Editor comments

Dear colleagues,

I am pleased to announce, that both reviewers agreed on the suitability of your manuscript for NHESS following minor revisions.

I kindly ask you to follow strictly the reviewers comments as you already pointed out in your respective responses and to resubmit the revised manuscript as soon as possible.

I am looking forward to receiving your revised contribution.

R: Dear Editor.

thank you very much for your support in reviewing the manuscript. In the following we attach a point-by-point response to the reviews, including all relevant changes made in the manuscript, and a marked-up manuscript version.

Reviewer #1

Nice study. Nicely and well written. Everything's fine.

R: we thank Reviewer #1 for his/her kind comments.

Very small comments:

- The first two lines in both abstract and introduction theoretically could be left out. These sentences are found in almost every rockfall publication.

R: We thank for this comment. Nevertheless, we would prefer to leave these generic statements to introduce the topic and the potential hazard of rockfalls in the wider framework of rainfall-induced mass movements.

- Line 88: you could add a reference to Fig. 1 here as well.

R: Done.

- Section 3.2: how were the occurance times of the rockfalls been detected? Was it the date of the road inspection? How often took the inspection place? etc.

R: We thank the reviewer for the question that allowed us to clarify the item. We added now the following text in Section 3.2: "After a rockfall occurrence, a quick action is required by local authorities to remove boulders from road, and to repair the damage. The average response time for these emergencies is down to one day, and therefore the date of the rockfall occurrence is the same of the road inspection."

Apart from scheduled inspections, a survey is carried out whenever a rockfall occurrence is reported.

- Section 5: is there a reasonable approach to see rainfall intensity changing due to climate change and therefore to use this study to predict rockfall activity?

R: This is an interesting topic, indeed. Following the indications of the World Meteorological Organization, which reads: "a 30-year period is long enough to filter out any inter annual variation or anomalies, but also short enough to be able to show longer climatic trends", we believe that the number of years available to us to assess a reliable climatic trend is inadequate.

- References: You could add Contino et al, NHESS 17(12), 2017.

R: Done.

Reviewer #2

This manuscript applies the ED rainfall threshold on rockfall and discusses the link between the change of seasonal rainfall amount and the occurrence of rockfall through great amounts of rockfall cases and rainfall data. The contents, figures, tables are all well-organized. The results also provide a new insight for the early warning of rockfall. Here are some suggestions for the authors that may take into account.

R: We thank the Reviewer #2 for his/her kind comments.

1. Since there are lots of rockfall records in GC and TEN, is it possible to discuss the influence of geology and geomorphology on rainfall threshold? Besides, the relationship between the volume of rockfall and the cumulated event rainfall may also be discussed.

R: We thank the Reviewer #2 for the interesting suggestion, but from the geological point of view, all Canary Islands are located inside the Africa Plate, in the transitional zone between the Continental Crust and the Oceanic Crust, and are characterized by the volcanic origin that is still active. As a consequence, there aren't significant differences in the geology. Moreover, rockfalls in our catalogue occurred on the road network of GC and TEN, and therefore the dependence of the thresholds on the geology and geomorphology is likely not relevant here.

Regarding the second question, besides a qualitative classification in small, medium, and large boulders provided by the Civil Protection Authorities, the volume associated to each rockfall is unfortunately unavailable. We agree with the Reviewer that this relationship would have been very useful in defining the magnitude of the event as a function of the rainfall.

2. [Section 2.2, Line 128~130] Comparing to the last paragraph in section 2.1, the information of average annual rainfall in TEN is missed.

R: We now added the required information for TEN, and we corrected the upper average annual rainfall.

3. [Figure 4 & Line 227] There are 29 cases with hourly rainfall data in GC, but they are not analyzed in Figure 4 c and d. Too less to be analyzed?

R: The Reviewer #2 is right. The low number of cases precludes the calculation of rainfall thresholds. We now added this statement in the text.

4. It is suggested to add some representative pictures of these rockfalls in order to make readers understand the environment of study area more.

R: We thank Reviewer #2 for this suggestion. We now added two pictures of rockfalls on GC-200 road and on a road in Tenerife (shown below).





Minor issues:

[Line 219~220] "light green dots in Fig. 2" These are no green dots in Figure 2. Is it Figure 4?

R: Light green dots are in Fig.1 We corrected the citation.

[Figure 2 and Figure 3] The Y-axis of Fig. 2 c, d and Fig. 3 d, e, f should be "Rockfall" instead of "Rock fall."

R: We thank Reviewer, and we noticed that the same error was also in Fig. 1. We corrected it.

- 1 RAINFALL AND ROCKFALLS IN THE CANARY ISLANDS: ASSESSING A SEASONAL
- 2 LINK

12

14

- 3 Massimo Melillo^a, Stefano Luigi Gariano^a, Silvia Peruccacci^a, Roberto Sarro^b, Rosa Marìa Mateos^c,
- 4 Maria Teresa Brunetti^a
- 6 a CNR IRPI, via Madonna Alta 126, 06128, Perugia, Italia
- 7 massimo.melillo@irpi.cnr.it, stefano.luigi.gariano@irpi.cnr.it, silvia.peruccacci@irpi.cnr.it
- 8 ^b IGME, c/ Alenza, 1, 28003, Madrid, España
- 9 r.sarro@igme.es
- 10 ° IGME, Urb. Alcázar del Genil, 4. Edificio Zulema, bajos, 18006, Granada, España
- 11 rm.mateos@igme.es
- 13 Correspondence to: Maria Teresa Brunetti maria.teresa.brunetti@irpi.cnr.it

Abstract

15

16

17

18

19

20

21

22

23

24

25

26

27

28

29

30

31

32

33

34

Rockfalls are frequent and harmful phenomena occurring in mountain ranges, coastal cliffs and slope cuts. Albeit several natural processes concur in their formation and triggering, rainfall is one of the most common causes. The prediction of rock failures is of social significance for civil protection purposes and can rely on the statistical analysis of past rainfall conditions that caused the failures. The paper describes the analysis of information on rainfall-induced rockfalls in Gran Canaria and Tenerife, Canary Islands (Spain). An analysis of the monthly rainfall versus the monthly distribution of rockfalls reveals that they are correlated for most of the year, except in summer, when other triggers act to induce collapses. National and regional catalogues with hourly and daily rainfall measurements are used to reconstruct the cumulated amount (E) and the duration (D) of the rainfall responsible for the rock failures. Adopting a consolidated statistical approach, new ED rainfall thresholds for possible rockfall occurrence and the associated uncertainties are calculated for the two test sites. As far as is known, this is the first attempt to predict this type of failure using the threshold approach. Using the rainfall information, a map of the mean annual rainfall is obtained for Gran Canaria and Tenerife, and it is used to assess the differences between the thresholds. The results of is study are expected to improve the ability to forecast rockfalls in the Canary Islands, in view of implementing an early warning system to mitigate the rockfall hazard and reduce the associated risk.

Keywords: Rockfall, rainfall threshold, Canary Islands.

1 Introduction

35

36

37

38

39

40

41

42

43

44

45

46

47

48

49

50

51

52

53

54

55

56

57

58

59

60

61

62

63

64

65

66

Rockfalls are instability processes affecting mountainous regions, coastal cliffs and slope cuts. Being very rapid, they are extremely dangerous and life-threatening, especially when they occur in populated areas, along roads and railways. The most frequent triggering factors of rockfalls are rainfall, cycling thermal stress, and seismic activity (Wieckzorek and Jaeger, 1996; Keefer, 2002; Mateos, 2016; Ansari et al., 2015; Collins and Stock, 2016; Contino et al., 2017; Sarro et al., 2018; Saroglou, 2019; González de Vallejo et al., 2020). At regional and global scales, empirical approaches to forecast the occurrence of rockfalls may contribute reducing risk. Generally, for rainfall-induced slope failures the forecast can rely upon the definition of rainfall thresholds, i.e. the rainfall conditions that when reached or exceeded are likely to trigger the failure. Rainfall thresholds are calculated through the statistical analysis of historical rainfall conditions that have resulted in landslides (e.g., Guzzetti et al., 2007, 2008; Cepeda et al. 2010; Sengupta et al., 2010; Ruiz-Villanueva et al. 2011; Berti et al., 2012; Staley et al., 2013; Zêzere et al., 2015; Palenzuela et al., 2016; Rosi et al., 2016; Peruccacci et al., 2017; Segoni et al., 2018; Valenzuela et al., 2018, 2019). The definition of reliable empirical rainfall thresholds relies on the use of objective procedures for (i) the reconstruction of the rainfall events responsible for the failures and (ii) the calculation of the thresholds. For the purpose, Melillo et al. (2018) have proposed an algorithm that reconstructs rainfall events, identifies the rainfall conditions that have resulted in slope failures, and calculates probabilistic cumulated event rainfall-rainfall duration (ED) thresholds at different non-exceeding probabilities and their associated uncertainties (Peruccacci et al., 2012). The obtained thresholds are a set of parallel power-law curves in a log-log (D,E) plane, which are characterized by a slope and an intercept, the last being a function of the non-exceeding probability value (Brunetti et al., 2010). In this work, a relationship between the amount of rainfall and the occurrence of rockfalls is assessed and empirical rainfall thresholds are defined for two test sites in Gran Canaria and Tenerife, Canary Islands (Spain). The possible prediction of rainfall-induced rock failures is of fundamental importance primarily for the safety of the inhabitants and for preserving infrastructures such as roads and buildings. An increasing level of safety against this type of hazard is also important for the local economy, one third of which is based on tourism. As far as is known, this is the first attempt to predict rock failures triggered by rain using the threshold approach. Recently, in Italy it has been observed that the slope of the power-law curve is dependent on the mean annual rainfall (MAR). In particular, the higher is the MAR the steeper is the threshold (Peruccacci et al., 2017). This relationship is explained assuming that where the landscape has been shaped over long time periods by landslides

- triggered by a given minimum amount of rainfall, it is likely necessary at least as much rainfall to
- trigger the next landslides (Chen, 2015). For improving the discussion of the results, it has been
- 69 considered worthwhile producing a map of the MAR for the islands of Gran Canaria and Tenerife
- vi using the available rainfall data sets.
- 71 The manuscript is organized as follows. After a description in Section 2 of the general settings of the
- two test sites, Section 3 describes the rainfall and rockfall datasets, and the methods used to determine
- 73 ED rainfall thresholds and the map of the MAR. Section 4 illustrates in detail the relationship between
- 74 the rainfall regime and the occurrence of rock failures, and presents the rainfall thresholds for the
- possible rockfall occurrence in the two test sites. Finally, in Section 5, the main findings of the work
- are summarised and discussed.

2 Test site description

77

89

- 78 The Canary Islands (Spain) are one of the major volcanic chain in the oceans. The archipelago
- 79 consists of eight islands in the Atlantic Ocean, aligned along a W-SW to E-NE direction: El Hierro,
- 80 La Palma, La Gomera, Tenerife, Gran Canaria, Fuerteventura, La Graciosa and Lanzarote. The
- geological origin of the Canary archipelago (800 km in length) is still under debate, but it has been
- traditionally interpreted as a hotspot track (Fullea et al., 2015).
- 83 The steep topography and the geological complexity of the archipelago influence the activation of an
- 84 intense slope failures activity. Rockfalls are the most frequent landslide type in the Canary Islands,
- causing damage on built-up areas and communication networks.
- 86 Two test sites are selected for assessing the relationship between the rainfall and the occurrence of
- 87 rockfalls. The first site (GC) is located in the north-western part of Gran Canaria island, and the
- second site (TEN) is the entire Tenerife island (Fig. 1).

2.1 Gran Canaria Island (GC-200 road)

- 90 Gran Canaria is the third island in size of the Canarian archipelago. With an area of 1560 km² and a
- 91 maximum altitude of 1956 m a.s.l., the island is approximately circular in shape (Fig. 1). The origin
- of Gran Canaria can be dated about 15 million years ago (Miocene) with the first submarine building
- 93 stages of the Gran Canaria Volcano. From the geological point of view, the island presents the greatest
- 94 variability of igneous rocks of the entire archipelago. Besides the distinctive lavas of the basanite
- 95 basalt to trachyte phonolite series, Gran Canaria presents also other types of magma, such as tholeiitic

- 96 basalts and rhyolites (Troll and Carracedo, 2016). Massive flank failures and erosion give place to chaotic deposits that cover large areas. 97 98 The test site is the GC-200 road located in the north-western extreme of Gran Canaria, and specifically 99 between the localities of Agaete and Aldea. The road constitutes the main transportation corridor 100 between the two localities. With a length of 34 km, the road path is very tortuous following the 101 contour of the coast, a very step coastline with some of the highest cliffs in Europe. The road has 102 heavy traffic estimated on average at 1500 vehicles per day. The geology of the test site area is within 103 the domain of the basaltic shield stage, Middle Miocene in age. Along the road, an alternance of 104 alkaline basaltic deposits and piroclastic flows can be observed. In some parts, gravitational deposits 105 (mainly colluvial) also outcrop covering wide areas. 106 Regarding climatological conditions, Gran Canaria is located in a transitional zone between temperate 107 and tropical conditions. The conical morphology of Gran Canaria retains the humidity of the 108 predominant N-NE trade winds of the subtropical Azores anticyclone on the north side of the island. 109 As a result, the northern flanks are humid and vegetation is vigorous, while the south part of the island 110 is very dry and the conditions are very arid and desert-like. Annual rainfall ranges between 100 and 111 900 mm on average, increasing with altitude. In the test site the climate is very dry, with low average 112 annual rainfall (< 100 mm) and high average annual temperature (~ 20°C). 113 2.2 Tenerife island 114 Tenerife (Fig. 1) is the largest (2057 km²) and the most populated (950,000 inhabitants and 13.2 million visitors in 2019) island of the archipelago. It is home to the third largest volcano in the world 115 116 (Pico del Teide, 3718 m a.s.l.). 117 From a geological point of view, Tenerife was constructed via Miocene-Pliocene shields that now form the vertices of the island. The shields were unified into a single edifice by later volcanism that 118 119 continued in central Tenerife from approximately 12 to 8 million years ago and was followed by a 120 period of dormancy. Rejuvenation at approximately 3.5 Ma is recorded by the central Las Cañadas 121 Volcano, and long residence times of magmas during this period favoured magmatic differentiation 122 processes to produce an episode of felsic and highly explosive felsic volcanism (Troll and Carracedo,
 - The steep orography of the island and the climate variety have resulted in a diversity of landscapes and geographical formations. Very impressive coastal cliffs (till 500 m in height) are present in the

124

125

2016).

- northern corner of Tenerife. This area is also characterized by narrow and deep ravines which
- determine an intense slope activity.
- 128 The climate of Tenerife is subtropical oceanic; the minimum and maximum annual average
- temperatures are about 15°C in winter and 24°C in summer. Similarly to Gran Canaria, annual rainfall
- ranges between 100 and 900 mm. Tenerife offers a large variety of micro-climate zones controlled
- by the altitude and the winds.

136

143

3 Data and methods

- The availability of rainfall measurements and landslide information is fundamental to define reliable
- rainfall thresholds. For the selection of the rain gauges, the data quality and the location of the rain
- gauges are assessed, given that these features are crucial to characterize the spatial-temporal variation
 - of the precipitations. Similarly, the calculation of the MAR relies on the availability of sufficiently
- long rainfall series (at least 30 years). This is difficult to achieve for a dense network of rain gauges,
- where sensors may exhibit different operating time periods. The World Meteorological Organization
- 139 (WMO) guidelines on the calculation of the annual standard normal, specifically the MAR,
- recommend at least 10 years to define at least provisional MAR maps (WMO, 1989). This is the case
- in the test sites, where a lot of rainfall information is limited to short time periods (the average is 15.6
- 142 years), thus hampering the calculation of the MAR with a detailed space resolution.

3.1 Rainfall data

- In the GC test site, hourly rainfall data (purple triangles in Fig. 1) from the Spanish National
- Meteorological Service (AEMET) network (in total 25 stations among which 4 are close to the study
- area) are used for the calculation of rainfall thresholds. Moreover, daily rainfall data (orange triangles
- in Fig.1) are provided from the Consejo Insular de Aguas de Gran Canaria (CIAGC) regional rain
- gauge network (13 stations) and from AEMET (92 stations, among which 7 are close to the study
- area). Some of the sensors of the AEMET network provide both hourly and daily rainfall, in different
- time periods. Details of the rainfall series are reported in Table 1.
- 151 For the TEN test site, rainfall measurements are provided by AEMET with the contribution of
- regional networks. As for the GC test site, the rainfall analysis is performed using both hourly and
- daily data. The two networks in the TEN test site are composed by 34 rain gauges recording hourly
- data (purple triangles in Fig. 1) and 66 rain gauges recording daily data (orange triangles in Fig. 1).

- To calculate the MAR for the two test sites, yearly and monthly rainfall data provided by AEMET
- and by Sistema de Información Agroclimático y de Regadíos (SIAR), respectively, are used (Table
- 157 1). In particular, in order to obtain homogeneous maps, data recorded in the 20-year period from
- January 2000 to December 2019 in both test sites are selected. Following WMO guidelines (WMO,
- 159 1989), only stations with at least 10 years of data are included in the analysis. Overall, 72 (one every
- 160 22 km²) and 67 (one every 31 km²) rain gauges are used to calculate MAR in Gran Canaria and
- Tenerife, respectively. The average number of sensors operating per year in the considered period is
- 162 56 (84%; one every 28 km²) in Grand Canaria and 47 (65%; one every 43 km²) in Tenerife. The used
- rain gauges are homogeneously distributed over the test site areas.
- Using the monthly and annual rainfall data recorded by the 103 rain gauges in the two islands, the
- MAR for the period 2000-2019 was calculated for each station. Moreover, the coefficient of variation
- of the MAR is calculated by dividing its standard deviation by the MAR. This coefficient represents
- the variability of the MAR in the considered time interval. The map of the MAR and of its coefficient
- of variation are calculated using the tension spline tool in ESRI ArcMAP 10.7.1.

3.2 Rockfall data

- 170 The information on the rockfalls was collected by the Road Maintenance Service in the GC test site
- and by the Canarian Civil Protection Authorities in the TEN test site. After a rockfall occurrence, a
- quick action is required by local authorities to remove boulders from road, and to repair the damage.
- 173 The average response time for these emergencies is down to one day, and therefore the date of the
- 174 rockfall occurrence is the same of the road inspection. Figure 2 portrays an example of rockfalls
- occurred in GC and TEN test sites.
- For the GC test site a total of 8174 rockfall events occurred from January 2010 to March 2016 was
- documented. A catalogue was prepared defining accurately the location of each impact along the road
- using orthophotos available for the region and technical reports. The information for each event
- includes kilometre point, number of events, date, and boulder size. In GC only 535 rockfalls
- characterized by medium to large size are included in the analysis for the thresholds, whereas small
- and very small rockfalls ($< 10^{-3} \text{ m}^3$) are discarded. Analogously, a catalogue of 1898 rockfalls that
- impacted along Tenerife roads from January 2010 to November 2017 was prepared. For each event,
- the information includes rockfall localization, geographic accuracy, occurrence day, month, year, and
- time (if available), and temporal accuracy.

The influence of the rainfall on the occurrence of rockfalls is assessed analysing the distribution of monthly rainfall (Figs. 3a,b) and monthly number of rockfalls (Figs. 3c,d) on the two test sites. As expected, an increase of the rainfall in the autumn-winter period, between October and March, is observed in both islands, with a maximum in November.

The monthly distribution of rockfalls in Gran Canaria (Fig. 3c) is coherent with the rainfall values in the period January-April, with a maximum in February (~ 130; Fig. 3a). For the remaining dry (May to September) and wet (October to December) months the number of rock failures decreases and becomes almost flat (below 50). This behaviour suggests the presence of triggering mechanisms other than the rainfall. For the TEN test site, the number of rockfalls per month (Fig. 3b) is similar to the rainfall distribution, confirming anyway the presence of one or more additional triggers as evidenced by the abundance of failures between May and September (Fig. 3d) when the rainfall is irrelevant.

3.3 Empirical rainfall thresholds

Empirical ED thresholds are represented by the following power law curve:

198
$$E = (\alpha \pm \Delta \alpha) \times D^{(\gamma \pm \Delta \gamma)}$$
 (1)

where E is the cumulated event rainfall (in mm), D is the duration of the rainfall event (in hours or in days), α and γ are the intercept and the slope of the curve, respectively, and $\Delta\alpha$ and $\Delta\gamma$ are the uncertainties associated with them. Thresholds at different non-exceedance probabilities are calculated adopting the frequentist approach and the bootstrap nonparametric statistical technique (Brunetti et al., 2010; Peruccacci et al., 2012), using 5000 randomly selected synthetic series of DE pairs. A threshold at 5% non-exceedance probability should leave 5% of the empirical DE pairs below itself. The parameter uncertainties depend mostly on the number and the distribution of the rainfall conditions. The minimum number of DE pairs needed for having stable mean values of the parameters α and γ (i.e. reliable thresholds) depends on the distribution and dispersion of the empirical data points in the DE domain.

3.4 The CTRL-T algorithm for threshold calculation

The quantitative identification of the rainfall responsible for slope failures and the definition of reliable thresholds are fundamental steps towards a well-founded event prediction (Peruccacci et al., 2017; Melillo et al., 2018). The use of standardized procedures for the reconstruction of the rainfall conditions able to trigger past failures and for the definition of thresholds is necessary for enhancing

the objectivity and reproducibility of the curves. The tool named CTRL-T (Calculation of Thresholds for Rainfall-induced Landslides - Tool) proposed by Melillo et al. (2018) is exploited to calculate *ED* thresholds for the two test sites. CTRL-T reconstructs rainfall events starting from continuous rainfall series. For each rockfall, the algorithm: 1) identifies automatically the representative rain gauge; 2) identifies multiple (*D*,*E*) rainfall conditions responsible for the failure; 3) selects among them the maximum probability rainfall conditions (MPRCs). Then, analysing the distribution of the MPRCs it calculates probabilistic rainfall thresholds at different non-exceeding probabilities and their associated uncertainties. In order to avoid using wrong temporal information (i.e., incorrect dates for the occurrence of rockfalls) in the definition of the thresholds, the rainfall conditions having a delay longer than 48 hours between the rainfall ending time and the rockfall occurrence are discarded.

Using CTRL-T, 82 rockfalls occurred between 2012 and 2016 in GC test site and 626 rockfalls occurred between 2010 and 2016 in the TEN test site are selected (light green dots in Fig. 1). The remaining records are discarded due to the: 1) absence of rainfall data in the period including the collapse occurrence time; 2) absence of rain gauges within a buffer of 15 km radius centred on the rockfall; 3) lack of an evident correlation with the rainfall. The definition of rainfall thresholds relies only upon rainfall conditions that triggered the first failure in each event. As a consequence, numerous rockfalls (106, 39% in GC and 271, 30% in TEN) which occurred at the same date and in the same location, and which are associated with the same rainfall event are discarded. In GC among the remaining rockfalls 53 are analysed with daily and 29 with hourly rainfall data, respectively. The low number of rock failures associated to hourly-based rain gauges is to be ascribed to the low density of the sensors in the area, and precludes the definition of rainfall thresholds. In TEN 245 rockfalls are reconstructed with hourly data and 381 with daily rainfall data. Note that for 83 failures it was possible reconstructing the rainfall conditions using sensors from both the two rain gauge networks. As a consequence, the reconstructed (*D*,*E*) rainfall conditions have different temporal resolutions and are used to define both hourly-based and daily-based rainfall thresholds.

4 Results

A correlation between the rainfall and the observed failures is confirmed by the comparison between the monthly rainfall and the corresponding number of rockfalls both in GC and in TEN (Fig. 4). Figures 4a,b,c show the boxplots of cumulated monthly rainfall based on the data recorded in rain gauges used to reconstruct the rainfall responsible for rockfalls for GC and TEN test sites. Inspection of these figures reveals that the rainfall pattern in the two test sites is typically Mediterranean, with a maximum in winter (but also in October and November) and a minimum in summer, with practically

246 no rain in the warmest months. Analysing data from seven daily-based rain gauges in GC (GC-d), it 247 turns out that the rainiest months are February and November with an average rainfall of 52.2 mm 248 and 55.7 mm, a highest rainfall of 98.6 mm and 133.9 mm, and a median rainfall of 42.3 mm and 249 39.8 mm, respectively (Fig. 4a). A similar trend is found for Tenerife using both daily and hourly 250 data. Data from 40 daily-based rain gauges in TEN (TEN-d) are analysed finding an average rainfall 251 of 64.6 mm and 86.4 mm, a highest rainfall of 183.5 mm and 183.6 mm, and a median rainfall of 56.1 252 mm and 93.2 mm for February and November, respectively (Fig. 4b). Data from 21 hourly-based rain 253 gauges in TEN (TEN-h) are analysed finding an average rainfall of 88.8 mm and 82.0 mm, a highest 254 rainfall of 169.8 mm and 190.8 mm, and a median rainfall of 97.5 mm and 66.4 mm for February and 255 November, respectively (Fig. 4c). 256 Figures 4d,e,f portray the monthly number of rockfalls associated with rainfall events for GC-d, TEN-257 d and TEN-h. The GC catalogue lists 53 collapses occurred in the period from November 2012 to 258 October 2016, with the majority of the failures in 2015 (22). The month with the largest number of 259 rockfalls (14) is February, followed by January (8) and November (7). The least number of failures 260 is reported in September (1) and no rainfall-induced rockfalls are reported in May and July (Fig. 4d). 261 The 245 rock failures in the TEN-d catalogue cover the period from September 2010 to February 262 2016, with the majority of records in 2014 (66). The month with the largest number of rockfalls (80) 263 is November, followed by October (37) and December (36). The least number of failures is reported in May (1) and no rainfall-induced rockfalls are reported in June and July (Fig. 4e). The TEN-h 264 265 catalogue lists 381 rockfalls occurred in the period from September 2010 to November 2016, with 266 the majority of the failures in 2014 (90). The month with the largest number of rockfalls (115) is 267 November, followed by December (72) and October (64). The least number of failures is reported in 268 May (1) and no collapses are reported in July (Fig. 4f). 269 The rainfall that triggered the rockfalls is classified according to the method proposed by Alpert et al. 270 (2002), based on six daily rainfall (E_d) categories from "light" to "torrential" over the Mediterranean (Table 2). Using the procedure adopted by Melillo et al. (2016), each rainfall condition responsible 271 272 for rock failures (MPRC) is attributed to a specific category. In particular, for events lasting less than 273 24 hours, a category based on the total cumulated rainfall of the event is assigned. For events lasting 274 more than 24 hours, the maximum value of the cumulated rainfall in 24 hours in a moving window is 275 used. In GC, over 40% of the MPRCs responsible for the collapses are classified as moderate-high 276 (MH); in TEN, approximately 30% as high (H) and high-torrential (HT). No MPRCs are found in the 277 lowest Alpert's category (light, Table 2). Figures 4g,h,i show the cumulated percentage of rainfall 278 events per month grouped according to Alpert's classification. In GC-d, in February (Fig. 4g) 6 279 rockfalls (43%) are triggered by a rainfall classified as H, 3 (21%) as torrential (H), 3 (22%) as MH, 280 and 1 as light-moderate (LM) and HT each (14%). In TEN-d, in November, 29 rockfalls (36%) are 281 triggered by a rainfall classified as HT, 26 (33%) as MH, 22 (28%) as H, 2 (2%) as LM, and 1 (1%) 282 as T (Fig. 4h). In TEN-h, in November, 5 (4%), 26 (23%), 26 (23%), 31 (27%) and 27 (23%) rockfall 283 are triggered by a rainfall classified as LM, MH, H, HT and T, respectively (Fig. 4i). 284 Using the catalogues of rainfall events with rockfalls described above and the CTRL-T tool, ED 285 thresholds, and their associated uncertainties are calculated for GC and TEN test sites. Table 3 lists 286 the number of MPRC used to define the thresholds, the equations of the power law curves, and the 287 range of validity for the thresholds, expressed in hours or days. Note that D must be expressed in days 288 in the equations for the thresholds calculated with daily data, and in hours in the equations for the 289 thresholds calculated with hourly data (Gariano et al., 2020). 290 Figure 5a shows, in logarithmic coordinates, the distribution of the (D,E) rainfall conditions, 291 reconstructed with daily data, that have caused rockfalls in GC (53 blue dots) and in TEN (245 green 292 dots). In particular, the 53 daily rainfall conditions responsible for the rockfalls in GC have durations 293 in the range $1 \le D \le 11$ days (with an average value of 2 days) and cumulated rainfall in the range 294 $16.5 \le E \le 219.9$ mm (average value 51.6 mm). All the conditions were recorded in rain gauges 295 located at a maximum distance of 5.7 km from the failures, with a mean value of 2.8 km. The 245 296 daily-based rainfall conditions associated with the collapses in TEN have durations ranging from one 297 to 15 days, with a mean value of two days. The cumulated rainfall ranges from 15.4 to 235.0 mm, 298 with an average of 71.5 mm. The average distance between the rockfalls and their representative rain 299 gauges is 2.2 km, with a maximum distance of 5 km. Figure 5a portrays also the 5% ED thresholds 300 for GC (T_{5,GC-d}, blue curve) and TEN (T_{5,TEN-d}, green curve). The shaded areas around the threshold 301 lines show the uncertainty regions associated to the thresholds (Table 3). Figure 5b portrays the same 302 $T_{5,GC-d}$ and $T_{5,TEN-d}$, in linear coordinates, in the range $1 \le D \le 7$ days. 303 Figure 5c shows, in logarithmic coordinates, the distribution of the (D,E) rainfall conditions, 304 reconstructed with hourly data, that have triggered rock failures in TEN (381 purple dots). The hourly 305 rainfall conditions associated to rockfalls have durations ranging from 2 to 712 hours and mean value 306 of 111 hours. The cumulated rainfall ranges from 10.6 to 433.9 mm, with an average of 105.6 mm. 307 The average distance between the rockfalls and the representative rain gauges is 6.7 km, with a 308 maximum distance of 14.9 km. In the log-log plot the purple curve is the 5% threshold for TEN

(T_{5,TEN-h}) obtained with hourly data. Figure 5d portrays the same T_{5,TEN-h}, in linear coordinates, in the

- range $1 \le D \le 120$ hours. The uncertainty associated with the threshold (purple shaded area in Figs.
- 311 5c,d) is also shown.

- The difference between the $T_{5,GC-d}$ and $T_{5,TEN-d}$ thresholds can be ascribed to the different MAR in the
- 313 two test sites. Figure 6 portrays the maps of the MAR and of its coefficient of variation, which is the
- 314 percentual variability (standard deviation) of the MAR in the considered time interval. The
- 315 geographical distribution of the MAR values exhibits the highest values in the northern parts of both
- islands, where it overcomes 800 mm (Fig. 6a). On the contrary, the highest values of the coefficient
- of variation (i.e. an index of the MAR variability) are localized in the southern part of the islands,
- 318 where the rain gauge density is lower (Fig. 6b).

5 Discussion and conclusions

- In Canary Island rainfall is the most important triggering factor for rockfalls (Fig. 3). Nevertheless,
- there are other factors that predispose directly or indirectly the trigger of the failure (Temiño et al.,
- 322 2013a). Factors that greatly accentuate this hazard in the two test sites are wind, geomorphological
- 323 characteristics (e.g., slope, aspect), type of soil and seismic activity. Regarding the wind many
- 324 collapses are caused by strong gusts of wind that affect the northern side of Tenerife Island and the
- road GC-200 from Agaete to Aldea in Gran Canaria. (Temiño et al., 2013b). Regarding the
- 326 geomorphology, the existence of many sections of road running through the old basaltic massifs with
- 327 significant sub-vertical jointing, makes the area very susceptible to rock failures. In addition, the
- 328 action of the trade winds on the higher altitude areas, produces an increase in the relative humidity,
- as large masses of water vapor are retained by steep slopes resulting in an intense weathering (and
- weakening) of the rock masses. Finally, the large flank instability of the two test sites (especially in
- the northwest sector of the Gran Canaria island) could be related to structural control and to seismic
- activity connected to the dynamic geologic condition that characterizes them. (Masson et al., 2002;
- 333 Temiño et al. 2013b; Urgeles et al., 2001).
- 334 By selecting the subset of rockfalls triggered by rainfall it can be observed that their monthly
- frequency is linked to the monthly distribution of the rainfall measured in nearby rain gauges (Figs.
- 4a-f). For GC-d (Figs. 4a,d) the correlation is apparently weaker in fall than in winter, but this could
- be ascribed to a statistical fluctuation and should be confirmed by increasing the number of events.
- Conversely, for TEN-d (Figs. 4b,e) the monthly number of rock failures well reflects the monthly
- rainfall amount, suggesting that rainfall is the only triggering cause. Hourly rainfall data in TEN-h

340 (Figs. 4c,f) confirm partially this outcome, since even with a median lower amount of rainfall, a 341 higher number of rock failures is expected to occur from October to December than in February. 342 The number of rockfalls for which it has been possible to reconstruct the rainfall conditions (MPRCs) 343 using daily and hourly data in the TEN test site (Figs. 4e,f) is different. This is mostly due to the worst 344 temporal resolution of the TEN-d dataset. 345 In the two test sites, the majority of the rainfall responsible for rockfalls belongs to the Alpert's MH 346 category (Figs. 4g,i). In TEN-h, 31 events belong to the most severe category T, whereas in TEN-d 347 only one event is found in the T category. This result could be ascribed to the time step of the moving 348 window used to assign the Alpert's category. For a rainfall event lasting more than one day, the 349 Alpert's category varies depending on the data temporal resolution, since the time step is one hour or 350 one day for the hourly and daily data, respectively. In TEN-d, the total amount of rainfall responsible 351 for the failure is shared in two or more consecutive days, causing a lowering of the Alpert's category, 352 as confirmed by the paucity of T events in TEN-d. 353 Figure 5 shows that T_{5,TEN-d} is higher and steeper than T_{5,GC-d}. This means that, at increasing values 354 of D, a smaller amount of rainfall (E) is necessary to trigger the collapses in GC than in TEN. 355 Comparing Figures 1 and 5, the recorded rockfalls in the TEN test site are localized in areas including 356 several classes of MAR (ranging from 100 to 800 mm), while in GC test site they fall in the area 357 characterized by the lowest class of MAR (≤ 100 mm). The different ranges of MAR values in the 358 two test sites are able to explain the observed differences in the two daily ED thresholds (Fig. 5a). 359 This finding confirms that where the MAR is higher, the minimum rainfall conditions able to trigger 360 a failure, specifically a rockfall, are also higher. 361 Moreover, the threshold defined for the TEN test site has an uncertainty smaller than the threshold 362 for GC test site. Peruccacci et al. (2012) observed that the parameter uncertainty reduces as the 363 number of MPRC used to calculate the threshold increases. In particular, as derived from Table 3, the relative uncertainty of the intercept, $\Delta\alpha/\alpha$ is 9.8% for T_{5,GC-d} and 4.9% for T_{5,TEN-d}. Regarding the 364 365 slope of the curves, $\Delta \gamma / \gamma$ is 16.1% for $T_{5,GC-d}$ and 6.7% for $T_{5,TEN-d}$. Given the lower uncertainty range 366 and relative uncertainties of both parameters, T_{5,TEN-d} has a reliability higher than that of T_{5,GC-d}. The 367 same analysis for the $T_{5,TEN-h}$ threshold gives $\Delta\alpha/\alpha = 9.3\%$ and $\Delta\gamma/\gamma = 4.2$. Thresholds with an hourly 368 temporal resolution and having relative uncertainties of the parameters α and γ lower than 10% could 369 be implemented in an operative system for the prediction of rainfall-induced failures (Peruccacci et 370 al., 2012; 2017). The thresholds for different non-exceedance probabilities obtained for TEN test site using hourly rainfall data are suited for the design of probabilistic schemes for the operative prediction of rainfall-induced rockfalls. An improvement in the number of rain gauges providing hourly

measurements, as well as in the number of recorded rock failures, would be necessary in GC test site

in order to reduce the uncertainty of the threshold.

implemented in a prototype early warning system.

375 Currently, neither prototype nor operative early warning systems for rainfall-induced failures are 376 present in the Canary Islands (Guzzetti et al. 2020). The findings of this work can contribute to the 377 understanding of the rainfall conditions that can trigger rainfall-induced rockfalls in Tenerife and in 378 the western part of Gran Canaria, and their relationship with mean annual rainfall regime. These 379 findings have scientific and social implications given that in both test sites also spring and autumn 380 are characterized by a moderately occurrence of rock failures, with relevant impacts on the population, tourism activities, and local economy. As long as a sufficient amount of empirical data 381 382 will be available in both test sites (and also in other islands of the archipelago), the method adopted 383 in this work for the definition of reliable rainfall thresholds can be replicated, and the results can be

6 Code availability

373

374

384

385

391

395

- The code of CTRL-T was written using the R open-source software and can be freely downloaded at:
- 387 http://geomorphology.irpi.cnr.it/tools/rainfall-events-and-landslides-thresholds/ctrl-t-algorithm/ctrl-
- 388 code/ctrl_t_code.r/view. An example of the input files required by the algorithm can be freely
- downloaded at: http://geomorphology.irpi.cnr.it/tools/rainfall-events-and-landslides-thresholds/ctrl-
- 390 t-algorithm/input-demo/INPUT.zip/view.

7 Author contribution

- 392 All authors designed the experiment and carried it out. M.M., S.P. and M.T.B. developed the model
- 393 code, R.S. and R.M.M. provided the data, and M.M. and S.L.G. performed the simulations. All
- authors analyzed the results and prepared the manuscript.

8 Competing interests

396 The authors declare that they have no conflict of interest.

9 Acknowledgements

- 398 Research conducted within the framework of the U-Geohaz project (Geohazard Impact Assessment
- 399 for Urban Areas) funded by the European Commission, Directorate-General Humanitarian Aid and
- 400 Civil Protection (ECHO), under the call UCPM-2017-PP-AG. This work was also funded by the
- 401 Salvador de Madariaga Mobility Program from the Spanish Ministry of Science, Project:
- 402 PRX18/00020 and by the University of Alicante. Authors would also like to thank Dr. Luis E.
- 403 Hernández Gutiérrez, from the Regional Service of Studies and Ecological Impacts of the Canary
- 404 Islands Government.

397

405

10 References

- 406 Alpert, P., Ben-Gai, T., Baharan, A., Benjamini, Y., Yekutieli, D., Colacino, M., Diodato, L., Ramis,
- 407 C., Homar, V., Romero, R., Michaelides, S., and Manes, A.: The paradoxical increase of
- 408 Mediterranean extreme daily rainfall in spite of decrease in total values, Geophy. Res. Lett., 29(11),
- 409 31-1–31-4, https://doi.org/10.1029/2001GL013554, 2012.
- 410 Ansari, M.K., Ahmed, M., Rajesh Singh, T.N., and Ghalayani, I.: Rainfall, a major cause for rockfall
- 411 hazard along the roadways, highways and railways on hilly terrains in India, in: Lollino, G., Manconi,
- 412 A., Clague, J., Shan, W., and Chiarle, M. (Eds.) Engineering Geology for Society and Territory –
- 413 Vol. 1, pp. 457–460, Springer, Cham, https://doi.org/10.1007/978-3-319-09300-0_87, 2015.
- Berti, M., Martina, M.L.V., Franceschini, S., Pignone, S., Simoni, A., and Pizziolo, M.: Probabilistic
- rainfall thresholds for landslide occurrence using a Bayesian approach, J. Geophys. Res, 117, F04006,
- 416 https://doi.org/10.1029/2012JF002367, 2012.
- Brunetti, M.T., Peruccacci, S., Rossi, M., Luciani, S., Valigi, D., and Guzzetti, F.: Rainfall thresholds
- 418 for the possible occurrence of landslides in Italy, Nat. Hazards Earth Syst. Sci., 10, 447–458,
- 419 https://doi.org/10.5194/nhess-10-447-2010, 2010.
- 420 Cepeda, J., Höeg, K., and Nadim, F.: Landslide-triggering rainfall thresholds: a conceptual
- 421 framework, Q. J. Eng. Geol. Hydroge., 43, 69–84, https://doi.org/10.1144/1470-9236/08-066, 2010.
- 422 Chen, C., Saito, H., and Oguchi, T.: Rainfall intensity-duration conditions for mass movements in
- 423 Taiwan, Prog. Earth Planet. Sci., 2, 14, https://doi.org/10.1186/s40645-015-0049-2, 2015.

- 424 Collins, B.D., and Stock, G.M.: Rockfall triggering by cyclic thermal stressing of exfoliation
- 425 fractures. Nat. Geosci., 9, 395–400, https://doi.org/10.1038/ngeo2686, 2016.
- 426 Contino, A., Bova, P., Esposito, G., Giuffré, I., and Monteleone, S.: Historical analysis of rainfall-
- 427 triggered rockfalls: the case study of the disaster of the ancient hydrothermal Sclafani Spa (Madonie
- 428 Mts, northern-central Sicily, Italy) in 1851, Nat. Hazards Earth Syst. Sci., 17, 2229–2243,
- 429 https://doi.org/10.5194/nhess-17-2229-2017, 2017.
- 430 Fullea, J., Camacho, A.G., Negredo, A.M., and Fernández, J.: The Canary Islands hot spot: new
- insights from 3D coupled geophysical-petrological modelling of the lithosphere and uppermost
- 432 mantle, Earth Planet. Sci. Lett., 409, 71–88, https://doi.org/10.1016/j.epsl.2014.10.038, 2015.
- 433 Gariano, S.L., Melillo, M., Peruccacci, S., and Brunetti, M.T.: How much does the rainfall temporal
- 434 resolution affect rainfall thresholds for landslide triggering?, Nat. Hazards, 100, 655-670,
- 435 https://doi.org/10.1007/s11069-019-03830-x, 2020.
- 436 González de Vallejo, L.I.; Hernández-Gutiérrez, L.E.; Miranda, A.; Ferrer, M. Rockfall Hazard
- 437 Assessment in Volcanic Regions Based on ISVS and IRVS Geomechanical Indices. Geosciences, 10,
- 438 220–239, https://doi.org/10.3390/geosciences10060220, 2020.
- 439 Guzzetti, F., Peruccacci, S., Rossi, M., and Stark, C.P.: Rainfall thresholds for the initiation of
- landslides in central and southern Europe, Meteorol. Atmos. Phys., 98(3), 239–267,
- 441 https://doi.org/10.1007/s00703-007-0262-7, 2007.
- 442 Guzzetti, F., Peruccacci, S., Rossi, M., and Stark, C.P.: The rainfall intensity-duration control of
- 443 shallow landslides and debris flows: an update, Landslides 5(1), 3–17,
- 444 https://doi.org/10.1007/s10346-007-0112-1, 2008.
- Guzzetti, F., Gariano, S.L., Peruccacci, S., Brunetti, M.T., Marchesini, I., Rossi, M., and Melillo, M.:
- 446 Geographical landslide early warning systems, Earth-Sci. Rev., 200, 102973,
- 447 https://doi.org/10.1016/j.earscirev.2019.102973, 2020.
- 448 Keefer, D.K.: Investigating landslides caused by earthquakes A historical review, Surv. Geophys.,
- 449 23, 473–510, https://doi.org/10.1023/A:1021274710840, 2002.

- 450 Masson, D.G., Watts, A.B., Gee, M.J.R., Urgeles, R., Mitchell, N.C., Le Bas, T.P., and Canals, M.:
- 451 Slope failures on the flanks of the western Canary Islands, Earth-Sci. Rev., 57, 1–35,
- 452 https://doi.org/10.1016/S0012-8252(01)00069-1, 2002.
- 453 Mateos, R.M., García-Moreno, I., Reichenbach, P., Herrera, G., Sarro, R., Rius, J., and Aguiló, R.:
- 454 Calibration and validation of rockfall modelling at regional scale: application along a roadway in
- 455 Mallorca (Spain) and organization of its management, Landslides, 13, 751-763,
- 456 https://doi.org/10.1007/s10346-015-0602-5, 2016.
- Melillo, M., Brunetti, M.T., Peruccacci, S., Gariano, S.L., and Guzzetti, F.: Rainfall thresholds for
- 458 the possible landslide occurrence in Sicily (southern Italy) based on the automatic reconstruction of
- 459 rainfall events, Landslides, 13(1), 165–172, https://doi.org/10.1007/s10346-015-0630-1, 2016.
- 460 Melillo, M., Brunetti, M.T., Peruccacci, S., Gariano, S.L., Roccati, A., and Guzzetti, F.: A tool for
- the automatic calculation of rainfall thresholds for landslide occurrence, Environ. Model. Softw., 105,
- 462 230–243, https://doi.org/10.1016/j.envsoft.2018.03.024, 2018.
- Palenzuela, J.A., Jiménez-Perálvarez, J.D., and Chacón, J.: Assessing critical rainfall thresholds for
- landslide triggering by generating additional information from a reduced database: an approach with
- 465 examples from the Betic Cordillera (Spain), Nat. Hazards, 84, 185–212,
- 466 https://doi.org/10.1007/s11069-016-2416-8, 2016.
- Peruccacci, S., Brunetti, M.T., Luciani, S., Vennari, C., and Guzzetti, F.: Lithological and seasonal
- 468 control of rainfall thresholds for the possible initiation of landslides in central Italy, Geomorphology,
- 469 139–140, 79–90, https://doi.org/10.1016/j.geomorph.2011.10.005, 2012.
- 470 Peruccacci, S., Brunetti, M.T., Gariano, S.L., Melillo, M., Rossi, M., and Guzzetti, F.: Rainfall
- 471 thresholds for possible landslide occurrence in Italy, Geomorphology, 290, 39–57,
- 472 https://doi.org/10.1016/j.geomorph.2017.03.031, 2017.
- 473 Rosi, A., Peternel, T., Jemec-Auflič, M., Komac, M., Segoni, S., and Casagli, N.: Rainfall thresholds
- 474 for rainfall-induced landslides in Slovenia, Landslides, 13, 1571–1577,
- 475 https://doi.org/10.1007/s10346-016-0733-3, 2016.
- 476 Ruiz-Villanueva, V., Bodoque, J.M., Díez-Herrero, A., and Calvo, C.: Triggering threshold
- 477 precipitation and soil hydrological characteristics of shallow landslides in granitic landscapes,
- 478 Geomorphology, 133(3), 178–189, https://doi.org/10.1016/j.geomorph.2011.05.018, 2011.

- 479 Saroglou, C.: GIS-based rockfall susceptibility zoning in Greece, Geosciences, 9(4), 163,
- 480 https://doi.org/10.3390/geosciences9040163, 2019.
- 481 Sarro, R., Riquelme, A., García-Davalillo, J., Mateos, R., Tomás, R., and Pastor, J.: Rockfall
- 482 Simulation Based on UAV Photogrammetry Data Obtained during an Emergency Declaration:
- 483 Application at a Cultural Heritage Site, Remote Sens., 10(12), 1923,
- 484 https://doi.org/10.3390/rs10121923, 2108.
- 485 Segoni, S., Piciullo, L., and Gariano, S.L.: A review of the recent literature on rainfall thresholds for
- 486 landslide occurrence, Landslides, 15, 1483–1501, https://doi.org/10.1007/s1034 6-018-0966-4, 2018.
- Sengupta, A., Gupta, S., and Anbarasu, K.: Rainfall thresholds for the initiation of landslide at Lanta
- 488 Khola in north Sikkim, India, Nat. Hazards, 52, 31–42, https://doi.org/10.1007/s11069-009-9352-9,
- 489 2010.
- 490 Staley, D.M., Kean, J.W., Cannon, S.H., Schmidt, K.M., and Laber, J.L.: Objective definition of
- 491 rainfall intensity—duration thresholds for the initiation of post-fire debris flows in southern California,
- 492 Landslides, 10, 547–562, https://doi.org/10.1007/s10346-012-0341-9, 2013.
- 493 Temiño, J.Y., Rodrìguez-Peces, M.J., Marchesini, S., Leyva, S., and Dìaz-Hernàndez, J.L.:
- 494 Amplification of the destructive effects of rock falls by sliding on volcanic soils: examples from the
- 495 Anaga Massif (Tenerife Island, Spain), in: Margottini, C., Canuti, P., and Sassa, K. (Eds.) Landslide
- 496 Science and Practice, Springer-Verlag Berlin Heidelberg, Vol. 1, pp. 191-195,
- 497 https://doi.org/10.1007/978-3-642-31325-7_25, 2013a.
- 498 Temiño, J.Y., Rodrìguez-Peces, M.J., Sànchez, N., Galindo, I., and del Potro, R.: Geomorphologic
- 499 evidences of flank instabilities in the eastern sector of the Tejeda volcano (Canary Islands, Spain)
- during the Quaternary, in: Margottini, C., Canuti, P., and Sassa, K. (Eds.) Landslide Science and
- Practice, Springer-Verlag Berlin Heidelberg, Vol. 7, pp. 65–72, https://doi.org/10.1007/978-3-642-
- 502 31313-4_9, 2013b.
- 503 Troll, V.R., and Carracedo, J.C.: The Geology of the Canary Islands, Elsevier, pp. 636. ISBN 978-0-
- 504 12-809663-5, https://doi.org/10.1016/C2015-0-04268-X, 2016.
- 505 Urgeles, R., Canals, M., and Masson, D.G.: 2001 Flank stability and processes off the western Canary
- 506 Islands: a review from El Hierro and La Palma, Sci. Mar., 65(1), 21-31,
- 507 https://doi.org/10.3989/scimar.2001.65s121, 2001.

- Valenzuela, P., Domínguez-Cuesta, M.J., Mora García, M.A., and Jiménez-Sánchez, M.: Rainfall
- thresholds for the triggering of landslides considering previous soil moisture conditions (Asturias,
- 510 NW Spain), Landslides, 15, 273–282, https://doi.org/10.1007/s10346-017-0878-8, 2018.
- Valenzuela, P., Zêzere, J.L., Domínguez-Cuesta, M.J., and Mora García, M.A.: Empirical rainfall
- 512 thresholds for the triggering of landslides in Asturias (NW Spain), Landslides, 16, 1285–1300,
- 513 https://doi.org/10.1007/s10346-019-01170-2, 2019.
- Wieczorek, G.F., and Jaeger, S.: Triggering mechanisms and depositional rates of postglacial slope
- 515 movement processes in the Yosemite Valley, California, Geomorphology, 15, 17–31,
- 516 https://doi.org/10.1016/0169-555X(95)00112-I, 1996.
- 517 WMO World Meteorological Organization: Calculation of Monthly and Annual 30-Year Standard
- 518 Normals, WMO/TD No. 341, WCDP-No. 10, Geneva, 1989.
- Zêzere, J.L., Vaz, T., Pereira, S., Oliveira, S.C., Margues, R., and Garcia, R.A.C.: Rainfall thresholds
- for landslide activity in Portugal: a state of the art, Environ. Earth Sci., 73(6), 2917–2936,
- 521 https://doi.org/10.1007/s12665-014-3672-0, 2015.

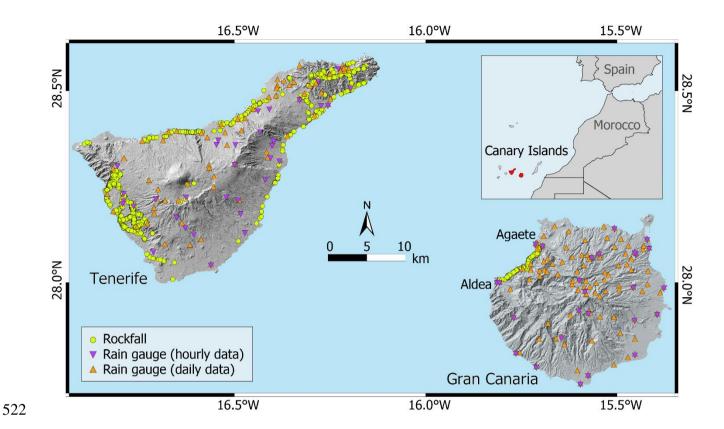


Figure 1. GC and TEN test sites. Location of the rain gauges providing hourly (purple triangles) and daily (orange triangles) rainfall measurements, and of rockfalls used for threshold calculations (light green dots). Hillshade derived from MDT05 2009 CC-BY 4.0 scne.es.





Figure 2. Photos of rockfalls in GC (a) and TEN (b) test sites (credits R.M. Mateos and S. Leyva Campos).

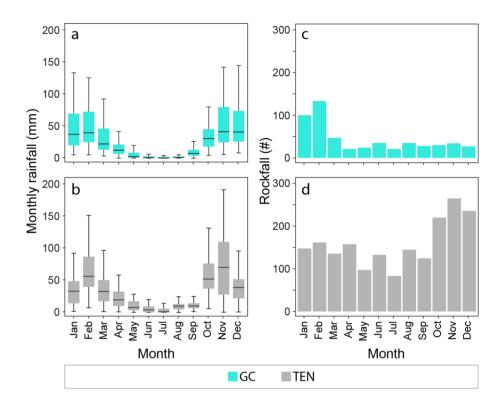


Figure 3. Comparison between monthly rainfall and rockfall occurrence. (a, b) Annual variation of monthly rainfall measures in GC (cyan) and TEN (grey). The whiskers show 1.5 times the interquartile range. (c, d) Number of rockfalls per month in the two test sites.

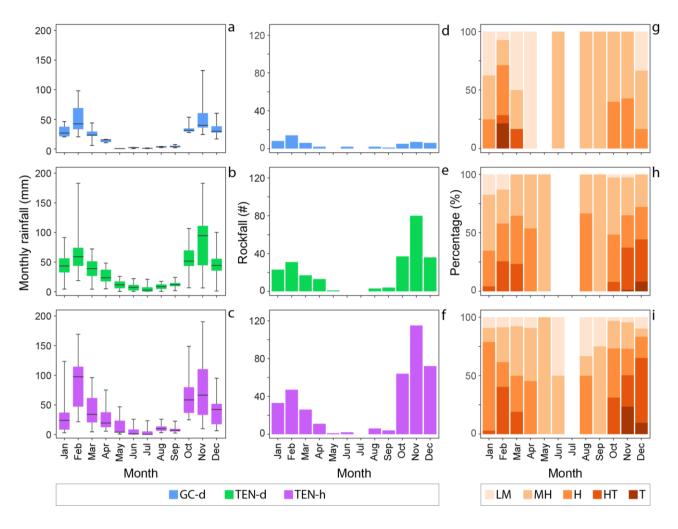


Figure 4. Comparison between monthly rainfall and rainfall-induced rockfalls and Alpert classification. (a, b, c) Annual variation of monthly rainfall measures in the test sites. Legend: GC-d, daily rainfall data in GC test site; TEN-d, daily rainfall data in the TEN test site; TEN-h, hourly rainfall data in TEN test site. (d, e, f) Number of rainfall-induced rockfalls per month. (g, h, i) Cumulated percentage of rainfall events per month classified according to Alpert et al. (2002). Legend: LM, light-moderate ($4 < E_d \le 16$ mm); MH, moderate-heavy ($16 < E_d \le 32$ mm); H, heavy ($32 < E_d \le 64$ mm); HT, heavy-torrential ($64 < E_d \le 128$ mm); T, torrential ($E_d > 128$ mm).

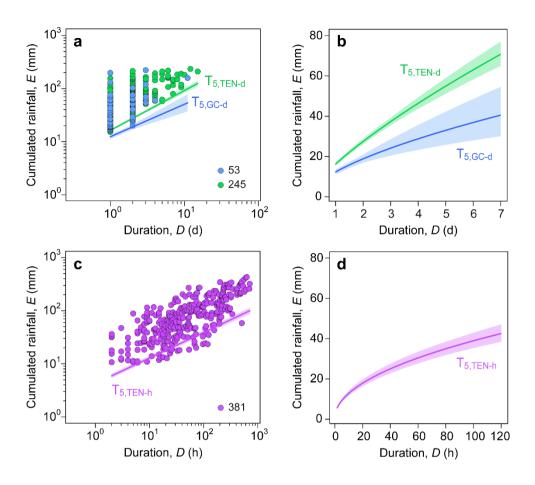


Figure 5. Rainfall thresholds for the possible rockfall occurrence in the two test sites. (a) Rainfall duration D (x-axis, in days) and cumulated event rainfall E (y-axis, in mm) conditions that have produced rockfalls in GC (53 blue dots) and TEN (245 green dots) test sites, respectively. Green and blue curves are the 5% power law thresholds ($T_{5,TEN-d}$, $T_{5,GC-d}$). (b) 5% daily ED thresholds for GC and TEN in linear coordinates, in the range of durations $1 \le D \le 7$ days. (c) Rainfall duration D (x-axis, in hours) and cumulated event rainfall E (y-axis, in mm) conditions that have produced rockfalls in TEN (381 purple dots) test site. Purple curve is the 5% power law threshold ($T_{5,TEN-h}$). (d) 5% hourly ED thresholds for GC and TEN in linear coordinates, in the range of durations $1 \le D \le 120$ hours.

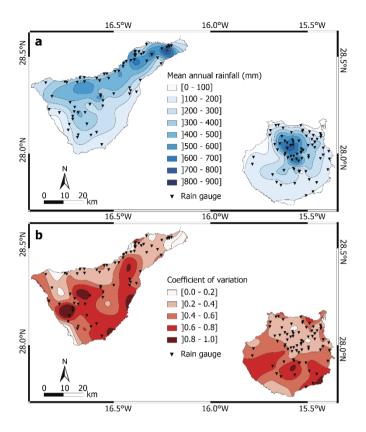


Figure 6. Maps of (a) mean annual rainfall and (b) of its coefficient of variation. The rain gauges used for these analysis (cf. Table 1) are also shown.

Table 1. Summary of the three available rain gauge networks (CIAGC, AEMET, SIAR) in the two test sites (GC and TEN) i.e., network name, network operating time period, temporal resolution, test site, number of used rain gauges, their average operating time, and the use of data.

Network	Period	Temporal resolution	Test site	Rain gauges (#)	Average operating time (year)	Data application	
CIAGC	Jan 2010 - Dec 2017	daily	GC	13	8.0		
AEMET	Jan 1951 - May 2019	uarry	GC	92	41.8	Thresholds	
	Oct 1997 - May 2019	hountr	GC	25	16.5		
	Jan 2010 - Mar 2018	hourly	TEN	34	5.5		
	Jan 2010 - May 2018	daily	TEN	66	8.2		
	Jan 2000 - Dec 2019	yearly	GC	67	15.2		
			TEN	58	13.8	MAR	
SIAR	Jan 1999 - Dec 2019	monthly	GC	5	18.2	IVIAK	
			TEN	9	15.1		

Table 2. Summary of the number (#) and percentage (%) of MPRC in the categories proposed by Alpert et al. (2002), in the two test sites.

Catagory	F. (mm)	GC-d		TEN-d		TEN-h	
Category	E _d (mm)	#	%	#	%	#	%
Light (L)	$E_{\rm d} \leq 4$	0	0	0	0	0	0
Light-moderate (LM)	$4 < E_{\rm d} \le 16$	11	20.7	11	4.5	28	7.3
Moderate-heavy (MH)	$16 < E_{\rm d} \le 32$	23	43.4	92	37.5	86	22.6
Heavy (H)	$32 < E_{\rm d} \le 64$	14	26.4	80	32.7	117	30.7
Heavy-torrential (HT)	$64 < E_{\rm d} \le 128$	2	3.8	58	23.7	116	30.5
Torrential (T)	$E_{\rm d} > 128$	3	5.7	4	1.6	34	8.9

Table 3. *ED* rainfall thresholds at different non-exceedance probabilities (1%, 5%, 10%, 20%, 35% and 50%) for the GC and TEN test sites. The number of MPRC and the duration range of each threshold are also reported.

Threshold	Number of	Threshold equation	Duration		
name	MPRC	Tineshold equation	range		
T _{1,GC-d}		$E = (8.3 \pm 1.0) \times D^{(0.62 \pm 0.10)}$			
$T_{5,GC\text{-}d}$		$E = (12.3 \pm 1.2) \times D^{(0.62 \pm 0.10)}$			
$T_{10,GC\text{-}d}$	52	$E = (15.1 \pm 1.4) \times D^{(0.62 \pm 0.10)}$	1 11 days		
$T_{20,GC\text{-}d}$	53	$E = (19.5 \pm 1.8) \times D^{(0.62 \pm 0.10)}$	1-11 days		
$T_{35,GC\text{-}d}$		$E = (25.5 \pm 2.5) \times D^{(0.62 \pm 0.10)}$			
$T_{50,GC\text{-}d}$		$E = (31.9 \pm 3.6) \times D^{(0.62 \pm 0.10)}$			
T _{1,TEN-d}	*	$E = (11.6 \pm 0.6) \times D^{(0.75 \pm 0.05)}$			
$T_{5,TEN\text{-}d}$		$E = (16.3 \pm 0.8) \times D^{(0.75 \pm 0.05)}$			
$T_{10, TEN\text{-}d}$	245	$E = (19.6 \pm 0.8) \times D^{(0.75 \pm 0.05)}$	1 15 days		
$T_{20,TEN\text{-}d}$	243	$E = (24.4 \pm 1.0) \times D^{(0.75 \pm 0.05)}$	1-15 days		
$T_{35,TEN\text{-}d}$		$E = (30.6 \pm 1.4) \times D^{(0.75 \pm 0.05)}$			
$T_{50, TEN\text{-}d}$		$E = (37.1 \pm 1.8) \times D^{(0.75 \pm 0.05)}$			
T _{1,TEN-h}		$E = (2.8 \pm 0.3) \times D^{(0.48 \pm 0.02)}$			
$T_{5,TEN\text{-}h}$		$E = (4.3 \pm 0.4) \times D^{(0.48 \pm 0.02)}$			
$T_{10, TEN\text{-}h}$	201	$E = (5.3 \pm 0.5) \times D^{(0.48 \pm 0.02)}$	2.712 hove		
$T_{20, TEN\text{-}h}$	381	$E = (6.9 \pm 0.6) \times D^{(0.48 \pm 0.02)}$	2-712 hours		
$T_{35,TEN\text{-}h}$		$E = (9.1 \pm 0.7) \times D^{(0.48 \pm 0.02)}$			
T _{50,TEN-h}		$E = (11.4 \pm 1.0) \times D^{(0.48 \pm 0.02)}$			