Letter of response to comment on nhess-2021-18

Dear Sigrid Roessner,

We thank you for your valuable comments on our manuscript and appreciate the time and the efforts you have invested. Your feedback has helped us to see and clarify ambiguous areas to further improve our work.

Based on your suggestions we have restructured the entire manuscript, especially introduction, study site description, discussion and conclusion. In addition, we have specified many conceptual and methodological concerns according to your more specific remarks. We have also rephrased several ambiguous paragraphs.

Please find below the following colour coding for the review and your comments in black; our responses to the review are in blue and the changes made to the manuscript are in green (following RC2), orange (following RC1) and in blue by the authors. Reference to line numbers are based on the original preprint.

General comments
The paper represents an interesting contribution to process oriented remote sensing based monitoring of complex landslides with the aim of making a conceptual contribution to early warning. The paper is well written in language and structure and the figures are of good quality. Despite the overall good scientific relevance and presentation quality, in the current form the paper lacks a coherent scientific goal justifying the use of approach. This problem already becomes apparent in L40 where the authors state that the study presents a new concept to systematically evaluate remote sensing techniques to optimize lead time for landslide early warning’. Although the presented work is very interesting, it does not fit the stated goal for the following reasons:

- Concept of lead time and need for best possible reduction is not new.
  While we agree that the concept itself may not be new, we find that using multispectral remote sensing products to assess and increase lead time to ensure the timely prediction of landslide early warning systems represents an important research gap that so far has rarely been addressed. We evaluate the capabilities of remote sensing to identify hot-spots and detect process behaviour changes based on the local conditions. Thus, the landslide process is the precondition. We want to estimate, based on the assumption that the particular sensor is able to deliver the necessary information, the time demand of each sensor for time to warning.
  We have now replaced the phrase optimising lead time with a more precise description of what we have done. Please see revision of the conclusion further below.

L10–11: We introduce a novel conceptual approach for comprehensive to structure and quantitatively assess lead time assessment and optimisation for LEWS.

[...]

L39–41: This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for to providing early warnings.

[...]

L34: Lead time as defined in the context of LEWS is the interval between the issue of a warning (i.e. dissemination) and the forecasted landslide onset (Pecoraro et al. 2019) and thus crucially depends on time requirements in phases
The success of an EWS therefore requires measurable pre–failure motion (or slow slope displacement) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014; Hungr et al., 2014).

- Remote sensing techniques themselves are not the bottleneck for shortening the lead time.
  
The goal of our concept is not to refine remote sensing as a technique itself but to provide a tool for choosing the appropriate sensors based on time required for the time to warning phase. We thereby increase lead time.
  
  We do not agree with your objection to the word “bottleneck” especially given your comment below which says “In remote sensing based approaches lead time mostly depends on the available imaging constellation and data distribution to the end user.”

L39–61: This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for providing early warnings. This approach reduces the number of suitable remote sensing products to a small number with high temporal and spatial resolution. With these constraints, we investigated the application of data from satellites and unmanned aerial systems (UAS) to allow the assessment of the data, after a spaceborne area-wide but low–resolution acquisition, into a downscaled detailed image recording. In so doing, we analysed the capability of these different passive remote sensing systems focusing on spatiotemporal capabilities for ground motion detection and landslide evolution to provide early warnings.

[…]

L94–102: In recent years, data provision for users has increased and today data hubs provide easy accessibility to rapid, pre–processed imagery. Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Nonetheless, technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non–existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Timely information extraction and interpretation are critical for landslide early warnings yet few studies have so far explicitly focused on time criticality and the influence of the net temporal resolution of remote sensing data.

- In remote sensing based approaches lead time mostly depends on the available imaging constellation and data distribution to the end user and in case of optical data on the atmospheric conditions (clouds). Both factors are only to a very limited extent in control of the authors - only in case of the UAV data acquisitions.
  
  Thank you for your comment. We agree that the limitation of meteorological conditions including effects such as cloud shadow and snow are important constraints as we described in L45–55 and L158. We took this into consideration when estimating
the number of available PlanetScope images (Sect. 4.2.) and discussed atmospheric affected images with regard to displacement derivation results in L477–481.

You are right that for UAS campaigns, most of the control is on the user side and only to a very limited part for other satellites. Today, some data providers promise new images daily, sometime even more frequently (e.g. PlanetScope).

But this is the point we want to highlight with our study. In a real world situation, we wish to determine which satellites can provide useful timely information in terms of an effective repetition rate and real availability in the data hub (provider). In addition, the natural conditions such as atmospheric and site specific constraints can reduce the net image number. For this reason, we assess the capabilities of optical remote sensors in a spatiotemporal context for given circumstances to detect hot spots and identify possible changes in slope processes.

L52–55: Previously, high spatial resolution satellite data was obtained at the expense of a reduction in the revisit rates (Aubrecht et al., 2017). Consequently, the return period between two images increased, limiting ground displacement assessment and the range of observable motion rates. The number of useful images was further reduced due to natural factors such as snow cover, cloud cover and cloud shadows.

[...] L86–91: In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019). However, Aa flexible, cost–effective alternative to spaceborne optical data are airborne optical images taken by UASs (unmanned aerial systems). Freely selectable flight routes and acquisition dates prevent enable avoiding shadows from clouds and topographic obstacles, and as well as allow avoiding unfavourable weather conditions and summer time snow cover, all of which frequently impair satellite images (Giordan et al., 2018; Lucieer et al., 2014).

L96–102: [...] technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non–existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Timely information extraction and interpretation are critical for landslide early warnings yet few studies have so far explicitly focused on time criticality and the influence of the net temporal resolution of remote sensing data.

- The used data sources (planet and UAV) do not allow optimization of lead time in the context of early warning because of the scarcity of their availability which is reflected in the small number of only three multitemporal data takes between July and September analyzed in this study (Table 3). Thank you. With regard to this comment we assume this needs further clarification. First, we have changed the entire phrase on “optimising lead time” to be more precise in the description of our approach (see previous comment). Regarding the data takes, yes, we do have three UAS acquisitions but over the course of more than one year (7/2018–9/2019). For the purpose of this comparison we selected PlanetScope data at a similar time to UAS acquisitions, whereby one Planet image (02.07.2018, see Table 5) showed low quality results why the time interval was excluded (see caption Fig. 4). In both UAS and PlanetScope DIC results we can see the general distinctive hot–spot
identification as well as changes in motion behaviour indicating an acceleration for the time intervals I and II. Second, we can obtain a higher frequency of UAS acquisitions if necessary. We have revised our conclusion to be more concise in our work with regard to both, the term optimisation as well as the total number of data takes.

L567–569: This paper presents an innovative concept to compare the lead time for landslide early warnings, utilising of two optical remote sensing systems. We tested this temporal concept by applying UAS and PlanetScope images of temporal proximity as these are currently the sensors with the best spatiotemporal resolution.

[…] L573–580: Our findings derived from DIC for this steep high–alpine case study show that high resolution UAS data (0.16 m) can be employed to identify and demarcate the main landslide process and reveal its heterogeneous motion behaviour as confirmed by single block tracking. Thus, validated total displacement ranges from 1–4 m and up to 14 m for 42 days. PlanetScope Ortho Scenes (3 m) can detect the displacement of the landslide central core, however, cannot accurately resolve represent its extent and internal behaviour. The signal–to–noise ratio, including multiple false–positive displacements, complicates the detection of hotspots at least in this very steep and heterogeneous alpine terrain.

Coarse temporal data resolution, such as in the case study investigated here, represents an important restriction to the use of optical remote sensing data for landslide early warning applications. Acceleration (and the resulting failure) over short periods of time will likely go unnoticed due to large data acquisition intervals. However, for prolonged acceleration periods, such as observed at the Sattelkar slide and many other relevant hazard sites, the chosen data sources have been demonstrated to represent a formidable early warning approach capable of contributing to an improved risk analysis and evaluation in steep high–alpine regions.

[…] L589–594: For continuous monitoring and early warning, the warning time window could be shortened by on–site drone ports with autonomous acquisition flights and automatic processing. Our systematic evaluation of the sensor potency capability can be applied and transferred to other optical remote sensing sensors, and the same is true for our conceptual approach optimising which extending the lead time. Future studies should focus on the applicability of complementary optical data to confirm the detection of landslide displacement and adjust UAS output resolution as this significantly increases the validity of DIC internal ground motion behaviour.

- The missing sound conceptual approach is also reflected in the introduction in form of a lengthy summary of in principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 1 is very general and not specific to landslide and does not qualify as a novelty in the current form.

1. Introduction
We revised the abstract and the introduction, to be more precise with regard to our goal and implementation. In so doing we more clearly defined our approach to lead time and early warning systems for landslides. Further we did our best to improve the line of arguments and to show the historic limitations of optical remote sensing for LEWS up to the recent developments when it comes to options such as high spatiotemporal products and their usage for monitoring, early warning and time-series displacement analyses.
2. The conceptual approach

We decided to keep this concept general, to employ it for other remote sensing techniques and maybe even other kind of instrumentation as well as different use cases of other time challenging issues. We revised and added some sentences to emphasise our approach/idea. Even after intense research we did not find good conceptual approaches challenging remote sensing in the direct context of landslide early warning systems. We therefore consider our approach novel. This concept forms the basis to employ this for the setup of ‘a real early warning system’.

L21–102: Landslides are a major natural hazard leading to human casualties and socio-economic impacts, mainly by causing infrastructure damage (Dikau et al., 1996; Hilker et al., 2009). They are often triggered by earthquakes, intense short–period or prolonged precipitation, and human activities (Hung et al., 2014; Froude and Petley, 2018). In a systematic review Gariano and Guzzetti (2016) report in a review study that 80% of the papers examined papers show causal relationships between landslides and climate change. The ongoing warming of the climate (IPCC, 2014) is likely to decrease slope stability and increase landslide activity (Huggel et al., 2012; Seneviratne et al., 2012), which indicates a vital need to improve the ability to detect, monitor and issue early warnings of landslides and thus to reduce and mitigate landslide risk.

Early warning, as defined by the UN International Strategy for Disaster Reduction (UNISDR), refers to a set of capacities for the timely and effective provision of warning information through institutions, such that individuals, communities and organisations exposed to a hazard are able to take action with sufficient time to reduce or avoid risk and prepare an effective response (UNISDR, 2009). According to UNISDR (2006), an effective early warning system consists of four elements: (1) risk knowledge, the systematic data collection and risk assessment; (2) the monitoring and warning service; (3) the dissemination and communication of risk as well as early warnings; and (4) the response capabilities on local and national levels. Incompleteness or failure of one element can lead to a breakdown of the entire system (ibid.). Lead time as defined in the context of LEWS is the interval between the issue of a warning (i.e. dissemination) and the forecasted landslide onset (Pecoraro et al. 2019) and thus crucially depends on time requirements in phases (1)–(3). The success of an EWS therefore requires measurable pre–failure motion (or slow slope displacement) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014; Hungr et al., 2014).

While remote sensing has been established for early warnings, remote sensing is not yet used for real early warnings of the onset of landslides in steep-alpine terrain (with a few exceptions), where geotechnical instruments are still preferred. Exceptions include terrestrial InSAR (Pesci et al., 2011; Walter et al. 2020) and terrestrial laser scanning with high repetition rates. However, repeated UAS (unmanned aerial systems) and optical satellite images (PlanetScope) with high repetition rates have so far not been applied for landslide early warning in steep-alpine catchments. In this regard, knowledge of sensor capabilities and limitations is essential, as it determines which rates and magnitudes of pre–failure motion can potentially be identified (Desrues et al., 2019). Our proposed framework refers to mass movements in steep-alpine catchments with significant pre–failure motion operating over a sufficient time periods and thus excludes instantaneous events triggered by processes such as heavy rainfalls or earthquakes.

This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote–sensing data for providing early warnings. This approach reduces the to a small number the of suitable remote sensing products
to a small with high temporal and spatial resolution. With these constraints, we investigated the application of data from satellites and unmanned aerial systems (UAS) to allow the assessment of the data, after a spaceborne area-wide but low-resolution acquisition, into a downscaled detailed image recording. In so doing, we analysed the capability of these different passive remote sensing systems focusing on spatiotemporal capabilities for ground motion detection and landslide evolution to provide early warnings.

Until recently, the spatial and temporal resolution of optical satellite imagery has significantly improved requirements for accurate early warning purposes have not been met by optical satellite imagery (Scaiouoni et al., 2014) and has allowed substantial advances in the definition of displacement rates and acceleration thresholds to approach requirements for early warning purposes. This is essential since spatial and temporal resolution determines whether landslide monitoring is possible with the detection allows defining displacement rates and approximation acceleration thresholds, both of which are lacking if information is based solely on post-event studies (Reid et al., 2008; Calvello, 2017). Landslide monitoring offers the potential to significantly advance landslide early warning systems (LEWS) (Chae et al., 2017; Crosta et al., 2017). Previously, high spatial resolution satellite data was obtained at the expense of a reduction in the revisit rates (Aubrecht et al., 2017). Consequently, the return period between two images increased, limiting ground displacement assessment and the range of observable motion rates. The number of useful images was further reduced due to natural factors such as snow cover, cloud cover and cloud shadows. High-resolution remote sensing data was long restricted due to high costs and data volume (Goodchild, 2011; Westoby et al., 2012). Today commercial very high resolution (VHR) optical satellites exist, but tasked acquisitions make them inflexible and very cost intensive, thus limiting research (Butler, 2014; Lucieer et al., 2014). There is a vast spectrum of available remote sensing data with high spatiotemporal resolution (Table 1). Complementary use of different remote sensing sources can significantly improve landslide assessment as demonstrated by Stumpf et al. (2018) and Bontemps et al. (2018), who draw on archive data and utilise different sensor combinations to analyse the evolution of ground motion.

Table 1 Overview of different optical multispectral remote sensors with their corresponding resolution [m] and revisit rate [days]. The sensors are categorised into commercial and free data policy. 1Free quota via Planet Labs Education and Research Program. 2PlanetScope Ortho Scene Product, Level 3B/Ortho Tile Product, Level 3A (Planet Labs, 2020b). 3Reached end of life, 3/2020, archive data usable, 45 m Ortho Tile Level 3A (Planet Labs, 2020a), 50.5 m colour pansharpened, 6self-acquired. Source: (ESA, 2020).

<table>
<thead>
<tr>
<th>Sensor</th>
<th>Temporal resolution [d]</th>
<th>Spatial resolution [m]</th>
<th>Free/Commercial</th>
</tr>
</thead>
<tbody>
<tr>
<td>UAS</td>
<td>flexible</td>
<td>0.08</td>
<td>F(^6)</td>
</tr>
<tr>
<td>WorldView 2</td>
<td>1.1</td>
<td>1.84</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 3</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 4</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye 2</td>
<td>5</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>SkySat</td>
<td>1</td>
<td>1.5</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye-1</td>
<td>3</td>
<td>1.64</td>
<td>C</td>
</tr>
<tr>
<td>Pléiades 1A/B</td>
<td>1</td>
<td>2.0 (0.5)(^5)</td>
<td>C</td>
</tr>
<tr>
<td>PlanetScope</td>
<td>1</td>
<td>3.0/3.125(^2)</td>
<td>C/F(^1)</td>
</tr>
<tr>
<td>RapidEye(^3)</td>
<td>5.5</td>
<td>5(^4)</td>
<td>F</td>
</tr>
<tr>
<td>Sentinel-2 A/B</td>
<td>5</td>
<td>10</td>
<td>F</td>
</tr>
<tr>
<td>Landsat 8</td>
<td>16</td>
<td>30</td>
<td>F</td>
</tr>
</tbody>
</table>

The latest developments in earth observation programs include both the new Copernicus’ Sentinel fleet operated by the ESA, and a new generation of micro cube satellites, sent into orbit in large numbers by Planet Labs Inc. These PlanetScope micro cube satellites, known as Doves/PlanetScope (from now on referred to as PlanetScope
satellites), and Sentinel-2 a/b offer very high revisit rates of 1–5 days and high spatial resolutions from 3–10 m, respectively (Table 1), for multispectral imagery (Drusch et al., 2012; Butler, 2014; Breger, 2017). This opens up unprecedented possibilities based on these high spatiotemporal resolutions to study a wide range of landslide velocities and natural hazards through remote sensing. Future data access is fostered by PlanetLabs and by Copernicus (via its open data policy) providing affordable or free data for research. This leads to unprecedented possibilities for studying natural hazards through remote sensing. Examples of landslide activity studies employing multi-temporal datasets of landslide activities based on this access to high spatiotemporal data include Lacroix et al. (2018), using Sentinel–2 scenes to detect motions of the 'Harmalière' landslide in France, and Mazzanti et al. (2020), who applied a large stack of PlanetScope images for the active Rattlesnake landslide, USA.

As forecasted, landslides tend to accelerate beyond the deformation rate observable with radar systems before failure, we concentrate on optical image analysis (Moretto et al., 2016). One advantage of optical imagery is its temporally dense data (Table 1) compared to open data radar systems with sensor visit repeat frequency more than every six days and revisit frequency between three days at the equator, about two days over Europe and less than one day at high latitudes (Sentinel–1, ESA). Optical data allows direct visual impressions from the multispectral representation of the acquisition target and the option to employ this data for further complementary and expert analyses. While active radar systems overcome constraints posed by clouds and do not require daylight, data voids can be significant due to layover or shadowing effects in steep mountainous areas (Mazzanti et al., 2012; Plank et al., 2015; Moretto et al., 2016). Moreover, north/south facing slopes are less suitable, thus limit the range of investigation (Darvishi et al., 2018). In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019).

However, a flexible, cost-effective alternative to spaceborne optical data are airborne optical images taken by UASs (unmanned aerial system). Freely selectable flight routes and acquisition dates prevent enable avoiding shadows from clouds and topographic obstacles, and as well allow avoiding as unfavourable weather conditions and summer time snow cover, all of which frequently impair satellite images (Giordan et al., 2018; Lucieer et al., 2014). UAS–based surveys provide accurate very high resolution (few cm) orthoimages and digital elevation models (DEM) of relatively small areas, suitable for detailed, repeated analyses and geomorphological applications (Westoby et al., 2012; Turner et al., 2015).

In recent years, data provision for users has increased and today data hubs provide easy accessibility to rapid, pre-processed imagery. Knowledge of the most useful remote sensing data options is vital for complex, time-critical analyses such as ground motion monitoring and landslide early warning. Nonetheless, technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non-existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time-critical analyses such as ground motion monitoring and landslide early warning.
• L140: General applicability to optical data: This subheading does not fit the content of this section comprising a compilation of rather basic and general steps of remote sensing data processing. Thank you for your comment. We agree that it describes general steps of the data processing chain; however, these steps are applied within each phase of the ‘time to warning’ of our proposed concept. Otherwise the steps would not be explained and thus the basis for the concept would be lacking. We have revised the subheading to “Practical implementation of multispectral data in the concept” which more accurately describes the content of this section.

2.2. Practical implementation of multispectral data in the concept

The study site (starting at L175) represents a very complex landslide case leading to rather erratic mass movements in form of debris flows initiated by changing slope water conditions related to increased atmospheric precipitation. This situation is another obstacle for an early warning approach which is solely based on optical remote sensing data and thus making it impossible to make full use of the in principle daily temporal resolution of the planet data. Taking into account these natural conditions and the constraints introduced by the used imaging constellations, leaves no room for true optimization of lead time in the sense as stated in the overall scientific goal of this paper.

We agree with your assessment and have replaced the term “optimisation” with a description that hopefully is more accurate in the entire manuscript. The chosen Sattelkar slide is one of the most relevant high-alpine geohazards in Austria and thus represents a compelling study site for natural hazard studies. While we agree that its complexity represents an obstacle, we nonetheless believe that the Sattelkar slide is well-suited for an investigation based on optical remote sensing because (i) we were clearly able to detect significant displacement and (ii) we were able to identify patches of increasing motion. In any case an increase in frequency of UAS flights is possible.

L39–41: This study presents a new concept to systematically evaluate remote sensing techniques to estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for providing early warnings.

• Any sensible early warning approach for slope movements requires a continuous and reliable high temporal resolution input of observation data related to parameters which are relevant for triggering the potential mass movements. Such information are mostly provided by ground based measurements. In this context, it is surprising that no relevant ground based monitoring information seem to be available to this study despite the longterm history of scientific work at this study site. The mentioned temperature loggers need to be explained in their function for early warning. The GPS measurements seem to only support the remote sensing based analysis. The described setting does not seem to be suitable for identification of precursory signs of ‘slope preparation’ related to the triggering of potential mass movements at this site in a way which would be required in the context of early warning.

Thank you for your feedback. We understand your arguments, yet we are not trying to
create an all-encompassing landslide early warning study that includes all state-of-the-art methods. We have chosen the Sattelkar due to its scientific and societal relevance and its high-alpine location with very limited vegetation. This site was not selected to evaluate a wide range of remote sensing applications. Our goal was to determine if and how our conceptual approach is applicable to this highly complex study site. Due to its topographical characteristics no ground based technique can be implemented. Therefore, only air- and spaceborne sensors can be employed which we believe is the case for numerous potentially hazardous slides/creeps in mountain ranges worldwide. However, we have considered installing a camera on the opposite slope but currently the distance is a problem (3.5 km, selection of camera).

We agree that the temperature data mentioned in the manuscript is not absolutely necessary to understand our conceptual approach. We still think that the (brief) inclusion of the temperature data makes sense as it suggests local permafrost presence/degradation which may be one of the main drivers of the Sattelkar slide. To clarify the role of the temperature data we amended the relevant sections in the study site section.

L175 et seq. […] massive volumes of glacial and periglacial debris as well as rockfall deposits (Fig. 2b, c). Near surface temperature data indicates sporadic permafrost distribution in the upper part of the cirque. […] allowing visual block tracking and delimiting the active process area. High displacement was measured between 2012 and 2015 with up to 30 m a⁻¹.

L200 et seq.: In the Sattelkar cirque, several monitoring components are installed to provide ongoing and long-term monitoring. Nine permanent ground control points (GCPs) measured with a dGPS to provide stable and optimal conditions to derive orthophotos from highly accurate UAS images (GeoResearch, 2018). A total number of 15 near surface temperature loggers (buried at 0.1 m depth) recorded annual mean temperatures slightly above the freezing point (1–2 °C) in the period 2016 to 2019. Ground thermal conditions at depth react with significant lag times to recent warming and therefore are primarily determined by climatic conditions of the past (Noetzli et al., 2019). Significantly cooler climatic conditions in previous decades and centuries (Auer et al., 2007) thus likely contributed to the formation of (patchy) permafrost at the Sattelkar cirque. Recent empirical-statistical modelling of permafrost distribution in the Hohe Tauern Range confirms possible permafrost presence at the study site (Schrott et al., 2012).

These components include 30 near surface temperature logger (NSTL) nine permanent ground control points (GCP) measured with a dGPS to provide stable and optimal conditions for the derivation of orthophotos from highly accurate UAS images (GeoResearch, 2018). Field-based mapping and measurements help to delimit the active process area.

Correct, the dGPS measurements are only used for repeated UAS campaigns and their data derivation. As described earlier, with our technical approach we were able to not only detect hot spots of total displacement but also to see changes in motion and thus certain areas of accelerating behaviour.

- L210: The complete dismissal of radar data is not justifiable in the current form since the authors only take into account InSAR based deformation analysis and neglect that
the technique of pixel offset tracking can be also be applied to the intensity component of radar data. For the mainly rainfall driven processes at the study site, the integration of radar data seems to be mandatory into any sensible remote sensing based early warning approach, since a combination of optical and radar data is required to establish an as continuous as possible time series of remote sensing observations. Thank you for mentioning radar data. We have described the application of InSAR/DInSAR in the introduction (L86–91) and placed the argument in section “4.1. Optical Imagery”.

For this particular site radar data is not practical. Even if foreshortening and layover effects are a minor issue for this site, the main reason to not include this kind of data is the fact that the velocity shows rates exceeding the limits of radar data leading to a loss of coherence.

L78 et seq.: As forecasted, landslides tend to accelerate beyond the deformation rate observable with radar systems before failure, we concentrate on optical image analysis (Moretto et al., 2016). One advantage of optical imagery is its temporally dense data (Table 1) compared to open data radar systems with sensor visits repeat frequency more than every six days and revisit frequency between three days at the equator, about two days over Europe and less than one day at high latitudes (Sentinel–1, ESA). Optical data allows direct visual impressions from the multispectral representation of the acquisition target and the option to employ this data for further complementary and expert analyses. While active radar systems overcome constraints posed by clouds and do not require daylight, data voids can be significant due to layover or shadowing effects in steep mountainous areas (Mazzanti et al., 2012; Plank et al., 2015; Moretto et al., 2016). Moreover, north/south facing slopes are less suitable, thus limit the range of investigation (Darvishi et al., 2018). In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019).

Moreover, taking into account the goal of lead time optimization, I consider it crucial to also include ground-based live-streamed time-lapse imagery in the proposed remote sensing based early warning approach (for an example see the Khan et al. (2021) paper ‘Low-Cost Automatic Slope Monitoring Using Vector Tracking Analyses on Live-Streamed Time-Lapse Imagery’ published in Remote Sensing). Thank you for this idea and forwarding the information on the article of this useful approach for the ‘Rest and Be Thankful slope’, Scotland, with PIV on time–lapse imagery. For the Sattelkar we conducted preliminary investigations regarding the installation of a camera on the opposite slope. Due to the steep slope the camera would have to be mounted at the same altitude. This means a camera would have to be able to cover a horizontal distance of about 3.5 km. There is a higher chance of mobile network signal which is otherwise unavailable beginning at the entrance of the valley. Nevertheless, the power supply and issues such as rain drops and general pollution on the lense pose problems as Khan et al. (2021) also acknowledge.

The materials and methods section (4.) as well as the result section (5) are sound and well written. Since reviewer 1 has already focused on this part of the paper as well as the accuracy assessment and made detailed suggestions for improving these parts, I only have a few comments left to make on these aspects of the paper.

L355: The authors state that core areas of the landslide are surrounded by wide fringes with no data. In this context the meaning of the term ‘no data’ is not clear to me.
Please, explain, what do you mean by ‘no data’ – either missing results or zero deformation.

Thank you for pointing this out. Here by ‘no data’ we mean that there is zero deformation and we have revised the text accordingly.

L354 et seq.: No motion was present in a fringe zone along the landslide front (west boundary), similar to results in Fig. 5a and Fig. 5b. In general, the displacement patterns are less smooth than at 0.16 m input resolution. Outside the landslide significant displacements exist at the eastern image border (Fig. 5e) and towards the west (h, i) (Fig. 5f). In comparison, total displacement rates derived from PlanetScope cover in large parts the active area for Ib (Fig. 5c); however, for II only the core area of the landslide shows displacement. In both results the core areas of the landslide are surrounded by wide fringes with zero deformation.

- L370: Fig 6. The obtained deformation results show a very different degree of detail throughout the landslide. For better evaluation of the reasons for these differences the inclusion of an RGB UAV image of the same area would be helpful in order to be able to include surface texture properties in the evaluation of the obtained differences in the deformation patterns.

Thank you for your good suggestion. We added the corresponding master and slave image below the presented DIC result. The caption has been adjusted accordingly.
Conclusions related to the results presented until L370: The presented specific deformation results obtained from the analyzed planet and UAV data, represent a valuable contribution towards an improved area-wide process understanding of so far unprecedented detail for this study site. Conceptually, such investigations mainly contribute to the preparedness phase within the disaster management cycle. Continuation of monitoring of the study site using the described approach would represent a very valuable prerequisite for developing and setting up a true early warning system for this site combining ground based and remote sensing observations. However, the results presented in this paper do not allow optimization of lead times within an early warning approach being stated being as the goal of this paper. Our approach is not to set up a comprehensive early warning system, which includes all four elements defined by the UNISDR (2006) (see L35–38).

We agree that optimisation of lead time does not accurately represent what we have done in our study. Thus we have revised our manuscript to make it more precise (see...
changes to the manuscript here on p. 1, 3–4). Our concept enables us to evaluate lead time based on our proposed structure.

Introduction, L10–11: We introduce a novel conceptual approach for comprehensive to structure and quantitatively assess lead time assessment and optimisation for LEWS.

[...] L39–41: This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for early warning.

[...] Conclusion, L578 et seq.: Coarse temporal data resolution, such as in the case study investigated here, represents an important restriction to the use of optical remote sensing data for landslide early warning applications. Acceleration (and the resulting failure) over short periods of time will likely go unnoticed due to large data acquisition intervals. However, for prolonged acceleration periods, such as observed at the Sattelkar slide and many other relevant hazard sites, the chosen data sources have been demonstrated to represent a formidable early warning approach capable of contributing to an improved risk analysis and evaluation in steep high–alpine regions.

- L375: 5.3 Time required for collection, processing and evaluation. The presented analysis is rather meaningless, since the scarcity of the available time steps does not allow the detection of critical process stages. Taking into account the big temporal gaps between the data acquisitions, the time needed for handling the planet and UAV imagery is not really relevant for lead time optimization. The obtained times only allow a relative comparison between planet and UAV based data acquisition within the narrow limits of the chosen approach. However, true early warning would require setting up a semi-automated processing chain including automated download and screening of available remote sensing data as well as semi-automated subsequent deformation analysis reducing data handling time to a minimum. Under such conditions, primary remote sensing data availability becomes the crucial decisive factor determined by the data distribution procedures of the satellite data providers and the atmospheric conditions in case of optical imagery. In conclusion, it needs to be stated that the used parameter of time to warning is only applicable under the condition of a near real time continuous data stream of input information which is not available within the presented study.

Thank you for your comment which helps to clarify your understanding of our text. We did not intend to create a ‘true early warning’ as you described. This was not the goal of our study. The repeated measurements allow the detection of spatial and temporal acceleration patterns and we believe the repeated measurements can be scaled to early warning demands. With regard to your comment on a semi-automated processing chain we do not fully agree. Based on our knowledge, even in case of most geotechnical investigations, the data is analysed by experts prior to issuing an early warning (e.g. https://www.bgu.tum.de/landslides/alsense/projekt/, Leinauer et al. (2020): DOI: 10.1002/geot.202000027).

- L390: In the current form of the paper the points raised in the discussion (6.) are only relevant in the frame of a process-oriented study and not for early warning purposes
since the latter one requires the identification of precursors for critical process stages – tipping points – which are likely to trigger substantial complex mass movements later turning into potentially catastrophic debris flows.

It is our understanding, we can only provide early warnings for processes we understand. The processual understanding is key to anticipating the magnitude, timing, and reach of alpine hazards, thus processual understanding and early warning cannot be separated.

- **L490:** Estimating time to warning (6.3). This part of the discussion also suffers from the conceptual limitations which have already been pointed out earlier in this review. A comparison of lead times between the different example landslides would only be meaningful in case of continuous high resolution temporal information on deformation allowing the identification of precursory events which is usually only possible using ground based observations. The presented comparison between potential repeat rates of remote sensing data acquisitions and retrospectively derived lead times is too simplistic (Fig. 8), since the main remaining question is, whether the relevant deformation (cracks etc.) can be first, resolved by the used imagery and second, distinguished from other surface disturbances by the used analysis methods. In this paper, in contrast to remote sensing papers, the time scale required for effective early warnings is given by nature, i.e., the typical acceleration patterns of particular landslides.

With regard to the comparison of historic events, we referred to their natural landslide processes which delimits the possible lead time. Unfortunately, a comparison to these historic examples is limited to a retrospective view. We agree with you regarding the detection of relevant deformations. If the sensors evaluated here could have identified the motion excluded disturbances, then in this temporal concept UAS and PlanetScope would have been able to show an acceleration in a timely fashion. We want to keep this concept simple to allow the transfer for required processing times from other sensors. The main question is, if the time is sufficient for the whole processing prior to landslide release.

L:148–149 Natural processes and their developments constantly take place independently, thus dictate the technical approaches and methodologies researchers must can and must apply within a certain time period.

**Overall recommendation:**
The presented results comprise a very interesting process-oriented study evaluating the use of planet and UAV imagery for the derivation of spatiotemporally differentiated deformation information for a rather large and topographically pronounced terrain affected by complex mass wasting processes. I consider these findings well worth being published in this journal. However, the publication of these specific results requires a major conceptual reframing of the work which is targeted at the real potential usability of these results which cannot be early warning because of the reasons already stated in this review.

However, the work presented in this study has the potential to form an important basis for the development of a true early warning concept / approach in the future combining remote sensing and ground based observations targeting at the same parameters allowing a multi-scale assessment of surface deformation related to triggering potential catastrophic mass movements at the study site.
Letter of response to comment on nhess-2021-18

Dear Jan Blöthe,

We thank you and appreciate your valuable comments on our manuscript. Your feedback has helped us to improve our work and pointed to areas which were ambiguous and therefore needed clarification.

Please find below the following colour coding for the review and your comments in black; our responses to the review are in blue and the changes made to the manuscript are in green.

General comments

A) Description of digital image correlation method and error assessment

In my view, digital image correlation is not a trivial method and deserves a more detailed description in section 4.3. Especially because the conceptual approach presented here grounds on the detection of significant movement (or even acceleration) from optical imagery, the authors should elaborate the exact processing steps and include a detailed accuracy assessment. This can easily be achieved by:

- The quantification of a level of detection between images, i.e. the residual mismatch of stable surfaces outside the landslide between consecutive images after image correlation, beyond which significant displacement can be detected with a given confidence.
- Excluding spurious matching results (displacement vectors) on the basis of a correlation threshold.

The description of section 4.3., Data Acquisition and Processing, has been modified by adding more details.

The attached Online Supporting Material (OSM) contains the variety of results which show our approach to selecting the appropriate combination of UAS input data (orthophotos, DSM and hillshade derivates) and displacement vectors (see OSM Figs. 7, 8 and 9). In addition, signal to noise results and volume calculations are provided (see OSM Figs. 3, 5, 8, 9 and 11, 12). The distribution of GCPs combined with DIC total displacement results of UAS are also presented (see OSM Fig. 1 and 4).

In terms of the selection of appropriate parameter settings, we decided to use:

- for a step size of one, as larger step sizes smoothed the velocity pattern, did not obviously improve the matching while decreasing the spatial resolution. Computation time would decrease if larger step sizes are employed.
- UAS 128 x 32, as an initial window of 256 returned a general decrease in velocity. Furthermore, the smaller initial window of 64 matching was only partially successful with very low velocities. The final window size is important to detect small scale features. If set too large, features could be smoothed out. In our case there were no distinct differences, which is why we selected the smaller final window option: to necessary small scale features.

However a detailed accuracy assessment requires comparable data which is not available such as in the verification process of DEM production based on stable surfaces. Therefore, we...
added the signal to noise results as you requested in the OSM. An accuracy assessment similar to Travelletti et al. (2012) having GCPs within the active landslide cannot be conducted, as in contrast our GCPs are located on stable positions outside of the active landslide (see OSM Fig. 1 and 4). Our approach to this study is to compare manual block tracking with the calculated velocities from DIC as part of the data evaluation.

B) Result of image correlation

As stated above, digital image correlation and the extraction of displacement from correlated imagery is not a trivial task and many pitfalls can lead to spurious results (the authors term these decorrelated). I will outline my doubts regarding the validity of the obtained displacement values referring to Fig. 5, but have given many detailed comments on the respective text positions in the specific comments below. In large areas, the image correlation returns areas that are “decorrelated”, such as the western part of the landslide in (a) and (b), but also positions in (e) and (f) are affected by this. In my experience, such a pattern indicates that matching between images did not work, which should be visible by adjacent vectors having very different magnitudes and directions. Furthermore, the patchy nature of displacement values in the western part of (c) is very surprising. Here very high total displacement of ~18 m is located in the vicinity of displacement on the order of 4-8 m. From an image matching procedure, I would expect a rather smooth picture here, such as in (d). But also from a geomorphic perspective, I am unsure how this pattern could be explained by a natural process. Finally, the results obtained from the downsampled UAS DEMs predominantly show high rates (16-18 m) that are interrupted by areas of no movement or very slow movement. My impression would be that these results are least reliable, because a) they show a completely different picture as (a) and (b), while being computed with the same data (just a different resolution), b) the displacement values are nearly the same for two very different time intervals (e = 376 days, f = 42 days), c) they are not matching the values obtained from manually tracking boulders (again, based on the same data), and d) I am unsure if such a pattern can be produced by a natural process.

Having outlined my reservations regarding the image correlation results, let me suggest a couple of strategies to improve the results:

- Use a hillshade not a DEM for tracking (not clear if this was done)
  Originally we used UAS orthoimages. Please see the OSM Fig. 8 for calculations using DSM and OSM Fig. 9 using hillshades.
- Resample the DEM to a slightly coarser resolution (0.5 m?)
  We have tried a 0.5 m resolution for the UAS orthophotos with different parameter settings showing overall better matching with still some decorrelation. However, with this input resolution and the best suited parameter settings of 128 x 32 the extent is already decreased in its size to a smaller displacement area.
- Try a different software for image correlation, there are many and all have their advantages and disadvantages
  This was done with DIC–FFT and IMCOOR (please see OSM Fig. 10 for results of DIC–FFT).
• Have a detailed look into the correlation coefficients and the bearings of the displacement vectors and exclude spurious results. Please see OSM Fig. 3 (b) and (h) with displacement vectors and signal–to–noise maps in the OSM Fig. 3 (c), (i) and (d), (j), Fig. 5 (f) and (i) and a cross profile cutting the DIC total displacement for both intervals I and II, Fig. 6.

Yes, indeed a mismatch of the initial and final search windows, i.e. a decorrelation, is visible for many areas but especially obvious in the western part of our DIC results. The current literature states among others that there is an upper limit regarding velocities of ground motion (Delacourt et al. 2007; Travelletti et al. 2012). In this area very high motion clusters of this complex landslide exhibit debris slide characteristics. We observed that acceleration of the landslide body takes place here. In contrast, in the eastern part of our DIC results, there are correlated areas and smooth motion patterns indicating that matching took place and the method was successful with the applied parameter settings.

Additionally, in our case, the terrain surface is altered rapidly; big blocks with edge lengths of up to 10 m rotate and cause significant surface changes, which could be a further reason for decorrelation (see OSM Fig. 9 for results of DIC–FFT) (Lewis 2001; Stumpf et al. 2018). The geomorphic causes for the observed acceleration are unknown but could be related to permafrost degradation and increased infiltration of rain- and meltwater.

In the OSM we support the result from DIC with the corresponding displacement vectors (OSM Fig. 3 (b) and (h)). With regard to the 3 m downsampled UAS orthophotos we are aware that these results are less trustworthy in terms of delineated velocities. Here, our purpose was to compare two different sensors in order to see how accurate PlanetScope data are for high alpine displacement calculations. Please see here our comment further below.

• L22/23: While this is certainly true, the authors should elaborate in the introduction that events instantaneously triggered by earthquakes or heavy precipitation are beyond what their proposed framework can deliver an early warning for. The necessity of gathering and evaluating data prior to issuing a warning limits the analysis to mass movements that indeed show a pre-failure acceleration on the order of days. Thank you for highlighting this. We totally agree that this has to be mentioned in the beginning to complement our explanations in the discussion, L561/562.

L31: This definition of an early warning system (EWS) contains a time component but includes no exact time scale reference. ‘Early’ suggests that events are detected before harm or damage occurs and thus stands in contrast to events which are only detected once they have begun (e.g. snow avalanches). Thus, it is necessary to know sensor capabilities and limitations for pre–event mass movement observations (Desrues et al., 2019). The success of a warning requires that information is provided with enough lead time for decisions on reactions and counter measures (Grasso, 2014). The success of an EWS therefore requires measurable pre-failure motion (or slow transport velocities) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014). In this regard, knowledge on sensor capabilities and limitations is essential, as it determines which rates and magnitudes of pre-failure motion can potentially be identified (Desrues et al., 2019). Our proposed framework refers to mass movements with significant pre-failure motion operating over a sufficient time periods and thus excludes instantaneous events triggered by processes such as heavy rainfalls or earthquakes.
L25/26: Is this really just attributable to the warming of the climate?
To the best of our understanding and following Gariano and Guzzetti in their review (2016) the global climate warming directly and indirectly impacts natural and human induced factors which can again directly or indirectly condition landslide activity, abundance and frequency of events. Other reasons for landslide triggers are included in L22/23, earthquakes, rainfall events and human interaction.

L47/50: I would think that also the rate of landslide movement defines whether or not it can be detected by optical imagery.
Thank you for pointing out that detection is not restricted to sensor characteristics. This is very important to say, of course.

L79/80: This is the maximum revisit time at the equator, right? For the study area shown here, revisit time should be shorter.
Yes, thank you for mentioning this. We will differentiate here between revisit frequency and repeat frequency, with the latter of importance for coherence.

L121: What do you mean by “natural developments” and how are these conditioned or different from natural processes?
Thank you for this comment. We are sorry that this was not specific. We meant the development of natural processes.

Figure 1: While I like the idea behind this conceptual figure, I would recommend the authors add a time axis and limit the area of “significant acceleration” to a vertical line that coincides with $t = 0$. In the present form, the conceptual figure contradicts statements in the text, such as “The forecasting window is started […] following significant acceleration […]” (L126), or “Simultaneously with the forecasting window, time to warning ($t_{\text{warning}}$) starts (grey outline)” (L128/129).
Thank you, you are right. We changed it to our best understanding of your feedback.
- L133/134: This also does not match what Fig. 1 is showing.
  “The lead time is the difference between the forecasting window and the time to warning.”

  We want to express that $t_{lead}$ is the rest/remainder of the subtraction as follows
  $t_{forecasting\ window} - t_{warning} = t_{lead}$

  Please let us clarify this as it seems to be some sort of misunderstanding here. As a suggestion, this could be replaced with L133/134, if you prefer “Lead time is the forecasting window minus the time to warning.”

- L139: This also does not match what Fig. 1 is showing. In Fig. 1, $t_{lead} < t_{react}$.
  “An imperative for an effective EWS, the required time to take appropriate mitigation and response measures has to be within the lead time interval ($t_{lead}$) (Pecoraro et al., 2019) with $t_{lead} \geq t_{react}$”.

  Please let us try to clarify this: in best case, the lead time is longer than the time needed to take responsive measures and react to the impending event ($t_{react}$), this is indicated by the shorter solid grey arrow. However, if the reaction time is as long as the lead time, see dashed extension of the grey arrow, then it is a coincident ending of both, $t_{react}$ and $t_{lead}$ prior the release and impact.

- L215/127: In theory yes, but as you show later (Tab. 2), the effective revisit time of optical imagery might in fact be very similar.
  Unfortunately, we do not understand what you are referring to in L127.

  L215: Sentinel–1 does have a revisit time of about every second day over Europe. However, the repeat frequency for coherence to generate interferograms is every six days. This is the shortest possible temporal baseline.
In terms of optical satellite images, yes, this is what the author team finally wants to lead to. PlanetLabs claim to have daily acquisitions and thus can provide daily imagery supply. But upon a closer look the practitioner knows the reality is different. This has to be kept in mind if this kind of data is employed for the purpose of a reliable monitoring and process observation. For this reason, Table 1 and Table 2 have different and contradicting statements, in this case for PlanetScope.

You are right in some way: free satellite images by Sentinel–2 are, at five days, very close to the six days for interferograms by Sentinel–1, given that both sensors are suitable for the given characteristics by the acquisition target (motion velocity, exposition). Apart from open data providers, there are many others providing even sub–daily acquisitions such as WorldView 3/4.

- L242/248: It might be worth mentioning here that on average, only 11% of the images were usable, significantly reducing the theoretical revisit time, as you also outline in the discussion. 
  Thank you, indeed this is worth to be mentioned and we changed accordingly.
  In this seven–month period, 43 images (20.1 %) had data voids or did not cover the AoI, thus the overall usability is limited to about 11 %.

- L267/269: Please elaborate how you filtered for “errors of location, shift and spectral colour problems” (are the latter spectral differences between images?). 
  We used QGIS software to manually select the satellite images with the reference UAS images at the base and the visual “show/hide” of the satellite slave images on top. Similarly, the application Map Swipe Tool plugin was employed by dragging the slider across the images. 
  Spectral colour problems are shifts in the individual r, g and b bands within one single image:

![Image]

The other shifts which might occur cannot be corrected for. The first time these can be detected is in a GIS software with the visual check previously described:
Thereafter, a second selection (visually with the Map Swipe Tool plugin) from the downloaded images was filtered for errors of location, shift and spectral colour problems which were previously not clearly discernible in the online data hub.

- L281/285: Please specify the accuracy of dGPS coordinates as measured for the GCPs and also include an accuracy information for the DEMs and their derivatives that were produced from UAS surveys. The accuracy of dGPS coordinates, which were employed for the processing of UAS data and DEM/orthophoto generation, range between 5 cm horizontally and 10 cm vertically. All UAS model calculations are based on the same dGPS measurements. The RMS errors from UAS image processing in Pix4Dmapper range between 4 and 8 cm. If generation reports are necessary, they can be provided on request later (due to current office access difficulties). These were repeatedly (1000 measurements/position) registered with the TRIMBLE R5 dGPS and corrected via the baseline data of the Austrian Positioning Service (APOS) provided by the BEV (Bundesamt für Eich- und Vermessungswesen). Horizontal root–mean–squared errors (RMSE) range from 0.05 m to 0.10 m for vertical RMSE. These GCPs were employed for georeferencing and further rectification of all UAS surveys.

- L285/286: Please elaborate how image co-registration was achieved and state here the residual mismatch between co-registered images. DIC methods for estimating terrain movements require accurate geo-referencing of consecutive satellite images avoiding falsely detected systematic drifts. Although the investigated satellite sensors are equipped with high–quality geo-localization sensors, subtle deviations in the absolute geo-referencing rates are expected for different acquisition times. Therefore, a fine–registration between satellite image patches in the AoI was conducted based on a Matlab script (by Tobias Koch) applying a state–of–the–art image registration technique (Lowe 2004). Since radiometric differences between the different acquisition times and image distortions (e.g. clouds) could remain in the images, feature–based registration methods are preferable over correlation–based registration methods due to their ability to match local feature points instead of entire image areas.
To ensure that actual terrain movements in the AoI do not cause undesired shifts in the registration, the AoI was excluded from the feature point detection step. The remaining feature points were used for estimating a geometric similarity transformation between the reference and all target images including a statistical outlier removal (RANSAC). This transformation was finally used to accurately register a target image towards the reference image.

Regarding the registration quality in the test site, a satisfying amount of feature matches of at least 500 after outlier removal could be found for all reference (master) and target (slave) image pairs and for all investigated sensors. The mean distance of transformed inlier feature points of the target image to their corresponding feature matches in the reference image ranged between 0.6 and 0.8 pixels, confirming the high registration accuracy (see OSM Fig. 14).

- L288/289: Usually matching between consecutive images is not achieved by matching “common pixels”, but by maximizing the correlation between pixel-value distributions of patches of pixels (i.e. your windows of different sizes in Tab. 6).

Yes, you are correct it estimates first the pixelwise displacement between two patches based on correlation peaks and second, the final correlation is performed to retrieve the subpixel displacement.

We added this information and reordered the processing steps according to the COSI–Corr manual (Ayoub et al. 2009).

There are two correlators; in the frequency domain based on FFT algorithm (Fast Fourier Transformation) and a statistical one. Applying the more accurate frequential correlator engine, recommended for optical images, different parameter combinations of window sizes, direction step sizes and robustness iterations were tested.

Parameter settings include the initial window size for the estimation of the pixelwise displacement between the images and the final window size for subpixel displacement computation in \( x, y \); a direction step in \( x, y \) between the sliding windows; and several robustness iterations (Fehler! Verweisquelle konnte nicht gefunden werden.).

[...] The results of each correlation computation returns a signal–to–noise ratio map (SNR) and displacement fields in east–west and north–south directions. These results were exported from ENVI classic as GTiff, and the total displacement was then calculated with QGIS.

- L304/305: What is the uncertainty of these east-west and north-south displacement estimates? Did you check whether the bearing of the displacement matches the general slope of the Sattelkar?

In the OSM we are providing the results of the correlation computations for our published results (east–west and north–south displacement fields as well as signal–to–noise maps). The results are consistent. We further provide total displacement results of other parameter combinations.

Yes, we checked the overall orientation of the correlation based on computed directional vectors (with SAGA GIS software). We provide these vectors in the OSM, too (OSM Fig. 3 (b) and (hi)).
L307/308 and L440/442: This seems a bit arbitrary. How did you determine a cutoff–value of 4m displacement? How did you distinguish outliers from non-outliers? What is the confidence of your estimates?

We determined the cutoff–value employing several criteria. First based on field experience we know the landslide extent and displayed the results in combination with the demarcation displayed as ‘Active area’ in Figs. 2, 3, 5 and 7. Then we checked the value distribution in the histograms for both the calculated total displacement as well as the signal–to–noise maps. These maps were further used to visually compare the total displacement results. This allowed us to identify outliers and unlikely displacement. Based on the histograms and the acquired experience for the results, the thresholds were tested and set for transparency and to display values. Please see the OSM (Fig. 13).

L308/309: This contradicts the descriptions of Fig. 5a, where you point out that “ambiguous, small-scale patterns with highly variable displacement rates” (L332/333) dominate the western part of the mass movement.

Here we would like to differentiate between inconsistencies which we understand as artefacts and noise due to snow, vegetation, clouds, cloud shadows and terrain shadows. De–correlation with its salt–and–pepper appearance due to velocities exceeding the correlation capability of DIC have a different origin and reason.

However in the results, section 5, we described the appearance of these ambiguous signals, while in the discussion section they are explained.

L311/312: I am not convinced that manually tracking boulders in the same images that were used for image correlation can verify the results of this correlation. You can use these data to check if manual and automated tracking give consistent results. Comparing manually tracked boulders from UAS imagery could however be used to compare against the displacement estimates from satellite imagery.

We are certain that the direct measurements of travelling distances from blocks of 10 m size for consecutive orthoimages, which were also employed for the DIC method, are a valid method to underpin the total displacement results by the DIC. Comparing these tracks with satellite imagery might be useful keeping in mind that the difference between UAS orthoimages of 0.16 m and PlanetScope satellite images of 3 m spatial resolutions is substantial and sensor type, image processing etc. can introduce further inaccuracies.

L320: As you present total displacement for different time intervals here, not rates in distance per unit time, I would suggest changing the title here. Same is true for L326, L346 and L361.

Yes, thank you for pointing this out. We changed the section title (see below) and in the text accordingly (L326, L346, L350, L354, L357 and L361).

Section Title: 5.1. Total displacements

L335/336 and L366: Did you check the direction of displacement for the areas of smallscale patterns of ambiguous signals? I would suspect that these are very heterogenous here as well. It would also be worth looking into the quality information (correlation coefficients) for these regions.
Yes, this is a good point. Indeed, we checked the direction based on displacement vectors as well as signal–to–noise maps. They both give the same indication of heterogeneous and ambiguous signals with no correlation for exactly the same areas with ambiguous signals in the total displacement calculation. Please see our OSM (OSM Fig. 3).

- L397: For a comparison (and also for a better readability) you could convert your total displacement to average rates of m yr⁻¹ or cm d⁻¹.
  Yes, converting them into averaged rates is a good suggestion for the discussion section, see below. If you recommend this conversion for the results section 5 too, then the section title should be kept “Displacement rates” as before (see your previous comment for L320). For the (old) L417, 418 and 420, the values were added with yearly rates in brackets: trajectories up to 4 m (34.8 m yr⁻¹) (d); a 16 m (139 m yr⁻¹) trajectory (a); approximately 10 m (86.9 m yr⁻¹).

- L398/399 and L402/404: Given the large differences in total displacement between sensors and resolutions used for image cross-correlation, I do not think that you can make this claim. Please use an appropriate measure to quantify the agreement between manual boulder tracking and the three different approaches used for digital image correlation. These lines refer to the results of the total displacement derived from UAS orthophotos. Regarding L398/399 the parameters were tested and selected independent of others’ recommendations, but we arrived at the similar conclusions. With regard to L402/404 we believe that the travel distance measurements of field mapped boulders based on the same data (UAS orthophotos) are comparable to DIC derived total displacements.

- L419/422: This might be the case, though you tested larger patch sizes (Tab. 6) that should have given you consistent results for this region then.
  In the OSM we provide results of our parameter tests for larger final window sizes (see OSM Fig. 5 and 7).

- L433/434: This should be backed by a statistical measure. From a close look to Fig. 5, I rather get the impression that the only patches you can make this statement for is location a in Fig. 5 (b) and (d) and location c in Fig. 5 (a) and (c), but to a lesser extent.
  Thank you for pointing this out. In our opinion the first time interval with slightly more than one year of accumulated displacement, the frontal area and core body of the landslide are reflected in both DIC results of UAS and PlanetScope (locations (a) and (c), as well as (d) and slightly (e) and (f) in Fig. 5 a) for UAS and c) for PlanetScope I). In contrast to the second interval of 42 days, it seems that there is not enough accumulated displacement to be captured by PlanetScope DIC, as the middle to rear landslide body are only reflected in the UAS DIC result (locations (b)–(d) Fig. 5 c) and remain free of signal for these locations in Fig. 5 d) for PlanetScope.

- L445/447: The size of the snow patches does not play an important role. The presence of snow in one image hampers correlation between images and leads to false patchmatching results.
  Yes, we absolutely agree and this is also described by Leprince et al. (2007; 2008), noting that variations, thus the difference in snow cover, limit the technology. In addition, they say that in images with high gains, the areas of snow coverage are saturated too, and as a result, do not allow for any correlation (Scherler et al. 2008).
Regarding the displacement for (j) as identified in both sensor combinations (see Fig. 5), there is a patch of snow (1–2 m height, length ~ 25 m, see OSM Fig. 10) in the UAS and PlanetScope images on 24.7.2019 while for the images on 13.7.2018/19.7.2018 (UAS/PlanetScope) and 4.9.2019 (UAS and PlanetScope) there is no snow (see OSM Fig. 2 and 11). Thus, in this case, the existence of snow in one image but not in the other explains this false correlation and indication of displacement.

Minor snow fields as visible in the images from 24.07.2019 for both, UAS and PlanetScope, likely explain the big cluster of incorrect displacement southeast of the lobe (j); nonetheless, in the satellite image they are smaller than the resulting DIC displacement.

- L457/462: To be frank, I do not see much similarity between Fig. 5 (c) and (e) nor (d) and (f). I would be very cautious in interpreting these results as is. This is especially true for the resampled UAS results.

Thank you for pointing this out. Yes, we agree in some part. Our purpose was to compare our high accuracy UAS orthophotos to PlanetScope satellite images, in order to estimate the goodness of fit and limitations of the latter.

We are aware that this downsampling factor is large, and therefore the resulting displacement rates and inherent velocities have to be viewed with reservations.

However, in terms of noise outside our defined active landslide area and the overall detection to the landslide boundary as delineated based on the 0.16 m UAS data: for the first, the noise is low to moderate, and there is generally a good fit for the 3 m downsampled UAS data similar to DIC results of UAS at 0.16 m, respectively. In contrast, DIC results of PlanetScope neither show likewise noise–free areas outside the active landslide regions nor do they reach the same extent total displacement extent as the downsampled UAS data.

- L463/464: As the GCPs for referencing the UAS data are probably located close to the landslide, it is not surprising, but neither disturbing, that false displacement clusters appear outside the area of interest.

Please see our map of GCP distribution as well as images thereof in the OSM (OSM Fig. 1). Some GCPs are close to the landslide area, but installed on stable bedrock and to best of our knowledge, they are not moving and thus provide continuous usability and comparability.

False displacement is indicated for a cluster outside of the boundary to the image border in the east for UAS interval I (Fig. 5e) and in the north western area (h, i) for interval II (Fig. 5f) contributing to changes in shading and illumination.

- L468/470: Again, I would not trust the displacement estimates of the resampled UAS data. While it is true that your manual boulder tracking identified 2 boulders with displacement of 10 or more meters, the remaining 34 boulders show something different.

Yes, you are right that not all of the 34 boulders are exactly reflecting the DIC total displacement result. However there are more than two which are in the same range of displacement, and others are very close to it, keeping in mind that there are some uncertainties and limitations when it comes to the threshold of identification of small ground motions in the DIC method. Please see here the section 6.1, discussion. We are happy to revise this further.
L471/476: While it might be true that the results obtained from image correlation of resampled 3m UAS data are better (internally) correlated and show a more homogeneous deformation pattern, this does not mean that the result is correct. As I outlined above, I have serious doubts regarding the interpretability of this data, as there is no agreement with the manually tracked boulder velocities (except 2 boulders). Also, from a geomorphic perspective, I am not sure how you would explain a velocity pattern where high velocities dominate throughout the entire landslide, but are speckled with lower to zero movement within (Fig. 5 e and f).

We agree with the 3 m resolution to some extent. Please see comment above for L457/462 the comparability of manual block tracking to UAS DIC result. The ‘speckled’ pattern, is due to decorrelation resulting from velocities too high to be captured with the DIC method; this combined with an observation period of 42 days delay (Delacourt et al. 2007; Travelletti et al. 2012) may be exceeding the accumulated displacement to be captured by the method, which could contribute to this pattern and explain the resulting limitation to some extent. In addition, we know that the surface changes significantly in the frontal part and these strong alterations also limit the DIC method (Lewis 2001; Travelletti et al. 2012). For more please see section 6.1, discussion, too. If this is not clear enough in the discussion, we would be happy to further revise this.

L485/488: Did you evaluate the proportion of false-positive displacements to true-positive displacements and if so, how did you do this and can you please include this data? Based on the image correlation results shown here, you can make this statement, but I would be cautious to make a general claim on the usability of the data.

We approached our results by testing of different parameter settings and combinations based on visual comparison as is common in the field (Bontemps et al. 2018). The PlanetScope DIC results presented here are the most suitable master–slave image combinations. We could provide the other intervals of DIC results which are not meaningful for comparison if wished.

L552/554 / Table 7 / Figure 9: I do like the idea behind this, where the authors show that their proposed workflow would enable a timely warning in the case of historic landslides. However, in the case of Vajont, I think you should include a critical factor. While it is theoretically true that a “forecasting window” would allow for your workflow to be completed well before the failure, the slow deformation of Vajont (35 mm d-1) in the 30 days will be well below the level of detection of your image correlation analysis, if you collect an image directly after the onset of “significant acceleration”. In order to be detectable, movement must have accumulated a critical distance before data collection of your workflow can set in (30 days = 1.05 m total displacement) – a factor that in my view would be important to include here.

Thank you for mentioning this, you are absolutely right. We added the following sentence below to emphasise this critical detection capability limit of the DIC method.

We assume that approximately 30 days before failure Vajont would have displayed a signal exceeding the noise at modern standards and would have become predictable.

For Vajont, the 1/velocity plot by Petley and Petley (2006) (based on data from Semenza and Ghirotti (2000)) shows an increase in movement at about day 60 along with a transition from a linear to an asymptotic trend at
approximately day 30, defined as a transition from ductile to brittle. Therefore, we assumed 30 days of forecasting window for twarning and tlead until the impact of the hazardous event on 09.10.1963. However, it has to be kept in mind that velocities of about 35 mm d\(^{-1}\) are still low and at the minimum of the displacement recognition capability for the digital image correlation method.

**Technical corrections:**

- L1: Landslide
  
  Here we are referring to landslides in general, not to a specific landslide.

- L103/105: Check grammar
  
  We did not add a comma as the text is in BE; in AE, however, a comma could be added (In this investigation,…). We added quotation marks to improve readability.

- L185: Is this really the source the authors need to cite for the location map?
  
  Thank you, we modified in response to comment by RC1 (J. Blöthe) by changing Vienna to Wien. Otherwise this is according to the publishing company and the copyright statement from the online map.

  **Figure 1** (a) Overview map Austria (Österreichischer Bundesverlag Schulbuch GmbH & Co. KG and Freytag–Berndt & Artaria KG, Wien).

- L229: beginning of April
  
  Thank you, we inserted missing word.

- L257: UgCS-Software?
  
  Further information on the flightplanning Software UgCS can be found here: https://www.ugcs.com/photogrammetry-tool-for-land-surveying

- L299/300: I guess this is only relevant if you explicitly mention the image–processing times.
  
  Thank you, however we think this is relevant as the duration of image processing and DIC calculation are an important part of our temporal concept in the results section 5.3. and discussion section 6.3.

- L398: can be compared
  
  We think that the repetition of ‘compared’ is not necessary; it follows from the logic of the sentence.

- L409/410: resulting from significant morphological changes?
  
  Thank you for pointing this out. After a detailed verification of volumetric calculations, we can confirm changes of about 1 m. Please see our calculations and visualisations in the OSM.

  In Fig. 5a, the large southern patch (g) shows clear displacement values for the rear part and decorrelation for the front region resulting from morphological changes within the image pair of interval I.

- L443: bracket missing?
  
  Yes you are right, thank you.
- L 460: check figure reference
  Thank you for pointing on this auto-correction mistake.
References


Delacourt, Christophe; Allemand, Pascal; Berthier, Etienne; Raucoules, Daniel; Casson, Bérangère; Grandjean, Philippe et al. (2007): Remote-sensing techniques for analysing landslide kinematics: a review. In: Bulletin de la Societe Geologique de France 178 (2), S. 89–100. DOI: 10.2113/gssgfbull.178.2.89.


Leprince, Sébastien; Berthier, Etienne; Ayoub, Francois; Delacourt, Christophe; Avouac, Jean-Philippe (2008): Monitoring Earth Surface Dynamics With Optical Imagery. In: Eos 89. DOI: 10.1029/2008EO010001.


Letter of response to comment on nhess-2021-18

Dear Sigrid Roessner,

We thank you for your valuable comments on our manuscript and appreciate the time and the efforts you have invested. Your feedback has helped us to see and clarify ambiguous areas to further improve our work.

Based on your suggestions we have restructured the entire manuscript, especially introduction, study site description, discussion and conclusion. In addition, we have specified many conceptual and methodological concerns according to your more specific remarks. We have also rephrased several ambiguous paragraphs.

Please find below the following colour coding for the review and your comments in black; our responses to the review are in blue and the changes made to the manuscript are in green (following RC2), orange (following RC1) and in blue by the authors. Reference to line numbers are based on the original preprint.

General comments

The paper represents an interesting contribution to process oriented remote sensing based monitoring of complex landslides with the aim of making a conceptual contribution to early warning. The paper is well written in language and structure and the figures are of good quality. Despite the overall good scientific relevance and presentation quality, in the current form the paper lacks a coherent scientific goal justifying the use of approach. This problem already becomes apparent in L40 where the authors state that the study presents a new concept to systematically evaluate remote sensing techniques to optimize lead time for landslide early warning’. Although the presented work is very interesting, it does not fit the stated goal for the following reasons:

- Concept of lead time and need for best possible reduction is not new.
  While we agree that the concept itself may not be knew, we find that using multispectral remote sensing products to assess and increase lead time to ensure the timely prediction of landslide early warning systems represents an important research gap that so far has rarely been addressed. We evaluate the capabilities of remote sensing to identify hot-spots and detect process behaviour changes based on the local conditions. Thus, the landslide process is the precondition. We want to estimate, based on the assumption that the particular sensor is able to deliver the necessary information, the time demand of each sensor for time to warning. We have now replaced the phrase optimising lead time with a more precise description of what we have done. Please see revision of the conclusion further below.

L10–11: We introduce a novel conceptual approach for comprehensive to structure and quantitatively assess lead time assessment and optimisation for LEWS.

[…]

L39–41: This study presents a new concept to systematically evaluate remote sensing techniques to optimize estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for warning-early warnings.

[…]

L34: Lead time as defined in the context of LEWS is the interval between the issue of a warning (i.e. dissemination) and the forecasted landslide onset (Pecoraro et al. 2019) and thus crucially depends on time requirements in phases
The success of an EWS therefore requires measurable pre–failure motion (or slow slope displacement) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014; Hungr et al., 2014).

- Remote sensing techniques themselves are not the bottleneck for shortening the lead time.
  The goal of our concept is not to refine remote sensing as a technique itself but to provide a tool for choosing the appropriate sensors based on time required for the time to warning phase. We thereby increase lead time.
  We do not agree with your objection to the word “bottleneck” especially given your comment below which says “In remote sensing based approaches lead time mostly depends on the available imaging constellation and data distribution to the end user.”

L39–61: This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for providing early warnings. This approach reduces the to a small number of suitable remote sensing products with high temporal and spatial resolution. With these constraints, we investigated the application of data from satellites and unmanned aerial systems (UAS) to allow the assessment of the data, after a spaceborne area-wide but low–resolution acquisition, into a downscaled detailed image recording. In so doing, we analysed the capability of these different passive remote sensing systems focusing on spatiotemporal capabilities for ground motion detection and landslide evolution to provide early warnings.

[...] 

L94–102: In recent years, data provision for users has increased and today data hubs provide easy accessibility to rapid, pre–processed imagery. Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Nonetheless, technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non–existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Timely information extraction and interpretation are critical for landslide early warnings yet few studies have so far explicitly focused on time criticality and the influence of the net temporal resolution of remote sensing data.

- In remote sensing based approaches lead time mostly depends on the available imaging constellation and data distribution to the end user and in case of optical data on the atmospheric conditions (clouds). Both factors are only to a very limited extent in control of the authors - only in case of the UAV data acquisitions.
  Thank you for your comment. We agree that the limitation of meteorological conditions including effects such as cloud shadow and snow are important constraints as we described in L45–55 and L158. We took this into consideration when estimating
the number of available PlanetScope images (Sect. 4.2.) and discussed atmospheric affected images with regard to displacement derivation results in L477–481. You are right that for UAS campaigns, most of the control is on the user side and only to a very limited part for other satellites. Today, some data providers promise new images daily, sometime even more frequently (e.g. PlanetScope). But this is the point we want to highlight with our study. In a real world situation, we wish to determine which satellites can provide useful timely information in terms of an effective repetition rate and real availability in the data hub (provider). In addition, the natural conditions such as atmospheric and site specific constraints can reduce the net image number. For this reason, we assess the capabilities of optical remote sensors in a spatiotemporal context for given circumstances to detect hot spots and identify possible changes in slope processes.

L52–55: Previously, high spatial resolution satellite data was obtained at the expense of a reduction in the revisit rates (Aubrecht et al., 2017). Consequently, the return period between two images increased, limiting ground displacement assessment and the range of observable motion rates. The number of useful images was further reduced due to natural factors such as snow cover, cloud cover and cloud shadows. […]

L86–91: In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019). However, as a flexible, cost–effective alternative to spaceborne optical data are airborne optical images taken by UASs (unmanned aerial systems). Freely selectable flight routes and acquisition dates prevent enable avoiding shadows from clouds and topographic obstacles, and as well as allow avoiding unfavourable weather conditions and summer time snow cover, all of which frequently impair satellite images (Giordan et al., 2018; Lucieer et al., 2014).

L96–102: […] technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non–existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Timely information extraction and interpretation are critical for landslide early warnings yet few studies have so far explicitly focused on time criticality and the influence of the net temporal resolution of remote sensing data.

- The used data sources (planet and UAV) do not allow optimization of lead time in the context of early warning because of the scarcity of their availability which is reflected in the small number of only three multitemporal data takes between July and September analyzed in this study (Table 3)

Thank you. With regard to this comment we assume this needs further clarification. First, we have changed the entire phrase on “optimising lead time” to be more precise in the description of our approach (see previous comment). Regarding the data takes, yes, we do have three UAS acquisitions but over the course of more than one year (7/2018–9/2019). For the purpose of this comparison we selected PlanetScope data at a similar time to UAS acquisitions, whereby one Planet image (02.07.2018, see Table 5) showed low quality results why the time interval was excluded (see caption Fig. 4). In both UAS and PlanetScope DIC results we can see the general distinctive hot–spot
identification as well as changes in motion behaviour indicating an acceleration for the time intervals I and II. Second, we can obtain a higher frequency of UAS acquisitions if necessary. We have revised our conclusion to be more concise in our work with regard to both, the term optimisation as well as the total number of data takes.

This paper presents an innovative concept to compare the lead time for landslide early warnings, utilising of two optical remote sensing systems. We tested this temporal concept by applying UAS and PlanetScope images of temporal proximity as these are currently the sensors with the best spatiotemporal resolution.

Our findings derived from DIC for this steep high–alpine case study show that high resolution UAS data (0.16 m) can be employed to identify and demarcate the main landslide process and reveal its heterogeneous motion behaviour as confirmed by single block tracking. Thus, validated total displacement ranges from 1–4 m and up to 14 m for 42 days. PlanetScope Ortho Scenes (3 m) can detect the displacement of the landslide central core, however, cannot accurately resolve represent its extent and internal behaviour. The signal–to–noise ratio, including multiple false–positive displacements, complicates the detection of hotspots at least in this very steep and heterogeneous alpine terrain.

Coarse temporal data resolution, such as in the case study investigated here, represents an important restriction to the use of optical remote sensing data for landslide early warning applications. Acceleration (and the resulting failure) over short periods of time will likely go unnoticed due to large data acquisition intervals. However, for prolonged acceleration periods, such as observed at the Sattelkar slide and many other relevant hazard sites, the chosen data sources have been demonstrated to represent a formidable early warning approach capable of contributing to an improved risk analysis and evaluation in steep high–alpine regions.

For continuous monitoring and early warning, the warning time window could be shortened by on–site drone ports with autonomous acquisition flights and automatic processing. Our systematic evaluation of the sensor potency capability can be applied and transferred to other optical remote sensing sensors, and the same is true for our conceptual approach optimising which extending the lead time. Future studies should focus on the applicability of complementary optical data to confirm the detection of landslide displacement and adjust UAS output resolution as this significantly increases the validity of DIC internal ground motion behaviour.

- The missing sound conceptual approach is also reflected in the introduction in form of a lengthy summary of in principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 1 is very general and not specific to landslide and does not qualify as a novelty in the current form.

1. Introduction

We revised the abstract and the introduction , to be more precise with regard to our goal and implementation. In so doing we more clearly defined our approach to lead time and early warning systems for landslides. Further we did our best to improve the line of arguments and to show the historic limitations of optical remote sensing for LEWS up to the recent developments when it comes to options such as high spatiotemporal products and their usage for monitoring, early warning and time-series displacement analyses.
2. The conceptual approach
We decided to keep this concept general, to employ it for other remote sensing techniques and maybe even other kind of instrumentation as well as different use cases of other time challenging issues. We revised and added some sentences to emphasise our approach/idea. Even after intense research we did not find good conceptual approaches challenging remote sensing in the direct context of landslide early warning systems. We therefore consider our approach novel. This concept forms the basis to employ this for the setup of ‘a real early warning system’.

L21–102: Landslides are a major natural hazard leading to human casualties and socio-economic impacts, mainly by causing infrastructure damage (Dikau et al., 1996; Hilker et al., 2009). They are often triggered by earthquakes, intense short-period or prolonged precipitation, and human activities (Hung et al., 2014; Froude and Petley, 2018). In a systematic review Gariano and Guzzetti (2016) report in a review study that 80% of the papers examined papers show causal relationships between landslides and climate change. The ongoing warming of the climate (IPCC, 2014) is likely to decrease slope stability and increase landslide activity (Huggel et al., 2012; Seneviratne et al., 2012), which indicates a vital need to improve the ability to detect, monitor and issue early warnings of landslides and thus to reduce and mitigate landslide risk.

Early warning, as defined by the UN International Strategy for Disaster Reduction (UNISDR), refers to a set of capacities for the timely and effective provision of warning information through institutions, such that individuals, communities and organisations exposed to a hazard are able to take action with sufficient time to reduce or avoid risk and prepare an effective response (UNISDR, 2009). According to UNISDR (2006), an effective early warning system consists of four elements: (1) risk knowledge, the systematic data collection and risk assessment; (2) the monitoring and warning service; (3) the dissemination and communication of risk as well as early warnings; and (4) the response capabilities on local and national levels. Incompleteness or failure of one element can lead to a breakdown of the entire system (ibid.). Lead time as defined in the context of LEWS is the interval between the issue of a warning (i.e. dissemination) and the forecasted landslide onset (Pecoraro et al. 2019) and thus crucially depends on time requirements in phases (1)–(3). The success of an EWS therefore requires measurable pre-failure motion (or slow slope displacement) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014; Hungr et al., 2014).

While remote sensing has been established for early warnings, remote sensing is not yet used for real early warnings of the onset of landslides in steep-alpine terrain (with a few exceptions), where geotechnical instruments are still preferred. Exceptions include terrestrial InSAR (Pesci et al., 2011; Walter et al. 2020) and terrestrial laser scanning with high repetition rates. However, repeated UAS (unmanned aerial systems) and optical satellite images (PlanetScope) with high repetition rates have so far not been applied for landslide early warning in steep-alpine catchments. In this regard, knowledge of sensor capabilities and limitations is essential, as it determines which rates and magnitudes of pre-failure motion can potentially be identified (Desrues et al., 2019). Our proposed framework refers to mass movements in steep-alpine catchments with significant pre-failure motion operating over a sufficient time periods and thus excludes instantaneous events triggered by processes such as heavy rainfall or earthquakes.

This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for providing early warnings. This approach reduces the to a small number the of suitable remote sensing products...
to a small with high temporal and spatial resolution. With these constraints, we investigated the application of data from satellites and unmanned aerial systems (UAS) to allow the assessment of the data, after a spaceborne area–wide but low–resolution acquisition, into a downscaled detailed image recording. In so doing, we analysed the capability of these different passive remote sensing systems focusing on spatiotemporal capabilities for ground motion detection and landslide evolution to provide early warnings.

Until Recently, the spatial and temporal resolution of optical satellite imagery has significantly improved requirements for accurate early warning purposes have not been met by optical satellite imagery (Scaiionti et al., 2014) and has allowed substantial advances in the definition of displacement rates and acceleration thresholds to approach requirements for early warning purposes. This is essential since spatial and temporal resolution determine whether landslide monitoring is possible with the detection allows defining of displacement rates and the approximation approximate acceleration thresholds, both of which are lacking if information is based solely on post–event studies (Reid et al., 2008; Calvello, 2017). Landslide monitoring offers the potential to significantly advance landslide early warning systems (LEWS) (Chae et al., 2017; Crosta et al., 2017). Previously, high spatial resolution satellite data was obtained at the expense of a reduction in the revisit rates (Aubrecht et al., 2017). Consequently, the return period between two images increased, limiting ground displacement assessment and the range of observable motion rates. The number of useful images was further reduced due to natural factors such as snow cover, cloud cover and cloud shadows. High–resolution remote sensing data was long restricted due to high costs and data volume (Goodchild, 2011; Westoby et al., 2012). Today commercial very high resolution (VHR) optical satellites exist, but tasked acquisitions make them inflexible and very cost intensive, thus limiting research (Butler, 2014; Lucieer et al., 2014). There is a vast spectrum of available remote sensing data with high spatiotemporal resolution (Table 1). Complementary use of different remote sensing sources can significantly improve landslide assessment as demonstrated by Stumpf et al. (2018) and Bontemps et al. (2018), who draw on archive data and utilise different sensor combinations to analyse the evolution of ground motion.

Table 1 Overview of different optical multispectral remote sensors with their corresponding resolution [m] and revisit rate [days]. The sensors are categorised into commercial and free data policy. 1 free quota via Planet Labs Education and Research Program, 2PlanetScope Ortho Scene Product, Level 3B/Ortho Tile Product, Level 3A (Planet Labs, 2020b), 3reached end of life, 3/2020, archive data usable, 45 m Ortho Tile Level 3A (Planet Labs, 2020a), 50.5 m colour pansharpened, 6self–acquired. Source: (ESA, 2020).

<table>
<thead>
<tr>
<th>Sensor</th>
<th>Temporal resolution [d]</th>
<th>Spatial resolution [m]</th>
<th>Free/Commercial</th>
</tr>
</thead>
<tbody>
<tr>
<td>UAS</td>
<td>flexible</td>
<td>0.08</td>
<td>F</td>
</tr>
<tr>
<td>WorldView 2</td>
<td>1.1</td>
<td>1.84</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 3</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 4</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye 2</td>
<td>5</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>SkySat</td>
<td>1</td>
<td>1.5</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye–1</td>
<td>3</td>
<td>1.64</td>
<td>C</td>
</tr>
<tr>
<td>Pléiades 1A/B</td>
<td>1</td>
<td>2.0 (0.5)</td>
<td>C</td>
</tr>
<tr>
<td>PlanetScope</td>
<td>1</td>
<td>3.0/3.125</td>
<td>C/F</td>
</tr>
<tr>
<td>RapidEye 1</td>
<td>5.5</td>
<td>5</td>
<td>F</td>
</tr>
<tr>
<td>Sentinel–2 A/B</td>
<td>5</td>
<td>10</td>
<td>F</td>
</tr>
<tr>
<td>Landsat 8</td>
<td>16</td>
<td>30</td>
<td>F</td>
</tr>
</tbody>
</table>

The latest developments in earth observation programs include both the new Copernicus’ Sentinel fleet operated by the ESA, and a new generation of micro cube satellites, sent into orbit in large numbers by PlanetLabs Inc. These PlanetScope micro cube satellites, known as ‘Doves’/PlanetScope (from now on referred to as PlanetScope
satellites), and Sentinel–2 a/b offer very high revisit rates of 1–5 days and high spatial resolutions from 3–10 m, respectively (Table 1), for multispectral imagery (Drusch et al., 2012; Butler, 2014; Breger, 2017). This opens up unprecedented possibilities based on these high spatiotemporal resolutions to study a wide range of landslide velocities and natural hazards through remote sensing. Future Continuing data access is fostered by PlanetLabs and by Copernicus (via its open data policy) providing affordable or free data for research. This leads to unprecedented possibilities for studying natural hazards through remote sensing. Examples of landslide activity studies employing multi-temporal datasets of landslide activities based on this access to high spatiotemporal data include Lacroix et al. (2018), using Sentinel–2 scenes to detect motions of the ‘Harmalière’ landslide in France, and Mazzanti et al. (2020), who applied a large stack of PlanetScope images for the active Rattlesnake landslide, USA.

As forecasted, landslides tend to accelerate beyond the deformation rate observable with radar systems before failure, we concentrate on optical image analysis (Moretto et al., 2016). One advantage of optical imagery is its temporally dense data (Table 1) compared to open data radar systems with sensor visits repeat frequency more than every six days and revisit frequency between three days at the equator, about two days over Europe and less than one day at high latitudes (Sentinel–1, ESA). Optical data allows direct visual impressions from the multispectral representation of the acquisition target and the option to employ this data for further complementary and expert analyses. While active radar systems overcome constraints posed by clouds and do not require daylight, data voids can be significant due to layover or shadowing effects in steep mountainous areas (Mazzanti et al., 2012; Plank et al., 2015; Moretto et al., 2016). Moreover, north/south facing slopes are less suitable, thus limit the range of investigation (Darvishi et al., 2018). In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019).

However, as a flexible, cost-effective alternative to spaceborne optical data are airborne optical images taken by UASs (unmanned aerial system). Freely selectable flight routes and acquisition dates prevent enable avoiding shadows from clouds and topographic obstacles, and as well allow avoiding as unfavourable weather conditions and summer time snow cover, all of which frequently impair satellite images (Giordan et al., 2018; Lucieer et al., 2014). UAS–based surveys provide accurate very high resolution (few cm) orthoimages and digital elevation models (DEM) of relatively small areas, suitable for detailed, repeated analyses and geomorphological applications (Westoby et al., 2012; Turner et al., 2015).

In recent years, data provision for users has increased and today data hubs provide easy accessibility to rapid, pre-processed imagery. Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning. Nonetheless, technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non–existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time–critical analyses such as ground motion monitoring and landslide early warning.
2.2. Practical implementation of multispectral data in the concept

The study site (starting at L175) represents a very complex landslide case leading to rather erratic mass movements in form of debris flows initiated by changing slope water conditions related to increased atmospheric precipitation. This situation is another obstacle for an early warning approach which is solely based on optical remote sensing data and thus making it impossible to make full use of the in principle daily temporal resolution of the planet data. Taking into account these natural conditions and the constraints introduced by the used imaging constellations, leaves no room for true optimization of lead time in the sense as stated in the overall scientific goal of this paper.

We agree with your assessment and have replaced the term “optimisation” with a description that hopefully is more accurate in the entire manuscript. The chosen Sattelkar slide is one of the most relevant high-alpine geohazards in Austria and thus represents a compelling study site for natural hazard studies. While we agree that its complexity represents an obstacle, we nonetheless believe that the Sattelkar slide is well-suited for an investigation based on optical remote sensing because (i) we were clearly able to detect significant displacement and (ii) we were able to identify patches of increasing motion. In any case an increase in frequency of UAS flights is possible.

L39–41: This study presents a new concept to systematically evaluate remote sensing techniques to optimise estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for providing early warnings.

Any sensible early warning approach for slope movements requires a continuous and reliable high temporal resolution input of observation data related to parameters which are relevant for triggering the potential mass movements. Such information are mostly provided by ground based measurements. In this context, it is surprising that no relevant ground based monitoring information seem to be available to this study despite the longterm history of scientific work at this study site. The mentioned temperature loggers need to be explained in their function for early warning. The GPS measurements seem to only support the remote sensing based analysis. The described setting does not seem to be suitable for identification of precursory signs of ‘slope preparation’ related to the triggering of potential mass movements at this site in a way which would be required in the context of early warning.

Thank you for your feedback. We understand your arguments, yet we are not trying to
create an all-encompassing landslide early warning study that includes all state-of-the-art methods. We have chosen the Sattelkar due to its scientific and societal relevance and its high-alpine location with very limited vegetation. This site was not selected to evaluate a wide range of remote sensing applications. Our goal was to determine if and how our conceptual approach is applicable to this highly complex study site. Due to its topographical characteristics no ground based technique can be implemented. Therefore, only air- and spaceborne sensors can be employed which we believe is the case for numerous potentially hazardous slides/creeps in mountain ranges worldwide. However, we have considered installing a camera on the opposite slope but currently the distance is a problem (3.5 km, selection of camera).

We agree that the temperature data mentioned in the manuscript is not absolutely necessary to understand our conceptual approach. We still think that the (brief) inclusion of the temperature data makes sense as it suggests local permafrost presence/degradation which may be one of the main drivers of the Sattelkar slide. To clarify the role of the temperature data we amended the relevant sections in the study site section.

L175 et seq. [...] massive volumes of glacial and periglacial debris as well as rockfall deposits (Fig. 2b, c). Near surface temperature data indicates sporadic permafrost distribution in the upper part of the cirque. [...] allowing visual block tracking and delimiting the active process area. High displacement was measured between 2012 and 2015 with up to 30 m a⁻¹.

[...]

L200 et seq.: In the Sattelkar cirque, several monitoring components are installed to provide ongoing and long-term monitoring. Nine permanent ground control points (GCPs) measured with a dGPS to provide stable and optimal conditions to derive orthophotos from highly accurate UAS images (GeoResearch, 2018). A total number of 15 near surface temperature loggers (buried at 0.1 m depth) recorded annual mean temperatures slightly above the freezing point (1–2 °C) in the period 2016 to 2019. Ground thermal conditions at depth react with significant lag times to recent warming and therefore are primarily determined by climatic conditions of the past (Noetzli et al., 2019). Significantly cooler climatic conditions in previous decades and centuries (Auer et al., 2007) thus likely contributed to the formation of (patchy) permafrost at the Sattelkar cirque. Recent empirical–statistical modelling of permafrost distribution in the Hohe Tauern Range confirms possible permafrost presence at the study site (Schrott et al., 2012).

These components include 30 near surface temperature logger (NSTL) nine permanent ground control points (GCP) measured with a dGPS to provide stable and optimal conditions for the derivation of orthophotos from highly accurate UAS images (GeoResearch, 2018). Field-based mapping and measurements help to delimit the active process area.

Correct, the dGPS measurements are only used for repeated UAS campaigns and their data derivation. As described earlier, with our technical approach we were able to not only detect hot spots of total displacement but also to see changes in motion and thus certain areas of accelerating behaviour.

- L210: The complete dismissal of radar data is not justifiable in the current form since the authors only take into account InSAR based deformation analysis and neglect that
the technique of pixel offset tracking can be also be applied to the intensity component of radar data. For the mainly rainfall driven processes at the study site, the integration of radar data seems to be mandatory into any sensible remote sensing based early warning approach, since a combination of optical and radar data is required to establish an as continuous as possible time series of remote sensing observations. Thank you for mentioning radar data. We have described the application of InSAR/DInSAR in the introduction (L86–91) and placed the argument in section “4.1. Optical Imagery”.

For this particular site radar data is not practical. Even if foreshortening and layover effects are a minor issue for this site, the main reason to not include this kind of data is the fact that the velocity shows rates exceeding the limits of radar data leading to a loss of coherence.

L78 et seq.: As forecasted landslides tend to accelerate beyond the deformation rate observable with radar systems before failure, we concentrate on optical image analysis (Moretto et al., 2016). One advantage of optical imagery is its temporally dense data (Table 1) compared to open data radar systems with sensor visits repeat frequency more than every six days and revisit frequency between three days at the equator, about two days over Europe and less than one day at high latitudes (Sentinel–1, ESA). Optical data allows direct visual impressions from the multispectral representation of the acquisition target and the option to employ this data for further complementary and expert analyses. While active radar systems overcome constraints posed by clouds and do not require daylight, data voids can be significant due to layover or shadowing effects in steep mountainous areas (Mazzanti et al., 2012; Plank et al., 2015; Moretto et al., 2016). Moreover, north/south facing slopes are less suitable, thus limit the range of investigation (Darvishi et al., 2018). In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019).

• Moreover, taking into account the goal of lead time optimization, I consider it crucial to also include ground-based live-streamed time-lapse imagery in the proposed remote sensing based early warning approach (for an example see the Khan et al. (2021) paper ‘Low-Cost Automatic Slope Monitoring Using Vector Tracking Analyses on Live-Streamed Time-Lapse Imagery’ published in Remote Sensing).

Thank you for this idea and forwarding the information on the article of this useful approach for the ‘Rest and Be Thankful slope’, Scotland, with PIV on time–lapse imagery. For the Sattelkar we conducted preliminary investigations regarding the installation of a camera on the opposite slope. Due to the steep slope the camera would have to be mounted at the same altitude. This means a camera would have to be able to cover a horizontal distance of about 3.5 km. There is a higher chance of mobile network signal which is otherwise unavailable beginning at the entrance of the valley. Nevertheless, the power supply and issues such as rain drops and general pollution on the lense pose problems as Khan et al. (2021) also acknowledge.

The materials and methods section (4.) as well as the result section (5) are sound and well written. Since reviewer 1 has already focused on this part of the paper as well as the accuracy assessment and made detailed suggestions for improving these parts, I only have a few comments left to make on these aspects of the paper.

• L355: The authors state that core areas of the landslide are surrounded by wide fringes with no data. In this context the meaning of the term „no data” is not clear to me.
Please, explain, what do you mean by ‘no data’ – either missing results or zero deformation.

Thank you for pointing this out. Here by ‘no data’ we mean that there is zero deformation and we have revised the text accordingly.

L354 et seq.: No motion was present in a fringe zone along the landslide front (west boundary), similar to results in Fig. 5a and Fig. 5b. In general, the displacement patterns are less smooth than at 0.16 m input resolution. Outside the landslide significant displacements exist at the eastern image border (Fig. 5e) and towards the west (h, i) (Fig. 5f). In comparison, total displacement rates derived from PlanetScope cover in large parts the active area for Ib (Fig. 5c); however, for II only the core area of the landslide shows displacement. In both results the core areas of the landslide are surrounded by wide fringes with zero deformation.

- L370: Fig 6. The obtained deformation results show a very different degree of detail throughout the landslide. For better evaluation of the reasons for these differences the inclusion of an RGB UAV image of the same area would be helpful in order to be able to include surface texture properties in the evaluation of the obtained differences in the deformation patterns.

Thank you for your good suggestion. We added the corresponding master and slave image below the presented DIC result. The caption has been adjusted accordingly.
Conclusions related to the results presented until L370: The presented specific deformation results obtained from the analyzed planet and UAV data, represent a valuable contribution towards an improved area-wide process understanding of so far unprecedented detail for this study site. Conceptually, such investigations mainly contribute to the preparedness phase within the disaster management cycle. Continuation of monitoring of the study site using the described approach would represent a very valuable prerequisite for developing and setting up a true early warning system for this site combining ground based and remote sensing observations. However, the results presented in this paper do not allow optimization of lead times within an early warning approach being stated being as the goal of this paper. Our approach is not to set up a comprehensive early warning system, which includes all four elements defined by the UNISDR (2006) (see L35–38). We agree that optimisation of lead time does not accurately represent what we have done in our study. Thus we have revised our manuscript to make it more precise (see
changes to the manuscript here on p. 1, 3–4). Our concept enables us to evaluate lead
time based on our proposed structure.

Introduction, L10–11: We introduce a novel conceptual approach for comprehensive
structure and quantitatively assess lead time assessment and optimisation for LEWS.

L39–41: This study presents a new concept to systematically evaluate remote sensing
techniques to optimise estimate and increase lead time for landslide early warnings
in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing
data for to providing early warnings.

Conclusion, L578 et seq.: Coarse temporal data resolution, such as in the case study investigated here, represents an important restriction to the use of optical remote sensing data for landslide early warning applications. Acceleration (and the resulting failure) over short periods of time will likely go unnoticed due to large data acquisition intervals. However, for prolonged acceleration periods, such as observed at the Sattelkar slide and many other relevant hazard sites, the chosen data sources have been demonstrated to represent a formidable early warning approach capable of contributing to an improved risk analysis and evaluation in steep high–alpine regions.

- L375: 5.3 Time required for collection, processing and evaluation. The presented
analysis is rather meaningless, since the scarcity of the available time steps does not allow the detection of critical process stages. Taking into account the big temporal gaps between the data acquisitions, the time needed for handling the planet and UAV imagery is not really relevant for lead time optimization. The obtained times only allow a relative comparison between planet and UAV based data acquisition within the narrow limits of the chosen approach. However, true early warning would require setting up a semi-automated processing chain including automated download and screening of available remote sensing data as well as semi-automated subsequent deformation analysis reducing data handling time to a minimum. Under such conditions, primary remote sensing data availability becomes the crucial decisive factor determined by the data distribution procedures of the satellite data providers and the atmospheric conditions in case of optical imagery. In conclusion, it needs to be stated that the used parameter of time to warning is only applicable under the condition of a near real time continuous data stream of input information which is not available within the presented study.

Thank you for your comment which helps to clarify your understanding of our text. We did not intend to create a ‘true early warning’ as you described. This was not the goal of our study. The repeated measurements allow the detection of spatial and temporal acceleration patterns and we believe the repeated measurements can be scaled to early warning demands. With regard to your comment on a semi-automated processing chain we do not fully agree. Based on our knowledge, even in case of most geotechnical investigations, the data is analysed by experts prior to issuing an early warning (e.g. https://www.bgu.tum.de/landslides/alsense/projekt/, Leinauer et al. (2020): DOI: 10.1002/geot.202000027).

- L390: In the current form of the paper the points raised in the discussion (6.) are only relevant in the frame of a process-oriented study and not for early warning purposes
since the latter one requires the identification of precursors for critical process stages –
tipping points – which are likely to trigger substantial complex mass movements later
turning into potentially catastrophic debris flows. 
It is our understanding, we can only provide early warnings for processes we
understand. The processual understanding is key to anticipating the magnitude, timing,
and reach of alpine hazards, thus processual understanding and early warning cannot
be separated.

- L490: Estimating time to warning (6.3). This part of the discussion also suffers from
the conceptual limitations which have already been pointed out earlier in this review.
A comparison of lead times between the different example landslides would only be
meaningful in case of continuous high resolution temporal information on deformation
allowing the identification of precursory events which is usually only possible using
ground based observations. The presented comparison between potential repeat rates
of remote sensing data acquisitions and retrospectively derived lead times is too
simplistic (Fig. 8), since the main remaining question is, whether the relevant
deformation (cracks etc.) can be first, resolved by the used imagery and second,
distinguished from other surface disturbances by the used analysis methods.
In this paper, in contrast to remote sensing papers, the time scale required for effective
early warnings is given by nature, i.e., the typical acceleration patterns of particular
landsides.
With regard to the comparison of historic events, we referred to their natural landslide
processes which delimits the possible lead time. Unfortunately, a comparison to these
historic examples is limited to a retrospective view. We agree with you regarding the
detection of relevant deformations. If the sensors evaluated here could have identified
the motion excluded disturbances, then in this temporal concept UAS and PlanetScope
would have been able to show an acceleration in a timely fashion.
We want to keep this concept simple to allow the transfer for required processing
times from other sensors. The main question is, if the time is sufficient for the whole
processing prior to landslide release.
L:148–149 Natural processes and natural developments constantly take place independently, thus
dictate the technical approaches and methodologies researchers must apply within a certain
time period.

Overall recommendation:
The presented results comprise a very interesting process-oriented study evaluating the
use of planet and UAV imagery for the derivation of spatiotemporally differentiated
deformation information for a rather large and topographically pronounced terrain affected
by complex mass wasting processes. I consider these findings well worth being published
in this journal. However, the publication of these specific results requires a major conceptual
reframing of the work which is targeted at the real potential usability of these results which
cannot be early warning because of the reasons already stated in this review.

However, the work presented in this study has the potential to form an important basis for the
development of a true early warning concept / approach in the future combining remote
sensing and ground based observations targeting at the same parameters allowing a multi-
scale assessment of surface deformation related to triggering potential catastrophic mass
movements at the study site.
Challenging the timely prediction of landslide early warning systems with multispectral remote sensing: a novel conceptual approach tested in the Sattelkar, Austria

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Abstract
While optical remote sensing has demonstrated its capabilities for landslide detection and monitoring, spatial and temporal demands for landslide early warning systems (LEWS) were not met until recently. We introduce a novel conceptual approach to comprehensively assess lead time, and optimisation for LEWS. We analysed "time to warning" as a sequence; (i) time to collect, (ii) to process and (iii) to evaluate relevant optical data. The difference between "time to warning" and "forecasting window" (i.e. time from hazard becoming predictable until event) is the lead time for reactive measures. We tested digital image correlation (DIC) of best-suited spatiotemporal techniques, i.e. 3 m resolution PlanetScope daily imagery, and 0.16 m resolution UAS derived orthophotos to reveal fast ground displacement and acceleration of a deep-seated, complex alpine mass movement leading to massive debris flow events. The time to warning for UAS and PlanetScope totals 31h/21h and is comprised of (i) time to collect 12/14h, (ii) process 17/5h and (iii) evaluate 2/2h, which is well below the forecasting window for recent benchmarks and facilitates lead time for reactive measures. We show optical remote sensing data can support LEWS with a sufficiently fast processing time, demonstrating the feasibility of optical sensors for LEWS.

1 Introduction
Landslides are a major natural hazard leading to human casualties and socio-economic impacts, mainly by causing infrastructure damage (Dikau et al., 1996, Hilker et al., 2009). They are often triggered by earthquakes, intense short-period or prolonged precipitation, and human activities (Hungert al., 2014; Froude and Petley, 2018). In a systematic review, Gariano and Guzzetti (2016) report in a review study, that 80% of the papers examined show causal relationships between landslides and climate change. The ongoing warming of the climate (IPCC, 2014) is likely to decrease slope stability and increase landslide activity (Huggel et al., 2012; Seneviratne et al., 2012), which, thus indicates a vital need to improve the ability to detect, monitor and issue early warnings of landslides and thus to reduce and mitigate landslide risk.

Early warning, as defined by the UN International Strategy for Disaster Reduction (UNISDR), refers to a set of capacities for the timely and effective provision of warning information through institutions, such that individuals, communities and organisations exposed to a disaster are able to take action with sufficient time to reduce or avoid risk and prepare an effective response (UNISDR, 2009). According to UNISDR (2006), an effective early warning system consists of four elements: (1) risk knowledge, the systematic data collection and risk assessment; (2) the monitoring and warning service; (3) the dissemination and communication of risk as well as early warnings; and (4) the response capacities on local and national levels. Incompleteness or failure of one element can lead to a breakdown of the entire system (L20-10). Lead time, as defined in the context of LEWS is the interval between the issue of a warning (i.e. dissemination) and the forecasted landslide onset (Pecoraro et al. 2019) and thus crucially depends on time requirements in phases (1)- (3). The success of an EWS therefore requires measurable pre-failure motion (or slow slope displacement) to allow for sufficient lead time for decisions on reactions and counter measures (Grasso, 2014; Hunger et al., 2014).

1 Commentaries
Kommentiert [DH1]: Modified following comment by RC2 (S. Roessner)
RC2 (S. Roessner) wrote:
The used data sources (planet and UAv) do not allow optimization of lead time in the context of early warning because of the scarcity of their availability which is reflected in the small number of only three multitemporal data taken between July and September analyzed in this study (Table 3).

Kommentiert [DH2]: Answer to comment by RC1 (J. Blöthe) in response
RC1 (J. Blöthe) wrote:
L1: Landslide
RC1 (J. Blöthe) wrote:
L22/23: While this is certainly true, the authors should elaborate in the introduction that events instantaneously triggered by earthquakes or heavy precipitation are beyond what their proposed framework can deliver an early warning for. The necessity of gathering and evaluating data prior to issuing a warning limits the analysis to mass movements that indeed show a pre-accuracy acceleration on the order of days.

Kommentiert [DH3]: Modified following comment by RC1 (J. Blöthe) author below
RC1 (J. Blöthe) wrote:
L22/23: Is this really just attributable to the warning of the climate?

Kommentiert [DH4]: Author’s revision in order to be more concise
RC1 (J. Blöthe) wrote:
L25/26: Is this really just attributable to the warning of the climate?

Kommentiert [DH5]: Answer to comment by RC1 (J. Blöthe) in response
RC1 (J. Blöthe) wrote:
L1: Landslide
RC1 (J. Blöthe) wrote:
L22/23: Is this really just attributable to the warning of the climate?

Kommentiert [DH6]: Author’s revision in order to be more concise
RC1 (J. Blöthe) wrote:
L1: Landslide
RC1 (J. Blöthe) wrote:
L22/23: Is this really just attributable to the warning of the climate?

Kommentiert [DH7]: Modified and moved this quite following comment by RC2 (S. Roessner)
RC2 (S. Roessner) wrote:
The missing warning conceptual approach is also reflected in the introduction in form of a lengthy summary of principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 3 is very general and not specific to landslide and does not qualify as a novelty in the current form.

Kommentiert [DH8]: Author’s revision in order to be more concise
RC2 (S. Roessner) wrote:
The missing warning conceptual approach is also reflected in the introduction in form of a lengthy summary of principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 3 is very general and not specific to landslide and does not qualify as a novelty in the current form.

Kommentiert [DH9]: Added definition following comment by RC2 (S. Roessner)
RC2 (S. Roessner) wrote:
Concept of lead time and need for best possible reduction is not new and... The missing warning conceptual approach is also reflected in the introduction in form of a lengthy summary of principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 3 is very general and not specific to landslide and does not qualify as a novelty in the current form.

Kommentiert [DH10]: Modified following comment by RC2 (S. Roessner)
RC2 (S. Roessner) wrote:
The missing warning conceptual approach is also reflected in the introduction in form of a lengthy summary of principle available remote sensing methods and data showing no clear line of arguments (L20-100). Moreover, the new conceptual approach presented in Fig. 3 is very general and not specific to landslide and does not qualify as a novelty in the current form.
This definition of an early warning system (EWS) contains a time component but includes no exact time scale references. “Early” suggests that events are detected before harm or damage occurs and thus stands in contrast to events which are only detected once they have begun (e.g., snow avalanches). Thus, it is necessary to know sensor capabilities and limitations for pre-event mass movement observations (Dresner et al., 2019). The success of a warning requires that information is provided with enough lead time for decisions on reaction and counter measures (Grasso, 2014). The success of an EWS therefore requires measurable pre-failure motion or slow-transport velocities to allow for sufficient lead time for decisions on reaction and counter measures (Grasso, 2014). While remote sensing has been established for early warnings, remote sensing is not yet used for real early warnings of the onset of landslides in steep–alpine terrain (with a few exceptions), where post-earthquake instruments are still preferred. Exceptions include terrestrial InSAR (Pesci et al., 2011; Walter et al., 2020) and terrestrial laser scanning with high repetition rates. However, repeated UAS (unmanned aerial systems) and optical satellite images (PlanetScope) with high repetition rates have so far not been applied for landslide early warning in steep–alpine catchments. In this regard, knowledge of sensor capabilities and limitations is essential, as it determines which rates and magnitudes of pre-failure motion can potentially be identified (Dezsö et al., 2019). Our proposed framework refers to mass movements in steep–alpine catchments with significant pre-failure motion operating over sufficiently long periods and thus excludes instantaneous events triggered by processes such as heavy rainfall or earthquakes.

According to UNISDR (2006), an effective early warning system consists of four elements: (1) risk knowledge, i.e., the systematic collection and risk assessment of the monitoring and warning service; (2) the dissemination and communication of risk, i.e., early warnings; and (3) the response capacities on local and national levels. Incompleteness or failure of one element can lead to a breakdown of the entire system (ibid.).

This study presents a new concept to systematically evaluate remote sensing techniques to optimize estimate and increase lead time for landslide early warnings in these catchments. We do not start from the perspective of available data; instead, we define necessary time constraints to successfully employ remote-sensing data for early warnings. This approach reduces the to a small number of highly suitable remote sensing products with high temporal and spatial resolution. With these constraints, we investigated the application of data from satellites and unmanned aerial systems (UAS) to allow the assessment of the data, after a spaceborne area-wide but low-resolution acquisition, into a downscaled detailed image recording. In so doing, we analysed the capability of these different passive remote sensing systems focusing on spatiotemporal capabilities for ground motion detection and landslide evolution to provide early warnings.

Finally, recently, the spatial and temporal resolution of optical satellite imagery has significantly improved requirements for accurate early warning purposes have not been met by optical satellite imagery (Scaioni et al., 2014) and has allowed substantial advantages in the definition of displacement rates and acceleration thresholds to approach requirements for early warning purposes. This is essential since spatial and temporal resolution determines whether landslide monitoring is possible with the detection allowing defining of displacement rates and approaching approximate acceleration thresholds, both of which are lacking if information is based solely on post-event studies (Reid et al., 2008; Calvello, 2017). Landslide monitoring therefore not only deepens the understanding of landslide processes but also has potential to significantly advance landslide early warning systems (LEWS) (Chac et al., 2017; Crosta et al., 2017). Previously, high spatial resolution satellite data was obtained at the expense of a reduction in revisit rates (Aubrecht et al., 2017). Consequently, the return period between two images increased, limiting ground displacement assessment and the range of observable motion rates.

The number of useful images was further reduced due to natural factors such as snow cover, cloud cover and cloud shadows. High-resolution remote sensing data was long restricted due to high costs and data volume (Goodchild, 2011; Westoby et al., 2012).

Today commercial very high resolution (VHR) optical satellites exist, but tasked acquisitions make them inflexible and very costly, thus limiting research (Butler, 2014; Lucier et al., 2014). There is a vast spectrum of available remote sensing data with high spatiotemporal resolution (Table 1). Complementary use of different remote sensing sources can significantly improve the potential to significantly advance landslide early warning systems (LEWS) (Chac et al., 2017; Crosta et al., 2017).
improve landslide assessment as demonstrated by Stumpf et al. (2018) and Bontemps et al. (2018), who draw on archive data and utilise different sensor combinations to analyse the evolution of ground motion.

### Table 1
Overview of different optical multispectral remote sensors with their corresponding resolution (m) and revisit rate (days). The sensors are categorised into commercial and free data policy. 1Free quota via Planet Labs Education and Research Program. 2PlanetScope Ortho Scene Product, Level 3B/Ortho Tile Product, Level 3A (Planet Labs, 2020a), 3reached end of life, 3/2020, archive data usable, 45 m Ortho Tile Level 3A (Planet Labs, 2020a), 70.5 m colour pansharpened, 5self-acquired. Source: (ESA, 2020).

<table>
<thead>
<tr>
<th>Sensor</th>
<th>Temporal resolution (d)</th>
<th>Spatial resolution (m)</th>
<th>Free/Commercial</th>
</tr>
</thead>
<tbody>
<tr>
<td>UAS</td>
<td>flexible</td>
<td>0.08</td>
<td>F^*</td>
</tr>
<tr>
<td>WorldView 2</td>
<td>1.1</td>
<td>1.84</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 3</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>WorldView 4</td>
<td>&lt;1</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye 2</td>
<td>5</td>
<td>1.24</td>
<td>C</td>
</tr>
<tr>
<td>SkySat</td>
<td>1</td>
<td>1.5</td>
<td>C</td>
</tr>
<tr>
<td>GeoEye 1</td>
<td>3</td>
<td>1.64</td>
<td>C</td>
</tr>
<tr>
<td>Pléiades 1A/B</td>
<td>1</td>
<td>2.0 (0.5)^3</td>
<td>C</td>
</tr>
<tr>
<td>PlanetScope</td>
<td>1</td>
<td>3.0/0.125^2</td>
<td>C/F^2</td>
</tr>
<tr>
<td>RapidEye</td>
<td>5.5</td>
<td>5^5</td>
<td>F</td>
</tr>
<tr>
<td>Sentinel-2 A/B</td>
<td>5</td>
<td>10</td>
<td>F</td>
</tr>
<tr>
<td>Landsat 8</td>
<td>16</td>
<td>30</td>
<td>F</td>
</tr>
</tbody>
</table>

The latest developments in earth observation programs include both the new Copernicus’ Sentinel fleet operated by the ESA, and a new generation of micro cube satellites, sent into orbit in large numbers by PlanetLabs Inc. These PlanetScope micro cube satellites, known as ‘Doves’ (PlanetScope, from now on referred to as PlanetScope satellites), and Sentinel-2 a/b offer very high revisit rates of 1–5 days and high spatial resolutions from 3–10 m, respectively (Table 1), for multispectral imagery (Drusch et al., 2012; Butler, 2014; Breger, 2017). This opens up unprecedented possibilities based on these high spatiotemporal resolutions to study a wide range of landslide velocities and natural hazards through remote sensing. Future/Continuing data access is fostered by PlanetLabs and by Copernicus (via its open data policy) providing affordable or free data for research. This leads to unprecedented possibilities for studying natural hazards through remote sensing. Examples of landslide activity such as multi-temporal studies employing multi-temporal datasets of landslide activities based on this access to high spatiotemporal data are include Lacroix et al. (2018), using Sentinel-2 scenes to detect motions of the Harmalière landslide in France, and Mazzanti et al. (2020), who applied a large stack of PlanetScope images for the active Rattlesnake landslide, USA.

As forecasted, landslides tend to accelerate beyond the deformation rate observable with radar systems before failure, we concentrate on optical image analysis (Moretto et al., 2018). One advantage of optical imagery is its temporally dense data (Table 1) compared to open data radar systems with sensor revisit frequency more than every six days and revisit frequency between three days at the equator, about two days over Europe and less than one day at high latitudes (Sentinel-1, ESA). Optical data allows direct visual impressions from the multispectral representation of the acquisition target and the option to employ this data for further complementary and expert analyses. While active radar systems overcome constraints posed by clouds and do not require daylight, data voids can be significant due to layover or shadowing effects in steep mountainous areas (Giordan et al., 2018; Mazzanti et al., 2012; Plank et al., 2015; Moretto et al., 2016). Moreover, north/south facing slopes are less suitable, thus limit the range of investigation (Darvishi et al., 2018).

In general, sensor choice depends on the landslide motion rate with radar at the lower and optical instruments at the upper motion range (Crosetto et al., 2016; Moretto et al., 2017; Lacroix et al., 2019).

However, a flexible, cost-effective alternative to spaceborne optical data are airborne optical images taken by UASs (unmannad aerial systems). Freely selectable flight routes and acquisition dates prevent enable avoiding shadows from clouds and topographic obstacles and as well avaliaing unfavourable weather conditions and summer time snow cover, all of which frequently impair satellite images (Giordan et al., 2018; Lucieer et al., 2014). UAS-based surveys provide accurate
very high resolution (few cm) orthoimages and digital elevation models (DEM) of relatively small areas, suitable for detailed, repeated analyses and geomorphological applications (Westoby et al., 2012; Turner et al., 2015). In recent years, data provision for users has increased and today data hubs provide easy accessibility to rapid, pre-processed imagery. Knowledge of the most useful remote sensing data options is vital for complex, time-critical analyses such as ground motion monitoring and landslide early warning. Nonetheless, technological advances can be misleading as they promise high spatiotemporal data availability, which frequently does not reflect reality (Sudmanns et al., 2019). One key problem is the realistic net temporal data resolution which is often significantly reduced due to technical issues, such as image errors and non-existent data (i.e. data availability, completeness, reliability). Other problems include data quality and accuracy in terms of geometric, radiometric and spectral factors (Batini et al., 2017; Barsi et al., 2018). Knowledge of the most useful remote sensing data options is vital for complex, time-critical analyses such as ground motion monitoring and landslide early warning. Timely information extraction and interpretation are critical for landslide early warning; yet few studies have so far explicitly focused on time criticality and the influence of the net temporal resolution of remote sensing data.

In this investigation we propose both a conceptual approach to evaluating lead time as a time difference between the “time to predict” and the “forecasting time” and assess the suitability of UAS sensors (0.16 m) and PlanetScope (3 m) imagery (the latter with temporal proximity to the UAS acquisition) for LEWS. For this we have chosen the ‘Sattelkar’, a steep, high-alpine cirque located in the Hohe Tauern Range, Austria (Anker et al., 2016). We estimate times for the three steps (i) collecting images, (ii) pre-processing and motion derivation by digital image correlation (DIC) and (iii) evaluating and visualizing. The results from the Sattelkar site – and from historic landslide events – will be discussed in terms of usability and processing duration for critical data source selection which directly influences the forecasting window. Accordingly, we try to answer the following research questions:

1. How can we evaluate lead time as a time difference between the “time to predict” and the forecasting time for high spatiotemporal resolution sensors?
2. How can we quantify “time to warning” as a sequence of (i) time to collect, (ii) to process and (iii) to evaluate relevant optical data?
3. How can we practically derive profound “time to warning” estimates as a sequence of (i), (ii) and (iii) from UAS and PlanetScope high spatiotemporal resolution sensors?
4. Are estimated “times to warning” significantly shorter than the forecasting time for recent well-documented examples and able to generate robust estimations of lead time available to enable reactive measures and evacuation?

2 Lead time – a conceptual approach

2.1. The conceptual approach

Natural processes and natural developments constantly take place independently, thus dictate the technical approaches and methodologies researchers must can and must apply within a certain time period. For that reason, we hypothesise the forecasting window $t_{\text{real}}$ is externally controlled, consequently the applicability of LEWS methods ($t_{\text{real}}$) is restricted because they must be shorter than $t_{\text{natural}}$. This approach is the framework of our time concept (Fig. 1).
The forecasting window starts \((t_{\text{lead}})\) following significant acceleration exceeding a set displacement threshold, leading to a continuous process. Simultaneously with the forecasting window, warning \((t_{\text{warning}})\) starts (grey outline). Time to warning is divided into a three-phase-process to allow time estimations for a comparative assessment of different types of remote sensing data. This process consists of the phases (1) time to collect, (2) time to process and (3) time to evaluate, with their individual durations. Confidence in the forecasted event increases with time as process acceleration becomes more certain. Once a warning is released (orange box), the lead time begins \((t_{\text{lead}})\) and is terminated by the following release and subsequent impact (red box). The lead time is the difference between the forecasting window and the time to warning. During the lead time, reaction time \((t_{\text{react}})\) starts when appropriate counter measures are taken to prepare for and reduce risks ahead of the impending event, and ends with the final impact.

The time to warning period \((t_{\text{warning}})\) is defined by the time necessary to systematically collect data, analyse the available information and to evaluate it. Hence, the greater the lead time, the more extensive countermeasures can be implemented prior to the event. An imperative for an effective EWS, the required time to take appropriate mitigation and response measures has to be within the lead time interval \((t_{\text{lead}})\) (Pecoraro et al., 2019) with \(t_{\text{lead}} \geq t_{\text{react}}\).

### 2.2. Practical implementation of multispectral data in the conceptual approach: General applicability to optical data

The time to warning consists of a three-phase-process (see Sect. 2.1. and Fig. 1) to allow rough time estimations for a comparative assessment of different types of remote sensing data. Nevertheless, to realise this temporal concept an established, operating system is required, which includes reference data (DEM, previous results), experience from past field work and ready UAS flight plans with preparation for a UAS flight campaign, satellite data access, experience in the single software processing steps including final classification and visualisation templates and, if utilised for UAS, installed and measured ground control points.
The first phase includes the collection of data starting from the acquisition by the sensor, the data transfer, image pre-processing and provision to the end user. The user selects images online from the data hub, downloads and organises them. For a UAS campaign, the user must obtain flight permits, check flight paths and conduct the UAS flight. The second phase encompasses time to process for the complete data handling from the downloaded data to final analysis-ready image stacks in a GIS or a corresponding software. These preparatory steps may include image selection and renaming, atmospheric correction, co-registration, resampling and translation to other spatial resolutions and geographic projection systems, adjustments such as clipping, stacking of single bands into one multispectral image or the division into single bands, calculation of hillshade from DEM among others, depending on the requirements. Following this preparation, the data is processed with the appropriate software tools to derive ground motion, calculate total displacement and derive surface changes, e.g. volume calculations or profiles. In the third and last phase, time to evaluate, the results are compared to inventory data and, if available, ground truth data, displacement results of other sensors or different spatial resolutions, different time interval variations to observe changes in sensitivity to meteorological conditions. Additionally, filters may be applied to eliminate noise. Finally, the results are analysed and evaluated. In each phase quality management is carried out for data access and pre- and post-processing. In time to collect, the images must be selected manually prior to any download from the data hub, as its filter tool options on cloud and scene coverage are of limited help. Accordingly, the areal selection may be misleading as the region of interest (RoI) might not be fully covered, though the sought-for, smaller area of interest (AoI) is covered but not returned from the request.

Concerning cloud filters, first, the filter refers to the RoI as a whole in terms of percentage of cloud coverage. The AoI can still be free of clouds or else be the only area covered by clouds in the total RoI. Therefore, an image is either not returned although usable, or returned but not useable. Second, clouds can create shadows for which no filter is available. As a result, affected images have to be manually removed by the user. Images which are of low quality due to snow cover have to be discarded, too. These actions indirectly represent first quality checks in the collection phase. In the following processing phase, the images in a GIS, are checked for quality and accuracy. Depending on the data provider, some pre-processing such as radiometric, atmospheric and geometric corrections may have been conducted. During this phase, additional user-based steps will be checked if necessary. Finally, the results are compared to other data (e.g. DEM, dGPS), reviewed for their validity and may be supplemented by statistical evaluation.

3 Study Site

The Sattelkar is a high-alpine, deglaciated west-facing cirque at an altitude of between 2,130–2,730 m asl in the Obersulzbach valley, Großvenedigergruppe, Austria (Fig. 2a). Surrounded by a headwall of granitic gneiss, the cirque infill is characterised by massive volumes of glacial and periglacial debris as well as rockfall deposits (Fig. 2b, c). Heat–surface temperature data indicate sporadic permafrost distribution in the upper part of the cirque. Since 2003 surface changes have taken place as evidenced by a massive degradation of the vegetation cover and the exposure and increased mobilisation of loose material. A terrain analysis revealed that a deep-seated, retrogressive movement in the debris cover of the cirque had been initiated (Anker et al., 2016; GeoResearch, 2018). High water (over)saturation is assumed to be causing the spreading and sliding of the glacial and periglacial debris cover on the underlying, glacially smoothed bedrock cirque floor forming a complex landslide (Hungr et al., 2014). Detailed aerial orthophoto analyses, witness reports and damage documentations indicate a steady increase in mass movement and debris flow activity over the last decade (Anker et al., 2016).
In August 2014, heavy ongoing precipitation triggered massive debris flow activity of 170,000 m³ in volume, of which approximately 70,000 m³ derived from the catchment above 2,000 m. A further 100,000 m³ was mobilised in the channel within the cone. The consequence was that the Obersulzbach river was blocked leading to a general flooding situation in the catchment, resulting in substantial destruction in the middle and lower reaches (Fig. 3).

The Sattelkar has been the focus of international research projects such as “PROJECT Sattelkar” (GeoResearch, 2018) and AlpSenseBench (TUM, Chair of Landslide Research, 2020) since 2018. In 2015 preliminary findings revealed a mass movement coverage of 130,000 m² with approximately 1 mio. m³ of debris and displacement rates of more than 10 m a⁻¹. The debris consists of boulders up to 10 m in diameter (Fig. 2c, d) allowing visual block tracking and delimiting the active process area. High displacement was measured between 2012 and 2015 with up to 30 m a⁻¹.

In the Sattelkar cirque, several monitoring components are installed to provide ongoing and long-term monitoring. Nine permanent ground control points (GCPs) measured with a dGPS to provide stable and optimal conditions to derive orthophotos from highly accurate UAS images (GeoResearch, 2018). A total number of 15 near surface temperature loggers (buried at 0.1 m depth) recorded annual mean temperatures slightly above the freezing point (1–2 °C) in the period 2016 to 2019. Ground thermal conditions at depth react with significant lag times to recent warming and therefore are primarily determined by climatic conditions of the past (Noetzli et al., 2019). Significantly cooler climatic conditions in previous decades and centuries (Auer et al., 2007) thus likely contributed to the formation of (patchy) permafrost at the Sattelkar cirque. Recent empirical
statistical modelling of permafrost distribution in the Hohe Tauern Range confirms possible permafrost presence at the study site (Schrott et al., 2012).

These components include 3D near-surface temperature logger (NSTL) and nine permanent ground control points (GCPs) measured with a DGPS to provide stable and optimal conditions for the derivation of orthophotos from highly accurate UAS images (GeoResearch, 2018). Field-based mapping and measurements help to delimit the active process area.

The Satellkar is a suitable case study as it is in the early stages of the landslide development and thus fits best to this conceptual approach. Here, processes take place on time scales appropriate for long-term observation to provide sufficient warning time. The active part of the cinqueparte has accelerated in recent years allowing the analysis of EWS concepts based on multispectral remote sensing data supported by complementary block tracking.

4 Materials and Methods

4.1. Optical imagery

Optical satellite imagery is more appropriate for high deformation studies than radar applications due to the high spatial resolution as well as the short time span between acquisitions (Delacourt et al., 2007). Although the west-facing slope is favourable for the application of radar derivatives (InSAR/DInSAR), the choice to use optical imagery is based on the observed high displacement rates, which cause decorrelation when using radar technologies as they are more sensitive than optical technologies.

Complementary to the use of edge-enhanced radar products, optical imagery offers excellent spatial resolution and accuracy at the centimetre scale (Turner et al., 2015) and can be used for geomorphological processes in steep terrain (Lacroix et al., 2019). Here we employ DIC to compare the spatiotemporal resolution of multispectral optical imagery (UAS and PlanetScope) and to assess its suitability for early warning purposes. UAS images offer excellent spatial resolution and accuracy at the centimetre scale (Turner et al., 2015) and complement large scale satellite or airborne acquisitions (Lucieer et al., 2014).

PlanetScope imagery provides the highest temporal resolution among available sensors with daily acquisitions, guaranteed data availability, and free and open access for research purposes. In this study the PlanetScope Analytic Ortho Scene SR (surface reflectance) imagery (16-bit, geometric-, sensor- and radiometric corrections) was employed (Planet Labs, 2020b) and was supported by the Planet Labs Education and Research Program.

4.2. Data availability of PlanetScope

Research on the availability and usability of PlanetScope imagery was conducted on the Planet Explorer data hub for the time span from the beginning of April to the end of October in 2019, as during these months snow cover should be negligible. Filter parameters were solely set for 4-band PlanetScope Ortho Scenes and the Satellkar Aol. In order to obtain all available images, no filters (e.g. sun azimuth, off nadir angle) were applied. We defined four categories i) meteorological constraints due to snow cover, cloud cover and cloud shadow; ii) image (coverage) errors made by the provider; iii) no data availability and iv) the remainder of usable data (Table 2). The output request was evaluated according to the defined categories and was compared to the provider’s guaranteed daily image provision, which is comprised of 213 days for the time period (01.04.2019–31.10.2019). We calculated percentages for the above categories based on days per month as well as a seven–month sum and percentage average. The availability analysis did not include an examination of the data with regard to its spatial usability: positional accuracy and/or image shifts.
Unfavourable meteorological influences of cloud cover/shadow and snow cover affected up to 32.3 % and up to 33.3 %, respectively, on all 213 days; on average 14.5 % and 7 % of the days were not usable (Table 2). For 10 days in June snow influence had the greatest negative share (33.3 %), for April there were three days of snow coverage and the months September and October each had one day of snow coverage. Cloud cover/shadow exerted a higher impact on data usability by 14.5 %. Problems on the part of PlanetLabs made much of the data unusable due to image errors, between four and nine images per month were not usable (21 %). On average for 26.2 % of the analysed time period no image data was available. In this seven-month period, 43 images (20.1 %) had data voids or did not cover the AoI, thus the overall usability is limited to about 11.5 %.

### 4.3. Data Acquisition and Processing

In line with the concept in Fig. 1 (Sect. 1), the following processing steps are categorised and described.

1. **UAS data acquisition**
   - Extended: UAS data acquisition was preceded by detailed flight route planning and checks of local weather and snow conditions. UAS flights were carried out with a DJI Phantom4 UAS on 13.07.2018, 24.07.2019 and 04.09.2019 (see Table 3, Fig. 4).

<table>
<thead>
<tr>
<th>Acquisition set</th>
<th>UAS</th>
<th>PlanetScope</th>
</tr>
</thead>
</table>

For each acquisition, the total area was covered by four flights which were started on different elevations (Table 4). Flight planning was done with UgCS maintaining a high overlap (front: 80 %, side: 70 %) and a target ground sampling distance (GSD) of 7 cm. The area covered was approximately 3.4 km² and with a flight speed of about 8 m/s total flight time took 3.5 hours. The images were captured in RAW format. In the Planet Explorer Data Hub, PlanetScope Ortho Scenes were selected for usability; imagery affected by snow cover, cloud cover, cloud shadow and partial AoI coverage was discarded (Table 5).

### Table 4 UAS Flight plans.

<table>
<thead>
<tr>
<th>Flight plan parts</th>
<th>Length of flightpath [km]</th>
<th>Flight time [min]</th>
<th>Passes</th>
<th>No. of images</th>
<th>GSD [cm]</th>
<th>Altitude start point [m]</th>
<th>Highest flight position [m]</th>
<th>Lowest terrain point [m]</th>
</tr>
</thead>
<tbody>
<tr>
<td>Top</td>
<td>6.8</td>
<td>17</td>
<td>6</td>
<td>121</td>
<td>7</td>
<td>2630</td>
<td>3120</td>
<td>2365</td>
</tr>
<tr>
<td>Middle</td>
<td>7.5</td>
<td>19</td>
<td>6</td>
<td>135</td>
<td>7</td>
<td>2200</td>
<td>2682</td>
<td>1820</td>
</tr>
<tr>
<td>Low 1</td>
<td>7.3</td>
<td>17</td>
<td>6</td>
<td>130</td>
<td>7</td>
<td>1768</td>
<td>2115</td>
<td>1620</td>
</tr>
<tr>
<td>Low 2</td>
<td>5.6</td>
<td>14</td>
<td>6</td>
<td>81</td>
<td>7</td>
<td>1768</td>
<td>2110</td>
<td>1620</td>
</tr>
</tbody>
</table>
to measure the

Thereafter, a second selection

The final selection of images was made

based on the temporal proximity to the UAS data to guarantee the best comparability. For acquisition set (1), there are two PlanetScope images (02.07.2018 and 19.07.2018) which differed from the UAS acquisition date (13.07.2018) by 11 and 6 days, respectively. For acquisition sets (2) and (3), PlanetScope and UAS acquisition dates were identical (24.07.2019 and 04.09.2019). The acquired data sets were categorised in chronological intervals Ia/Ib and II (see Fig. 4). The PlanetScope images (19.07.2018–24.07.2019 and 03.09.2019) were taken between 11:35 and 11:42 local time.

Figure 4 Acquisition dates of UAS and PlanetScope images within the investigated time period. Calculated interval I for UAS images (13.07.2018–24.07.2019, 376 d) and interval Ib for PlanetScope images (19.07.2018–24.07.2019, 370 d), interval II for UAS and PlanetScope images (24.07.2019–04.09.2019, 42 d). Note: Ia PlanetScope interval was discarded.

The UAS images in RAW format were modified using Adobe Exposer to improve contrast, highlights, shadows and clarity. Thereafter, they were exported as JPEG (compression 95 %) and processed with Pix4Dmapper to 0.08 m resolution and orthorectified based on nine permanent ground control points (GCP, 30 x 30 cm). These were repeatedly (1000 measurements/position) registered with the TRIMBLE R5 dGPS and corrected via the baseline data of the Austrian Positioning Service (APOS) provided by the BEV (Bundesamt für Eich- und Vermessungswesen). Horizontal root-mean-squared errors (RMSE) range from 0.05 m to 0.10 m for vertical RMSE. These GCPs were employed for georeferencing and further rectification of all the UAS images.

Next, the data was clipped to a common area of interest (AoI) and resampled with GDAL, and the cubic convolution method to 0.16 m to enhance processing time and increased reliability of image correlation. PlanetScope Satellite images were co-registered in Matlab relative to a reference image (https://gitlab.lrz.de/tobi.koch/satelliteregistration.git). We used digital image correlation (DIC) to measure the displacement for the active landslide body of the Sattelkar and to assess the suitability of the PlanetScope and UAS data. This method employs optical and elevation data and calculates the distance between an image pair, based on the spatial distance of highest correlation peaks between an initial search and final reference window location changes of common pixels. The result provides displacement and ground deformation in 2 D on a sub-pixel level. COSI-Corr (Co-registration of Optically Sensing Images and Correlation), a widely used software in landside and earthquake studies was used for sub-pixel image correlation (Stumpf, 2013; Lacroix et al., 2015; Rosu et al., 2015;...
Bozzano et al., 2018). COSI-Corr is an open source software add-on developed by CALTECH (Leprince et al., 2007), for ENVI classic. There are two correlators, in the frequency domain based on FFT algorithm (Fast Fourier Transformation) and a statistical one. Applying the more accurate fractional correlator engine, recommended for optical images, different parameter combinations of window sizes, direction step sizes and robustness iterations were tested. Parameter settings include the initial window size for the estimation of the pixelwise displacement between the images and the final window size for subpixel displacement computation in x, y; a direction step in x, y between the sliding windows; and several robustness iterations (Table 6). There are two correlators, in the frequency domain based on FFT (Fast Fourier Transformation) and a statistical one. Applying FFT, different parameter combinations of window sizes, direction step sizes and robustness iterations were tested. We utilised recommended window sizes as suggested by Leprince et al. (2007) and Bickel et al. (2018). Step size one showed good results while keeping the original spatial resolution for the output; robustness iterations of two to four were sufficient for our purposes. Initial and final window sizes were systematically tested (see Table 6). For computing a state-of-the-art powerstation was employed (AMD Ryzen 9 3950X 16-core processor, 3.70 GHz, 128 GB RAM).

### Table 6 COSI-Corr input parameters for intervals of UAS and PlanetScope

<table>
<thead>
<tr>
<th>Sensor Resolution</th>
<th>Input interval</th>
<th>Initial window [pix]</th>
<th>Final window [pix]</th>
<th>Robustness iteration</th>
<th>Step size</th>
</tr>
</thead>
<tbody>
<tr>
<td>UAS [0.16 m]</td>
<td>I: 13.07.2019–24.07.2019</td>
<td>128x128</td>
<td>32x32</td>
<td>2</td>
<td>1x1</td>
</tr>
<tr>
<td></td>
<td>II: 24.07.2019–04.09.2019</td>
<td>128x128</td>
<td>32x32</td>
<td>2</td>
<td>1x1</td>
</tr>
<tr>
<td>UAS [3.0 m]</td>
<td>I: 13.07.2019–24.07.2019</td>
<td>32x32</td>
<td>16x16</td>
<td>2</td>
<td>1x1</td>
</tr>
<tr>
<td></td>
<td>II: 24.07.2019–04.09.2019</td>
<td>32x32</td>
<td>16x16</td>
<td>2</td>
<td>1x1</td>
</tr>
<tr>
<td>PlanetScope [3.0 m]</td>
<td>I: 19.07.2018–24.07.2019</td>
<td>64x64</td>
<td>32x32</td>
<td>4</td>
<td>1x1</td>
</tr>
<tr>
<td></td>
<td>II: 24.07.2019–04.09.2019</td>
<td>64x64</td>
<td>32x32</td>
<td>4</td>
<td>1x1</td>
</tr>
</tbody>
</table>

The results of each correlation computation returns a signal-to-noise ratio map (SNR) and displacement fields in east–west and north–south directions. These results signal-to-noise ratio (SNR), east–west and north–south displacements were exported from ENVI classic as GTiff and the total displacement was then calculated with QGIS.

Additional analyses were performed to estimate the DIC outputs of both, the UAS orthophotos and PlanetScope satellite imagery. Visual tracking of 36 single blocks, identifiable in the UAS orthophoto series allowed deriving direction and amount of movement; this supported the verification process of the total displacement. We employed this approach for the time interval I. In order to assess the information value and validity of the satellite imagery, UAS orthophotos were downsampled to 3 m (cubic convolution) for comparison purposes prior image correlation.

### 5. Results

In Sect. 5.1, we present ground motion results from DIC for the original input resolution for i) UAS, 0.16 m and ii) PlanetScope, 3 m input resolution based on parameters in Table 6. Second, for iii) DIC results of UAS downsampled to 3 m and of PlanetScope are compared. In Sect. 5.2. DIC results for UAS, 0.16 m are analysed with regard to displacement of visual single block tracking. Finally, in Sect. 5.3. required times for 

### Kommentar [DH58]:
Answer to comment by RC1 (J. Blöthe) in response
RC1 (J. Blöthe) wrote: L304/305: I guess this is only relevant if you explicitly mention the image-processing times.

### Kommentar [DH59]:
Modified following comment by RC1 (J. Blöthe) on Table 3
RC1 (J. Blöthe) wrote: Table 3: Here you use a different date format than in the text.

### Kommentar [DH60]:
Modified following comment by RC1 (J. Blöthe)
RC1 (J. Blöthe) wrote: L303/304: What is the uncertainty of these east–west and north–south displacement estimates? Did you check whether the bearing of the displacement matches the general slope of the Sattelkar?

### Kommentar [DH61]:
Answer to comment by RC1 (J. Blöthe) in response
RC1 (J. Blöthe) wrote: L307/308 and L440/442: This seems a bit arbitrary. How did you determine a cutoff value for displacement? How did you distinguish outliers from non-outliers? What is the confidence of your estimates?

### Kommentar [DH62]:
Answer to comment by RC1 (J. Blöthe)
RC1 (J. Blöthe) wrote: L306/309: This contradicts the descriptions of Fig. 5a, where you point out that "ambiguous, small-scale patterns with highly variable displacement rates" L322/333 dominate the western part of the mass movement.

### Kommentar [DH63]:
Answer to comment by RC1 (J. Blöthe) in response
RC1 (J. Blöthe) wrote: L311/312: I am not convinced that manually tracking boulders in the same images that were used for image correlation can verify the results of this correlation. You can use these data to check if manual and automated tracking give consistent results. Comparing manually tracked boulders from UAS imagery could however be used to compare against the displacement estimates from satellite imagery.
5.1. Total displacement/Displacement Rates

Figure 5 Results of DIC total displacement derived from UAS orthophotos at 0.16 m resolution for time intervals I and II (see Table 6). Apart from several minor displacement patches, no motion is visible outside the active body in either period. Time interval I (376 d) (Fig. 5a) shows mean displacement values from 6 to 14 m for a coherent area in the eastern half of the lobe from the centre (c) to the eastern boundary of the active area. The highest displacement rates (up to 20 m) are observed within small high-velocity clusters in the northwest sector (d). Lower velocities occur along the southern boundary (e, f), ranging from zero to 6 m with smooth transitions. Ambiguous, small-scale patterns with highly variable displacement rates are present in the western half (a) and along the northern boundary (b). No motion is detected along the western fringe (i.e., at the landslide head) which is 20 m in width. South of the landslide (g) there is a small patch of minor displacement with continuous (up to 3.5 m) and ambiguous signals. Furthermore, we observed small-scale patterns of ambiguous signals in the east (j) and in the west of the active area in the drainage channels (h, i).

Time interval II (42 d) (Fig. 5b) shows great similarity to time interval I with ambiguous signals in the same areas such as the drainage channels (h, i) and within the western half of the active area (b). In contrast to interval I (Fig. 5a), within the active area a homogenous higher velocity patch (up to 6 m) near the landslide head is evident (a). In the eastern half large homogenous...
patches extend from the landslide centre (c) to the root zone (d) showing coherent displacement values of zero to 4 m. During this shorter time interval II, no displacement is detected along the south eastern boundary (e) and for large parts of the root zone (f) previously covered in I. Similar to I, the landslide head has a 20 m rim free of signal (also see Fig. 6 x, y). In the central part of the lobe (c) total displacement values are significantly reduced.

Figure 5c and Fig. 5d demonstrate total displacement for similar time intervals to UAS (see Table 3 and Fig. 4). For interval Ib (370 d) (Fig. 5c) wide fringes with no motion were detected around an actively moving core area, which consists of small-scale clusters with variable total displacement in the western part, coherent high velocities in the middle, and coherent low velocities east of this core area. Outside the landslide, northeast and immediately south (j), high-velocity patches are observed. In interval II (42 d) (Fig. 5d) the detected displacement is restricted to the western half of the (a) and shows the same significant fringes with no motion as in I. Compared to interval I the motion pattern of this core area is more homogeneous with increasing displacement towards the east. Outside the active area several patches show medium to high total displacement, the largest of which is located 300 m northwest of the landslide (i).

In Fig. 5e and Fig. 5f displacement results of UAS downsampled to 3 m are compared to PlanetScope at 3 m for both time intervals, I and II. Overall, the results demonstrate high total displacements (~ 18–20 m) across the entire landslide interrupted by scattered speckles of low to medium total displacement (a, b, c). No motion was present in a fringe zone along the landslide front (west boundary), similar to results in Fig. 5a and Fig. 5b. In general, the displacement patterns are less smooth than at 0.16 m input resolution. Outside the landslide significant displacements exist at the eastern image border (Fig. 5e) and towards the west (h, i) (Fig. 5f). In comparison, total displacement values derived from PlanetScope cover in large parts the active area for Ib (Fig. 5c); however, for II only the core area of the landslide shows displacement. In both results the core areas of the landslide are surrounded by wide fringes with no data zero deformation.

5.2. Single Block Tracking

Figure 6 illustrates the total displacement values derived from the UAS data at high resolution (0.16 m) for interval II (42 d). UAS orthomages were used to manually measure single block displacement for 36 clearly identifiable boulders on the landslide surface. Block displacements of 1 m are visible in the eastern part (f), whereas DIC does not reveal any displacement below 1 m. Boulder tracks longer than 2 m in the central and western part of the landslide are reflected by DIC-derived displacement values. Near the front a 6 m displacement of one block (a) is represented in the DIC result. The highest values (6 m, 10 m, 16 m) were observed in regions where DIC delivered ambiguous, small-scale patterns of highly variable displacements.
Figure 6 (a) Displacement derived from UAS data at 0.16 m resolution for interval II (24.07.2019–04.09.2019, 42 d) combined with boulder trajectories (in metres) manually measured in the UAS orthophotos in the same time period. The solid black line represents the boundary of the active landslide based on field mapping. Background: UAS hillshade, 24.07.2019 (0.08 m), orientation -3° from north, UAS orthophotos at 0.16 m resolution for the master (b) and slave image (c) of the corresponding time interval.

5.3. Time required for collection, processing and evaluation

In Sect. 2 we introduced a novel concept to extend lead time, consisting of three phases within the warning time window (see Fig. 1Figure 1). This concept is based on DIC results, thus every step comprised in each phase has been previously undertaken.

On this basis, knowledge of required time for a further process iteration of the three phases is given.

Time required for collection, processing and evaluation of UAS and PlanetScope data are estimated and summed in Fig. 7. PlanetLabs specifies 12 hours from image acquisition to the provision in the data hub, which includes to a large amount data pre-processing (Planet Labs, 2020b). Adding two hours for the selection, order and download process, we assume that time required for the collection phase is approximately the same for both sensors, with 14 hours for PlanetScope and 12 hours for UAS. With regard to the time needed for the processing phase, the sensors differ with UAS requiring 17 hours and PlanetScope...
five hours. Time for the evaluation phase is estimated to be about two hours. In sum, $t_{\text{warning}}$ for UAS is approximately 31 hours compared to 21 hours for PlanetScope.

6 Discussion

To systematically analyse the predictive power of the UAS and PlanetScope data, we will (i) evaluate error sources and output performance, (ii) assess obtainable temporal and spatial resolution and (iii) derive a systemic estimate of the minimum predictable power of the UAS and PlanetScope data.

6.1. Error sources and output performance

To evaluate error sources and output performance, we compared results of digital image correlation results from optical data with (i) mapped mass movement boundary, (ii) visual block tracking for UAS and (iii) 3 m downscaled UAS orthophotos. The approximately one year evaluation period encompassed all seasons, hence freezing/thawing conditions and a wide range of meteorological influences, e.g. thunderstorms and heavy rainfall, are included. The two investigated time intervals are I/Ib and II, covering 376/370 days and 42 days (typical high-alpine summer season), respectively (Fig. 4). Interval II exclusively covers (high-alpine) summer conditions, with negligible to no contribution from freezing conditions. As these inclusion periods are inconsistent, the amount of total displacement cannot be directly compared; however the relative motion patterns can be compared. Accordingly, we can confirm the suggested parameter settings of earlier studies on window sizes, steps and robustness settings (Ayoub et al., 2009; Bickel et al., 2018).

In terms of the mass movement boundary, the total displacement derived from the DIC of the UAS data generally matches the field–mapped landslide boundary for both intervals (I, II) (Fig. 5a, b), and is supported by the absence of significant noise outside the AoI. Mapped boulder trajectories for interval II (see Fig. 6) are consistent with the calculated total displacement and thus confirm COSI-Corr as a reliable DIC tool to derive ground motion for this study site and UAS orthophotos as suitable input data. Nevertheless, there are several areas with ambiguous signals. Leprince (2008) describes snow cover, vegetation cover and alluvial processes, among others, as potential explanations for these decorrelations. In our study, the decorrelated
areas include to a large degree the landslide head (a), the drainage channel (b) (Fig. 5a, b), a larger patch south of the active area boundary (g) (Fig. 5a), and some smaller ones in little depressions (g) (Fig. 5a) and (j) (Fig. 5a, b). Most patches are identified as snow fields in the orthophotos and the noise results from decorrelation. In Fig. 5a, the large southern patch (g) shows clear displacement values for the rear part and decorrelation for the front region resulting from significant morphological/temporal changes within an the image pair of interval 1 limiting the ability to measure ground displacements. 

The decorrelation in the drainage channel (b) could stem from massive changes in pixel values, similar to the decorrelation on the basis of alluvial processes, as described by Leprince et al. (2007). Decorrelations in the areas with the fastest ground motions also lead to high pixel changes (Stumpf et al., 2016). These are observable in the active landslide area within the lobe, where large areas of decorrelation may be explained by high displacements in the leading part (a) with redetected, hence correlated pixels in the trailing part (c, d, e, f). These findings can be transferred to the landslide interior area (a, b), the frontal western regions and the northern margin (b). The observation is confirmed by geomorphological mapping and measured boulder block trajectories from the orthophotos (Fig. 6). Several patches of correlation (c, f) with corresponding boulder trajectories up to 4 m (34.8 m yr$^{-1}$) (d) are detected in the rear part. A correlated patch with a 16 m (34.8 m yr$^{-1}$) trajectory (a) is in flow direction behind the foremost boulder. In this case the method was able to capture the displacement partially as the distinct boulder block supported the detection, hence correlation. This allows us to conclude that displacements exceeding approximately 10 m (88.9 m yr$^{-1}$) for the calculated time period, thus 63 pixels or more at a resolution of 0.16 m, are definitely outside of a possible correlation and no pixel matching is possible. With a correlation window smaller than the displacement, the algorithm is not able to capture the displacement (Stumpf et al., 2016). Field observations provide evidence that the surface alters due to the high mobility and rotational behaviour of some boulder blocks, which leads to changed pixel values and spectral characteristics. Similar results were observed by Luceier et al. (2014), who described a loss of recognisable surface patterns if revolving and rotational displacements occur, causing decorrelation and a noise as output. These results show that with COSI-Corr and UAS orthophotos of 0.16 m, it is possible to detect the total displacement of the landslide in both extent and internal process behaviour even in this steep, heterogeneous terrain. Nevertheless, high displacement rates and rotational surface behaviour in the cirque limit the DIC method. A decrease of the time interval for this particularly highly mobile study site would likely reveal an enhanced correlation since for shorter time periods the total displacement decreases, and surface changes are reduced, which can be controlled by shortening the temporal baseline.

6.2. Comparison of temporal and spatial resolution

We compared the COSI-Corr total displacement results of PlanetScope (Ib and II, Fig. 5c, d) and UAS images (I and II, Fig. 5a, b) for the same time periods at different spatial resolutions (see Table 6). For the PlanetScope DIC result the main part of the landslide is detected, and its area is generally consistent with the results of the UAS DIC, which is additionally confirmed by boulder trajectories. The frontal part (a) reveals correlation signals (I and II); while for the same time intervals and parts, the UAS DIC results show a decorrelation (Ib and II). The correlation is likely to be attributable to the coarser spatial resolution of 3 m input data, hence a smaller number of pixels to be captured at this site with the DIC method. Similar texture of rock clast surfaces could lead to false positives resulting in correlation as patches appear similar in matching windows. However, in contrast to the UAS result (Fig. 5a, b), the outcome on a large scale fails to detect the entire active area (b), (f) as well as its internal motion behaviour. Nevertheless, for the visualisation and analysis of the PlanetScope results, the range of total displacements had to be restricted to values equal to and greater than 4 m due to noise and outliers over large areas, as applied and described by Bontemps et al. (2018). Even then, noise and several misrepresented displacement patches are observed for (i, j) and in the northeast image corner (Fig. 5). We can identify several reasons for these large clusters of high motion values. 

Massive cloud and snow coverage hampered both first images of interval Ib (19.07.2018) (Fig. 5c) and II (24.07.2019) (Fig. 5d), leading to a 20 m fringe of false displacements in the north–eastern part of the image. Minor snow fields are visible in the images from 24.07.2019 for both, the UAS and PlanetScope, could likely explain the big cluster of incorrect displacement
south east of the lobe (j); nonetheless, in the satellite image they are smaller than the resulting DIC displacement. High cloud coverage in two input images with large areas of white pixels may exert an influence leading to high gains due to sensor saturation (Leprince, 2008). Illumination changes in interval II (Fig. 5d) may cause unrealistic displacements outside the boundary with slightly darker colours due to shadows in the first satellite image (24.07.2019) and large parts within the second image (04.09.2019) are also in the shade. A comparison of the acquisition times and true sun zenith, e. g. for the second image, reveals a difference of 01:34 h between the image acquisition at 11:36 LT (local time) and the true local solar time at 13:10 LT. As the study site is located in a high-alpine terrain with a west facing cirque, at this time of day there are shadows of considerable length which have a significant influence on the result of digital image correlations. One clear advantage of the UAS images is that their acquisition is plannable according to the best illumination conditions with the sun at its zenith. Moreover, the UAS flight path as well as the system itself remained the same for all three acquisitions, while PlanetScope employs various satellites. Despite similar input resolutions and time intervals (Ib vs. I and II vs. II, see Table 3) with different sensors (UAS, PlanetScope), considerably divergent DIC outputs (Fig. 5c vs. e, d vs. f) are returned. To a large degree the active ground motion inside the mapped landslide boundary is represented by the 3 m UAS DIC result, while the same fringe remains free of signal for both UAS DIC results at different input resolutions [Fig. 5(e) and (f)]. This similarity with overall good agreement indicates that the displacement is restricted to a smaller area than the previously demarcated boundary, based on our field investigations. The satellite image detects large parts of the main active core area but widths of 50–80 m from the boundary show no displacement. False displacement is indicated for a cluster outside of the boundary to the image border in the east for UAS interval I (Fig. 5e) and in the north western area (h, i) for interval II [Fig. 5(e) and (f)]. Contributing to changes in shading and illumination, apart from these false signals, there is minor noise compared to false large clusters of high displacement within the PlanetScope result interval I for (j) and northeast image corner (Fig. 5c) and interval II (i) (Fig. 5d).

However, two striking differences with correlation/decorrelation and ground motion values are observed for the two UAS input resolutions; the coarser resolution of 3 m returns a correlation signal with values typically exceeding 18 m of displacement as the value range is extended, due to previous high factor downsampling. Measured ground motion of block tracking and PlanetScope results indicate and support existing high ground motions. This observation might be the explanation for the observed decorrelation at the finer resolution of 0.16 m for the landslide head. For this reason, the previous assumption using a shorter time interval leading to improved detection of inherent process behaviour (see Sect. 6.1.), can be complemented with a coarser resolution showing a clear improvement in the form of better correlations and returned signals. Generally, with high resolution images, such as UAS, we recommend first calculating displacements based on a coarser input resolution (1–3 m) to examine the overall situation and detect changes; and second to calculate displacements at a finer resolution in order to focus on relevant details of the Aoi. With regard to PlanetScope data, a 3 m resolution seems to be in a good spatial range to assess ground displacements even of this steep and heterogeneous study site with its high motion. Nonetheless, constraints such as illumination due to early daytime acquisitions leading to shadows, meteorological influences by clouds, cloud shadows and snow decrease the quality of the satellite images and reduce their applicability. Sensor saturation, shadow length, size and direction as well as changes in snow, cloud or vegetation cover impose limitations (Delacourt et al., 2007; Leprince et al., 2008) and accord with our observations. The authors identify additional limitations such as radiometric noise, sensor aliasing, man-made changes and co-registration errors (ibid.). All these limitations have a negative impact on the input image, which leads to impaired DIC calculations and results, and partially or wholly inaccurate analysis of the displacement. These might have played a role in our results. In our experience, the usability of the DIC result may be influenced by the input image quality. This restricts the application of PlanetScope images to a certain degree. They can be employed as input data to detect displacements, but as there are in the present setting too many signals of false-positive displacements, which can solely be discarded on the basis of field evidence, this data is currently of limited use. It should be handled with caution and we recommend, and combining it with complementary data and ground truth is recommended.

Kommmentiert [DH92]: Modified following comment by RC1 (J. Blöthe)

RC1 (J. Blöthe) wrote: L445/447: The size of the snow patches does not play an important role. The presence of snow in one image hampers correlation between images and leads to false patch/matching results.

Kommmentiert [DH93]: Modified following comment by RC1 (J. Blöthe)

RC1 (J. Blöthe) wrote: L460/461: Check figure reference.

Kommmentiert [DH94]: Answer to comment by RC1 (J. Blöthe) in response

RC1 (J. Blöthe) wrote: L473/474: To be frank, I do not see much similarity between Fig. 5 (i) and (n) (i) and (j). I would be very cautious in interpreting these results as is. This is especially true for the resampled UAS results.

Kommmentiert [DH95]: Correction by the authors.

Kommmentiert [DH96]: Modified in response by RC1 (J. Blöthe)

RC1 (J. Blöthe) wrote: L486/487: As for referencing the UAS data are probably located close to the landslide, it is not surprising, but neither disturbing, that false displacement clusters appear outside the area of interest.

Kommmentiert [DH97]: Answer to comment by RC1 (J. Blöthe) in response

RC1 (J. Blöthe) wrote: L489/490: Again, I would not trust the displacement estimates of the resampled UAS data. While it is true that your manual boulder tracking identified 2 boulders with displacement of 10 or more meters, the remaining 14 boulders show something different.

Kommmentiert [DH98]: Answer to comment by RC1 (J. Blöthe) in response

RC1 (J. Blöthe) wrote: L491/492: While it might be true that the results obtained from image correlation of resampled 3m UAS data are better (internally) correlated and show a more homogeneous deformation pattern, this does not mean that the result is correct. As I outlined above, I have serious doubts regarding the interpretability of this data, as there is no agreement with the manually tracked boulder velocities (except 2 boulders). Also, from a geomorphic perspective, I am not sure how you would explain a velocity pattern where high velocities dominate throughout the entire landslide, but are speckled with lower to zero movement within (Fig. 5 e and f).

Kommmentiert [DH99]: Modified and restructured this sentence to emphasise our recommendation following comment by RC1 (J. Blöthe). In addition, please see answer in response.

RC1 (J. Blöthe) wrote: L504/505: Did you evaluate the proportion of false-positive displacements to true positive displacements and if so, how did you do this and can you please include this data? Based on the image correlation results shown here, you can make this statement, but I would be cautious to make a general claim on the usability of the data.
6.3. Estimating time to warning

Early warning is essentially defined as being earlier than the event and thus puts high external time constraints on observation and decision. The time window between the detection of an accelerating movement preparing for final failure and the final failure itself is determined by the environment. Therefore, two sensors with the highest available spatiotemporal resolution were evaluated and compared with regard to their applicability to the early warning of landslides. We made rough assumptions and assessed the time needed for the phases of time (i) to collect, (ii) to process, and (iii) to evaluate relevant data (summarised in the time in warning window, see Fig. 7).

Despite different underlying technologies the time required for the collection phase is approximately the same for both sensors. For UAS, we estimated about 12 hours under ideal circumstances, while for PlanetScope 12 hours (Planet Labs, 2020b) plus two hours for image selection, download and initial analysis, adding up to 14 hours in total (see Sect. 5.3.). In the second phase, time to process, deriving orthophotos from raw UAS images is time consuming. The subsequent DIC calculations demand significantly more processing time for the UAS images than for lower resolution PlanetScope images. The final phase, time to deliver, takes about two hours for each sensor. In our case study, the estimated time to warning (\(t_{wa44m}\)) was 10 h longer for the UAS approach (31 h) in comparison to the Planet Scope approach (21 h). These time calculations are based on ideal environmental conditions and data availability. Assuming good conditions exist to conduct the UAS flight and no constraints limit the utilisation of satellite images, in theory a daily deployment is possible. In reality, unfavourable weather conditions, cloud and snow cover as well as limited data availability will increase the actual \(t_{wa44m}\) significantly. From the available images in the Planet Data hub (besides other exclusions) meteorological influences reduced for April–October 2019 the usability by 14.5 % and 7 % for cloud cover and snow cover, respectively (Table 2). The flexibility of a UAS can serve as a practical remote sensing tool for the investigation of ground motion behaviour in a spatiotemporal context. Nonetheless, weather influences can make a UAS flight impossible or impractical as the result might be useless. Depending on the level of illumination, the same may apply for satellite images. Regardless of any meteorological constraints, the promised daily availability by PlanetScope is unrealistic, due to data gaps and provider issues, our study showed that for the Sattelkar from April to October 2019 only 11 % of the captured images during this time were usable. In time–critical early warning scenarios, when time is running out, all available even partly usable images will be utilised and fieldwork may be conducted, even if the prevailing conditions are suboptimal but will increase data availability. The comparison of two selected remote sensing options demonstrates that the comprehensive knowledge on the available remote sensing data sources and their respective time requirements can substantially reduce the time to warning (\(t_{wa44m}\)) and to extend the lead time (\(t_{wa44d}\)).

Significant observations of the temporal evolution of historic landslides are presented in Table 7 and described below. These include (i) the Preonzo rock slope failure, CH (Satellite et al., 2016; Loew et al., 2017), (ii) the Vajont rock slide, ITA (Petley and Petley, 2006) and (iii) the Sattelkar complex slide, AUT (Anker et al., 2016). These landslides have specific evolution histories, e.g. early observed crack developments, increased movement and minor events like Preonzo (2002 and 2010) (Sattel et al., 2016); Sattelkar, with large volume mass wasting processes since 2005 and a debris slide event in 2014 (see Sect. 3 Study Site) (Anker et al., 2016); and Vajont, with ductile failures in 1960 and 1962 and a transition from ductile to brittle behaviour in 1963 (Petley and Petley, 2006; Barla and Paronuzzi, 2013).

Table 7 Relevant dates for historic failures of Vajont (ITA), Preonzo (CH) and Sattelkar (AUT). Time period in italics–bold used for Fig. 9. Time intervals in days (~ for rough estimations) and years in square brackets; sum of days based on the first day of the month, if only month as reference is available from literature (Petley and Petley, 2006; Anker et al., 2016; Satellite et al., 2016; Loew et al., 2017). Further explanation below.

<table>
<thead>
<tr>
<th>Vajont [days/year]</th>
<th>Preonzo [days/year]</th>
<th>Sattelkar [days/year]</th>
</tr>
</thead>
</table>

Komentiert [HD100]: Anmerkung zu den von RCI (J. Blüthe) in Beantwortung: L552/554/Table 7 / Figure 9: “I do like the idea behind this, where the authors show that their proposed workflow would enable a timely warning in the case of historic landslides. However, in the case of Vajont, I think you should include a critical factor. While it is theoretically true that a “forecasting window” would allow for your workflow to be completed well before the failure, the slow deformation of Vajont (35 m d-1) in the 30 days will be well below the level of detection of your image correlation analysis, if you collect an image directly after the onset of “significant acceleration”. In order to be detectable, movement must have accumulated a critical distance before data collection of your workflow can set in (30 days = 1.05 m total displacement) – a factor that I believe would be important to include here.”
Figure 9 is the extension of our concept (see Sect. 1, Fig. 1) systematically supplemented with our estimated time to warning (UAS, PlanetScope), and compared to the few data series predating larger slope failures.

Following a significant acceleration, the forecasting window is opened and $t_{\text{warning}}$ starts, which is composed of phases (i) time to collect, (ii) time to process and (iii) time to evaluate. To ascertain a significant acceleration one further observation is required. Hence, one complete cycle of the three phases, previous analyses and processing iterations are given. Our analysis showed that UAS and Planet Scope can approach times as short as 31/21 h, as a result $t_{\text{lead}}$ is increased and so is $t_{\text{react}}$.

Assuming both sensors reliably estimate ground motion, solely based on their time requirement, this concept was applied to the temporal development of historic landslide events, thus from measured increased displacements and/or massive accelerations to the final event (Table 7). On this basis we simplified the graph and what we defined as “significant acceleration” using dates of observations such as increased crack opening (Vajont), critical displacement (Preonzo) and the beginning of active ground motion (Sattelkar). Therefore, the opening of $t_{\text{warning}}$ and forecasting window are concrete observations of the particular site, independent of any intensity described by the corresponding authors and allows more freedom for temporal evaluations without going into details.

For the Preonzo case, the entire 2012 spring period was characterised by high displacement rates. We defined the first of May 2012, when geologists operating the warning system informed local authorities and assembled a crisis team, as the onset or ‘increased movement’ and the 15.05.2012 with 300 000 m³ as the impact (Sättele et al., 2016), in total approximately 15 days. For Vajont, the 1/velocity plot by Petley and Petley (2006) (based on data from Semenza and Ghirotti (2000)) shows an increase in movement at about day 60 along with a transition from a linear to an asymptotic trend at approximately day 30, defined as a transition from ductile to brittle. Therefore, we assumed 30 days of forecasting window for $t_{\text{warning}}$ and $t_{\text{lead}}$ until the impact of the hazardous event on 09.10.1963. However, it has to be kept in mind that velocities of about 35 mm d⁻¹ are still low and at the minimum of the displacement recognition capability for the digital image correlation method.

For the Sattelkar...
site, the observed mass displacement increase is presumed to have started in 2005 with the 170 000 m³ debris flow event on 31.07.2014 as the impact, thus about 3-498 days (Anker et al., 2016).

Even for the Peenmoor event, with its short forecasting window of 15 days, the ground motion assessment based on the evaluated optical remote sensing images, would have been possible under the assumption of reasonably good UAS flying conditions and the provision of usable PlanetScope images. For t_{\text{lead time}} there is enough temporal leeway to repeat at least three to four successive measurements comprising the three phases. However, as single accelerations are possible in very short time intervals of less than two days, it is impossible to capture these accelerations by means of optical remote sensing methods, given a time requirement of 31 hours for UAS and 21 hours for PlanetScope. Nevertheless, this comparison shows that for larger and long-preparing slope failures the technical lead time may well be shorter than the forecasting window starting at the time at which the process became predictable.

7 Conclusions and outlook

This paper presents an innovative concept to compare the lead time for landslide early warning with high-alpine sensing systems. We tested this temporal concept by applying UAS and PlanetScope images of temporal proximity as these are the sensors with the best spatiotemporal resolution. We assessed the sensors’ capability to identify hot spots and to recognize behaviour by delineating ground motion employing digital image correlation (DIC). In so doing, knowing the necessary processing time enabled us to estimate the time requirement and finally to incorporate it into the concept to evaluate sensors with regard to ongoing landslide processes of the Sattelkar as well as historic landslide events. Our findings derived from DIC for this high-alpine case study show that high resolution UAS data (0.16 m) can be employed to identify and delimitate the main landslide process and reveal its heterogeneous motion behaviour as confirmed by single block tracking. Thus, validated total displacement ranges from 1–4 m and up to 14 m for 42 days. PlanetScope Ortho Scenes (3 m) can detect the displacement of the landslide central core, however, cannot accurately represent its extent and internal behaviour. The signal-to-noise ratio, including multiple false-positive displacements, complicates the detection of hotspots at least in this very steep and heterogeneous alpine terrain.

Coarse temporal data resolution, such as in the case study investigated here, represents an important restriction to the use of optical remote sensing data for landslide early warning applications. Acceleration (and the resulting failure) over short periods of time will likely go unnoticed due to large data acquisition intervals. However, for prolonged acceleration periods, such as observed at the Sattelkar slide and many other relevant hazardous sites, the chosen data sources have been demonstrated to represent a formidable early warning approach capable of contributing to an improved risk analysis and evaluation in steep high-alpine regions.

With regard to the temporal aspect for early warning purposes, PlanetScope satellite images require less time compared to UAS for the time phases of collection, processing and analysing. As a consequence, when time is of the essence, the UAS acquisition cannot compete with the high frequency of PlanetScope daily revisit rates. In general, both are limited in their use as they are passive optical sensors dependent on favourable weather conditions. Nevertheless, with a realistic 10% of usable data for our study site, PlanetScope cannot provide daily data as promised.

To conclude, in methodological terms DIC is a reliable tool to derive total displacement of gravitational mass movements even for steep terrain. Given the high reliability of UAS data, its temporal resolution is the key in future attempts to overcome decorrelation due to high ground motions. In addition, a slightly coarser resolution reduces the time needed for total processing, enhances correlation while maintaining spatial accuracy and reliability. PlanetScope is especially interesting as a complementary sensor when UAS employment is restricted e.g. inaccessible and/or dangerous sites or for areas too extensive to be covered. For continuous monitoring and early warning, the warning time window could be shortened by on-site drone ports with autonomous acquisition flights and automatic processing. Our systematic evaluation of the sensor capability of PlanetScope DATA

Kommentiert [DH102]: Author’s revision to be more specific.

Kommentiert [DH103]: Author’s revision in order to be more concise and in line with the introduction and study site section.

Kommentiert [DH104]: Author’s revision to be more clear.

Kommentiert [DH105]: Modified following comment by RC2 (S. Roessner):

RC2 (S. Roessner):
The total data sources (planet and UAV) do not allow optimization of lead time in the context of early warning because of the scarcity of their availability which is reflected in the small number of only three multitemporal data sets between July and September analyzed in this study (Table 3) and the study site (starting at L175) represents a very complex landslide case leading to rather erratic mass movements in form of debris flows initiated by changing slope water conditions related to increased atmospheric precipitation. This situation is another obstacle for an early warning approach which is solely based on optical remote sensing data and thus making it impossible to make full use of the in principle daily temporal resolution of the planet data. Taking into account these natural conditions and the constraints introduced by the used imaging constellations, leaves no room for true optimization of lead time in the sense as stated in the overall scientific goal of this paper.
can be applied and transferred to other optical remote sensing sensors, which extends the lead time. Future studies should focus on the applicability of complementary optical data to confirm the detection of landslide displacement and adjust UAS output resolution as this significantly increases the validity of DIC internal ground motion behaviour.
Data availability

PlanetScope data are not openly available as and PlanetLabs Inc. is a commercial company. However, scientific access schemes to these data exist.

Author contribution

Doris Hermle developed the study together with Markus Keuschnig and Michael Krautblatter, analysed the data and wrote the paper. Markus Keuschnig and Michael Krautblatter supported the writing and editing of the paper. Ingo Hartmeyer provided critical proof reading with valuable suggestions. Robert Delleske is responsible for UAS flight campaigns and processing the images.

Competing interests

The authors declare that they have no conflict of interest.

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