



GEOPHYSICAL AND GEOTECHNICAL FACTORS IN URBAN PLANNING: BURSA (NİLÜFER, OSMANGAZI, AND YILDIRIM) CASES

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ABSTRACT

The study area covers the Central of Bursa, Osmangazi, Yildirim, Gürsü, Kestel and Nilüfer District boundaries in Bursa. The seismic process deals with the occurrence of an earthquake event and the process of wave propagation from the source to the site. Local amplification caused by surficial soft soils is a significant factor in destructive earthquake motion. In the first phase of this study, it is investigated the ground motion level and soil amplifications for Bursa city. For his aim, probabilistic and deterministic earthquake hazard analysis (including acceleration estimations) will be carried out for the region. Local amplification caused by surficial soft soils is a significant factor in destructive earthquake motion. In the first phase of this study, it is investigated the ground motion level and soil acharacterization for the region. For his aim, probabilistic earthquake hazard analysis (including acceleration estimations) was carried out for the region. Then, soil shear wave velocities were estimated from data obtained by MASW measurements. Soil liquefaction is a natural event in which the strength and stiffness of a soil are reduced by earthquake vibrations or other dynamic loadings. As it is known, liquefaction occurs in saturated soils, that is, soils in which the space between individual particles is completely filled with water. One of liquefaction evaluation methods is based on the cyclic stress approach. In this method, a safety factor is defined as CRR / CSR . CRR is a cyclic resistance ratio that represents soil liquefaction susceptibility, and CSR is the cyclic stress ratio that represents the earthquake effect. In the second phase of this study, possible soil potential index (PL) and ground induced settlements were estimated by using Isihara ve Yoshimine (1990) approach. All results on liquefaction potential index (Pi), liquefaction induced settlements and soil shear wave velocities in Bursa (Turkey) City were



38 compared with each other. Finally, a seismic microzonation map was prepared by the
39 integration of geophysical and geotechnical data for urban planning purposes.

40

41

42 1. INTRODUCTION

43

44 The vast majority of seismic microzonation studies have been carried out in earthquake-prone
45 areas around the world (Marcellini et al., 1982, 1998; Astroza and Monge, 1991; Lasterico
46 and Monge, 1972; Faccioli et al., 1991; Chavez-Garcia and Cuenca, 1998; Lungu and et al.,
47 2000; Faccioli and Pessina, 2001; Fah et al., 1997). In terms of countries, Italy (Marcellini et
48 al., 1998); Spain (Jimenez et al., 2000), Greece (Lachet et al., 1996) and Japan (Abeki et al.,
49 1995) can be given as examples.

50

51 The first micro-zoning study in Turkey was conducted by the Construction and Housing
52 Ministry and Earthquake Research Department in 1968. In this study, geological, geophysical
53 and is used as an all geotechnical data and its first land-use planning as reflected in the
54 physical planning of a city or in other words for estimation of local earthquake risk in Turkey.
55 This city is the town of Kuyucak in Aydın province (Gülkan ve Ergünay, 2002, Kozacı ve
56 diğ., 1969). Later, microzonation study was made for the town of Gediz (Kozacı, 1970;
57 Ergünay, 1971; Tabban, 1972, Tokay and Doyuran, 1978). Similar studies were carried out
58 for Adapazarı in 1970, for İzmit in 1974, and cities such as Bolu, Gerede, and Erzincan. Erdik
59 et al. (2000) carried out a microzonation study of İzmir city in the context of earthquake
60 master plan. The first microzonation work with international partners was initiated under the
61 name of “Disaster Prevention / Mitigation Including Istanbul City Seismic Microzonation”
62 following the agreement signed between Istanbul Metropolitan Municipality and JICA (Japan
63 International Cooperation Agency). If we take a look at the academic studies carried out
64 within the scope of microzonation in our country, it seems that some articles on this subject
65 were made by Ergünay (1973), Özçep et al. (2003) and Yılmaz et al. (2003).

66

67 The plan is the process to achieve a goal and perform a job with the simplest expression.
68 Planning, on the other hand, is an act of planning and can be defined as a means of rational
69 use of resources, a correct and rational decision-making process for the future to reach the
70 expected goals. The most important goal of today's city/region planning is to provide people



71 with a healthy and safe life. Regional planning can be defined as multidimensional planning,
72 which is carried out by considering the mutual effects of the natural and artificial physical
73 environment and economic and social activities to achieve the goals adopted by the
74 community in urban areas (Aydemir, 1999, Özçep, 2005; Özçep ve diğ., 2003).

75

76 In this study, Multi-Channel Analysis of Surface Wave (MASW) analysis measurements were
77 made at 100 points carried out to obtain shear wave velocities (V_s) in the site-effect
78 investigation. Using MASW, shear wave velocity was obtained at 100 points up to 50 meters
79 deep from the surface. As is known, MASW is derived from the classical seismic search
80 method based on the multi-receiver system lined up along a line. Its main advantage is its
81 ability to recognize different types of seismic waves. This technique is important in obtaining
82 the shear wave velocity (Park et al., 1999; Park et al., 2007). Uniform Building Code,
83 Eurocode, etc). With this method, both active and passive source measurements were made
84 and both measurement results were evaluated together.

85

86 The first task for soil liquefaction analysis is to determine the ground motion level for the
87 selected region. After earthquake hazard analysis for the region is undertaken, in the second
88 stage of the study, soil liquefaction and settlements for the selected region were calculated for
89 various earthquake accelerations and magnitudes.

90

91

92 **2. IMPORTANCE OF GEOTECHNICAL ANF GEOPHYSICAL FACTORS**

93

94 One of the first known settlements of the world with its urban character and structure is
95 Çatalhöyük located in Anatolian lands (Shane III and Küçük, 1998). Interestingly, the
96 relationship between this city (where the first urban planning was made) and natural disasters
97 was reflected in the art of the period. A wall painting from 6500 BC in Çatalhöyük can be
98 seen in Figure 1. Here, an active volcano and city plan were drawn together. This picture is
99 probably the oldest known geophysical figure in the world. The active volcano in the picture
100 is thought to be Hasandağı.

101

102 The most important goals of today's urban / regional planning and urban transformation
103 planning are to provide people with a healthy and safe life. Regional planning can be defined



104 as multi-dimensional planning, which is carried out by considering the mutual effects of the
105 natural and artificial physical environment and economic and social activities to achieve the
106 goals adopted by the community in urban areas. Two different factors play a role in planning
107 studies. The first factor is needed to be taken under control of the big cities emerging as a
108 result of industrialization. The second factor is to develop the underdeveloped regions of a
109 country and to use natural resources that are not evaluated effectively. These are driving force
110 of planning studies (Aydemir, 1999, Özçep, 2005; Özçep et al., 2003; Özçep and Zarif, 2006,
111 Korkmaz and Özçep, 2020, Özçep et al, 2020).

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113

114

115 **Figure 1. A wall painting from 6500 BC in Çatalhöyük (Anatolian Civilizations**
116 **Museum, Ankara)**

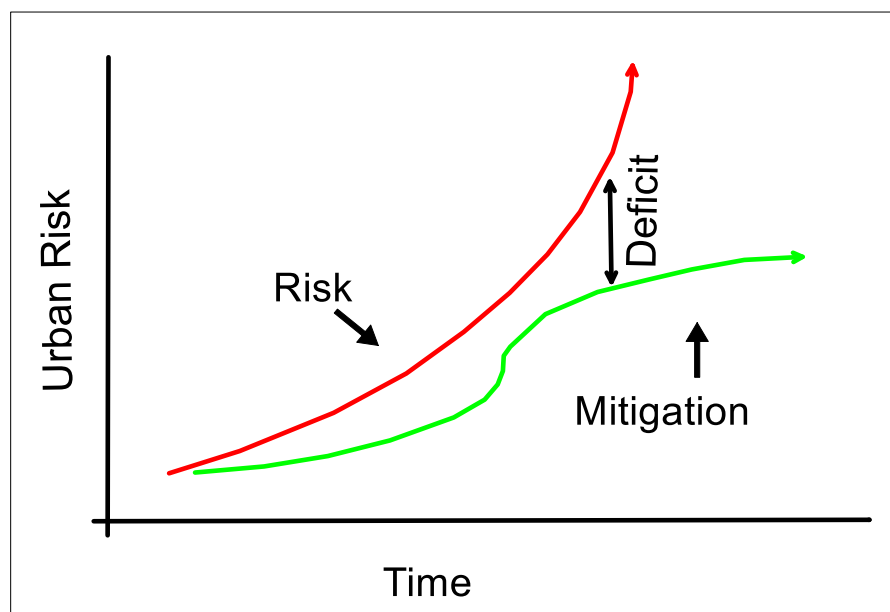
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118

119 Vast social changes and dramatic growth of the urban population led to the need for
120 improvement / development of urban areas to meet new demands. In this context, urban
121 transformation and microzonation studies were carried out in many large and small
122 settlements of the country with city or metropolitan status. Urban transformation can be
123 defined as multidisciplinary studies created to serve the improvement of unqualified and
124 suburb regions for a legal and healthy life in the border regions of industrial cities receiving
125 immigration (Nalkaya, 2006). Not only regions with historic texture or slum areas but also
126 regions under threat from earthquakes and other natural disasters should be evaluated within
127 the scope of these urban transformation studies. It should be noted that while factors such as



128 irregular urbanization and natural disasters increase the risk curve (Figure 2), our efforts to
 129 reduce risk remain insufficient.
 130



131
 132 **Figure 2.** Urban Risk-Time Relationship (Wenzel and Bendimerad, 2004)
 133

134
 135 From a general perspective, risk is defined by Beck (1999) as “human actions caused by the
 136 unplanned consequences of radicalized modernization”. According to Beck, we are members
 137 of the world risk community. Risk is a result of the impact of modern technologies. Risk has
 138 globalized and overcomes the power of the national state, and the modern age has entered a
 139 new phase called Beck's “second modernity”. The “risk society” view is defined as a state of
 140 radicalized modernity. Beck divides modern society into three periods: pre-industry, industry
 141 period and global risk society. The downside of globalization is that economic expansion and
 142 technological risks are beyond our control. According to Beck, this should be reacted with a
 143 unique form of globalization: Cross-border alliances and global networks should be
 144 established to tackle global problems on a global level. Only then can you have the chance to
 145 turn “world risk society” into “world citizenship society” as Kant says.

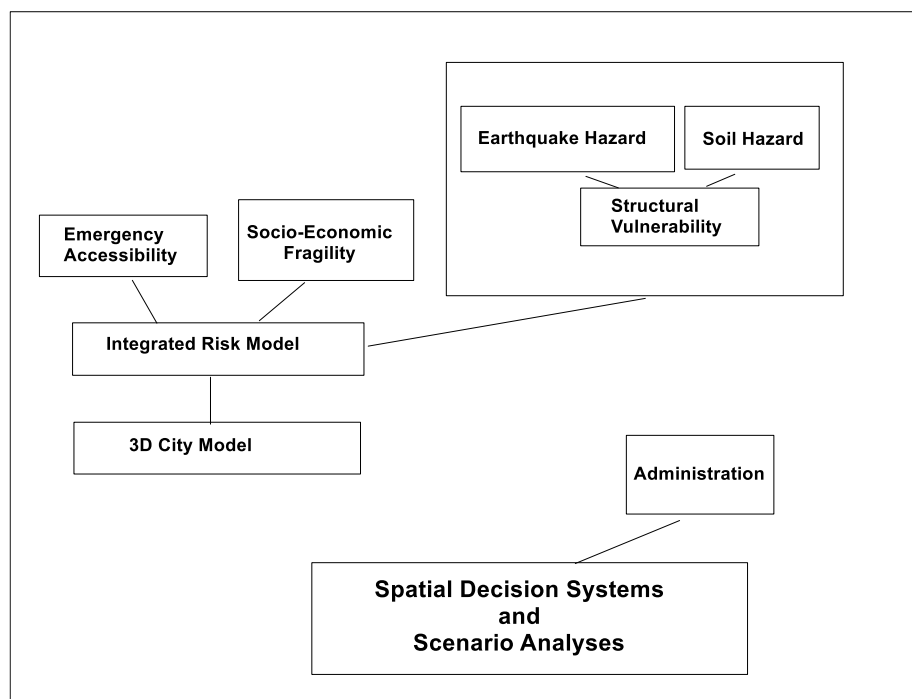
146
 147 The factors that create the risk society by Beck (1992) are defined as



148
149 • After the modern, consumption of nature / society / individual and transform them
150 into risks
151 • As science and technology are out of democratic control
152
153 If we take a look at one of the most important natural risks (earthquakes), an average of 700
154 damage-causing earthquakes occur annually around the world. Earthquakes occurring in
155 urban areas are among the most devastating natural disasters.
156 140,000 people died in the 1923 Kanto (Japan) earthquake and 240,000 people in the 1976
157 Tangshan (China) earthquake, and the total losses in the 1995 Kobe (Japan) earthquake
158 exceeded 200 billion dollars. Although it is not yet possible to predict earthquakes, studies in
159 this direction continue using all the possibilities of science and technology. On the other hand,
160 it is possible to reduce the material losses and socio-economic losses that will occur as a result
161 of earthquakes to reasonable levels by taking measures against earthquake damages. In this
162 regard, the city's local administrators have important duties in urban planning, land use and
163 control of buildings.
164
165 The intense and direct relationships of local administrators with the people of the city make it
166 possible to transfer information about earthquake damage reduction and to raise the awareness
167 of the public against earthquakes in general (Erdik et al., 2000). Düzgün (2006) proposed a
168 model to determine the spatial vulnerability of an urban area to the earthquake disaster in
169 terms of various factors (accessibility to structure, soil, social, economic, administrative and
170 critical services) using a selected city area as a pilot application area (Figure 3).
171



172



173

174

175 Figure 3. Spatial vulnerability model in terms of various factors (accessibility to structure,
 176 earthquake, soils, social, economic, administrative and critical services) to the
 177 earthquake disaster of the urban area using a selected urban area as a pilot
 178 application area (Düzgün, 2006).

179

180

181 Necessary items to estimate the effects of earthquakes in urban areas are historical earthquake
 182 information, geological, geophysical and geotechnical data, probabilistic or deterministic
 183 evaluation of earthquake hazard, determination of ground motion according to soil conditions
 184 and preparation of these studies in the context of microzonation studies.

185

186 Urban planning became prominent as an important study in expanding urban areas as a result
 187 of population growth. The balance between human settlements and the natural environment is
 188 disrupted by urbanization (De Mulder, 1996). Urban planning practices must reduce these
 189 conflicts / imbalances to improve the general well-being and quality of life of society (Bell,



190 1998; Bell et al., 1987). Such planning must be in a multi-disciplinary approach for various
 191 human needs (De Mulder, 1996). However, geological, geophysical and geotechnical data in
 192 urban planning have become increasingly important in identifying, controlling and mitigating
 193 natural disasters (Bell et al., 1987; Legget, 1987; Hake, 1987).

194

195 Many earthquakes in the past and present have shown that local soil conditions (including
 196 topography effects) have a major impact on damage distribution. When earthquake hazard is
 197 evaluated on a regional or local scale, it is very important to identify and take into account
 198 these possible local site effects. Seismic microzonation has been an economically useful
 199 element of the earthquake risk reduction process (Roca and Oliviera, 2001). Microzonation
 200 against earthquake hazards has been defined as the division of a geographical region into
 201 small areas according to the behavior of the soils underground motion or slope stability
 202 (Hays, 1980; Sharma and Kovacs, 1980)

203 Nigg (1982) stated that microzonation aims to divide the risky areas into small pieces in order
 204 to implement the right plans and policies that can minimize the damage that may occur after
 205 the earthquake. Finn (1991) has defined microzonation as procedures involving the
 206 development of measures for seismic hazards for building design, taking into account local
 207 soil conditions.

208

209 In 1993, a guideline study describing the principles was made on microzonation for "soil
 210 amplification", "slope stability" and "liquefaction" by the Earthquake Geotechnical
 211 Engineering Committee of the International Soil Mechanics and Foundation Engineering"
 212 (later "Geotechnical Engineering") Union (ISSMFE, 1993). (ISSMFE, 1993). This work was
 213 revised in 1999 (ISSMGE / TC4 1999).

214

215

216 During earthquakes, seismic waves, especially under the effect of shear waves, spread in
 217 loose soils that are generally saturated and without drainage, creating shear forces relative to
 218 each other, causing soil particles to displace. Under these conditions, saturated and loose soil
 219 particles tend to converge on each other. The stress at the contact points of the particles in this
 220 case is transmitted to the surrounding water. Since the seismic waves cause sudden and very
 221 short movements during the earthquake, it does not allow sufficient time for the water
 222 between the particles to drain. Therefore, the pressure of pore water, which cannot move away

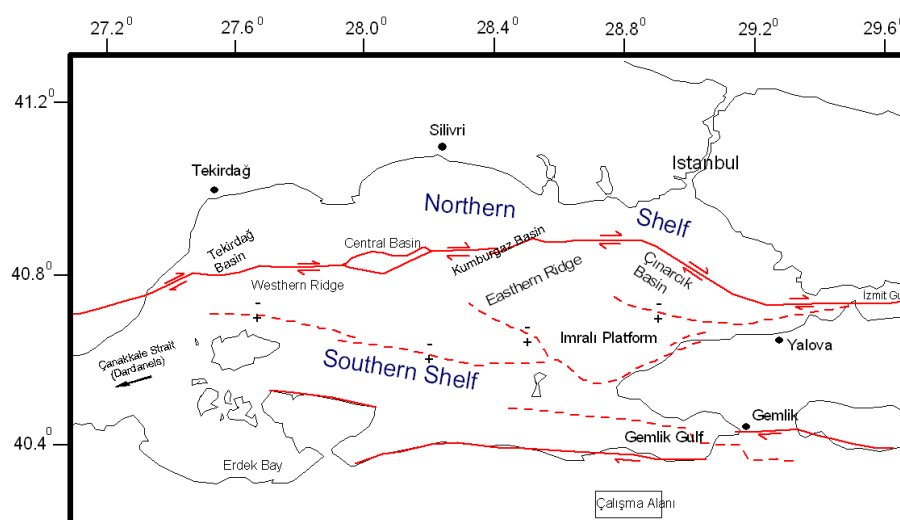


from the environment, increases suddenly. This sudden increase in pore water removes the contact forces that hold the soil particles together and removes the particles from each other. Thus, the soil loses its strength. Under these conditions, where the effective vertical pressure is zero, the soil acts as a liquid instead of a solid. This behaviour, which the soil presents as a result of dynamic loads, is known as liquefaction.

Both soil amplification and soil liquefaction developing on the soils due to earthquake effects are critical studies for microzonation studies of urban areas.

3. GEOLOGY TECTONICS OF THE REGION

The study area covers the central districts of Bursa province. These districts are connected to Bursa Metropolitan Municipality; Osmangazi, Kestel, Yıldırım, and Kestel districts. Regarding the historical earthquake records, the district has a remarkable feature as it is located in an area with high earthquake potential. The geology of the study area was prepared by Zetaş (2000). It is located in an ellipsoidal basin, known as the Bursa basin, which started its formation in the Neogene and continued during the Quaternary period and is in the direction of D-B. This basin is one of Turkey's major tectonic structures, formed in the effect of the North Anatolian Fault Zone. The main faults surrounding the study area are given in Figure 4.





245 **Figure 4. Main faults surrounding the study area (Le Pichon et al., 2001, Redrawn)**

246

247

248 Karakaya group rocks outcrop in the south and southeast of the study area, where the old-
 249 based (paleozoic) relationship cannot be seen. The conglomerate, micro conglomerate,
 250 sandstone, feldspar sandstone, micaceous sandstone, grovak, claystone, basic volcanic and
 251 limestone alternation are observed. It also contains a large amount of limestone (Permian)
 252 olistoliths in various sizes. It contains radiolarite, diabase and spilitic intermediate (Zetaş,
 253 2000). Alluvium: Bursa plain, which is a tectonic depression basin, is completely carried by
 254 rivers; blocks with detritic material such as gravel, sand, and silt. Alluvium thickness varies
 255 between 80m-200m. In the central part of the Bursa plain, the alluvium thickness varies
 256 between 140m-200m. It consists of clay with fine elements, silt, sand, and gravel. It is
 257 difficult to separate alluvium and Neogene units (Zetaş, 2000).

258

259 In order to explain the current features of the Marmara region, it is necessary to mention plate
 260 tectonic movement of Turkey and the surrounding area. Turkey is located on a mountain
 261 formation system we call the Alpine-Himalayan Belt (Ketin, 1977). Crustal movement occurs
 262 in the Turkish plate due to the young and active tectonics as a result of (2.0 - 2.5 cm)
 263 northward of the Arabian Peninsula.

264

265 **4. GEOTECHNICAL AND GEOPHYSICAL ANALYSES FOR REGION**

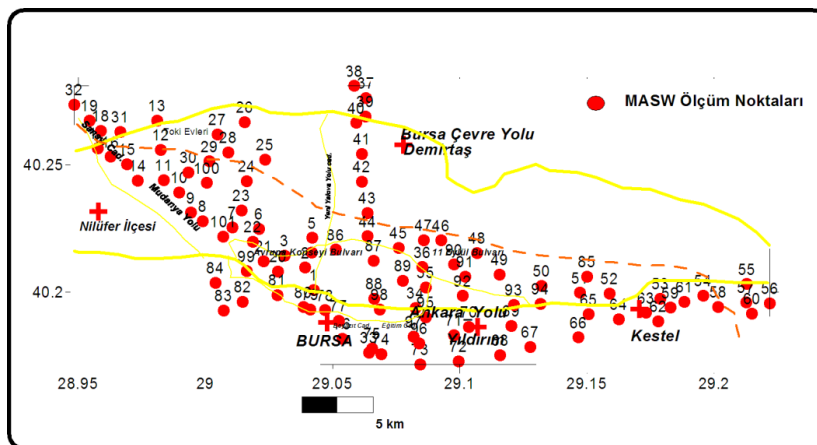
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267

268 For geophysical and geotechnical characterization, the view of the measurement points taken
 269 within the boundaries of the district is given in Figure 5a and b.



270

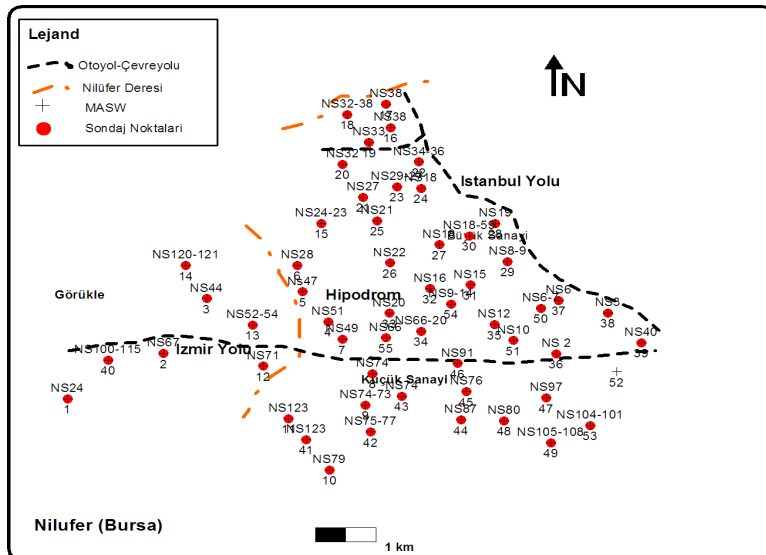


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272

273 **Figure 5a.** Location of the study area Geophysical (MASW) measurement points at
 274 Osmangazi ve Yıldırım

275



276

277 **Figure 5b.** Location of the study area and boreholes and the Geophysical (MASW)
 278 measurement points at Nilüfer District

279

280

281 4.1. Geophysical Factors

282

283

284 There are several geophysical measurements and evaluations used in field or site
 285 investigations. In this context, geophysical measurements can be divided into two groups as



exploration geophysics methods and engineering seismology studies. The most important exploration geophysics methods are seismic methods that are the main method to estimate V_s velocity. Engineering seismology studies include evaluating probabilistic and deterministic seismic hazards to estimate the ground motion level.

4.1.1. Seismic Hazard Analysis of the Region

Turkey is a region where there are interactions of Arab-Africa plates in the south and Eurasian plate in the North. Many important fault zones are effective within Turkey. The closest fault zone that will affect the study area is the North Anatolian fault zone. The approximate length of this fault zone is around 1200 km. In recent years, the city of Erzincan in eastern Turkey has seen very severe earthquakes (1939 and 1992). While more than 30.000 people died in the 1939 earthquake, 700 people died in 1992. With the collapse of several buildings and infrastructure, a high rate of property loss was observed. On August 17, 1999, an earthquake disaster called Kocaeli earthquake occurred around Izmit and Adapazarı, 110 km east of Istanbul. This earthquake, the magnitude of which was estimated to be between 7.4 and 7.6 by various researchers, damaged human life and property at a very high rate in the region. Another magnitude 7.2 earthquake occurred on November 12, 1999, in the northern Anatolian fault zone. Over a thousand people were dead or seriously injured. There are some estimations that the centres of these strong earthquakes are shifting from east to west along the North Anatolian Fault zone and draw attention to the possibility that a major earthquake could shake Istanbul, located at the western end of the North Anatolian Fault.

Probabilistic Earthquake Hazard Analysis for the region

The probabilistic earthquake hazard is defined as the probability that the damaging ground motion will occur in a certain place and within a certain period of time. For the analysis, it was obtained 4.5 and larger earthquakes during the instrumental period within an area of 100 km radius, focusing on the region from a database of Kandilli Observatory and Earthquake Research Institute. Earthquake hazard analysis was carried out using “SoilEngineering” software developed by Ozcep (2010). The results of this analysis are given in Figure 6 and in Table 1.a, 1.b, 1.c.



Table 1.a. Poisson Probability Distribution and Earthquake Hazard Analysis Data and Results.

Year Interval	98				
Magnitude Range	$4.5 \leq M < 5.0$	$5.0 \leq M < 5.5$	$5.5 \leq M < 6.0$	$6.0 \leq M < 6.5$	$7.0 \leq M < 7.5$
Ni (Frequency)	47	21	11	4	1
Mean Magnitude M or (Xi)	4,7	5,2	5,7	6,2	7,2

.a	3,02
.b	-0,66

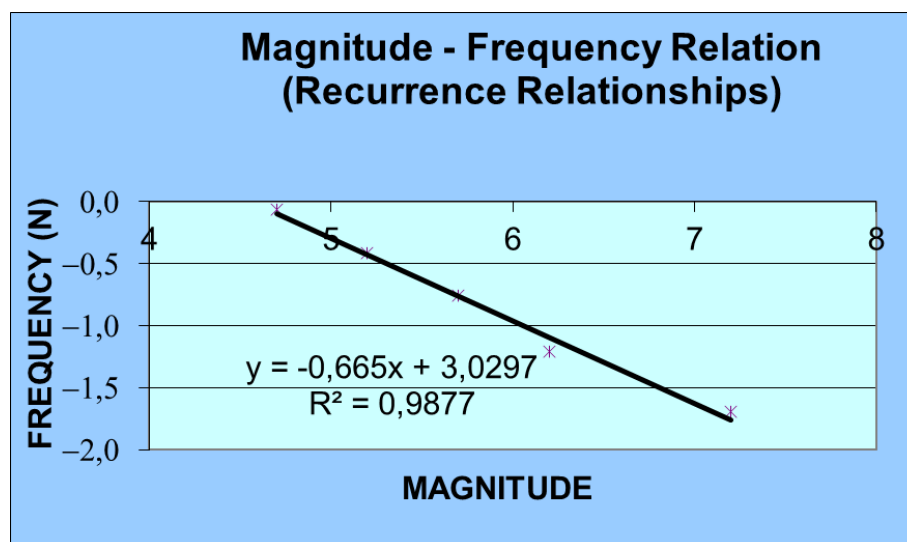


Figure 6. Magnitude – Frequency Relationship for Poisson probability model

Tablo 1.b Earthquake probability for various years and magnitudes estimated by Poisson Probability distribution

Magnitud	D = Probability for 10 years (%)	D = Probability for 50 years (%)	D = Probability for 75 years (%)	D = Probability for 100 years (%)	Recurrence Period (Year)
5	99.4	100.0	100.0	100.0	2
5.5	87.8	100.0	100.0	100.0	5
6	57.4	98.6	99.8	100.0	12
6.5	29.3	82.3	92.6	96.9	29
7	13.1	50.5	65.2	75.5	71
7.6	4.7	21.2	30.0	37.9	210



Tablo 1.c. Acceleration values in the epicentral distance for Probability of Exceedance (%) and design year

D (year)	Probability of Exceedance (%)	M (magnitude)
50	30	7,4

Δ , Epicentral Distance (km)	H, Focal depth (km)
25	15

Researcher	Donavan(1973c)	Oliviera (1974)	Joyner ve Boore (1981)	Campbell (1997)	Average
Acceleration (a)	0.23	0.16	0.39	0.42	0.30

Geophysical Analysis based on Field Studies

Within the coverage of this study, geophysical study (surface wave analysis (MASW) and shear wave velocity estimation) was carried out at 100 different points in the districts of Bursa Metropolitan Municipality. The average shear wave velocity (V_{s30}) map for the 30 meters obtained as a result of the measurement is given in Figure 7 for Osmangazi district and its surroundings (Region 1) and Figure 8. (for Nilüfer District (Region 2), soil fundamental period and soil amplification maps were given in Figure 9a, b, c, d, e. Some locations of the study area are not suitable to acquired geophysical data especially central parts of Osmangazi district and its surrounding area.

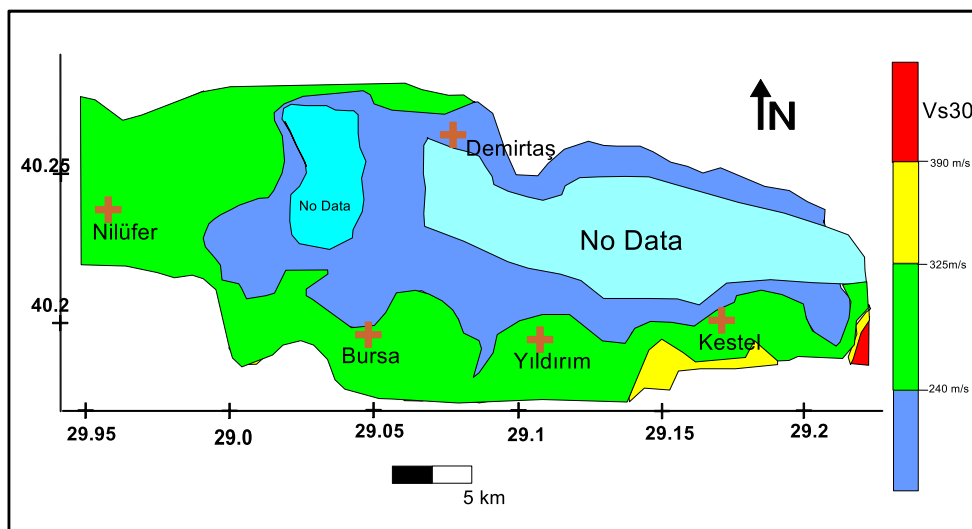


Figure 7. Bursa (Osmangazi, Yildirim, Gürsü, Kestel) Districts and VS30 map of Soils and measurement locations (black dots with numbers) acquired on the Osmangazi district.

It was observed there were three different velocity zones from the Vs30 distribution map of Osmangazi district and its surroundings (Fig. 7). On the basis of Vs 30 values of the study area can be classified as high (380m/sn - 450m/sn) , medium (260-380 m/sn-) and lower (< 260). S wave values related to regional topography. While lower values correspond to the basin, high and medium S wave values related to hill zones. The eastern part of Region 1, located in the part of the Nilüfer District, and in a narrow zone in the south of Region 1 have medium and higher shear - wave velocity values within the range 260m/sn – 440m/sn . The Vs 30 values are then decreasing to the west and to the north towards the central of Region 1. The Vs30 map in Region 1 shows very good agreement that low velocities in the central part correspond to alluvium in Bursa Plain.

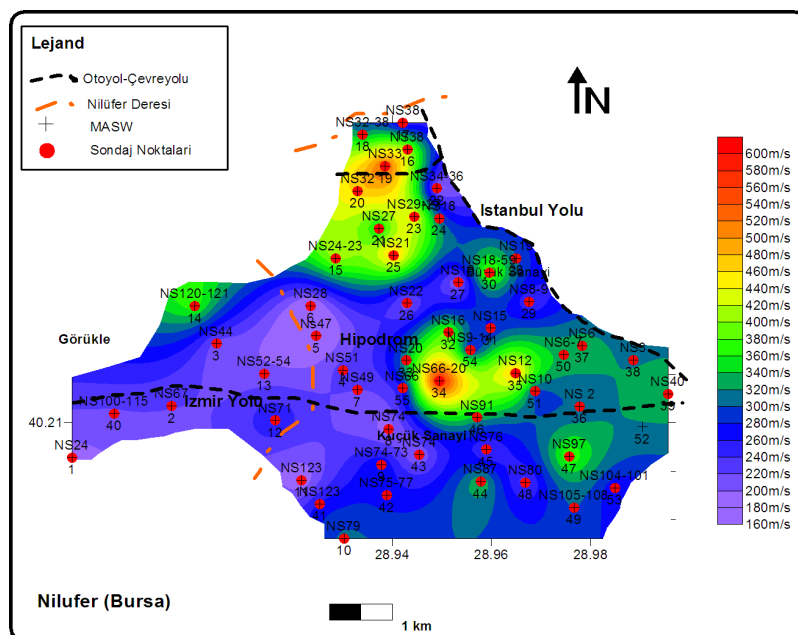


Figure 8. Bursa (Nilüfer) District and VS30 map of Soils

The Vs 30 distribution map of Region 2 shows three different velocity zones as Region 1. In contradistinction to, while higher velocity was obtained, ranging from 480-600m/sn, medium velocity values within range 340-480m/sn. In this region, only two local areas have high Vs velocities. While One of these local areas is the northern part of Region 2 that has 580m/sn shear velocity maximum, the other is the center of the region towards the east that has a maximum of 600m/sn higher Vs30 values. The lower Vs 30 velocities are calculated in all other areas. Lower values within range 160-280m/sn.

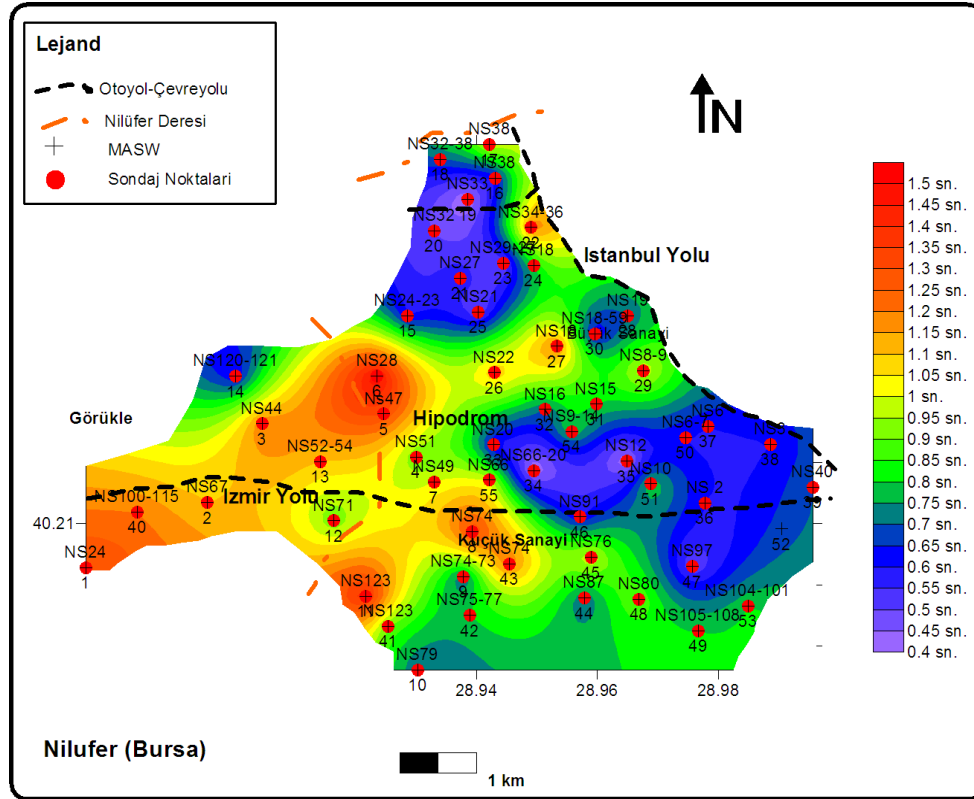


Figure 9a. Soil Fundamental Period Map in Nilüfer (Bursa) District

The contour map of the Soil Fundamental Period Map in Nilüfer (Bursa) District is computed from MASW and microtremor measurements. The fundamental period values of the sites located within areas with low $V_s 30$ seismic velocities are generally higher than areas with higher $V_s 30$ values. The soil fundamental period map in Nilüfer (Bursa) District, especially the lowest fundamental period values zones, coincides with the previous results of the $V_s 30$ distribution map. In the northern part of the area, the fundamental periods are computed between 0.4 - 0.7 sn, whereas at the sites in the alluvial basin, the values increase to 1.0 - 1.5sn. In addition, low fundamental periods of 0.4.- 0.7sn are also observed in the eastern part of the area.

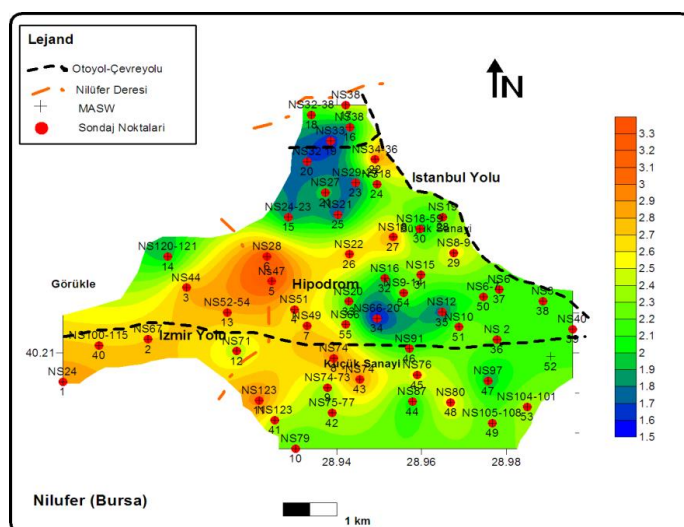


Figure 9b. Soil amplification maps accordin to Midorikawa (1987) Approach
 in Nilüfer (Bursa) District

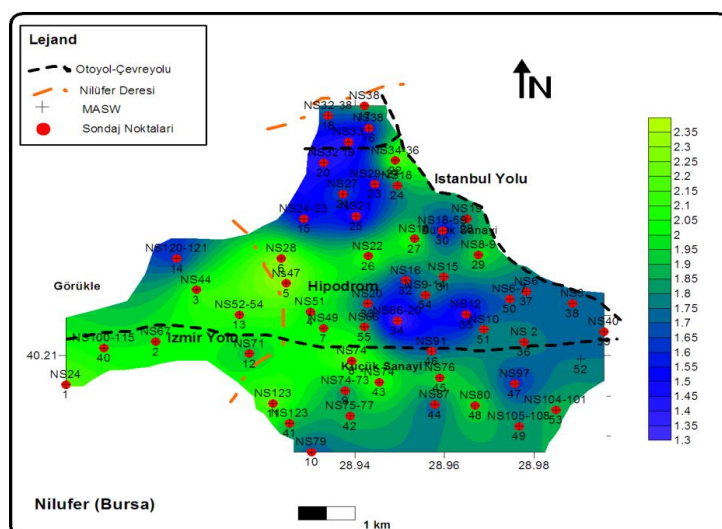


Figure 9c. Soil amplification maps according to Joyner and Fumal (1984) Approach in Nilüfer
 (Bursa) District

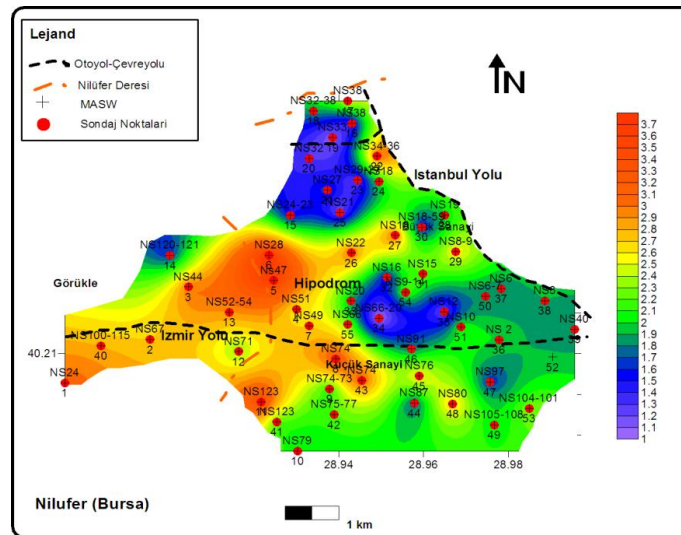


Figure 9d. Soil amplification maps according to Borchardt et al (1991) approach for strong motion in Nilüfer (Bursa) District

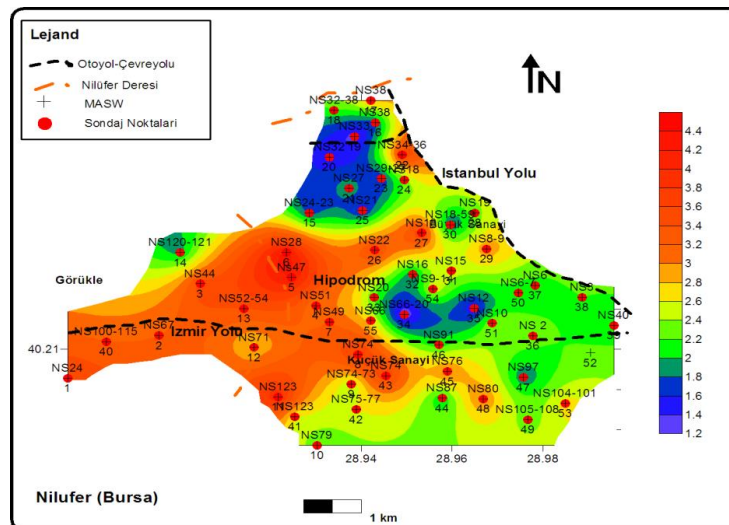


Figure 9e. Soil amplification maps according to Borchardt et al (1991) approach for weak motion in Nilüfer (Bursa) District



Figure 9d and e show comparisons between the soil amplification map according to various approaches (Joyner and Fumal (1984) ; Midorikawa (1987); Borchardt et al (1991) for strong and weak motion) . According to the map obtained from Joyner and Fumal's (1984) approach (Fig. 9d), while high soil amplification values are observed at the western and central part of the study area, low values are observed at the northern and eastern part of the area. This approach is very well compatible with the Vs 30 map and the fundamental period map. The high values calculated from other approaches also coincides with previous results of geophysical investigations. If the approaches are compared to each other, maximum soil amplification values are calculated from Borchardt et al (1991) approach for weak motion with these values ranging from 4.4 -3.0.

430

In other approaches, these maximum values within range 2.0- 3.7. According to Joyner and Fumal's (1984) approaches, soil amplification values are computed within the range of 1.3 - 2.35. These values are the minimum values in all approaches.

434

435

Soil-Earthquake Interaction Maps

437

Soil-earthquake interactive acceleration maps were made by taking into account Vs30 values obtained by surface wave velocity (MASW) analysis in Nilüfer District. The measurement points on the Vs30 map are divided into 3 regions considering that the NAF zone passes through the north. The shortest perpendicular distance to the measurement points in the northern part is 15 km, the measurement points in the central part are 20 km and the measurement points in the south are 25 km and the size chosen for the design is M: 7.6. Boore et al. (1997) spectral acceleration map was made with approach. Besides, mean acceleration map and spectral acceleration maps were made for periods of 0.0-0.2 and 1.0 seconds. These are given in Figure 9 f,g, and h.

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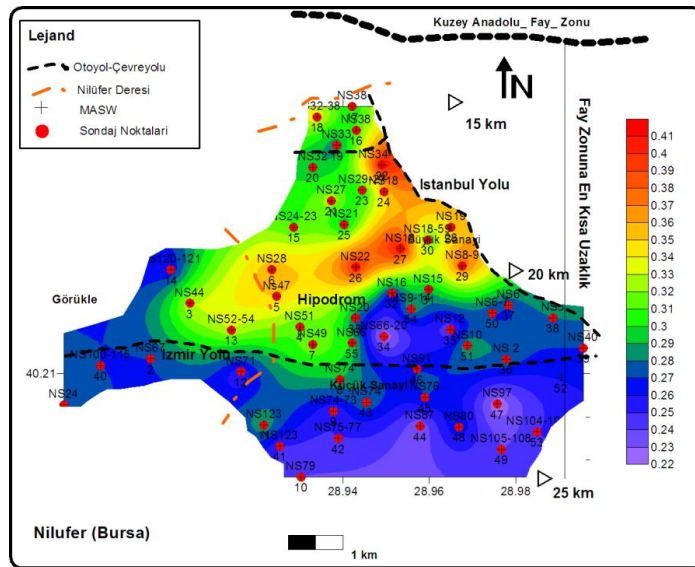


Figure 9 f. Spectral Acceleration Map with soil-earthquake interaction for 0.0 sec period in Nilüfer (Bursa) District

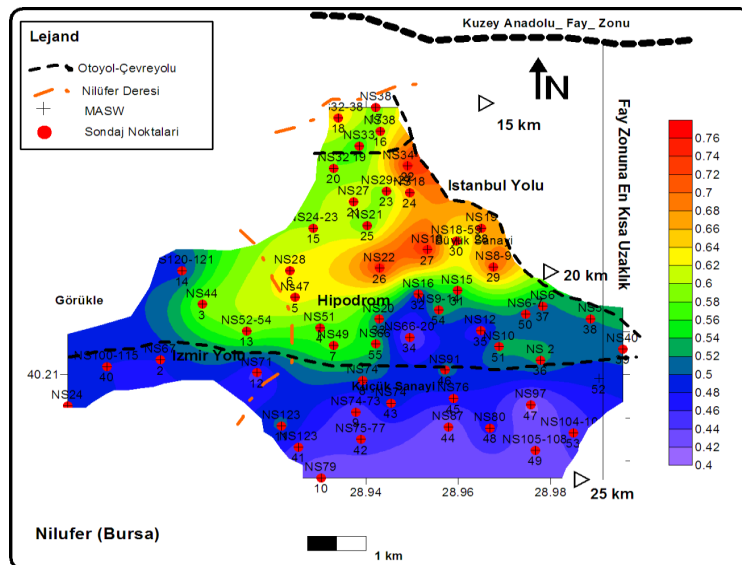
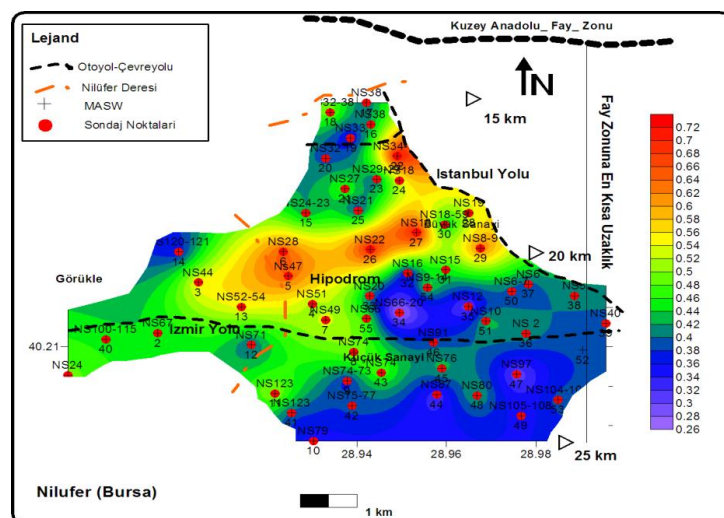


Figure 9 g. Spectral Acceleration Map with soil-earthquake interaction for 0.2 sec period in Nilüfer (Bursa) District



457



458

459 **Figure 9 h.** Spectral Acceleration Map with soil-earthquake interaction for 1.0 sec period in Nilüfer
 460 (Bursa) District

461

462 4.2. GEOTECHNICAL FACTORS

463

464 The shear strength of a soil is the highest shear stress that can be applied to the soil mass. The
 465 shear stresses that cause the shear stress along the shearing plane are the highest shear stress
 466 that the soil can bear, hence the shear strength. Accordingly, the shear strength is the limit
 467 value. The shear strength is the value that corresponds to the plastic equilibrium state, that is,
 468 it has no return and is a failure condition. Shear strength on the soil is formed by two
 469 properties of the soil:

470 (1) Internal Friction Angle (ϕ)

471 (2) Cohesion (c)

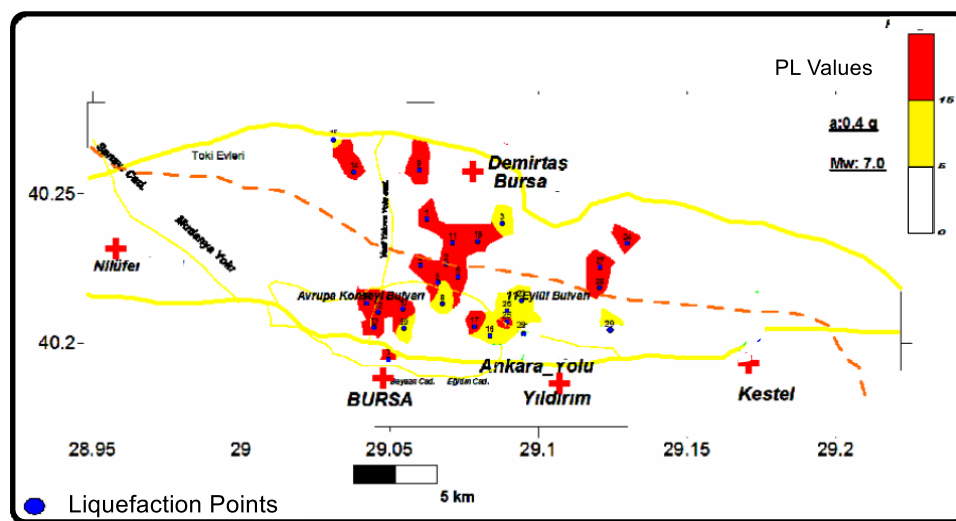
472 The important shear strength parameter of the soil is ϕ . Because it is a parameter that covers
 473 the behaviour for general and all soils compared to c . Friction is caused by the soil grains that
 474 are in contact with each other and the stress transfer between each other. The soil is a three-
 475 phase environment. So, soils are the three components of water, air, soil particles. In the soil



476 environment where these three components are mixed, the ratios of these components relative
 477 to each other are effective on soil behavior.

478 As geotechnical data for this study, over 200 boreholes (geotechnical) were provided from
 479 Bursa City Municipality. From this data, sand and silty sand units were extracted. Soil
 480 liquefaction analysis was performed based on geotechnical data. For this analysis,
 481 acceleration values (a : 0.40–0.45) and magnitude (M : 7.0–7.5) were selected. As a result,
 482 liquefaction analysis and liquefaction settlement analysis were estimated. Liquefaction
 483 Potential Index (PL) map is given according to several accelerations and magnitudes for
 484 Bursa (Osmangazi, Yıldırım, , Gürsü, Kestel) districts in Figure 10a and 10b and a map of
 485 settlements related to liquefaction, for Bursa (Osmangazi, Yıldırım, , Gürsü, Kestel) city (a :
 486 0.40- M : 7.0) is given in Figure 10c. For Nilüfer district, the same maps were given in Figure
 487 11a, b, c, d and e. The map of groundwater level for Nilüfer district was given in Figure 12.

488



489

490 Figure 10a Liquefaction Potential Index (PL) map according to acceleration and magnitude (a :
 491 0.40- M : 7.0) for Bursa (Osmangazi, Yıldırım, Gürsü, Kestel) city

492

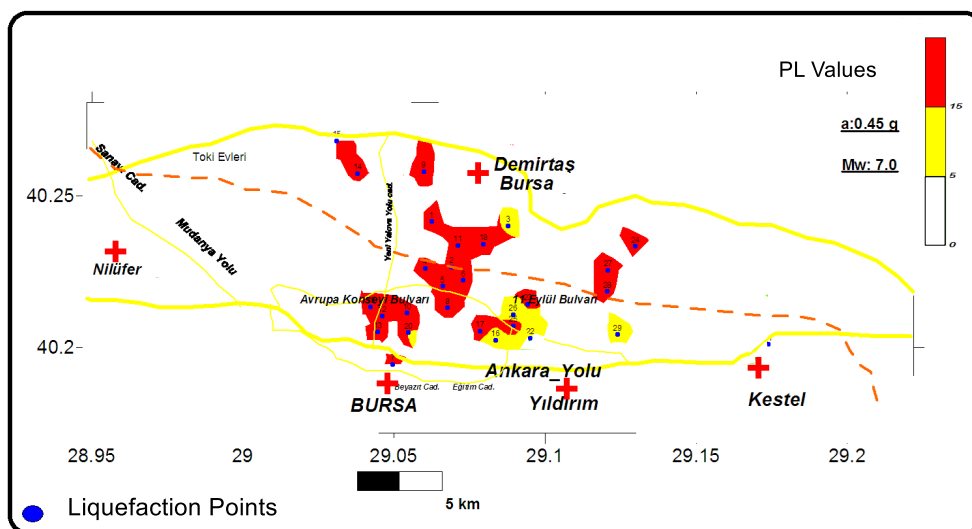
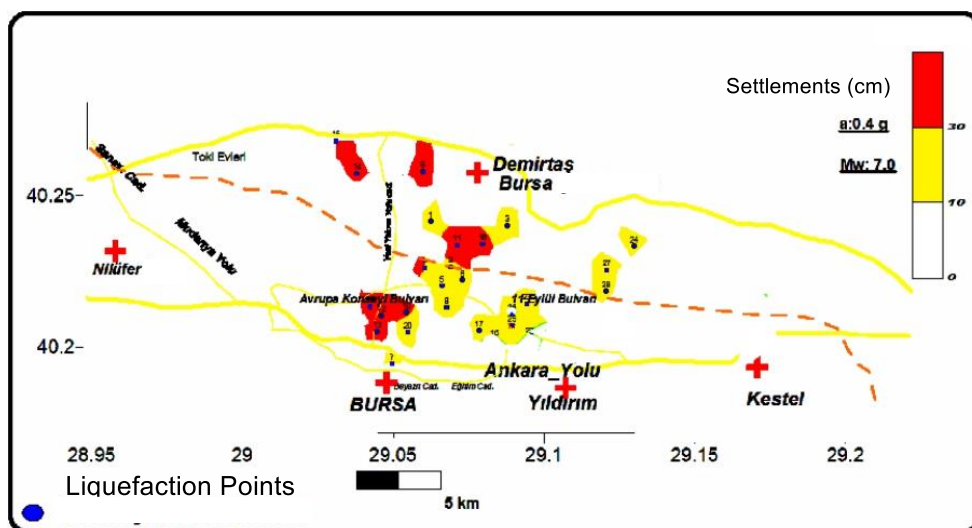
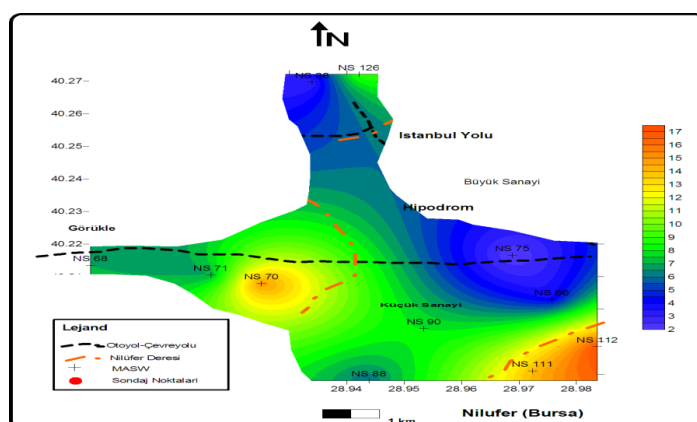


Figure 9b Liquefaction Potential Index (PL) map according to acceleration and magnitude ($a: 0.45$ - $M: 7.0$) for Bursa (Osmangazi, Yıldırım, Gürsü, Kestel) city



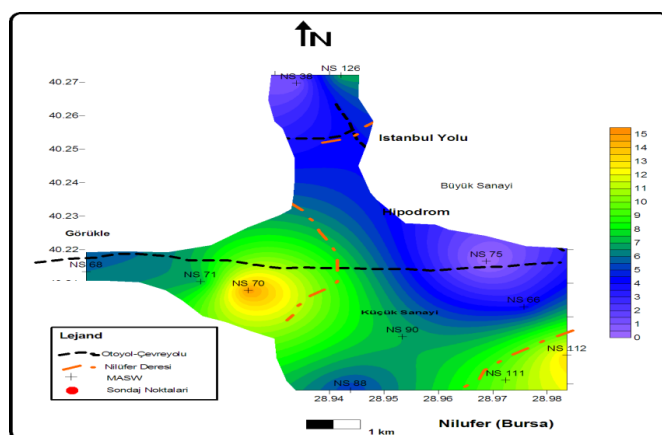


503 Figure 10c. Map of settlements related to liquefaction, for Bursa (Osmangazi, Yıldırım,
 504 Gürsü, Kestel) city (a: 0.40-M: 7.0).
 505



506

507 Figure 11a Liquefaction Potential Index (PL) map according to acceleration and magnitude (a:
 508 0.45-M: 7.6) in Nilüfer (Bursa) District.



509

510 Figure 11b Liquefaction Potential Index (PL) map according to acceleration and magnitude (a:
 511 0.40-M: 7.4) in Nilüfer (Bursa) District.

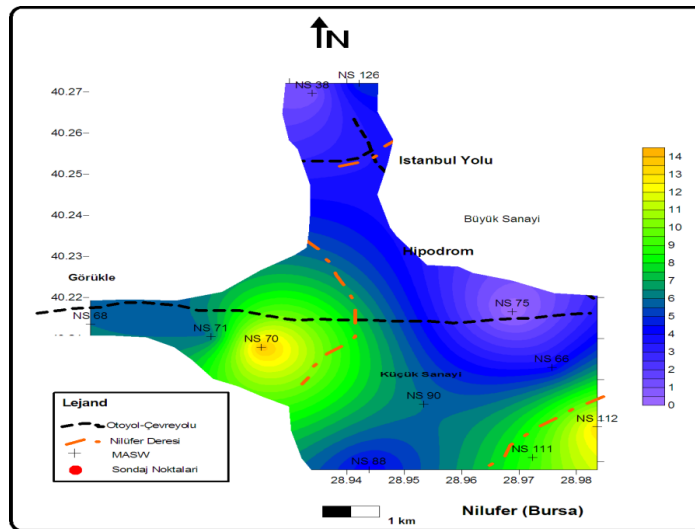


Figure 11c Liquefaction Potential Index (PL) map according to acceleration and magnitude (a: 0.40-M: 7.2) in Nilüfer (Bursa) District.

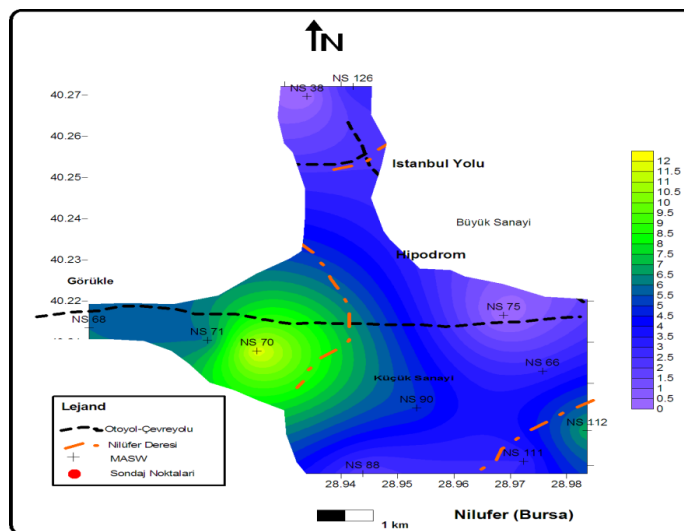
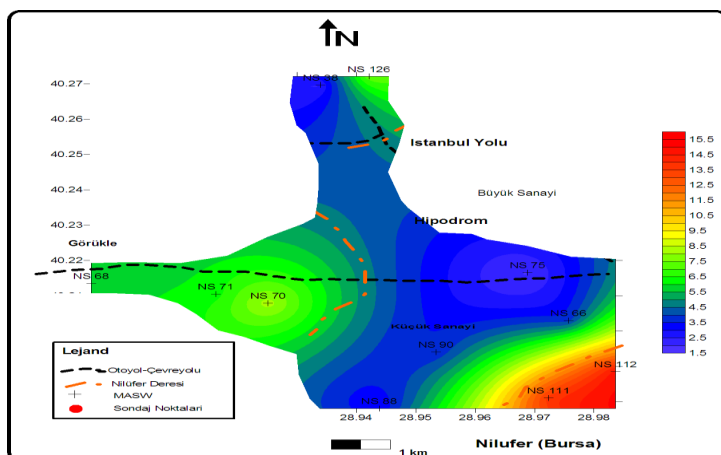
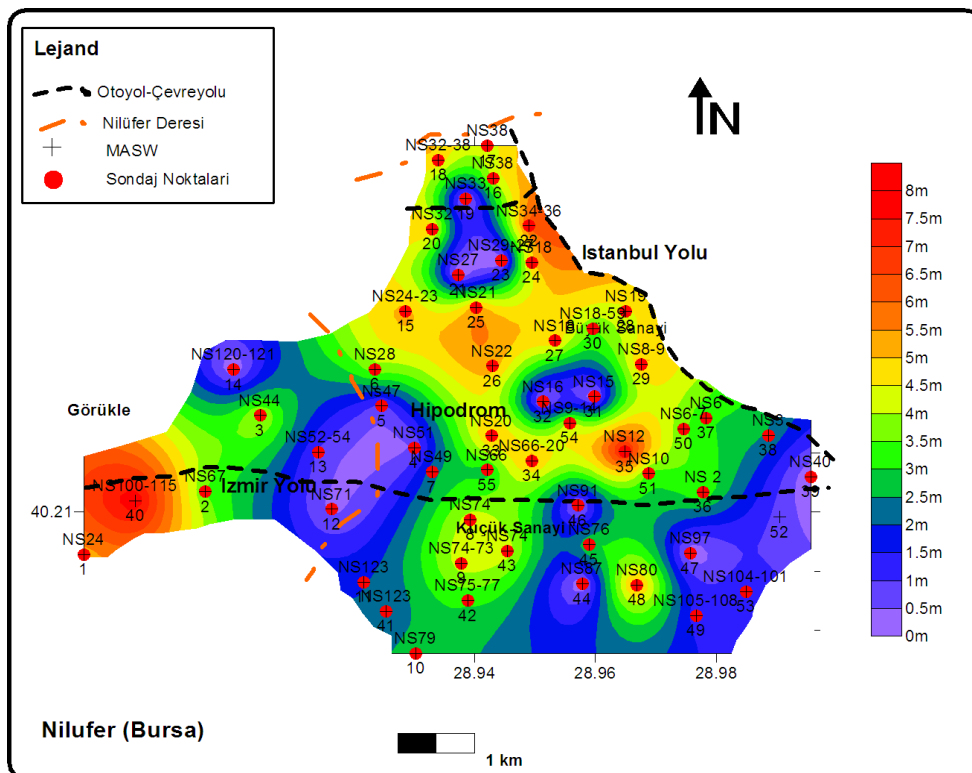


Figure 11d Liquefaction Potential Index (PL) map according to acceleration and magnitude (a: 0.35 M: 7.0) in Nilüfer (Bursa) District.



521

522 Figure 11f. Map of settlements related to liquefaction, in Nilüfer (Bursa) District



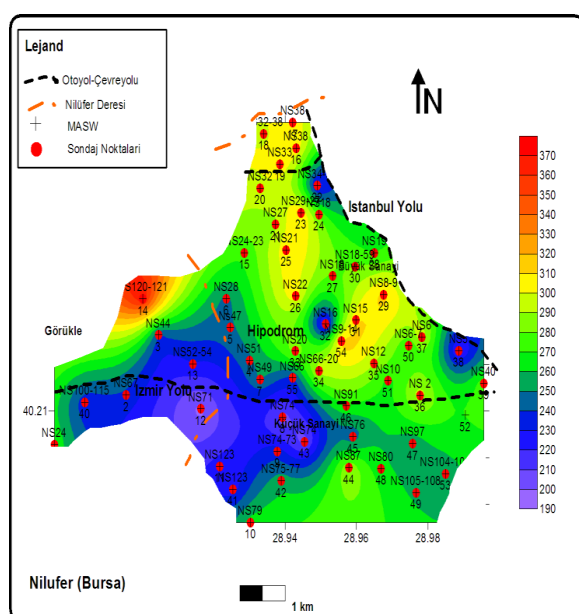
523

524 Figure 12. Map of Ground water level for Nilüfer (Bursa) District



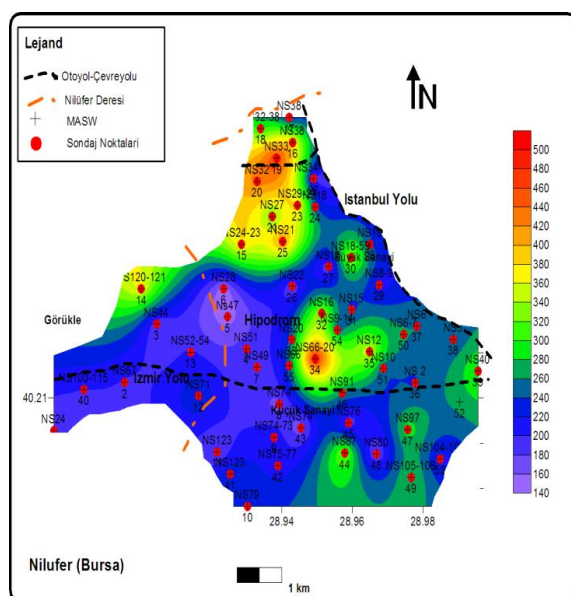
525

526 Vs15 values was calculated from SPT Data and comparisons of real (MASW) geophysical
 527 data. This and comparisons were given in Figure 13a and b.



528

529 Figure 13 a. Vs15 values, estimated from SPT Data in Nilüfer (Bursa) District

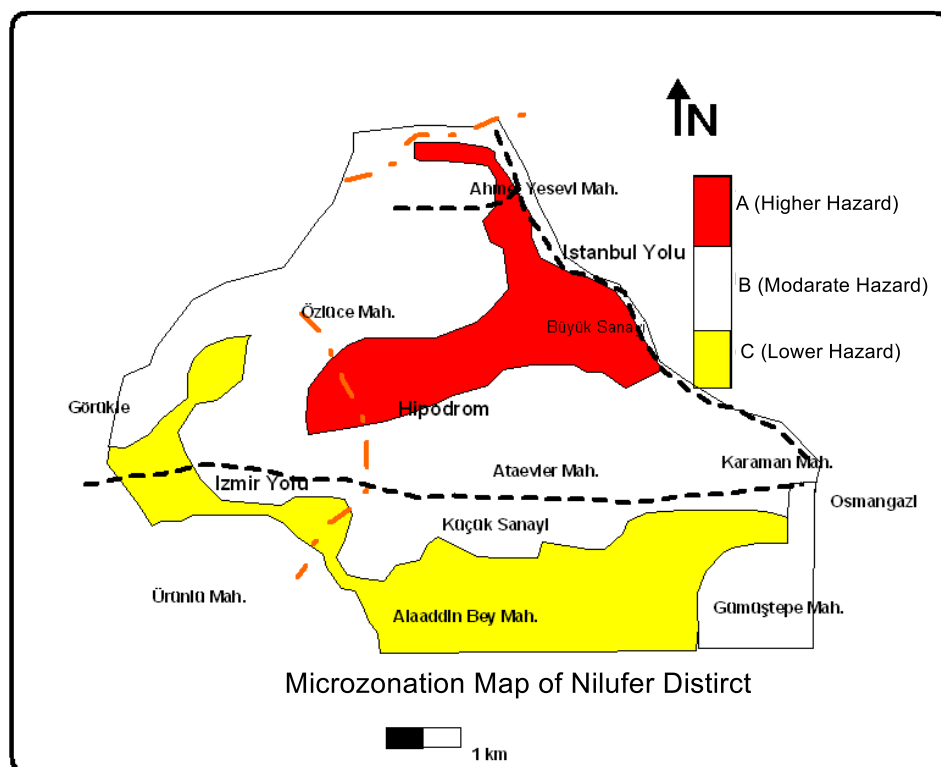


530

531 Figure 13b. Vs15 values, measured from Geophysical (MASW) Data in Nilüfer (Bursa)
 532 District

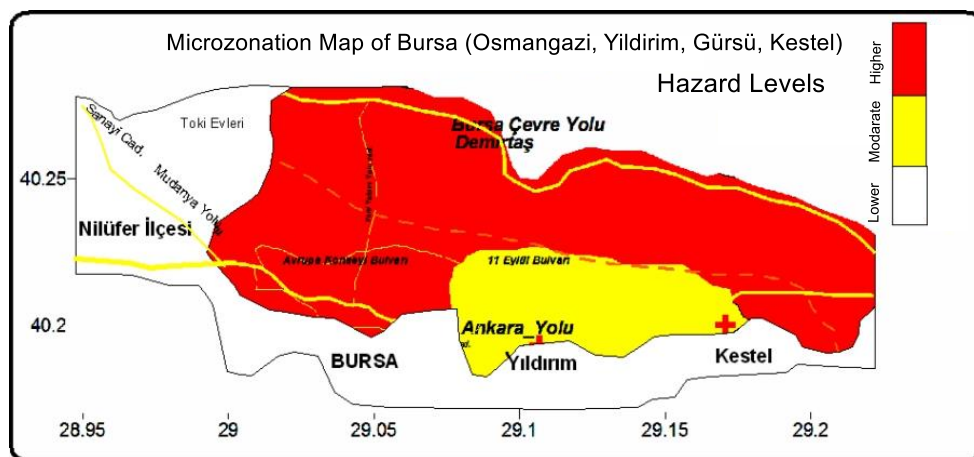
533 4.4. Preparation of Seismic-Micro-Zoning Map for Urban Planning

534 As a result of the geophysical and geotechnical studies, a microzonation map was prepared
 535 for a basis for urban planning (Figure 14a and b). As a result of the calculations, the hazard
 536 level for Bursa city is classified into 3 categories.



537

538 Figure 14a Microzonation map for a basis for urban planning in Bursa(Nillüfer) District



539

540 Figure 14b Microzonation map for a basis for urban planning in Bursa (Osmangaz, Yıldırım,
 541 Gürsü, Kestel) Districts



542
543

544 **5. RESULTS**

545
546

547 Geophysical measurements were made to estimate the shear wave velocity at 150 points in the
548 study area by using the MASW method. The shear wave velocity map was prepared for the
549 region. When we look at these maps, it was observed that Vs30, the type C type soils
550 (According to NEERP) were dominant and the other parts were represented by the D-E type
551 floors.

552

553 Earthquake hazard was determined probabilistically by using the Poison approach for the
554 study area. In the probabilistic analysis, the magnitude of the project earthquake has been
555 calculated for a certain time period for a certain exceeding rate. In the second phase of the
556 study, soil liquefaction and possible settlements due to liquefaction were estimated and
557 mapped by using geotechnical data. In the last phase, a final seismic microzonation map has
558 been created. In this study, the main purpose is to reveal the use of geophysical and
559 geotechnical data together in the context of microzonation.

560

561 In Bursa city, the levels of hazard for the problems that may occur interactively with the soils
562 were determined in the event of an earthquake. In new studies to be carried out in the context
563 of urban planning, urban transformation, and urban renewal, these hazard levels must be taken
564 into consideration.

565

566

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570

571

572 **6. References**

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