UAV survey method to monitoring and analysing geological hazards: The Case study of a mud volcano of Villaggio Santa Barbara, Caltanissetta (Sicily)

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Abstract. All active geological processes determine effects on the soil due to different deformation processes: surface uplift and subsidence, shear lineaments with differential kinematics in relation to the source and the soils involved. Nowadays the remote sensing represents a key tool for the evaluation and monitoring of the hazard. The use of Unmanned Aerial Vehicles (UAVs) in contexts of natural hazard presents three main steps for risk assessment and monitoring: pre-post event data acquisition, emergency support and monitoring. The mud volcano of Santa Barbara (Municipality of Caltanissetta, Italy), represents a dangerous site because on 11th August 2008 a paroxysmal event caused serious damage to infrastructures for a range of about 2 km. The main clues of mud volcano paroxysmal events are the uplift and the development of structural features with dimensions ranged from centimetre to decimetre. Here we present a methodology for monitoring of deformation processes that may be precursors of paroxysmal events on the Santa Barbara mud volcano. This methodology is based on i) the data collection, ii) the Structure for Motion (SfM) processing chain and iii) the M3C2-PM algorithm for the comparison between point clouds and uncertainty analysis with a statistical approach. The objective of this methodology is to detect hazard precursors by monitoring of deformation processes with centimetre precision and a temporal frequency of 1 - 2 months.

1 Introduction

In recent decades, both high-resolution digital photographs and Structure for Motion (SfM) software allowed to generate high-quality topographic information. In geosciences many studies have been dedicated to morphological processes (Amici et al. 2013b; Casella et al., 2014; Castillo et al., 2012; Deng et al. 2019; Dietrich, 2016; Eltner et al., 2015; Gomez-Gutierrez et al., 2014; James and Robson, 2012; James and Robson, 2014b, James and Varley, 2012; Javernick et al., 2016; Lucieer et al., 2014; Ryan et al., 2015; Smith et al., 2016; Walter et al., 2018b; Westoby et al., 2015; Woodget et al., 2015). Applications include runoff laboratory trials (Morgan et al., 2017), applied geology (Niethammer et al., 2012; Russell, 2016;

Saito et al., 2018), geomorphology (Bakker and Lane, 2016; Balaguer-Puig et al., 2017; Dietrich, 2015; Dietrich, 2016 a/b; Heindel et al., 2018; Javernick et al., 2014; Marteau et al., 2016; Mercer and Westbrook, 2016; Pearson et al., 2017; Prosdocimi et al., 2017; Seitz et al., 2018; Smith and Vericat, 2015; Snapir et al., 2014; Vinci et al., 2017), glaciology (Immerzeel et al., 2017; Piermattei et al., 2016), coastal morphology (Casella et al., 2016; Brunier et al., 2016; James and Robson, 2012), volcanology (Andaru and Rau, 2019; Bretar et al., 2013; Carr et al., 2018; De Beni et al. 2019; Favalli et al., 2018; Giordan et al., 2017, 2018; James and Robson, 2012; Müller et al. 2017; Witt et al. 2018) and geophysics (Amici et al., 2013a; Di Felice et al., 2018; Federico et al., 2019; Greco et al., 2016; Zahorec et al., 2018). SfM is commonly used in the cultural heritage field for the 3D reconstruction (Jalandoni et al., 2018; Sapirstein, 2016, 2018; Sapirstein and Murray, 2017). The monitoring of active geological processes is a preventive action in risk mitigation (Deng et al. 2019; Diefenbach et al., 2018, Rosa et al., 2018; Stöcker et al., 2017; Turner et al., 2017b). Disasters occur when two factors - hazard and vulnerability - coincide. The risk is proportional to the magnitude of the hazards and the vulnerability of the affected population. Among the deformation monitoring systems, the photogrammetry technique from Unmanned Aerial Vehicles (UAVs) is spreading thanks to the high efficiency in data acquisition, the low cost compared to traditional techniques and in the acquisition of high resolution images (Fonstad et al., 2013; Harwin and Lucieer, 2012; James and Robson, 2012; James et al., 2017/a/b; James et al., 2020; Javernick et al., 2013; Johnson et al., 2014; Westoby et al., 2012). This technique is important to study the catastrophic natural events such as floods, earthquakes, landslides, subsidence, etc. Different acquisition methods and the ability to obtain high spatial and temporal resolutions (Boccardo et al. 2015) allow to acquire detailed information on the evolution of the landscape, therefore UAVs are an effective and complementary tool for field investigations. Furthermore, UAVs have other advantages including: (i) the ability to fly at low altitudes, (ii) the ability to reach remote locations, (iii) the ability to host multiple sensors (cameras, Lidar, thermal imaging cameras, navigation / inertial sensors, etc.), (iv) the ability to capture images at different angles, and (v) the flexibility to carry out monitoring operations on a small, medium and large scale (Jordan et al. 2017). Ground Control Points (GCPs) are used to improve the accuracy of the resulting data. Therefore, recognisable points on the UAV imagery are measured with a high-precision surveying device to georeferenced the data. In this process a correct number of GCPs is required which lead to a greater accuracy of the resulting data (point clouds, 3D grid, orthomosaic or Digital Surface Model (DSM)). The precision of the resulting data is also controlled by other variables, such as: the focal distance of the camera; flight path and flight altitude; the orientation of the camera; the picture quality; the processing chain and the category of UAV system (fixed or rotary wings).

In this paper, we present the results and analysis of the surface deformation monitoring of the mud volcano of Santa Barbara (Caltanissetta, central Sicily) (Fig.1). We have applied the statistical analysis of significant changes with 95% Level of Detection (LoD 95%). In detail, we used precision maps and the M3C2-PM (Lague et al., 2013; James et al., 2017b) algorithm to determine the surface variations. The statistical analysis allows to verify i) the uncertainty between the different surveys, ii) the spatial variability of the accuracy in the surveys (James et al., 2017/b), iii) the quality of the georeferencing of the surveys based on the number of GCPs.

Figure 1: Aerial view of the Santa Barbara mud volcano. Photo taken by the UAV from the east side of the mud volcano.

The mud volcano of Santa Barbara is located within the Caltanissetta foredeep basin of the Apennines-Maghrebian collisional chain which developed from the Late Miocene to the Quaternary, along the border of the converging Eurasia-Nubia plate (Catalano et al., 2008; Dewey et al., 1989; Serpelloni et al., 2007). This structural domain is formed by a foreland fold and a thrust belt involving the deposition of clastic sediments which were gradually contracted and moved from the late Miocene to the Pleistocene (Monaco and Tortorici, 1996; Lickorish et al., 1999, and references therein).

According to Madonia et al., (2011) mud volcanoes are most of the time in stasis, but they are a preferential way for rising fluids rich in methane and sludge, therefore they can be considered a risk to urbanized areas or sites with an economy dedicated to natural attractions.

On the mud volcanoes (e.g. Ayaz-Akhtarma and Khara Zira Island mud volcanoes in Indonesia), geomorphic/structural features have been observed from one year to shortly before the paroxysmal event (Antonelli et al., 2014; Madonia et al., 2011). Also the area of the Santa Barbara volcano was affected in 2008 by paroxysmal mud eruption which was preceded by deformation features (Fig.2). Moreover, the surface of the mud volcanic cone is incised by a drainage system (Fig.1) characterised by hydrographic basins with elongated dendritic geometry arranged to a centrifugal development from the areas of the summit craters towards the lower slopes of the volcano complex. The higher order of the thalwegs present deep recessed meanders (landscape rejuvenation process). This morphometric structure is typical of uplifting areas and therefore relative decrease of the base level. This suggests an inflection process of the volcano ground surface induced by increase of fluid pressure inside of shallower stagnation chamber which is located at a depth of about 30 m (Imposa et al., 2018). The stagnation chamber has a "sill-like" geometry, a radius of about 50 m and a thickness of about 30 m. This morphostructural configuration supported by geophysical data configures the active geological structure as a high potential geological hazard.
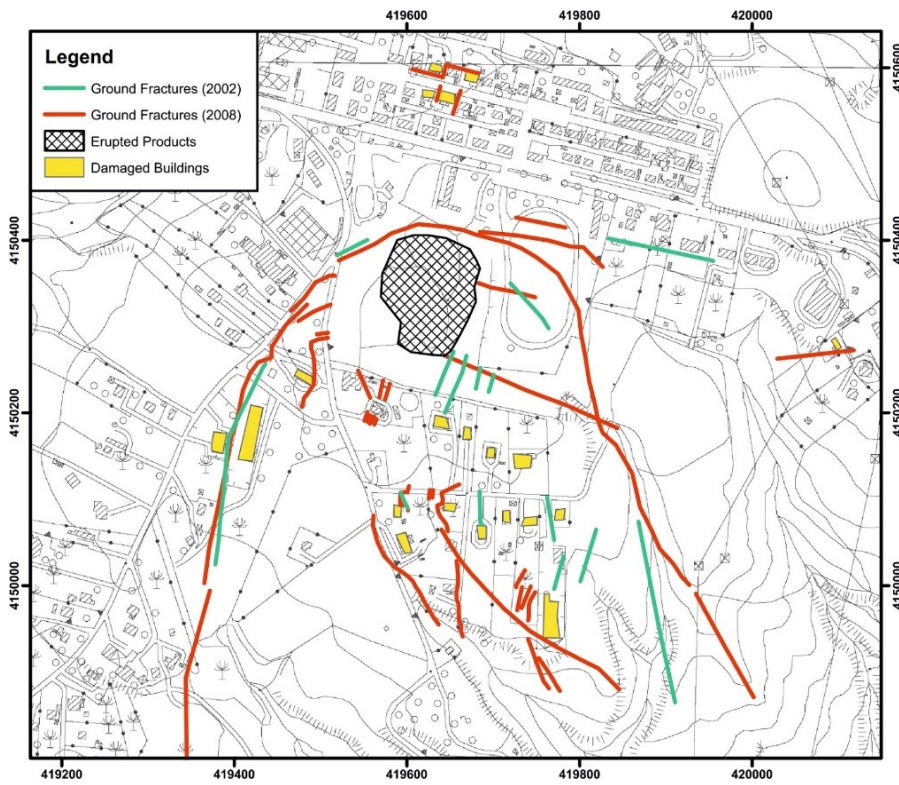
Figure 2: Cartographic extract with the location of the major fractures (in red) detected on the ground and the damaged structures (in yellow) related to the paroxysmal event of 2008. In green fractures detected in 2002. The base map is Carta Tecnica Regionale 2008 (CTR). The reference system is WGS 84 / UTM zone 33N.

On the surface, fractures and shear lineaments extend outside the areas of the volcano, (Madonia et al., 2011; Bonini et al. 2012; INGV, 2008a; Regione Siciliana, 2008) and they highlight the high stress and strain environment induced by the mud volcano (Fig.2). Such structures have been detected in the past and we speculate that they are still active. This development has often been a precursor of paroxysmal events such as the 11th August 2008 event (INGV, 2008a; Regione Siciliana, 2008).

2 Methods

2.1 Local Network

In order to monitor active deformation in the mud volcano area, a local GNSS network was created according to the criteria described by De Guidi et al (2017), in particular: i) ensuring the basic requirement of spatial and temporal stability ii) absence of possible gravitational instabilities in both static and dynamic conditions at sites and iii) a panoramic and elevated position for the Theodolite Total Station (TST).

According to these criteria two GNSS benchmarks were created: CTN0 and CTN1, located on the roof of a building on the northern sector of the studied area (Fig.3).

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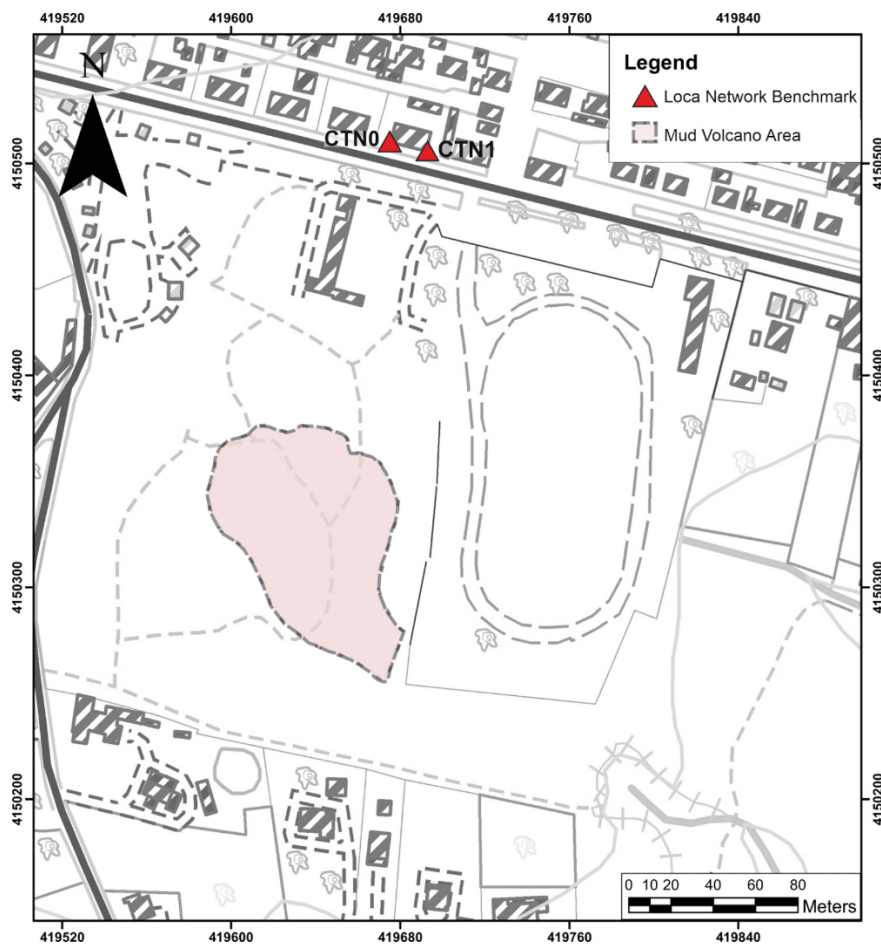
Figure 3: Chorography of the eastern periphery of the inhabited area of Caltanissetta (Santa Barbara Village). The base map is Carta Tecnica Regionale 2008 (CTR). The reference system is WGS 84 / UTM zone 33N.

The benchmarks were surveyed using double frequency (L1/L2) receivers (TOPCON Hiper V and HiPer SR) in static mode.

115   Once the stability of the benchmarks has been assessed, we surveyed CTN0 and CTN1 5 and 2 times respectively.

Post processing of GNSS data was carried out by AUSPOS Online service (Geoscience Australia, 2011; Jia, M et al., 2014).

To process the CTN0 data of 2018/02/28 survey, AUSPOS used 15 IGS stations to compute the baselines: ANKR, BOR1, BRUX, BUCU, GANP, GRAS, GRAZ, LROC, MAT1, MEDI, SOFI, TLSE, VILL, YEBE and ZIM2 with an average of ambiguity resolution of 90.0%, position uncertainty (95% C.L.) are respectively of 0.005 m, 0.005 m, 0.016 for East, North

120   and Ellipsoidal height.

To process the CTN1 data of 2018/11/26 survey, AUSPOS used 15 IGS stations to computed the baselines: ANKR, BOR1, BRUX, BUCU, GANP, GRAS, GRAZ, LROC, MAT1, MEDI, SOFI, TLSE, VILL, YEBE and ZIM2 with an average of ambiguity resolution of 84.5%, position uncertainty (95% C.L.) are respectively of 0.007 m, 0.007 m, 0.024 for East, North and Ellipsoidal height.

125   Finally, the optimal ITRF2014-UTM33N coordinates have been definitively assigned to CTN0 and CTN1.

2.2 Ground Control Points (GCPs)

Various acquisition methods of Ground Control Points (GCPs) were tested in order to define the most suitable one. The main function of GCPs is to geo-reference outcomes from SfM.

130   Initially (2016-2018) we used only GNSS receivers in different configurations: Real Time Kinematics (RTK) and Static Ultra Rapid. The GPCs were made of 50 cm x 50 cm alveolar polypropylene square targets. Using these configurations (Tab. 1), errors ranging from centimetre to decimetre were recorded. In these early phases, errors were only computed by SfM software PhotoScan (v 1.4.5.7554).

PhotoScan provides different types of error estimation: XY error (m) - root mean square error for horizontal coordinates for
135   a GCP location; Z error (m) - error for elevation coordinate for a GCP location; Error X, Y and Z (m) - root mean square error for X, Y, Z coordinates for a GCP location; Error Img (pix) - root mean square error for X, Y coordinates on an image for a GCP location averaged over all the images; Total Error (m) - implies averaging over all the GCP locations.

Using Static Ultra Rapid mode, the total error was about 18 cm, whereas using the RTK configuration the total error were reduced up to about 4 cm. (Brighenti et al, 2018).

140   From 2018 the use of the TopCon DS-103 TST was introduced, obtaining lower values on the GCP errors than those detected with the GNSS technique (Tab.1). With this measurement technique the total error has been reduced to about 3 cm. On the first 3 campaigns, we used to georeferenced the cloud points 6 GCPs (Fig.4). In the following section, these first 3 campaigns will be not considered in the computation of the significant changes for their incomparable to the last campaigns. Since 2019 only TST has been used for the GCP survey and according to Tahar et al. (2013) the number of GCPs has been
145   increased (Tab.1). Considering these two improvements, the total error has been reduced to about 1.4 cm and in the last campaign about 0.7 cm.

DATE	METHOD OF SURVEY	GCPs NUMBER	X ERROR (CM)	Y ERROR (CM)	Z ERROR (CM)	TOTAL ERROR (CM)	IMG ERROR (PIX)
2018/02/28	Static Ultra Rapid	6	12.06	8.85	10.02	18.01	0.587
2018/04/16	RTK	6	2.18	1.79	3.34	4.29	0.326
2018/04/16	TST	6	2.04	1.39	1.68	2.99	0.325
2019/07/29	TST	29	0.78	0.95	0.72	1.43	0.244
2019/09/13	TST	30	0.79	0.84	0.86	1.45	0.316
2019/10/14	TST	31	0.78	0.66	0.78	1.29	0.311
2020/01/13	TST	31	0.95	0.82	0.77	1.48	0.237
2020/06/15	TST	26	0.43	0.39	0.45	0.73	0.249

Table 1: Average RMSE of GCPs obtained for each survey and for each technique used to determinate the GCPs coordinates: X (Easting), Y (Northing), Z (Altitude) and the Total Error. The image residual (IMG ERROR, in the table) is shown.

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Count	X error (cm)	Y error (cm)	Z error (cm)	XY error (cm)	Total (cm)
6	8.84501	6.28639	9.78311	10.8514	14.6103

Count	X error (cm)	Y error (cm)	Z error (cm)	XY error (cm)	Total (cm)
30	0.793059	0.848473	0.868639	1.1614	1.4503
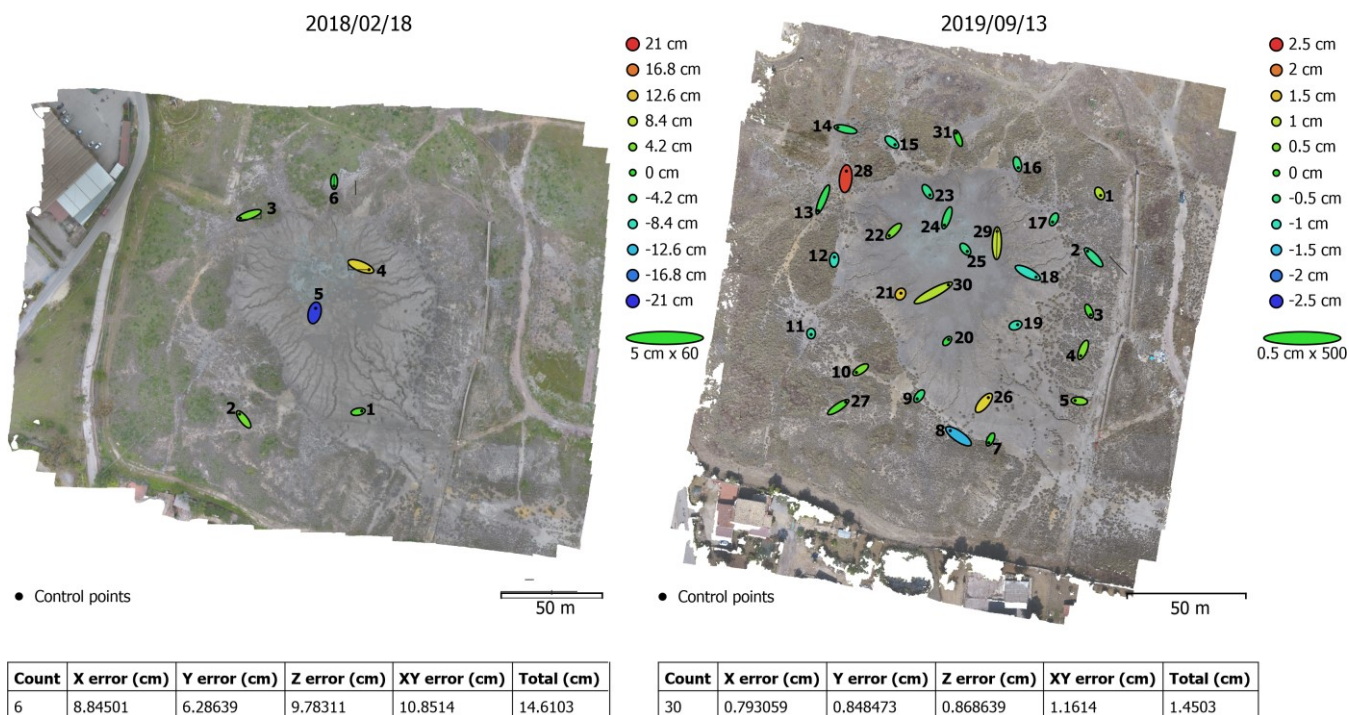
Figure 4: GCPs locations and error estimates. Z error is represented by the colour of the ellipse. X, Y errors are represented by ellipse dimensions. GCP locations are marked with a dot. Note that the different scale of the error ellipse in green: in the left image the ellipse in X direction is enlarged 60 times, in the right image it is enlarged 500 times. The reference system is WGS 84 / UTM zone 33N.

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In the last two campaigns we used TST to obtain the coordinates of the GCPs. The TST was positioned on the CTN0 point of the local network, (Fig.3) which coincides with the roof of the nearby private houses in the northern part (Fig.5 A). A classic celerimetric survey was carried out.

160 The measurements of the GCPs were carried out with a surveying ranging rod equipped with a reflecting prism (offset of -30 mm) assisted by a tripod with a bubble level (Fig.5 B), to ensure the verticality and stability of the measurement.

Figure 5: A) TST DS 103 placed on the fixed metal base in coincidence with point CTN0. B) Surveying ranging rod on the mud volcano during the survey phase.

Based on the CTN0 point and directing on the CTN1 point, we assumed an instrumental error of 5 mm (named in PhotoScan "Marker Accuracy"). This value has been assumed due to the uncertainty of the CTN0 and CTN1 point coordinates, obtained through GNSS measurements, considering that the uncertainty derived from the TST is negligible.

To validate GCPs data we have done an analysis (compared with the SfM software data) using the python script "Monte_Carlo_BA.py", with a statistical iterative approach (Monte Carlo approach) (James et al. 2017/a). For clarity, the following terminology will be used in the text:

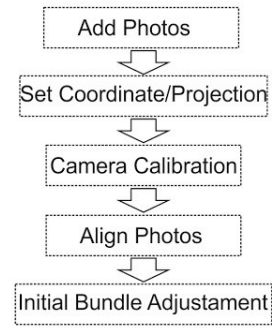
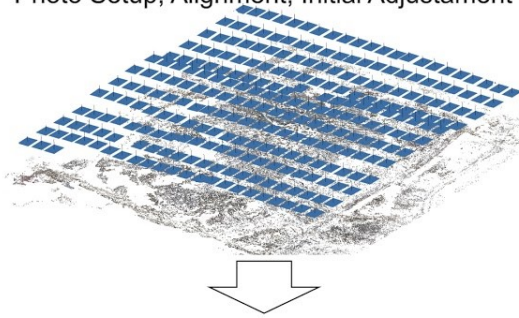
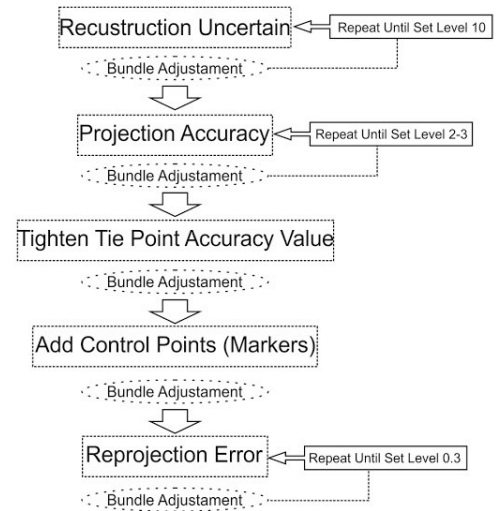
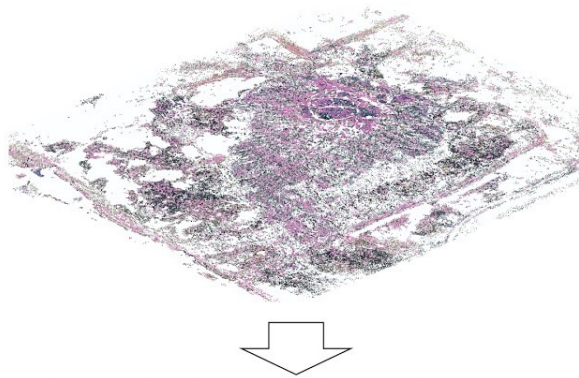
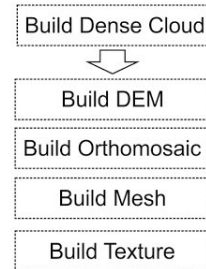
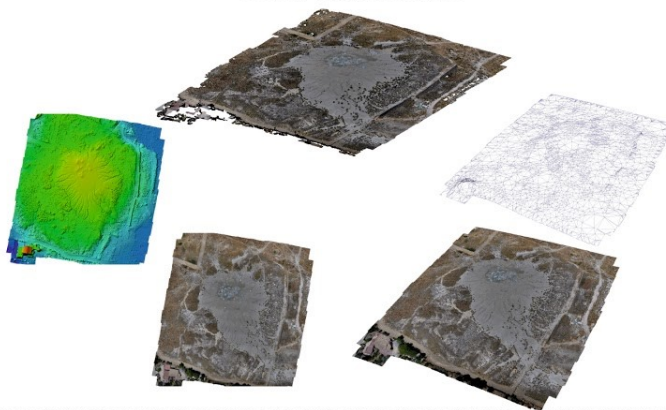
- GCPs are the points measured in the field. They can be used as control points or check points within the bundle adjustment (James et al. 2017/a).
- Control points are GCPs when they are tied to the model during in the bundle adjustment.
- Check points are GCPs when they are not tied to the model in the bundle adjustment.

This script modifies the percentage of GCPs which are used as check points or control points and applies to the check points random variations (James et al. 2017/a). To be more precise values ranging from 10% up to 80% of GCPs have been set.

2.3 Photo Acquisition

We have performed five measurement campaigns in approximately one year. The same flight plan was used for all five of them. The AOI (Area Of Interest) was captured by a DJI Phantom 4 Standard, a quadcopter UAV, at a flight height of 33 m above the ground. The sensor size of the UAV's digital camera is 6.17 mm by 4.55 mm, capable of shooting images with a resolution of 12 MP (4000 × 3000 pixels) with a mechanical shutter. Each flight planning was carried out with the Pix4D Mapper software, adopting a frontal and side overlapping of 80% and 70% respectively. The camera was set-up in a nadiral orientation to capture vertical imagery. The flight was carried out in a single grid (simple geometric flight patterns without intersections). An average of 280 images for each survey were acquired with about 1.1 cm Ground Sampling Distance (GSD).

2.4 Data Processing SfM

The photogrammetric processing was performed using the commercial software Agisoft PhotoScan. The photogrammetric processing is based on the workflow formulated by the USGS (2017) with steps of the processing chain of Tie Point Accuracy and Marker Accuracy recommended in James et al. (2017) (Fig.6).

195   **Figure 6: Simplified block diagram of the photogrammetric processing chain.**

The procedure above is conventional, but we modified same parameters during the cleaning procedure of sparse point cloud. We adopted the optimal values of "Reconstruction Uncertain" and "Projection Accuracy" in the "Gradual Selection" option. Finally, after that the model has been georeferenced, we adopted the optimal value of the "Reprojection Error" (USGS, 2017).

The first parameter, "Reconstruction Uncertain", improves the geometric reconstruction of the cloud point.

The second one, the "Projection Accuracy", improves pixel mismatches in images.

After these operations, the GCPs can be tied to the sparse point cloud.

After the georeferencing procedure, we have to perform the "Gradual Selection" option, adjusting the "Reprojection Error" values to reduce the Residual Pixel Errors.

Furthermore, appropriate pixel values of tie point accuracy and marker accuracy are set, following the suggestions of James, et al. (2017 / b).

After that, we applied the "precision_estimate.py" script (James, et al. 2017 / b). The script operates, by an interactive Monte Carlo approach, to estimate the accuracy of SfM surveys through photogrammetric and georeferencing parameters, which are then used to provide spatially variable confidence limits for the detection of surface variations.

The precision estimates are calculated through multiple "bundle adjustments" ("Optimize Camera Alignment" in PhotoScan) with different pseudo-random offsets (in this case 4000 pseudo-random offsets) (Fig.7), applied to each image and checkpoint. The pseudo-random offsets are derived from normal distributions with standard deviations representative of the appropriate accuracy within the survey (James et al. 2017 / b).

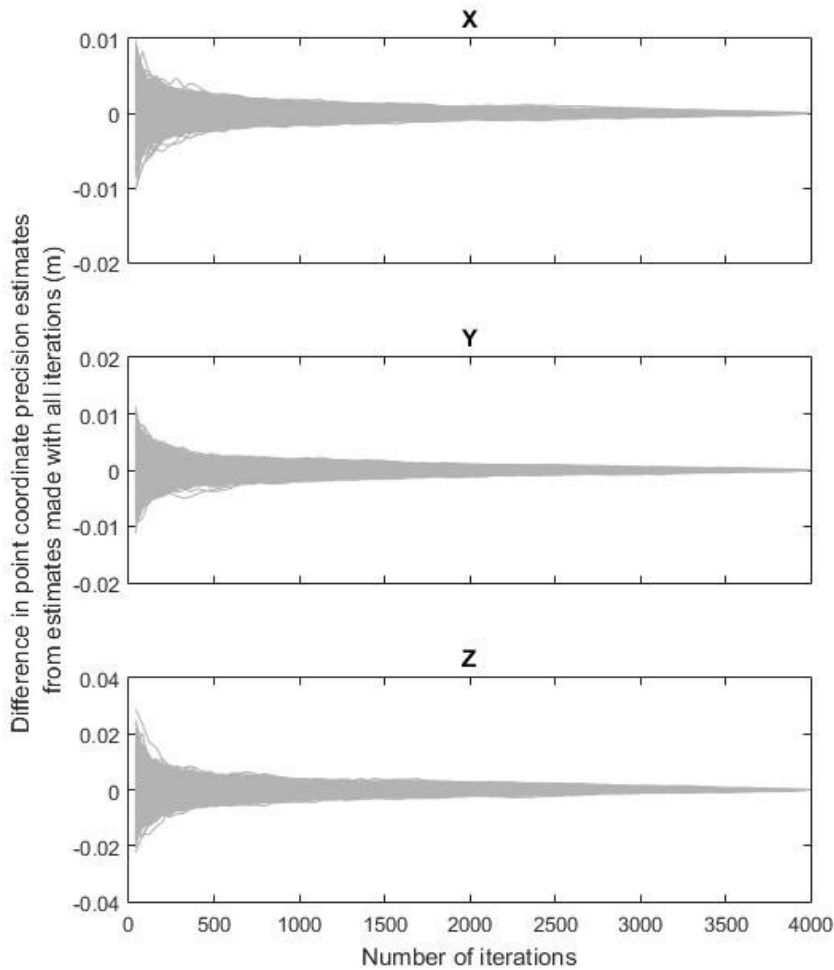
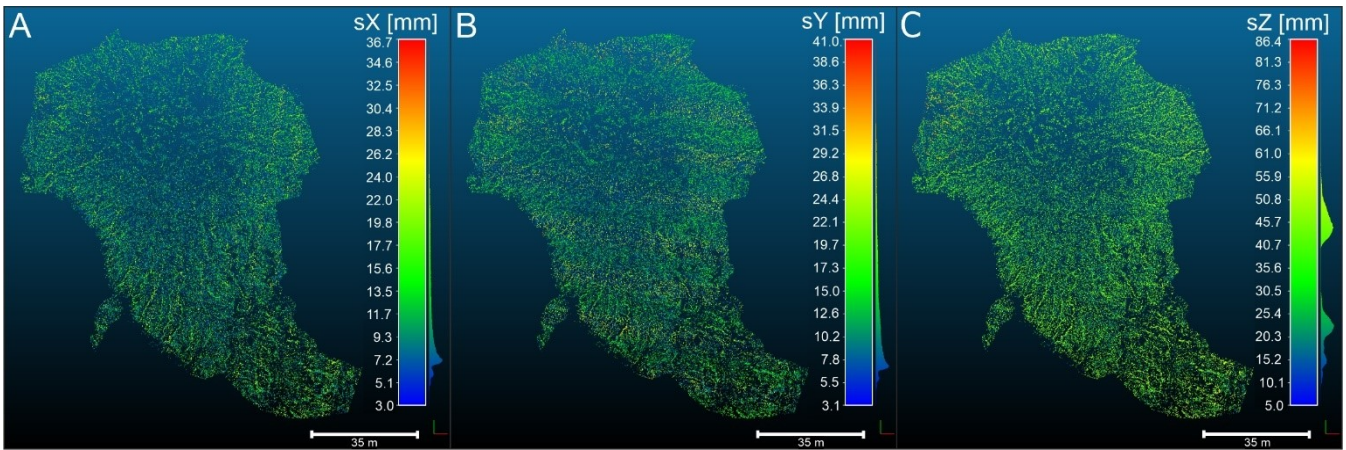



Figure 7: Example of the iterative approach in the 2020/07/29 survey. Estimates of uncertainties on X, Y and Z improve as number of iterations increase.

After that, the Sfm-Georef software (James and Robson, 2012) reads the output given by the Monte Carlo python script, setting to each point of the sparse cloud different values of precision on the three spatial components. The results of the script, read by Sfm-Georef, are estimates of the error of the individual points of the sparse cloud point in the three different spatial dimensions (Fig.8).

 Figure 8: 3D error estimates of each point of the sparse cloud of the 19/09/13 survey. The error (sX, sY, sZ) is in mm. A) and B) The horizontal errors (X and Y component) are shown. C) The vertical errors (Z component) are shown. The reference system is WGS 84 / UTM zone 33N.

The 3D topographic change is usually detected from sparse cloud points that have been cleaned to exclude the vegetation that interferes with the comparison techniques. Next step is to link those points to the precision estimates of the sparse cloud; this has been done in CloudCompare v. 2.11.

Through CloudCompare the sparse cloud is interpolated (with the relative precision values on the three-dimensional components) with the dense cloud point. In this phase we decide which interpolation technique is the more suitable, in this case the best is "nearest neighbors" that consider the three closest points to the sparse cloud, using the median value (better outliner mitigation) to assign the error values to the dense cloud points (Fig.9). This methodology has been chosen due to the heterogeneous distribution of the points in the sparse cloud, avoiding point with null values.

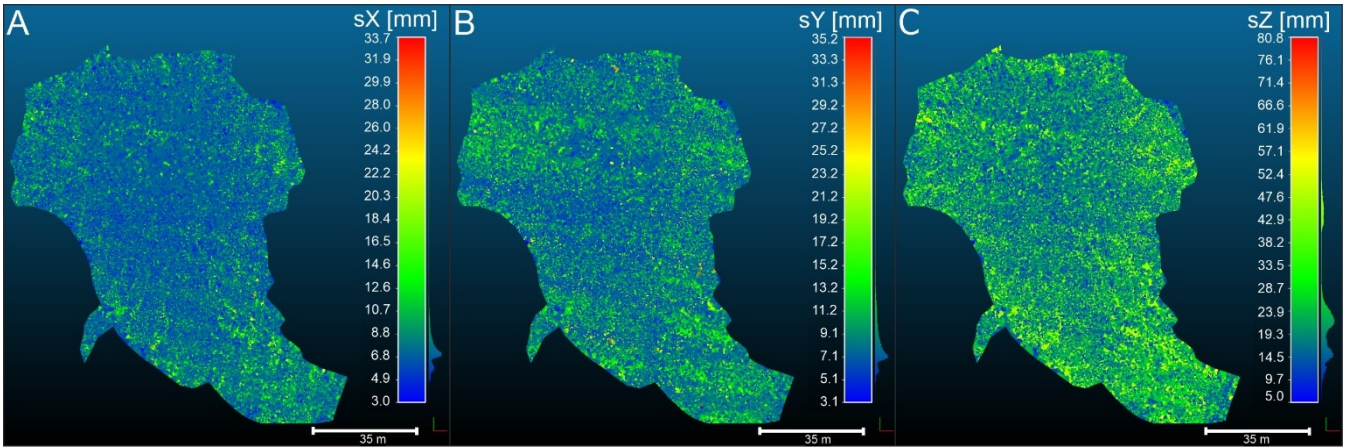



Figure 9: Estimates of the 3D error of the dense cloud points obtained by interpolation with the sparse cloud of the 19/09/13 survey. The error (sX, sY, Sz) is in mm. A) and B) The horizontal errors (X and Y component) are shown. C) The vertical errors (Z component) are shown. The reference system is WGS 84 / UTM zone 33N.

Once that the precise dense cloud points and their error have been obtained, it is possible to compare the surveys in order to get the changes between them. Comparisons between the surveys were performed on CloudCompare by the M3C2-PM plugin (James et al., 2017B; Lague et al., 2013) that identifies a statistically significant change where the topographical differences exceed a value of spatially variable uncertainty. According to James et al. (2017 / b), the M3C2-PM is particularly suited for point clouds derived from SfM. The M3C2-PM (James et al., 2017 / b) uses estimates of precision of the coordinates of points (3D precision map) that we have previously calculated.

The outputs of M3C2-PM are scalar values applied to the cloud:

- significant change (Fig.10 A).
- M3C2 distance (Fig.10 B).
- distance uncertainty (Fig.10 C).

The first output (Fig.10 A) shows the changes which exceeds the uncertainty values in both point clouds. It represents a confidence interval constrained by values of Level of Detection $_{95\%}$ (LoD$_{95\%}$) which are spatially variable. This is applicable in any morphological setting, providing a reliable 3D analysis of topographical change.

The second output (Fig.10 B) shows the calculated distances between the two clouds.

The third output (Fig.10 C) shows the uncertainty values of the distances and their spatial variation between the two clouds. Once the uncertainty value is defined, the changes are significant when they overtake the value of the uncertainty.

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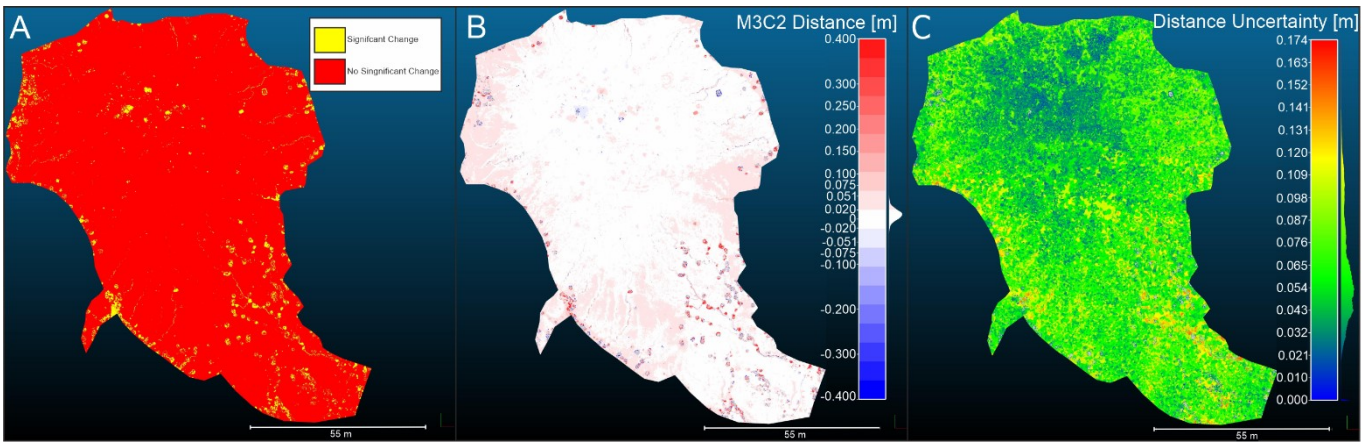
Figure 10: Point clouds resulting from M3C2-PM processing, comparison between 2019/07/29 and 2020/06/15; A) significant changes between the two point clouds with LOD$_{\%95}$; B) distances between the two point clouds; C) uncertainty of the distance between the two point clouds. The reference system is WGS 84 / UTM zone 33N.

265   In the 2020/06/15 survey, we used callipers to test and validate the method. Five numbered callipers, with steps of increasing height of about 2 cm, were positioned on the mud dome. The heights are 2, 4, 6, 8 and 10 cm.

These were used to obtain an instrumental sensitivity of the measures (Fig.11). All callipers are detected, we show as example the smallest (Fig.11).
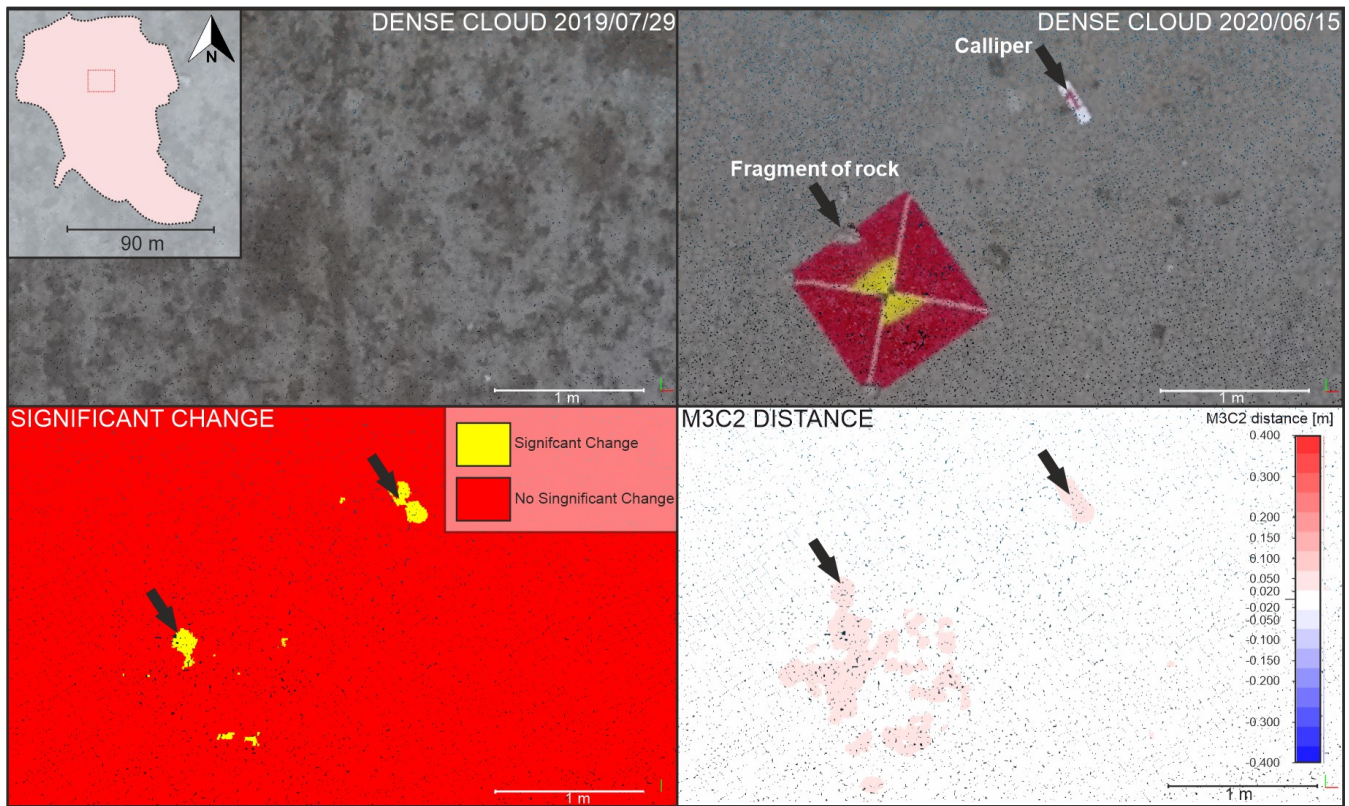
Figure 11: On the upper part of the figure the dense point clouds of the two surveys carried out on 2019/07/29 (left) and 2020/06/15 (right) are shown. On the lower part the clouds with the significant change (left) and the M3C2 distance (right). The arrows indicate the significant changes between the surveys. The calculated changes between the two clouds estimate an altitude increase of about 6.3 cm for the fragment of rock used to maintain the target in position and an increase in height of about 2.5 cm for the calliper which is 2 cm thick. The reference system is WGS 84 / UTM zone 33N.

3 Result

As illustrate on the section 2.2, the results show that using a percentage of GCPs (considered as control points) between 40 and 60, the RMSE value has a slight minimal variation. Thus, the optimal minimum number of GPCs is between 12 and 18. Exceeded the threshold of 60% there are no significant improvements (Fig.12). This result has been confirmed in all campaigns.
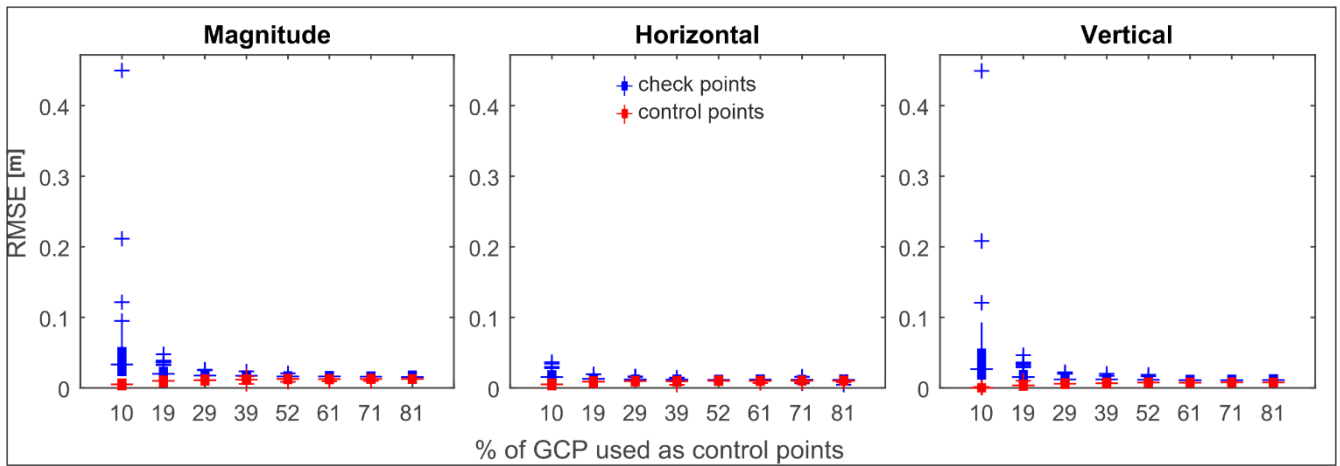
Figure 12: The boxes represent the distribution of the RMSE of GCPs (2019/09/13 survey) on the three components: Magnitude (3D), Horizontal and Vertical according to the % of the GCP used as control points. RMSE is calculated on 50 self-calibrating bundle adjustments for each % of GCPs used as control points. The GCPs are randomly selected for each self-calibrating bundle adjustments. The central bars indicate the median RMSE values, which are included in the boxes that extend from 25th to 75th percentile and the outliers are indicated by the + symbols.

The results obtained from the photogrammetric comparisons supported by geodetic topographic survey, have an average uncertainty of about 6.4 cm, with a minimum of about 2 cm and a maximum of about 12 cm, relative to an area of 42700 m² (Fig.13 A). The uncertainty of the central area has an average value of about 3.9 cm, with a minimum of about 2 cm and maximum of 10 cm, on an extension of 360 m² (Fig.13 B).
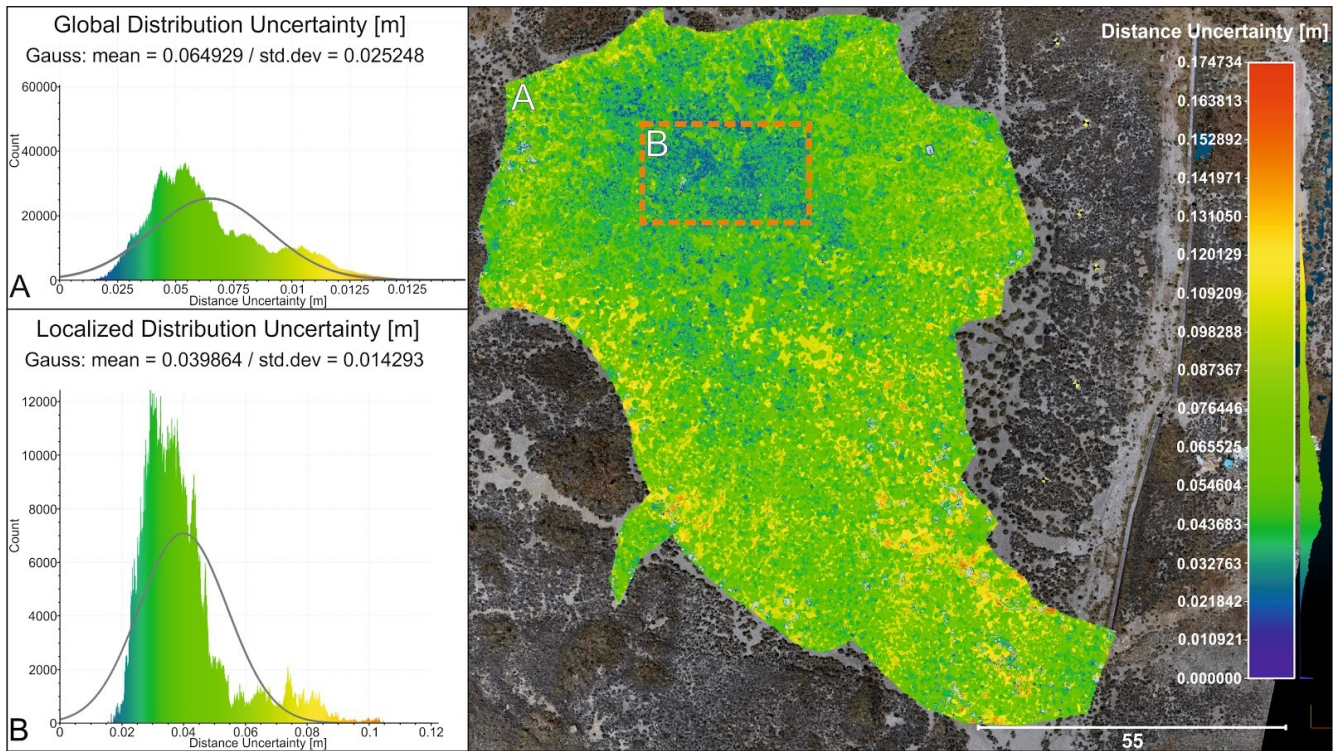
Figure 13: Points Cloud with uncertainty distance values between 2019/07/29 and 2020/06/15 (right). On the left, (A) the Gaussian distribution of the distance uncertainty over the whole area is shown, the mean is about 6.4 cm and the standard deviation is about 2.5 cm. (B) the Gaussian distribution of the distance uncertainty of the central emission area is shown, the mean is about 3.9 cm and the standard deviation is about 1.4 cm. The reference system is WGS 84 / UTM zone 33N.

During the monitoring period, in addition to natural changes, we recorded anthropogenic action, due to the relocation of objects (garbage) in the area after unauthorized access. These changes have a decimetric order of magnitude and are easily detected by the technique used (Fig.14).
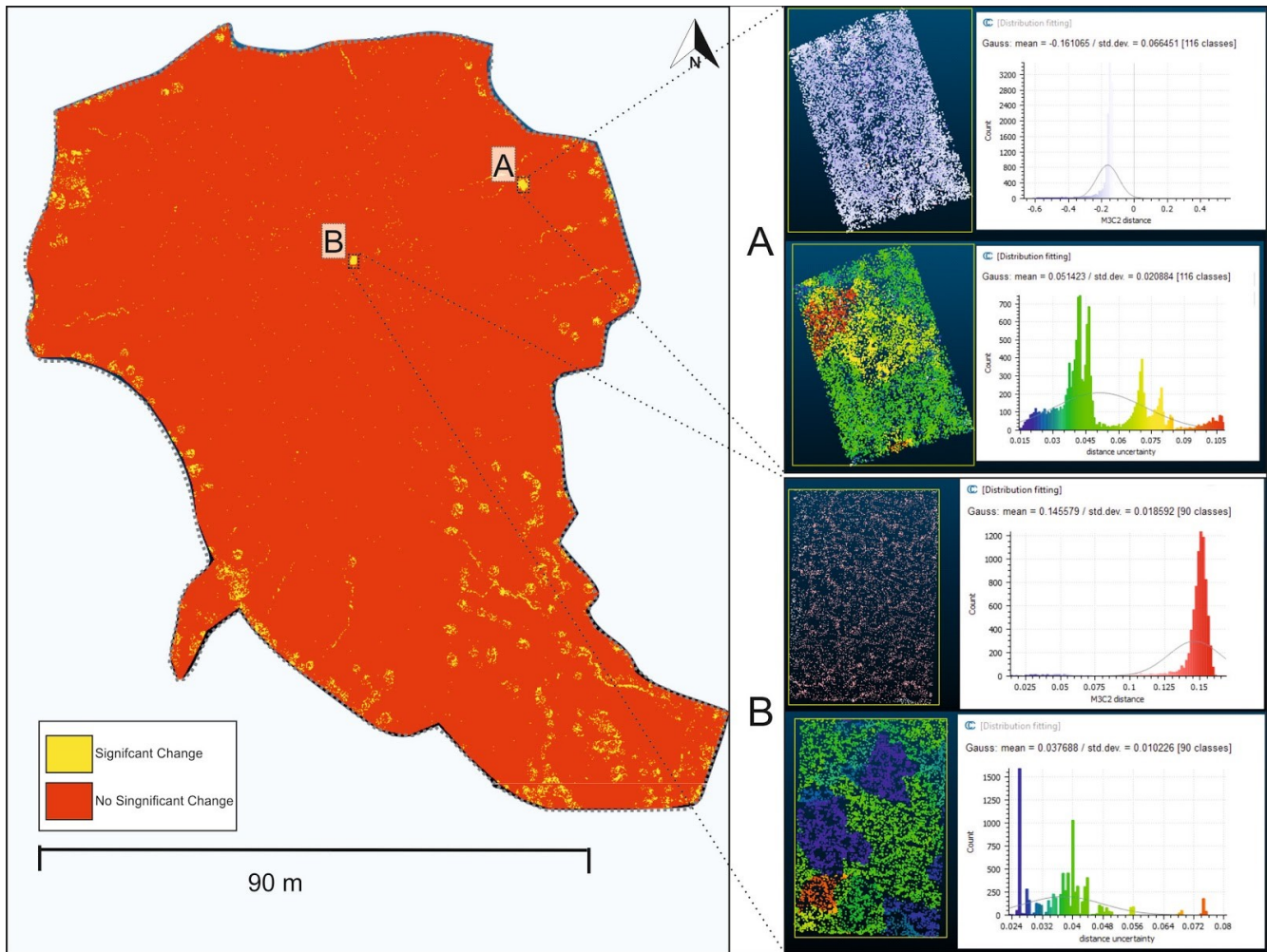
Figure 14: Enlargement of the comparison between the point clouds of the 2019/09/13 and 2019/10/14 campaigns with significant change. Two significant changes (A and B) are shown, which highlight an anthropogenic action, i.e. the movement of a wooden platform about 15 cm high. (A, upper part) a lowering of an average height of 16 cm is detected, the Gaussian distribution of the M3C2-PM distances with the mean and the standard deviation are shown. (A, lower part) the Distance Uncertain varies spatially with an average value of about 5 cm, the Gaussian distribution of the M3C2-PM distances with the mean and the standard deviation are shown. (B, upper part) an increase of height of 14 cm is recorded, the Gaussian distribution of the M3C2-PM distances with the mean and the standard deviation are shown. (B, lower part) the distance uncertain has an average value of about 3 cm, the Gaussian distribution of the M3C2-PM distances with the mean and the standard deviation are shown. The reference system is WGS 84 / UTM zone 33N.

Two types of analysis were carried out: semi-quantitative and quantitative. The surveys 2019/07/29 and 2020/06/15 (time interval of about 1 year) were chosen to perform the semi-quantitative analysis. This analysis was carried out on the whole of the mud volcano area, the objective is to detect important deformations (in the order of decimetres). To visualized important deformation (Fig.15), the values of M3C2 distance ranged between -2 and 2 cm were excluded, according to the minimum value of distance uncertainty. This range was verified instrumentally with the use of callipers.

Legend
M3C2 DISTANCE [m]

■	-0.39 - -0.15
■	-0.15 - -0.05
■	-0.05 - -0.02
⊠	-0.02 - 0.02
■	0.02 - 0.05
■	0.05 - 0.15
■	0.15 - 0.39

0 5 10   20   30   40
Meters

320
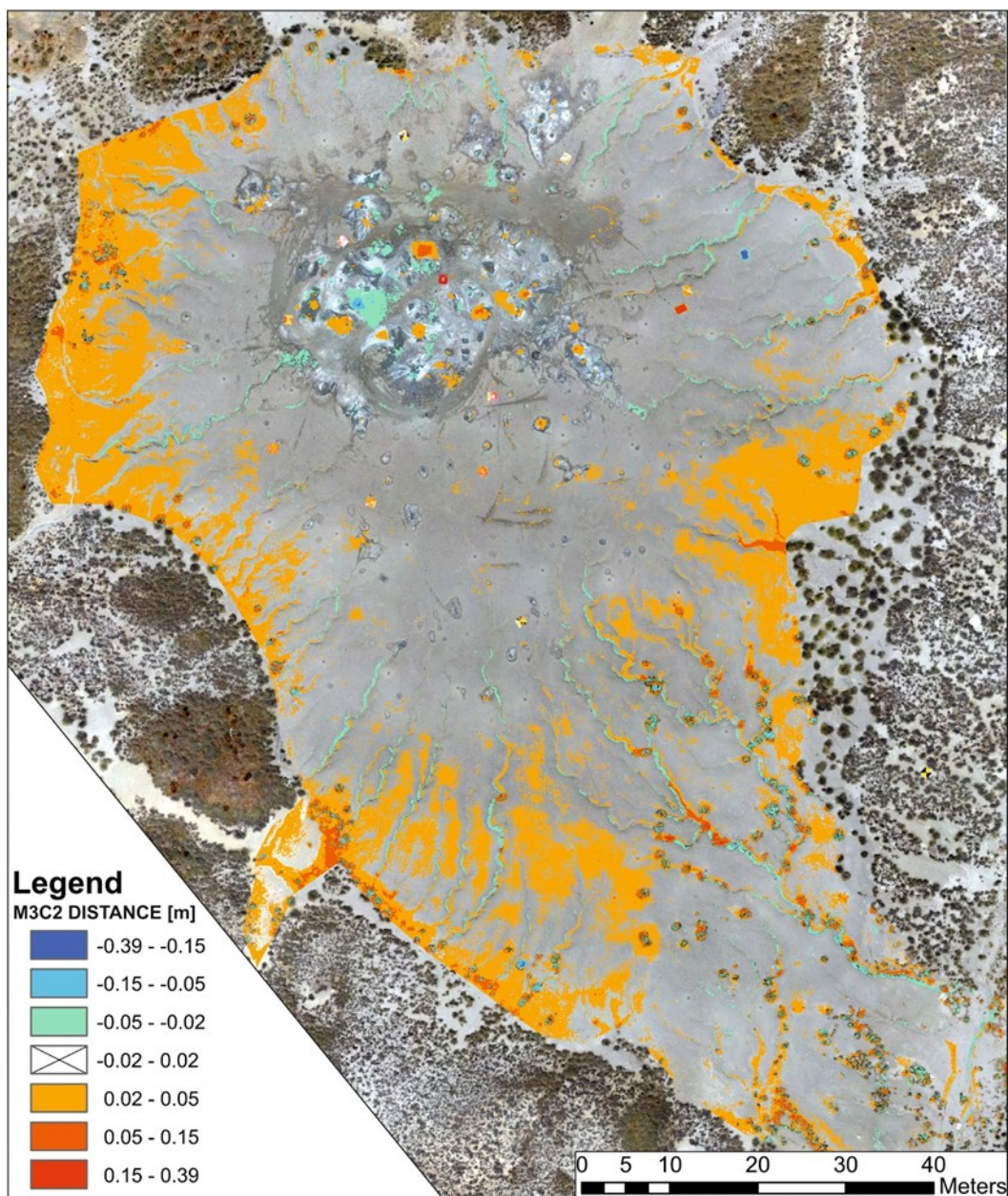
Figure 15: Orthophoto with M3C2 Distance. The distance scale has been adjusted to exclude values between -2 and +2 cm. The positive distances, observed at the margin of the mud dome, are interpreted as mud flow or sediment deposition coming from peripheral griffon vents eruptions or erosion of summit area. The reference system is WGS 84 / UTM zone 33N.

325 Considering the 2019/07/29 and 2020/06/15 surveys (Fig. 15) we observe that the total surface of the volcanic cone can be considered not affected by deformations. Only morphological changes in the volcanic structures can be underlined, such as small eruptive cones, griphons, sauces and mud pools.

The quantitative analysis was performed on small central portions of the mud volcano. The aim of the quantitative analysis is to estimate the trend and evolution of the deformation. To assess the deformation and the local morpho-structural evolution,

330 two temporal series have been developed in two areas: Z1 (collapse zone) and Z2 (uplifting zone) (Fig.16 and Fig.17). The zones were chosen related to the significant change. On the selected zone, the average distances between points with significant change were computed by M3C2-PM. Z1 has an extension of 1 m² (Fig.16) and Z2 has an extension of 0.84 m² (Fig.17).

As master was chosen the campaign of 19/07/29 (T0) (Table 2).

335

T_0	2019/07/29
T_1	2019/09/13
T_2	2019/10/14
T_3	2020/01/13
T_4	2020/06/15

Table 2: Surveys used to the generation of time series.
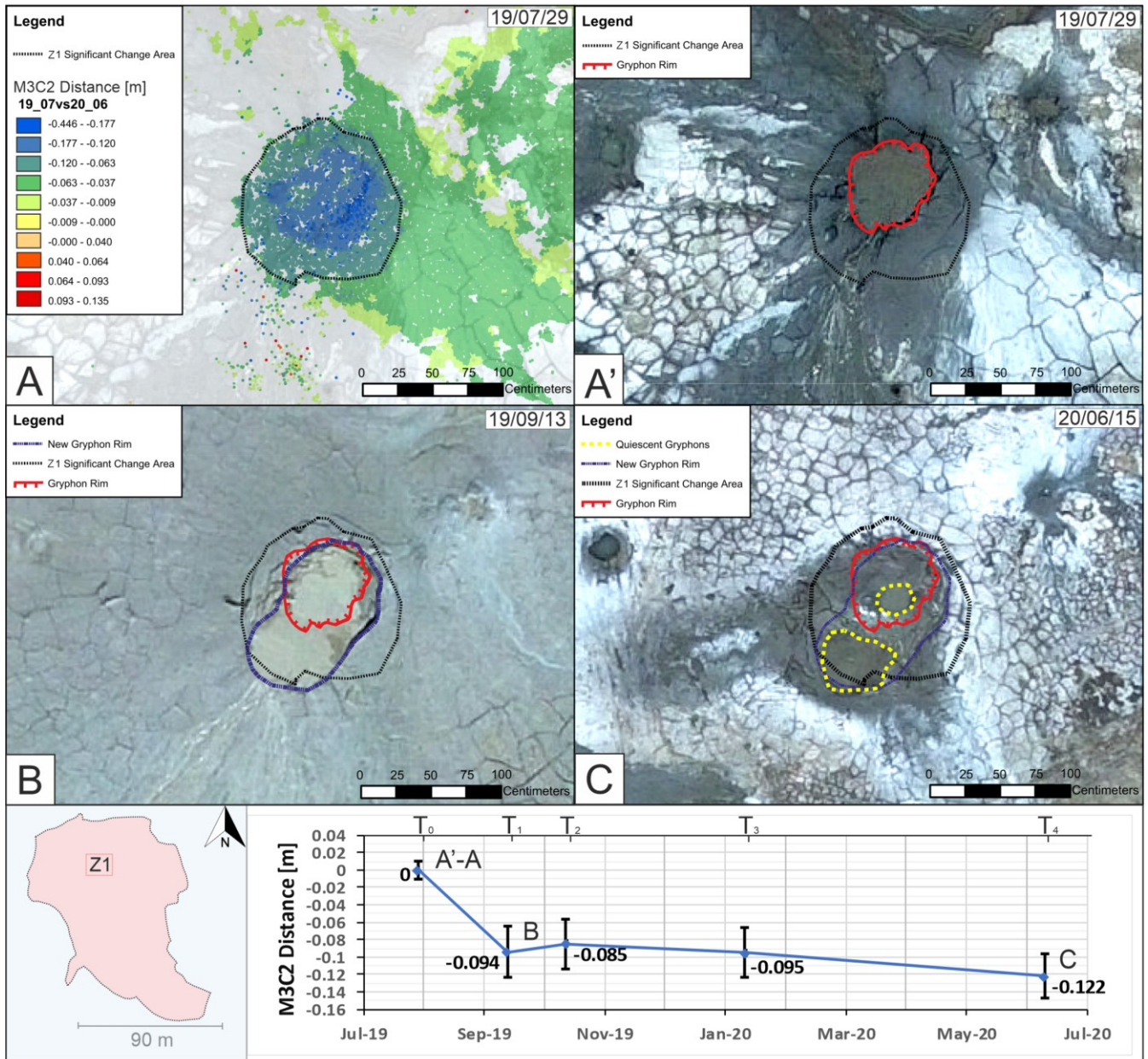
Figure 16: (A) M3C2 Distance (2019/07/29 vs 2020/06/15) of Z1; the area used to create the time series is delimited with a dashed line (selected by significant change). **(A')** Orthophoto of 2019/07/29 with the gryphon border in red. **(B)** The orthophoto of 19/09/13 highlights the expansion of the gryphon border (blue). **(C)** Orthophotos of 2020/06/15, gryphon during the quiescence phase and the split of the channel in yellow. Below the time series with a decreasing trend is shown. The reference system is WGS 84 / UTM zone 33N.

345 In figure 16 the subsidence (about 10 cm / 60 days) of gryphon is represented by the southwest migration of the mud pool edge. In figure 17, the uplift (about 4 cm / 60 days) of a blind gryphon is represented by the development of radial fractures on the cone surface. Close to the growing gryphon, another indication of deformation is the deviation of a mudslide flowing in the N-S direction towards the southern sector of the analysed area. Finally, the sudden appearance of the gryphon is well highlighted in the ortophoto of the last survey (Fig. 17).

350
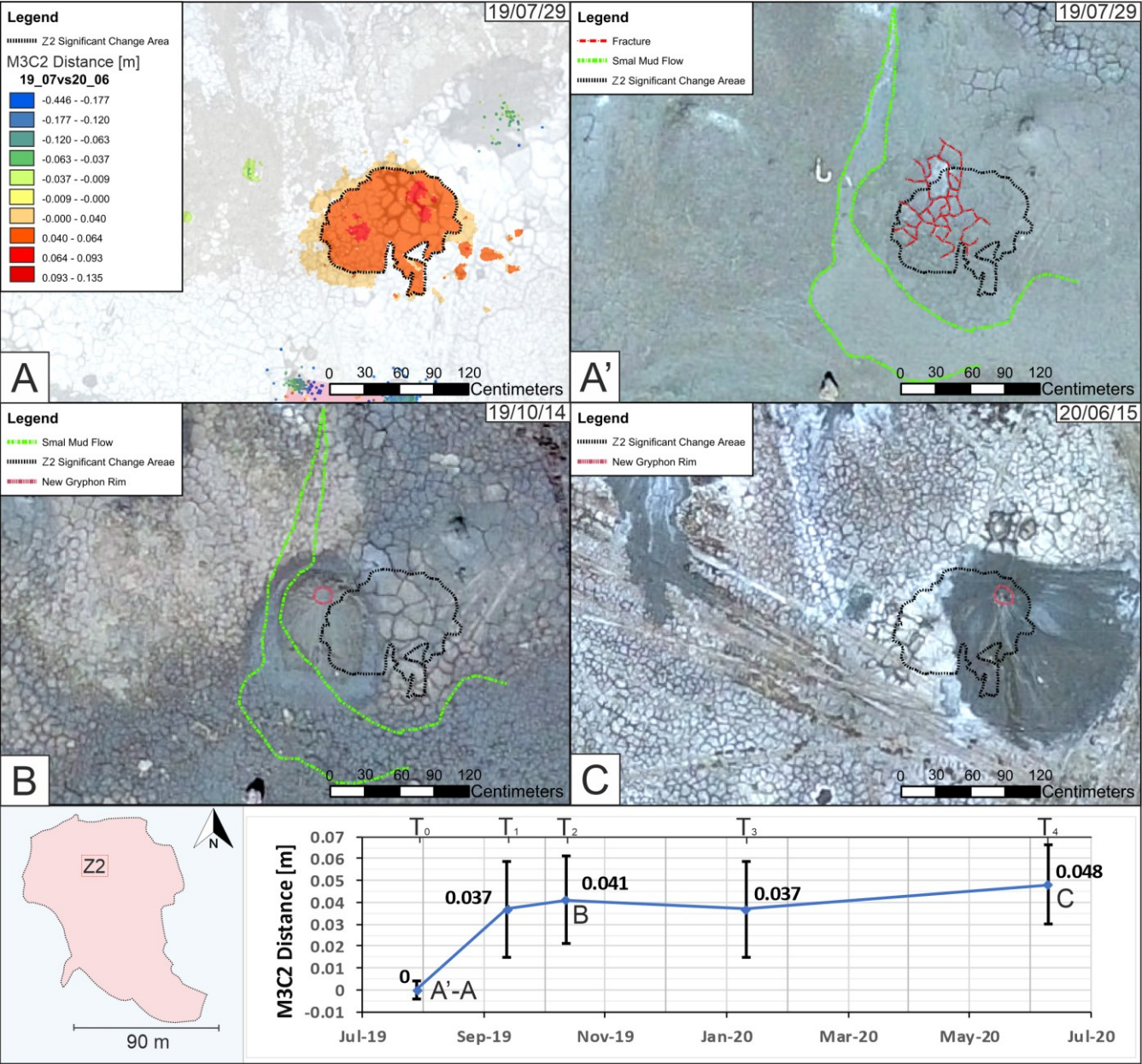
Figure 17: (A) M3C2 Distance (2019/07/29 VS 2020/06/15) of the Z2, the area used to create the time series is delimited with a dashed line (selected by significant change). (A') Orthophoto of 2019/07/29, the margin of the deflected flow, in green, and the radial fractures, in red. (B) Orthophoto of 19/10/14, formation of a new emission gryphon in red and the previous flow in green. (C) Orthophoto of 2020/06/15, formation of a new emission gryphon, in red. Below the time series with upward trend is shown. The reference system is WGS 84 / UTM zone 33N.

4 Discussion

UAV technology combined with SfM is a valuable tool in geological risk assessment and monitoring, however, some questions must be considered.

The first one is the quantity and quality of the acquired GCPs. According to the results, the minimum optimal number for geo-referencing is 12 GCPs. A slight improvement is observable up to 18 GCPs, while no appraisable improvement is detected for a higher number of GCPs (Fig.12). In addition, the method of acquiring GCPs reduces their error by using the Total Station Theodolite (Tab.1). The combined use of high-precision topographic instruments with an optimal number of GCPs improves the reliability of the datasets.

The second question is the evaluation of distance uncertainty when two surveys are compared. The distance uncertainty between two data sets (surveys) can be considered as an estimate of the sensitivity of the methods to detect measurable topographic changes. The results of the M3C2-PM show that the average distance uncertainty between the first and the last surveys is about 6.5 cm over the whole area at the 95% confidence level (Fig.13 A). Furthermore, considering a smaller portion of the area, the uncertainty decreases to about 3.9 cm at the 95% confidence level (Fig.13 B). This allows us to analyse certain morphological changes and anthropogenic activity on the volcano's shell (Fig.14). The anthropogenic activity determined a height decrease of about 16 cm where a wood platform was previously located. As well as, a height increase of about 14 cm is recorded at the place where the wood platform has been relocated. The values recorded are consistent with the real thickness of the object detected, differing by 2 cm (Fig.14). The callipers show that it is possible to measure changes at least 2 cm (Fig.11).

After that the uncertainty and sensitivity of the surveys were computed, the time series were made. Considering the significant changes and their errors, we computed the trend of deformation of two areas in order to reconstruct the evolution of the phenomena which generated these changes (Fig.16 and Fig.17).

Considering the results obtained by SfM, we propose a monitoring system of Santa Barbara mud volcano based on a semi-quantitative approach. We defined three states by setting a space-time range:

- Normal state if the deformation value does not exceed 5 cm in the range between 1 to 2 months.
- Pre-Alert state if the deformation value is between 5 cm and 10 cm during a period of between 1 to 2 months.
- Alert state if the deformation value exceeds 10 cm in the range between 1 to 2 months.

Obviously, these values must be considered in relation to the surface involved by deformation. To define the state of the activity of the mud volcano, the area affected by deformation must be a significant area of the total surface of the volcano. The definition of the activity is an important aspect of the monitoring of hazard, because there are many natural phenomena that can produce deformation in a little area (like a gryphon) or in a big area (like the sedimentation occurred at the border of the mud volcano), so the significant area affected by deformation must be correctly defined by an operator (public agency, university, institute of research or Dipartimento della Protezione Civile) which really knows the natural phenomena that can occur on the Santa Barbara mud volcano.

The methodological approach is valuable and efficient from the point of view of quantity and quality of data collected in relation to the work and time spent. This monitoring system of the deformations is a useful tool to detect the early unrest phase of the mud volcano usually induced by changes in pressure and volume of fluid raising from the stagnation chamber. The pre-eruptive deformation consists of a marked uplift and occasional small subsidence which are probably related to the redistribution of the subsoil of the pressurized fluids (Antonelli et al. 2014). According to Antonelli et al. (2014), soil uplifting can occur up to a year before the eruption.

5 Conclusion

In the management of the hazard, the SfM technique, (Gomez et al., 2016, Kaab, 2000, Fugazza et al., 2018; Giordan et al., 2017, 2018) starts to be largely used by the scientific community. In the monitoring of potentially dangerous active sites, the UAVs are very advantageous because they are not used only as a support in the post-disaster events (Rokhmana and Andaru, 2017; Hisbaron et al., 2018), but also for the pre-event monitoring.

According to Kopf (2002), Antonelli et al., (2014), Madonia et al., (2011), INGV (2008a) and Regione Siciliana (2008), the deformations of the surface shell of mud volcanoes can occur to a year before the paroxysmal event with doming and developing of structural lineaments with order of magnitude from centimetres to decimetres.

The results allow us to define the criteria for monitoring and analysing the study area. For the mud volcano of Santa Barbara, the monitoring criteria are:

- Monitoring interval between 1 and 2 months.
- The optimal number of GCPs is between 12 and 18.
- Acquisition of GCPs by high-precision topographical instrumentation (TST).
- The processing chain of the sparse point cloud according to workflow of USGS (2017) was enhanced by the correct value of "Tie Point Accuracy" and "Marker Accuracy" as suggested by James et al. (2017).
- An assessment of the state of activity of the mud volcano based on a semi-quantitative approach.

The frequencies of the campaigns depend on the status of activity, while the other criteria depend on the object / structure of the monitoring.

415 These criteria allow us to detect events with deformation of at least 2 cm. In case of anomalous values were detected the monitoring campaigns must be improve. This involves extensive monitoring, such as: i) developing time series localized in key areas, ii) combining different methodologies, e.g. micro-seismicity monitoring and three-dimensional geophysical prospecting (Imposa et al., 2016) to improve the monitoring system of the active geological process.

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