1   **CHILDA − Czech Historical Landslide Database**

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8   **Abstract**: National and regional historical landslide databases are increasingly viewed as providing empirical

9   evidence for the geomorphic effects of ongoing environmental change and for supporting adaptive territorial

10   planning. In this work, we present the design and current content of the Czech Historical Landslide Database

11   (CHILDA), the first of its kind for the territory of Czechia (the Czech Republic). We outline the CHILDA system,

12   its functionality and technical solution. The database was established by merging and extending the fragmented

13   regional datasets for highly landslide prone areas in Czechia. Currently, the database includes 699 records (619

14   landslides, 75 rockfalls, and 5 other movement types) encompassing the period from the oldest determined records

15   (1132) up to 1989 which represents an important cultural, political and socioeconomic divide. g with

16   characterizing the content of the database, we discuss its further developments and applications.

17   **1 Introduction**

18   Historical landslide inventories and databases are among the key challenges within landslide risk reduction efforts

19   as they fill the gap, on one hand, between the landslide occurrence in the past environments studied with the use

20   of various proxies and the present-day landslides, for which different monitoring and mapping techniques may be

21   used, on the other (e.g., Glade et al., 2001; Raška et al., 2015; Piacentini et al., 2018). In light of the recent global

22   climate change, the historical landslide databases contribute a better understanding to changes in various

23   meteorological triggers of landslides in different environmental settings (Gariano and Guzzetti, 2016; Caracciolo

24   et al., 2017). Given the severe impacts of landslides on society (Froude and Petley, 2018), these databases also

25   make it possible to identify changes in hot spots of landslide occurrence and the character of their impacts (Klose

26   et al., 2016). In both these respects, the historical databases complement current landslide inventories that can be

27   used to trace the spatial patterns in landslide occurrence and their causative factors (Van Den Eeckhaut and Hervás,

28   2012; Herrera et al., 2018; Marc et al., 2018). Within the landslide hazard and risk assessments, the historical

29   landslide databases provide evidence as it is generally assumed that past landslide occurrence frequencies may be

30   used to describe the probabilities of landslide occurrence in the near future (Remondo et al., 2008; Van Den

31   Eeckhaut et al., 2009; Wu and Yeh, 2020).

32   Connecting these directions, increasing attention has been also paid to revealing the vulnerabilities and adaptive

33   behaviours of past societies regarding landslides (Tropeano and Turconi, 2004; Caloiero et al., 2014; Klose et al.,

34   2016; Raška, 2019; Klimeš et al., 2020). These studies argue that historical landslide databases − if approached

35   critically − may inform current efforts for adaptive management of landslide risks (Klose et al., 2016; Raška and

36   Dubišar, 2017). Historical landslide databases have been recently established for various countries and regions,

37   for instance, in Italy (Guzzetti et al., 1994; Piacentini et al., 2018), Nicaragua (Devoli et al., 2007), USA (Elliott

38   and Kirschbaum, 2007), Norway (Hermanns et al., 2013), the United Kingdom (Taylor et al., 2015), Germany

39   (Damm and Klose, 2015), Portugal (Pereira et al., 2014), most of them covering ca. the last 150 years, but some

40   databases also including scarce records dating back as early as the twelfth century.

41   The aim of this work is to present the Czech Historical Landslide Database (CHILDA), a project that unified and

42   extended the fragmented existing regional databases and established an open access and concurrently updated

43   map inventory of historical landslides in Czechia. The presented database thus further fills in the gap of missing

44   historical databases for Central-European mid-mountain environments (Damm and Klose, 2015). Within this

45   paper, *landsliding* is used as a generic term covering all major types of rapid mass movements (cf. Hungr et al.,

46   2014) that are usually recorded in the documentary data. Slow slope deformations are not studied here since they

47   usually did not cause rapid harm to society and have not been registered by past societies. For CHILDA content,

48   we only differentiate the three following groups according to major mechanism: (a) landslides sensu stricto (also

49   including spreading and flows), (b) rock falls (including topples) in ~~solid~~ bedrock, while (c) all remaining mass

50   movements are grouped as 'others'. This rough classification is used since the documentary data often do not

51   allow for detailed and reliable identification of the mass movement type.

52   In the following sections, we will first review the previous studies on historical landslides in Czechia with

53   emphasis given to attempts to establish systematic historical landslide databases. We will then outline the

54   availability of the documentary sources and present a design of the CHILDA. Finally, the current content of this

55   database and its completeness will be presented in a comparative perspective along with discussion of its future

56   directions. Although CHILDA is an open database, the last analysed year was set to 1989 for the purpose of this

57   study. The year is considered an important cultural, political and socioeconomic divide in the recent history of

58   Czechia, turning the country into a democratic regime and thus influencing production, diversity and

59   accessibility of documentary data and the data from ongoing landslide monitoring.

60   **2. Landslides in Czechia**

61   **2.1 Landslide Predispositions in Czechia**

62

63   Despite the fact that Czechia can be generally considered a low-risk country, given the relatively low landslide

64   frequencies and impacts (Klimeš et al.,   2017), the country displays high spatial variability in landslide

65   occurrence with some highly landslide-prone regions due to their predisposition and presence of causative

66   factors. Among the most affected by landslides are the Outer Western Carpathians (OWC), NW Czechia (České

67   středohoří Mts., Děčínská vrchovina) and several of the scattered spatially limited areas across the country (see

68   Figure 1).

69   The eastern part of Czechia, OWC, is built up of Tertiary and Mesozoic flysch rocks which are particularly

70   susceptible to landsliding. As a consequence, for example, of the 1997 landslide period as many as 3,700

71 individual landslides were mapped in that region (Krejčí et al., 2002). High numbers of reactivated landslides

72 were also further identified during the periods of intense landsliding which followed, specifically in 2006 (Bíl

73 and Müller, 2008), and 2010 (Pánek et al., 2011). The Registry of Slope Deformations of the Czech Geological

74 Survey (www.geology.cz) contains in all approximately 14,500 landslides in this area of the Czech part of the

75 OWC (7,200 km2), which was 82 % of all the landslides registered within Czechia (Bíl et al., 2016).

76 In NW Czechia, two major areas display high landslide susceptibility, namely the České středohoří Mts. (CS)

77 and the Děčínská vrchovina Highland (DV). The lithology of CS is built up of Neogene volcanic rocks,

78 including basalts and phonolites and their volcaniclastics, surrounded or underlain by Mezosoic weak

79 sedimentary rocks (Cajz, 1999). In a rugged terrain, such conditions often result in landsliding accelerated during

80 the snow-melt season and heavy rainfalls (Hroch et al., 2002; Raška et al., 2014a; CHMI, 2020). DC is formed in

81 an uplifted and dissected Mesozoic sandstones, which are prone to rockfalls and toppling (Kalvoda and Balatka,

82 1995; Zvelebil et al., 2005). These mainly occur in the tectonically predisposed and deeply eroded valley of the

83 Labe/Elbe River and its tributaries. Weathering, bioturbation as well as human alterations of the terrain act as the

84 most frequent causative factors.

85 Other parts of Czechia are not as susceptible to landsliding to the extent comparable to the above-mentioned

86 three primary landslide areas (OWC, CS, DV). Landslide activity has been long recorded in the Neogene and

87 Quaternary sediments of the Most basins (MB, see Figure 1), along the banks of the Ohře/Eger River and in the

88 anthropogenic landscape at the edges of the open-pit brown coal mines (e.g., Burda et al., 2013). Another area

89 prone to landslides is in central-eastern Czechia in the Mesozoic sandstones of the Bohemian Paradise (BP)

90 which form steep elevations and rockfall-prone areas (e.g., Forczek, 2008). Rockfalls occur along a number of

91 deeply incised valleys in the Bohemian Massif (e.g., along the Vltava/Moldau river valley south of Prague, VR)

92 as well as some transportation corridors, particularly along rail tracks (as documented in a database of road and

93 railway blockages due to natural processes, www.rupok.cz (Bíl et al., 2017)).

94

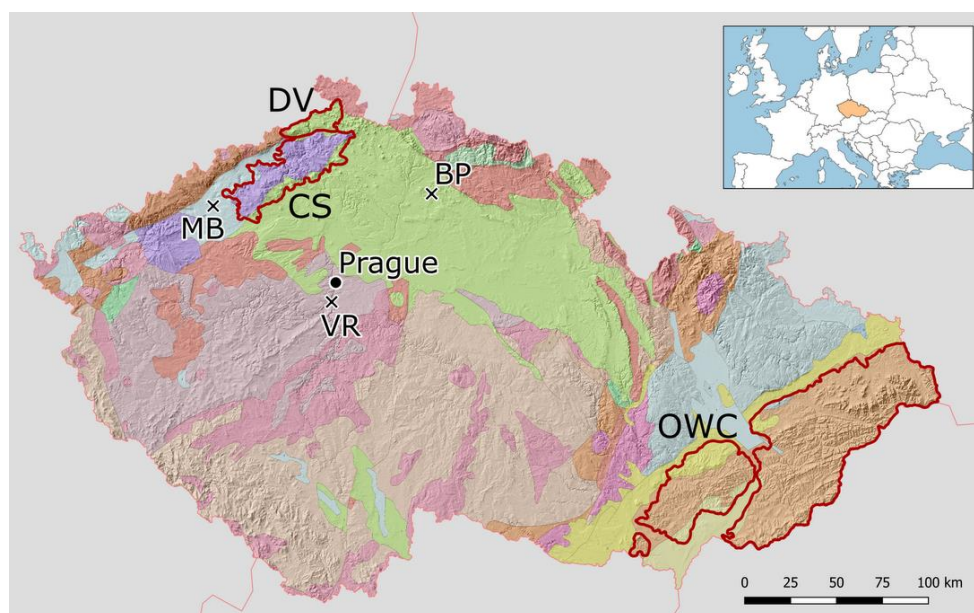



Figure 1: Delimitation of the primary areas where landsliding concentrates in Czechia. Neogene volcanic rocks
(CS – České středohoří Mts.), Mesozoic sandstones (DV – Děčínská vrchovina Mts.), Neogene and Quaternary
sediments (MB – Most Basin) on the west, Mesozoic sandstones in central parts of Czechia (BP – Bohemian
Paradise sandstones) and Mesozoic and Tertiary flysch belt (OWC – Outer Western Carpathians) in the east of
Czechia represent the most susceptible parts to landsliding. VR – concentration of rockfalls along the Vltava river
https://mapy.geology.cz/arcgis/rest/services/Inspire/GM2_5mil/MapServer. © Czech Geological Survey.

2.2 Historical Landslide Research in Czechia

2.2.1 The Beginning of Landslide Research in Czechia

The first protoscientific works that also included reports on landsliding in Czechia are dated to the eighteenth
century (e. g. Strnad, 1790), followed by geological and engineering geological scientific works emerging as of
the end of the nineteenth century (Zahálka, 1890; Dědina, 1896; Woldřich, 1899) and at the beginning of the
twentieth century (Čermák, 1912, Dědina, 1916). More systematic landslide research which described landsliding
in the area of modern Czechia started, however, in the 1920s with the work of Záruba (1922, 1923, 1926, 1938).
Particular attention was paid to a general description of historical landslide areas in the Pavlovské vrchy Hills
situated in OWC (Jüttner 1931, 1937) as well as historical landslides themselves in the same area (Stejskal, 1931;
Woldřich and Stejskal, 1934). Záruba and Myslivec (1942) documented landslides related to transportation
infrastructure in OWC. Landsliding in the broader area of the city of Zlín (OWC) was described by Krejčí (1943).
The first Czech modern landslide classification was published by Záruba and Mencl (1954).

115

2.2.2 Systematic Works Describing Landslide Occurrence Based on Historical Data

117   We present here an overview of works from Czechia which created at least a regional landslide chronology based
118   on documentary data analyses. The only systematic studies of landslide occurrence, based on a range of historical
119   sources in Czechia, were conducted by Špůrek (1967, 1972, 1985). These studies were mainly based on the
120   investigation of articles published in national newspapers covering the territory of former Czechoslovakia (and
121   also mentions landsliding all over the world). The recorded information includes each landslide date, location, type
122   and amount of damage as well as the bibliographic source.

123   There have not been any other attempts to prepare an overview of dated landslides for all of Czechia which
124   would follow up on Špůrek's work from 1970 onwards. Concerning debris avalanches, which are only located in
125   the highest parts of Czechia, Pilous (1973) presented their overview from the Krkonoše Giant Mnts. The efforts
126   have been re-established since the 2000s and have focused on the major Czech areas prone to landsliding.

127   For the Outer Western Carpathians, Bíl et al. (2014) studied historical landsliding in an area around the village of
128   Halenkovice (central part of OWC). They analysed documentary data, chronicles and interviewed eyewitnesses.
129   They determined dates for 120 individual landslides. The oldest records were found in local chronicles and
130   described landsliding in two villages (Jankovice and Košíky) in the close vicinity of Halenkovice in 1915. Bíl
131   al. (2020) created an overview of the chronology of landsliding in the Pavlovské vrchy Hills, an area at the
132   Czech-Austria border belonging to the Western Carpathian Flysch Belt. They determined dates for 30 historical
133   landslides. The first written resource relates to a landslide record dating back to 1663.

134   In NW Czechia, Raška et al. (2015) established a regional historical landslide database which was compared
135   with the central part of OWC in respect to the data availability and content. The multihazard database for the
136   latter half of the nineteenth century (Raška and Dubišar, 2017) allowed for an exploration of the relative direct
137   impacts of landslides on society. Finally, Raška (2019) used the landslide database to suggest five phases in the
138   evolution of community-based landslide risk reduction and the various approaches and mechanisms employed in
139   landslide mitigation measures.

140   Apart from these works, several studies have been published which used historical data to understand current
141   local landslide hazards (Krejčí et al., 2017) and which explored the availability of documentary sources (Kozák
142   and Rybář, 2003

143   **3. Design of the CHILDA Database**

144   **3.1 Data availability**

145   Czechia disposes of an extremely diverse and extensive range of documentary data that may be explored to build
146   historical landslide inventories. The number of these sources have been, however, subjected to academic scrutiny
147   mostly in historical climatology and hydrology (Kjeldsen et al., 2014; Brázdil et al., 2018a). Similarly, Bíl et al.
148   (2020) combined different documentary and archaeological data to compile a chronology of landsliding in the
149   Pavlovské vrchy Hills (Czechia, OWC) and described the basic historical landslides terminology. The conceptual
150   differences in hydrometeorological and geomorphologic hazards do not allow for uncritical transposition of the

151 climatological insights into the historical landslide research and therefore call for new insights into the potential

152 of the documentary data (Crozier and Glade, 1999; Raška et al., 2014b).

153 The present landslide database is based on a systematic data search in documentary data (both written and

154 iconographic) of several types with varying content, coverage and availability (Figure 2). Most of these data are

155 available in local archives in a printed or hand-written form (e.g., municipal and school chronicles, official

156 correspondence, photos), while some were found in private collections. The primary documentary data were also

157 complemented by a search in secondary (published) literature (e. g. newspapers, historical synthesis, historical

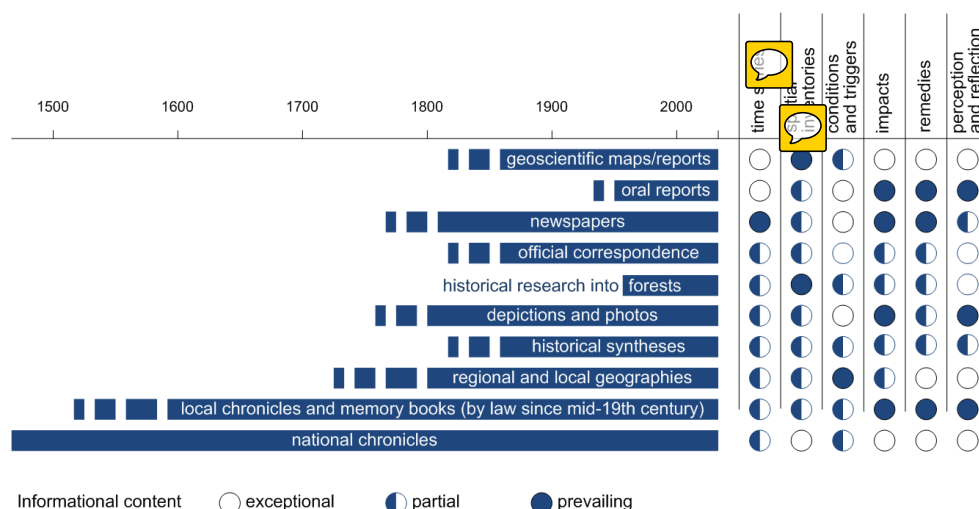
158 research into forests) and in some areas also with oral inquiry. Within the data search, the lower time boundary for

159 the database has not been set. The upper (recent) boundary, for the aim of the analysis in this work, was set at

160 1989, ~~representing a significant socio-economic and political turn in Czech history and thus marking the change~~

161 ~~in public data availability as well as in approaches to scientific inquiry regarding landslides~~. We have allowed,

162 however, the database to remain open in order to allow users to add new landslides as well.

163 Informational content ○ exceptional ◐ partial ● prevailing

164 *Figure 2 An overview of available documentary data applicable for historical landslide research in Czechia with*

165 *respect to their time coverage and informational content*

166

167

168   **3.2 Database Structure**

169   Attributes related to each database record are presented in Table 1. Some of the attributes are added by users via a

170   form. The items with an asterisk are mandatory and the items which are not part of the input form are underlined.

171   They are processed automatically, within the system.

172

173   *Table 1: Structure of records in Childa.*

Name *	Meaning	Values
ID	Unique identifier of a landslide	Number
Type *	Determination of kind of landsliding	Landslide; rockfall; earthflow; debris flow; human-induced landslide
Position *	Latitude and longitude of the record inserted via a click on the map	WGS 84 coordinates
Locality *	Description of the locality	String
Accuracy *	Describes the spatial precision of landslide localization by DB user (not a precision in an original source)	List: Metres; tens of metres; hundreds of metres; kilometres
MASL	Height above mean sea level - Landslide highest point elevation	Metres
Count	The number of landslides related to the particular location, given its accuracy; default value 1	String
Start *	The earliest possible date for the beginning of landsliding indicated by a record	Date
End *	The latest possible date for the beginning of landsliding as indicated by a record	Date
Period	Shows an interval during which the landslide originated. It is computed as End – Start.	String
Causes	Description of landslide cause, more causes can be selected	List: Earthquake; lithology; flooding; precipitation; mining; snow thaw; storm; artificial cause
Magnitude	Degree of landslide damage in three categories	List: 1 - Negligible or no losses; approx. 1ha or road infrastructure; 2 - hundreds of metres, >1ha; 3 - large volumes and areas, deep seated landslides
Impact	List of losses caused by landsliding, more impacts can be selected	List: Human fatality; human injury; buildings; transport infrastructure; other infrastructure (mine, water tower, utilities, etc); landscape including old mines, etc.
Remedies	Kind of remediation if applied	String

Source *	Full citation of the source of the landslide record	String
Details	Additional information and original data availability and accessibility (e.g., museum, archive, private collection, etc.)	String
Notes	Other relevant information about landsliding	String
Photo	More than one graphics file can be attributed to a record, e.g., photo, map, a copy of a written source, etc.	Graphics file, pdf

174 *Mandatory attributes.*

175 *Note: The attributes which are underlined are automatically added by the system and are not part of the user*

176 *form for data entering.*

177

178 **3.3 Web-map Application**

179 The database can be accessed through a web-map application at https://childa.cz (Figure 3). CHILDA is

180 administrated and hosted on the CDV – Transport Research Centre servers. The software requirement is as

181 follows: PostgreSql/PostGIS, php, php NetteFramework, HTML, CSS, JQuery.

182

183 *Figure 3: Title page of CHILDA (www.childa.cz).*

184 CHILDA users can select a background map (see Figure 4, A). There are the following possibilities: base map,

185 orthophoto, Open Street Map, Mapy.cz map and a hiking map. A municipality or other geographical feature with

186 a conventional name (mountains, etc.) can be selected via a form (B). The full-screen mode is launched when a

187 user clicks on the icon (C). Standard map control features such as full extent (D), backward and forward screen

188 (E), Google Street View (G) and zoom in/out (H) are also available here. A detailed landslide description is

189 shown in the pop-up window after clicking on the landslide map symbol (F). A new landslide record can be

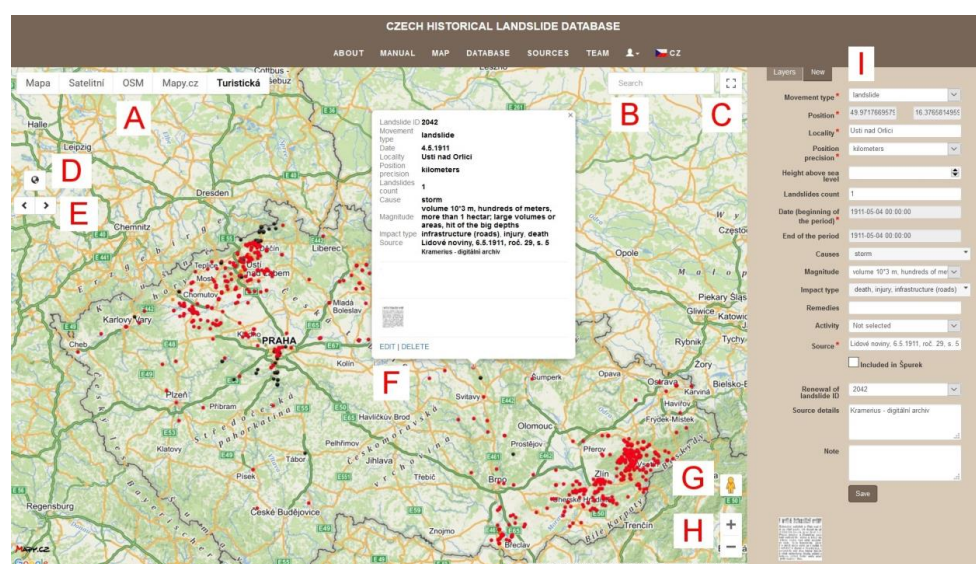
190 added through a form (I). This panel contains two bookmarks. In the *Layer,* the symbology and the time filter

191 can be seen. In the *New*, a new landslide record can be added, or an existing one can be edited. Only registered

192 users can edit their own landslides. The edit button is visible in the pop-up window in the map or in the Database

193 after log in. There is currently no automatic validation of data entered by users to have the app as accessible and

194 user friendly as possible. Three roles of users are defined: non-registered users, registered users (who are

195 allowed to edit, delete their own data) and administrators (editing, deleting all data).

196

197 *Figure 4: Additional information on a landslide record can be obtained after clicking on a point in the map*

198 *view. A new landslide can be added through the form (right side). A – background map, B – municipality, C –*

199 *full-screen mode, D – full extent control, E – backward/forward screen, F - landslide description, G – Google*

200 *Street View, H – zoom in/out, I – new landslide button. © Mapy.cz*

201

202 The landslide database is also accessible via a table (Figure. 5). Four filters are available: movement type,

203 magnitude, only selection of records added by a particular user and time period. Full text searching is also

204 possible. The given filter is kept when the map is selected unless it is reset by the user.

Figure 5: A database table visualizes information on landslides and allows for filtering data according to a number of parameters

4. Results

4.1 Landslide Records in Childa

We present below an overview of data contained in CHILDA for the 1132–1989 period. The database contains 699 records, 619 of them categorised as landslides, 75 as rockfalls and five as 'other' kind of mass movements (earth flow, rockslide or human-induced landslide). As regards temporal accuracy, 231 records were determined exactly at single day precision, 17 records are known with a week and 88 with a month precision. In total, 363 records were only attributed to a given year. Concerning the location accuracy, 111 records were localized precisely, 71 records with a precision of "*tens of metres*", 260 records to "*hundreds of metres" and* 478 to kilometres (mostly between 1–2 km, exceptionally up to 5 km in the mountain terrains of Czechia).

Table 2 presents database completeness that was determined based on several non-mandatory fields (cause of landsliding, magnitude, impact, etc.). The relative number [%] of particular fields always refers to all records in CHILDA (i.e., a proportion of 100%). The average number of database completeness amounted to 33 %. An increase, however, in database completeness in the next years, based on the incorporation of new results arising from future research, is assumed.

Table 2: Database completeness on the basis of particular non-mandatory fields.

Filed name	Number of records	Relative number [%]
Cause of landsliding	120	17.2
Magnitude	181	25.9
Impact	354	50.6
Remedies	38	5.4
Source details	645	92.3
Photo	53	7.6

Our data can also be compared to the previous landslide chronology compiled by Špůrek (1972) for the area of Czechia. Špůrek's last data on landsliding come from 1970. Our comparison consequently also ends this year. As for 1970, CHILDA currently (April 2021) contains 667 records. In total, 359 of these records were also part of Špůrek's overview. This indicates that our new investigation constitutes almost 50% of new records in the

228 database, i.e., records which were not previously covered by Špůrek's catalogue. A major source of Špůrek's

229 information about landsliding was newspapers. We focused in our research, however, apart from newspapers on

230 the primary documentary data available in archives (e.g., chronicles, memory books, official correspondence) as

231 well. This approach implies huge potential for revealing new and unique landslide events in the future

232 As Figure 6 indicates, the highest density of historical landslide records in Czechia in the studied period

233 concentrates on three primary landslide areas (OWC, CS, DV). Tens of landslide records come, however, from

234 the area of the capital Prague (VR) represented mainly by rockfalls. A higher ~~occurrence~~ of records is also

235 ~~typical~~ for Bohemian Paradise sandstones (BP).

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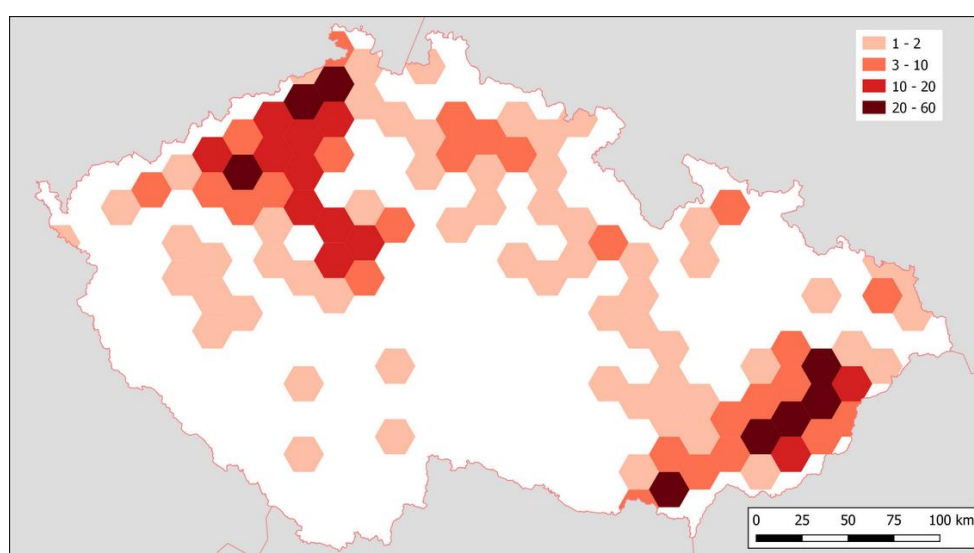
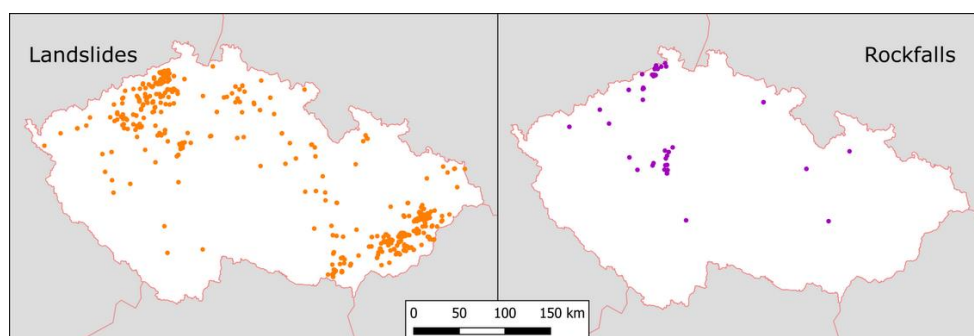

237 *Figure 6: Landslide density across Czechia.*

238 While landslide records come from all three core landslide areas, the rockfall records were typical for the west

239 part of Czechia (DV area) and along the Vltava/Moldau river valley (VR) south of Prague (Figure 7).

240
241 *Figure 7: Landslide and rockfall distribution.*

242

243   The temporal variability of the landslide records, represented by their numbers within centuries in the whole

244   studied period, is shown in Table 3. The only one (and at the same time the oldest) record is known from the

245   twelfth century (19 January 1132, see chapter 4.2). No reports were detected between the thirteenth and fifteenth

246   centuries and only five and two reports were recorded in the sixteenth and seventeenth centuries, respectively.

247   While only 8 records were found between the twelfth and seventeenth centuries, and 30 records in the eighteenth

248   century, the majority of the records are evidenced in the last two centuries. Landslide records covering the

249   nineteenth century account for 27.8 % whereas records from the twentieth century embrace 66.8% of all reports.

250   *Table 3: The number of landslide records within centuries in Czechia between the twelfth and twentieth*

251   *centuries*

Century	12th	13th	14th	15th	16th	17th	18th	19th	20th*
n	1	0	0	0	5	2	30	194	467

252   ** up to 1989 including*

253   Due to a rapid increase in landslide records after 1770, a decadal frequency of landslide records was created

254   starting in 1761. The records in the period 1761–1989 represent 98% of all records in CHILDA (Figure 8). The

255   highest numbers of records were registered during the 1891–1900 decade (19 %) followed by 1931–1940 (14 %)

256   and 1961–1970 (14 %) decades. No records were detected, however, between 1791–1810. A significant rise in

257   the number of records is apparent since the 1891–1900 decade and the course over the rest of the twentieth

258   century is more or less variable. The variability has been in all probability influenced by important rainfall

259   and/or snow thaw events.

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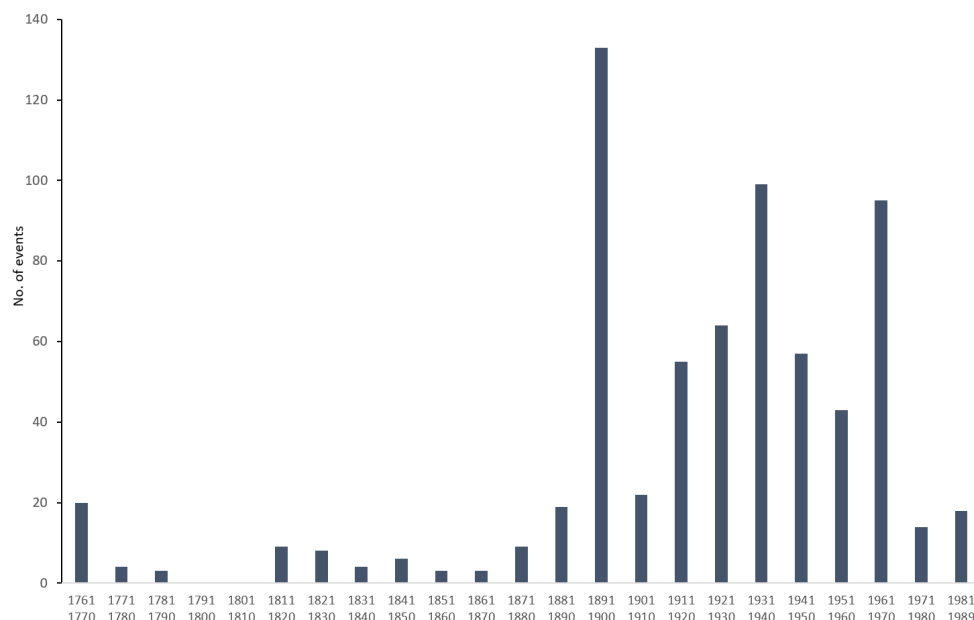

261   *Figure 8: Decadal frequency of landslide records in Czechia between 1761–1989*

262 The first most important landslide year occurred in 1770. A minimal number of landslide records to determine a
263 particular year as the most important landslide year was set at 10 reports. From 1770, 15 such years were
264 revealed encompassing 55% of all records on mass movements (385 out of 699) (Table 4). The highest number
265 of landslide records was found in 1900 (61), 1941 (51), and 1939 (46), representing 15.8%, 13.2, and 11.9% of
266 the most important 15 landslide years, respectively.

267 *Table 4: The most important 15 landslide years (1770–1989) when at least 10 records were found*

Year	1770	1898	1899	1900	1915	1919	1926	1937
n	17	39	13	61	10	10	32	10
Year	1938	1939	1940	1941	1965	1967	1970	
n	11	46	19	51	38	14	14	

268

4.2 The Oldest Records on Landsliding

270 The issue of precise determination of dates of landslide activity is rising when we look back in ~~history,~~ The
271 oldest records describing landslides in Czechia suffer from spatiotemporal inhomogeneity. CHILDA currently
272 contains 14 reports of mass movements which took place before 1770, the first most important landslide year
273 (see Tab. 4). It is important to mention that in 50 % of all recorded cases, citations referring to historical
274 landslides in this study came from the Špůrek landslide catalogue (1972). Because of our efforts at maximum
275 authenticity, we took over these citations although we were not able to study some referred citations personally
276 in certain cases.

277 The oldest known written report describes a rockfall in Praha – Chuchle (VR area) on 19 January 1132 (Strnad,
278 1790). More detailed information about three landslide events in spring 1531 is described by chroniclers from
279 Litoměřice and Bílina (CS area). Landsliding was reported from the beginning of April until the middle of May
280 and affected Radobýl Hill near Litoměřice and Holý vrch Hill near Zahořany village. Vineyards planted on the
281 hills and slopes slid and two great parts of Radobýl Hill slipped down, including trees and plants. Similarly, a
282 large portion of Holý vrch Hill slid at that time and a series of other landslides were observed (Smetana, 1978).
283 In the wet spring of 1531, several landslides also occurred in the surroundings of the nearby Bílina River after a
284 flood (Hutter, 1891).

285 A day before Christmas Eve of 1595 a landslide occurred near Vraclav – Domoradice village (Špůrek, 1967).
286 According to Kárník et al. (1957), seismic activity preceded this event. In March 1599, extraordinary damage
287 was described in Litoměřice (CS area) as a consequence of a great deal of snow and wet weather (Brázdil et al.,
288 2013b). The chronicler described the situation as follows: "*A piece of town wall near St. Laurentius* [church] *fell
289 and collapsed* [...] *On 16 March* [...] *at the cemetery* [...] *a huge section slipped down too so the graves opened
290 and the dead bodies were thrown out* [...] *That same year, in the month of March, extensive damage to vineyards
291 was experienced by many* [people]*, the walls caved in and one vineyard after another slid and all of this was
292 happening due to great wetness*". In addition, a number of springs appeared on the surface, the cellars were full
293 of water which had to be pumped out and one house even slipped away (Smetana, 1978).

294 The oldest landslide report from OWC, based on a written record, occurred in the Pavlovské vrchy Hills in
295 Pavlov village in 1663 (Maca, 1994; Kryčer). In contrast to previous mass movement events, continuing

296 landsliding was also recorded on the same street in 1667 (Štefková-Vajayová, 2001), 1715, 1730, and 1763
297 (Maca, 1994; Kryčer). A detailed description of landslide damage before 1763 is missing, but the particular
298 landslide consequences in the area of the Pavlovské vrchy Hills in the following years were described by Bíl et
299 al. (2020).

300 A brief report of landslide occurrence comes from Děčín – Chrochvice (DV area) in 1736 (Mauder, 1931). This
301 event was the first one from a series of consequent landslides in 1823, 1850 and 1914 which affected an area of
302 16 ha and damaged buildings and local infrastructure (Špůrek, 1972). A similar lack of information relates to a
303 landslide in Ústí nad Labem (DV area) in 1767 (Špůrek, 1972). In contrast the sliding down of a parish cellar in
304 Líbeznice village in 1769 is relatively well described by Třebízský (1885) who mentioned great wetness and
305 surfeit of water this year.

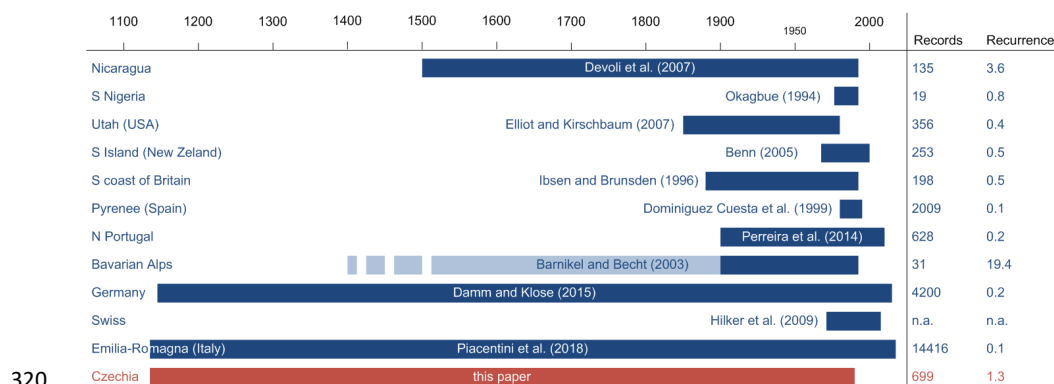
306

307 **5. Discussion**

308 **5.1 CHILDA and Other Historical Landslide Databases**

309 We presented an overview of the CHILDA database where as many as possible records on historical landsliding,
310 which took part in the area of modern Czechia, were collected. This database currently contains 699 records
311 (between 1132 and 1989) and can be compared to other similar databases which have been completed in other
312 countries (Figure 9).

313 It should be noted, however, that similarly to other databases CHILDA displays high asymmetry in the number
314 of recorded events over time. First, the sole oldest record dated to 1132 extends the span by four centuries as
315 further records are only dated to 1531. Second, the majority of records (93%) relates of landsliding that occurred
316 since 1850. In contrast, and unlike the other databases, CHILDA also records only a few increased landslide
317 frequencies in the pre-industrial periods, namely the 1770 landslide year (17 landslides) resulting from the
318 Central-European adverse climate (Raška et al., 2016), and 1817 (with five landslides) possibly influenced by
319 the Tambora eruption in 1815 (Brázdil et al., 2016b).

320

321 *Figure 9: A comparison of CHILDA and other existing historical landslide databases in terms of the number of*
322 *records and recurrence.*

323   *Note: dark-blue stripes mark the periods under study. In case of compound (multi-hazard) databases, light-blue*

324   *marks complete the studied period for which geohazards other than landslides have been found, whereas dark-*

325   *blue stands for landsliding. Dashed stripes represent an unspecified beginning (e.g., since the fifteenth century*

326   *without explicit dating of the oldest record).*

327

5. 2 Limitations of the CHILDA Database

329   As with similar databases which focus on historical records and therefore depend on availability, accessibility

330   and reliability of original sources, CHILDA also has certain limitations that may be grouped in the following

331   kinds of uncertainties:

332   (a) Uncertainty is usually related to a description of events as landslides (mass movements in general) in the

333   documentary records which were named differently. The terminological inconsistency, closely connected with

334   exploiting the wide range of documentary data, lasted at least until the twentieth century and persists in media

335   reports up until recent days. This inconsistency was caused by gaps in scientific knowledge in the first Czech

336   modern landslide classification published in 1954 (Záruba and Mencl, 1954; compared with much earlier

337   attempts in English listed in Cruden, 2003), and by cultural and resulting language diversity in the Czech Lands

338   up until 1945 (i.e. Czech and German culture realms). Raška et al. (2015) analysed historical landslides in two

339   Czech regions and found seven different terms referring to landslides and three for rockfalls. In total, five

340   German different terms were used to describe the landslides in the Pavlovské vrchy Hills (OWC) between the

341   middle of the seventeenth–twentieth centuries (Bíl et al., 2020).

342   (b) Moreover, documentary evidence was not continuous, sometimes only depending on concrete persons.

343   Chroniclers were often not interested in this phenomenon as there were more dangerous ones in this area, such as

344   floods (Brázdil et al., 2011; Elleder et al., 2020), strong winds (Brázdil et al., 2004, 2017, 2018b), episodes of

345   drought (Brázdil et al., 2013a; Dolák et al., 2015; Řezníčková et al., 2015; Brázdil et al. 2016a, 2019), etc. There

346   is obviously also a significant growth of available sources especially since the nineteenth century, which may

347   distort the represented landslide frequencies over time. The availability of sources also displays a distinct

348   geographical variation due to different historical developments in individual regions. Several smaller

349   uncertainties related to historical landslides have thus appeared, e.g., a lack of narrative sources before 1920 and

350   in the Czech border areas after 1945 (Bíl et al., 2020) or difficulties with recognition of landslides and scoured

351   slopes during the floods in early documentary records. As regards the Czech border areas, chronicles were lost or

352   carried away mainly by German inhabitants displaced after 1945. This occurred, along with other regions, in the

353   landslide prone areas of the Pavlovské vrchy Hills or parts of NW Czechia (CS and DV). Only some of the lost

354   information from the chronicles became available again via a historical synthesis published in the second half of

355   the twentieth century (Bíl et al., 2020). In addition, keeping memory books was recommended by the state in

356   1836 and chronicles were made mandatory from 1920 onward, therefore any older events are either missing in

357   these sources or were recorded retrospectively with some risk of misinterpretation.

358   (c) Additional problems include uncertainties resulting from difficult to identify duplicities in records of different

359   coverage (i.e., the issue of upscaling and downscaling) and from false frequency peaks caused by a combination

of continual (regularly published) and stationary (published once as a collection of records) data (Raška et al., 2014b). Both these issues may result in false peaks and gaps in the landslide time series. Verification of data through checking of the data reliability based on comparison among more sources, or field research is not always possible and reliable, and therefore any database must always be considered a catalogue of records rather than events themselves. In the case of the catalogue of events utilisation, these uncertainties relate, e.g., to the example of 12 records in CHILDA that describe more than one event, but the real number is not known. Such events could therefore present events with higher intensities than other events where only single landslides were reported.

In addition, any quantification of landslide extents and their impact is also complicated. We therefore decided to only select from a few attributes generally describing impacts. Reactivations of landslide at the same place could also not be determined exactly. Sometimes, as documented in Bíl et al. (2014) from the village of Halenkovice, villagers from landslide prone areas were used to seeing the landslides often and as a result did not pay attention to them. In contrast, rare landsliding in other areas attracted the attention of the locals. Construction work, related to the first railways across OWC, also both found and caused some landslides (e.g., Záruba, 1938). The apparent lack of documented landslides before 1920 in the Carpathians (except for the Pavlovské vrchy Hills) was attributed by Bíl et al. (2014) to dispersed settlements built primarily from wood, the majority of the unpaved roads and relatively sparse inhabitation in the area. Limited spatial accuracies of historical records often influence any reliable evaluation regarding the possible structural or anthropogenic triggers.

Despite all these uncertainties, documentary evidence stands as a valuable and indispensable source of data describing the occurrence and as well as the consequences of landslides in Czechia during the last five centuries.

5. 3 Further Applications and Development of the Database

Data, currently contained in CHILDA, can be further analyzed in order to delineate and explain their temporal and spatial concentration. While the spatial extent of landsliding can easily be understood as an apparent relation to bedrock, temporal distribution has been influenced by important rainfall and/or snow thaw events. Information on landslides with known dates of activities, which is the case of CHILDA records, can therefore be used in such a determination of regional rainfall or total water content triggers. Bíl et al. (2016) have already utilized, for example, the information on landslide periods, defined for an area in the central part of OWC, to determine rainfall thresholds for this area. Further applications of the data will include analyses of the long-term changes in landslide risk reduction approaches, their effectiveness and efficiency (Caloireo et al., 2014; Klose et al., 2016) framed by disaster risk reduction strategies (DRR) (Bíl et al., 2014; UNDRR, 2015). Analysis of community responses to landslide risk in individual landslide-prone areas has already been published by Raška (2019) and Klimeš et al. (2020) and allowed for an exploration of both formal (planning, DRR administration) and informal (community help, mobilization of local knowledge) mechanisms in landslide risk reduction.

6. Conclusions

We presented the online landslide database CHILDA (Czech historical landslide database) which summarizes information about landslides which took place in the area of Czechia (the Czech Republic). The database is freely accessible via the https://childa.cz/ website, and currently includes 699 records (spanning the 1132–1989

period). The oldest record relates to a rockfall which took place in 1132. In total, the database doubled the

number of records known from the previous historical database in Czechia. We further described in detail

another eight of the oldest records (1531 to 1730) and analysed centennial and decadal frequencies of landslide

records. It was demonstrated that 55 % of all recorded landslide events occurred only within 15 years of the

extreme landslide incidence. Finally, the limitations of the documentary data sources have been summarized

pointing at uncertainties within the database. The future research direction should focus on analysing historical

landslide triggers and their thresholds, changes in spatiotemporal patterns of landslide impacts on society and on

narratives of societal adaptive management to landslide risk.

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