-dependent, neglecting the deepening and filling of the cyclone. The value of the pressure anomaly affects the simulated sea level mainly when the cyclone moves over the sea. In the Baltic Sea, cyclone passes the sea area in less than 24 hours, limiting the effect of the time-varying pressure anomaly on the simulated coastal sea level.

2.3   Cyclone parameters

105 The ~~depth and horizontal extent~~ horizontal extent and depth of the synthetic cyclone, related to ΔP and the radius R and ΔP, must be limited by physical considerations. ~~We set the maximum value of the pressure anomaly ΔP to -40 hPa and the minimum~~ According to Rudeva and Gulev (2007), the effective radius of cyclones over the ocean can be greater than 900 km, with an average of about 700 km. Therefore, we set the radius R to 1000 km. ~~With these values, the maximum~~, although we acknowledge that this value may be at the upper limit of what is plausible in the Baltic Sea region (Rudeva and Gulev, 2007).

110 The value of the pressure anomaly ΔP was constrained by the resulting geostrophic wind speeds. Somewhat subjectively, we decided that the surface wind speeds ~~reach~~ could not exceed 40 ~~m s^{-1} over the Baltic Sea. For~~ m/s. With this constraint, we set the pressure anomaly ΔP to -40 hPa. According to the geostrophic wind law, larger pressure anomalies ~~and smaller radii, even~~ would have resulted in higher wind speeds ~~are generated in the simulations, but such high speeds are not realistic~~, but wind speeds higher than 40 m/s are not plausible in the Baltic Sea ~~.~~ region.

115 The numerical sea level model describes how surface winds cause water flow through the wind friction component in the Navier-Stokes equations. The dependence of the wind friction on velocity is quadratic in our model, but in reality this dependence is likely different for extreme wind speeds (over 50 m s^{-1}). This is an additional argument for limiting the surface wind speeds to 40 m s^{-1} for the synthetic cyclones. We vary only four parameters in the simulations: the latitude and longitude
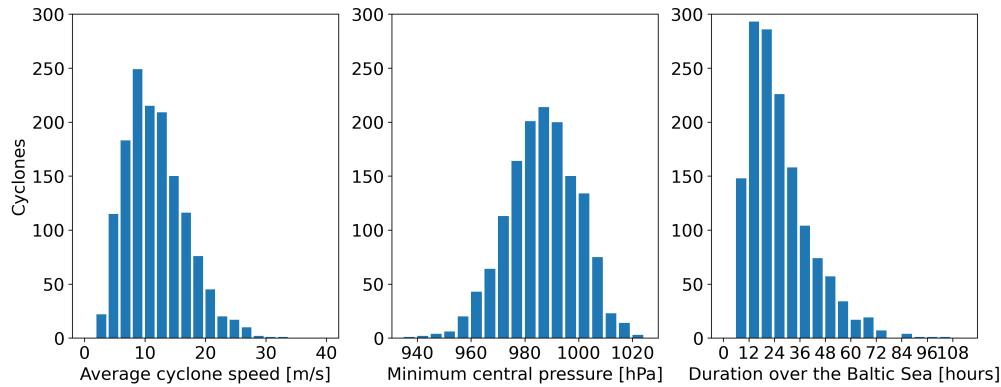
Figure 3. Distribution of mean speeds, minimum pressures and durations of cyclones that pass the Baltic Sea region in October-March 1979-2020 in ERA5 reanalysis.

of cyclone origin, plus the speed and direction of cyclone propagation (cyclone velocity and direction given by zonal speed u and meridional speed v). We estimate the distribution of mean cyclone velocities for storms that pass the Baltic Sea region during October-March 1979-2020 based on ERA5 reanalysis data (Hersbach et al., 2020) using the same cyclone track dataset as was used in Laurila et al. (2021). The smallest velocities from this analysis are around 5 m s^{-1} (Fig. 3a). Even smaller velocities of the cyclone lead to situations where the cyclone is almost stationary, with cyclone winds piling water towards the coast. As this scenario is likely unrealistic, we limit the minimum velocity of the cyclone to 5 m s^{-1} .

Genesis points of cyclones that pass the Baltic Sea region in October-March 1979-2020 in ERA5 reanalysis.

2.4 Numerical sea level simulation

In our numerical simulations we use a barotropic sea level model which describes the internal (intra-basin) fluctuations of the Baltic Sea with a set of two-dimensional hydrodynamic equations (Häkkinen, 1980) based on the Hansen sea level model (Hansen, 1956). We treat the Baltic Sea as a closed basin, with no connection through the Danish Straits to the North Sea. This one-layer This sea level model is based on shallow water equations, derived from the Navier-Stokes equations by integrating the vertical coordinate out, and uses surface winds and surface pressure as atmospheric forcing. As the Baltic Sea has a mean depth of ca. 50 m, the shallow water approximation is valid. Computational domain extends from 53 N to 66 N and from 9 E to 31 E. The spacing of the numerical grid is 0.1 degrees in the meridional direction and 0.2 degrees in the zonal direction. We treat the Baltic Sea as a closed basin, with no connection through the Danish Straits to the North Sea. The model has previously been validated with Finnish tide gauge data and with ERA-Interim forcing 1979-2012 (Särkkä et al., 2017). The correlation was between 0.91 and 0.94 for all 13 Finnish tide gauge sites, even if the spacing of the numerical grid was sparser (0.25 X 0.5 degrees) than in this study. The details of the sea level model were explained in Särkkä et al. (2017).

The variations in the water volume of the Baltic Sea are mainly caused by the flow of the water in and out of the Baltic Sea through the Danish Straits. The contribution of the water volume to the local sea level (also called preconditioning) is termed

140   the water balance component. This component can be estimated with a statistical model, based on wind speeds at a single coordinate point at (55° N, 15° E) (Johansson et al., 2014; Särkkä et al., 2017). The progression of a cyclone over the Baltic Sea region in the high sea level cases takes only one or two days, and the change in the water balance within this time is small (less than 10 cm). Thus we do not include changes in the water balance in our simulations. Instead, we estimate the extreme value of the water balance component to be 100 cm (Särkkä et al., 2017) and add it to the simulated local sea level maximum.

145   This describes the situation when preceding weather conditions have raised the water volume in the Baltic Sea, and then the arriving cyclone causes fluctuations on top of the high mean sea level.

The tides are small in the Baltic Sea, with largest tidal ranges (over 20 cm) at the eastern end of the Gulf of Finland and near the Danish Straits in the southwestern Baltic Sea. For a large part of the Baltic coast, the tidal range is only a few cm (Medvedev et al., 2016), and thus we do not include tides in the model.

150   ## 2.5   Simulation parameters

We The sea level model used in this study enables us to choose any location on the Baltic Sea coast to investigate what are the cyclones that cause high sea levels on that location. As we are interested in the highest sea levels on the entire Baltic Sea coast, we select three sites to represent different bay areas of the Baltic Sea, including coastal cities where the sea level variability is large and where the effects of flooding are significant - the latter due to dense inhabitation with substantial infrastructure on

155   low-lying coastlines. As the predominant direction of the cyclone tracks is from the west(Fig. ??), we choose to study sites at the ends of bays, on the eastern coast of the Baltic Sea. At these locations, the effect of wind-driven water transport is largest, as the water piles towards the coastline. We select the locations so that the Gulf of Bothnia is represented by Oulu, situated at its northern end (Fig. 1). The Gulf of Finland is represented by St. Petersburg, situated at its eastern end. The smaller sub-basin, the Bay of Riga is represented by Riga. In the Gulf of Bothnia, we also simulated sea levels at Kemi, but found that the highest

160   simulated sea levels at Oulu exceeded the ones found at Kemi.

We use the surface wind and pressure fields of synthetic cyclones as atmospheric forcing for the sea level model. We compare simulation results with the observed extremes in tide gauge data from Fig. 1 of Averkiev and Klevannyy (2010). As we calculate only short-term sea levelchangesthe initial condition for the Baltic Sea level in the simulation is a uniform constant sea level, our results represent the sea level fluctuation with respect to the sea level prevailing mean sea level in the

165   Baltic Sea preceding the arrival of the cyclone. In Table 1 we give the parameters that yield the highest sea levels at each of the three sites. The highest sea levels are reached, as expected, for the cyclones with the maximal values of the pressure anomaly (40 hPa) and cyclone radius (1000 km)origins of the cyclones are all on the longitude 10 E. The effect of the winds on the sea level fluctuation with the distance close to the radius from the cyclone center is small, thus is was found that it is not necessary to set the cyclone origin further in the west. A common factor for the cyclone tracks of these maxima was that the propagation

170   speeds were close to the minimum speed limit 5.0 m s^{-1} : 5.7 m s^{-1} for Oulu, 5.4 m s^{-1} for St. Petersburg and 5.0 m s^{-1} for Riga. Due to the low propagation speed, high wind speeds prevail for a longer time over the Baltic Sea, enhancing the water flow towards the ends of bays. For even lower propagation speeds, the extreme sea levels at the studied sites would likely be still higher.

Location	Origin	~~Press. anomaly Radius (km)~~ zonal vel.	meridional vel.	total velocity (m s^{-1})	Max. sea level (cm)
Oulu	(61.0, 10.0)	~~-40 1000~~ 4.3	3.7	5.7	294
St. Petersburg	(57.5, 10.0)	~~-40 1000~~ 4.8	2.5	5.4	346
Riga	(67.0, 10.0)	~~-40 1000~~ 3.3	-3.7	5.0	265

Table 1. Cyclone parameters for the highest simulated sea levels at Oulu, St. Petersburg and Riga.
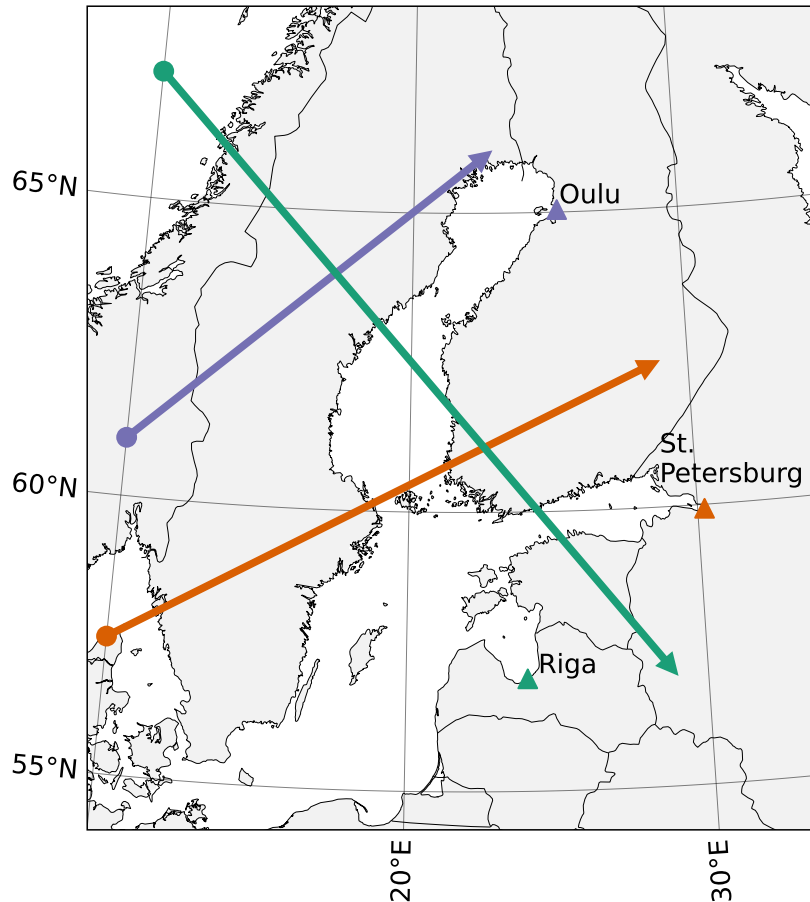
Figure 4. Cyclone tracks of the highest simulated sea levels at Oulu, St.Petersburg and Riga.

3 Results

175 The differences in vertical reference systems between different countries complicate our comparison of simulated extremes against observed extremes. As our aim is only to estimate the extremes to an accuracy of tens of centimeters, we use a fixed value of 100 cm for the water balance component, even where this value differs between the reference systems.
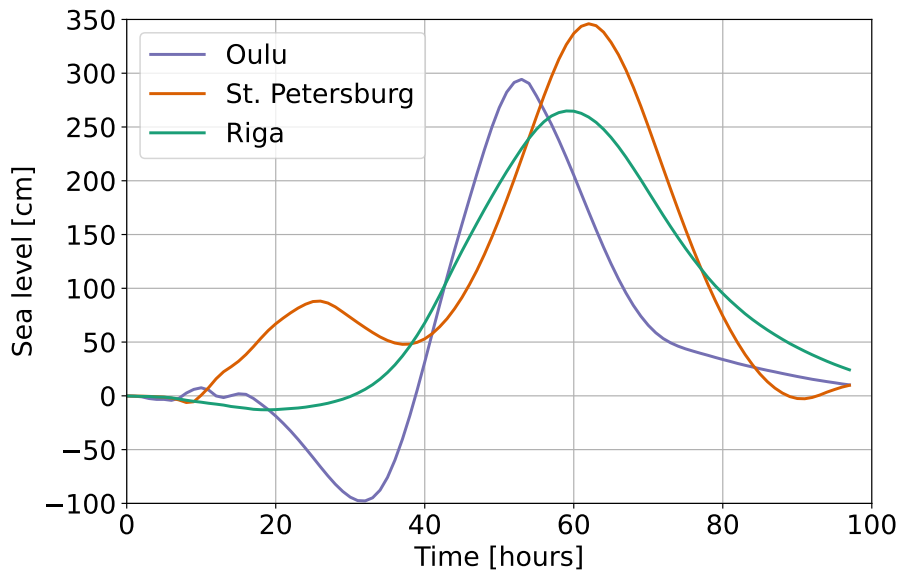
Figure 5. Highest simulated sea levels at Oulu, St.Petersburg and Riga.

The simulated sea level extreme at **Oulu** is caused by a cyclone progressing northeastwards from southern Norway towards Finnish Lapland (Fig. 4). The strong winds on the right hand side of the propagation direction generate water flow from the central Baltic Sea towards the northern end of the Bothnian Bay. At first, the wind direction in front of the cyclone center is towards the northwest in the Bothnian Bay, moving water away from Oulu and causing a decrease in sea level down to -100 cm. Then, as the cyclone center moves northwest of the Bothnian Bay, the wind direction is towards the northeast, piling water towards the northeastern end of the bay, which causes a rapid increase in sea level at Oulu. Finally, when the cyclone is north of the Bothnian Bay, the wind direction is from west to east over the northern Bothnian Bay. This pushes yet more water towards the Oulu coastal region, leading to a very high sea level maximum of 294 cm (Fig. 5). The highest observed sea level at Oulu is 183 cm from January 14 1984. This event was caused by a deep low pressure (950 hPa) having its center in Sweden, inducing southern winds that both increased the water amount in the Baltic Sea and pushed water towards the end of the Bothnian Bay. Taking into account the maximum water balance component (100 cm), a sea level at Oulu of 394 cm could be reached, which is significantly higher than the observed maximum value. In the study of Andrée et al. (2022) with uniform wind fields (wind speed 22 m s^{-1}), the highest simulated sea levels in the Bothnian Bay were at Kalix and Kemi (both situated west of Oulu), where extreme sea levels between 275 and 300 cm were obtained. These levels are of the same magnitude as our result for the extreme sea level at Oulu without the water balance component.

Räty et al. (2022) used Bayesian hierarchical modeling to analyse return level estimates and theoretical upper limits for the sea level on the Finnish coast from tide gauge data. At Oulu, the 1000-year return level is approximately 200 cm and the theoretical upper limit for the sea level using Bayesian modeling is approximately 250 cm. These estimates include the water

balance component, and are considerably lower than the nearly 4 meters estimate in our analysis. This indicates that the high values that we obtain from our synthetic cyclones method can not always be derived from the statistical extreme value analysis.

The cyclone track that causes the highest simulated sea level extreme at **St. Petersburg** (346 cm, Fig. 4) is interesting. The cyclone originates from northern Denmark and moves over southern Finland. Winds cause water to flow from the central Baltic Sea to the Gulf of Finland, which accumulates water at its eastern end. The highest observed sea level occurred in 1824, reaching 421 cm at St. Petersburg. The simulated sea level at St. Petersburg by Averkiev and Klevannyy (2010) was 590 cm. Taking into account the maximum water balance component of 100 cm, the highest estimated sea level in our simulations at St. Petersburg is 446 cm. The higher spatial resolution Averkiev and Klevannyy (2010) use near St. Petersburg is likely the reason for the 1.5 meter difference between their simulated maximum and our result. The narrow shape of the Gulf of Finland and its west-east orientation leave the eastern end of the Gulf of Finland especially vulnerable to sea flooding when weather conditions induce water transport towards St. Petersburg.

The highest simulated sea level maximum at **Riga** is due to a cyclone originating in central Norway, then moving to the southeast over the Bothnian Sea and southwestern Finland (Fig. 4). Northerly winds push water from the Bothnian Sea and central Baltic Sea towards the Bay of Riga. As the cyclone moves over the Gulf of Finland and Estonia to Russia, the north-westerly winds cause water accumulation in the vicinity of Riga. The simulated maximum at Riga is 265 cm. The highest observed sea level is 229 cm from November 1969. By adding the water balance component of 100 cm, the highest estimated sea level is 365 cm, considerably higher than the observed maximum.

In summary, the simulated extremes clearly exceed the observed ones at Oulu and Riga, even accounting for a tens of centimeters ambiguity in the value of the water balance component from the reference level of the tide gauge location. At St. Petersburg, however, the simulated extreme is of similar height to the observed extreme in 1824.

4   Conclusions

The results of our study show that a numerical sea level model combined with synthetic cyclones can be used to study the characteristics of those cyclones that induce extreme sea levels on the Baltic coast. Comparing the three studied sites, we find the highest simulated sea levels at St. Petersburg, - around 3.5 meters, - so that, adding the water balance component, sea levels over 4 meters are possible. These values are comparable with the values observed during the extreme flood of 1824 (421 cm in St. Petersburg). As the flood in 1824 was the highest in the 300-year history of St. Petersburg, the simulated case could correspond to the extreme weather conditions that contributed to the sea levels observed in 1824.

The high sea level at Oulu, approximately 4.0 meters, is due to a cyclone causing water flow from the entire Gulf of Bothnia towards the northern Bothnian Bay, and finally the change in wind direction from westwards to eastwards in the Bothnian Bay. One could expect that an even higher sea level would be found in Kemi, situated at the northern end of the Bothnian Bay, as the highest sea level in Finland (201 cm in 1982) was observed there. In our simulations at Kemi (results not discussed in this publication), the highest maxima are all under 4.0 meters. The shape of the coastline and bathymetry in the Bothnian Bay

apparently favor high sea levels at Oulu more than at Kemi. Our simulation results indicate that there is potential for a flood exceeding four meters at Oulu, far surpassing the observed maximum in 1984 (183 cm).

At Riga, the highest simulated sea level is obtained with a cyclone passing the Baltic Sea with a track from central Sweden towards northwestern Russia. The simulated extreme of 265 cm (with 1 m added to include the water balance component) is considerably higher than the observed maximum (229 cm) in 1969. The cyclone related to the maximum in 1969 had a similar track compared to the one used in our simulation. This indicates that a similar storm track can occur and that a cyclone with such a track and slow velocity could cause high sea levels in the Riga region.

Our method has certain limitations. The spatial description of the simulated cyclone (Fig. 2) is not completely realistic, as the cyclone is symmetrical and lacks frontal structures which are typical for extratropical cyclones. The pressure anomaly of the simulated cyclone is not time-dependent, neglecting the deepening and filling of the cyclone. Comparing to the description The effect of waves on the extreme sea levels on the coast was not considered in this study, as this effect strongly depends on the bathymetry and on the location. Our aim was not to evaluate rigorously the highest sea level taking into account all factors, but to evaluate what is the contribution of the synthetic cyclone in (Averkiev and Klevannyy, 2010), we find that the deviation from the symmetric structure is not large in their model. The value of the pressure anomaly affects the simulated sea level mainly when the cyclone moves over the sea . Therefore including the deepening and filling of the pressure anomaly would have only a small effect on the simulated sea levels on the coast, as the simulated cyclones move in a short time over the the sea level maximum, which already exceeds observed sea level extremes in many locations on the Baltic Sea coast. For detailed studies of extreme sea levels for a single site, simultaneous sea level and wave simulations are needed, and such studies have already been made in the Baltic Sea (Apukhtin et al., 2017; Gordeeva and Klevannyy, 2020).

Our results indicate that synthetic cyclones are a useful tool for studying the sea level response to cyclones with different characteristics. In particular, our method offers the opportunity to study those extreme sea levels which can be expected on the Baltic Sea coast if all the components that raise the water level are present, namely a high water balance paired with a slow and deep cyclone moving along an optimal track. Our simulations of extreme sea levels reveal those areas of the Baltic Sea which, due to shoreline geometry and coastal bathymetry, are especially prone to extreme coastal flooding.

Author contributions. All authors contributed to the design of the study. JR and MK implemented the method for the generation of synthetic cyclones. JS implemented the numerical sea level model. MR studied the properties of cyclones from reanalysis data. JS performed the sea level simulations with synthetic cyclones and conducted the analysis presented in the paper, and wrote the manuscript with critical comments from the other authors.

Competing interests. The authors declare that they have no conflict of interest.

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