



1 **Intra-annual variability of the Western Mediterranean Oscillation (WeMO)**  
2 **and occurrence of extreme torrential rainfall in Catalonia (NE Iberia)**

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9

10 **Abstract**

11 In previous studies the Western Mediterranean Oscillation index (WeMOi) at daily  
12 resolution has proven to constitute an effective tool for analysing the occurrence  
13 of episodes of torrential rainfall over eastern Spain. The Western Mediterranean  
14 region is therefore a very sensitive area, since climate change can enhance these  
15 weather extremes. In the present study we selected the extreme torrential  
16 episodes ( $\geq 200$  mm in 24 hours) that took place in Catalonia (NE Iberia) during  
17 the 1951-2016 study period (66 years). We computed daily WeMOi values and  
18 constructed WeMOi calendars. Our principal results reveal the occurrence of 50  
19 episodes (0.8 cases per year), mainly concentrated in the autumn months. We  
20 inferred a threshold of WeMOi  $\leq -2$  to define an extreme negative WeMO phase  
21 at daily resolution. Most of the 50 episodes (60%) in the study area occurred on  
22 days presenting an extreme negative WeMOi value. Specifically, the most  
23 negative WeMOi values are detected in autumn, during the second 10-day period  
24 of October (11th-20th), coinciding with the highest frequency of extreme torrential  
25 events. On comparing the subperiods, we observed a statistically significant  
26 decrease in WeMOi values in all months, particularly in late October, and in  
27 November and December. No changes in the frequency of these extreme  
28 torrential episodes were observed between both subperiods; in contrast, a  
29 displacement of the episodes is detected from early to late autumn.

30 **Keywords**

31 Catalonia, climate variability, torrential rainfall, WeMO.

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## 33 1. Introduction

34 The Mediterranean seasonal rainfall regime is characterised by rainy winters and  
35 dry summers, linked to the westerly atmospheric circulation in winter and to the  
36 subtropical anticyclone belt in summer. Nevertheless, in some regions of the  
37 Mediterranean basin, the seasonal rainfall regime differs from the typically  
38 Mediterranean one; for example, most of eastern Iberia (Spain) displays a  
39 seasonal precipitation maximum in autumn, and a secondary one in spring (De  
40 Luis *et al.*, 2010; González-Hidalgo *et al.*, 2011). This bimodal rainfall pattern is  
41 recorded in few regions of the world. It only occurs over approximately 7% of the  
42 global land surface, and is commonly associated with locations within the tropics  
43 (Knoben *et al.*, 2019). This bimodal behaviour in eastern Spain is mainly due to  
44 the physical geographic complexity of the Iberian Peninsula, which comprises  
45 several mountain ranges, all of which present different slope orientations.  
46 Furthermore, the Mediterranean Sea is practically cut off from other water bodies,  
47 which favours higher surface sea temperatures (SST) than in the Atlantic at the  
48 same latitude in summer and autumn. This contributes to the development of high  
49 vertical gradients of air temperature in some months over the Mediterranean  
50 basin (Estrela *et al.*, 2008; Pérez-Zanón *et al.*, 2018). These physical  
51 geographical factors give rise to a high concentration of daily precipitation in the  
52 Mediterranean basin, i.e. torrential rainfall events, above all, in the Western  
53 Mediterranean (Beguería *et al.*, 2011; Cortesi *et al.*, 2012); all this reveals the  
54 need for water management in Spain to be based upon rainfall variability rather  
55 than on the rainfall mean (Lopez-Bustins, 2018). Heavy rainfall in the Western  
56 Mediterranean is mainly centred in eastern Spain, the south of France and the  
57 region of Liguria (NW Italy) (Peñarrocha *et al.*, 2002). These torrential events can  
58 cause dangerous floods and have an important social and economic impact on  
59 the Mediterranean regions, e.g. in eastern Spain (Olcina *et al.*, 2016; Kreibich *et al.*,  
60 2017; Nakamura and Llasat, 2017; Martin-Vide and Llasat, 2018) and  
61 particularly in southern Spain (Gil-Guirado *et al.*, 2019). Climatological studies on  
62 torrential rainfall frequency and intensity are therefore relevant with regard to  
63 improving emergency plans and mitigating flood damage. Extreme precipitation  
64 is expected to increase with global warming as a result of a greater atmospheric  
65 water content (Papalexiou and Montanari, 2019); for instance, extreme peak river



66 flows are predicted to increase in Southern Europe during the current century  
67 (Alfieri *et al.*, 2015), and the frequency of heavy rainfall events is projected to be  
68 higher for the 2011-2050 period (Barrera-Escoda *et al.*, 2014).

69 Previous studies have associated extreme daily rainfall events in Spain with  
70 synoptic patterns (Martin-Vide *et al.*, 2008; Peña *et al.*, 2015); these studies have  
71 addressed several different tropospheric levels (Romero *et al.*, 1999; Merino *et al.*  
72 *et al.*, 2016; Pérez-Zanón *et al.*, 2018). Furthermore, many studies have also  
73 statistically correlated several teleconnection indices (El Niño Southern  
74 Oscillation, North Atlantic Oscillation, Arctic Oscillation, Mediterranean  
75 Oscillation, Western Mediterranean Oscillation, etc.) with rainfall series for the  
76 Iberian Peninsula at different timescales (Rodó *et al.*, 1997; Rodríguez-Puebla *et al.*  
77 *et al.*, 2001; Trigo *et al.*, 2004; Lopez-Bustins *et al.*, 2008; González-Hidalgo *et al.*,  
78 2009; Ríos-Cornejo *et al.*, 2015a; Merino *et al.*, 2016). Among these, the Western  
79 Mediterranean Oscillation (WeMO) was found to be the index most statistically  
80 and significantly correlated with annual, monthly and daily precipitation on the  
81 littoral fringe of eastern Spain (Martin-Vide and Lopez-Bustins, 2006; González-  
82 Hidalgo *et al.*, 2009). The daily timescale of the WeMO index (WeMOi) could  
83 constitute a potential tool for analysing the frequency of torrential events in some  
84 regions of the Western Mediterranean basin.

85 The present study provides an exhaustive inventory of the heaviest rainfall events  
86 in Catalonia (NE Iberia) over the last few decades (1951-2016) in order to furnish  
87 a better understanding of their temporal distribution. Moreover, we will analyse  
88 changes in frequency according to subperiods, since the Western Mediterranean  
89 basin is a global warming hotspot, where a decrease in mean annual precipitation  
90 is expected for the following decades, particularly in summer, together with a  
91 potential rise in storm-related precipitation and drought duration (Christensen *et al.*  
92 *et al.*, 2013; Barrera-Escoda *et al.*, 2014; Cramer *et al.*, 2018; Greve *et al.*, 2018).  
93 The main aim of the study involves establishing a period of high potential  
94 torrentiality in Catalonia at daily resolution, considering the catalogue of extreme  
95 torrential events in the area analysed. Most studies delimit the wet season of a  
96 region within one or several months (Kottek *et al.*, 2006), and do not employ a  
97 smaller timescale than the monthly one. Therefore, the present research attempts  
98 to go beyond the monthly timescale in order to determine the period with the



99 highest accumulation of heavy rainfall according to fortnights and 10-day periods.  
100 The intra-annual variability of the daily WeMOi values may help to establish the  
101 period with the highest propensity for torrential events in Catalonia.

102 In section 2, we describe the main orographic and pluviometric features of the  
103 study area. The methods followed to calculate daily WeMOi values and construct  
104 the WeMOi calendar are explained in section 3. In section 4, the results of the  
105 intra-annual variability of torrential episodes and WeMOi values are analysed and  
106 discussed; additionally, we have demonstrated the usefulness of the WeMOi  
107 calendar for pinpointing the time of year presenting the heaviest rainfall events in  
108 another region of the Western Mediterranean, the south of France. Finally, in  
109 section 4 we derive the conclusions.

## 110 **2. Study area**

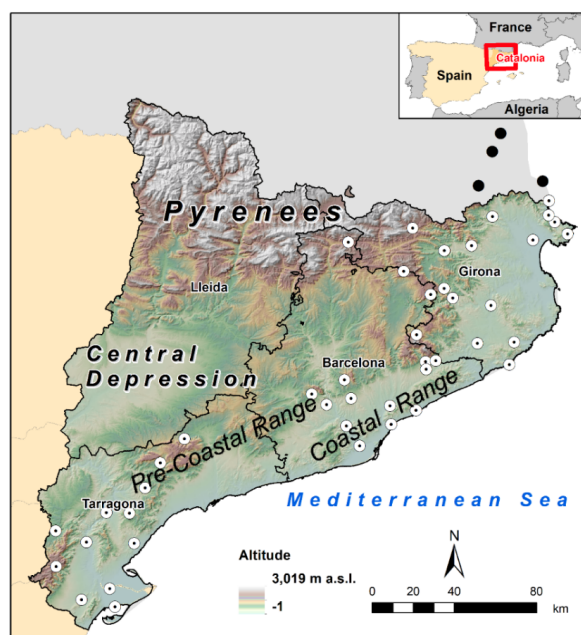
111 Catalonia covers an area of 32,100 km<sup>2</sup> in northeast Spain; it is physically  
112 separated from France by the Pyrenees (Figure 1). Altitude ranges from 0 (littoral)  
113 to 3,100 m (northwestern Pyrenees) a.s.l. The Coastal and Pre-Coastal ranges,  
114 with an altitude ranging from 500 m.a.s.l. to 1,700 m.a.s.l., present a SW-NE  
115 orientation. On the western border, the Central Depression is approximately 200-  
116 300 m.a.s.l., constituting the driest part of the study area (350 mm annual-mean  
117 precipitation) (Figure 2a). The wettest part of Catalonia is located in the Pyrenees,  
118 with an annual-mean rainfall over 1,200 mm. In general terms, southern Lleida  
119 and Barcelona, as well as almost the entire province of Tarragona, make up the  
120 dry part of Catalonia (<700 mm). The rainy part of Catalonia (≥700 mm)  
121 comprises the province of Girona and the northern halves of the provinces of  
122 Lleida and Barcelona.

123 Catalonia's complex orography, as well as the fact that it falls under the influence  
124 of the Atlantic Ocean and the Mediterranean Sea, endow it with a highly  
125 heterogeneous spatial distribution of seasonal precipitation regimes throughout  
126 the study area. Using 70 monthly precipitation series (1950-2015) provided by  
127 the Meteorological Service of Catalonia (SMC, 2016), we ascertained that, of the  
128 total of 24 possible permutations between winter, spring, summer and autumn as  
129 dominant and subdominant precipitation seasons, 7 of these are detected in  
130 Catalonia (Figure 2b) (Martin-Vide and Raso-Nadal, 2008). A clear predominance



131 of autumn rainfall can be observed, followed by spring precipitation, especially in  
132 the coastal zone. The driest season on the coast is summer; however, inland the  
133 driest time of year is winter. Many areas of the Pyrenees, above all, in the east,  
134 exhibit their maxima in summer as a result of convective rainfall.

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137 Figure 1. Location of Catalonia (NE Spain) within Europe, altitude and 4 provinces.  
138 The white dots indicate the 43 different weather stations that have recorded the  
139 highest rainfall amount during an extreme torrential event at least once in  
140 Catalonia. The black dots show the four weather stations located in the south of  
141 France. The base map was provided by Cartographic and Geological Institute of  
142 Catalonia.

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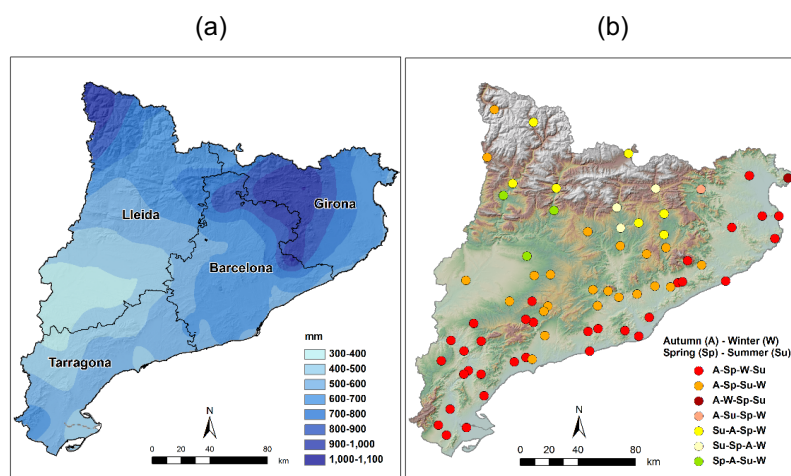
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Figure 2. (a) Average annual precipitation (mm) in Catalonia for the 1950-2015 period. (b) Seasonal precipitation regimes for 70 weather stations during the 1950-2015 period. The base map was provided by Cartographic and Geological Institute of Catalonia.

### 161 3. Data and Methods

#### 162 3.1. Selection of torrential events

163 Several studies have selected the torrential rainfall events in Spain based on the  
164 threshold of 100 mm in 24 h (Pérez-Cueva, 1994; Martin-Vide and Llasat, 2000;  
165 Armengot, 2002; Riesco and Alcover, 2003; Martin-Vide *et al.*, 2008). Herein we  
166 chose the extreme torrential episodes ( $\geq 200$  mm in 24 h) (Lopez-Bustins *et al.*,  
167 2016) that took place over Catalonia during the 1951-2016 study period (66  
168 years). We consider the threshold of 200 mm in 24 h to present a natural risk in  
169 most cases, with significant consequences. Episodes involving  $\geq 100$  mm in 24 h  
170 are more frequent, but sometimes have no direct impact, or quite a negligible  
171 effect, because other factors are the main drivers of floods, e.g. rainfall duration  
172 (Jang, 2015), initial soil moisture conditions and hydrological parameters  
173 (Norbiato *et al.*, 2008; Martina *et al.*, 2009).



174 In order to select the extreme torrential events, we considered all available rainfall  
175 data sources in Catalonia (Meteorological Service of Catalonia, Spanish National  
176 Meteorological Agency, Catalan Water Agency and Ebro Hydrographic  
177 Confederation). Data sources were furnished both by automatic (semi-hourly  
178 data) and manual (one register per day) weather stations. We considered the  
179 pluviometric day as 7-7 UTC in both types of observatories in order to ensure a  
180 homogeneous criterion when selecting episodes along the whole study period  
181 and analysing any temporal changes in their frequency. Manual weather stations,  
182 which record the pluviometric day as 7-7 UTC, were the only precipitation data  
183 source in Catalonia until the end of the 1980s. Therefore, we adjusted the  
184 episodes provided by the automatic stations to the pluviometric day 7-7 UTC. An  
185 exhaustive spatial and temporal check was performed among the extreme  
186 torrential episodes identified. We tested the reliability of the events considering  
187 the daily precipitation recorded in neighbouring stations and examining the  
188 original handwritten observation cards. Furthermore, we fixed some episodes  
189 registered the day after the pluviometric day from manual weather stations, and  
190 we eliminated events derived from rainfall accumulation of over one day. The  
191 additional pluviometric data used in section 4.4. were provided by Météo-France.

192 The catalogue of extreme torrential events in Catalonia addresses the following  
193 columns: date, maximum precipitation in 24 h, location, province and daily  
194 WeMOi value. Several observatories in Catalonia can occasionally register  $\geq 200$   
195 mm in 24 h on one same date, but only the highest amount was taken into  
196 account. Finally, we obtained 50 extreme torrential events for consideration in the  
197 present study (Table 1). 32 out of the 50 episodes (64%) have a decimal place of  
198 0, and 10 out of the 50 episodes (20%) present a decimal place of 5. Most of  
199 these episodes were registered by manual weather stations prior to the 1990s.  
200 This is known as the rounding effect (Wergen *et al.*, 2012): a weather observer  
201 rounds off the daily precipitation accumulation value during heavy rainfall events.  
202 This effect has no influence on the results of the present research.

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| Date                     | Max RR (mm)  | Location                           | Province         | WeMOi value  |
|--------------------------|--------------|------------------------------------|------------------|--------------|
| <b>13 October 1986</b>   | <b>430.0</b> | <b>Cadaqués</b>                    | <b>Girona</b>    | <b>-2.22</b> |
| <b>11 April 2002</b>     | <b>367.5</b> | <b>Darnius</b>                     | <b>Girona</b>    | <b>-3.85</b> |
| 20 September 1971        | 308.0        | Esparreguera                       | Barcelona        | -1.75        |
| 20 September 1972        | 307.0        | Sant Carles de la Ràpita           | Tarragona        | -1.58        |
| <b>09 October 1994</b>   | <b>293.0</b> | <b>Cornudella de Montsant</b>      | <b>Tarragona</b> | <b>-2.88</b> |
| 03 October 1987          | 291.0        | Castelló d'Empúries                | Girona           | -1.96        |
| <b>22 September 1971</b> | <b>285.0</b> | <b>Cadaqués</b>                    | <b>Girona</b>    | <b>-2.19</b> |
| <b>19 October 1977</b>   | <b>276.0</b> | <b>Cadaqués</b>                    | <b>Girona</b>    | <b>-2.80</b> |
| <b>21 September 1971</b> | <b>275.0</b> | <b>Santa Maria de Palautordera</b> | <b>Barcelona</b> | <b>-2.21</b> |
| <b>18 October 1977</b>   | <b>271.8</b> | <b>Camprodon</b>                   | <b>Girona</b>    | <b>-2.21</b> |
| <b>21 October 2000</b>   | <b>270.0</b> | <b>Falset</b>                      | <b>Tarragona</b> | <b>-2.26</b> |
| <b>07 November 1982</b>  | <b>266.0</b> | <b>la Pobla de Lillet</b>          | <b>Barcelona</b> | <b>-5.56</b> |
| 12 October 2016          | 257.0        | Vilassar de Mar                    | Barcelona        | -1.86        |
| <b>05 March 2013</b>     | <b>253.5</b> | <b>Darnius</b>                     | <b>Girona</b>    | <b>-5.32</b> |
| <b>29 November 2014</b>  | <b>253.5</b> | <b>Parc Natural dels Ports</b>     | <b>Tarragona</b> | <b>-4.54</b> |
| <b>16 February 1982</b>  | <b>251.2</b> | <b>Amer</b>                        | <b>Girona</b>    | <b>-2.41</b> |
| 25 September 1962        | 250.0        | Martorelles                        | Barcelona        | -1.52        |
| <b>04 November 1962</b>  | <b>248.5</b> | <b>SantLlorenç del Munt</b>        | <b>Barcelona</b> | <b>-2.79</b> |
| 02 September 1959        | 246.5        | Cadaqués                           | Girona           | -0.84        |
| <b>10 October 1994</b>   | <b>245.0</b> | <b>Beuda</b>                       | <b>Girona</b>    | <b>-2.33</b> |
| <b>22 October 2000</b>   | <b>240.0</b> | <b>Tivissa</b>                     | <b>Tarragona</b> | <b>-2.50</b> |
| <b>12 November 1999</b>  | <b>233.5</b> | <b>Castellfollit de la Roca</b>    | <b>Girona</b>    | <b>-3.00</b> |
| <b>06 January 1977</b>   | <b>233.0</b> | <b>Girona</b>                      | <b>Girona</b>    | <b>-2.22</b> |
| <b>20 December 2007</b>  | <b>230.2</b> | <b>Parc Natural dels Ports</b>     | <b>Tarragona</b> | <b>-3.54</b> |
| 06 October 1959          | 230.1        | Tossa de Mar                       | Girona           | -1.36        |
| 03 October 1951          | 230.0        | Cornellà de Llobregat              | Barcelona        | -1.02        |
| 20 September 1959        | 230.0        | Gualba de Dalt                     | Barcelona        | -1.49        |
| 11 October 1970          | 230.0        | Riudabella                         | Tarragona        | -1.61        |
| <b>23 October 2000</b>   | <b>229.0</b> | <b>Horta de Sant Joan</b>          | <b>Tarragona</b> | <b>-2.41</b> |
| <b>26 September 1992</b> | <b>226.4</b> | <b>Amposta</b>                     | <b>Tarragona</b> | <b>-2.22</b> |
| <b>04 April 1969</b>     | <b>226.0</b> | <b>Rupit</b>                       | <b>Barcelona</b> | <b>-2.21</b> |
| <b>12 November 1988</b>  | <b>225.0</b> | <b>Corbera de Llobregat</b>        | <b>Barcelona</b> | <b>-2.76</b> |
| 11 October 1962          | 223.0        | Sils                               | Girona           | -1.20        |
| 20 November 1956         | 221.0        | Cornellà de Llobregat              | Barcelona        | -0.45        |
| <b>06 November 1983</b>  | <b>220.0</b> | <b>Terrassa</b>                    | <b>Barcelona</b> | <b>-2.34</b> |
| <b>19 October 1994</b>   | <b>220.0</b> | <b>el Port de Llançà</b>           | <b>Girona</b>    | <b>-2.36</b> |
| 31 July 2002             | 218.2        | Badalona                           | Barcelona        | -0.13        |
| 13 September 1963        | 217.5        | l'Ametlla de Mar                   | Tarragona        | -1.14        |
| 19 September 1971        | 217.0        | Xerta                              | Tarragona        | -0.97        |
| 17 September 2010        | 216.8        | l'Ametlla de Mar                   | Tarragona        | -0.60        |
| <b>17 October 2003</b>   | <b>213.0</b> | <b>Vidrà</b>                       | <b>Girona</b>    | <b>-2.48</b> |
| 09 June 2000             | 210.0        | el Bruc                            | Barcelona        | -0.23        |
| 31 August 1975           | 208.5        | Santa Agnès de Solius              | Girona           | -0.15        |
| <b>29 January 1996</b>   | <b>206.5</b> | <b>Fogars de Montclús</b>          | <b>Barcelona</b> | <b>-2.37</b> |
| 09 October 1971          | 204.0        | Miravet                            | Tarragona        | -0.86        |
| <b>26 December 2008</b>  | <b>202.5</b> | <b>Darnius</b>                     | <b>Girona</b>    | <b>-2.84</b> |
| <b>07 May 2002</b>       | <b>200.8</b> | <b>Godall</b>                      | <b>Tarragona</b> | <b>-2.47</b> |
| <b>07 October 1965</b>   | <b>200.0</b> | <b>les Planes d'Hostoles</b>       | <b>Girona</b>    | <b>-2.12</b> |
| 27 October 1989          | 200.0        | el Port de la Selva                | Girona           | -1.90        |
| <b>01 November 1993</b>  | <b>200.0</b> | <b>Portbou</b>                     | <b>Girona</b>    | <b>-2.57</b> |





210 Table 1. Catalogue of extreme torrential events ( $\geq 200$  mm in 24 h, 7-7 UTC) in  
211 Catalonia (NE Iberia) during the 1951-2016 period. Max RR is the highest rainfall  
212 accumulation of the episode. The events are classified according to the extreme  
213 negative Western Mediterranean Oscillation (WeMO) phase (**bold**), the negative  
214 WeMO phase and the slight negative WeMO phase (*italics*).

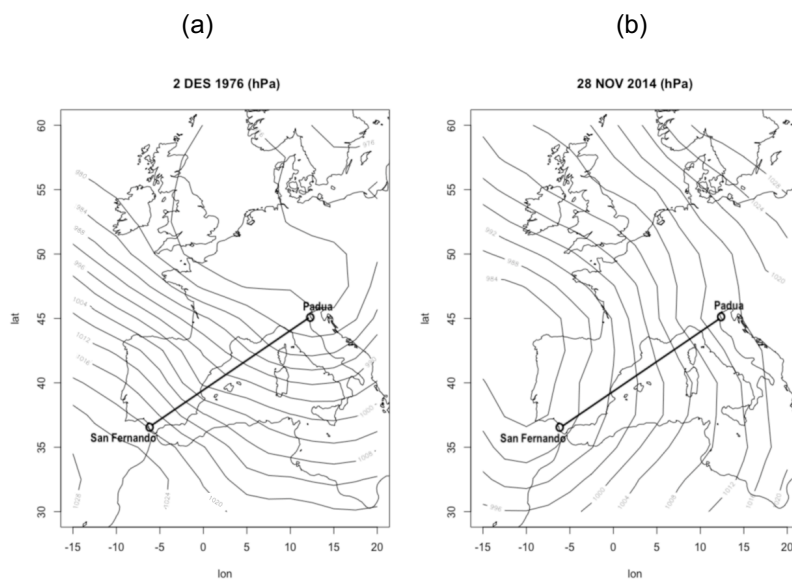
### 215 3.2. Daily WeMO<sub>i</sub> values

216 The WeMO<sub>i</sub> is a regional teleconnection index defined within the Western  
217 Mediterranean basin (Martin-Vide and Lopez-Bustins, 2006) and already used in  
218 a wider range of studies (Azorin-Molina and Lopez-Bustins, 2008; Vicente-  
219 Serrano *et al.*, 2009; Caloiero *et al.*, 2011; El Kenawy *et al.*, 2012; Coll *et al.*,  
220 2014; Ríos-Cornejo *et al.*, 2015b; Lana *et al.*, 2017; Jghab *et al.*, 2019). WeMO<sub>i</sub>  
221 values are computed by means of surface pressure data from the San Fernando  
222 (SW Spain) and Padua (NE Italy) weather stations (Figure 3); the synoptic  
223 window 30°-60°N - 15°W-20°E is found to best represent WeMO phases (Arbiol-  
224 Roca *et al.*, 2018). The positive phase of the WeMO corresponds to the  
225 anticyclone over the Azores encompassing the southwest quadrant of the Iberian  
226 Peninsula and low pressures in the Gulf of Genoa (Figure 3a); its negative phase  
227 coincides with an anticyclone located over Central or Eastern Europe and a low-  
228 pressure centre, often cut off from the northern latitudes, within the framework of  
229 the Iberian southwest (Figure 3b). Martin-Vide and Lopez-Bustins (2006) found  
230 that the WeMO<sub>i</sub> was significantly and statistically correlated with rainfall over  
231 areas that were weakly influenced by the North Atlantic Oscillation (NAO): in the  
232 northernmost and easternmost parts of Spain; precipitation over the Cantabrian  
233 fringe (northern Spain) was strongly and positively correlated with the WeMO<sub>i</sub>,  
234 and precipitation over the Spain's eastern façade was strongly and negatively  
235 correlated with the WeMO<sub>i</sub>.

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250 Figure 3. (a) Most extreme positive phase of the Western Mediterranean Oscillation  
251 (WeMO) in a daily synoptic situation during the 1951-2016 period (2nd December  
252 1976). (b) Most extreme negative WeMO phase in a daily synoptic situation during  
253 the 1951-2016 period (28th November 2014). Data source: NCEP Reanalysis data  
254 provided by the NOAA/OAR/ESRL PSD, Boulder, Colorado, USA.

255 Application of the daily WeMOi is a methodological contribution by Martin-Vide  
256 and Lopez-Bustins (2006). It converts the low-frequency feature of the  
257 teleconnection patterns into a high-frequency mode. It is suitable for application  
258 both to the regional scale of the WeMO teleconnection pattern and the lesser  
259 variability of atmospheric pressure at Mediterranean latitudes. Patterns have  
260 rarely been used at daily resolution (Baldwin and Dunkerton, 2001; Beniston and  
261 Jungo, 2002; Azorin-Molina and Lopez-Bustins, 2008; Liu *et al.*, 2018). The  
262 method selected consists of previously standardising each series of the dipole. It  
263 is necessary to use the mean and standard deviation of the 1961-1990 reference  
264 period of all days of the year (from 1st January 1961 to 31st December 1990).



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266 For example, the WeMOi on January 1<sup>st</sup> 1981

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$$268 \quad Z \text{ WeMOi 1st Jan 1981} = \frac{P \text{ 1st Jan 1981 SF} - \bar{X} \text{ 1961\_1990 SF}}{S \text{ 1961\_1990 SF}} - \frac{P \text{ 1st Jan 1981 PD} - \bar{X} \text{ 1961\_1990 PD}}{S \text{ 1961\_1990 PD}},$$

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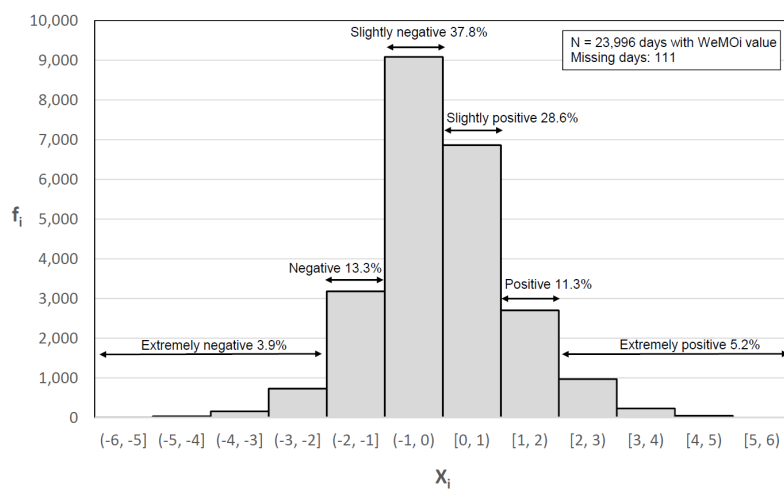
270 where P is pressure, SF, San Fernando, PD, Padua,  $\bar{X}$ , mean, and S, standard  
271 deviation.

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273 This calculation method, which considers all days of the year in the reference  
274 period, enables all Mediterranean flows (negative WeMO phase) to be detected,  
275 even if they are very weak. Otherwise, these moderate Mediterranean winds  
276 would not be detected in autumn, since the WeMOi means are clearly negative  
277 during this season. Likewise, the weak Mediterranean flows would be  
278 overestimated in winter due to the high WeMOi mean during the coldest months.  
279 According to previous studies (Martin-Vide and Lopez-Bustins, 2006; Azorin-  
280 Molina and Lopez-Bustins, 2008), in the histogram of daily WeMOi frequencies,  
281 WeMOi values between -1.00 and 1.00 are considered to constitute a neutral  
282 WeMO phase, values ranging from 1.00 to 1.99 are considered as a positive  
283 WeMO phase, those between -1.99 and -1.00 as a negative WeMO phase,  
284 values  $\geq 2.00$  are deemed to represent an extreme positive WeMO phase and  
285 those  $\leq -2.00$  to indicate an extreme negative WeMO phase. The most positive  
286 WeMOi value (+5.99) of the 1951-2016 study period refers to December 2nd  
287 1976 (Figure 3a), when an intense rainfall episode was recorded in the Basque  
288 Country (northern Spain) according to ECA & Dataset (Klein Tank *et al.*, 2002;  
289 Cornes *et al.*, 2018). The most negative WeMOi value (-5.97) during the 1951-  
290 2016 period corresponds to November 28th 2014 (Figure 3b), when 253.5 mm  
291 was registered in the *Parc Natural dels Ports* (Tarragona) during the following  
292 day (Table 1). Lana *et al.* (2016) studied the statistical complexity and  
293 predictability of the WeMOi and demonstrated the Gaussian distribution of this  
294 index. Most daily WeMOi values are positive (55%) and two thirds of the 23,996  
295 days displaying WeMOi values correspond to a neutral WeMO phase (Figure 4).  
296 The positive (negative) WeMO phase was detected in 16.5% (17.2%) of the total  
297 days presenting a WeMOi value. The extreme WeMOi values, both positive



298 (5.2%) and negative (3.9%), represent less than 10% of the total number of days  
299 for which WeMOi values are available.  
300



301  
302 Figure 4. Frequency histogram of all daily WeMO index (WeMOi) values during  
303 the 1951-2016 study period.

### 304 3.3. Construction of calendars

305 Constructing calendars is a common procedure in climatological studies (Soler  
306 and Martin-Vide, 2002; Azorin-Molina and Lopez-Bustins, 2008; Meseguer-Ruiz  
307 *et al.*, 2018). They enable the intra-annual variability of the climate variable to be  
308 visualised. We computed daily WeMOi values for the 1951-2016 (66 years) study  
309 period, constructing two WeMOi calendars based upon the mean values obtained  
310 for each month, a 15-day period (i.e. a fortnight) and a 10-day period; the latter  
311 period corresponds approximately to the baroclinic prediction period (Holton,  
312 2004). The first climate calendar will show the annual cycle of the WeMOi values  
313 according to months (12 values), the second will display a more detailed intra-  
314 annual oscillation with 24 values and, finally, the 36 WeMOi values derived from  
315 the 10-day calendar will enable the slightest intra-annual variations in the WeMOi  
316 to be detected. We will add to these calendars all the extreme torrential events in  
317 order to observe correspondences between WeMOi values and heavy rainfall  
318 events along the year. In order to detect any changes in the calendars throughout  
319 the study period, we consider two subperiods for the construction of two



320 additional calendars: 1951-1983 (33 years) and 1984-2016 (33 years). The mean  
321 WeMOi values according to subperiods were statistically tested to detect  
322 significant differences. This statistical significance is computed by means of a  
323 Normal distribution test according to several levels of confidence: 95.0%  
324 ( $z=1.960$ ), 99.0% ( $z=2.576$ ) and 99.9% ( $z=3.291$ ).

#### 325 **4. Results and discussion**

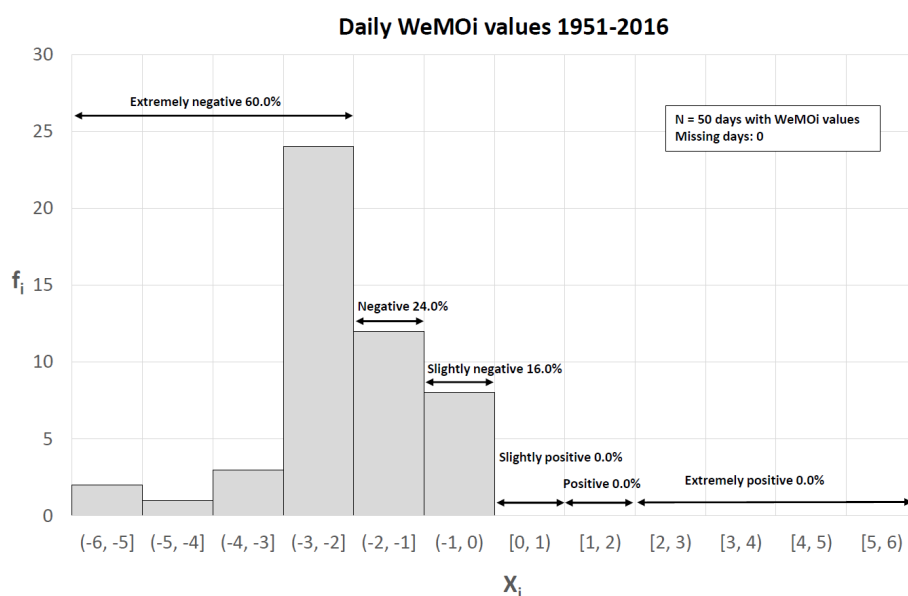
##### 326 *4.1. Frequency and temporal evolution of the extreme torrential events*

327 During the 1951-2016 period, 50 episodes presenting  $\geq 200$  mm in 24 h were  
328 detected (0.8 cases per year) in Catalonia (Table 1); these were mainly  
329 concentrated in the Eastern Pyrenees (Girona) and southern Catalonia  
330 (Tarragona) (Figure 1). In the province of Lleida no maximum values for  
331 precipitation episodes have been recorded, because this province is less  
332 influenced by easterly flows as a result of its continental features. Other parts of  
333 Iberia register a higher frequency of extreme torrential events, e.g. in the Valencia  
334 Region, eastern Spain, there were 2 cases per year during the 1971-2000 period  
335 (Riesco and Alcover, 2003). The highest frequency of torrential events ( $\geq 100$  mm  
336 in 24 h) over the Iberian Peninsula also corresponds to the Valencia Region,  
337 where more than one case per year can be recorded by one same observatory  
338 (Pérez-Cueva, 1994) and approximately 11 cases per year by all the stations in  
339 the Valencia Region (Riesco and Alcover, 2003). Catalonia exhibits a lower  
340 frequency of these torrential events (i.e.  $\geq 100$  mm in 24 h), 5-6 cases per year for  
341 the whole region (Martin-Vide and Llasat, 2000; Lopez-Bustins *et al.*, 2016). The  
342 highest rainfall amount during 7-7 UTC ever recorded in Catalonia is 430 mm.  
343 This occurred in Cadaqués (Cape Creus, in the easternmost part of the Iberian  
344 Peninsula) on October 13th 1986. It was an extraordinary episode which also  
345 affected the region of Pyrénées-Orientales (S France) (Vigneau, 1987), albeit  
346 with a lower amount of rainfall than that produced by other extreme torrential  
347 events of over 800 mm in Liguria Region (NW Italy), Valencia Region (E Spain)  
348 and this region of Pyrénées-Orientales (Peñarrocha *et al.*, 2002).

349 Most of the episodes in Catalonia (60%) (30 events) took place in an extreme  
350 negative ( $\leq -2.00$ ) WeMOi phase (Figure 5), whereas less than 4% of the total  
351 number of days with WeMOi data showed a value equal to or lower than -2.00



352 (Figure 4). Moreover, 24% (12 events) of the episodes occurred in a negative (-  
 353 2.00, -1.00] WeMO phase. The remaining 8 events (16%) took place in a slightly  
 354 negative (-1.00, 0.00) WeMOi phase. No extreme torrential episodes presenting  
 355 a positive WeMOi value occurred in Catalonia during the study period. In addition,  
 356 Martin-Vide and Lopez-Bustins (2006) found no positive daily WeMOi values for  
 357 torrential episodes ( $\geq 100$  m in 24 h) in Tortosa (south Catalonia) during the 1951-  
 358 2000 period.



359

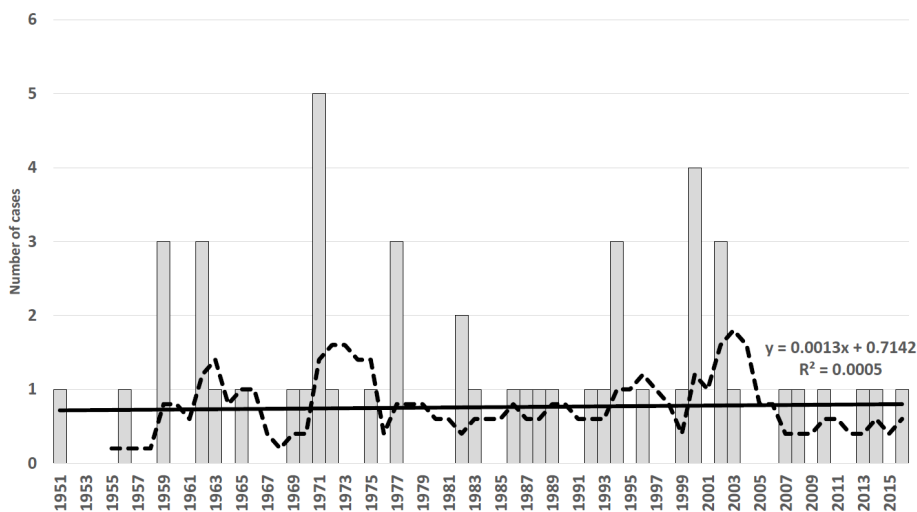
360 Figure 5. Frequency histogram of the daily WeMOi values of the 50 extreme  
 361 torrential events recorded in Catalonia from 1951 to 2016.

362 No statistical temporal trend is observed in the annual frequency of the 50  
 363 extreme torrential episodes during the 1951-2016 period (Figure 6). Most of the  
 364 years in the study period present no episodes, or only one; in six years there were  
 365 2 or 3 episodes, and in just two years (1971 and 2000) we detected over 3  
 366 episodes in one year. The greatest accumulation of cases can be observed in  
 367 1971, when a long-lasting torrential episode exceeded the threshold of 200 mm  
 368 in 24 h during four consecutive days. This is one of the most noteworthy episodes  
 369 recorded in Catalonia (Llasat, 1990; Martin-Vide and Llasat, 2000) in the last few  
 370 decades. It started on September 19th in southern Catalonia and ended on  
 371 September 22nd in the northeast of the study area (Llasat *et al.*, 2007). During



372 the last decade, there has been no more than one episode in one single year.  
373 However, for torrential events ( $\geq 100$  mm in 24 h) in Catalonia, Lopez-Bustins *et*  
374 *al.* (2016) detected a 45% increase in cases between the 1950-1981 and 1982-  
375 2013 subperiods. In accordance with this rise in torrential rainfall events, many  
376 studies on Iberian precipitation are showing an increase in rainfall of  
377 Mediterranean origin in eastern Spain (Miró *et al.*, 2009; Lopez-Bustins *et al.*,  
378 2008; De Luis *et al.*, 2010); this contributes to an increase in rainfall variability  
379 over the Western Mediterranean (Hartmann *et al.*, 2013). On the other hand,  
380 extreme torrential rainfall (i.e.  $\geq 200$  mm in 24 h) in Catalonia does not show any  
381 statistical increase along the study period (Figure 6). This is in line with Llasat *et*  
382 *al.* (2016), who found non-statistical temporal trends in extreme daily precipitation  
383 in Catalonia.

384



385

386 Figure 6. Temporal evolution of the annual frequency of extreme torrential events  
387 ( $\geq 200$  mm in 24 h) throughout the 1951-2016 period. The figure shows the linear  
388 regression (black line) and 5-year running mean (dashed line).

#### 389 4.2. Calendars of the daily WeMOi values

390 The lowest WeMOi values are detected in autumn, especially in October (-0.38)  
391 (Figure 7a), when the humid easterly flows from the Mediterranean Sea are  
392 usually expected. This explains why autumn and October are the wettest season



393 and month, respectively, on most of Spain's eastern façade (De Luis *et al.*, 2010).  
394 The greatest accumulation of extreme torrential events in Catalonia is in October,  
395 with 19 events (38% of all cases). September also shows a remarkable  
396 accumulation of events (11 cases), displaying the second lowest WeMOi monthly  
397 value (-0.29). Positive WeMOi values are observed from December to March,  
398 with very few events occurring. Although negative WeMOi values are detected  
399 from April to November, very few episodes are registered in late spring and  
400 summer; the predominance of atmospheric stability during the warm season  
401 reduces the chances of extreme torrential events occurring over the study area.  
402 At the fortnightly timescale, we detected the minimum WeMOi value (-0.39)  
403 during the second half of October (Figure 7b). The greatest accumulation of  
404 episodes, however, is in the first half of October. The lowest WeMOi values are  
405 found from September 16th to October 31st. This short period of the year (46  
406 days) accumulates over a half of the total amount of extreme torrential events (28  
407 cases, 56%). The most positive WeMOi values are detected in the winter months,  
408 particularly from January 1st to February 15th, and only 2 episodes are  
409 registered.

410 At the 10-day timescale, we observed the WeMOi minimum value (-0.45) from  
411 October 11<sup>th</sup> to 20<sup>th</sup> (Figure 7c). This 10-day period also presents the largest  
412 accumulation of extreme torrential events in Catalonia (8 cases; 16% of the total  
413 number of cases). At least 4 cases are registered in each 10-day period from  
414 September 11th to November 10th. This period of the year (61 days) accumulates  
415 two thirds (33 cases, 66%) of all extreme torrential events. WeMOi values are  
416 lower than -0.20 from August 1st to November 10th, fitting well with the period of  
417 highest frequency of extreme torrential events in Catalonia. From August 1st to  
418 September 10th, only 2 cases are registered due to the above-mentioned  
419 atmospheric conditions in summer. From September 11th to November 10th,  
420 favourable conditions can arise for the occurrence of extreme torrential events in  
421 Catalonia: a higher SST in the Western Mediterranean Sea and the early cut-off  
422 of subpolar lows travelling to Mediterranean latitudes (Estrela *et al.*, 2008; Lopez-  
423 Bustins, *et al.*, 2016; Pérez-Zanón *et al.*, 2018). The positive WeMOi values are  
424 observed from December to March and each 10-day period presents either no  
425 episode or only a single one. The most positive WeMOi value is observed from





426 January 1st to 10<sup>th</sup> (+0.38); this indicates the total predominance of the positive  
427 phase of the teleconnection during these days, according to the study period  
428 (1951-2016) (Figure 8a). During this 10-day period, the occurrence of extreme  
429 torrential events in eastern Iberia is highly inhibited by the NW atmospheric  
430 circulation over the study area, and the Genoa low is well represented. The  
431 remaining 10-day periods in winter also present a predominance of the western  
432 circulation over the Iberia Peninsula. This pattern causes positive pressure  
433 differences between the Gulf of Cadiz (at lower latitude) and the North of Italy (at  
434 higher latitude), which means positive WeMOi values and the inhibition of  
435 precipitation in eastern Iberia because of its location in the lee of the westerlies.  
436 On the other hand, the mean sea level pressure (SLP) map from October 11th to  
437 20th shows a predominance of the negative WeMO phase, with humid easterly  
438 flows over Iberia, low pressure usually located in the Western Mediterranean  
439 basin, and a blocking anticyclone over Central and Eastern Europe (Figure 8b).

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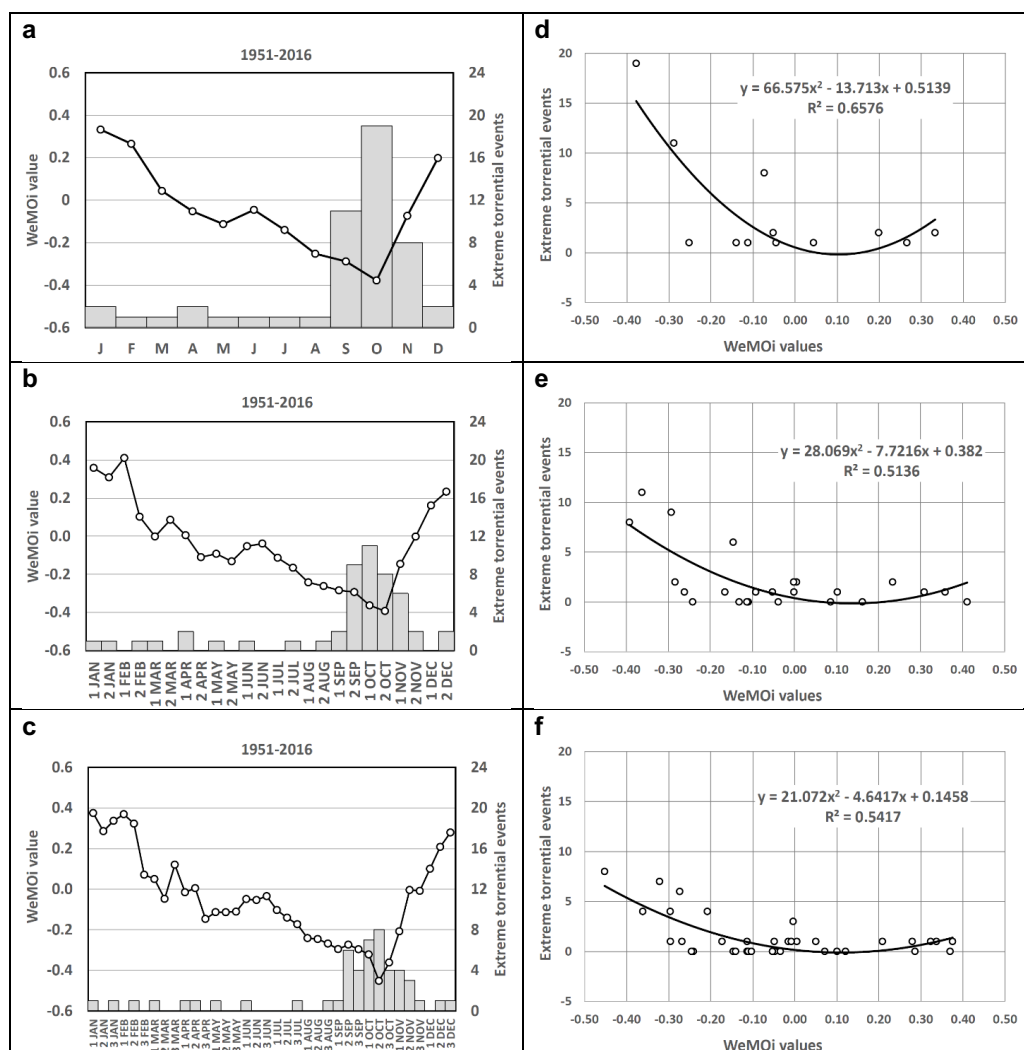
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451 Figure 7. WeMOi calendars and frequency of the extreme torrential episodes at  
 452 several timescales: monthly (a), fortnightly (b) and 10-day (c). Scatter plot of the  
 453 quadratic function of the relationship between extreme torrential events and  
 454 WeMOi values at several timescales: monthly (d), fortnightly (e) and 10-day (f).

455 The fitness of the second-order polynomial is statistically significant at all  
 456 timescales, obtaining an  $R^2$  of 0.66 (Figure 7d), 0.51 (Figure 7e) and 0.54 (Figure  
 457 7f), respectively. Fitness is especially significant at monthly resolution. There is  
 458 an increase in the occurrence of events as the WeMOi value decreases.

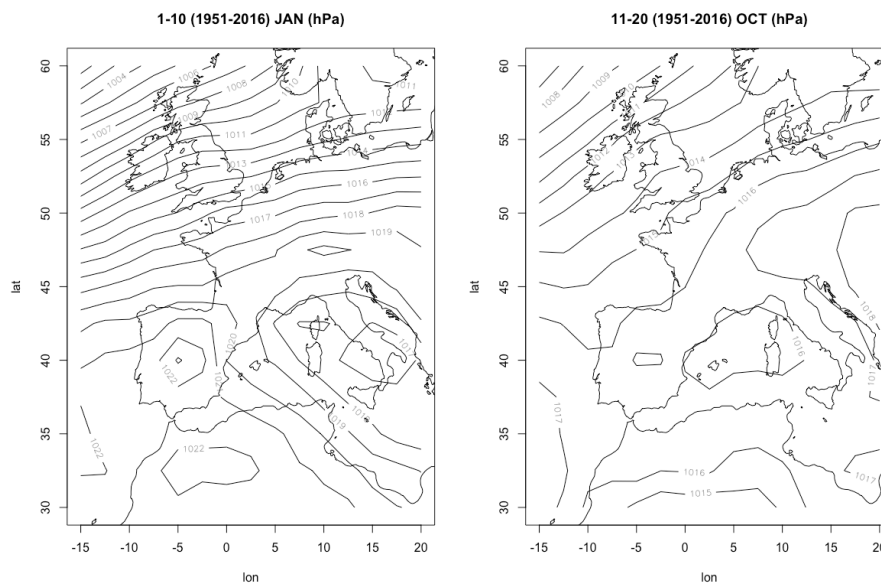


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(a)

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(b)



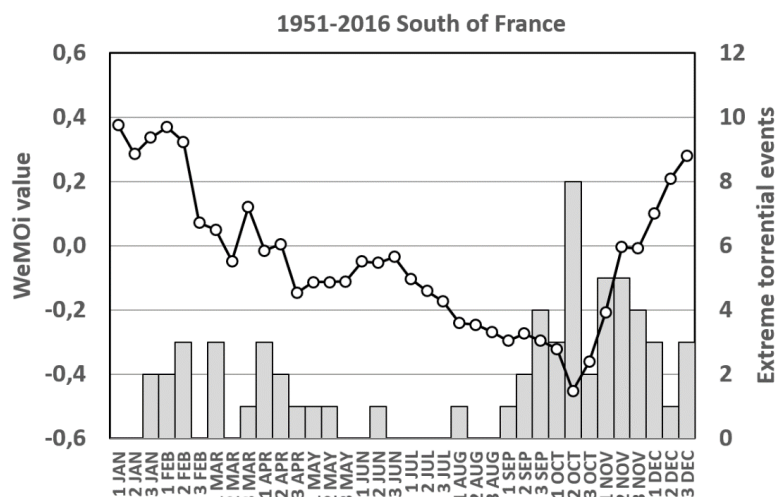
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462 Figure 8. Sea level pressure (SLP) mean of the synoptic window 30°N-60°N and  
463 15°W-20°E from January 1st to 10th (a) and from October 11th to 20th (b) during  
464 the 1951-2016 study period. Data source: NCEP Reanalysis data provided by the  
465 NOAA/OAR/ESRL PSD, Boulder, Colorado, USA.

466 The WeMO teleconnection pattern can exert its influence upon precipitation  
467 variability in other regions of Southern Europe (Caloiero *et al.*, 2011; Milosevic *et al.*,  
468 2016; Mathbout *et al.*, 2019). This occurs in the south of France, where four  
469 weather stations located in the southernmost region of continental France,  
470 Pyrénées-Orientales, are analysed for the 1951-2016 study period (Figure 1).  
471 They recorded 62 torrential events (i.e.  $\geq 100$  mm in 24 h) during these 66 years,  
472 almost one case per year provided by these four weather stations. Figure 9 shows  
473 a remarkable column of episodes on the calendar, which corresponds to the  
474 period from October 11th to 20th; 8 cases are accumulated in the south of France  
475 in this 10-day period, i.e. 13% of all cases. This central period of October is  
476 confirmed to be the most prone to torrential events over many regions of the  
477 western Mediterranean due to presenting the lowest WeMOi values of the year.  
478 On the Iberian Peninsula, the Almanzora river (SE Spain) suffered 2 of the 4 most



479 catastrophic floods in the last 450 years within this central interval in October  
480 (Sánchez-García *et al.*, 2019). Moreover, the deadliest torrential episodes in the  
481 Valencia Region (E Spain) occurred on October 13-14th 1957 and October 19-  
482 20th 1982 (Olcina *et al.*, 2016; Miró *et al.*, 2017).



483

484 Figure 9. The WeMOi calendar and frequency of torrential episodes in the south  
485 of France at the 10-day scale during the 1951-2016 period.

#### 486 4.3. Subperiods and differences in the calendars

487 In relation to the calendars, and according to subperiods, we observed an overall  
488 decrease in WeMOi values throughout the year (Figure 10). On the other hand,  
489 no change was observed in the frequency of episodes between both subperiods;  
490 exactly 25 extreme torrential events occurred in each subperiod. At the monthly  
491 timescale, the extreme torrential period takes place in September and the first  
492 half of October (1951-1983). For the second half (1984-2016), the maximum  
493 accumulation of cases shifts from September-October to October-November,  
494 with the highest concentration of cases in October, whilst new cases occur during  
495 the very late autumn (December). All WeMOi values are statistically and  
496 significantly lower during the second subperiod than during the first one in all  
497 months, especially from October to December. In the summer months, the  
498 decrease in WeMOi values is moderate, albeit statistically significant due to the  
499 low variability of the WeMOi values during the warm months.



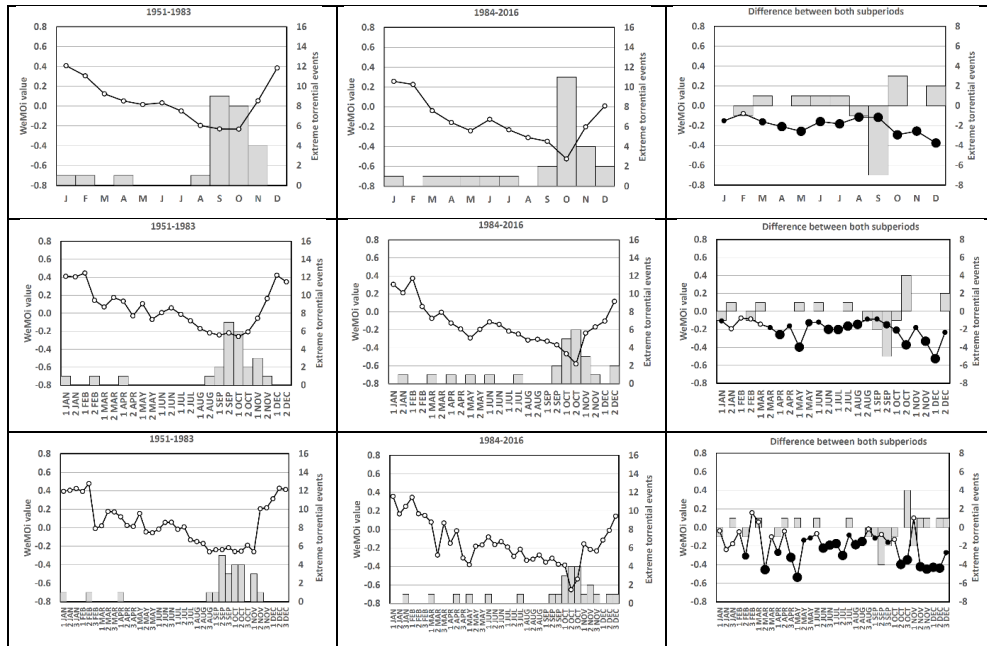
500 At the fortnightly timescale, a shifting of maximum torrentiality is observed from  
501 September 16th–October 15th to October 1st–October 31st. The lowest WeMOi  
502 value of the calendar from 1951 to 1983 was in the first fortnight of October (-  
503 0.26); however, this value is observed in the second fortnight of October during  
504 the 1984–2016 period (-0.58). All WeMOi values according to fortnights displayed  
505 a statistical and significant decrease during the second period, except from  
506 January 16th to March 15th. The sharpest decline in WeMOi values is in the first  
507 fortnight of May, the second fortnight of October, the second fortnight of  
508 November and the first fortnight of December. The lowest WeMOi value during  
509 the second subperiod is detected in the second fortnight of October, when the  
510 greatest increase in extreme torrential events is observed.

511 At the 10-day timescale the lowest WeMOi values remain relatively constant from  
512 the end of August to the beginning of November during the first subperiod, which  
513 corresponds well with the occurrence of extreme torrential events. During the  
514 second subperiod, the lowest WeMOi values are found from October 11th to 31st,  
515 with an accumulation of 8 cases (32% of the total number of cases of the second  
516 subperiod). A continuous and statistically significant decrease in WeMOi values  
517 (at the 99.9% confidence level) is observed from October 16th to December 20th  
518 during the second subperiod, except for the first 10-day period of November. The  
519 increase in torrential events is especially concentrated from October 21st to 31st.  
520 From August 21st to October 10th there is an overall decline in extreme torrential  
521 events, which might be associated with the fact that the WeMOi values hardly  
522 decrease over these 10-day periods of the year during the second subperiod.

523 In general terms, no more cases of extreme torrential events are observed during  
524 the 1984–2016 period in comparison with the 1951–1983 period. Nonetheless, a  
525 greater accumulation of cases can be observed during late autumn and a lesser  
526 accumulation in early autumn during the second subperiod, in comparison with  
527 the first one. A sharp and continuous drop in WeMOi values is observed at the  
528 very end of autumn, which might indicate a shift in the seasonality of the extreme  
529 torrential period from September–October to October–November. This seasonal  
530 displacement might be caused by a recent increase in SST in the Western  
531 Mediterranean basin, particularly in November (Lopez-Bustins, 2007; Estrela *et al.*,  
532 2008; Lopez-Bustins *et al.*, 2016; Arbiol-Roca *et al.*, 2017).



533



534 Figure 10. WeMOi calendars and frequency of the extreme torrential episodes in  
 535 Catalonia at several timescales: monthly (above), fortnightly (middle) and 10-day  
 536 (below) for the 1951-1983 (left) and 1984-2016 (central) subperiods. The right-  
 537 hand column shows the difference in the number of episodes and WeMOi values  
 538 between both subperiods (for WeMOi values: white dots indicate not statistically  
 539 significant differences, and small-, medium- and large-sized black dots show  
 540 statistically significant differences at the 95.0%, 99.0% and 99.9% confidence  
 541 levels, respectively).

542 **4. Conclusions**

543 The present research confirms the usefulness of the WeMOi at daily resolution  
 544 as an effective tool for analysing the occurrence of episodes of torrential rainfall  
 545 over eastern Spain and the south of France. October is the rainiest month in most  
 546 regions of the Northwestern Mediterranean basin and can account for the lowest  
 547 values of the year on the WeMOi monthly calendar. Moreover, most torrential  
 548 episodes take place during a very short period in the middle of this month.

549 Catalonia is located in the Northwestern Mediterranean basin and its extreme  
 550 precipitation is highly dependent upon the atmospheric circulation over the



551 Mediterranean. The present study considers the threshold of 200 mm in 24 h for  
552 extreme torrential episodes, due to the fact that this rainfall accumulation in one  
553 day can cause serious widespread damage. Following thorough review of several  
554 databases, and contrasting these results with the original files and nearby  
555 weather stations, we confirmed that Catalonia registered 0.8 cases per year of  
556 extreme torrential episodes during the 1951-2016 period in accordance with the  
557 7-7 UTC pluviometric day.

558 The 10-day period from October 11th to 20th exhibits both the highest  
559 accumulation of extreme torrential episodes in Catalonia and the lowest intra-  
560 annual WeMOi value. This 10-day period has been demonstrated to be the most  
561 prone to torrential events in the Northwestern Mediterranean area according to  
562 the WeMOi values; and this is also confirmed by the accumulation of episodes in  
563 the same 10-day period in the south of France. The most torrential event in  
564 Catalonia ever recorded by an official weather station is in Cape Creus (the  
565 easternmost part of the Iberian Peninsula) within the period most susceptible to  
566 torrential rainfall (October 13th 1986) with a total amount of 430 mm. The most  
567 positive WeMO phase of the year usually takes place in January, especially from  
568 January 1st to 10th, when the synoptic conditions of this time of the year inhibit  
569 torrential events.

570 No extreme torrential episodes in Catalonia occurred in a positive WeMOi phase.  
571 Additionally, 60% of the cases occurred in an extreme negative WeMO phase,  
572 i.e. a WeMOi value equal to or lower than -2.00. In the present study this threshold  
573 is considered to constitute the onset of a rainstorm favoured by a strong  
574 Mediterranean flow. The lower WeMOi value is related to an increase in extreme  
575 torrential events at all timescales.

576 On comparing both study subperiods (1951-1983 and 1984-2016), an overall  
577 statistically significant decrease is detected in most WeMOi values of the year,  
578 especially at the end of October and some periods in November and December.  
579 On the other hand, extreme torrential events show no changes in frequency  
580 between both subperiods; no temporal trend is observed, either, over the 1951-  
581 2016 study period. The most notable change involves the displacement of  
582 extreme torrential episodes from early autumn to late autumn; this is in  
583 accordance with the lower WeMOi values detected in the last three months of the



584 year during the second subperiod. Further research on this theme is required and  
585 SST temporal trends might provide a better understanding of these changes in  
586 extreme torrential events and WeMOi calendars.

#### 587 **Data availability**

588 WeMOi data can be downloaded from the Climatology Group website  
589 <http://www.ub.edu/gc/en/>

#### 590 **Author contributions**

591 JALB carried out the analysis and wrote the paper. LAR updated WeMOi data  
592 and plotted pressure maps. JMV discussed results. ABE elaborated the inventory  
593 of the episodes and discussed results. MPD discussed results.

#### 594 **Competing interests**

595 The authors declare that they have no conflict of interest.

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600 Industry, and Competitiveness. Our research benefited from the daily  
601 precipitation data provided by the Meteorological Service of Catalonia and by  
602 Météo-France.

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