# Landslide susceptibility assessment in the rocky coast subsystem of

Essaouira – Morocco

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#### Abstract

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15 During the last fewIn recent decades, many-multiple researchers have produced landslide susceptibility maps using different techniques and models, including the information value method, which is a statistical model that is widely applied to various coastal environments. This study aims aimed to evaluate the susceptibility tofor the occurrence of landslides in the Essaouira coastal area using the bivariate statistical method methods. In this study Here, 588 landslides of distinct landslide types were identified, inventoried, and mapped. They mostly primarily result from 20 the observation and interpretation of different data sources, namely, high-resolution satellite images, aerial photographs, topographic maps, and extensive field surveys. The rocky coastal system of Essaouira is located in the middle part of the Morocco Atlantic coast of Morocco. The study area was split into 1534 cliff terrain units of 50 m in width. For training and validation purposes, the landslide inventory was divided into two independent groups: 70% for training and 30% for validationvalidating. Twenty-two layers of landslide-conditioning factors were prepared, 25 includingnamely: elevation, slope angle, slope aspect, plan curvature, profile curvature, cliff height, topographic wetness index, topographic position index, slope over area ratio, solar radiation, presence of faulting, lithological units, toe lithology, presence and type of cliff toe protection, layer tilt, rainfall, streams, land-use patterns, the normalized difference vegetation index (NDVI), lithological material grain size, and presence of springs. The statistical relationship between the conditioning factors and the different landslide types of landslides waswere 30 calculated using the bivariate information value method, in a pixel and in the elementary terrain unit-s-based base model. Coastal Validation of the coastal landside susceptibility maps were validated using was done using the landslide training group partitions. The receiver operating characteristic curve (ROC curve) and area under the curve were used to assess the accuracy and prediction capacity of the different coastal landslide susceptibility models. Two methodologies, considering a pixel-based approach andor using coastal terrain units, were adopted to 35 evaluate the coastal landslide susceptibility. The results allowed maps allowed classifyingfor the classification of 38

% of the rocky coast subsystem with high susceptibility to landslides, which were mostly located in the southern part

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Commenté [A2]: It is only necessary to abbreviate a term in the Abstract when it is used more than once.

Commenté [A3]: The unnecessary repetition of the same article in a series tends to cause wordiness. Identify the repeated words and eliminate them.

For example,

Original: We verified the samples using the source, the original, and the final images. Revised: We verified the samples using the source, original, and final images.

of the Essaouira coastal area. -These susceptibility maps would-will be useful for general-future planned development activities in the future as well as for environmental protection.

**Keywords**: Coastal landslide susceptibility mapping, coastal landslide inventory, conditioning factors, Information Value, Essaouira coastal area, Morocco

#### 1. Introduction

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Landslides are common processes in the rocky coastal system of Essaouira province. They essentially result from the interaction of sub-aerial, marine, and anthropogenic processes (Trenhaile 1987; Sunamura 1992; Hampton & and Griggs 2004; Greenwood & and Orford, 2007), making this system exposed more susceptible than any other natural system to Anthropicanthropogenic activities pressure and erosional processes more than any other natural system. Consequently, In consequence, the fast dynamic evolution imposes restrictions on the way the human humans occupy coastal areas (Teixeira 2006; Marques 2009; Teixeira 2015; Moore and Davis 2015; Gilham *et al.*, -2018).

The processes process of building creating landslide susceptibility maps generally involve-involves several qualitative
or quantitative approaches, started\_starting\_withby landslides\_landslide\_inventory as the first step for assessing landslide\_susceptibility, hazard, and risk (e.g., Aleotti & and Chowdury 1999; Dai & and Lee 2002; Van Westen et al., 2008; Corominas et al., 2014; Oliveira et al., 2017; and Meena et al., 2018). For rocky coastal areas, landslide susceptibility/hazard assessment\_assessments\_mainly addresses\_address the evaluation of the cliff retreat (Oliveira et al., 2008; Rocha et al., 2007; Oliveira et al., 2017), landslide inventorying, and or susceptibility mapping
(e.g., Marques et al., 2011). The identification of factors controlling the rocky coast system is a critical step to for

- better <u>understand understanding</u> how this system is evolving and <u>to</u> predict its future evolution (Neves <u>&and</u> Ramos Pereira 1999). Landslides are responsible for significant erosion in rocky <u>coast systemcoastal systems</u> (Andriani <u>&and</u> Walsh 2007; Violante 2009<u>; and</u> Sunamura 2015). Therefore, by knowing the set of predisposing factors that <u>conditioned-condition</u> the landslide occurrence, it is possible to spatially predict where future landslides will occur
- 60 (Varnes, 1984). <u>ManyThere are many</u> different landslide-conditioning factors <u>play</u>, with an important role in the preparation of the landslide susceptibility maps (e.g., Zêzere 2002). These factors, <u>although-which</u> are dependent of the analysis scale and <u>landslide type</u> of <u>landslides</u> generally include: elevation, slope, aspect, plan– and profile curvature, topographic wetness factor index (TWI), topographic position index (TPI), slope over area ratio (SOAR), solar radiation, faulting, lithology, lithological layers tilt, precipitation, streams, land-use patterns, <u>normalized</u>
- 65 difference vegetation index (NDVI) or vegetation density factor, grain size, and spring presence. (e.g., Van Westen  $\underline{et}$   $al_{\underline{n}}$ , 2008; Reichenbach  $\underline{et}$   $al_{\underline{n}}$  2018; Pereira  $\underline{et}$   $al_{\underline{n}}$  2020). When related to with sea cliffs susceptibility assessmentassessments of sea cliffs, landslide-conditioning factors also include the cliff edge height and, coastal slope toe protection (e.g., Marques  $\underline{et}$   $al_{\underline{n}}$ , 2011, 2013; Marques 2018; Guilham  $\underline{et}$   $al_{\underline{n}}$  2018; Letortu  $\underline{et}$   $al_{\underline{n}}$  2019; Queiroz and Marques 2019). In this studywork, we followed the classification of by Cruden and Varnes (1996), Varnes
- 70 (1978), WP/WLI (1993), and Dikau *et al.* (1996), to differentiate the types of landslides that may occur in coastal cliffs: falls, slides, topples, lateral spreads, and flows. Identifying landslide types remains challenging, even when

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**Commenté** [A4]: Collocations are combinations of words often used together. When certain expressions do not sound "natural" or "right", consult a dictionary for usage.

#### For example,

Original: We *arrived on* the same conclusion. Revised: We *arrived at* the same conclusion.

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supported by intensive fieldwork, which that often faces the lack of clear evidence associated with the degradation of landslide features degradation or inaccessibility to cliff face faces (Neves <u>et al.</u> 2012). Datasets To overcome these field limitations, datasets of aerial photographs ca-can be used to overcome these limitations (Oliveira <u>et al.</u> 2017).

- 75 Many Multiple bivariate and multivariate statistical models are used to analyzeanalyse the landslide susceptibility, and most of of these models require a subdivision of territory in into terrain units and the selection of the appropriate type of terrain mapping units (e.g. grid cells, slope units, geo-hydrological units, unique condition units, and administrative units [{Van Den Eeckhaut *et al.*, 2009-; Marques *et al.*, 2011, 2013; Epifânio 2014-; Corominas *et al.*, 2014-; Zêzere *et al.*, 2017])).
- B0 Data-driven approaches are the most <u>commonly</u> used for landslide susceptibility and hazard zonation (Kanungo *et al.*<sub>2</sub>-2006; Girma *et al.*<sub>2</sub>-2015; Hamza and Raghuvanshi 2017; Mengistu *et al.*<sub>2</sub>-2019; Shano <u>et al.</u><sub>2</sub>-2020) whereas other approaches, such as the bivariate, multivariate, and active learning statistical <u>methods</u>, are also suitable for assessing to assess susceptibility (Corominas <u>et al.</u>, 2014). The bivariate Bivariate statistical methods use on an inductive logic, which assumes that the combination of conditions pertaining to various conditioning factors, <u>analyzedanalyzed</u> separately, may lead to landslide prediction in a given area. The evaluation of <u>the</u>-conditioning factors and their relationship with the past landslides in the study area form forms the basis for the prediction of places where <u>future</u> landslides may occur in the future (Varnes *et al.*, 1984; Van Westen *et al.*, 1997; Dai *et al.*, 2002-; Lan *et al.*, 2004-; Girma *et al.*, 2015-; Chimidi *et al.*, 2017-; Shano *et al.*, 2020).

The information value (IV) method (Yin and Yan 1988), is considered appropriate foras an appropriate method to
 evaluate evaluating landslide susceptibility (Corominas et al., 2014), it lt has been widely used worldwide within different geomorphological backgrounds (Yin and Yan, 1988; Jade and Sarkar, 1993; Lin and Tung, 2003; Yalcin, 2008; Balasubramani and Kumaraswamy, 2013; Zêzere et al., 2017; Mengistu et al., 2019). The IV model is based on the weighted presence or absence of drivers of slope instability. -Thus, the landslide density for conditioning factor classes can be determined by overlaying both maps of both conditioning factors and inventoried landslides (Mengistu

- 95 *et al.*, 2019; Shano *et al.*, 2020)., If the resulted resulting information value V is positive, the causative factor class represents <u>a</u> strong interdependence with the landslides in the area (Yin and Yan 1988; Shano *et al.*, -2020), and the weighted value of a conditioning factor class can be represented as the natural logarithm of <u>the landslide</u> density-of landslide in a factor class, divided by <u>the landslide</u> density in the total map area (Van Westen *et al.*, 1997; Shano *et al.*, 2020).
- The validation Validation of the landslide susceptibility map is an essential step for the evaluating evaluation of the predictive capacity of the model. It can be <u>consideredseen</u> as a test of the <u>model's</u> ability of the model to reflect the real environment trough and evaluation of evaluate its accuracy and predictive capacity (Beguería 2006; Frattini *et al.*, 2010; Shano *et al.*, 2020; Mateus *et al.*, 2021). The Receiver receiver operator characteristic (ROC) is a recognized recognised technique used in statistical approaches approach validations to check the performance of the prediction ability of the probability of with correctly identified landslides, against the probability of incorrectly identified landslides (Gorsevski *et al.*, 2006a; Shano *et al.*, 2020).

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**Commenté [A5]:** Concise writing means using the fewest words necessary. One way to achieve conciseness is to use a direct verb rather than its noun form.

For example, Original: We *performed an evaluation* of all methods used previously. Revised: We *evaluated* all methods used previously.

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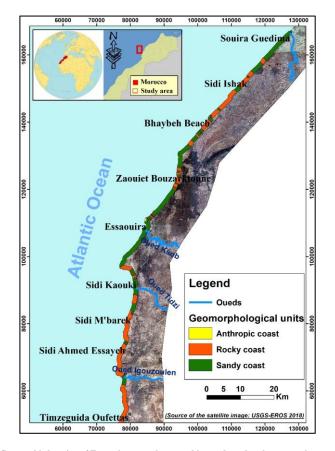
However, the The dynamics of the Essaouira coastal area have been is poorly studied. A beach granulometric technical study was <u>conductedearried out in the</u> Essaouira bay Bay in 1955 by the hydraulic laboratory of Neyrpic (El Mimouni A. and Daoudi L. 2012). Other existing studies have focused on the general morphology of the sandy dunes in the upper part of the beach and on the mainland (Gentile, 1997; Simon, 2000; and Lharti *et al.*, 2006). Alternatively, This this workstudy, on the other hand, hims toat: i) to define the type and emplacement of each landslide by an inventory validated by using a field survey; ii) to identify the most important predisposing variables that control the spatial distribution of different landslide types; iii) to assess landslide susceptibility in Essaouira coastal cliffs for different landslide types of landslides and to classify susceptible areas to the occurrence of landslides; and, finally, v) to validate the susceptibility map.

#### 2. Study Area

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The Essaouira coastal area is located along the middle section of the Atlantic coast of Morocco (Fig. 1), and which extends over 134\_-km-long. It has high coastal systems\_system\_diversity\_ including estuaries, bays, beaches, sandy spits, cliffs\_ and rock-rocky shore platforms (Weisrock 1980; Simon 2000; Lharti *et al.*, 2006), where we adopted a <u>A</u> classification was applied based on tree three\_subsystems; sandy coast, rocky coast, and anthropic coast. The study site (Fig. 1) is characterized characterised by stretches of sandy coast (48%), rocky coast (51%), and anthropic coast (1%, the Port of Essaouira port), delimited toon the north by the Tensift estuary, toin the south by Timzguida Ouftas village, toin the east by Essaouira province municipalities\_ and toin the west by the Atlantic Ocean and the island of Mogador in front of Essaouira City (Fig. 1). This coastal area has a predominantly-a semi-natural landscape which is locally interrupted by heavily anthropized coastal areas, especially particularly inat the city of Essaouira City (Fig. 1). **Commenté [A6]:** The phrase "on the other hand" is a part of a correlative conjunction pair (on the one hand...on the other hand). It is not used without its preceding pair. Moreover, it is meant to indicate a strong contrast which is not the case here. Hence, I have replaced with a more suitable connective word.



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# Figure-Fig. 1: Geographic location of Essaouira coastal area and its sandy and rocky coast subsystems (Coordinate Coordinate systems: Lambert Zone I projection)

Geologically, the study area is located in the Atlantic Atlas, which is – considered to beas the westernmost part of the High Atlas mountains Mountains (Weisrock 1980),- with its northern and largest plateau (Haha and Chiadma) part dropping gently from SE to NW, in accordance with the overall structural framework. However, the The landscape is however varied, crossed by cuestas and vigorous crests, turned towards the SE, and associated with the frequent alternations of sandstone, dolomitic, limestone, or-marl, clay, and gypsum layers. The landscape is interrupted by sudden isolated anticlinal folds, such as the Jbel Hadid (725 m high), quite to the north, or the Jbel Ouamsitten (900 m high) to the south. Towards to the west, the abundance of consolidated dunes and sandstones with oblique stratification and conglomeratic levels is relevantgain relevance the (Weisrock 1980). To the south, a coastal basin with original sedimentary material known as the "Haha Basin" (Dufaud *et al.*, 1966) is associated with the opening of the North Atlantic, which is generally consistent with the end of the Triassic (Choubert *et al.*, 1971; Hallam, 1971;

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adaptations (Weisrock 1980).

Le Pichón; 1971; Weisrock 1980). It mainly consists mainly of sandstones, pelites, conglomerates, and red salt clays, with essentially continental facies. Deep marine sedimentation was successful from From the Lower Liassic to the Upper Cretaceous succeed deep marine sedimentations. During these long periods, the sedimentation of the coastal basin has constantly oscillated between an epicontinental regime, with terrigenous deltaic or alluvial contributions and marine organogenic or evaporitic deposits, and a more openly open marine regime with neritic limestones and marls. Towards the north, the coastal platform is largely developed, also called the "Moghrebian platform", from the name attributed to the sandy and sandstone deposits that cover it (Choubert and Ambrogi; 1953;; Weisrock 1980), and thus tapers off at the southern mountainous part.

- From a structural point of view, the study area is <u>characterized characterised</u> by a double structural division marked
  by a close adaptation <u>towith</u> the hydrographic network (Weisrock 1980). The first one is linked to the opening of the Atlantic, <u>includes including</u> the extensional faults fundamentally oriented NNE-SSW, <u>this</u> for the <u>entirewhole</u> Atlantic Atlas and its northern edge. This second direction may have the same origin as the first; these Hercynian breaks in the basement influenced sedimentation and then reappeared, affecting the cover<sub>7</sub> during the Atlasic phases (Saadi<sub>7</sub> 1972). This second direction, WNW-ESE, related to the opening of the Atlantic, is <u>increasinglymore and more</u> evident (Oued
  Ksob, northern fallout anticlines, Oued Tensift). This direction is attenuated towards the S, while in the central region (Tamanar <u>plateauPlateau</u>), a W-E direction appears as a result of the ancient Hercynian direction (Saadi<sub>7</sub> 1972). In addition to these two fundamental systems, <u>the Essaouira region</u>, and because of the thickness of the saliferous clay layer, <u>the Essaouira region</u> is marked by the development of <del>a</del>-diapiric style tectonics, well represented <u>especially particularly</u> in the SE of Essaouira (Weisrock 1980).
- 160 From a geomorphological point of view, landform distribution in the study area is asymmetrical: all the plateaus dominate to the N and NW sectors, and are almost absent to the S and SE, occupied by the mountain. In; in accordance with the general layout of the High Atlas, the altitudes rise-increase towards the south and east. Thus, The the morphogenesis of the Atlantic Atlas thus-depends on general physical geography, in addition to the structural morphology of the folded chains, the phenomena of encrustation, coastal eolian constructions, and glaciation (Weisrock 1980). The Atlantic Atlas is open to oceanic influenceinfluences. This area is particularly characterized characterized by its dual character as a mountainous and coastal region, which makes it possible to link continental and marine morphology<sub>25</sub> the latter offering offers the advantage of being able to establish a solid chronological base from the Pliocene onwards by comprising a whole series of stepped fossil beaches. The coastal area haspart takes on a uniform appearance from north to south. On average, the Mesozoic bedrock disappears under a sandy cover shaped into innumerable encrusted hills all-along the ocean (Weisrock 1980).

From a hydrological point of view, we note the presence of two large watersheds, Oued Tensift and Oued Ksob, were noted, to which are added coastal Oueds are added: Oued Tidzi and Oued Igouzoullen. These hydrographic networks present are an important source of sediment supply and, they are characterized characterised by a flow which is roughly carried out very roughly from E to W, rather faithfully adapted to the topographic framework; however, the courses of the valleys, more often monoclinal or orthoclinal than cataclinal, reveal a long evolution and successive re-

**Commenté [A7]:** Compound adjectives jointly modify the noun they precede. For clarity, hyphenate the compound adjectives.

For example.

Original: There was no correlation with *butyric acid producing bacteria*. Revised: There was no correlation with *butyric acidproducing bacteria*.

**Commenté [A8]:** Try to avoid referencing previously mentioned information with "former" and its counterpart with "latter." Using such vague references forces your audience to read backward, which should be avoided as much as possible in academic writing.

The Essaouira cliffed coastal sector is eharacterized characterised by the presence of many-multiple landslide types of landslides, which are being the dominant hazards responsible for the constraint constraints of human activities and a safe land use (e.g., Moore and Griggs, 2002). The seismic context shows that the coast between Safi and Essaouira have has landslide activity that is probably likely related to the seismic events (Elmrabet *et al.*, 1989). The most significant of which, capable of causing disproportionate effects on a highly unstable cliff, occurred in 1757, on 7th March 1930 (32° N, 11.5° W, M = 5.1, felt in Casablanca, Safi and Essaouira, intensity IV)<sub>2</sub> and on 2nd August 1963 (34.7° N, 8.9° W, M = 4.1, felt in Casablanca and Mohammedia, intensity IV). In the 1757 event, the landslide could also have been conditioned by an aftershock of the earthquake on 1st November 1755 affecting the eliff natural instability of the cliff, which had been enhanced by the effects of the tidal wave (Elmrabet *et al.*, -1989).

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Climatically, the Atlantic Atlas is located <u>atin</u> a relatively low latitude (<u>approximatelyaround</u> 31° parallel), which places it under the predominant influence of subtropical anticyclonic cells, at the limit of the great displacements of polar air masses. It is a position sensitive to the slightest deviations of these <u>centreseenters</u> of action; and thus, it is particularly interesting to reconstitute the possible conditions of <u>the</u>-past climatic oscillations, identified by their morphogenetic marks (Weisrock 1980).

The Essaouira province is characterized characterised by a steppe climate of type BSh according to the Köppen-Geiger classification, whit with low rainfall, the an average annual temperature in Essaouira city is of 18.7 °C, and the average annual rainfall reaches of 295 mm (Climate-data.org). The dominant climate in the Essaouira region is semi-arid, with a diversity diverse of both temperature and precipitation values. This is because due of to the oceanic (Atlantic) setting 195 on one side and to the height of the mountains on the other. The Essaouira region is an area where hot summer winds and humid winter winds change. The "Chergui" (the hot wind from the Sahara), and the northeast wind that blows almost all year round. It is characterized characterised by a mild climate throughout the year all year round. The average temperatures are 16.4 °C in January and 22.5 °C in August. The As for the annual rainfall, it is around approximately 280 mm. Two main seasons can be distinguished: i) thea wet season that includes winter and autumn, with a monthly 200 maximum fluctuating between December and November. Precipitation peaks are clearly marked in autumn and winter, before gradually decreasing from February to May,; and ii) a-the dry season from April to September. This season is marked characterised by scarce rainfall. July and August are the driest months throughout the year, with almost no rainfall. About Regarding the spatial distribution, both the precipitation and the humidity are higher in the coastal zone, for this last it is and they are always higher than 275-%. Summer fog is particularly important at in Essaouira, 205 and other sites that are exposed to maritime influences (Hander, 1988).

Using the rainfall data from stations of Adamna, Chichaoua, Talmest, Abadla, and Igrounzar, which were provided to us by the Tensift Water Basin Agency, we analyzed the average monthly variability of rainfall was analysed for the period 1965–2015, and main results shows show the existence of a rainy season between October and April with a maximum in March for the two stations-Abadla and Chichaoua stations and a maximum in December and November for the stations-Talmest, Igrounzar, and Adamna stations. The dry season extends between June and September, with where the lowest rainfall is-recorded in July and August. The monthly variation in rainfall shows-showed an average

of 15.3 mm for Chichaoua and 14.4 mm for Abadla. The rainfall Rainfall wasis similar over the same period for

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**Commenté [A9]:** For ranges, an "en dash" (–) should be used instead of a hyphen.

Chichaoua and Abadla. The values observed <u>fromin the months of</u> October to April <u>exceed exceeded</u> the average rainfall for each of these two stations, with a maximum in March (27 mm) and a minimum in July (0.5 mm) and August (1 mm). Thus, the evolution of monthly precipitation <u>is-was</u> the same <u>atfor</u> these two stations. It argues in <u>favorfavour</u> of a simple hydrological regime <u>characterized characterised</u> by a regular annual alternation of high and low water. The monthly rainfall <u>atof</u> the three stations Adamna, Talmest, and Igrounzar <u>show showed</u> that the maximum rainfall <u>is-was</u> recorded in the months of November and December, while the average rainfall <u>is-was</u> approximatelyabout 20 mm for Igrounzar and Talmest and 26 mm for Adamna.

Regarding the annual variations inof rainfall at the five stations, the first three concern the stations of Adamna, Igrounzar<sub>a</sub> and Talmest stations have mean annual rainfall of 322 mm, 229.1 mm<sub>a</sub> and 255.4 mm, respectively. At the Adamna station of Adamna, there are were several rainy years with values that greatly exceed exceeded the interannual average, namely 1987/88, 1988/89, 1994/95 to 1996/97, 2008/09 to 2010/11, with a maximum rainfall of approximatelyabout 718 mm in 1995/96 and a minimum of approximatelyabout 136 mm in 2006/06. For the station of -Igrounzar\_station, the highest rainfall was observed during the years-1987/88, 1988/89, 1994/95 to 1996/97, 2008/09<sub>a</sub> and 2009/10. However, the least rainy years arewere: 1968/69, 1976/77, 1991/92<sub>a</sub> and 2014/15. For the Talmest station, the wettest year is-was 1995/96 which recorded with 559.5 mm and the least rainydriest year is-was 2014/15. For this station, we do did not have data for the years 1971/72 until 1975-76.

From a hydrogeological point of view, the Essaouira basin Basin and its coastal zone constitute a set of independent
 but very similar hydrogeological systems that correspond to synclinal basins. Within these systems, groundwater Groundwater exists only in very localized localised areas within these systems. The water Water generally circulates at depth in different limestone or sandstone levels by karstic pathways, and the compartmentalization of the effects of tectonics and diapirism have has caused the compartmentalisation compartmentalization of the basin into several aquifer systems.

For example, the The piezometry of the Plio-Quaternary aquifer, for example, shows an overall flow direction from E-SE to W-NW, conditioned by the straightening of its bedrock to the east following the uplift of the Tidzi diapir (Mennani; 2001). There arelt presents significant fluctuations between periods of high and low water (Fekri; 1993; Mennani; 2001; Bahir *et al.*, 2002; Bahir *et al.*, 2017), which These are related to precipitation which thus controls the regime of the phreatic aquifer. Several problems related to water scarcity and long recurrent periods of drought, have been observednoticed in the Essaouira region during the lastin recent decades (Bahir *et al.*, 2002; Chkir *et al.*, 2008; Chamchati and Bahir; 2013; Bahir *et al.*, 2017). For this reason, the piezometric level in the study area tends to a generalized decline with the inability of some other wells to recover their initial water level, aggravated by under the combined effect of the year 1995, the driest year that Morocco has experienced during the 20th century (Bahir *et al.*, 2002), and overexploitation (Chkir *et al.*, 2008; Bahir *et al.*, 2017).

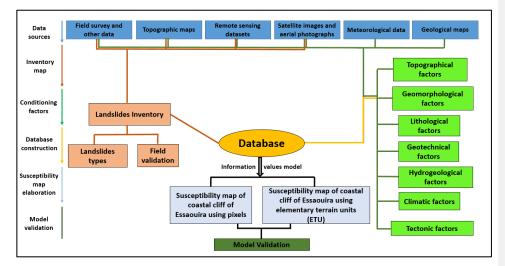
3. Methodology

**Commenté [A10]:** The word "about" is commonly used when rough estimations are made. I suggest using the word "approximately" with numerical values, as it indicates that it is very close to the actual value. Moreover, "approximately" is a more formal alternative. The use of either word, however, is acceptable.

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The current research <u>uses-used</u> different data sources for landslide susceptibility analysis, and their preparation was supported by field <u>survey-surveys</u> and validation. The methodological steps considered for training and validation <u>of</u> the coastal landslide susceptibility models are <u>showed shown</u> in the Fig. 2 and follow this sequence: i) elaborate the landslide inventory, classifying the landslides by type and depth of the rupture surface (shallow and deep);- ii) prepare a set of – 22 conditioning factors grouped <u>in-into 7seven</u> categories (topographical, geomorphological, lithological, geotechnical, hydrological, climatic, and tectonic);- iii) model coastal <u>landslides landslide</u> susceptibility <u>usingwith</u> the <u>information valueIV</u> method for the Essaouira coastal area, using pixels and elementary terrains-terrain units (ETU<u>s</u>);- and iv) independently validate the predictive susceptibility models using ROC curves and <u>area under</u> the curve (AUC).



#### Fig.Figure 2: Schematic diagram of the used methodology used

<u>Coastal</u>A classification of coastal systems were classified into sandy and rocky subsystems was done-according to a morphometric and operational eriterion<u>criteria</u>, and the ETU waswere defined based on the methodology proposed by (Marques *et al.* -(2011); the upper and lower limits of the terrain units, are were defined by the bottom and the top of the cliff, respectively, while the lateral limits were geometrically drawn perpendicular to the contour lines of the topography, and defined by the segmentation of the ridge line into 50 m wide sections. In total, we there werehad 1534 terrain units of on the rocky coast. Each terrain unit was classified as stable or unstable based on the quantification of the percentage of the unstable area of each slope unit.

#### 265 3.1. Landslide inventory

The most essential part of <u>the</u> landslide susceptibility assessment framework is the landslide inventory, <u>which</u> <u>including includes</u> the identification of their location, size<sub>2</sub>-and type<sub>2</sub> and depth<sub>7</sub> to understand the relationship between landslide occurrence and the dataset of predisposing factors (Ercanoglu and Gokceoglu 2004; van Westen *et al.*<sub>2</sub>

 $2006_{27}$  Petley  $2008_{27}$  Epifânio *et al.*<sub>2</sub>- 2013). <u>The Landslide landslide</u> inventory is-was of the historical type, with no past date of occurrence limits, <u>andi</u>t was based in-on the interpretation of different data sources <del>covered\_covering-all</del> the <u>entire</u> study area (<u>Tab.Table</u>-1), such as historical records, 10 m resolution Sentinel satellite imagery, <u>Highhigh</u>-resolution <u>Orthoortho</u>-imagery analysis, and <del>an</del>-intensive field investigation.

# Table 1: Data sources Table <u>table</u>

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Data type	Data denomination	Source	Scale / resolution / Duration	
	Sidi Ishaq 2008			
	Berrakat Erradi 2008			
	Sebt Akermoud 2008			
	Bir Kaouat 2008			
	Moulay Bouzarqtoune 2008			
	Jbel lahdid 2008			
Topographic	Essaouira 2008	National Agency of Land Conservation, Cadastre and	1/25000	
maps	Chicht 2008	Cartography (ANCFCC)	1/25000	
	Ras Sim 2008			
	Essaouira El Jadida 2008			
	Sidi Kaouki 2008			
	Tidzi 2008			
	Sidi Ahmed Essayeh 2009			
	Tafdna 2009			
	Tamanar map	Ministry of Energy and Mines,	1/100000	
Geological maps	Taghazout map	Water and Sustainable	1/100000	
	Marrakech map	Development	1/500000	
Aerial photographs	Mission TAMANAR 07/2016	National Agency of Land Conservation, Cadastre and Cartography (ANCFCC)	1/7500	
	Adamna station		1977_2015	
Meteorogical	Igrounzar station		1968_2015	
Meteorological	Talmest station	Hydraulic basin agency of Tensift (ABHT)	19842015	
data	Chichaoua station		1965_2014	
	Abadla station		1969_2014	
	Sentinel	https://scihub.copernicus.eu/ (Copernicus 2021)	10 m	
Satellite images	High resolution Ortho-imagery	https://earthexplorer.usgs.gov/ (USGS-EROS 2018)	0.3 m	
	Digital elevation model	https://search.asf.alaska.edu/ (JAXA/METI 2020)	12.5 m	

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The identification of the landslides was based on the interpretation of their specific morphological features <u>that are</u> noticeable in high-resolution imagery, including <u>the</u> crown, main scarp, flanks, body, and toe (Pawluszek 2019). Other features <u>includewere detected by</u> the presence of flow materials along gullies, streams with different erosional features, flow tracks, scars along <u>the</u> cliff face, and block deposits on the cliff base (Epifânio *et al.*, 2013, Elkadiri *et al.*, 2014). In addition to these, extensive field observations were used to validate the inventory and add new landslides that were not observed in satellite <u>image images</u> or identified in other data sources.

#### 3.2. Conditioning factors

Conditioning factors describe terrain conditions that are directly or indirectly associated with landslide occurrence, andthey are essential for landslide susceptibility mapping based on data-driven methodologies. Different types of variables (conditioning factors) were compiled and/or generated in a geographical information system (GIS) for susceptibility analysis. According to Marques *et al.* -(2011, 2013), all conditioning factors influencing the stability of coastal cliffs and eoastal slopes should be considered because they may contribute to predict predicting the spatial occurrence of future instability. It is important to mention thatNotably, the selection of conditioning factors associated with these processes eeems appears to be a difficult task becauses these these factors usually typically work in combination; in a multivariate system (e.g. Epifânio *et al.*: 2013; Reichenbach *et al.*: 2018).

Solar radiation was used as a proxy variable for slope aspect, because it enables the quantification of the weight of trivial qualitative quadrant-quadrants (Epifânio, et al., -2013). Slope angle is the most important predisposing factor for the occurrence of landslides (Mancini et al., 2010); but-however, in our study area, the slope angle does not have the same importance for all type of landslides landslide types, and plan and profile curvatures can be associated withto the acceleration and deceleration of the flow, as well as the convergence or divergence of the flow, and can influence the local drainage systems and the kinematics of landslides (Mancini et al., 2010).

The land use map and the Normalized Difference Vegetation Index (NDVI) were extracted from Sentinel images 2021
 (10 m resolution, Tab:Table 1). The lithology, toe lithology and faulting data were obtained from the compilation of a bibliographical review and from three digitised digitalized geological maps: Tamanar and Taghazout 1/100000-scale in the southern section, and Marrakech 1/500000-scale for the northern section, completed with the field survey.

The meteorological Meteorological data and the historical rainfall records are were used for extracting the rainfall factor, using the arithmetic mean method (Smaij, 2011), which consists of calculating the annual arithmetic mean of the values obtained at the weather stations, and projecting them using the Inverse inverse Distance distance Weighting

weighting(<del>IDW</del>) interpolation. While; field surveyField surveys, topographic maps (1:10,000), and <u>digital elevation</u> <u>models (DEM)</u> were used <u>tofor</u> identify and map the stream networks. <u>The Presence presence</u> and type of cliff toe protection, lithology tilting, and <u>the presence of springs</u> were extracted from the observation of satellite <u>images image</u> <u>observations</u> and <del>by</del> field <u>surveysurveys</u>.

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Field work revealed that most landslides occur above the-weak or friable layers, therefore, making geotechnical properties a factor to account for. Moreover, the grain size was added to the variables variable list after data extraction from 16 samples using the BetterSize Lazer Particle Size <u>Analyzer Analyser 9300S</u> (Tab.Table- 2). The Grain-grain size sizes of clay, silt, and sand (Tab.Table- 2) are were spatially identified asim the same predisposing factorfactors. The sampling sites are showed shown in the Fig. 6 (red arrowarrows). The organic Organic matter content analysis is was also applied toon those-the samples using the loss on ignition (LOI) method (Heiri *et al.*- 2001), as an important factor that has a strong relationship with the presence of vegetal cover (Tab.Table 2), and who says the presence of vegetation, sayswhich indicates that the presence of water that-promotes the occurrence of landslides.

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Conditioning Factor	Number of classes	Minimum value	Maximum value	Variable Type
Elevation (m)	13	0	261	numerical
Aspect	10 or 9	Flat (-1) North 337.5–22.5°		numerical
Slope (°)	11	0 75		numerical
Curvature profile	3	-17.81 (concave)	21.1 (convex)	numerical
Curvature plan	3	-9.82 (convergent)	11.35 (divergent)	numerical
Height (m)	13	0	254	numerical
TPI	6	-88	69.37	numerical
TWI	6	-1.55	29.35	numerical
SOAR	6	0	4.72	numerical
Solar radiation (kWh/m2)	6	400	1000	numerical
Land-use	6	Bare ground, Light vegetation, Breakwater area, Dense vegetation, Cultivated areas, Roads and habitation		categorical
NDVI	5	Water-, Bare soil, Sparse vege Dense ve		categorical
Layers tilt	2	Towards sea tilting, S	Sub horizontal tilting	categorical
Grain size <del>Clay_clay_</del> (% Clays < 2 μm)	6	3	35	numerical
Grain size <del>Silt_silt_(</del> % Silt 2 μm < <del>&lt; 6</del> 3 μm)	6	6	72	numerical
Grain size <del>Sand_sand_</del> (% Sand 63_µm < ←2 mm)	6	0 91		numerical

Table 2: Input conditioning factors

Organic <del>Matter <u>matter</u> (LOI%)</del>	6	0.94	7.41	numerical
Precipitation (mm)	5	252	306	numerical
Drains network	2	0	1	categorical
Spring	2	0	1	categorical
Faulting	2	0	1	categorical
Lithology	20	See the rest	categorical	
Toe lithology	5	Grey <u>_Mm</u> arls, Marley <u>_Ss</u> andstone, Dolomitic_ limes	categorical	
Toe Protection	4	Rock platform protection, Slo protection, N	categorical	

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#### 3.3 Susceptibility modelling and validation

The method used to evaluate the susceptibility to the occurrence of coastal landslides is the Information ValueIV (Yin & and Yan, 1988; Zêzere, 2002), which is a bivariate statistical method particularly suited to study relationships between the dependent variable (landslides) and the set of independent conditioning factors. This method has been was successfully applied into coastal areas worldwide (Marques *et al.*, 2011, 2013; Epifâneo *et al.*, 2013, 2014).

Using this bivariate statistical method, it is possible to weight each class of each predisposition factor of slope instability in an objective and quantified mannerway.

The <u>Informational ValueIV</u> score (Ii) for any class Xi of an independent variable (X) was determined, for each <u>landslides landslide</u> type Y using, by the following equation:

(1)

$$Ii = ln \frac{Si/Ni}{S/N}$$

Where:

- $\rightarrow$  Si = n° of cells with landslides- and variable Xi<sub>7</sub> in the Essaouira coastal area;
- $\rightarrow$  Ni = n° of cells with variable Xi in the Essaouira coastal area;
- $\rightarrow$  S = total n<sup>o</sup> of cells with landslides in the Essaouira coastal area;
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 $\rightarrow$  N = total n<sup>o</sup> of cells in the Essaouira coastal area.

When a class of the conditioning factor defactors does not have registers of landslides (Si\_=\_0), the Ii score is not calculated <u>becausedue ofto the impossibility</u> of logarithmic normalizationnormalisation, than and it was is assumed that that class has an Ii score lower than the minimum registered. For example, the minimum IV index was -5.7014031 for <u>Slope-slope Aspect aspect</u> Class 1 (Flat flat areas) for deep translational slides; therefore, so we took -5.702 for variable classes without any landslide.

The final value of susceptibility to landslides calculated for each cell j corresponds to the sum of <u>the li</u> scores present in that unit, given by the following equation:

$I_{j} = \sum_{i=1}^{m} X_{ij} I_{i}$	(2)

Where:

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> m = number of variables;

> Xij is equal to 1 or 0, depending on whether variable Xi is present or not in cell j, respectively.

To assess coastal landslide susceptibility, 15 predictive models are-were individually developed for each inventoried landslide type in this coastal area, considering the landslide partitions defined on Tab-in Table 3, and the standard 355 model procedures defined inpreviously on section Section 3. Tab.Table- 3 shows the 15 different landslide partitions according to the landslide type used for assess assessing landslide susceptibility: total landslides, deep landslides, shallow landslides, deep rotational slides, shallow rotational slides, deep translational slides, shallow translational slides, rock toppletopples, rock fallfalls, rock slides, debris fallfalls, debris flows, and debris slides, with those these landslide dataset partitions, we expect to better understand better the different drivers responsible for the 360 occurrence of the different landslide types of landslides in this coastal area. -Each landslide\_-type inventory dataset was then sub-divided into a training and a validation group groups (Remondo et al., 2003). The Training training group containing 70% of the inventory was used in the model building and the validation group containing 30-% of the inventory was used to carry outconduct an independent cross validation process over the model first results;, The the 70/30 partition was selected randomly, because it agrees with the commonly used partitions used for landslide 365 susceptibility models model training and validation (e.g., Chen et al., 2020). We also adopted also a sensitive approach to eliminate of eliminating some landslide conditioning factors, that have little or no or less contribution toin landslides landslide occurrence basing based on the IV score results.

Additionally, to assess the importance of the representativeness of the inventory, the susceptibility modelling was also considered also, for some landslide types, splitting them in into two2 subgroups considering the depth of the rupture surface: shallow and deep-seated, for rotational and translational slide types.

Table 3: Predictive susceptibility models model strategy and landslide inventory dataset partitions

Description of the landslide		Training – 70%			Validating – 30%		
Model ID	partition dataset used for assess susceptibility	Area	Slides number	ETU number	Area	Slides number	ETU number
Model 1	All landslides (no landslide type or depth of the rupture surface differentiation)	3149643	412	682	1349847	176	292

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Model 2	Deep-seated landslides (no landslide type differentiation)	2570471	92	371	1101630	40	159
Model 3	Shallow landslides (no landslide type differentiation)	208086	75	180	89180	32	77
Model 4	Rotational slides (no depth of the rupture surface differentiation)	553238	100	281	237102	43	120
Model 5	Deep-seated rotational slides	490737	67	207	210316	29	89
Model 6	Shallow rotational slides	64840	34	74	27789	14	32
Model 7	Translational slides (no depth of the rupture surface differentiation)	2222341	67	270	952432	29	116
Model 8	Deep-seated translational slides	2082644	26	165	892562	11	71
Model 9	Shallow translational slides	143551	41	106	61522	18	45
Model 10	Rock topple (source areas)	41086	85	136	17608	36	58
Model 11	Rock fall (source areas)	175529	104	219	75227	45	94
Model 12	Rock slides	21920	11	26	9394	5	11
Model 13	Debris fall (source areas)	39314	4	21	16849	2	9
Model 14	Debris flow (source areas)	204500	33	67	87643	14	29
Model 15	Debris slide	14206	8	20	6088	3	8

For the pixel terrain unit approach, susceptibility was assessed for the different landslide types, and all dependent and independent variables were transformed into a spatial grid database withby 12.5\_×\_12.5 m resolution following the
 DEM pixel size, and all the data are-were projected in the lambert Lambert conformal conic Zone 1 coordinate system with Merchich datum.

For the ETU approach, in order-to assess landslide susceptibility, the application of any statistical method, requires the partitionpartitioning-of the study area into smaller terrain units. In the present workstudy, the main modelling is was developed on a pixel-based base-model, and the conditioning factors-factor\_layers were transformed into elementary terrains units (ETUs), considering the weight of each factor in each ETU, in order to apply the terrain units unit method and make a comparison betweencompare the two approaches (pixel and ETU).

However, susceptibility results are harmonized harmonised in elementary terrains units (ETU<sub>S</sub>). The ETU use is done because: i) they have a strong relationship with the morphology and geometry of the system that we are trying to model; ii) they are fitting to the most used land\_use planning formats as they are mostly vector approaches and system\_ based, either that is as a physical system or a human settlement; and iii) and they are also a factor of uniformity and help dealing deal with heterogeneous data (e.g. Calvello *et al.* 2015). Additionally, for planning purposes, ETU areit is easier to clearly identify the ETU in the territory when compared to pixelpixels.

4. Results and discussion

#### 4.1. Landslides in cliffs and coastal slopes of Essaouira

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The detected landslides were assigned according to the classifications of Varnes (1978)-classification<sub>2</sub>; WP/WLI's (1993); Cruden and Varnes (1996); and Dikau *et al.*- (1996); and 10 landslide types were identified: debris fall, debris flow, debris slide, rock fall, rock slide, rock topple, deep rotational slidesslide, shallow rotational slidesslide, deep translational slidesslide, and shallow translational slidesslide.

Expert and fieldwork inventory validation allowed for-landslide limit corrections and threw landslide identification of new landslides. Some examples of landslides examples are shownpresented in Fig. 3.

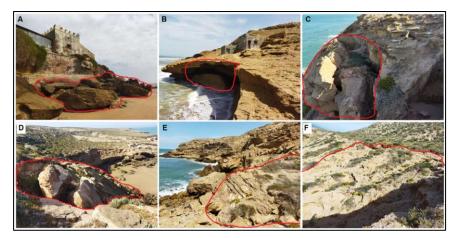


 Fig.Figure 3: Some landslide types type
 examples from the study area; A, B:-Rock-rock falls, C: Rock-rock topple, D:

 Translational translational slide, E: Rotational rotational slide with back tilting, and F: Debris debris flow 

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The final inventory of the study area is composed bycomprised 588 landslide records (Fig. 4). Rock\_falls are-were the most frequent slope instability phenomena in the study area, with 149 records, followed by rotational slides, while the least frequent landslide type is-was the debris fall, with only six <u>6</u> records. Most of the study area is-was occupied by translational slides (68-%), followed by rotational slides. These landslide types have usually-typically have bigger larger area areas per landslide, have-deeper rupture surfaces, and frequently occur along the entirewhole cliff-/-coastal slope profile.

405 Slope instability is-was present along the whole study area, resulting in the identification of 974 elementary terrain unitsETUs with landslides (63.5-%), and 28797 unstable pixels (46.5-%).

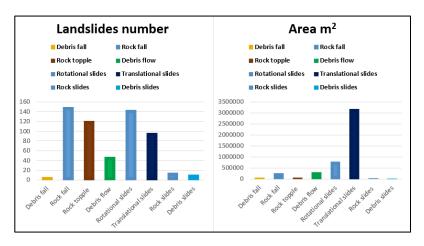


Fig.Figure 4: The relative distribution of landslides by type and area in the ETUs of the study area

Nevertheless, the heterogeneity of the spatial distribution of landslide types (Fig. 5) over the study area shows-was higher concentration in the southern section becausedue ofto the higher concentration concentrations of rotational and translational slides.

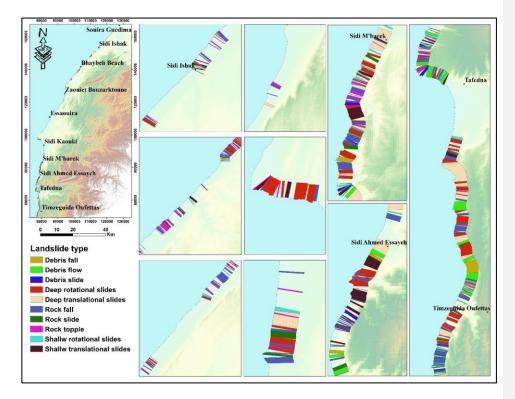


Fig.Figure 5: Spatial distribution of landslide types in the study area

#### 4.2. Driving forces of instability in Essaouira

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Lithology\_and structure, and landslide deposits are important conditioning factors for susceptibility analysis. These can be proxies <u>forof</u> permeability, shear strength and propensity for physical and chemical weathering of rock and soil materials (Varnes; 1984, Epifânio *et al.*, 2013). Twenty main lithological units were <u>found-identified</u> in the study area-including: (1) calcareous crusting; (2) clay and sandstone; (3) conglomerate and dune sediments; (4) conglomerate with sandy matrix; (5) dolomitic limestones; (6) dolomitic sandstones; (7) dune sandstone with oblique stratification; (8) Essaouira <u>Sandstonesandstone</u>/Calcarenitecalcarenite; (9) friable sandstone layer; (10) graygrey clays; (11) graygrey marls; (12) heterogeneous conglomerate; (13) limestone bar; (14) lumachelic clayey limestones; (15) marls; (16) marly limestones; (17) pudding conglomerate; (18) sandstone dolomites; (19) sequence of marls and marly limestone; and (20) -terrigenous red deposit (<u>Tab.Table</u>-4).

The spatial distribution of <u>the lithological units</u> (Tab.<u>Table</u>- 4), shows that, in general, <u>the limestone units</u> are more
 frequent in the southern sector, <u>often-frequently</u> combined with grey marls and clays of the Hauterivian and Aptian (Cretaceous). Calcareous crusting, friable sandstone layers, and terrigenous deposits <u>arecan be</u> found in all coastal

area<u>areas</u>. <u>The Conglomerate conglomerate</u> and sandstone units are more concentrated in the northern sector, where consolidate consolidated dunes can also be found.

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Regarding the number of  $ETU_{\underline{S}}$  per lithology type, calcareous crusting and Essaouira Sandstone-calcarenite are the two lithological formations most funded in the majority of  $ETU_{\underline{S}}$ , present in 1216 and 1270  $ETU_{\underline{S}}$ , respectively. This can be explained is because, in the encrustation phenomena, coastal eolian constructions become dominant in the study area as we mentioned in geological settings.

The most lithological formations occupied by the instabilities are: <u>Dune\_dune\_sandstone</u> with oblique stratification, <u>Friable\_friable\_sandstone</u> layers, <u>Gray\_grey\_Marlsmarls</u>, <u>Heterogeneous\_heterogeneous\_conglomerate</u>, <u>Limestone</u> <u>limestone</u> <u>barre</u>, <u>Marlsmarls</u>, <u>Sequence sequence</u> of <u>Marlsmarls</u>, and <u>Marlymarly</u> limestone.

Lithology	Predominance area	Number of ETU <u>s</u>	Number of unstable ETU <u>s</u>	% Of unstable ETU <u>s</u>	IV Results
Calcareous crusting	All coastal area	1216	240	19.74	-0.01
Clay and Sandstone	Southern coastal area	33	25	75.76	-1.68
Conglomerate and dune	All coastal area	1340	782	58.36	-0.31
Conglomerate with sandy matrix	Northern coastal area	33	3	9.09	-5.70
Dolomitic limestone	Southern coastal area	320	183	57.19	-1.03
Dolomitic Sandstonessandstones	Southern coastal area	13	4	30.77	-5.70
Dune sandstone with oblique stratification	Southern coastal area	284	243	85.56	0.64
Essaouira Sandstone sandstone - calcarenite	All coastal area	1270	628	49.45	-1.67
Friable sandstone layers	All coastal area	479	343	71.61	0.39
Gray Grey Clays	Southern coastal area	50	23	46.00	-2.96
Gray-Grey Marls	Southern coastal area	229	167	72.93	-1.01
Heterogeneous conglomerate	Southern coastal area	147	119	80.95	1.03
Limestone barre	Southern coastal area	159	154	96.86	0.56
Lumachelic clayey limestone	Southern coastal area	50	32	64.00	0.24

Table 4: Predominance lithology by area and ETU

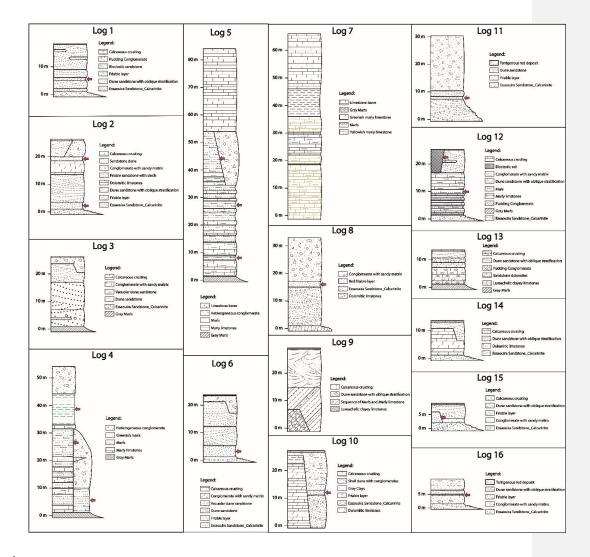
Marls	Southern coastal area	69	60	86.96	0.61
Marly limestone	Southern coastal area	67	63	94.03	0.27
Pudding Conglomerateconglomerate	Northern coastal area	152	33	21.71	-2.23
Sandstone dolomites	Southern coastal area	50	28	56.00	-0.35
Sequence of Marls-marls and Marly-marly limestone	Southern coastal area	282	275	97.52	0.70
Terrigenous red deposit	All coastal area	48	12	25.00	-1.18

Stratigraphic profiles (Figs. 6 and 7) show detailed lithological ehange-changes over the study area, and allow for a better understanding of cliff lithology-lithological variations and, the emplacement of friable layers that have a direct influence on the occurrence of landslides.

In the southern section, we noted a largebig variation in the lithological units was noted with respect to regarding the spatial distributionin therefore, the majority of stratigraphic logs are-were concentrated in the southern section (from log 1 to log 13), while there was little is no much variation in the northern section (from log 14 to log 16). RegardingAbout the lithological materials, we note the presence of friable layers or weak layers (it could be friable sandstone, sand, clays, and marls) were noted in all logs except log logs 3, log 7, log 9, log 13, and log 14.

As tilting layers are more favorablefavourable to instabilities because of the gravitational forces, the predominant subhorizontal layering has also has a contribution, while the majority of those layers are deposited on weak or friable layers, which are stimulated the instability in many-multiple locations in the study area referring to the field survey. Those These friable layers are usually typically placed between the impermeable or competent layers, which they are 450 the result of: either the different diagenesis degrees or compaction.; or the high clay\_content of elays according to grain size analysis,- whichthat makes them more friable than adjacent layers. According to the field survey, those these layers are usually generally behind the occurrence of many numerous landslides, which that is why we consider them they are considered important, especially particularly because some of them are in contact with springs, and others are in the bottom part of the cliff, which means more lithostatic pressure, and thus more susceptibility toof 455 landslides landslide occurrence.

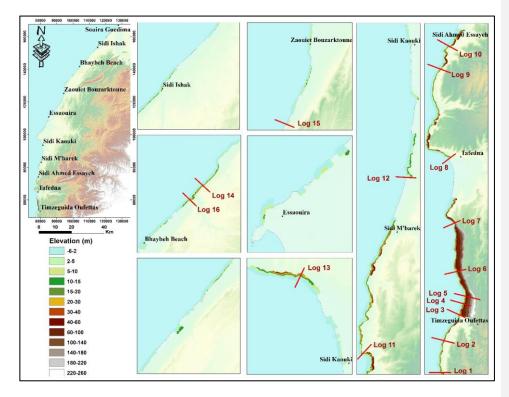
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# Fig.Figure 6: Stratigraphic columns for the Essaouira coastal area

460 <u>Elevation is Another another important factor infor</u> landslide susceptibility mapping-is elevation factor. The Fig. 7 shows the spatial distribution of these factor<u>factors</u>, and it can be seen that we can remark the southern section cliffs

present higher elevation because, for the reason that those area areas are more closed closer to the feet of the High Atlas Mountains-feet.



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Fig.Figure 7: -The spatial distribution of elevation factor in the study area with the profilesprofile emplacement.

Others Other conditioning factors are provided by the fieldwork: i) the presence and type of cliff toe protection, as it showed shown in the Fig. 8 A, B<sub>2</sub> and C, either rock platform, slope deposit, or beach protection; ii) lithology tilting, whichthat has a big impact on the landslides-landslide occurrence, as shown inwe remark in the Fig. 8 D and E; iii) 470 the presence of stream networks and springs in the cliff face which stimulate the landslides landslide occurrence; and iv) the presence of springs.25 we localized 9Nine springs were localised, four4 of which are them concentrated around Timzeguida Oufettas village which has a locally-a visible impact on landslides-landslide occurrence, especially particularly considering the presence of marls, which are becomingbecome more sliding when in contact with the water. The Other other springs are in the southern section, except for one in the north between Bhaybeh beach and Sidi Ishak village. There are They considerably affect the mechanical processes that lead to slope failure and to the subsequent post-failure movements, especially-particularly in the case of where we have marls or clays.

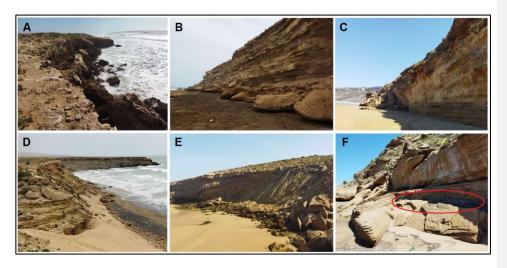


Fig.Figure 8: Examples of some conditioning factors; A: Absence absence of toe protection, B: Rock-rock platform protection, C: Beach-beach protection, D and E: tilted layers towards sea, and F: Cliff-cliff toe lithology effect-

- 480 For the rainfall factor, the interpolation of rainfall records from <u>four</u>4 meteorological stations, from 1968 to 2015, <u>waswere</u> used to assess the spatial distribution of this conditioning factor. <u>The, the</u> results <u>shows-showed</u> that the maximum average <u>of</u> 306 mm of precipitation <u>falls-fell</u> around Essaouira city, while the precipitation values <del>decrease</del> <u>decreased</u> towards the two extremities of the study area, reaching a minimum average precipitation of 252 mm.
- 485 Finally, the NDVI and land-use map werewas prepared from the Sentinel satellite images image analysis, and six land-use types were extracted, including bare ground, cultivated areas, light vegetation, dense vegetation, roads and habitation, and breakwater areaareas.

### 4.3. Coastal landslide susceptibility assessment

490 Coastal landslide susceptibility using the Information ValueIV method, as mentioned in the objectives, was produced considering two different susceptibility zonation approaches: susceptibility assessed at the pixel scale and considering elementary terrains units: ETUs.

# 4.3.1 By Pixel:

Table S1 represent-presents the information values IV scores obtained for each class of each landslide conditioning
 factor used in the construction of each susceptibility model for 15 landslide inventory partitions defined according to their classification into shallow and deep-seated landslides, landslide type, or type of affected material (debris or rock).

The information value<u>IV</u> scores represent a clear contrast between the most <u>favorable and least favourable</u> areas and the less favorable areas for the <u>occurrence of</u> different landslide type's <u>occurrence</u>, and we will describe the most important conditioning factors for each landslide type:

For all landslides landslide types (Model 1)<sub>2</sub>— The the most relevant conditioning factor forto the occurrence of all inventoried landslides are areas with slope angles >\_45 (IV score =\_1.377), followed by the solar radiation factor between 400 and 600 kWh/m2 (IV score =\_1.322) and anthe elevation factor of 60\_-100 m (IV score =\_1.320). The minimum value was obtained for the aspect class Flat-flat (IV score =\_-3.845). Those The results pointed outrevealed, considering no landslide type or depth of the rupture surface differentiation, that slope angle and elevation are the most influent factors for landslide occurrence especially particularly in dry climate areas like such as the Essaouira coastal cliff area, except for model 10 (rock topple), in which the slopes >\_15° have negative scores.

-Deep-seated landslides (Model 2)<sub>s</sub> —in the Essaouira coastal area, occurred more in areas with 400\_600 kWh/m2 solar radiation (IV score = 1.536), in slope areas >\_45-° (IV score = 1.494), and in the high areas between 60 and 100 m (IV score = 1.480), where the minimum was in the same class as previous results. AlthoughHowever, shallow mass
510 movements occurred more in friable layers with an IV score = of 3.011, in 600\_700 kWh/m2 solar radiation (IV score = 2.072), and in areas with 35\_45-° slopes.

-Rotational slides (<u>Model-Models</u> 4, <u>Model-5</u>, and <u>Model-6</u>) <u>generally</u>- occur <u>in general</u>, in <u>Sandstone sandstone</u> dolomites and dune sandstone with oblique stratification lithologies. For deep rotational slides, the grain size factor 38\_51 (% <u>Sandsand</u>) presented the highest value <u>of</u> 1.550, followed by slope angle factor class 30\_40° with <u>an</u> IV score <u>of</u>= 1.441. <u>ForWhile for</u> shallow rotational slides, the grain size factor <u>was stronglypresented a strong</u>

independence-independent of with the occurrence of this landslide type, with an IV score of= 2.323.

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-Translational slides (<u>Model-Models</u> 7, <u>Model 8</u>, and <u>Model 9</u>),- deep and shallow <u>ones-slides</u> in the Essaouira coastal area, <u>occurs-occur</u> more in areas with 400\_-700 kWh/m2 solar radiation and in slope areas >40-°.

-Rock topple (Model 10).— The Grain-grain size factor, especiallyparticularly; classes 0.—11% Silt\_silt (IV score = 2.092)-, 66.—91% Sand\_sand (IV score = 2.037), and 0.—7% Clay\_clay (IV score = 2.016), are-contribute more contributing into the occurrence of Rock\_rock topples, as because they are usually happenedtypically occur next to friable layers in the Essaouira coastal cliff area.

-Rock falls (Model 11) - occursoccur more in the "dune sandstone with oblique stratification" class of lithology factor, while the minimum IV value was -4.978 Heterogeneous heterogeneous conglomerate, which is normal as rock falls does do not happen occur in this lithology type.

-Rock slides (Model 12). <u>the The</u> lumachelic clayey <u>limestones\_limestone</u> lithology class presented a strong dependence <u>on with</u> rock slides, with <u>an</u> IV score = <u>of</u> 3.253, while the <u>Flat flat</u> (-1) areas for <u>the</u> aspect factor presented the <u>a</u> minimum IV score = <u>of</u> -3.960.

Essaouira coastal area, and the <u>Slope slope</u> angle factor class  $0_{-2}^{-2}$  is less <u>favorable favourable</u> with <u>an IV</u> score <u>-of</u> -4.822.

- Debris slides (Model 15)— presented a strong dependence <u>on with the Terrigenous terrigenous</u> red deposit class, lithology factor, while the minimum was <u>an</u> IV score =<u>of</u>-3.565 for <u>the Flat flat</u> (-1)—\_\_\_\_\_class aspect factor, which is normal <u>because</u> this landslide type occurs more in <u>t</u>Terrigenous lithologies and in <del>as</del>lope areas.

To represent landslide susceptibility for each model, we reclassify the final Information ValueIV scores were reclassified into four classes<sub>2</sub>; Very-very low susceptibility (IV score < -1), low susceptibility (-1 < IV score < 0), moderate susceptibility (0 < IV score < 1)<sub>a</sub> and high susceptibility (IV score > 1). The Fig. 9 present presents the spatial distribution of susceptibility classes for pixel-based landslide susceptibility Model 1. It can be observed is possible to observe that <u>a</u> very low susceptibility class appeared more in the northern section of the study area, whereaswhile the southern section present presented higher susceptibility to the occurrence of landslides, especiallyparticularly, becausedue ofto the weight of the translational and rotational slides in those areas.

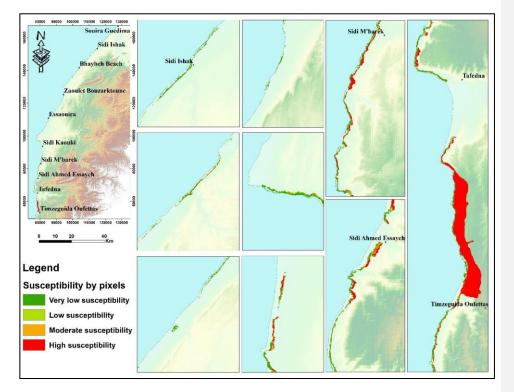


Fig.Figure 9: Landslides Landslide susceptibility map by using the pixels pixel method

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545 The information value<u>IV</u> model allowed elassifying for the classification of 38-% of our study area with high susceptibility to the occurrence of all landslides landslide types occurrence, while the very low susceptibility class is was present in 56-% of the study area (Tab.Table - 5).

All other landslide types-type susceptibility models presented high percentages for the very low susceptibility class, with a maximum of 89.76-% for debris slideslides. The exception is for debris flow, where the highest percentage was for high susceptibility with 53.85-% of the study area.

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		Very low	Low	Moderate	High
		susceptibility	susceptibility	susceptibility	susceptibility
Model 1	All landslides	55.45	2.55	2.66	39.35
Model 2	Deep-seated landslides	60.22	2.32	2.22	35.25
Model 3	Shallow landslides	72.58	4.10	3.80	19.52
Model 4	Rotational slides	52.71	7.02	6.55	33.72
Model 5	Deep rotational slides	55.03	5.84	5.95	33.18
Model 6	Shallow rotational slides	71.29	3.75	4.55	20.40
Model 7	Translational slides	61.08	2.42	2.07	34.43
Model 8	Deep translational slides	63.99	1.42	1.44	33.15
Model 9	Shallow translational slides	74.35	3.41	3.02	19.21
Model 10	Rock topple	67.41	5.52	5.95	21.12
Model 11	Rock fall	71.39	3.21	3.65	21.75
Model 12	Rock slides	80.02	2.72	2.56	14.70
Model 13	Debris fall	59.75	5.82	5.32	29.10
Model 14	Debris flow	39.15	3.04	3.96	53.85
Model 15	Debris slide	89.76	1.67	1.50	7.07

#### Table 5: Percentage of landslides landslide susceptibility classes

# 4.3.2 By elementary terrain units (ETUs)

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In general, the susceptibility assessment is carried outconducted by classifying the elementary terrain units<u>ETUs</u> into two classes: stabilized\_stabilised (37% of ETUs) and non-stabilized\_stabilised (63% of ETUs). The approach was <u>performed</u> individually for each type of landslide studied, and shows that, for all type of landslides<u>landslide types</u>, the unstable areas (classified as non-stabilizedstabilised) are located more to the <u>south-southern</u> units of <u>the</u> study area.

To represent the ETU landslide susceptibility results, we present a zoomed section of the southern section of the study area next to Timzeguida Oufettas is presented (Fig. 10), for which is possible to observe landslide susceptibility zonation can be observed for the elementary terrains units ETUs. This map presents the same allure or same variation

as the susceptibility map produced by <u>the pixels pixel approach</u>, except that, <u>in the second ETU</u> approach <u>of</u>, <u>ETU</u>, <u>we can use</u> ETU ID <u>can be used</u> to define the susceptible area in situ.

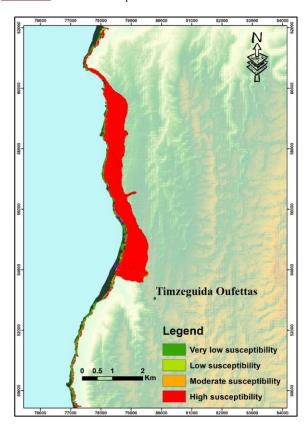


Fig.Figure 10: Landslide susceptibility map by using the ETU method for Model 1

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# 4.4. <u>Validation of Coastal coastal landslide susceptibility models validation</u>

All coastal landslide susceptibility models were validated by spatial confrontation, with the-independent landslide partitions defined as validating subsets. ROC curves (Linden 2006; and Remondo *et al.*, 2013) (Tab; Table- 6) of the predictive models were computed, and the respective Area Under Curve (AUC) value values were calculated; Tab; Table- 6 shows the AUC values obtained in the validation process for all models, as we We can remark that all landslide susceptibility models presented AUC values > 0.7 AUC values, and Model-Models 1, -Model 4, Model-10, Model-13, and Model-14 (0.7 to 0.8) are were considered acceptable. Model Models 2, Model 5, Model 6, Model 6, Model-10, Model-10, Model-10, Model-14 (0.7 to 0.8) are were considered acceptable.

7, Model-8, and Model-9 (0.8 to 0.9) are-were considered excellent, and Model-Models 3, -Model-11, Model-12, and Model-15 (more than> 0.9) are-were considered outstanding.

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Table 6: AUC values obtained in the validation process for all models.

Models	Landslide type	AUC Low	AUC High	AUC values
Model 1	All landslides	0.751	0.842	0.798
Model 2	Model 2 Deep-seated landslides		0.858	0.815
Model 3	Shallow landslides	0.735	1	0.92
Model 4	Rotational slides	0.694	0.872	0.794
Model 5	Deep rotational slides	0.709	0.889	0.813
Model 6	Shallow rotational slides	0.438	1	0.817
Model 7	Translational slides	0.759	0.854	0.809
Model 8	Deep translational slides	0.795	0.893	0.847
Model 9	Shallow translational slides	0.728	0.976	0.895
Model 10	Rock topple	0.25	1	0.75
Model 11	Rock fall	0.755	1	0.961
Model 12	Rock slides	0.827	1	0.948
Model 13	Debris fall	0.44	0.92	0.72
Model 14	Debris flow	0.561	0.878	0.731
Model 15	Debris slide	0.898	0.998	0.972

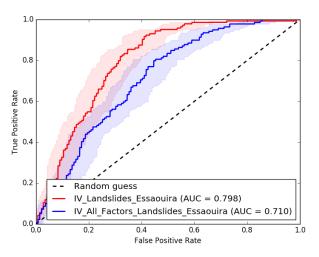
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For total landslides (Model 1) with all factors, we obtained 0.710 (Fig. 11) was obtained., then we eliminatedNext, the topographic wetness factor and the rainfall factor factors were eliminated due to the we are in dry climate of the area, and those factors did not a't present a strong dependence on with the occurrence of landslides; we obtained a value of than 0.798 (Fig. 11) was obtained, which means that the performance of model Model 1 performance was improved in term terms of prediction, adopting this sensitive approach especially particularly when we get low values of AUC values were obtained.

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We presented also the AUC graph graphs were plotted for translational slides (Model  $7_{2}$ ) 0.809; (Fig. 12) and rotational slides (Model  $4_{2}$ ) 0.794; (Fig. 13), as these two landslide types occupied <u>approximatelyabout</u> 85% of the unstable area in the <u>pixels pixel</u> model approach. These results shows show that susceptibility models have a good predictive skill and highlight the higher performance of predictive models when built for each type of landslide in comparison with the model built for the total landslides.

**Commenté [A11]:** Contractions refer to two words combined in casual contexts (e.g., "It's", "I'd", and "can't"). Avoid these in academic writing (use "It is," "I would," and "cannot", respectively).



590 Fig.Figure 11: ROC curves of the susceptibility model for all landslides with all factors (AUC = 0.710) and without TWI and rainfall factor (AUC = 0.798).

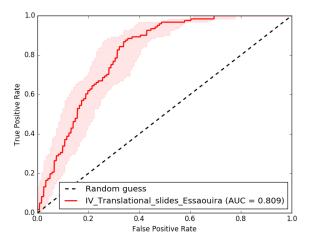


Fig.Figure 12: ROC curves of the susceptibility model for translational slides (AUC = 0.809)

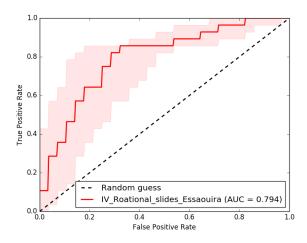


Fig.Figure 13: ROC curves of the susceptibility model for rotational slides (AUC = 0.794)

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#### 5. Conclusion

The information value <u>IV</u> bivariate statistical approach to assess landslide susceptibility assessment in the 134 km of coastal area of Essaouira, based in <u>on</u> geological, <u>and</u> morphological <u>analysis analyses</u> (interpretation of aerial photos, satellite images, and field survey), allowed elassifying for the classification of 38-% of <del>our the</del> study area with high susceptibility to <u>landslides landslide</u> occurrence (using <u>the pixelpixels</u> approach).

The translational slides followed by rotational slides are occupying aboutoccupied approximately 85-% of the landslide area, we-which\_can explain that as a matter of be explained by the fact that; the conditioning factors that are contributingcontribute more in-to\_the occurrence of those landslides, namely >\_45° slope angle, 400\_-700 kWh/m<sup>2</sup> solar radiation, and some certain lithological formations,; are occupyingwere all present in the study area, especially particularly the southern section. Another reason is that those landslides type are usually occupied these landslide types typically occupy large areas.

Landslides are distributed along the entire study area, with <u>a highermore</u> concentration in the southern section because <u>of it</u> to study and <u>it</u> to study next to study mere to study area, study area, study area, with <u>a highermore</u> concentration in the southern section <u>sections</u> of <u>it</u> to study area, with <u>a highermore</u> concentration in the southern section <u>sections</u> of <u>it</u> to study area, with <u>a highermore</u> concentration in the southern section <u>sections</u> of <u>it</u> to study area, with <u>a highermore</u> concentration in the southern section <u>sections</u> of <u>the</u> Essaouira coastal area.

For all <u>landslides\_landslide</u> types, the most important <u>explanatoryexplaining</u> drivers are; slope factor, <u>especially</u> <u>particularly > 45°</u>, solar radiation factor class 400\_-600 kWh/m2, and elevation class 60\_-100 m<sub>2</sub>, those These factors are <u>have</u> already <u>been</u> highlighted by <u>many-multiple</u> authors as important conditioning factors of <u>many-several</u> landslides landslide types. Most of the landslide susceptibility models (10 models out of 15) presented a strong interdependence with lithological factor factors or factors extracted from lithology, <u>such</u> as grain size and organic matter, which means that the landslides\_landslide\_occurrence is highly <u>affected by lithologicalimpacted by lithology</u> variations.

In the study area, precipitation <u>wasdoes notn't</u> present in our study as a decisive conditioning factor, as a consequence
 of the spatial distribution of rainfall, since the highest values are concentrated around Essaouira eity<u>City</u>, <u>which is</u> more related to sandy coast subsystems.

To define in deep-detail the spatial distribution of <u>the</u> most susceptible areas to the different landslide types along the Essaouira coastal area, <u>especially-particularly</u> in the southern section, next to Timzeguida Oufettas village, more <u>in-depthdeep</u> studies are recommended.

630 Both <u>the pixel and ETU models are holding approximativelyhold approximately</u> the same <u>valueallure</u> in all the study <u>areaareas</u>. <u>Basing Based</u> on those these models, this study <u>presented presents an essential material for spatial planning</u> and civil protection emergency actions, in <u>the Essaouira coastal area</u>, <u>especially particularly in the rocky coast</u> subsystem.

BecauseSince ETUs are closer to the morphometry of the area, there is a more "guided" analysis in this approach<sub>T</sub>
 comparing compared with pixel-based analysis that is no-not related to with a particular morphology on the cliff area. Both approaches have advantages and are inconvenient to use. It is true that The ETU approach takes more into account considers the cliff morphometry more and it's more useful for territorial management interventions, but however, it also leads to loss of susceptibility classification detail comparing compared with the pixel approach, which is more relevant in term-terms of resolution.

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