

5



# Assessing the impact of explosive eruptions of Fogo volcano (São Miguel, Azores) on the tourism economy

Joana Medeiros<sup>1,2</sup>\*, Rita Carmo<sup>1,2</sup>, Adriano Pimentel<sup>1,2,3</sup>, José Cabral Vieira<sup>4</sup>, Gabriela Queiroz<sup>1,2</sup>

<sup>1</sup>Centro de Informação e Vigilância Sismovulcânica dos Açores (CIVISA), 9501-801 Ponta Delgada, Portugal

<sup>2</sup>Instituto de Investigação em Vulcanologia e Avaliação de Riscos (IVAR), Universidade dos Açores, 9501-801 Ponta Delgada, Portugal

<sup>3</sup>Centro de Investigação em Biodiversidade e Recursos Genéticos (CIBIO), InBIO Laboratório Associado, Pólo dos Açores, 9501-801 Ponta Delgada, Portugal

<sup>4</sup>Faculdade de Economia e Gestão da Universidade dos Açores, Universidade dos Açores, 9501-801 Ponta Delgada, Portugal

10 Correspondence to: Joana Medeiros (Joana.FC.Medeiros@azores.gov.pt)

Abstract. The Azores is an active volcanic region that offers exceptional conditions for nature-based tourism, one of the main axes of economic growth in this region. A future volcanic eruption in the archipelago may have long-term consequences to this economic sector. Therefore, it is fundamental to assess its vulnerability to volcanic hazards in order to try to mitigate the associated risk. This study proposes a new approach to assess the economic impact of explosive eruptions

- 15 on the tourism sector. We considered two eruptive scenarios for Fogo volcano (São Miguel Island), the most probable (VEI 4 sub-Plinian eruption) and the worst-case (VEI 5 Plinian eruption), both producing tephra fallout and PDCs. The results of numerical simulations were overlaid with tourism-related buildings and infrastructure of Vila Franca do Campo municipality to identify the elements at risk. The Loss Present Value method was used to estimate the benefits generated by the accommodation units over 30 years for different economic scenarios. The assessment of the economic impact using 2018
- 20 indicators reveals that in a near total destruction scenario the economic loss is approximately 145 million euros. Such approach can also be adopted to other volcanic regions, other geologic hazards and other economic sectors.

Keywords: eruptive scenarios; economic impact; ocean island volcano, tephra fallout, pyroclastic density currents

# 25 1 Introduction

Among all natural phenomena on Earth, explosive volcanic eruptions are one of the most destructive and can cause major socio-economic impacts. Explosive eruptions can affect large areas of land, ocean, and airspace, threatening people, animals, buildings, infrastructure, transportation, communications, agriculture land, and water resources (e.g. Guffanti et al., 2005; Spence et al., 2005; Wilson et al., 2012; Scaini et al., 2014; Wilson et al., 2014; Craig et al., 2016; Brown et al., 2017;

30 Kueppers et al., 2019). In more extreme cases, explosive eruptions may also affect global climate (e.g. Rampino and Self, 1982; Hansen et al., 1992).





Despite the numerous hazards related to explosive volcanism, in several places volcanoes are regarded as attractions and not as a potential sources of problems. For some communities, nature-based tourism plays a significant role in the development of the local economy. However, the potential of some volcanoes to produce hazardous explosive eruptions is often overlooked or underestimated.

35

The 1995 eruption of Soufrière Hills volcano, on Montserrat Island (Lesser Antilles), is a dramatic example of the major impact that a volcanic eruption can have on a local community. This long-lasting eruption led to the island's largest migratory outbreak, with approximately 70 % of the population leaving Montserrat (Kokelaar, 2002; Annen and Wagner, 2003; Hicks and Few, 2015). Before 1995, Montserrat had a prosperous tourism industry, with revenues accounting for

- 40 approximately 25 % of the island's gross domestic product (GDP) (Caribbean Community Secretariat, 2009 in Pacheco and Lewis-Cameron, 2010). The eruption had a significant impact in this sector, with a decrease of roughly 50 % in arrivals between 1995 and 1996, reaching an all-time low in the following year, with a decrease of 44 % relative to 1996. The stabilization of the volcanic activity in 1998 led to a 50 % increase in arrivals compared to 1997. Although the eruption caused the destruction of critical infrastructure, including airport and harbour, and the capacity to accommodate tourists, this
- 45 sector has recovered steadily, as shown by the increase of tourist arrivals in 1997 and 1999, from 5000 visitors to approximately 10 000, respectively (Pacheco and Lewis-Cameron, 2010). Despite the last significant activity has occurred in 2010, officially it is still considered an on-going eruption (Wadge et al., 2014; Hicks and Few, 2015), yet tourism continues to grow with Soufrière Hills volcano representing one of the island's ex-libris (Pacheco and Lewis-Cameron, 2010). Taking another perspective, the 2010 eruption of Eyjafjallajökull volcano (Iceland) drew the attention to the vulnerability of
- 50 modern society to the atmospheric dispersal of ash plumes. The peculiar characteristics of this eruption combined with specific meteorological conditions resulted in the dispersal of volcanic ash over large areas of the North Atlantic Ocean and Europe (e.g. Gasteiger et al., 2011; Gudmundsson et al., 2012). Despite the moderate size and duration of the eruption, it caused unprecedented disruption to civil aviation over European airspace, with more than 8.5 million passengers stranded (Alexander, 2013). The aviation sector is extremely important to global economy, since it represents 0.7 % of the world's
- GDP and 35 % of world trade (Pálsson, 2010). The overall GDP loss resulting from the long-term incapacity to move people 55 and assets was estimated at approximately US \$ 4.7 billion, including airline industry losses and loss at destination, as well as general productivity losses (Oxford Economics, 2010; Pallister and McNutt, 2015). Although locally this eruption had a reduced impact, statistical data estimate that between April and May 2010 the number of tourists decreased by approximately 17.5 % in Iceland (Jónsdóttir, 2011).
- Oceanic islands are particularly vulnerable to volcanic eruptions and other geological hazards due to their typical remote 60 location, small size, and rough topography, which combined with high population densities and weak economies make risk management and evacuation very challenging (e.g. Pelling and Uitto, 2001; Clare et al., 2018; Kueppers et al., 2019). The Azores islands (North Atlantic Ocean) have an extensive geological record of explosive volcanic eruptions that typically occur at trachytic central volcanoes (Pimentel et al., 2015). Although the frequency of large explosive eruptions is relatively
- low, the impact of a future explosive eruption could be devastating, with long-term consequences to the regional economy. 65



70



As the Azorean tourism suffered a significant boost since 2015 with the alteration of the accessibility and mobility model, the importance of this sector in the regional economy has grown over the years. Therefore, it is fundamental to assess its vulnerability to volcanic hazards in order to implement appropriate mitigation strategies.

São Miguel Island is the largest (744 km<sup>2</sup>) and the most populated (> 137 000 inhabitants) of the Azores archipelago, where most of the tourism industry is concentrated. The volcanic history of São Miguel shows that it is the most active island in the Azores with the highest eruptive frequency of explosive events. At least 34 explosive eruptions (sub Plinian and Plinian) are

- Azores, with the highest eruptive frequency of explosive events. At least 34 explosive eruptions (sub-Plinian and Plinian) are known to have occurred in the last 5000 years. Fogo volcano, located in the central part of São Miguel, was the only volcano that produced a Plinian eruption in this time frame, although it has the lowest eruptive frequency of the three active central volcanoes of the island (Gaspar et al., 2015a). Nonetheless, its potential to produce hazardous explosive eruptions must not
- 75 be underestimated and the economic impact of future eruptions should be properly assessed. In this framework, we aim to evaluate the impact of future explosive eruptions of Fogo volcano on the tourism of São Miguel, taking Vila Franca do Campo municipality as study area. To assess which areas and exposed elements are susceptible to be affected by tephra fallout and pyroclastic density currents (PDCs), two eruptive scenarios were considered: the most probable scenario, corresponding to a sub-Plinian eruption; and the worst-case scenario, corresponding to a Plinian
- 80 eruption. The assessment was carried out on buildings related to tourism and allowed to estimate which exposed elements could be affected, with physical damage and loss of functionality. The methodology proposed consists in evaluating the benefits generated by the tourism industry, restricted to the quantification of revenues generated by accommodation units, to determine the current loss of that revenue over 30 years. Such analysis is especially important as tourism is a growing and promising activity in the Azores in general (Vieira et al., 2019), in São Miguel Island (Vieira and Antunes, 2017) and
- 85 certainly in Vila Franca do Campo, which due to the existence of bathing areas, a moderate climate, and maritime and recreational infrastructures attracts many visitors. This study represents the first attempt to quantify the economic loss related to future explosive eruptions of Fogo volcano on the island of São Miguel. However, the present approach may be adopted to other active volcanic regions and economic activities.

#### 90 2 Geographical and geological setting

The Azores archipelago is located in the North Atlantic Ocean and comprises nine volcanic islands. From a geodynamic point of view, this region corresponds to the triple junction where the North American, Eurasian and Nubian lithospheric plates meet (Fig. 1a,b). Due to this particular geodynamic framework, the archipelago is subjected to frequent seismic and volcanic activity (e.g. Gaspar et al., 2015b).

95 São Miguel Island is formed by three active central volcanoes, Sete Cidades, Fogo and Furnas, linked by Picos and Congro fissure volcanic systems. The eastern part of the island comprises the older and inactive Povoação volcano and Nordeste volcanic system (Fig. 1c). All three active central volcanoes are truncated by summit calderas related to paroxysmal





explosive eruptions. In the last 5000 years at least 33 sub-Plinian eruptions and one Plinian eruption are known to have occurred in São Miguel (Pacheco et al., 2013; Gaspar et al., 2015a).

- 100 Located in the central part of São Miguel, Fogo volcano (also known as Água de Pau volcano) is the largest of the three active central volcanoes of the island, reaching a maximum altitude of 947 m above sea level. The 3.2 km-wide summit caldera was formed by several collapse events and is presently occupied by a lake. Several volcanic structures are present inside the caldera (lava domes, pumice and tuff cones), as well as on the northern and southern flanks of the volcano (lava domes, scoria and pumice cones) (Wallenstein, 1999; Wallenstein et al., 2015).
- 105 The stratigraphy of Fogo volcano is divided into two major lithostratigraphic units (Wallenstein, 1999; Wallenstein et al., 2015): the Lower Group, which comprises all volcanic products older than 40 ka; and the Upper Group, which includes all products emitted over the last 40 ka, including from the historical eruptions. The recent eruptive period was marked by at least two paroxysmal explosive eruptions of Plinian dimensions: the Ribeira Chã eruption (8000–12 000 years ago) and the Fogo A eruption (~4600 years ago). In particular, the last 5000 years were characterized by five sub-Plinian eruptions: Pisão,

110 Fogo B, C, D, and 1563 (historic).

Fogo A was one of the largest eruptions recorded on São Miguel. Its deposit was studied by several authors (e.g. Walker and Croasdale, 1971; Booth et al., 1978; Bursik et al., 1992; Wallenstein, 1999; Pensa et al., 2015a,b) and encompasses a complex and widespread succession of trachytic pyroclastic products emitted ~4600 years ago during a Plinian eruption from the summit caldera. The eruption started with a short-lived phase of hydromagmatic activity, producing widespread fine ash.

- 115 This was followed by a steady Plinian eruptive column that produced a major pumice fall deposit. The radial and almost symmetrical distribution of the fall deposit around the caldera indicates weak wind blowing from the W during the eruption. Then the eruptive column experienced partial collapses that generated PDCs, leading to the emplacement of small-volume ignimbrites inter-bedded with pumice fall deposits. The final phase of the eruption was marked by the total collapse of the eruptive column and consequent emplacement of a voluminous ignimbrite that covered the flanks of the volcano, reaching >
- 20 m thick in the main valleys and especially in Ribeira Grande graben (Wallenstein, 1999; Pensa et al., 2015a,b; Wallenstein et al., 2015).

The last sub-Plinian eruption of Fogo volcano occurred in CE 1563. The deposit was studied by several authors (e.g. Walker and Croasdale, 1971; Booth et al., 1978; Wallenstein, 1999; Aguiar, 2018) and corresponds to a stratified succession of trachytic pumice and ash fall layers. This eruption started on June 28<sup>th</sup> in the centre of the caldera, on the previously existing

- 125 cone of Pico da Lagoinha, also known as Pico das Berlengas (Frutuoso [1522-1591†], 1981; Walker and Croasdale, 1971; Wallenstein, 1999). The first phase of the eruption was hydromagmatic, producing a fine ash layer. This was followed by the development of an unsteady sub-Plinian eruptive column with frequent hydromagmatic pulses. It produced a stratified sequence dominated by ash layers at the base and pumice lapilli layers towards the top. Tephra was mostly dispersed to the eastern part of the island due to predominant WSW-blowing wind. On June 29<sup>th</sup>, the wind changed direction to the NE
- 130





the north flank of Fogo volcano, four days after the onset of the sub-Plinian eruption. A subsequent phreatic explosion was also reported from inside the caldera in February 1564 (Frutuoso [1522-1591†], 1981; Wallenstein et al., 2015; Aguiar, 2018).

135 The present study area, Vila Franca do Campo municipality, is located on the southern flank of Fogo volcano and it borders with Lagoa municipality to the west, Ribeira Grande to the north and Povoação to the east. To the south it is bounded by the ocean in a costal extension of 15.5 km. Vila Franca do Campo has an area of 77.9 km<sup>2</sup> and is divided in six parishes: Água d'Alto, São Pedro, São Miguel, Ribeira Seca, Ribeira das Tainhas and Ponta Garça (Fig. 2).

#### 140 3 Methodology

#### 3.1 Definition of the eruptive scenarios

The first step in the assessment of the impact of future explosive eruptions of Fogo volcano on São Miguel Island was to define the eruptive scenarios. Based on the recent geological record of Fogo volcano (last 5000 years) and the frequency and magnitude of past trachytic explosive eruptions, two scenarios were defined: the most probable scenario (i.e. the most likely

145 eruption) and the worst-case scenario (i.e. the largest magnitude eruption), in agreement with previous studies (e.g. Gaspar et al., 2015a). The most probable scenario is a sub-Plinian eruption with Volcanic Explosivity Index (VEI, Newhall and Self, 1982) of 4, similar to the Fogo 1563 eruption. Such an eruption would produce widespread tephra fallout and could also generate PDCs in proximal areas, although this was not the case of the 1563 eruption. The worst-case scenario is a Plinian eruption with VEI 5, similar to the Fogo A eruption. Such an eruption would produce thick widespread tephra fallout and 150 generate voluminous PDCs along the flanks of the volcano.

Taking into account that in the Azores region there are significant differences in the wind patterns of summer (May to September) and winter periods (October to April) (see wind statistical analysis in Pimentel et al., 2006; Cole et al., 2008; Gaspar et al., 2015a) different tephra fallout scenarios should also be considered for the two periods.

# **3.2 Numerical simulations**

To identify which areas of São Miguel are susceptible to be affected by trachytic explosive eruptions of Fogo volcano, the dispersion of tephra fallout and PDCs was simulated using **VORIS** (Volcanic Risk Information System) version 2.0.1 (Felpeto et al., 2007) implemented in a geographic information system (GIS) (ArcGIS 9.1 ESRI®). VORIS 2.0.1 is a tool used in the assessment of volcanic hazards that provides users with the necessary instruments for the production of scenarios and hazard maps.





# 160 3.2.1 Tephra fallout

Numerical simulations of tephra fallout were computed using an advection-diffusion model that assumes that above the vent the mass is distributed along a vertical line following the Suzuki approach (Suzuki, 1983). Far from the vent, the transport of particles is controlled by the advective effect of the wind, the diffusion due to atmosphere turbulence and the terminal settling velocity of the particles (see details in Folch and Felpeto, 2005; Felpeto et al., 2007).

- 165 The eruptive source parameters used in the simulations were obtained from the literature related to Fogo A and Fogo 1563 eruptions, and when unavailable, from published charper of similar explosive eruptions. For the most probable scenario (a VEI 4 sub-Plinian eruption), we considered a total tephra volume of 1 km<sup>3</sup> (Booth et al., 1978) and a column height of 18 500 m (Carey and Sparks, 1986). For the worst-case scenario (a VEI 5 Plinian eruption), we used a total tephra volume of 3.2 km<sup>3</sup> (Booth et al., 1978) and a column height of 27 000 m (Bursik et al., 1992). The simulations were conducted assuming a vent located in the centre of Fogo caldera. Eruptive input parameters are shown in Table 1.
- Wind parameters were compiled from the Integrated Global Radiosonde Archive dataset of the National Climatic Data Centre for the Lajes/Santa Rita station, in the neighbouring island of Terceira, between 1947 and 2003. Statistical analysis of 56 years of radiosonde data performed by Pimentel et al. (2006) revealed significant differences between summer (May to September) and winter periods (October to April). In the troposphere (and lower levels of the stratosphere, i.e. up to 17 000
- 175 m altitude), the most frequent directions in the summer period are west to northwest and north to northwest blowing winds, whereas in winter period the dominant trend is west to southwest blowing winds. At higher altitudes (above 17 000 m), a strong eastern direction prevails during the summer period, while in the winter period a western direction dominates (Pimentel et al., 2006; Gaspar et al., 2015a). The most probable combinations of wind direction and intensity were chosen for different vertical heights, according to the simulated column height. Wind input parameters are summarized in Table 1.
- 180 In total four tephra fallout scenarios were simulated: 1) VEI 4 sub-Plinian eruption during the summer period; 2) VEI 4 sub-Plinian eruption during the winter period; 3) VEI 5 Plinian eruption during the summer period; and 4) VEI 5 Plinian eruption during the winter period.

#### **3.2.2 Pyroclastic Density Currents**

Simulations of PDCs were performed with the energy cone model (Malin and Sheridan, 1982), which provides a fast and conservative approach to assess the maximum potential extent of these volcanic products (e.g. Alberico et al., 2002, 2008). The energy cone model relates the height of the column collapse (H) and the maximum run-out distance of the PDCs (L), through the Heim coefficient, which translates the rheological properties of the currents. The energy cone inclination is given by the angle (θ) defined by the arctangent of H/L (Felpeto et al., 2007; Toyos et al., 2007). This method compares the

elevation of the energy cone, such the potential energy of a given PDC, with the topographic elevation. If the energy cone is higher than the surrounding topography, the area may be inundated by the current. Otherwise, the PDC may encounter a topographic obstacle that blocks its propagation (Alberico et al., 2008).





The maximum potential extent of a PDC is directly related to the VEI of the eruption and the topography around the vent. A higher VEI implies that the PDC can reach larger distances (Alberico et al., 2008). Input parameters used in the simulations were collapse equivalent heights of 300 and 500 m (for VEI 4 sub-Plinian and VEI 5 Plinian scenarios, respectively) and a constant collapse equivalent angle of 6°, in agreement with previous studies (e.g. Alberico et al., 2002, 2008, 2011; Cole et al., 2008). Both simulations were preformed assuming a source area equal to the floor of the caldera, where each 50 m cell had an equal probability of generating PDCs.

#### **3.3 Exposed elements**

The vulnerability of buildings and infrastructure to different volcanic products depends on the type of construction materials, 200 the quality of the workmanship, the age and maintenance level, their shape and orientation (Pomonis et al., 1999), but also on the loss of benefits from the activity for which they are intended to. Thus, in the particular case of explosive eruptions, we considered the vulnerability of buildings to tephra fallout and PDCs. In this study, we do not intend to assess the damage to the buildings but instead the loss of functionality.

To identify the exposed elements, a detailed inventory of all buildings and infrastructure related to the tourism sector in Vila 205 Franca do Campo municipality was carried out during the summer of 2017. This followed a similar approach to the studies carried out by Pomonis et al. (1999) in Furnas parish and Gomes et al. (2006) in the parishes located on the flanks of Sete Cidades volcano (São Miguel Island), and by Cabral (2015) in Santa Catarina municipality (Fogo Island, Cape Verde). The inventory included accommodation (hotels, guest houses, rural tourism, local accommodations), restaurants (ice cream

shops, pastry shops, cafes, pubs), tourism animation/activities (travel agencies, rent-a-cars, nautical and terrestrial activities,
souvenirs shops) and culture (churches and places of cult, museums, libraries, theatres, marketplaces, cultural centres).

- Although this study does not intend to assess building damage but rather its loss of functionality, the inventory was based on a classification method developed by the Centre for Volcanology and Geological Hazards Assessment (CVARG) of the University of the Azores, now Research Institute for Volcanology and Risk Assessment (IVAR), to study building vulnerability to different geological hazards in the Azores. This method classifies buildings according to use, number of
- 215 floors, type of materials used in the construction, roof inclination, type of windows, etc. Details on the classification method can be found in Gaspar et al. (2004) and Gomes et al. (2006). The exposed elements (i.e. buildings and infrastructure associated with tourism in Vila Franca do Campo) were mapped in detail in a GIS.
  The investerial elements are then explained with the end of the elements of the different elements (i.e. buildings and infrastructure associated with tourism in Vila Franca do Campo) were mapped in detail in a GIS.

The inventoried elements were then combined with the maps resulting from the different numerical simulations of tephra fallout and PDCs. This allowed us to identify which buildings and infrastructure would be affected by a certain thickness of tephra, as well as those located within the maximum potential extent of PDCs.

#### 3.4 Economic value of tourism

The methodology used to assess the impact of future explosive eruptions of Fogo volcano on the tourism economy consists in evaluating the benefits generated by the accommodation units of Vila Franca do Campo municipality. With that aim, we





225

255

determined the current loss of revenue over a period of 30 years for different eruptive scenarios and discount rates, and considered the evolution of the occupancy rate of the accommodation capacity. This study follows the approach proposed by Vianna et al. (2012) that quantifies the value of a tourism industry based on shark diving in the Republic of Palau (Pacific Ocean), one of the main activities that contributes to the country's economy.

In the present case and following the building inventory, we calculated the annual income of 46 accommodation units by estimating the average price per night of each unit. To calculate this value, we considered the price of one night in March

230 2018 (low season) and the cost of one night in August 2018 (high season). Some of the costs were obtained directly, provided by a representative of the accommodation unit, while others were obtained indirectly through online booking and shopping platforms, as well as through own websites. In specific cases, when was not possible to obtain this information, an average cost was calculated considering the type of accommodation.

The annual revenue was obtained by multiplying the average cost per night by 365 days for the cases where the rent

235 corresponds to the entire accommodation building. For the cases where the rent corresponds to only one room (e.g. hotels, apartments or chalets), the average cost of one room was multiplied by 365 days and by the number of rooms of the accommodation unit. The total annual revenue of the tourism sector in Vila Franca do Campo at 2018 values was obtained by adding up the annual revenues of each of the 46 accommodation units. However, this total revenue presupposes an occupancy rate of 100% of the existing capacity. To consider a more realistic situation, the total annual revenue was

240 multiplied by a factor of 0.5 and 0.65 to assume an occupancy rate of 50 % and 65 %, respectively. This originates two alternative scenarios for the existing accommodation capacity (explained later on). To calculate the economic loss in Vila Franca do Campo municipality due to future explosive eruptions, we considered three economic scenarios resulting from the numerical simulations: 1) Destruction of accommodation buildings affected by  $\geq 20$ cm of tephra from a VEI 4 sub-Plinian eruption during the summer period; 2) Destruction of accommodation buildings

- 245 located within the maximum potential extent of PDCs generated by a VEI 4 sub-Plinian eruption; and 3) Destruction of accommodation buildings located within the maximum potential extent of PDCs generated by a VEI 5 Plinian eruption. Economic scenarios 1 and 2 (tephra fallout and PDCs, respectively) from a VEI 4 sub-Plinian eruption are more conservative and realistic in case of a future explosive eruption of Fogo volcano. Regarding tephra fallout, the summer period was chosen
- 250 portion of the study area. Economic scenario 3 (PDCs from a VEI 5 Plinian eruption) is not the most likely but represents the worst case possible, assuming near total destruction of the study area. The annual revenue of the accommodation units was calculated for the three scenarios.

as it is the high season of touristic activity, but also because it corresponds to the wind conditions that would affect a larger

The next step was to calculate the Loss Present Value (LPV) for each of the economic scenarios. The LPV corresponds to the value lost (in Euros) after a certain period of time, discounted to the initial period (2018), which corresponds to year zero *t*<sub>0</sub>, following the expression:

$$LPV = \sum_{t_0}^{t_i} \frac{R_t}{(1+r)^t} \tag{1}$$





Where t is the year (0 to i),  $R_t$  is the revenue of the accommodation units in year t, taking into account the occupancy rate, and r is the discount rate.

The occupancy rate (ro) is the percentage of occupied units or rooms in a certain area during a specific period. The discount rate (r) is an intertemporal preference rate that allows to convert future values into present values, since a given monetary amount does not have the same value in the future. A discount rate equal to zero is used when it is intended to give the same weight to future and present values in a long-term analysis. This rate is considered a critical element in cost-benefit analysis when costs and benefits differ over time. In Portugal, this rate has varied between 0 and 6 % over the years. According to the Official Journal of the European Union (2006), in 2006 the discount rate in Portugal varied between 3.70 and 4.62 %

265 (January to December, respectively).

In this study, we considered a period of 30 years for the analysis, following the example of the eruption of Soufrière Hills volcano (Montserrat, Lesser Antilles) that started in 1995. This eruption also occurred on an island and produced tephra fallout, PDCs and lahars, with devastating consequences that last until today, even after more than two decades. Currently Montserrat's tourism industry is still in a recovery phase (Pacheco and Lewis-Cameron, 2010). The LPV was calculated for

each economic scenario, considering two occupancy rates of the existing accommodation capacity: one more conservative, with an occupancy rate of 50 % and discount rates of 2 and 4 %; and another less conservative, considering that the Azores tourism shows a growing tendency, with an occupancy rate of 65 % and the same discount rates (2 and 4 %).

## **4** Numerical simulations results

# 4.1 Tephra fallout from VEI 4 sub-Plinian eruption

- 275 The VEI 4 sub-Plinian simulation for the summer period (Fig. 3a) shows that Vila Franca do Campo, located southeast of Fogo caldera, is the most affected municipality of São Miguel, with > 3 m of tephra fall deposited in the caldera and immediately to east-southeast of the vent. Almost all the municipality is affected by tephra deposition up to 2 m thick. The predominant winds from west-northwest also lead to the deposition of tephra on the western half of Povoação municipality (between 50 cm and 1 mm, from west to east) and on the southern part of Ribeira Grande municipality (up to 3 m close to 280 the celdere).
- the caldera).

For the winter period (Fig. 3b), the prevailing winds from the west and the higher wind intensities promote the deposition of tephra on the eastern part of the island. Although Vila Franca do Campo is located on the southeast flank of the volcano, the strong westerly winds cause most of the tephra to be deposited on the northern sector of the municipality, with > 3 m immediately to the east of the caldera. However, in this scenario the entire Povoação municipality is affected by tephra fall

that can reach maximum thicknesses of 1 to 2 m in much of its area. Ribeira Grande and Nordeste municipalities are also affected on their southern sectors, although with smaller thicknesses (1 m to 1 mm, from south to north).





# 4.2 Tephra fallout from VEI 5 Plinian eruption

The VEI 5 Plinian simulations show the same dispersal patterns of the VEI 4 sub-Plinian simulations described above, but with greater extents. For the summer period (Fig. 4a), Vila Franca do Campo municipality is the most affected, with > 3 m of tephra deposited on its northern part, east-southeast of the caldera. Much of the municipality is affected by tephra deposition between 1 and 3 m thick. The neighbouring Povoação municipality is affected by tephra fall up to 1 m thick, particularly on the western side. The southern part of Ribeira Grande municipality is also affected by thicknesses of > 3 m of tephra close to the caldera.

For the winter period (Fig. 4b), the predominant stronger winds blowing from west lead to the deposition of tephra on the eastern part of São Miguel. Vila Franca do Campo is affected by thick tephra deposition (> 3 m) on the northern sector of the municipality. Given the higher intensities of the wind, tephra thicknesses of 1 to 3 m are deposited on much of Povoação municipality, reaching > 3 m on the western part. Ribeira Grande and Nordeste municipalities are also affected on their southern sectors, with thicknesses of 2 m to 1 mm, from south to north.

# 4.3 Pyroclastic Density Currents

The simulations of PDCs for the VEI 4 sub-Plinian and VEI 5 Plinian eruptions (Fig. 5) show the maximum potential extent of these volcanic products generated from within Fogo caldera. In both cases, PDCs are not contained inside the caldera and flow down the flanks of the volcano. The entire central part of São Miguel may be affected by PDCs, which reach the sea on the north and south coasts. As expected, the VEI 4 and VEI 5 simulations differ in the maximum potential extent that the PDCs can achieve. The VEI 4 simulation (blue dotted line in Fig. 5) shows that the western half of Vila Franca do Campo, most of Lagoa and the central part of Ribeira Grande municipalities are affected by PDCs, which can reach maximum distances of 9 km from the caldera. In the case of the VEI 5 simulation (red dashed line in Fig. 5), more than two-thirds of Vila Franca do Campo, all of Lagoa and the central part of Ribeira Grande municipalities are affected by PDCs that can reach maximum distances of 11 km from the source.

#### 5 Exposed elements at risk

- 310 The inventory and characterization of buildings related to tourism in Vila Franca do Campo municipality (140 in total) revealed that 46 correspond to accommodation units (33 %), 51 to restaurants (36 %), 29 are destined to culture or cultural activities (21 %) and 14 buildings are related to tourism animation/activities (10 %) (Fig. 6). Infrastructure such as the marina, the fishing port and the water park were also considered.
- The identification of the elements at risk was attained by overlapping the mapped exposed elements (buildings and 315 infrastructure) with the results of the simulations (Figs. 7 to 9). Regarding tephra fallout scenarios, it is important to distinguish the type of destruction that the buildings can suffer depending on the tephra thicknesses. When a building (represented by a polygon) was overlapped by more than one thickness class we considered the most conservative option, by





choosing the class with the higher thickness. The impact on the buildings was assessed according to the expected type of destruction following critical thickness thresholds.

- 320 Buildings affected by 1 mm to 20 cm of tephra and subject to constant and careful cleaning of the roof are usually not damaged or may suffer only minor damage. For the summer period (Figs. 7a and 8a), 62.1 % and 61.4% (VEI 4 sub-Plinian and VEI 5 Plinian scenarios, respectively) of all buildings related to tourism are likely to suffer thicknesses of up to 20 cm. On the other hand, for the winter period (Figs. 7b and 8b) buildings in Vila Franca do Campo municipality will only be covered by a maximum of 5 cm of tephra. The affected buildings are located in Ponta Garça, the easternmost parish, and
- 325 correspond to 5 % and 15 % (VEI 4 sub-Plinian and VEI 5 Plinian scenarios, respectively) of all touristic buildings (see Table 2).

Buildings covered by  $\ge 20$  cm of tephra can suffer significant damage, such as roof collapse (Pomonis et al., 1999). For both summer scenarios (Figs. 7a and 8a), approximately 31 % of the buildings are in these conditions. In more extreme cases, buildings affected by  $\ge 1$  m of tephra are likely to suffer total collapse (Blong, 1984; Spence et al., 2005). For this case, the

330 percentage of affected buildings is higher in the VEI 5 Plinian scenario with 19 %, whereas in the VEI 4 sub-Plinian scenario is 9.3 % (Table 2).

On what concerns PDCs, it should be noted that the distinction between dense and dilute currents was not taken into account in this analyses and therefore only total destruction was considered. The percentage of tourism-related buildings located within the maximum potential extent of PDCs (Fig. 9) from a VEI 4 sub-Plinian eruption is 87.1 %, whereas from a VEI 5

335 Plinian eruption is 95 %.

In addition to buildings, infrastructure related to the tourism sector in Vila Franca do Campo municipality will suffer similar damage to the buildings when affected by tephra fallout (summer period) and PDCs. Like the other exposed elements, infrastructure will not be affected by tephra fallout in the winter period.

#### 6 Assessment of the economic impact on tourism

340 The economic impact of future explosive eruptions of Fogo volcano on the tourism sector of Vila Franca do Campo municipality was calculated considering three economic scenarios. The number of accommodation buildings affected in each economic scenario is shown in Table 3.

The sum of the annual revenue of each accommodation unit resulted in an estimated total annual revenue of approximately 9.5 million euros, assuming that the accommodation capacity was fully occupied. Combining the revenue data with the

345 chosen occupancy and discount rates allowed to estimate the loss present values for a period of 30 years for each of the economic scenarios (Fig. 10). Table 4 summarizes the occupancy rates (ro), discount rates (r) and LPV for the three economic scenarios considered.

The graphs of figure 10 show that in all scenarios the loss present value is greater for higher occupancy rates and lower discount rates. For example, in economic scenario 3 (PDCs from a VEI 5 Plinian eruption) the loss present value after 30





- 350 years is almost 145 million euros, when the applied occupancy rate is 65 % and discount rate is 2 %. On the other hand, if an occupancy rate of 65 % and a discount rate of 4 % are considered, the loss is approximately 113 million euros. The lowest value of the updated loss is 87 million euros for an occupancy rate of 50 % and a discount rate of 4 %.
- In terms of LPV per year, there is a decreasing trend over time in all scenarios. For economic scenario 3 and considering both discount rates, the loss value in year zero does not reach 5 million euros, for occupancy rates of 50 %, and does not exceed 6 million euros, for a higher occupancy rate of 65 % (Fig. 10c). However, over the years the loss value downward trend is more evident when a discount rate is applied. For an occupancy rate of 65 %, the loss value in year 30 is 1.9 and 3.4 million euros for discount rates of 4 % and 2 %, respectively, while for an occupancy rate of 50 %, this value is approximately 1.4 and 2.6 million euros, considering discount rates of 4 % and 2 %, respectively.

### 7 Discussion

# 360 7.1 Impact of tephra fallout and PDCs from Fogo volcano

Fogo volcano is considered one of the most hazardous volcanoes in the Azores archipelago. Despite the low eruptive frequency of sub-Plinian and Plinian eruptions, with a recurrence interval of 833 years (Gaspar et al., 2015a), Fogo volcano produced the largest eruption in the Azores in the last 5000 years (Fogo A) and one of the most recent explosive eruptions of the archipelago (Fogo 1563) (Walker and Croasdale, 1971; Booth et al., 1978; Wallenstein, 1999). Even in non-eruptive periods, Fogo volcano is a hazardous landform with significant seismicity (Silva et al., 2015; 2020) and geothermal activity, manifested in fumarolic fields, CO<sub>2</sub>-cold and thermal springs, and diffuse degassing areas (Viveiros et al., 2015). Several unrest episodes have occurred in the last 20 years, with intense earthquake swarms lasting for several months and accompanied by ground deformation (inflation episodes), such as in May-September 2005 (Wallenstein et al., 2007; Silva et al., 2012). Due to its location, in the central part of São Miguel, a future explosive eruption will severely impact the parishes

- 370 located on the flanks of the volcano, but also others in more distal areas depending on the wind conditions. Such an eruption will have significant long-term economic consequences for the island and even for the entire Azores archipelago. Next, we discuss the impact of tephra fallout and PDCs on São Miguel, following the eruptive scenarios defined in this study. The simulation of tephra fallout from a VEI 4 sub-Plinian eruption (most probable scenario) during the summer period,
- shows that the prevailing winds blowing from west-northwest promote the deposition of tephra within the caldera and to the 375 east-southeast of the vent, making Vila Franca do Campo municipality the most affected, with tephra deposition up to 2 m thick. For an eruption occurring during the winter period, the predominant wind blowing from the west and the higher wind intensities, lead to the deposition of tephra on the eastern part of the island. Validation of the simulation results can be made by comparison with the geological record, namely with the deposit of the Fogo 1563 eruption, which has a well-defined dispersal axis towards the east (Walker and Croasdale, 1971; Aguiar, 2018). Regarding tephra fallout from a VEI 5 Plinian
- 380 eruption (worst-case scenario), the simulation results show the same dispersion patterns as the VEI 4 sub-Plinian eruption





scenario, for summer and winter conditions, though with thicker tephra deposition, due to the larger volume of ejected material.

The impact of tephra fallout on buildings and infrastructure will depend on the thickness of accumulated tephra, which is translated as static load. In localities affected by tephra fallout, the accumulation of 20 cm or more of tephra would be

- sufficient to cause significant damage (Pomonis et al., 1999) and where the 1 m threshold is exceeded buildings would likely suffer total collapse, including constructions reinforced with concrete (Blong, 1984; Spence et al., 2005). However, if tephra is wet these critical thickness thresholds are substantially reduced (Spence et al., 2005). This possibility should not be overlooked given the rainy Azorean climate (Hernández et al., 2016).
- Other elements such as the ground transportation network would also be affected by tephra fallout. Roads are particularly vulnerable as the accumulation of only a few millimeters of tephra decreases drivability, due to traction loss and limited visibility. In areas where the accumulation of tephra fall deposits would be substantial, roads can become impassable until they are cleared. Vehicles are also vulnerable to tephra fallout, as ash enters the engines, causes filters and brake systems to clog, and corrode paintwork and exterior fittings (Wilson et al., 2012; Wilson et al., 2014). Marine transportation can also be affected by tephra fallout as ports and marinas become inoperable and vessels can sink. In addition, the deposition of tephra
- 395 into the ocean can form pumice rafts that may stay afloat for weeks to months creating a hazard to navigation (Wilson et al., 2012; Jutzeler et al., 2014). Particularly during the summer period, marine operations along the south coast of São Miguel would be severely affected and important fishing ports such as Vila Franca do Campo and Ribeira Quente would be brought to a halt.

The simulations of PDCs, for VEI 4 sub-Plinian and VEI 5 Plinian eruptions, show that these currents are not contained

- 400 inside the caldera but are able to overcome the caldera walls and flow down the flanks of the volcano. The eruption of such volcanic products would affect the entire central part of São Miguel, reaching the sea on the north and south coasts of the island. The maximum potential extent of the PDCs is in the order of 9 and 11 km from the caldera for VEI 4 and VEI 5 scenario, respectively. It should be noted that in case of a real event, the progression of PDCs is strongly controlled by topography and channelled through valleys and depressions along the flanks of the volcano. Validation of these simulations
- 405 can be done by comparison with the geological record of PDC deposits. Fogo A ignimbrites were emplaced quasi radially outward from the caldera, within narrow paleo-valleys on the southern flank and along Ribeira Grande graben on the northern flank, reaching the sea in both cases (Pensa et al., 2015a,b).

The impact of PDCs on buildings and infrastructure is mostly related to their dynamic pressure and high temperature. In this case we assumed a binary impact approach which considered the absence of damage or total destruction of buildings and

410 infrastructure by PDCs. Ground and marine transportation networks are also highly vulnerable to PDCs as roads, ports and marinas would become buried or suffer irreparable heat damage. Vehicles and vessels would also suffer permanent physical damage if impacted by PDCs (c.f. Wilson et al., 2014).





# 7.2 Impact of explosive eruptions on the economy of tourism of Vila Franca do Campo

The main economic activities in the Azores are public services, retail and wholesale trade, fishing, livestock ranching, and 415 production of dairy products (Vieira et al., 2019). Tourism is a growing and promising activity for job creation and for the development of this archipelago, where the landscape and marine-related activities constitute the main attractions (Calado et al., 2011; Torres et al., 2017; Vieira and Antunes, 2017). The number of visitors sharply expanded after the liberalization of airspace in two corridors between the mainland and the Azores since 2015, which allowed the entrance of low-cost airlines (Vieira et al., 2019). According to satellite tourism accounts, tourism accounted for 6.7 % of the gross value added (GVA)

420 and 10 % of the employment in 2015. At the same time, the consumption of goods and services by tourists amounted to 14.1% of the GDP (Azorean Statistical Office, 2018).

The location of Vila Franca do Campo municipality in the southern part of São Miguel Island, approximately 20 km east of Ponta Delgada city (the island's capital), together with its moderate climate and numerous bathing areas along the coast attract many tourists throughout the year. Therefore, tourism is a major source of revenue for this municipality. In 2018, the

425 estimated total annual revenue of the accommodation units on Vila Franca do Campo was approximately 9.5 million euros (assuming that the accommodation capacity would be fully occupied). This figure clearly shows the importance of this sector on the economy of the municipality and of the entire island of São Miguel.

Looking at the calculated LPV for a period of 30 years for each of economic scenario, we can conclude that this value is greater when the occupancy rate is higher and the discount rate is lower. Economic scenario 1 (tephra fallout from a VEI 4

- 430 sub-Plinian eruption during the summer period) has the smallest LPVs for the different rates used. The other two economic scenarios (PDCs from a VEI 4 sub-Plinian eruption and PDCs from a VEI 5 Plinian eruption) have similar LPVs because they comprise 96 % and 98 % of the accommodation buildings, respectively (Table 3). Still, and as expected, economic scenario 3 represents the higher loss, almost 145 million euros (Table 4). As seen in the graphs of figure 10, the LPV shows a decreasing trend over the years, as result of the applied methodology.
- 435 The method used in this study is a first attempt to quantify the economic loss of the tourism sector resulting from future explosive eruptions of Fogo volcano. It should be noted that both the annual revenue and the LPV obtained for the tourism economy were estimated using only the values from the accommodation units of Vila Franca do Campo municipality. To achieve a more realistic estimation of the total revenue related to the tourism economy, the benefits of other sectors, such as restaurants or tourism animation/activities, should also be considered. The tourism industry also contributes to the economy
- 440 by generating jobs in hotels, guest houses, local accommodations, restaurants, souvenir shops, etc. and therefore the workers' salaries should also be quantified. Nevertheless, the present methodology may be adopted to calculate the loss of revenue of other sectors related to tourism and also applied to different volcanic regions vulnerable to explosive volcanism.





#### **8** Conclusions

In this study we present a new approach to quantify the impact of explosive volcanic eruptions on the tourism industry. The methodology consists in determining the economic loss related to future explosive eruptions, in this case of Fogo volcano (São Miguel Island), by estimating the benefits generated by tourism, in particular, by the accommodation units of Vila Franca do Campo municipality. However, such approach may also be adopted to other active volcanic regions and economic sectors.

In the present case, we considered two eruptive scenarios for Fogo volcano, the most probable scenario (a VEI 4 sub-Plinian

- 450 eruption) and the worst-case scenario (a VEI 5 Plinian eruption), both producing tephra fallout and PDCs. Furthermore, we evaluated the vulnerability of tourism-related buildings and infrastructure in Vila Franca do Campo to tephra fallout and PDCs by analysing their loss of functionality. The LPV method was used to estimate the benefits generated by accommodation units over a period of 30 years for different economic scenarios.
- The simulation of tephra fallout from a VEI 4 sub-Plinian eruption during the summer period shows that tephra deposition occurs to the east-southeast of Fogo caldera, while during the winter period the deposition is to the east of the vent, making Vila Franca do Campo municipality the most affected. The same occurs for the simulation of a VEI 5 Plinian eruption, where the dispersion patterns are similar to the VEI 4 sub-Plinian eruption scenario, although with a larger dispersion area and thicker tephra deposition due to the larger erupted volume. The simulations of PDCs show that the central part of São Miguel is the most affected, as currents flow down the flanks of the volcano, reaching maximum potential extent of 9 and 11 km from the caldera for VEI 4 sub-Plinian and VEI 5 Plinian eruptions, respectively.
- The assessment of the economic impact on the tourism sector, taking into account 2018 indicators, shows that economic scenario 1 (tephra fallout from a VEI 4 sub-Plinian eruption) has the lowest LPV when compared to economic scenarios 2 and 3 (PDCs from a VEI 4 sub-Plinian eruption and from a VEI 5 Plinian eruption, respectively), which have similar LPVs. Although economic scenario 3 is not the most likely scenario, as it represents near total destruction of Vila Franca do Campo
- 465 municipality, it represents the higher economic loss, with approximately 145 million euros over 30 years. Tourism is a growing industry worldwide and particularly in the Azores it has had an increasing importance in the regional economy since 2015. However, the Azores and other active volcanic regions are vulnerable to future eruptions, which in some cases may have long-term economic consequences. Volcanic hazard and risk assessment studies are therefore essential in areas where people live side by side with active volcanoes. The information obtained allows the competent authorities to
- 470 implement appropriate strategies to try to mitigate volcanic risk, such as land use planning, emergency management and post-disaster economic recovery planning.

#### **Author Contribution**

Joana Medeiros was responsible for the investigation, conceptualization, data curation, formal analysis, writing and preparing the manuscript with contributions from all co-authors. Rita Carmo was responsible for the supervision,





475 conceptualization and project administration. Adriano Pimentel was responsible for conceptualization and supervision. José Cabral Vieira was responsible for the implementation of the economic methodology. Gabriela Queiroz was responsible for the funding acquisition and project administration.

#### Acknowledgements

480 This work is supported by the ERUPÇÃO Project, financed by the Azores 2020 Operational Program, funded by FEDER and FSE Azores Regional Government (ACORES010145FEDER000016).

#### References

Aguiar, S.: Contribuição para o estudo da estratigrafia dos vulcões Fogo e Furnas (S. Miguel, Açores): implicações para a sua história eruptiva recente, M.Sc. thesis, Universidade dos Açores, 158 pp., 2018.

Alberico, I., Lirer, L., Petrosino, P., and Scandone, R.: A methodology for the evaluation of volcanic risk in Campi Flegrei,
J. Volcanol. Geoth. Res., 116, 6378, 2002.
Alberico, I., Lirer, L., Petrosino, P., and Scandone, R: Volcanic hazard and risk assessment from pyroclastic flows at Ischia

Island (Southern Italy), J. Volcanol. Geoth. Res., 171, 118136, doi: 10.1016/j.jvolgeores.2007.11.014, 2008.

Alberico, I., Petrosino, P., and Lirer, L.: Volcanic hazard and risk assessment in a multi-source volcanic area: the example of
Napoli city (Southern Italy), Nat. Hazards Earth Syst. Sci., 11, 10571070, doi: 10.5194/nhess-11-1057-2011, 2011.

Alexander, D.: Volcanic ash in the atmosphere and risks for civil aviation: a study in European crisis management, Int. J. Disast. Risk Sc., 4, 919, doi: 10.1007/s13753-013-0003-0, 2013.

Annen, C. and Wagner, J.: The impact of volcanic eruptions during the 1990s, Nat. Hazards Rev., 4, 169175, doi: 10.1061/(ASCE)1527-6988(2003)4:4(169), 2003.

- Blong, R. J.: Volcanic hazards, Academic Press, Australia, 424 pp., 1984.
  Booth, B., Croasdale, R., and Walker, G. P. L.: A quantitative study of five thousand years of volcanism on S. Miguel, Azores, Philos. T. R. Soc. Lond., 228, 271319, 1978.
  Brown, S. K., Jenkins, S. F., Sparks, R. S. J., Odbert, H., and Auker, M. R.: Volcanic fatalities database: analysis of volcanic threat with distance and victim classification, Journal of Applied Volcanology, 6, 15, doi: 10.1186/s13617-017-0067-4,
- 500 2017.

Cabral, J.: Avaliação dos perigos vulcânicos e fenómenos associados na Ilha do Fogo (Cabo Verde): implicações para o planeamento de emergência e ordenamento do território, Ph.D. thesis, Universidade dos Açores, 281 pp., 2015.

Bursik, M. I., Sparks, R. S. J., Gilbert, G. S., and Carey, S. N.: Sedimentation of tephra by volcanic plumes: I. Theory and its comparison with a study of the Fogo A plinian deposit, Sao Miguel (Azores), B. Volcanol., 54, 329344, 1992.





505 Calado, H., Ng, K., Borges, P., Alves, F., and Sousa, L.: Climate change and coastal tourism in the Azores archipelago, in: Disappearing Destinations: Climate Change and the Futures Challenges for Coastal Tourism, edited by Jones, A. J., and Phillips, M., CAB Intl, Wallingford, 111–124, 2011. Carey, S. and Sparks, R.S.J.: Quantitative models of the fallout and dispersal of tephra from volcanic eruption columns. B.

Volcanol., 48, 109-125, 1986.

510 Clare, M. A., Le Bas, T., Price, D. M., Hunt, J. E., Sear, D., Cartigny, M. J. B., Vellinga, A., Symons, W., Firth, C., and Cronin, S.: Complex and cascading triggering of submarine landslides and turbidity currents at volcanic islands revealed from integration of high-resolution onshore and offshore surveys, Front. Earth Sci., 6, 223. doi: 10.3389/feart.2018.00223, 2018.

Cole, P. D., Pacheco, J. M., Gunasekera, R., Queiroz, G., Gonçalves, P., and Gaspar, J. L.: Contrasting styles of explosive

- 515 eruption at Sete Cidades, São Miguel, Azores, in the last 5000 years: hazard implications from modelling, J. Volcanol. Geoth. Res., 178, 574-591, 2008.
  Cole, P. D., Queiroz, G., Wallenstein, N., Gaspar, J. L., Duncan, A. M., and Guest, J. E.: An historic subplinian/phreatomagmatic eruption: the 1630 AD eruption of Furnas volcano, São Miguel, Azores. J. Volcanol. Geotherm.
- Res., 69, 117-135, 1995.
  520 Craig, H., Wilson, T., Stewart, C., Outes, V., Villarosa, G., and Baxter, P.: Impacts to agriculture and critical infrastructure
- in Argentina after ashfall from the 2011 eruption of the Cordón Caulle volcanic complex: an assessment of published damage and function thresholds, Journal of Applied Volcanology, 5, 7, doi: 10.1186/s13617-016-0046-1, 2016. Dibben, C. and Chester, D. K.: Human vulnerability in volcanic environments: the case of Furnas, São Miguel, Azores, J. Volcanol. Geoth. Res., 92, 133150, 1999.
- EMODnet Bathymetry Consortium: EMODnet Digital Bathymetry (DTM). http://doi.org/10.12770/18ff0d48-b203-4a65-94a9-5fd8b0ec35f6. The GEBCO\_08 Grid, version 20100927, http://www.gebco.net, 2018.
   Felpeto, A., Marti, J., and Ortiz, R.: Automatic GIS-based system for volcanic hazard assessment, J. Volcanol. Geoth. Res., 166, 106116, 2007.

Folch, A. and Felpeto, A.: A coupled model for dispersal of tephra during sustained explosive eruptions, J. Volcanol. Geoth. 530 Res., 145, 337349, doi: 10.1016/j.jvolgeores.2005.01.010, 2005.

Frutuoso, G.: Livro Quarto das Saudades da Terra, in: Saudades da Terra, edited by Frutuoso, G., 2<sup>a</sup> ed., Instituto Cultural de Ponta Delgada, Ponta Delgada, 1522-1591<sup>†</sup>, 1981.

Gaspar, J. L., Goulart, C., Queiroz, G., Silveira, D., and Gomes, A.: Dynamic structure and data sets of a GIS database for geological risk analysis in the Azores volcanic islands, Nat. Hazards Earth Syst. Sci., 4, 233-242, doi: 10.5194/nhess-4-233-2004. 2004

535 2004, 2004.

Gaspar, J. L., Guest, J. E., Queiroz, G., Pacheco, J., Pimentel, A., Gomes, A., Marques, R., Felpeto, A., Ferreira, T., and Wallenstein, N.: Eruptive frequency and volcanic hazards zonation in São Miguel Island, Azores, in: Volcanic Geology of

241255, doi: 10.1016/j.jvolgeores.2004.12.011, 2005.





São Miguel Island (Azores Archipelago), Memoirs, 44, edited by Gaspar J. L., Guest, J. E., Duncan, A. M., Barriga, F. J. A. S., and Chester, D. K., Geological Society, London, 155166, doi: 10.1144/M44.12, 2015a.

540 Gaspar, J. L., Queiroz, G., Ferreira, T., Medeiros, A. R., Goulart, C., and Medeiros, J.: Earthquakes and volcanic eruptions in the Azores region: geodynamic implications from major historical events and instrumental seismicity, in: Volcanic Geology of São Miguel Island (Azores Archipelago), Memoirs, 44, edited by Gaspar J. L., Guest, J. E., Duncan, A. M., Barriga, F. J. A. S., and Chester, D. K., Geological Society, London, 3349, doi: 10.1144/M44.4, 2015b.

Gasteiger, J., Groß, S., Freudenthaler, V., and Wiegner, M.: Volcanic ash from Iceland over Munich: mass concentration retrieved from ground-based remote sensing measurements, Atmos. Chem. Phys., 11, 22092223, doi: 10.5194/acp-11-2209-

2011, 2011.

Gomes, A., Gaspar, J. L., and Queiroz, G.: Seismic vulnerability of dwellings at Sete Cidades Volcano (S. Miguel Island, Azores). Nat. Hazards Earth Syst. Sci., 6, 4148, doi: 10.5194/nhess-6-41-2006, 2006.

Gudmundsson, M. T., Thordarson, T., Höskuldsson, A., Larsen, G., Björnsson, H., Prata, F. J., Oddsson, B., Magnússon, E.,

- Högnadóttir, T., Petersen, G. N., Hayward, C. L., Stevenson, J. A., and Jónsdóttir, I.: Ash generation and distribution from the April-May 2010 eruption of Eyjafjallajökull, Iceland, Sci. Rep., 2, 572, doi: 10.1038/ srep00572, 2012.
  Guffanti, M., Ewert, J. W., Swanson, G., Gallina, G., and Bluth, G.: The volcanic-ash hazard to aviation during the 2003-2004 eruptive activity of Anatahan volcano, Commonwealth of the Northern Mariana Islands, J. Volcanol. Geoth. Res., 146,
- 555 Hernández, A., Kutiel, H., Trigo, R. M., Valente, M. A., Sigró, J., Cropper, T., and Santo, F. E.: New Azores archipelago daily precipitation dataset and its links with large-scale modes of climate variability. Int J Climatol 36(14), 4439–4454. doi: 10.1002/joc.4642, 2016.

Hicks, A. and Few, R.: Trajectories of social vulnerability during the Soufrière Hills volcanic crisis, Journal of Applied Volcanology, 4, 10, doi: 10.1186/s13617-015-0029-7, 2015.

Hansen, J., Lacis, A., Ruedy, R., and Sato, M.: Potential climate impact of Mount Pinatubo eruption, Geophys. Res. Lett., 19, 215218, doi: 10.1029/91GL02788, 1992.
Jutzeler, M., Marsh, R., Carey, R. J., White, J. D. L., Talling, P. J., and Karlstrom, L.: On the fate of pumice rafts formed during the 2012 Havre submarine eruption, Nat. Commun., 5, 3660, doi: 10.1038/ncomms4660, 2014.

Jónsdóttir, A.: Impact of Eyjafjallajökull on tourism and internacional flights, B.S. thesis, Líf-ogumhverfisvísindadeild, 565 Verkfræði-ognáttúruvísindasvið, HáskóliÍslands, Reykjavík, 57 pp., 2011.

Kokelaar, B.: Setting, chronology and consequences of the eruption of Soufrière Hills Volcano, Montserrat (1995–1999), in:The Eruption of Soufrière Hills Volcano, Montserrat, from 1995 to 1999, Memoirs, 21, edited by Druitt, T. H. and Kokelaar,B. P., Geological Society, London, 143, 2002.

Kueppers, U., Pimentel, A., Ellis, B., Forni, F., Neukampf, J., Pacheco, J., Perugini, D., and Queiroz, G.: Biased volcanic

570 hazard assessment due to incomplete eruption records on ocean islands: an example of Sete Cidades Volcano, Azores, Front. Earth Sci., 7, 122, doi: 10.3389/feart.2019.00122, 2019.





Malin, M. C. and Sheridan, M. F.: Computer-assisted mapping of pyroclastic surges, Science, 217, 637640, doi: 10.1126/science.217.4560.637, 1982.

Newhall, C. G. and Self, S.: The Volcanic Explosivity Index (VEI): an estimate of explosive magnitude for historical volcanism, J. Geophys. Res., 87, 12311238, doi: 10.1029/JC087iC02p01231, 1982.

- Official Journal of the European Union: Commission notice on current State aid recovery interest rates and reference/discount rates for 25 Member States applicable as from 1 January 2007 Published in accordance with Article 10 of Commission Regulation (EC) 794/2004 of 21 April 2004 (OJ L 140, 30.4.2004, p. 1) and the Commission notice on the method for setting the reference and discount rates (OJ C 273, 9.9.1997, p. 3). OJ C 317, 23.12.2006: 22 (ES, CS, DA, DE,
- ET, EL, EN, FR, IT, LV, LT, HU, NL, PL, PT, SK, SL, FI, SV), 2006.
   Oxford Economics: The economic impacts of air travel restrictions due to volcanic ash, A report prepared for Airbus, 12 pp., 2010.

Pacheco, G. B. and Lewis-Cameron, A.: Weathering the storm – crisis marketing for small island tourist destinations, in: Marketing island destinations, edited by Lewis-Cameron, A. and Roberts, S., Elsevier, London, 149164, 2010.

- 585 Pacheco, J. M., Ferreira, T., Queiroz, G., Wallenstein, N., Coutinho, R., Cruz, J. V., Pimentel, A., Silva, R., Gaspar, J. L., and Goulart, C.: Notas sobre a Geologia do Arquipélago dos Açores, in: Geologia de Portugal, Vol. 2, edited by Dias, R., Araújo, A., Terrinha, P., and Kullberg, J. C., Escolar Editora, Vila Franca de Xira, 595690, 2013. Pallister, J. and McNutt, S. R.: Synthesis of Volcano Monitoring. in: The Encyclopedia of Volcanoes, Second Edition, edited
- by Sigurdsson, H., Houghton, B. F., McNutt, S. R., Rymer, H., and Stix, J., Academic Press, London, 11501171, 2015.
- 590 Pálsson, T.: Atlantic Conference on Eyjafjallajökull and Aviation Conference Summary and Conclusions, Keilir Aviation Academy, Keflavik Airport, Iceland, 15-16 September 2010, 2010.
   Delling, M. and Litte, L. L. Swell, island, developing, attack, natural disector, enlageshility, and alabel, shares. Clabel

Pelling, M. and Uitto, J. I.: Small island developing states: natural disaster vulnerability and global change, Global Environmental Change Part B: Environmental Hazards, 3, 4962, doi: 10.1016/S1464-2867(01)00018-3, 2001.

Pensa, A., Cas, R., Giordano, G., Porreca, M., and Wallenstein, N.: Transition from steady to unsteady Plinian eruption 595 column: the VEI 5, 4.6 ka Fogo A Plinian eruption, São Miguel, Azores, J. Volcanol. Geoth. Res., 305, 118, 2015a.

Pensa, A., Giordano, G., Cas, R., and Porreca, M.: Thermal state and implications for eruptive styles of the intra-Plinian and climactic ignimbrites of the 4.6 ka Fogo A eruption sequence, São Miguel, Azores, B. Volcanol., 77, 99, doi: 10.1007/s00445-015-0983-2, 2015b.

Pimentel, A., Pacheco, J. M., and Felpeto, A.: Influence of wind patterns on the dispersal of volcanic plumes in the Azores

region: test study of the 1630 eruption of Furnas Volcano (S. Miguel, Azores), EGU General Assembly, Vienna, Austria, 2–
7 April 2006, EGU06-A-04983, 2006.

Pimentel, A., Pacheco, J., and Self, S.: The ~1000-years BP explosive eruption of Caldeira volcano (Faial. Azores): the first stage of incremental caldera formation, B. Volcanol., 77, 42, doi: 10.1007/s00445-015-0930-2, 2015.

Pomonis, A., Spence, R., and Baxter, P.: Risk assessment of residential buildings for an eruption of Furnas volcano, São Miguel, the Azores, J. Volcanol. Geoth. Res., 92, 107131, 1999.





Rampino, M. R. and Self, S.: Historic eruptions in Tambora (1815), Krakatau (1883), and Agung (1963), their stratospheric aerosols, and climatic impact, Quaternary Research, 18, 127143, doi: 10.1016/0033-5894(82)90065-5, 1982. Scaini, C., Biass, S., Galderisi, A., Bonadonna, C., Folch, A., Smith, K., and Höskuldsson, A.: A multi-scale risk assessment for tephra fallout and airborne concentration from multiple Icelandic volcanoes Part 2: vulnerability and impact, Nat.

- Hazards Earth Syst. Sci., 14, 22892312, doi: 10.5194/nhess-14-2289-2014, 2014.
  Silva, R., Havskov, J., Bean, C., and Wallenstein, N.: Seismic swarms, fault plane solutions and stress tensors for São Miguel Island central region (Azores). Journal of Seismology, 16, 389–407, doi: 10.1007/s10950-012-9275-x, 2012.
  Silva, R., Ferreira, T., Medeiros, A., Carmo, R., Luís, R., Wallenstein, N., and Sousa, R.: Seismic activity on S. Miguel Island volcano-tectonic structures (Azores archipelago), in: Volcanic Geology of São Miguel Island (Azores Archipelago),
- 615 Memoirs, 44, edited by Gaspar, J. L., Guest, J. E., Duncan, A. M., Barriga, F. J. A. S. and Chester, D. K., Geological Society, London, 227238, doi: 10.1144/M44.17, 2015. Silva, R., Carmo, R., and Marques, R.: Characterization of the tectonic origins of historical and modern seismic events and their societal impact on the Azores Archipelago, Portugal, Geol. Soc. Spec. Publ., 501, doi: 10.1144/SP501-2019-106, 2020.

Spence, R. J. S., Kelman, I., Baxter, P. J., Zuccaro, G., and Petrazzuoli, S.: Residential building and occupant vulnerability to

- tephra fall, Nat. Hazards Earth Syst. Sci., 5, 477494, doi: 10.5194/nhess-5-477-2005, 2005.
  Suzuki, T.: A theoretical model for dispersion of tephra, in: Arc volcanism: physics and tectonics, edited by Shimozuru, D. and Yokoyama, I., Terra Scientific Publishing Company, Tokyo, 95116, 1983.
  Torres, P., Bulhão, N., da Cunha, R., Vieira, J., and Rodrigues, A.: Dead or alive: the growing importance of shark diving in the Mid-Atlantic region. Journal for Nature Conservation, 36, 20–28, doi: 10.1016/j.jnc.2017.01.005, 2017.
- Toyos, G. P., Cole, P. D., Felpeto, A., and Martí, J.: A GIS-based methodology for hazard mapping of small pyroclastic density currents, Nat. Hazards, 41, 99112, doi: 10.1007/s11069-006-9026-9, 2007.
  Vianna, G. M. S., Meekan, M. G., Pannell, D. J., Marsh, S. P., and Meeuwig, J. J.: Socio-economic value and community benefits from shark-diving tourism in Palau: a sustainable use of reef shark populations, Biological Conservation, 145, 267-277, doi: 10.1016/j.biocon.2011.11.022, 2012.
- 630 Vieira, J.C., Antunes, M.C.: Touristic big-game fishing in Saint Michael Island (Azores): evaluating anglers' profiles, perceptions about the destination and business revenues. Tourism Economics, 23, 1362–1368, doi:10.1177/1354816616686414, 2017.

Vieira, J., Câmara, G., Silva, F., and Santos, C.: Airline choice and tourism growth in the Azores, Journal of Air Transport Management, 77, 16, doi: 10.1016/j.jairtraman.2019.02.009, 2019.

635 Viveiros, F., Gaspar, J. L., Ferreira, T., Silva, C., Marcos, M., Hipólito, A.: Mapping of soil CO<sub>2</sub> diffuse degassing at São Miguel Island and its public health implications, in: Volcanic Geology of São Miguel Island (Azores Archipelago), Memoirs, 44, edited by Gaspar, J. L., Guest, J. E., Duncan, A. M., Barriga, F. J. A. S., and Chester, D. K., Geological Society, London, 185195, doi: 10.1144/M44.14, 2015.





Wadge, G., Voight, B., Sparks, R. S. J., Cole, P. D., Loughlin, S. C., and Robertson, R. E. A.: An overview of the eruption of

640 Soufrière Hills Volcano, Montserrat form 2000 to 2010, in: The Eruption of Soufrière Hills Volcano, Montserrat from 2000 to 2010, Wadge, G., Robertson, R. E. A., and Voight, B., Memoirs, 39, Geological Society, London, 140, doi: 10.1144/M39.1, 2014.

Walker, G. P. L. and Croasdale, R.: Two plinian-type eruptions in the Azores. Journal of Geological Society of London, 127, 1755, 1971.

- Wallenstein, N.: Estudo da história recente e do comportamento eruptivo do vulcão do Fogo (S. Miguel, Açores). Avaliação preliminar do "hazard", Ph.D. thesis, Universidade dos Açores, 266 pp., 1999.
  Wallenstein, N., Duncan, A., Guest, J. E., and Almeida, M. H.: Eruptive history of Fogo Volcano, São Miguel, Azores, in: Volcanic Geology of São Miguel Island (Azores Archipelago), Memoirs, 44, edited by Gaspar, J. L., Guest, J. E., Duncan, A. M., Barriga, F. J. A. S., and Chester, D. K., Geological Society, London, 105123, doi: 10.1144/M44.8, 2015.
- Wallenstein, N., Duncan, A., Chester, D. K., and Marques, R.: Fogo Volcano (São Miguel, Azores): a hazardous edifice. Géomorphologie: relief, processus, environnement, 3, 259–270, doi: 10.4000/geomorphologie.2853, 2007.
  Wilson, G., Wilson, T. M., Deligne, N. I., and Cole, J. W.: Volcanic hazard impacts to critical infrastructure: a review, J. Volcanol. Geoth. Res., 286, 148182, doi: 10.1016/j.jvolgeores.2014.08.030, 2014.
  Wilson, T. M., Stewart, C., Sword-Daniels, V., Leonhard, G. S., Johnston, D. M., Cole, J. W., Wardman, J., Wilson, G., and
- 655 Barnard, S. T.: Volcanic ash impacts on critical infrastructure, Phys. Chem. Earth, pt. A/B/C, 4, 523, doi: 10.1016/j.pce.2011.06.006, 2012.

Eruptive source para	meters	References		
VEI 4 sub-Plinian	Erupted volume (km <sup>3</sup> )	1	Booth et al. (1978)	
	Column height (m)	18 500	Carey and Sparks (1986)	
VEI 5 Plinian	Erupted volume (km <sup>3</sup> )	3.2	Booth et al. (1978)	
	Column height (m)	27 000	Bursik et al. (1992)	
Grain size	Mean diameter (Md $\Phi$ )	0	Cole et al. (1995)	
	Standard deviation ( $\sigma\Phi$ )	2	Cole et al. (1995)	
	Minimum $(\Phi)$	4	Walker and Croasdale (1971)	
	Maximum ( $\Phi$ )	-4	Walker and Croasdale (1971)	
Clast density (kg/m <sup>3</sup> )	Large ( $\Phi < 1$ )	800	Wilson and Huang (1979)	
	Medium $(1 \le \Phi \le 3)$	1200	Wilson and Huang (1979)	





	Small ( $\Phi > 3$ )		2300	Wilson and Huang (1979)		
Wind conditions						
VEI 4 sub-Plinian						
Summer period	Altitude (m)	1500	6125	10 750	15 375	20 000
	Direction (°)	270	315	315	270	90
	Intensity (m/s)	8	17	20	8	6
Winter period	Altitude (m)	1500	6125	10 750	15 375	20 000
	Direction (°)	270	270	270	270	270
	Intensity (m/s)	17	20	20	17	11
VEI 5 Plinian						
Summer period	Altitude (m)	1500	8125	14 750	21 375	28 000
	Direction (°)	270	315	270	90	90
	Intensity (m/s)	8	17	18	8	11
Winter period	Altitude (m)	1500	8125	14 750	21 375	28 000
	Direction (°)	270	270	270	270	270
	Intensity (m/s)	17	20	20	10	17

# 660 Table 1 - Input parameters used for the simulations of VEI 4 sub-Plinian and VEI 5 Plinian eruptions.

Trachas dh'alanaan	Summer per	iod scenarios	Winter period scenarios			
Tephra thickness	Unaffected	1 mm - 20 cm	20 cm - 1 m	>1 m	Unaffected	1 mm - 5 cm
VEI 4 sub-Plinian	7.1 %	62.1 %	21 %	9.3 %	95 %	5 %
VEI 5 Plinian	7.9 %	61.4 %	12 %	19 %	85 %	15 %

# Table 2 - Percentage of buildings affected by tephra fallout from VEI 4 sub-Plinian and VEI 5 Plinian eruption scenarios in the summer and winter period.

665

Economia coonoria	Number	of	accommodation	Percentage	of	accommodation
Economic scenario	buildings			buildings		





1 (tephra fallout VEI 4 sub- Plinian)	17	36 %
2 (PDCs VEI 4 sub-Plinian)	44	96 %
3 (PDCs VEI 5 Plinian)	45	98 %

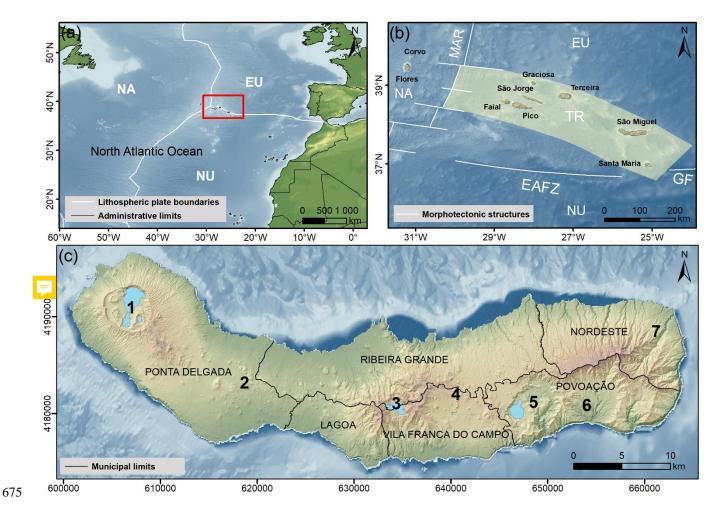
# Table 3 - Number and percentage of buildings in each economic scenario.

Economic scenario	Occupancy rate (ro)	Discount rate (r)	LPV (euros)
	50 %	2 %	13 765 834
1 (tankers fallout VEL 4 sub Divise)	30 %	4 %	10 762 532
1 (tephra fallout VEI 4 sub-Plinian)	(F. 0)	2 %	17 895 584
	65 %	4 %	13 991 291
	50.0/	2 %	109 100 634
2 (PDCs VEI 4 sub-Plinian)	50 %	4 %	85 298 067
	65 %	2 %	141 830 825
		4 %	110 887 487
	50.0/	2 %	111 491 752
	50 %	4 %	87 167 513
3 (PDCs VEI 5 Plinian)	6 <b>5</b> 0/	2 %	144 939 278
	65 %	4 %	113 317 767

670 Table 4 - Loss Present Value after 30 years for economic scenario 1 (tephra fallout VEI 4 sub-Plinian eruption), economic scenario 2 (PDCs VEI 4 sub-Plinian eruption) and economic scenario 3 (PDCs VEI 5 Plinian eruption), considering a discount rate of 2 % and an occupancy rate of 50 %; discount rate of 4% and occupancy rate of 50 %; discount rate of 2 % and occupancy rate of 65 %.





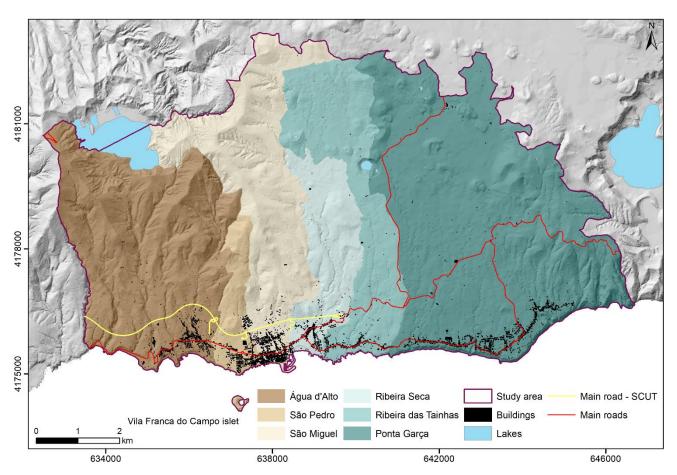


680

Figure 1 - (a) Location of the Azores archipelago in the North Atlantic Ocean and relation to the triple junction between the North American (NA), Eurasian (EU) and Nubian (NU) lithospheric plates (world bathymetry and topography from The GEBCO\_08 Grid; plate and country boundaries from ESRI). (b) Geodynamic setting of the Azores archipelago and main morphotectonic structures of the region. MAR - Mid-Atlantic Ridge; TR - Terceira Rift; EAFZ - East Azores Fracture Zone; GF - Gloria Fault (Azores bathymetry from EMODnet Bathymetry Consortium (2018); morphotectonic structures modified from Hipólito et al., 2010). (c) Digital elevation model of São Miguel Island showing the volcanic systems and administrative limits of the six municipalities. 1 - Sete Cidades volcano; 2 - Picos fissure volcanic system; 3 - Fogo volcano; 4 - Congro fissure volcanic system; 5 - Furnas volcano; 6 - Povoação volcano; 7 - Nordeste volcanic system (after Gaspar et al., 2015).







685 Figure 2 - Location of Vila Franca do Campo municipality (the study area) on São Miguel Island showing the six parishes, buildings, infrastructure and main roads.





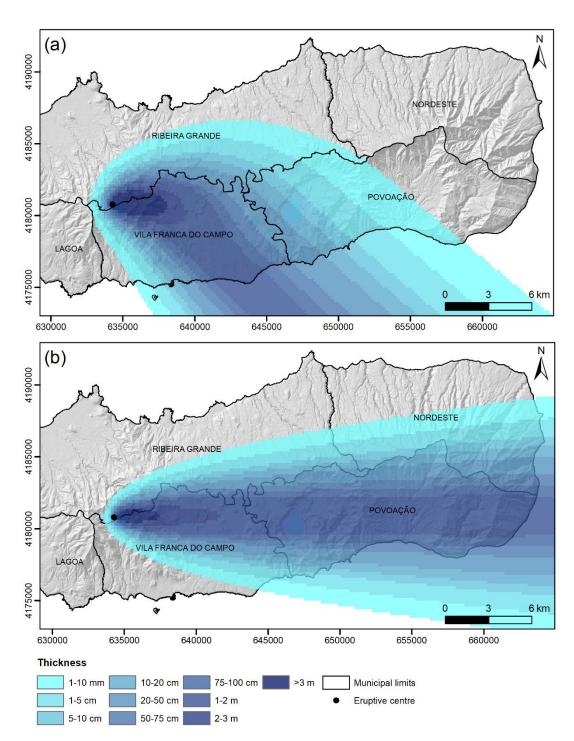


Figure 3 - (a) Tephra fallout deposition from a VEI 4 sub-Plinian eruption of Fogo volcano, considering the dominant winds for the summer period. (b) Tephra fallout deposition from a VEI 4 sub-Plinian eruption of Fogo volcano, considering the dominant winds for the winter period.





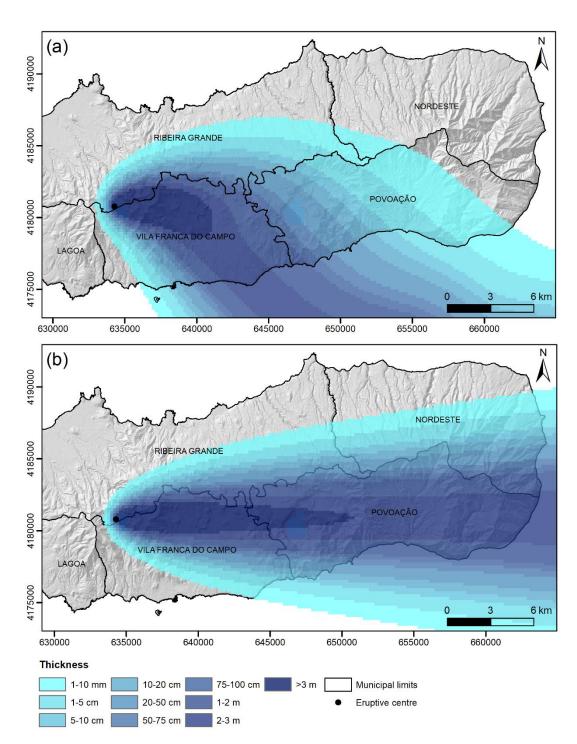
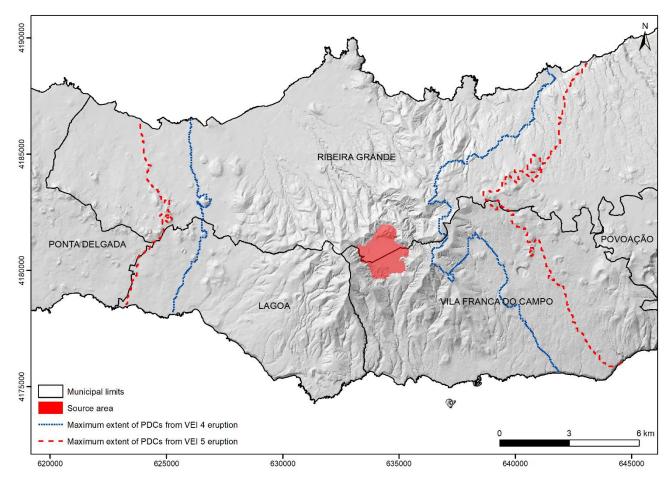


Figure 4 - (a) Tephra fallout deposition from a VEI 5 Plinian eruption of Fogo volcano, considering the dominant winds for the summer period. (b) Tephra fallout deposition from a VEI 5 Plinian eruption of Fogo volcano, considering the dominant winds for the winter period.







695

Figure 5 – Maximum potential extent of PDCs from a VEI 4 sub-Plinian eruption (blue dotted line; collapse height of 300 m) and from a VEI 5 Plinian eruption (red dashed line; collapse height of 500 m) of Fogo volcano.





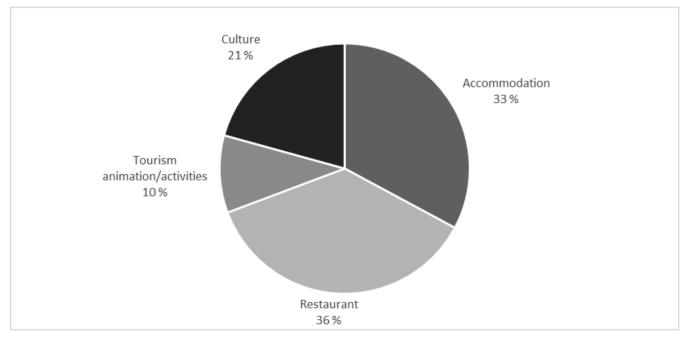
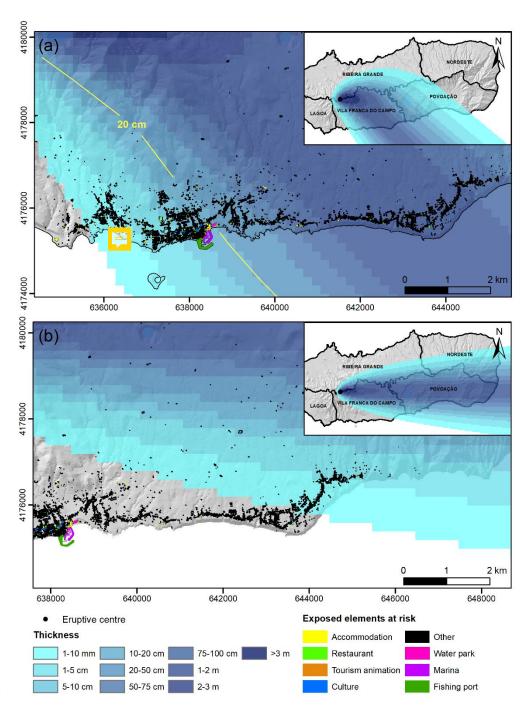


Figure 6 - Percentage of buildings of each typology located in Vila Franca do Campo municipality.





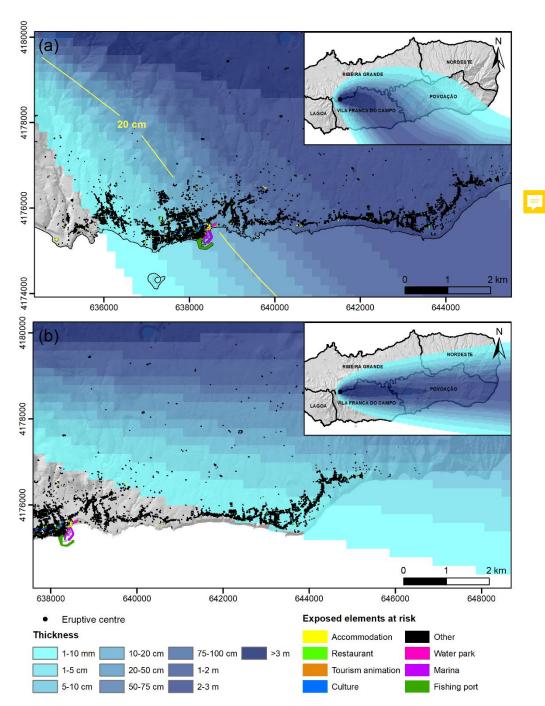


700

Figure 7 - (a) Spatial distribution of the elements at risk in Vila Franca do Campo municipality affected by tephra fallout resulting from a VEI 4 sub-Plinian eruption at Fogo volcano, considering the dominant winds for the summer period. (b) Spatial distribution of the elements at risk in Vila Franca do Campo municipality affected by tephra fallout resulting from a VEI 4 sub-Plinian eruption of Fogo volcano, considering the dominant winds for the winter period.







705

Figure 8 - (a) Spatial distribution of the elements at risk in Vila Franca do Campo municipality affected by tephra fallout resulting from a VEI 5 Plinian eruption of Fogo volcano, considering the dominant winds for the summer period. (b) Spatial distribution of the elements at risk in Vila Franca do Campo municipality affected by tephra fallout resulting from a VEI 5 Plinian eruption of Fogo volcano, considering the dominant winds for the winter period.





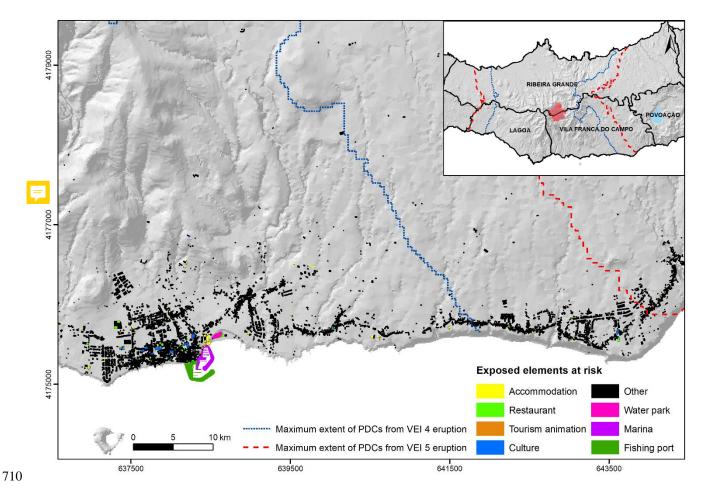


Figure 9 - Spatial distribution of the elements at risk in Vila Franca do Campo municipality affected by PDCs resulting from VEI 4 sub-Plinian (blue dotted line) and VEI 5 Plinian (red dashed line) eruptions of Fogo volcano.





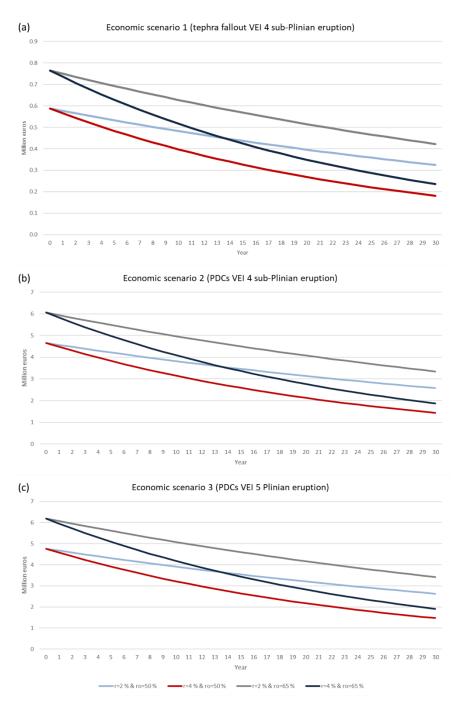


Figure 10 – (a) Loss Present Value after 30 years for economic scenario 1 (tephra fallout VEI 4 sub-Plinian eruption) considering a discount rate of 2 % and an occupancy rate of 50% (light blue line); discount rate of 4 % and occupancy rate of 50 % (red line); discount rate of 2 % and occupancy rate of 65 % (gray line); discount rate of 4 % and occupancy rate of 65 % (dark blue line). (b) Loss Present Value after 30 years for economic scenario 2 (PDCs VEI 4 sub-Plinian eruption) considering a discount rate of 2 % and an occupancy rate of 50 % (light blue line); discount rate of 4 % and occupancy rate of 50 % (red line); discount rate of 2 % and occupancy rate of 50 % (gray line); discount rate of 4 % and occupancy rate of 50 % (red line); discount rate of 2 % and occupancy rate of 65 % (gray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 50 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate of 65 % (logray line); discount rate of 4 % and occupancy rate





720 after 30 years for economic scenario 3 (PDCs VEI 5 Plinian eruption) considering a discount rate 2 % and an occupancy rate of 50% (light blue line); discount rate of 4 % and occupancy rate of 50 % (red line); discount rate of 2 % and occupancy rate of 65 % (gray line); discount rate of 4 % and occupancy rate of 65 % (dark blue line).