



1 Comparison of two nationwide lightning location systems and

- 2 characteristics of could-to-ground lightning in China
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8 Abstract

9 The lightning location system based on multiple sub-stations is an effective means 10 of lightning observation. This study first compares the two nationwide lightning 11 location systems in China, the Advanced TOA and Detection System (ADTD) and the 12 Three-Dimensional Lightning Location System (3D-LLS), using the observations in 13 2020. As a significantly updated version of ADTD, 3D-LLS has a cloud-to-ground (CG) 14 detection efficiency of nearly twice ADTD. However, its scarce distribution sites in 15 central Tibet, western Xinjiang, eastern Qinghai, and eastern Heilongjiang account for 16 some blind observation areas where ADTD could play a supplementary role. The ratio 17 of +CG and the distribution of currents indicate that 3D-LLS misjudges a certain 18 number of intracloud (IC) pulses as return strokes. Besides, the IC detection developed 19 by 3D-LLS is still inefficient. In the results, the IC flashes only account for half of CG 20 flashes, much smaller than the regular ratio (twice). According to the comparison with 21 the nearly no-miss optical observations in Guangzhou, the CG detection efficiency of 22 ADTD and 3D-LLS herein is 24.5% and 50.5%. Further improvements are expected in 23 terms of the above shortages. Still, the ground-based observation has better detection 24 efficiency and accuracy than the satellite, especially for CG lightning. The dataset from 25 ADTD in 2016-2021 is employed to analyze the lightning characteristics' temporal and 26 spatial distributions and the difference between +CG and -CG over China. It can be 27 concluded that low latitude, undulating terrain, seaside, and humid surface are favorable 28 factors for lightning occurrence. Thus the southeast coastland has the largest lightning 29 density, while the northwest deserts and basins and the western and northern Tibetan 30 Plateau have almost no lightning. For the period with high CG frequency (summer of a 31 year or afternoon of a day), the ratio of +CG and the discharge intensity is relatively 32 small. The Tibetan Plateau leads to the complexity of lightning activity in China and 33 lays the foundation for studying the impact of surface elevation on lightning. Results 34 indicate that the +CG ratio on the eastern and southern plateau is up to 15%, and the 35 west and north sides have a low percentage. The discharge intensity of +CG and -CG 36 on the Tibetan Plateau is approximate, while the +CG always has a larger current than





37 -CG on the plains.

38 Keywords: China, ADTD, 3D-LLS, Comparison, Lightning characteristics, +CG

39 1. Introduction

40 Most lightning is generated mainly through meso-small scale severe convective 41 weather, with few occurring in stratus clouds and tropical cyclones and occasionally 42 during volcanic eruptions, nuclear explosions, and dust storms (Rakov and Uman, 43 2003). Lightning, a violent long-distance transient discharge phenomenon, could cause 44 severe disasters such as human and animal casualties, forest fires, and electronic and 45 communication equipment interruptions. Therefore advanced lightning monitoring 46 technology is necessary for the development of lightning science and also scientific 47 protection against lightning disasters.

48 Lightning discharge emits electromagnetic spectrums with a broad range, 49 providing an essential avenue for lightning detection. The very low frequency / low 50 frequency (VLF/LF, 20-300 kHz) band radiation is mainly produced by the cloud-to-51 ground (CG) return strokes, intracloud (IC) K-changes, and other discharge processes 52 with a large spatial scale. VLF/LF electromagnetic waves could propagate along the 53 ground surface or be reflected between the surface and ionosphere propagation, with 54 superiority of long propagation distance (hundreds to thousands of kilometers) and slow 55 attenuation. This frequency range thus is suitable for large-scale lightning detection and 56 is currently the most commonly used target detection band for ground-based lightning 57 location systems.

58 Representative examples of modern lightning location systems working in 59 VLF/LF band are mainly the U.S. National Lightning Detection Network (NLDN), Los 60 Alamos Sferic Array (LASA), European Cooperation for Lightning Detection 61 (EUCLID), etc. The three nationwide detection networks in China are the Advanced 62 TOA and Detection System (ADTD) operated by the Meteorological Observation 63 Centre of China Meteorological Administration (CMA), the Lightning Location System 64 (LLS) of the State Grid Corporation of China, and the Three-Dimensional Lightning 65 Location System (3D-LLS) deployed by the Institute of Electrical Engineering of 66 Chinese Academy of Sciences (CAS). There are also small-scale and refined detection systems in local areas, such as the Beijing Lightning Network (BLNET) established by 67 the Institute of Atmospheric Physics of CAS, the Guangdong-Hongkong-Macao 68 69 Lightning Location System (GHMLLS) deployed by the meteorological departments 70 of Guangdong Province, Hongkong, and Macao.

Due to the differences in equipment types, instrument errors, calculation principles, and installation environments, there are certain deviations in the detection results of different lightning location systems. To improve the reliability, effective utilization rate, and practical application effect of the systems, and also to provide a theoretical basis for the optimization of future networking efficiency and the fusion of multi-source





76 lightning location data, it is crucial to evaluate and analyze the quality of the networks. 77 Currently, there are two dominant methods for assessing lightning location systems. 78 One is to use a limited number of rocket-triggered lightning cases and lightning current 79 measurement data or optical image data of high structures as references to statistically 80 evaluate the current or location deviation (Jerauld et al., 2005; Nag et al., 2011; Schulz 81 et al., 2016). Another way is the comparison between different systems (Bitzer and 82 Burchfield, 2016; Murphy and Said, 2020; Srivastava et al., 2017). The former gives 83 more accurate assessment results but is only available when the striking location can be 84 confirmed by other sources. For vast observation areas, the assessment of detection 85 accuracy and efficiency can only be obtained by comparing networks, often using one 86 with known better performance to assess another. If the detection efficiency (DE) of 87 both systems has not been evaluated before, a Bayesian algorithm can be used to find 88 the relative DE (Bitzer and Burchfield, 2016; Bitzer et al., 2016).

89 China spans a wide range of latitudes from north to south and significant terrain 90 changes from east to west, and the western and northern parts of the Tibetan Plateau have large uninhabited areas with altitudes above 4500 m. The above factors pose 91 92 challenges for the installation of lightning detectors and the improvement of the 93 accuracy of locating algorithms. Among the three nationwide systems, only ADTD and 94 3D-LLS are available for us. In reviewing the literature, comparative evaluation of 95 these two networks is lacking and mainly aimed at localized areas. In this paper, we use 96 the nationwide CG data in 2020 obtained from ADTD and 3D-LLS to compare their 97 capabilities in terms of lightning density, time difference, relative positioning accuracy, 98 and current peak value. The consequences could assess the quality and reliability of the 99 two datasets and their application in different scenarios. In addition, China's wide 100 latitude and longitude range and complex topography make for studying the 101 relationship between lightning and geographic factors. This study makes use of ADTD 102 data from 2016-2021 to analyze lightning distribution in China and the temporal and 103 spatial distribution difference between +CG and -CG.

104 2. Comparison of ADTD and 3D-LLS

105 2.1 Introduction to the two networks

ADTD was first developed by the National Space Science Center (NSSC) of CAS in 2007 and is currently operated by the Meteorological Observation Centre of CMA. The system consisted of 435 sub-stations (by 2020) equipped with lightning detectors and the central data processing station deployed at the National Meteorological Information Center. ADTD can generally achieve nationwide detection without blind areas and is nowadays the most widely used system by the meteorological departments in China.

3D-LLS has been in operation by the Institute of Electrical Engineering of CAS
 since 2013, which is an upgraded version of the equipment and algorithm of ADTD,





115 adding the function of detecting IC lightning and improving the efficiency of CG 116 lightning. 3D-LLS also went on to scale over 400 sub-stations by 2020, but thick 117 distributed in southern and eastern China and sparsely in western China. Part of the 118 station distribution in 2018 can be found in Wang et al. (2020). 3D-LLS has also 119 expanded to other Asian countries such as Sri Lanka, Myanmar, Cambodia, and Korea. 120 A time-of-arrival (TOA) technique with a GPS timing error of fewer than 20 ns 121 (clear sky) is used by both two networks, and the detecting targets are the VLF/LF 122 pulses of CG return strokes (and IC K-changes). A lightning flash might comprise 123 several CG strokes or IC K-changes, and both systems group single-point signals to a 124 flash event according to their separation in time and space. However, their rules of 125 categorization are different, as ADTD classified the return strokes within 500 ms time 126 interval and 10 km distance interval as a single CG flash, while 3D-LLS set the 127 thresholds as 1 s and 10-30 km to group a single CG or IC flash and distinguished the 128 two types by the discharge height. This is one factor leading to the inconsistency in the 129 number of lightning flashes between the two networks.

In this study, ADTD and 3D-LLS datasets were downloaded from the CMA big
data cloud platform. Time of occurrence, latitude, longitude, current peak value,
number of located stations, (type of lightning) for each flash was obtained.

133 2.2 Comparison of detection results

Since 3D-LLS had been adding sub-stations until 2020, this research only used the data in 2020 to compare the performance of the two networks. As 3D-LLS only retained lightning detected by five or more stations simultaneously this year, accordingly, this study did the same for ADTD data. Therefore, the retained cases were highly reliable, but it led to some missing cases where stations are sparse.

139 In 2020, ADTD detected a total of 5,898 thousand CG flashes in China, in which +CG flashes accounted for 8.5%. 3D-LLS detected 10,464 thousand CG flashes, with 140 141 a DE nearly twice that of ADTD and a +CG flash percentage of 21.4%. In addition, 3D-142 LLS also detected 4,250 thousand IC flashes. In general, IC flashes account for 2/3 of 143 the total number of lightning flashes (Rakov and Uman, 2003). It can be inferred that 144 the missed rate for IC is greater than CG, as the discharge intensity of IC K-change is much smaller than that of CG stroke, which is also the technical bottleneck of most 145 ground-based lightning location systems at present. 146

147 In our previous study, we calculated the lightning density in the vicinity of the Canton Tower in Guangdong using optical observation by the Tall-Object Lightning 148 149 Observatory in Guangzhou (TOLOG), which is currently regarded as a nearly no-miss observation (Wu et al., 2019; Jiang, 2021). The result was that the CG flash density was 150 151 20 km⁻² yr⁻¹ within a 3 km radius of the Canton Tower. The detection results in the same region for ADTD and 3D-LLS are 4.9 km⁻² yr⁻¹ and 10.1 km⁻² yr⁻¹, respectively. The 152 153 DE of the two networks correspondingly is only 24.5% and 50.5%, and improvement 154 is expected.

155 For further comparison, China's land area is divided into a 0.25°×0.25° grid





156 according to the latitude and longitude, and the annual CG flash density is calculated 157 for each grid. Fig. 1(a,b) shows the distribution results of ADTD and 3D-LLS in 2020. 158 Due to the huge regional differences, an exponential color scale is used here, with 159 regions of high density indicated by red and regions of low indicated by blue, with a 160 transition color in between. The overall distribution of the two networks is similar, and 161 the high-value and low-value regions correspond well. The specific distribution 162 characteristics will be detailed in section 3.1.



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- 177 Qinghai, eastern Heilongjiang, etc., is scarce, even leading to some blind detection areas.
- 178 Overall, the DE of 3D-LLS is about two times that of ADTD.
- 179 The current peak value of CG detected by the two networks is compared (removing 180 outliers above ± 300 kA), as shown in Fig. 2, with the orange histogram representing 181 ADTD and the slash histogram representing 3D-LLS. Most of the lightning discharged 182 with the current peak between -100 kA and 100 kA. The negative currents measured by ADTD are generally larger than 3D-LLS, with an average ratio of 1.46. For +CG, the 183 184 current peak distributions of ADTD and 3D-LLS are approximately the same when the 185 value is greater than 30 kA, with a ratio of 1.03. However, 3D-LLS has a large number 186 of outliers in the 0-30 kA range, which is presumed to result from misclassifying IC as 187 +CG, leading to a high percentage of +CG of 21.4%. This phenomenon is mostly 188 because that 3D-LLS distinguished IC and CG by the height of the radiation source. 189 According to the principle of the TOA algorithm, the positioning error in the vertical direction is much larger than that in the horizontal direction, so a significant number of 190 191 misjudgment cases appeared. Merge the electric field waveform identification to the 192 discrimination algorithm might help to improve the recognition accuracy. ADTD had 193 already eliminated the +CG with a current peak of 0-10 kA from the original data to 194 reduce misjudgment.





Fig. 2. Distribution of current peak value of CG flashes detected by the two networks in 2020

197 2.3 Match of the same CG flash from two datasets

198 Comparing the detection results for the same radiation source is necessary for 199 valuing the difference between the two networks. The spatial and temporal thresholds 200 for matching vary between studies, with most having a time difference (Δt) threshold 201 of a few hundred ms and a spatial distance (Δr) threshold of 10-30 km (0.1 ms & 1 km 202 (Pohjola and Mäkelä, 2013), 1 s & 15 km (Poelman et al., 2013), 200 ms & 20 km 203 (Bitzer and Burchfield, 2016), etc.). Different matching thresholds bring differences to 204 the results. In our study, when the Δt threshold is set to be 1 s, 3,219 thousand CG





205 flashes are matched if the Δr threshold is set to be 30 km, and the match number is 206 3,079 thousand if the Δr threshold is set to be 10 km. It means the Δr threshold has no 207 significant effect on the matching results, and the number of matched CG flashes is 208 about 60% in ADTD and 30% in 3D-LLS. The percentages imply that even though the 209 DE of 3D-LLS has been improved a lot, a large number of flashes is still undetected. 210 By expanding the Δt distribution in the 0-1 s interval, it can be found that Δt is evenly distributed in the 20 μ s-1 s interval, while a large proportion of Δt in the 0-20 μ s (0.002%) 211 212 of 1 s) interval, as shown in Fig. 3(a). There are 64 thousand (1.1% of CG flashes by 213 ADTD, 0.6% of CG flashes by 3D-LLS) flashes with Δt in 0-20 µs, and the 214 corresponding Δr distribution is shown in Fig. 3(b), with 86.5% within 2 km and 96.8% 215 within 4 km. It is reasonable to assume that these radiation sources came from the same 216 lightning strokes recorded by two systems. Bitzer and Burchfield (2016) indicated that 217 the Δt of the same stroke by different networks should be no more than 50 µs in case 218 they are timed by GPS. The two networks have different criteria for grouping flashes, 219 and if there was a missed stroke in a lightning flash and they did not use the same stroke 220 to represent the flash, the same lightning flash could not be matched in this case, leading 221 to the low matching ratio.



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Fig. 3. Difference between the matched lightning stroke of the two systems. (a) Δt distribution; (b) Δr
 distribution; (c) current comparison, green represents +CG, and blue represents -CG, same below.

227 The same stroke currents detected by the two systems are compared, and +CG and -CG are indicated by different colors, as shown in Fig. 3(c). For most of the strokes, the 228 229 current measurements are very close. However, there are some anomalies that the 230 ADTD currents are significantly higher than the 3D-LLS currents. By examining the 231 locations of these points, they are concentrated within a 50 km radius of the junction of 232 Guizhou, Guangxi, and Hunan. Since there is no special feature in this area, it is inferred 233 that some ADTD sub-stations might work abnormally. After eliminating these anomalies, the currents of +CG and -CG are fitted, respectively, and the relationship is 234 marked in Fig. 3(c) with r² of 0.90 and 0.81. The current ratios after matching strokes 235 236 are more accurate than the generalized current ratios calculated in Section 2.2. By matching and fusing the detection results of the two systems, the advantages can be 237 238 well complemented, and the existing resources can be fully exploited to achieve further 239 optimization of existing equipment.

240 3. CG characteristics of China

241 Although the DE of ADTD is less than that of 3D-LLS in the areas with dense site distribution, ADTD has a relatively more uniform site placement and a smaller blind 242 243 spot range. Moreover, the available years of 3D-LLS are short, and the misjudgment of +CG is common. Therefore, the ADTD dataset from 2016-2021 is utilized to analyze 244 245 China's lightning characteristics. The data are filtrated by the number of triggered 246 stations (\geq 3). Positive strokes with small peak currents (<10 kA) are likely to be 247 misclassified as CG discharges when those are more likely to be of intracloud nature 248 (Cummins et al., 1998), so they are discarded from the dataset.

249 3.1 CG distribution in China

250 Fig. 4 showed the surface height distribution of China based on the global surface 251 elevation data (resolution of 30 m) from NASA's new generation earth observation 252 satellite, Terra. The large latitudinal span, the great terrain disparity, and the complex 253 topography varied the climate features in China. As a fundamental meteorological 254 element, lightning mainly occurs in meso-small scale thunderstorms, and its long-time 255 accumulation characteristics are closely associated with the climate of China. 256 Atmospheric circulation, topography, distance from the sea, latitude, etc., jointly 257 determine the terrestrial spatial distribution of lightning in China.







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- 259 260

Fig. 4. Altitude distribution map of China with the location of each province and municipality (indicated by abbreviations, for details, refer to: *https://www.iso.org/obp/ui/#iso:code:3166:CN*)

261 China can be divided into four major geographical regions, namely, southern 262 China (Region-I), northern China (Region-II), northwestern China (Region-III), and the 263 Qinghai-Tibet region of China (Region-IV), as shown in Fig. 5. Qinling Mountains-264 Huaihe River line (roughly coinciding with the 0 °C isotherms and 800 mm annual precipitation line in January) is the dividing line between Region-I and Region-II. The 265 266 Daxing'an Mountains-Yinshan Mountains-Helan Mountains (dividing monsoon and non-monsoon and the 400 mm annual precipitation line) is the boundary between 267 Region-II and Region-III. The dividing line between Region-IV and Region-I-II-III is 268 roughly the line between China's terrain's first and second steps. 269

270 Fig. 5 shows the annual average CG flash density distribution from 2016-2021. 271 Most CG flashes are concentrated in Region-I with a density greater than 1 km⁻² yr⁻¹. 272 The leap line of lightning density corresponds well with the 0 °C isotherms in January, 273 the 800 mm annual equivalent precipitation line, and the eastern dividing line of the 274 first and second terrain steps. The climate in Region-I is mainly influenced by the 275 tropical/subtropical monsoon. The southeast monsoon from the Pacific Ocean and the 276 southwest monsoon from the Indian Ocean makes the summer hot and humid and prone to thunderstorms. In particular, the monsoon influence is more pronounced in the 277 278 coastal areas with abundant water vapor and thermal conditions. In the mountainous





- regions of Hainan, Guangdong, Fujian, and Zhejiang, where the rolling topography lifts
 the warm and humid air masses, thunderstorm activity is most frequent, so the lightning
 density is especially high. Although the Sichuan Basin and Yunnan are far from the
 coastline, they are located at the eastern and southern windward slopes of the Tibetan
- Plateau, which benefits the generation of thunderstorm activities due to the topographicuplift.



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Fig. 5. 2016-2021 annual average CG flash density distribution in China (Region-I: southern China; Region-II: northern China; Region-III: northwestern China; Region-IV: Qinghai-Tibet region of China)

288 Region-II has mainly the temperate monsoon climate, with summer influenced by 289 the southeast monsoon carrying temperate marine air mass or degenerate tropical 290 marine air mass, making summer warm and rainy. Most areas have CG flash density between 0.1-1 km⁻² yr⁻¹, slightly lower than Region-I. The lightning density in Region-291 292 II is also greater in the seaside area than inland areas. Shanxi is located in mountainous 293 region, and the undulating terrain makes it a high incidence area for thunderstorm 294 activity. Region-II has the most extensive plain, Northeastern China Plain, which is in 295 the shape of a trumpet, surrounded by the Daxing'an Mountains-Xiaoxing'an 296 Mountains-Changbai Mountains. The landform is conducive to the southeast monsoon 297 to reach the inland areas of Region-II and form summer thunderstorms. Jilin is only a 298 dozen kilometers from the Sea of Japan, facilitating the entry of Japanese warm air





currents. Therefore, thunderstorm activity is relatively intense in Region-II despite itshigh latitude.

301 Region-III, including Xinjiang, northern Gansu, and most of Inner Mongolia, has 302 a temperate continental climate. The southern and central parts of Region-III are mostly 303 vast deserts and gobies. The Tibetan Plateau blocks the humid South Asian monsoon, 304 and its arid surface cannot produce abundant water vapor, so almost no thunderstorms 305 are generated here. Although the Tianshan Mountains, Kunlun Mountains, Altay 306 Mountains, and Tarbahatai Mountains in Region-III are located in the hinterland of the 307 Eurasian continent, they are still provided with water vapor for thunderstorm generation 308 from the westerly circulation transporting evaporated water vapor from the Atlantic 309 Ocean and the Eurasian continent, making the northern mountainous areas occupy 310 almost all the lightning activity in Xinjiang. The southeastern monsoon flowing through 311 Region-II, reaching the eastern and central mountainous regions in Inner Mongolia in summer, can still bring some thunderstorm processes. 312

313 The main body of Region-IV is the Tibetan Plateau, mainly including Tibet, 314 Qinghai, southern Xinjiang, and western Sichuan. It has a highland mountain climate, 315 and the overall geomorphic distribution trend is increasing from west to east (Ma et al., 316 2021). The uninhabited areas above 4500 m in elevation in the west and north of 317 Region-IV are icy all year round, covered by snows and glaciers. The Qaidam Basin in 318 Qinghai is a closed, huge interrupted basin, with dry sinking airflow from the northern edge of the plateau in summer, leading to water shortage. There are few thunderstorm 319 activities in the areas mentioned above, and the distribution of sub-stations is sparse, 320 making them the regions with the lowest lightning density detected in China, with CG 321 flash density less than 10⁻³ km⁻² yr⁻¹. The eastern Himalayas, near the Yarlung Tsangpo 322 River Grand Canyon, has relatively low altitude, opening a "gap" for the influx of 323 324 abundant water vapor from the Bay of Bengal. The rapid climb of water vapor leads to 325 frequent thunderstorms. The thunderstorms around Nagqu in the central part of the plateau are mainly located between the east-west Himalayas Mountains and Tanggula 326 Mountains. The thunderstorms on the east side of the plateau are mainly influenced by 327 the low vortex and the shear line, which is usually stable at around 32.5 °N. The high 328 329 lightning density area is located precisely on the south side of the shear line.

330 The annual average CG flash density of each province and municipality is 331 calculated and ranked in Fig. 6, and their geographical locations are shown in Fig. 4. 332 Among them, Fujian (FJ) has the leading average CG flash density with more than 4.5 km⁻² yr⁻¹, followed by Hainan (HI), Zhejiang (ZJ), Jiangxi (JX) and Guangdong (GD) 333 with more than 3 km⁻² yr⁻¹. Guangxi (GX) and Guizhou (GZ) have a density of around 334 2.5 km⁻² yr⁻¹, while other inland provinces in Region-I are less than 2 km⁻² yr⁻¹. The 335 density of Shanxi (SX) in Region-II is the highest, close to 1 km⁻² yr⁻¹. In the other 336 337 provinces and municipalities in Region-II, Henan (HA), Shaanxi (SN), Hebei (HE), Beijing (BJ), Shandong (SD), Tianjin (TJ), Heilongjiang (HL), and Liaoning (LN) have 338 339 the close value between 0.2-0.4 km⁻² yr⁻¹. The density in Ningxia (NX), Qinghai (QH), Gansu (GS), Tibet (XZ), and Xinjiang (XJ) in Region-III and Region-IV is the lowest 340







344 3.2 Differences between +CG and -CG

345 According to the different polarities of neutralized charge in thunderclouds, CG 346 can be divided into +CG and -CG. Compared with -CG, +CG generally has a lower 347 probability of occurrence, only about 10%, but has a larger spatial scale and charge transfer, resulting in a more robust hazard (Preston and Tolver, 1989). Much work has 348 been done on the comparative study of +CG and -CG in local areas (Nag et al., 2014; 349 350 Rakov and V., 2003), and on their basis, this study further analyzes the spatial and temporal variability of +CG and -CG in China with the complex climatic and 351 352 geographical factors.

353 3.2.1 Comparison of the temporal distribution of +CG and -CG

354 Fig. 7(a) shows the monthly average CG flash frequency distribution for six years 355 from 2016-2021 in China. The frequency fluctuates significantly in different months, 356 with the most frequent occurrence in August (-CG flashes up to 1,345 thousand and +CG flashes up to 152 thousand). December has the least number, with only 5 thousand 357 358 -CG flashes and less than 400 +CG flashes. Lightning activity is also rare in November, 359 January, and February, with an abrupt jump in March and a transitional increase in 360 subsequent months. According to the seasonal classification, lightning activity is most 361 active in summer (June, July, and August), accounting for 70.7% of the year. In other 362 seasons, lightning is more frequent in spring (19.1%) than in autumn (9.8%) but much less than in summer. The main reason is that the summer monsoon affecting China has 363 364 started to form during April and May, while the cold and dry winter monsoon starts to build up and push southward from September, making thunderstorm activity in spring 365 and autumn mainly concentrated in southern areas, especially coastal areas. In winter, 366 China's mainland is controlled primarily by cold high pressure, when there is very little 367 368 lightning in most regions and only a small amount of lightning occurs in the southeastern coastal areas, accounting for only 0.4% of all year round. Overall, the 369 370 seasonal trend is that the lightning distribution advances from south to north and then





371 retreats southward, which is consistent with the trend of the summer monsoon. In 372 addition, the proportion of +CG flashes in different months is calculated, indicated by 373 the gray line in Fig. 7(a). The proportion of +CG flashes and the lightning frequency 374 have an obvious inverse trend. The proportion of +CG is low, less than 10%, in the 375 months with frequent lightning. In contrast, the ratio is very high in winter, when 376 thunderstorms are not easily generated, even up to more than 40% in December.

The current peak value of the two types of CG flashes in different months is 377 378 analyzed, as shown in Fig. 7(b). On the whole, the distribution range is wider in winter 379 than in other seasons. The average of each month is shown by the white cross in the 380 figure and the variation trends by red lines. From the results, the discharge intensity of 381 +CG flashes is greater than -CG flashes. The current of +CG and -CG flashes is the 382 largest in January and December. The difference between -CG flashes is not obvious in 383 other months, while the intensity of +CG flashes has an obvious trough in August. The discharge intensity and the proportion of +CG flashes have similar trends, with a higher 384 385 proportion and stronger discharge intensity in winter and a lower proportion and weaker 386 discharge intensity in summer.



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Fig. 7. Seasonal variation of (a) frequency and (b) discharge intensity of +CG and -CG flash. The gray line in (a) represents the percentage of +CG flashes, the red line in (b) represents the average current of each month, and the -CG flash currents are expressed in absolute values, the same below.

391 In addition, the hour-by-hour frequency and intensity variations of the two types 392 of CG during the day are shown in Fig. 8. The frequency of +CG and -CG flashes has 393 obvious fluctuation during the day. The variation trend of +CG and -CG is consistent, 394 and the active period is mainly concentrated in the late afternoon to midnight. Most of the summer thunderstorms in China derive from thermal effects. The solar zenith angle 395 rises after noon, radiative heating keeps enhancing, thermal conditions become 396 397 abundant and conducive to the development of convection, and lightning activity is 398 gradually activated. With the accumulation of water vapor, lightning frequency peaks 399 at 17:00. After nightfall, following the weakening of thermal conditions and unstable 400 energy, the lightning decreases continuously and drops to a low at 11:00 the next day. 401 The proportion of +CG flashes is inversely correlated with the amount of CG flashes, 402 but the phase difference lags 1-2 h. The ratio of +CG flashes is highest at 10:00, above





- 403 15%, and lowest at 15:00, only 2%, with two sub-peaks at 22:00 and 6:00, respectively.
- 404 The current distribution and average of +CG and -CG flashes are shown in Fig.
- 405 8(b). The daily variation is not apparent. The current intensity is slightly lower from
- 406 12:00 to 19:00 in the afternoon, and the lowest point is at 14:00, while the other periods407 are relatively stable.







410 3.2.2 Comparison of the spatial distribution of +CG and -CG

411 The geography of China is complex, and the ratio of +CG and -CG flashes have 412 obvious variability in spatial distribution. Fig. 9 shows the spatial distribution of the 413 proportion of +CG flashes. In order to reduce the anomalous values, the grid with less 414 than 10 CG flashes accumulated in 6 years is not included and is indicated in gray. As 415 shown, these grids are mainly distributed in the central, western, and northern parts of 416 the Tibetan and the western and southern parts of the Xinjian. The proportion of +CG 417 flashes in Region-I, where the density of CG flashes is the highest, is low, less than 418 10%. The other three regions have a higher proportion of +CG, especially the North 419 China Plain and its adjacent Inner Mongolia and parts of Region-III, where the +CG 420 ratio is up to 30-40%. Overall, lower CG density mostly means higher +CG proportion,





- 0 01 02 03 04 Ratio of +CG
- 421 and high latitude corresponds to high +CG proportion.



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Fig. 9. Distribution of the ratio of +CG flashes in China. The gray grids have a flash number of less than 10 in 6 years and thus are not calculated.

The proportion of +CG flashes in different altitude layers is calculated, as shown by the gray line in Fig. 10. Below 4500 m altitude, the proportion rises with the increase of the altitude, from 7% to 15%. There is a sub-peak at 1500 m, caused by the high ratio region of +CG flashes in Xinjiang and Inner Mongolia. Above 4500 m altitude, which is mainly the uninhabited area in the western and northern Tibetan Plateau, the percentage of +CG decreases rapidly. Only 91 CG flashes occurred above 6000 m altitude in 6 years, so they are not included in the statistics.

The current distribution of +CG and -CG at different altitudes are represented by the box plot in Fig. 10. The distribution of lightning current shrinks with increasing altitude. A fascinating phenomenon is that the average current of -CG flashes is slightly positively correlated with altitude, but +CG flashes are negatively correlated with altitude. The opposite trend of the two causes a large difference in the discharge intensity at low altitudes, but their discharge intensity is close to the same at high altitudes.









Fig. 10. Relationship between the distribution and altitude of +CG and -CG.

441 Part of the reason for the complexity of lightning activity in China comes from the 442 Tibetan Plateau, the "third pole" of the Earth, where the charge structure of 443 thunderstorm clouds has some special characteristics due to the high-altitude ground 444 surface. Qie et al. (2002) analyzed the lightning characteristics of thunderstorms on the 445 Tibetan Plateau at an altitude of 1600 m using the lightning location system installed in 446 Pingliang, Gansu, and found that the proportion of +CG flashes is 16% on average, 447 which is higher than that of conventional thunderstorms. Zhang et al. (2007) studied 448 the charge structure of 30 thunderstorms in Nagqu at 4500 m altitude and concluded 449 that although the plateau thunderstorms have a relatively strong lower positive charge 450 region, the lack of a negative charge region that excites the positive charge region to 451 discharge to the ground makes it difficult for +CG to occur. On the contrary, in our 452 study, the proportion of +CG flashes in Nagqu is higher than the average, see Fig. 9. 453 You et al. (2019) analyzed the characteristics of lightning activity in the 18-36°N Asia-454 Pacific region using lightning imager data from the TRMM satellite. The statistical 455 results showed the discharge intensity of lightning in the Tibetan Plateau is smaller than 456 in other regions at the same latitude. The results of different studies vary, and there is 457 still no clear and unified understanding of the activity characteristics of thunderstorms 458 on the plateau. Further detailed analysis is in demand in combination with topographic 459 and climatic features.

460 4. Conclusion

461 China is mainly located in the temperate and subtropical zones, under the 462 combined influence of cold and warm monsoon, the interaction of land and sea, and the 463 undulating terrain. The above factors lead to frequent convective weather combined 464 with abundant lightning activities. Compared with the previous studies on China's 465 nationwide lightning characteristics based on satellite or thunderstorm day observation





466 from meteorological stations, the lightning location system used in this paper has467 relatively higher DE and smaller location error for CG lightning.

468 ADTD and 3D-LLS are two of the three nationwide lightning location systems in 469 China, capable of CG lightning detection. 3D-LLS has an additional function of IC 470 detection. By comparing the observations of the two systems in 2020, it is found that 471 the overall DE of 3D-LLS is about twice that of ADTD. The DE distribution of ADTD is relative uniform, while 3D-LLS has extreme high DE in some parts of Yunnan, 472 473 Sichuan, Tibet, Inner Mongolia, Hebei, Jilin, etc., where the sites are dense and low DE 474 in central Tibet, western Xinjiang, eastern Qinghai, eastern Heilongjiang, etc., where 475 the sites are scarce. After matching the same radiation sources detected by the two 476 networks, it is found that their detection time difference for the same return stroke is no 477 more than 10 μ s, the distance difference is mostly within 2 km, and the current peaks 478 are almost equal. However, the DE of ADTD and 3D-LLS corresponds to only 24.5% 479 and 50.5% of the optical observation in the 3 km radius around the Canton Tower in 480 Guangdong. Besides, 3D-LLS has a high probability of misjudging IC flashes as +CG flashes, leading to the ratio of +CG flashes up to 21.4%, much higher than the ratio 481 482 (8.5%) of ADTD. The proportion of IC flashes detected by 3D-LLS is only 1/3, which 483 means the DE of IC flashes is still insufficient. Both two systems need to be further 484 improved compared with other international developed networks.

485 The ADTD dataset for the past six years combined with surface height data is utilized to analyze the CG lightning characteristics on the land of China. In general, 486 487 more in southern regions than northern regions, more in the mountains than plains at 488 the same latitude, more in humid regions than arid regions, and more in coastal regions 489 than inland regions within the same climate zone. The region with the highest CG 490 flashes density is the southeast coastland. The lowest density is in the northwest deserts 491 and basins and the east and north Tibetan Plateau. The monsoon is the main factor 492 affecting thunderstorms and lightning in southern and northern China, and the Tibetan 493 Plateau leads to the complexity of lightning activity in northwestern China and the Qinghai-Tibet region of China. Nationally, the lightning distribution is consistent with 494 495 the precipitation on a climatic scale. According to the average CG density of provinces 496 and municipalities, Fujian, Zhejiang, Jiangxi, Guangdong, etc., are in the leading 497 position. The provinces with the lowest density are Xinjiang, Tibet, Gansu, Oinghai, 498 Ningxia, etc.

499 Due to the different mechanisms of +CG and -CG flashes, their spatial and 500 temporal distribution characteristics differ greatly. In terms of time distribution, lightning activities are most active in summer (70.7%), followed by spring (19.1%) and 501 autumn (9.8%), and scarcest in winter (0.4%). Lightning in spring, autumn, and winter 502 503 is mainly concentrated in the southeastern coastal areas. In August, when CG lightning 504 is the most frequent, the proportion of +CG is the lowest, less than 10%, and its 505 discharge intensity is weak. In December, with the least lightning number, the 506 proportion of +CG rises to 40%, while the discharge intensity is stronger. Within a day, 507 17:00 is the peak period of the lightning frequency, and 11:00 is the trough. Similarly,





508 the proportion of +CG is inversely correlated with the lightning frequency, and the 509 discharge intensity of lightning is relatively weak in the afternoon when the lightning 510 is frequent. The +CG and -CG flashes ratio also show evident spatial distribution 511 variability. Southern China, with the highest CG density, has the lowest +CG ratio (less 512 than 10%). The high-latitude regions such as the North China Plain and its adjacent 513 parts of Inner Mongolia and northern and central Xinjiang have a 30-40% ratio. The ratio of +CG below 4500 m is positively correlated with the altitude and decreases 514 515 rapidly after exceeding 4500 m at the western and northern Tibetan Plateau. The 516 discharge intensity of +CG decreases slightly with the increase of altitude, while that 517 of -CG increases with the altitude. The discharge intensities of the two types have a vast 518 difference at low altitudes and tend to be the same at high altitudes. The above rules 519 have also been verified by the dataset from 3D-LLS. 520 The lightning location system sites cannot be evenly distributed due to geographic

521 factors, thus bringing about errors in the lightning distribution. The observation from 522 Lightning Mapping Imager (LMI) on the FY-4A satellite will be used to correct the

523 distribution deviations by ground-based data in our following research.

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