

1 **Spatial database and website for reservoir triggered seismicity in Brazil**

2
3 **Eveline Sayão¹,**
4 evelinesayao@unb.br

5 **George França¹**
6 georgesand@unb.br

7 **Maristela Holanda²**
8 mholanda@unb.br

9 **Alexandro Gonçalves²**
10 alexandror2@yahoo.com.br

11 ¹ Seismological Observatory - Institute of Geosciences – Universidade de Brasília

12 Campus Darcy Ribeiro –SG13 – Zip Cod -70910-900 –Brasília, Brasil

13 ² Department of Computer Science – Universidade de Brasília

14 Campus Darcy Ribeiro - SGAN – Zip Cod -70910-900 –Brasília, Brasil

15 **Abstract**

16 After confirming that impoundment of large reservoirs could cause earthquakes, studies on
17 reservoir-triggered seismicity (RTS) have had a considerable scientific incentive. Most of the
18 studies determined that the vertical load increase due to reservoir load, and the reduction of
19 effective force due to the increase in pore pressure, can modify the stress field in the reservoir
20 region, possibly triggering earthquakes. In addition, the RTS is conditioned by several factors
21 such as pre-existing tectonic stresses, reservoir height /weight, area-specific geological and
22 hydromechanical conditions, constructive interaction between the orientation of seismotectonic
23 forces, and additional load caused by the reservoir. One of the major challenges in studying RTS
24 is to identify and correlate the factors in the area of influence of the reservoir, capable of
25 influencing the RTS process itself. A spatial seismicity-triggered reservoir database was created
26 to facilitate the research in this field, based on the specifications of the national spatial data
27 infrastructure (INDE), and to assemble data pertinent to the RTS study in the area of reservoirs.

28 In this context, this work presents the procedures and results found in the data processing of
29 seismotectonic factors (dam height, reservoir capacity, lithology and seismicity) and compared
30 first to the dams that triggered earthquakes and secondly to the Brazilian dam list. The list has
31 been updated with 4 more cases, making a total of 30 cases. The results indicate that the
32 occurrence of RTS increases significantly with dam height since dams less than 50 m high cause
33 only 2% of earthquakes while those higher than 100 m cause about 54%. The reservoir volume
34 also plays a role and it was estimated that RTS occurrence requires a limiting minimum value of
35 $1 \times 10^{-4} \text{ km}^3$. There was no clear correlation between the geology and geological provinces with
36 RTS. The delayed response time of the reservoirs represents 43% of the total, that is, almost half
37 of them have a hydraulic behavior. The highest magnitude, 4.2, was observed at a reservoir with
38 a volume greater than 10^{-3} km^3 . As a practical outcome, to assist the analysis by the general
39 community, the web viewer RISBRA (Reservoir Induced Seismicity in Brazil) was developed to
40 serve as an interactive platform for BDSDR data.

41 **Introduction**

42 The reservoir-triggered seismicity (RTS) phenomenon was first observed during the filling of
43 Lake Mead at the Hoover Reservoir (United States) in the mid-1930s, and occurrences of RTS in
44 case of the following reservoirs: Hsinfenghiang (China), Kariba (Zambia), Kremasta (Greece),
45 and Koyna (India) in the late 1960s (Marza,1999). Currently, there are more than 150 identified
46 as RTS (Gupta, 2002; Wilson at al., 2017; Foulger at al., 2018) and the worst case may be the
47 major earthquake in May 2008 in Sichuan, China. The 7.9 magnitude earthquake killed about
48 80000 people, broke nearly 300 miles of fault and damaged 2380 dams, including the 156-meter-
49 high Zipingpu Dam (International Rivers, 2009). (Figure1). Filling large reservoirs, mining

50 underground mines, injecting high-pressure fluids into deep wells, removing fluids during oil
51 exploration, and the after-effects of large nuclear explosions can cause earthquakes (Simpson,
52 1986). Among these, we highlight the RTS phenomenon related to geoengineering works that
53 can have major social, economic, environmental, and legal impacts, among others.

54 In Brazil, the first RTS case was a 3.7 magnitude earthquake with intensity V-VI (MMI) recorded
55 at the reservoir of Carmo do Cajuru, MG, in 1971. Approximately 185 RTS cases are known
56 worldwide, of which 30 happened in Brazil (Foulger et al., 2017; Wilson et al., 2017) (Figure 1).
57 There are several studies on reservoirs capable of triggering earthquakes (Assumpção et al.,
58 2002; Ferreira et al., 2008; Veloso and Gomide., 1997), few of them, however, correlate the
59 physical and geological information as possible agents of the triggered earthquakes. Making this
60 correlation expands the ability to understand this phenomenon. Thus, this work presents the
61 procedures and results found in the data processing concerning the following parameters (height,
62 volume, area, geology, and local seismicity level) and comparing them with the dams that
63 triggered earthquakes and the Brazilian dam catalog. Finally, a spatial database model of the
64 reservoirs and their geological and geophysical characteristics was developed.

65 This work is based on the work developed by the Comissão Nacional de Cartografia (CONCAR,
66 2010) and the Technical Specification for the Structuring of Vector Geospatial Data of Defense
67 of the Earth Force - ET-EDGV (Brazil, 2015, 2016). Because these specifications are still being
68 developed, the diagrams of the dam systems are not yet adequately represented. The amount of
69 information and probable effects of RTS causing requires the standardization of information,
70 which was accomplished in according with the National Spatial Data Infrastructure (INDE).

71 The work is based on the OMT-G (Object Modelling Technique for Geographic Applications)

72 model (Davis Jr., 2000; Borges et al., 2001; Borges et al., 2005) also used in these
73 documentations. This model aims to be more faithful to the modeled reality by using a smaller
74 set of graphic objects than would be used in other models for geographic data.

75 **Database and web viewer**

76 The motivation for creating the Seismicity Database Triggered by Reservoir (BDSDR) arose
77 from the research in the cases that occurred in Brazilian Reservoirs when observing the lack of
78 cohesion of information, pertinent to the study, presenting only isolated cases or listing with the
79 locations of occurrences.

80 According to the NeDiMAH WG4 the use of digital collections for research has an impact on the
81 creation, management and long-term sustainability of digital data, and the use of digital
82 resources for the creation and publication of new knowledge is a vital part of the digital life
83 cycle, then we used this group on the basis of our database. (NeDiMAH Working Groups.
84 European Science Foundation, 2020). The purpose of the database is to gather all the available
85 information such as physical, structural, geological and geophysical data on each reservoir, and
86 to store in a standardized way while sharing and making it accessible so that the database can
87 assist in RTS studies.

88 **National Spatial Data Infrastructure (INDE/ NSDI)**

89 The body responsible for developing spatial data structures is the Comissão Nacional de
90 Cartografia (CONCAR) that is linked to the former Ministry of Budget and Management
91 Planning. CONCAR is responsible for elaborating the technical specifications related to the
92 spatial data that make up the Infraestrutura Nacional de Dados Espaciais (INDE), regulated by

93 Decree No. 6,666/2008. According to this decree, INDE is an integrated set of technologies,
94 policies, mechanisms and procedures for coordinating and monitoring standards and agreements,
95 necessary to facilitate the storage, access, sharing, dissemination and use of geospatial data that
96 belong to the federal, state, district and municipal spheres of government (Brazil, 2008).

97 The spatial data infrastructure defines the standards for the data composition and can be
98 presented as a Technical Specification. In 2006, CONCAR set up the Specialized Committee for
99 the Structuring of the Digital National Map (CEMND), which developed the Technical
100 Specifications for the Structuring of the Geospatial Vector Data (ET-EDGV) for application in
101 the National Cartographic System and INDE (CONCAR, 2017).

102 The specifications proposed for the EDGV (CONCAR, 2017) divide the Brazilian geographical
103 space into two groups. The first group consists of the object classes usually produced in the
104 Small-Scale Mapping (MapTopPE), elaborated in the Systematic Mapping of the SCN (scales
105 of 1: 25,000 and smaller). The second group consists of the object classes usually acquired in the
106 topographic mapping of large scales. This work will use only the small-scale topographic model.

107 MapTopPE is divided into 14 categories: Energy and Communications (ENC), Economic
108 Structure (ECO), Hydrography (HID), Boundaries/Limits and Localities (DML), Reference
109 Points (PTO), Relief (REL), Basic Sanitation (SAB), Vegetation (VEG), Transport System
110 (TRA), Transport System/Airport Subsystem (AER), Transport System/Duct Subsystem (DUT),
111 Transport System/Rail Subsystem (FER), Transport System/Hydro Subsystem (HDV), and
112 Transportation System/Road Subsystem (ROD).

113 In conceptual modeling, the object classes are grouped into categories with common functional
114 aspect. Among the categories, the hydrography package covering the dam class is the class of

115 interest for this paper. However, the other classes inserted in the proposed model do not have
116 definitions pre-established by the INDE. According to the INDE Action Plan (CONCAR, 2017),
117 the data or datasets associated with each of these EDGV classes are considered as reference
118 geospatial data in the INDE.

119 The Action Plan for implementing INDE classifies the data into thematic and reference data.
120 Thematic data are sets of data and information on a phenomenon or a theme, such as climate,
121 education, vegetation, industry, among others, in a region or across the country. Whereas,
122 according to CONCAR (2010), the reference data are defined as:

123 "Datasets that provide general information of non-particular use, elaborated as indispensable
124 bases for the geographic referencing information on the surface of the national territory and can
125 be understood as basic inputs for georeferencing and geographical contextualization of all the
126 specific territorial themes".

127 **Designing the Spatial Database**

128 To implement the data in the database management system, three phases are required: conceptual
129 modeling, logical modeling and physical modeling or implementation. This same method is used
130 for modeling spatial databases.

131 **First Phase: Conceptual Modeling**

132 Conceptual modeling is not directly linked to implementation, its main objective is to capture the
133 semantics of the problem and the needs of the study in question (Cardoso and Cardoso, 2012).

134 The OMT-G (Object Modelling Technique for Geographic Applications) data model was used to
135 create the conceptual model of the Reservoir-Triggered Seismicity Database (BDSDR). This
136 model was chosen following the NSDI specification.

137 It was analyzed which entities would compose this database and which attributes each one of
138 them. As well as, check which entities are related and define the cardinality of each relationship.
139 From the studies on the metadata archives of the seismological data, a model consisting of 20
140 entities it was initially defined: Stress Regime, Fault Orientation, Fault Mechanism,
141 Chronostratigraphy, Structure, Lithology, Reservoir, Dam, Federative Unit, Municipality,
142 Hydrometry, Magnetometry, Electromagnetometry, Gravimetry, Pluviometry, Regional Stress
143 Regime, Hydrography, Crustal Thickness, Seismic Event, and Seismographic Station.

144 In this process, it was observed that a reservoir (main entity) is related to hydrometric entities,
145 dams, municipalities, gravimetry, chemical events, crustal thickness, hydrography, the region's
146 stress regime, pluviometry, electromagnetism and magnetometry. So, as it is an entity with the
147 highest number of relationships, place it at the center of the model. Opt for the OMT-G model, to
148 elaborate the conceptual model, as Borges et al. (2005) models like this are better suited the
149 needs of geographical applications, both in the form of presentation and in the way of relating.
150 So, when using the OMT-G model, you can easily identify how conceptual or relational tables
151 are non-geographic data or a type of geographic data that the table represents. When creating
152 tables in the OMT-G model, the user applies the type of geographic data similar to the correct
153 mode or type of display for an entity. So, if an entity is implemented in the database, it will be a
154 table point other than the upper left corner, a star, if it is a polygon or a multipolygon, in its upper
155 left corner, and so on. The types of representation used in this modeling were: point - represented
156 by a star, polygon or multipolygon - represented by a square, line - represented by a line, and a
157 level variation - represented by isolines.

158 Figure 2 presents the conceptual model based on OMT-G, developed in the StarUML 5.0.2.1570
159 software while Table 1 explains each relationship of the OMT-G model.

160 **Second Phase: Logical Modeling**

161 Creating the Reservoir-triggered Seismicity database in a Database Management System
162 (DBMS) required transforming the conceptual model into an implementation model. This
163 transformation consists of converting the OMT-G model into the relational model (MR) that
164 represents the data in the database as a collection of relationships (tables).

165 At this stage, key attributes such as imposing relational integrity, creating unique indexes,
166 attributing data types, and the height of the fields to store information are defined and identified.

167 Finally, the relational model was implemented after completing the process of developing the
168 conceptual model. A creation of the relational model consists of taking as elaborated tables in the
169 conceptual model and inserting as primary key identifications in the key characters and inserting
170 the type of each attribute (integer, real, text, char, varchar, etc.). It can be seen in the figures both
171 the conceptual and the relational model that, like the screens are executed by dotted lines,
172 because the OMT-G model as entities with georeferenced characteristics are applied by dotted
173 lines. And if a simple (non-spatial) relationship is performed, an indication of this relationship is
174 made by continuous lines.

175 The logical model was created using the StarUML 5.0.21570 software.

176 Figure 3 shows the BDSDR relational model that was created from this conversion.

177

178 **Third Phase: Physical Modeling**

179 The last phase of the database design consists of creating a physical schematic, which depends
180 on the used Database Management System (DBMS) (Cardoso and Cardoso, 2012). DBMS is the
181 set of computer programs that can change the logical and physical structure of the database. The
182 degree of freedom of the data is higher than in the older systems (Teorey et al., 2014).

183 Database management systems use database management software (DBMS), for example:
184 Medeiros (2012). For the development of the spatial database, in Linux environment,
185 PostgreSQL 9.3 with raster extension was used, PostGIS 2.4, pgAdim III and Quantum GIS
186 (QGIS) version 3.12.

187 Most database management systems do not support the spatial data implementation natively,
188 requiring the use of spatial extensions. The extension used in the implementation of BDSDR was
189 PostGIS 2.4. The PostgreSQL is an open source object-relational database management system,
190 that allows anyone to study, modify and distribute the software free of charge for any purpose.
191 Object-relational refers to the spatial database system optimized for storing and querying data
192 related to objects in space, including points, lines, and polygons (Elmasri and Navathe, 2011).

193 **Web viewer**

194 A web viewer is an interactive map in an application that allows the user to interact with
195 elements on the map and obtain information on these elements.

196 The web viewer, named RISBRA (Reservoir Induced Seismicity in Brazil), was created using the
197 leaflet, Node.js and Redis libraries. The leaflet is an open source JavaScript library for interactive
198 maps that provides great tools for implementing map applications for browser interaction
199 (Leaflet, 2018). Redis is an open source network application, in-memory data structure store,
200 used as a database, cache and message broker (Redis, 2018). Finally, Node.js is an open source
201 JavaScript interpreter that focuses on migrating client-side JavaScript to the server side (Node.js,
202 2018).

203 We developed a menu, named LAYERS, which contains all the tables of the bank that can be
204 represented in the map. Figure 4 shows the RISBRA interface and the earthquake icon selected.

205 The image shows the table *layers*, where the data that can be accessed by the user at any time

206 (Reservoir, Dam, Crustal Thickness, Seismographic Station, Structure, Seismic Event,
207 Hydrography, Lithology, Fault Orientation, Pluviometry, Stress Regime, Triggered Earthquakes,
208 Chronostratigraphy, and Fault Mechanism). The data are arranged in the interactive map using
209 icons with the conventional symbology of different formats and colors. All elements are
210 georeferenced on the map of Brazil. The zoom tool in the lower right corner of the screen allows
211 the map to be expanded to the street level.

212 **Reservoir-Triggered Seismicity List updated for the database**

213 Data linked to geology and/or geophysics are dispersed, varying from reservoir to reservoir. The
214 Brazilian bibliography of dam studies presents isolated cases and general listing of the cases.
215 Marza et al. (1999) pioneered the creation of the Reservoir-Triggered Seismicity List, which was
216 later updated by Assumpção et al. (2002), França et al. (2010) and Barros et al. (2018). However,
217 a systematic database containing this information has not yet been established.

218 From 1966 to 2018, 626 events were classified as RTS using Geiger's method (data from the
219 seismic bulletin of the IAG-USP and SISBRA-Brazilian bulletin cataloged by SIS-UnB), with
220 seismic recurrence in several dams, the largest being 4.2 recorded in the dams of Porto Colombia
221 and Volta Grande, at the border between the states of Minas Gerais and São Paulo. Figure 5
222 shows a histogram for the 367 events with a magnitude greater than 1, according to the data from
223 the seismic bulletin of the IAG-USP and SISBRA (Brazilian bulletin cataloged by SIS-UnB).
224 This histogram clearly shows the seismic swarms in the Itapebi and Carmo Cajuru dams in 2003,
225 and Lajeado and Nova Ponte in 2006. These swarms were well monitored by local networks. The
226 histogram also shows the increased monitoring and dam construction since 2002 (Oliveira,
227 2018).

228 In this work, the RTS cases are compared using the unified list (Table 2), where the maximum
229 magnitude recorded in each dam is considered from the reviewed list of all Brazilian dams. The
230 objective is to calculate the potential for triggering an earthquake according to dam height,
231 reservoir capacity, lithology and seismicity. Therefore, we use the data available in the National
232 Register of Dams from the Brazilian Committee of Dams which lists a total of 1413 dams with
233 different purposes. We selected a total of 348 reservoirs, at least 20 m high, built for producing
234 electricity (hydroelectric), except for the Açú and Castanhão reservoirs that fight drought and
235 irrigation, respectively. Dams lower than 20 m high were discarded since these dams have low
236 probability of triggering earthquakes, refer to previous works (e.g. Assumpção et al., 2002).

237 Table 2 and Figure 6 present the updated RTS cases, which increased from 17 (Marza et al.,
238 1999) to a total of 30 cases. Table 2 is based on the work of Marza et al. (1999), to which we
239 added other data such as area of reservoirs, type of seismicity, maximum magnitude,
240 predominant geological type of the reservoir (Craton, Fold and Thrust Belt and Basins), location
241 of the event in relation to the reservoir, and the references.

242 **Results and Discussions**

243 The known RTS cases have significant common features, especially during the initial filling
244 phase of the reservoir, when reservoir-triggered earthquakes generally begin to occur. Factors
245 such as dam height, volume, area, local geology, maximum magnitude, and seismicity in the
246 region may interfere with RTS, each one of these factors is addressed below.

247

248 **RTS**

249 In general, from the total of 348 reservoirs, only 8.6% of those presented RTS, and only two

250 events with a maximum magnitude greater than or equal to 4.0 (Table 3 and Figures 7 and 8).
251 Regarding damage, the highest seismic intensity of VI-VII (MMI) or Peak Ground Acceleration
252 (PGA) of 0.08 - 0.25, was estimated in Porto Colombia and Volta Grande while the seismicity
253 type was mostly Initial (Table 2).

254 Geographically, Brazil is divided into five regions; North, Northeast, Southeast, South, and
255 Midwest. From the regional viewpoint, the southeastern region has the highest number of cases,
256 which is directly related to the high number of reservoirs in the region that accounts for 43% of
257 the country's reservoirs. Additionally, the southeast also has a concentration of the largest
258 number of reservoirs higher than 50 m (Table 3 and Figures 7 and 8) and the greatest occurrence
259 of natural earthquakes cataloged in Brazil, thus explaining the highest number of RTS in the
260 Southeastern region. However, compared to the number of RTS, 17.8% of the total number of
261 Reservoirs in the Northeast shows that although there are fewer cases in the region, the relative
262 value is comparatively higher. Surprisingly the North region also has a considerable percentage
263 indicating a potential region for RTS whereas the Midwest region has the lowest percentage.

264

265 **Correlation of RTS with geological characteristics**

266 The hydromechanical properties of the rocks related to the RTS phenomenon were discussed by
267 Snow (1972), Brace (1974), Howells (1974), Bell and Nur (1978) and Do Nascimento (2002).
268 Despite the laboratory test determining these properties, little progress has been made, especially
269 due to the great practical difficulties in mapping the huge number of rocks below and in the
270 vicinity of a reservoir in terms of porosity, permeability, existence of faults, cracks, etc.
271 (Assumpção et al., 2002). It is known that permeability determines the diffusion velocity of the
272 fluid pressure and controls the volume of affected rocks while possibly being one of the most

273 important factors in the change of seismicity level in the vicinity of a reservoir (Do Nascimento,
274 2002). The existence of fractures and faults, besides generating a weakness zone due to the low
275 resistance to rupture, also facilitates liquid penetration all the way to the deepest and most distant
276 reservoir zones, increasing the pressure in the pores. Thus, depending on the orientation of the
277 natural efforts in relation to the fault system, a small effort/stress, even a very small one, of the
278 reservoir may be sufficient to trigger earthquakes (Assumpção et al., 2002).

279 In order to correlate the probability of RTS with the geotectonic characteristics, the local number
280 of reservoir-triggered seismicity cases was compared with the local lithology (types of rocks):
281 igneous, metamorphic and sedimentary, as indicated in Figure 9a, and the geological province as
282 well. Baecher and Keeney (1982) were among the first to propose comparing the number of
283 cases of RTS with local lithology. The results we had with the same correlation show that
284 igneous rocks have a higher percentage of occurrence of RTS (10.1%) than on sedimentary
285 (8.4%) and metamorphic (8.1%) rocks. This is contrary, for example, to what Baecher and
286 Keeney (1982) estimated for deep, very deep or very large reservoirs (that is, height > 100 m or
287 volume > 10 km³): sedimentary rocks are slightly more likely (16%) compared to metamorphic or
288 igneous (about 10% each).

289 Thus, the RTS was also compared to the main geological provinces that are classified by the
290 CPRM (Figures 9b and 10) into three categories: Craton, Basins, and Fold and Thrust Belt. The
291 values were again very close, with the tendency of a higher number of RTS in the region of
292 basins (10.65%).

293 Although the results show a slight tendency toward igneous rocks in the geological context and
294 basins in geological provinces, it is impossible to determine with certainty the trend of these
295 parameters. Therefore, we suggest an in-depth study on the local structural geology of the dams

296 so that the geological influence can be determined more clearly.

297 **Dimensional physical properties and their correlations**

298 Simpson (1986) observed that the higher the dam the greater the probability of triggering an
299 earthquake, and that the most common RTS occurrence is observed in reservoirs with a
300 maximum height greater than or equal to 100 m. The tectonic, geological and hydrogeological
301 environment of the reservoirs is most affected by the increase of the vertical efforts, via its own
302 weight and/or via the increase of water pressure that infiltrates through pores, faults, and
303 fractures.

304 Thus, in Brazil, a comparison between the RTS cases and the dam heights indicates that dams
305 smaller than 50 m are only 2% likely to trigger seismicity while those higher than 100 m are
306 approximately 54% (Figure 11a) more likely to trigger earthquakes, confirming Simpson's
307 (1986) findings.

308 According to the CBDB databank, the volume parameter is available for only 256 reservoirs.
309 Figure 11b shows that 47% of the reservoirs with a volume greater than $1 \times 10^{-2} \text{ km}^3$ triggered
310 earthquakes, and since this percentage decreases linearly with volume, reservoirs with a volume
311 less than $1 \times 10^{-3} \text{ km}^3$ have a low estimated probability of triggering earthquakes. This result
312 demonstrates the influence of volume (pressure) that is clearly related to the type of RTS in
313 Brazil, which are mostly of the initial type (Table 2 and Figure 11b).

314 Figure 12 shows the correlation between volume and height for RTS cases. We observe that the
315 height does not have a limit between 20m and 209m, which is the height of the largest dam.
316 However, regarding volume, we estimate a minimum value of $1 \times 10^{-4} \text{ km}^3$ for generating a RTS,
317 which is represented by a black bar in Figure 12.

318 **Response Time**

319 Seasonal variations in the water level of the reservoir can trigger earthquakes. Simpson (1986)
320 and Talwani (1995) divided the seismic response of a reservoir into two categories, depending on
321 the spatial and temporal pattern of RTS: (i) initial seismicity and (ii) steady state/ initial or
322 delayed response seismicity.

323 The initial seismicity occurs with the initial damming/impounding of the water or large
324 oscillation of the water level in the lake, which is observed more frequently. Cases of steady state
325 or delayed response seismicity occur at a certain time after the filling/impoundment when the
326 steady-state is reached and presents a more lasting associated seismicity. These different
327 responses may correspond to two fundamental mechanisms by which a reservoir can modify the
328 force in the crust - one related to the rapid increase of elastic stress due to the reservoir load
329 (mechanical behavior) and the other to the more gradual diffusion of water from the reservoir to
330 hypocentral depths (hydraulic behavior). The force may decrease as a result of changes in the
331 elastic stress (decrease of normal stress or increase in shear stress) or reduction of effective
332 normal stress due to increased pore pressure. The pore pressure at hypocentral depths can
333 increase rapidly, from a coupled elastic response due to the pore compaction, or more slowly,
334 with the diffusion of surface water.

335 Of the 30 RTS cases, only 4 were considered as a delayed response while 17 cases had only an
336 initial response (Figure 13). These different responses may correspond to two fundamental
337 mechanisms by which a reservoir can modify the force in the crust - one related to the rapid
338 increase of elastic stress due to the reservoir load (mechanical behavior) and the other to the
339 more gradual diffusion of water from the reservoir to hypocentral depths (hydraulic behavior).

340 Figure 14 shows reservoir height, volume, and area versus the delay time. The dispersion of the
341 results indicates that correlating any of these parameters with time delay is impossible.

342 **Highest Magnitude**

343 It is known that in large reservoirs, the chance of pressure in the rock pores affecting the existing
344 seismic structures in the area below the reservoir increase; however, there are cases in the
345 literature of small reservoirs triggering earthquakes that released stresses with magnitudes far
346 exceeding the sum of all additional stresses resulting from the lake. As an example, in 1974 in
347 Brazil, the largest RTS event (4.2 mb magnitude) occurred near the Porto Colombia and Volta
348 Grande reservoirs, with heights of 40 and 55 m high and areas of 19.5 and 143 km² respectively
349 (number 24 in Table 2). Furthermore, small reservoirs such as Açu and Carmo Cajuru with dams
350 only 31 and 23 m high, have triggered earthquakes with magnitudes higher than 3.0 (Veloso and
351 Gomide, 1997; Ferreira et al., 1995).

352 Based on Klose (2013), the reservoir volume showed a small tendency to generate higher
353 magnitude events compatible with the affected area of the reservoir, depending on its
354 dimensions. Figure 15 shows that most of the events occur in reservoirs with volumes greater
355 than 10⁻³ and with a magnitude of 4.2 in most cases, events between 3 and 4 magnitudes occur in
356 dams lower than 100 m.

357 **The intensity and highest magnitude**

358 Several events were not felt, or there was no micro-seismic survey to define its intensity, for
359 these we consider Intensity I., these are designated as Intensity I. Figure 16 shows a linear
360 correlation between magnitude and intensity, disregarding the Intensity I data. Thus, a linear least
361 squares adjustment was performed and resulted in the equation below:

362 $I = 1.147M + 1.016$ (0.35 standard deviation)

363 The correlation coefficient of 0.66 reflects the small number of data available. It is characteristic
364 of the Intraplate Intensity that the value estimated for intensity is greater than that estimated for
365 magnitude.

366 **Conclusions**

367 The complete compilation of reservoir-triggered seismicity occurrences, including
368 spatial/temporal behavior, allow a better evaluation of the seismic risk of future reservoirs. Thus,
369 the database allows the systematic presentation in one location of all the pertinent data regarding
370 RTS cases in Brazil, including all the known parameters influence the RTS process.

371 The created web viewer, RISBRA, presents an interactive platform with easy access and great
372 potential to improve knowledge on the RTS in Brazil.

373 The histogram of the RTS cases reflects seismic swarms, greater monitoring and the construction
374 of dams since 2002. We highlight that, since the establishment of the Brazilian Seismographic
375 Network (RSBR), in 2011, the acquisition of seismic monitoring data has improved. The RSBR
376 is the joint work of four different institutions: Universities of São Paulo (USP), Brasília (UnB),
377 Rio Grande do Norte (UFRN) and National Observatory (ON). The network consists of more
378 than 90 stations (in January 2020) operated by these four institutions (Bianchi et al., 2018).

379 From the regional viewpoint, the considerable percentage of RTS in the Northern region
380 indicates a potential RTS region, considering the exploratory growth. Despite having a small
381 number of RTS, 5 cases, the Northeast region has a comparatively higher relative value of RTS
382 compared to other regions.

383 Although the results show a trend with higher numbers of RTS in igneous rocks (rock type) and
384 sedimentary basins (geological provinces) being more prone to RTS, such trends cannot be
385 backed up with the currently available data. Therefore, we suggest an in-depth analysis of the
386 structural geology at the dam sites in order to understand and identify in more detail the
387 geological influence.

388 The dam height has been confirmed as one of the main indicators of the dam capability of
389 triggering earthquakes. Dams less than 50 m high are only 2% likely to cause seismicity while
390 those more than 100 m high are about 54% more likely to cause an earthquake.

391 The reservoir volume also strongly influences its capability for causing an earthquake and we
392 estimate the limiting minimum value of $1 \times 10^{-4} \text{ km}^3$ for the occurrence of RTS.

393 The delayed response of the reservoirs represents 43% in total, indicating hydraulic behavior for
394 almost half of the reservoirs. For higher magnitudes (4.2, the highest recorded), we found that
395 most events occur in reservoirs with volumes larger than 10^{-3} km^3 .

396 An equation " $I = 1.147M + 1.016 (+ -0.35)$ " has been determined to describe the relationship
397 between intensity and highest magnitude. Where "I" is the estimated intensity and "M" is the
398 determined magnitude.

399 Practical difficulty of mapping soil layers below the dams hinders the evaluation of the seismic
400 risk of a reservoir and, therefore, it is essential to obtain key parameters such as local stresses,
401 rock mass permeability, and fracture system geometry. Thus, studies of previous cases are useful
402 when trying to assess the seismic risk posed by future reservoirs. Most importantly, this work
403 shows that the possibility of RTS occurrence in Brazil cannot be neglected and highlights the

404 importance of continuous monitoring, before, during and after the construction of a dam.

405 **Data and Resources**

406 The data used in this article was extracted from the seismic bulletin and SISBRA. Data and
407 information from the SISBRA can be downloaded from the Seismological Observatory of the
408 University of Brasília (SIS / UnB), Center of Seismology of the University of São Paulo (USP):
409 www.obsis.unb.br; www.sismo.iag.usp.br; (last accessed December 2018). Information on the
410 dams was taken from the Brazilian Committee of Dams (CBDB) <http://www.cbdb.org.br/> (last
411 accessed in October 2018).

412 **Acknowledgments**

413 The authors thank the Comitê Brasileiro de Barragens for providing the data from the Cadastro
414 Nacional de Barragens.

415 **References**

416 Assumpção, M., Freire, M. and Ribotta L. C.: Sismicidade Induzida no reservatório de Capivara:
417 resultados preliminares sobre localização de fraturas ativas, IV International Congress of the
418 Brazilian Geophysical Society, Rio de Janeiro, Brasil, 20-24 August 1995, 961-964, 1995.

419 Assumpção, M., Marza V. I., Barros L. V., Chimpliganond C. N., Soares J. E., Carvalho J. M.,
420 Caixeta D. F., Amorim A. and Cabral E.: Reservoir induced seismicity in Brazil, Pure Appl.
421 Geophys, 159, 597, <https://doi.org/10.1007/PL00001266>, 2002.

422 Barros, L. V. and Caixeta D. F.: Induced seismicity at Miranda Reservoir—A fine example of
423 immediate seismic response, 8th International Congress of the Brazilian Geophysical Society,

424 Rio de Janeiro, 14-18 September 200, Brazil, 5, 2003.

425 Barros, L. V., Caixeta D. F., Chimpliganond C. N. and Fontenele D. P.: Evolution of the
426 Areado/MG seismic sequence—Started in January, 2004, International Congress of the Brazilian
427 Geophysical Society, Salvador, Bahia, Brazil, 11-14 September 2005, 6, 2005.

428 Barros, L. V., Carvalho J. M., Ferreira V. M., Albuquerque D. F., Von Huelsen M. G., Caixeta D.
429 and Fontenele D. P.: Determination of source seismic parameters of micro-earthquakes with
430 epicenter in the south of Minas Gerais State-Brazil, 6th International Congress of the Brazilian
431 Geophysical Society, Porto Alegre, Brazil, 14 – 17 October 2014, 2014.

432 Barros, L. V., Assumpção, M., Ribotta, L. C., Ferreira, V. M., Carvalho, M. J., Bowen, M. D. B.
433 and Albuquerque, F. D.: Reservoir – Triggered Seismicity in Brazil: Statistical Characteristics
434 in a Midplate Environment, Bulletin of the Seismological Society of America, 20, 4-6,
435 <https://doi.org/10.1785/0120170364>, 2018.

436 Bell, M. L. and Nur, A.: Strength Changes Due to Reservoir- Induced Pore Pressure and Stresses
437 and Application to Lake Oroville, Journal of Geophysical Research, California, 83, 4469-
438 4483, <https://doi.org/10.1029/JB083iB09p04469>, 1978.

439 Berrocal, J., Assumpção, M., Antezana R., Dias Neto C., Ortega R., França H. and Veloso J. A.:
440 Sismicidade do Brasil, Instituto Astronômico e Geofísico, Universidade de São Paulo e
441 Comissão Nacional de Energia Nuclear, São Paulo, Brazil, Esperança, 320, 1984.

442 Berrocal, J. and Fernandes, C.: Estudo de Sismicidade Induzida na Área dos Reservatórios
443 Hidroelétricos da Chesf, Sessão Regular da Academia Brasileira de Ciências: Ciências da Terra e
444 o Meio Ambiente. Anais da Academia Brasileira de Ciências, São Paulo, Brasil, 68, 613–620.
445 1996.

446 Bianchi, M. B., Assumpção, M., Rocha, M. P., Carvalho, J. M., Azevedo, P. A., Fontes, S. L.,
447 Dias, F. L., Ferreira, J. M., Nascimento, A. F., Ferreira, M. V., Costa, I. S. L.: The Brazilian
448 Seismographic Network (RSBR): Improving Seismic Monitoring in Brazil. *Seismological*
449 *Research Letters*, 89, 452–457, <https://doi.org/10.1785/0220170227>, 2018.

450 Borges, K. A. V., Davis JR., C. A. and Laender, A. H. F.: OMT-G: an object-oriented data model
451 for geographic applications, *GeoInformatica*, 5, 221-260, 2001.

452 Borges, K. A. V., Davis JR., C. A. and Laender, A. H. F.: Modelagem Conceitual de Dados
453 Geográficos. In: Casanova, M. A., Câmara, G., Davis Jr., C. A., Vinhas, L. and Queiroz, G. R.:
454 Banco de Dados Geográficos. Curitiba, Editora MundoGeo, 2005. Available at:
455 <http://www.dpi.inpe.br/livros/bdados/cap3.pdf>, last access: 01 March 2017.

456 Brace, W. F.: Experimental Studies of Seismic Behavior of Rocks Under Crustal Conditions.
457 *Engineering Geology*, 8, 109-127, [https://doi.org/10.1016/0013-7952\(74\)90018-0](https://doi.org/10.1016/0013-7952(74)90018-0), 1974.

458 BRASIL. Decreto nº 6.666, de 27 de novembro de 2008. Institui, no Âmbito do Poder Executivo
459 Federal, a Infraestrutura Nacional de Dados Espaciais – INDE, e de outras providências.
460 Available at: http://planalto.gov.br/ccivil_03/_Ato2007-2010/2008/Decreto/D6666.htm, last
461 access: 01 March 2017.

462 BRASIL. Portaria nº 011 - DCT, de 22 de abril de 2015. Aprova a Norma da Especificação
463 Técnica para Estruturação de Dados Geoespaciais Vetoriais de Defesa da Força Terrestre (EB80-
464 N-72.002) – 1ª Parte – 1ª Edição – 2015. Available at:
465 http://www.geoportal.eb.mil.br/imagens/PDF/EDGV_Defesa-Forca_Terrestre_2015.pdf, last
466 access: 10 May 2017.

467 BRASIL. Portaria nº 007 - DCT, de 10 de fevereiro de 2016. Aprova a Norma da Especificação

468 Técnica para Estruturação de Dados Geoespaciais Vetoriais de Defesa da Força Terrestre (EB80-
469 N-72.002) – 1ª Parte – 2ª Edição – 2016. Available at:
470 http://www.geoportal.eb.mil.br/images/PDF/EDGV_DEFESA_F_Ter_2a_Edicao_2016_Aprova
471 [da_Publicada_BE_7_16.pdf](http://www.geoportal.eb.mil.br/images/PDF/EDGV_DEFESA_F_Ter_2a_Edicao_2016_Aprova) . last access: 15 June 2017.

472 Cardoso, V. and Cardoso, G.: Sistemas de Banco de Dados: uma abordagem introdutória e
473 aplicada, First Issue, Saraiva Publications, Brazil, 142, 2012.

474 CBDB - Comitê Brasileiro de Barragens, Available at: <http://www.cbdb.org.br/>, last access: 23
475 October 2018.

476 Centro de Sismologia da USP, Available at: <http://www.sismo.iag.usp.br/eq/bulletin/>, last access:
477 15 October 2018.

478 Chimpliganond, C. N.: Characterization of induced seismicity at the Nova Ponte Reservoir/MG,
479 Brasil, M.Sc. Dissertation, University of Brasilia, Brazil, 2002.

480 Chimpliganond, C., França G. S., Bandeira A. E. and Bevilaqua L.: Reservoir-triggered
481 seismicity at the highest Brazilian dam, AGU 2007 - Meeting of Americas Joint Assembly
482 Abstract, Acapulco, México, 22-25 May, 2007.

483 CONCAR – Comissão Nacional de Cartografia. Plano de ação para implantação da infraestrutura
484 nacional de dados espaciais (INDE). Rio de Janeiro, 2010. Available at:
485 <https://www.concar.gov.br/pdf/PlanoDeAcaoINDE.pdf> . last access: 03 April 2018.

486 CONCAR – Comissão Nacional de Cartografia. Especificações técnicas para estruturação da
487 infraestrutura nacional de dados espaciais digitais vetoriais. Edição 3.0, 2017. Available at:
488 https://www.concar.gov.br/temp/365@ET-EDGV_versao_3.0_2018_05_20.pdf . last access: 04
489 April 2018.

490 CPRM – Serviço Geológico Do Brasil. Available at: <http://www.cprm.gov.br/>, last access: 01
491 May 2018.

492 Davis JR., C. A.: Múltiplas Representações em Sistemas de Informação Geográficos, Doctoral
493 thesis, Federal University of Minas Gerais, 115, 2000.

494 Do Nascimento, A. F.: The role of pore pressure diffusion in a reservoir-induced seismicity site
495 in NE Brazil, Doctoral thesis, University of Edimburgo, 203, 2002.

496 Elmasri, R.; Navathe, S B: Fundamentals of database systems, Pearson Education, Inc.,
497 publishing as Addison-Wesley, 6, 2011.

498 Ferreira J., Oliveira, M., Assumpção M., Moreira, J. A. M., Pearce, R. G. and Takeya, M. K.,
499 Correlation of seismicity and water level in the Açú reservoir—an example from Northeast
500 Brazil, Bulletin of the Seismological Society of America, 85, 1483-1489, 1995.

501 Ferreira J., França, G. S., Vilar S., Assumpção M.: Induced seismicity in the Castanhão
502 Reservoir, NE Brazil - Preliminary results, Tectonophysics 456:1, 103-110, 2008.

503 Foulger, G. R., Wilson, M., Gluyas, J., Julian, B. R., and Davies, R.: Global review of human-
504 induced earthquakes, Earth-Science Reviews, 2017.

505 França, G. S., Assumpção M., Ribotta L. C., Von Huelsen M. G. and Chimpliganond E. C. N.,
506 Updated compilation of reservoir triggered seismicity in Brazil, 2010 The Meeting of the
507 Americas (AGU – American Geophysical Union), Foz do Iguaçu, Paraná, Brazil, 2010.

508 Gomide, L. C.: Nature and history of reservoir induced seismicity in Brazil, M.Sc. Dissertation,
509 University of South Carolina, 1999.

510 Gupta, H. K.: A review of recent studies of triggered earthquakes by artificial water reservoirs

511 with special emphasis on earthquakes in Koyna, India, *Earth-Science Reviews* 58, 279 – 310,
512 [https://doi.org/10.1016/S0012-8252\(02\)00063-6](https://doi.org/10.1016/S0012-8252(02)00063-6), 2002.

513 Howells, D. A.: Mechanical properties of rock at the depth of earthquake ignition, *Engineering*
514 *Geology*, 8, 129-134, 1974.

515 The Human - Induced Earthquake Database. Available at: <http://inducedearthquakes.org/>, last
516 access: 10 May 2020.

517 International Rives (2009) A Faultline Runs Through It: Exposing the Hidden Dangers of Dam-
518 induced Earthquakes. Available At:
519 https://www.internationalrivers.org/sites/default/files/attached-files/ris_final_lorez2.pdf, last
520 access: 10 May 2020.

521 Klose, C.D.: Mechanical and statistical evidence of the causality of human-made mass shifts on
522 the Earth's upper crust and the occurrence of earthquakes, *Journal of Seismology*, 17, 109 -135,
523 DOI 10.1007/s10950-012-9321-8, 2013.

524 Leaflet. Leaflet 1.3.4. Available at: <https://leafletjs.com/>, last access: 05 April 2018.

525 Marza, V., Veloso J. A. V., Carvalho J. M., Barros L. V. and Gomide L. C.: Reservoir induced
526 seismicity at Nova Ponte (MG): Revisited, 5th International Congress of the Brazilian
527 Geophysical Society, São Paulo, Brazil, 968–971, 1997.

528 Marza, V., Barros L. V., Soares J. E., Carvalho J. M., Fontenele D., Chimpliganond C., Caixeta
529 D., Gomes I. P., Furtado G. O., Carim A. L., Souza G. F., Caliman E. H., Barros J. B.: Aspectos
530 da Sismicidade Induzida por Reservatórios no Brasil, XXIII Semana Nacional de Grandes
531 Barragens Belo Horizonte – Minas Gerais, 199–211, 1999.

532 Medeiros, A. M. L.: Aplicações geográficas do postgresql e seu módulo postGIS. Revista
533 FOSSGIS Brasil, Coluna Banco de Dados Geográficos, 25-27, 2012.

534 Mendiguren, J. A.: A procedure to resolve areas of different source mechanisms when using the
535 method of composite nodal plane solution, Bulletin Seismological Society of America, 70, 985–
536 998, 1980.

537 Mioto, J. A., Ribotta L. C., Verdiani A. C.: Aspectos geológico estruturais da sismicidade
538 relacionada ao reservatório de Capivara (SP/PR), II International Congress of the Brazilian
539 Society of Geophysics, Salvador, Brazil, 1, 513–520, 1991.

540 NeDiMAH Working Groups (2020). European Science Foundation. Available:
541 [http://archives.esf.org/coordinating-research/research-networking-programmes/humanities-](http://archives.esf.org/coordinating-research/research-networking-programmes/humanities-hum/nedimah/thematic.html)
542 [hum/nedimah/thematic.html](http://archives.esf.org/coordinating-research/research-networking-programmes/humanities-hum/nedimah/thematic.html), last access: 10 June 2020.

543 Node.js. Available at: <https://nodejs.org/en/>, last access: 06 Abril 2018.

544 Oliveira, N. C. C.: A grande aceleração e a construção de barragens hidrelétricas no Brasil. Varia
545 Historia, Belo Horizonte, 34, 65, 315-346, DOI:10.1590/0104- 87752018000200003, 2018.

546 PgAdmin III. pgAdmin III 1.22.2 documentation. Available at:
547 <https://www.pgadmin.org/docs/pgadmin3/1.22/>, last access: 06 April 2018.

548 PostGIS. Available: <http://postgis.net/docs/manual-2.4/>, last access: 07 May 2018.

549 PostgreSQL. Available: <https://www.postgresql.org/docs/9.3/index.html>, last access: 09 April
550 2018.

551 QGIS. Quantum GIS Documentation. Available: <https://docs.qgis.org/2.14/en/docs/>, last access:

552 09 April.

553 Redis. Available: <https://redis.io/>, last access: 09 April 2018.

554 Ribotta, L. C.: Aspectos da sismicidade na área do reservatório de Paraibuna/ Paraitinga, Masters
555 dissertation, USP, São Paulo, 147, 1989.

556 Ribotta, L. C., Miotto J. A., Manuzzi J. L., Carvalho A. M. B. E. and Vinciprova G.: Sismicidade
557 na área do reservatório de Barra Grande, SC/RS, Anais do III Simpósio Brasileiro de Geofísica,
558 Belém, Pará, Brazil, 2008.

559 Ribotta, L. C., Assumpção M., Manuzzi J. L., Carvalho A. M. B. E. and Regina J. V. M.:
560 Seismicity induced in 4 deep reservoirs, southern Brazil, 2010, The Meeting of the Americas
561 (AGU - American Geophysical Union), Foz do Iguaçu, Paraná, Brazil, 2010.

562 Ribotta, L. C., Miotto J. A. and Regina J. V. M.: Sismicidade na área do reservatório de Itá,
563 SC/RS, Anais do II Simpósio Brasileiro de Geofísica da SBGf, Natal, Rio Grande do Norte,
564 Brazil, 2006a.

565 Ribotta, L. C., Miotto J. A. and Regina J. V. M.: Sismicidade na área do reservatório de
566 Machadinho, SC/RS, Anais do XLIII Congresso Brasileiro de Geologia, Aracaju, Sergipe, Brazil,
567 2006b.

568 Ribotta, L. C., Moreira L. D., Souza S. L. E. and Regina J. V.: Reservatório de Itá, SC/RS, 19
569 Anos de Sismicidade, Anais do II Simpósio Brasileiro de Sismologia, João Pessoa, Paraíba,
570 Brazil, 2017.

571 Simpson, D. W.: Triggered Earthquakes, Annual Review of Earth and Planetary Sciences, New
572 York, 14, 21-42, 1986.

573 Snow, D. J., Geodynamics of seismic reservoirs, Proceedings of the International Symposium on
574 Percolation Through Fissured Rocks,1-19, 1972.

575 ShakeMap Scientific Background. Rapid Instrumental Intensity Maps". Earthquake Hazards
576 Program. U. S. Geological Survey. Available <https://earthquake.usgs.gov/>,last access: 06 May
577 2020.

578 Talwani, P.: Two categories of reservoir induced seismicity, Proceedings of the International
579 Symposium on Reservoir-induced Seismicity (ISORI'95), 44-64, 1995.

580 Teorey, T. J., Lightstone, S. and Nadeau, T.: Projeto e modelagem de banco de dados. Tradução
581 Daniel Vieira, Elsevier, 2, 309, 2014.

582 Veloso, J. A. V., Assumpção M., Gonçalves E. S., Reis J. C., Duarte V. M. and Mota C. G.:
583 Registro de SIR em reservatórios da CEMIG e FURNAS, Anais do V Congresso Brasileiro de
584 Geologia de Engenharia, São Paulo, Brazil, 135-146, 1987.

585 Veloso, J. A. V., Carvalho, J. M., Fernandes, E. P., Blum, M. L. B. and Araújo, D. P.: Micro
586 earthquakes and the Balbina Lake, a possible case of induced seismicity in the Amazon are, 2th
587 International Congress of the Brazilian Geophysical Society, Salvador, Brazil, 2, 508-512, 1991.

588 Veloso, J. A. V.: Terremotos induzidos pelo homem, Ciência Hoje, 14, 269-273, 1992a.

589 Veloso, J. A. V.: Cases of RIS in the Brazilian Amazon area, Proceedings Tenth World
590 Conference on Earthquake Engineering, Madrid, Spain, 1, 269-273, 1992b.

591 Veloso, J. A. V. and Gomide. L. C.: Induced seismicity at Cajuru Reservoir, Minas Gerais, Brazil,
592 Proceedings 19th of the International Congress on Large Dams, Florence, Italy, 1211-1225,
593 1997.

594 Viotti, C. B., Gomide L. C. and Brito, S. N. A: Induced seismicity in CEMIG’s reservoir in
595 Minas Gerais - Brazil, Proceedings of the International Symposium on Reservoir-induced
596 Seismicity (ISORIS’95), Beijing, China, 205–212, 1995.

597 Viotti, C. B., Veloso J. A. V. and Gomide, L. C.: Induced seismicity at Cajuru Reservoir, Minas
598 Gerais, Brazil, 19th International Congress on Large Dams Proceedings, Italy, 1211–1225, 1997.

599 Wilson, M. P., Foulger, G. R., Gluyas, J. G., Davies, R. J. and Julian, B. R.: HiQuake: The
600 human-induced earthquake database, Seismological Research Letters, 88, 1560-1565,
601 <https://doi.org/10.1785/0220170112>, 2017.

602 **TABLES**

603 **Table 1.** Explanation of the OMT-G model for the Reservoir-triggered Seismicity Database.

Relationship	Description
Lithology and Structure	The structure is the fault characteristic that is associated with lithology.
Lithology and Chronostratigraphy	Lithology (rock type) has one or more chronostratigraphy data.
Reservoir and Lithology	The reservoir area has one or more types of lithology.
Structure and Stress Regime	The stress regime focuses on the structures
Structure and Fault orientation	Fault orientation refers to diving, direction and inclination information of the structure (fault).
Structure and Fault Mechanism	Failure mechanism refers to information on the characteristics of the structure.
Reservoir and Crustal Thickness	The area of the reservoir has information on Crustal thickness.
Reservoir and Seismic Event	The seismic event may occur in the area of reservoir influence.
Seismic Event and Seismographic Station	Seismic station detects seismic event.
Hydrometry and Reservoir	The reservoirs have daily hydrometric data.

Reservoir and Magnetometry	The reservoir has magnetometry information in its area of influence.
Reservoir and Electromagnetometry	The reservoir has Electromagnetometry information in its area of influence.
Reservoir and Gravimetry	The reservoir has gravimetric information in its area of influence.
Reservoir and Region Stress Regime	The area of reservoir influence has forces acting on the stress regime.
Reservoir and Hydrography	The reservoir is part of the hydrography.
Reservoir and Rainfall	The reservoir area is influenced by rainfall
Reservoir and Dam	The reservoir has a dam.
Municipality and State	Each municipality is located in a state.

604 **Table 2- Seismicity Cases triggered in Brazil.**

N°	Name	Federative unit	Height (m)	Volume (10 ⁻³ km ³)	Maximum water depth in the Reservoir (m)	Area of the reservoir (km ²)	Start of impoundment	Geological Province	Seismicity type	Largest Events							
										Date (YY/MM/DD)	Magnitude	Magnitude Type	Io (MMI)	PGA (g)	ΔT (year)	Location	References
1	Açu	RN	41	2,400	55,0	195.0	1985	Thrust and Folding Range	Delayed	1994/08/26	3.0	mR	IV*	0.03 and below	9.5	Inside	Do Nascimento (2002) and Ferreira et al. (1995)
2	Balbina	AM	31	9,755	51,0	2.36	10/1987	Basin	Initial	1990/03/25	3.4	mb	I	-	2.5	Margin	Assumpção et al. (2002) and Veloso et al. (1991)

							Largest Events								
13	12	11	10	9	8	N°	Initial	Initial *	Craton	Thrust and Folding Range.	Thrust and Folding Range.	Thrust and Folding Range.	Initial	Initial	Date (YY/MM/DD)
Irapé	Furnas	Funil	Emborcação	Castanhão	Capivari- Cachoeira	Name	mR	mI	IV-V	IV-V	IV-V	IV	VI	VI	MMI
MG	MG	MG	GO/MG	CE	PR/SP	Federative unit	mI	mI	IV-V	IV-V	IV	VI	VI	VI	MMI
209	127	50	158	85	60	Height (m)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
5,964	22,950	0,258	17,588	6,700	0,178	Volume (10 ⁻³ km ³)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
470,8	-	808,0	653,0	100,0	-	Maximum water depth in the Reservoir (m)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
137.0	1.44	33.46	473.0	458.00	13.1	Area of the reservoir (km2)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
12/2005	1963	2002	08/1981	2003	07/1970	Start of impoundment	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
Thrust and Folding Range.	Thrust and Folding Range	Craton	Thrust and Folding Range.	Thrust and Folding Range.	Thrust and Folding Range	Geological Province	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
Initial	Initial *	Delayed	Initial	Initial - Delayed	Initial	Seismicity type	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
2006/05/14	1966/11/15	2011/08/14	1982/05/20	2007/08/07	1971/05/21	Date (YY/MM/DD)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
3.0	3.2	3.2	1.6	2.3	3.9	Magnitude	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
mR	mI	mR	ML	mD	ML	Magnitude Type	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
III-IV	IV-V	IV-V	I	I	VI	Io (MMI)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
0.03 -below	0.03 -0.15	0.03 -0.15	-	-	0.08 - 0.15	PGA (g)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
0.01	~1?	8	~1	1?	~1	ΔT (year)	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
Inside	-	Margin	Inside	Margin - Inside	-	Location	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI
França et al.(2010)	Berrocal et al. (1984) and Barros et al. (2005)	Barros et al. (2014)	Viotti et al. (1997,1995)	Ferreira et. al. (2008)	Berrocal et. al.(1984) eand Mioto et.al. (1991)	References	mI	mI	IV-V	IV	IV	VI	VI	VI	MMI

											Largest Events						
N°	Name	Federative unit	Height (m)	Volume (10 ⁻³ km ³)	Maximum water depth in the Reservoir (m)	Area of the reservoir (km2)	Start of impoundment	Geological Province	Seismicity type	Date (YY/MM/DD)	Magnitude	Magnitude Type	I ₀ (MMI)	PGA (g)	ΔT (year)	Location	References
19	Machadinho	RS/SC	126	3.339	-	79.0	2001/09/08	Basin	Initial- Delayed	2001/09/08	1.8	ML	I	-	0.01	Inside -Margin	Ribotta et al.(2006a) e Ribotta et al. (2010)
18	Lajeado	TO	31	5,190	212,3	630.0	2002	Basin	Initial- Delayed	2012/04/01	2.2	mD	I	-	10	Margin	Technical Report of the UnB Seismological Observatory
17	Jirau	RO	62	2,746	90,0	361.6	2014	Basin	Initial	2014/11/07	3.2	mR	IV-V	0.03 -below	0.8	Inside	Barros et.al (2015)
16	Jaguari	SP	77	0,793	-	56.0	12/1969	Thrust and Folding Range	Delayed	1985/12/17	3.0	ML	V-VI	0.03 -0.15	16	Margin	Veloso et al. (1987)
15	Itapebi	BA	120	1,633	-	61.58	12/2002	Craton	Initial	2003/08/03	1.5	M _b	I	-	~0.01	Inside - Margin	Barros (2008)
14	Itá	RS/SC	125	5,100	370,0	141.0	12/1999	Basin	Initial- Delayed	1999/12/15	2.5	ML	III-IV	0.03 -below	0.01	Margin -Inside	Ribotta et al. (2006b,2010,2017)

24		23		22		21		20		N°	
Porto Colombia e Volta Grade		Paraibuna-Paraitinga		Nova Ponte		Miranda		Marimbondo		Name	
MG/SP		SP		MG		MG		MG/SP		Federative unit	
40	55	84/105		142		79		90		Height (m)	
2,3	1,525	1,270	2,636	12,792		1,120		6,150		Volume (10 ⁻³ km ³)	
-		-		-		-		-		Maximum water depth in the Reservoir (m)	
19.5	143.0	47.0	177.0	443.0		70.0		438.0		Area of the reservoir (km2)	
09/1973	04/1973	1976	1974	10/1993		08/1981		1975		Start of impoundment	
		Thrust and Folding Range.		Basin		Basin		Basin		Geological Province	
		Initial		Initial- Delayed		Initial- Delayed		Initial		Seismicity type	
1974/02/24		1977-11-16		1998/05/22		2000-05-06		1978/07/25		Date (YY/MM/DD)	
4.2		3.3		4.0		3.3		2.0		Magnitude	
mD		mb		mR		mR		ML		Magnitude Type	
VI-VII		IV		VI		V-VI		I		Io (MMI)	
0.08 - 0.25		0.03 and below		0.08 - 0.15		0.03 and 0.15		-		PGA (g)	
~1		~1				2.7		~3		ΔT (year)	
Margin?		Inside		Margin		Margin		Margin		Location	
Berrocal et al. (1984), Veloso (1992a) and Gomide (1999)		Mendiguren (1980) eand Ribotta (1989)		Chimpiganond (2002), Marza, Barros, Soares et al. (1999)		Barros e Caixeta (2003) e Assumpção et al. (2002)		Veloso et al. (1987)		References	

				Largest Events																			
20	28	27	26	25	N°	Name	Federative unit	Height (m)	Volume (10 ⁻³ km ³)	Maximum water depth in the Reservoir (m)	Area of the reservoir (km ²)	Start of impoundment	Geological Province	Seismicity type	Date (YY/MM/DD)	Magnitude	Magnitude Type	Io (MMI)	PGA (g)	ΔT (year)	Location	References	
Tucuruí	Três Irmãos	Quebra-Queixo	Sobradinho	Serra da Mesa																			
PA	SP	SC	BA	GO																			
95	82	75	43	154																			
45,500	13,800	0,136	34,116	54,400																			
-	-	549,0	-	-																			
2.43	785.0	5.6	4.12	1.78																			
Set/1984	1990	2002	1977	10/1996																			
Craton	Basin	Basin	Craton	Thrust and Folding Range																			
Initial	Initial	Initial	Initial	Initial																			
1998/03/02	1990/11/01	2003/03/01	1979/07/05	1999/06/13																			
3,6	0.5	0.1	1.9	2.2																			
-	mD	mD	ML	mD																			
IV-V	I	I	I	I																			
0.03-0.08	-	-	-	-																			
14	~0.1	-	~2	~3																			
Inside	-	-	Inside	Margin																			
Assumpção et. al (2002) and Veloso et al. (1992b)	Technical Report of the UnB Seismological Observatory	Technical Report of the UnB Seismological Observatory	Berrocand Fernandes (1996)	Veloso et al. (1987) and Assumpção et. al (2002)																			

N°	Name	Federative unit	Height (m)	Volume (10 ⁻³ km ³)	Maximum water depth in the Reservoir (m)	Area of the reservoir (km ²)	Start of impoundment	Geological Province	Seismicity type	Largest Events							
										Date (YY/MM/DD)	Magnitude	Magnitude Type	Io (MMI)	PGA (g)	ΔT (year)	Location	References
30	Xingó	AL/SE	150	3,800	-	60.0	06/1984	Thrust and Folding Range	Initial	1994/07/20	1,7	ML	III-IV	0.03 and below	~0.1	Margem	Berrocal and Fernandes (1996)

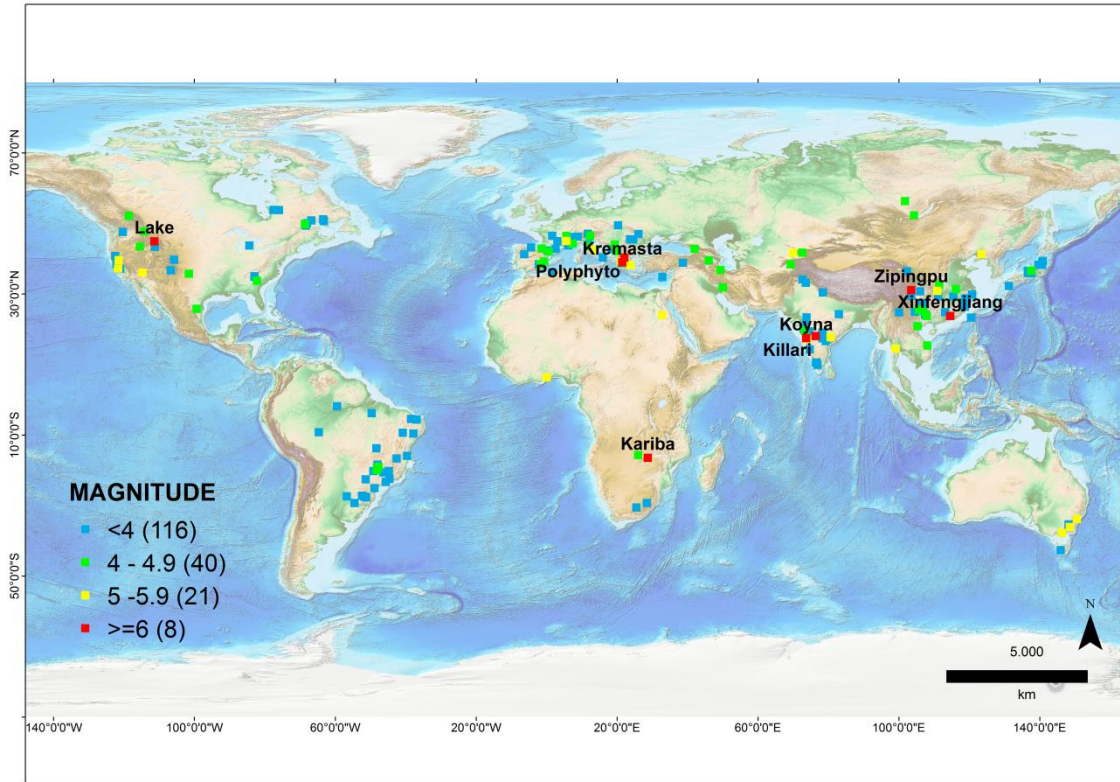
605 ΔT, Time interval (years) since the beginning of filling;MMI, Modified Mercalli Scale;
606 Doubtful cases; PGA(g), Peak Ground Acceleration; Adapted by Marza et al. (1999).
607 Table 3- Number of dams, RTSs and natural earthquakes by country regions.
608

Region	Total number of dams	RTSs	Percentage of RTS cases (%)	Number of natural earthquakes
Midwest	48	1	2 %	1821
Northeast	28	5	17.8%	2393
Southeast	167	14	8.4 %	3475
North	29	4	13.8 %	1814
South	76	6	8.9 %	139

609

610 **Figures**

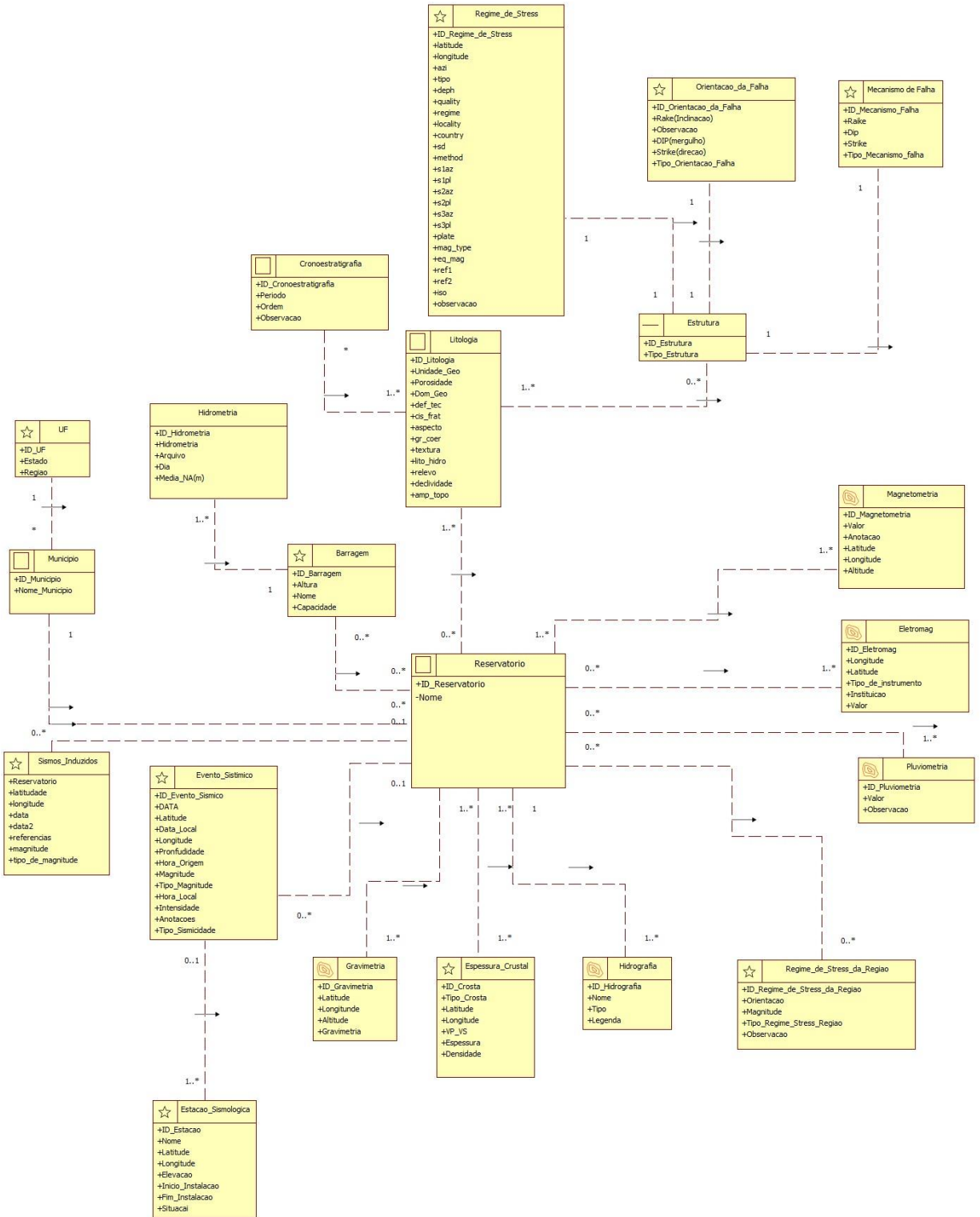
611



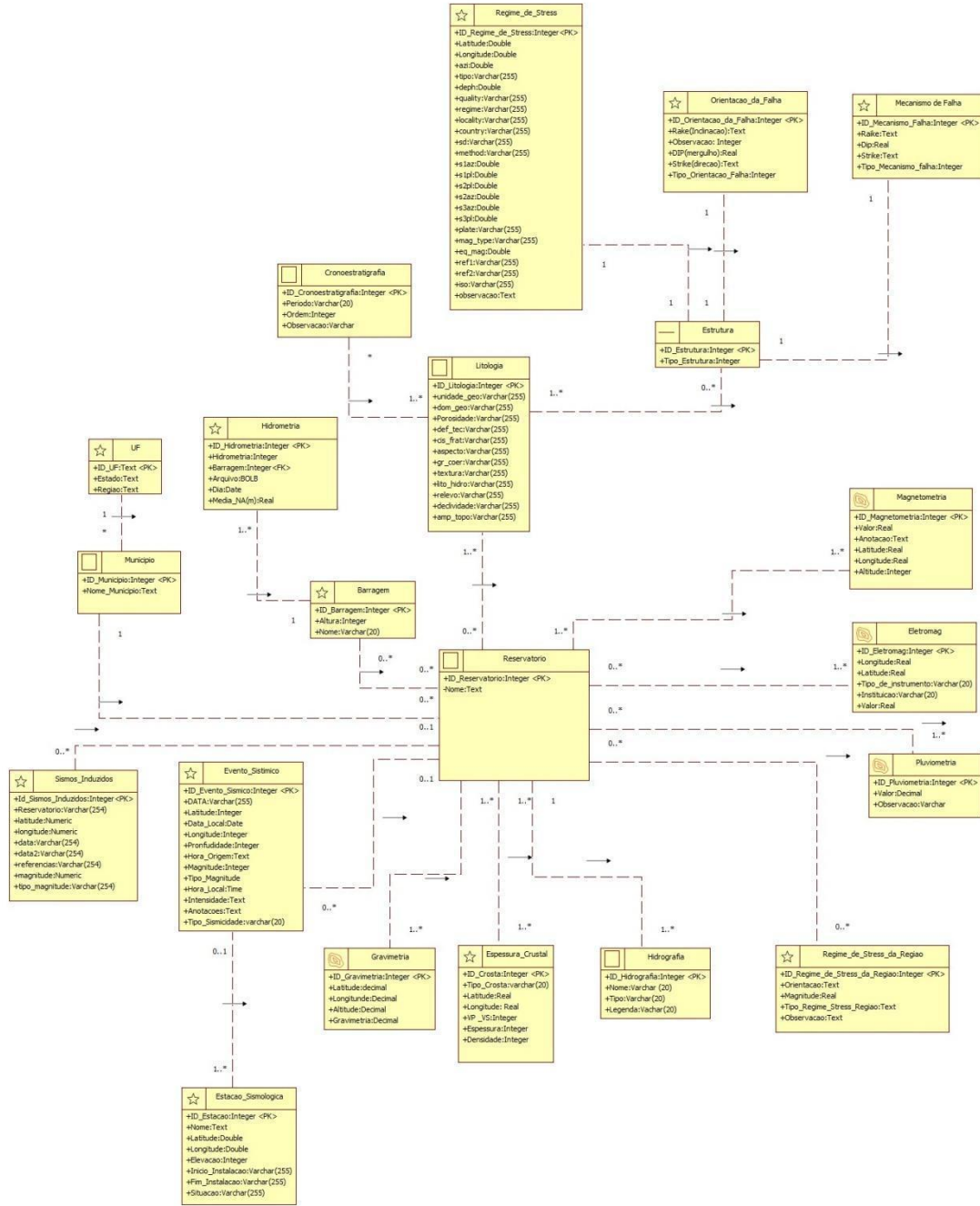
612 Fig 1 - World map of events triggered by reservoirs. (Data from the
613 www.inducedearthquakes.org. Last accessed 10 May 2020)

614

615

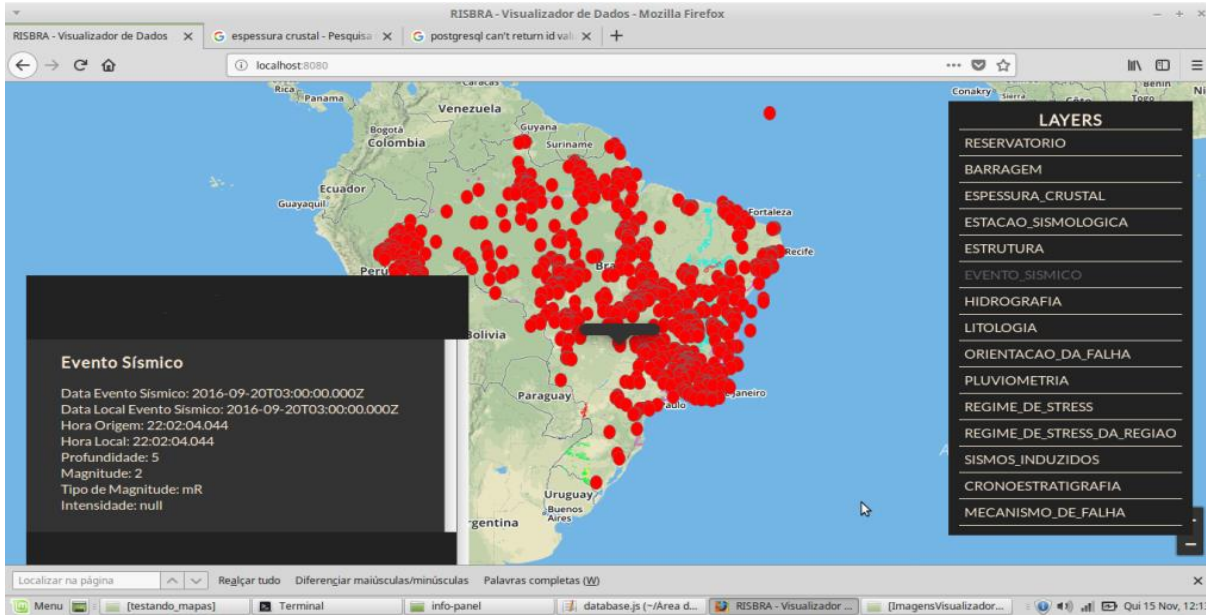


617 Fig 2- OMT-G Model of Reservoir-triggered Seismicity Database.



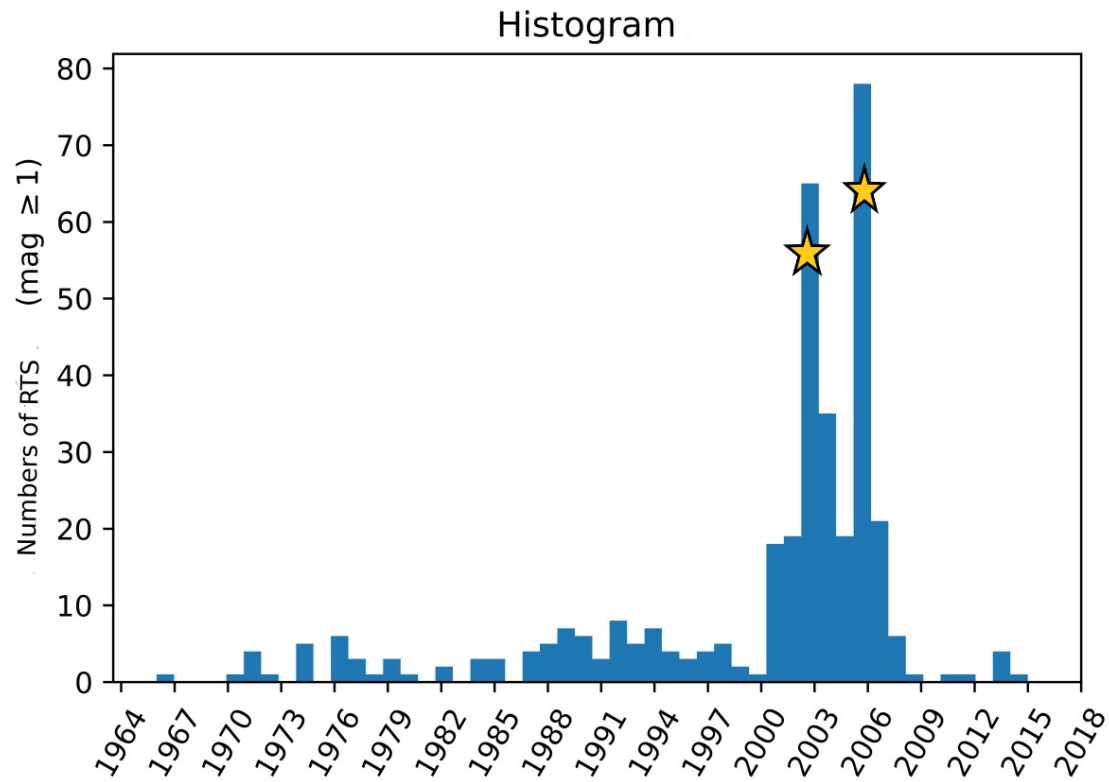
619 Fig 3- Relational model of Reservoir-triggered Seismicity Database.

622



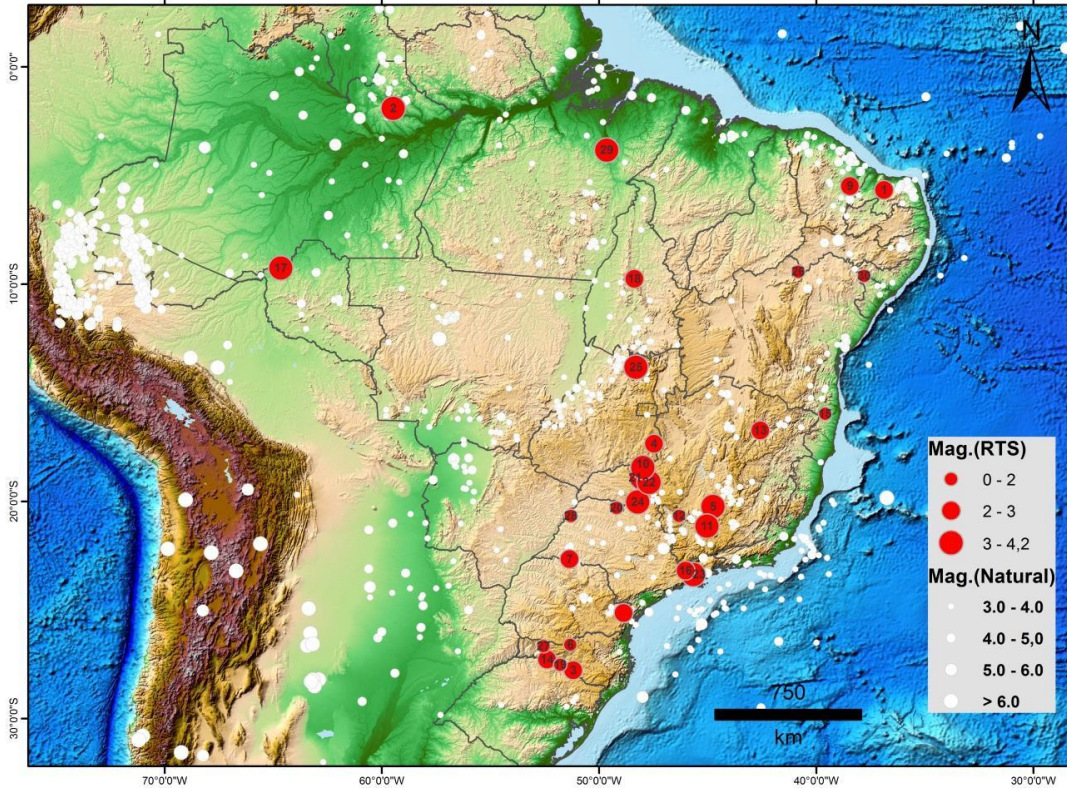
623 Fig 4- Example of researching Brazilian seismicity in RISBRA. The seismic events are
624 represented by red ball and table to the left with information regarding this seismic event layer.

625



627 Fig 5- Histogram of the RTS numbers with a magnitude greater 1, per year. The yellow stars
 628 highlight the seismic swarms at the Itapebi and Carmo Cajuru dams in 2003 and Lajeado and
 629 Nova Ponte in 2006.

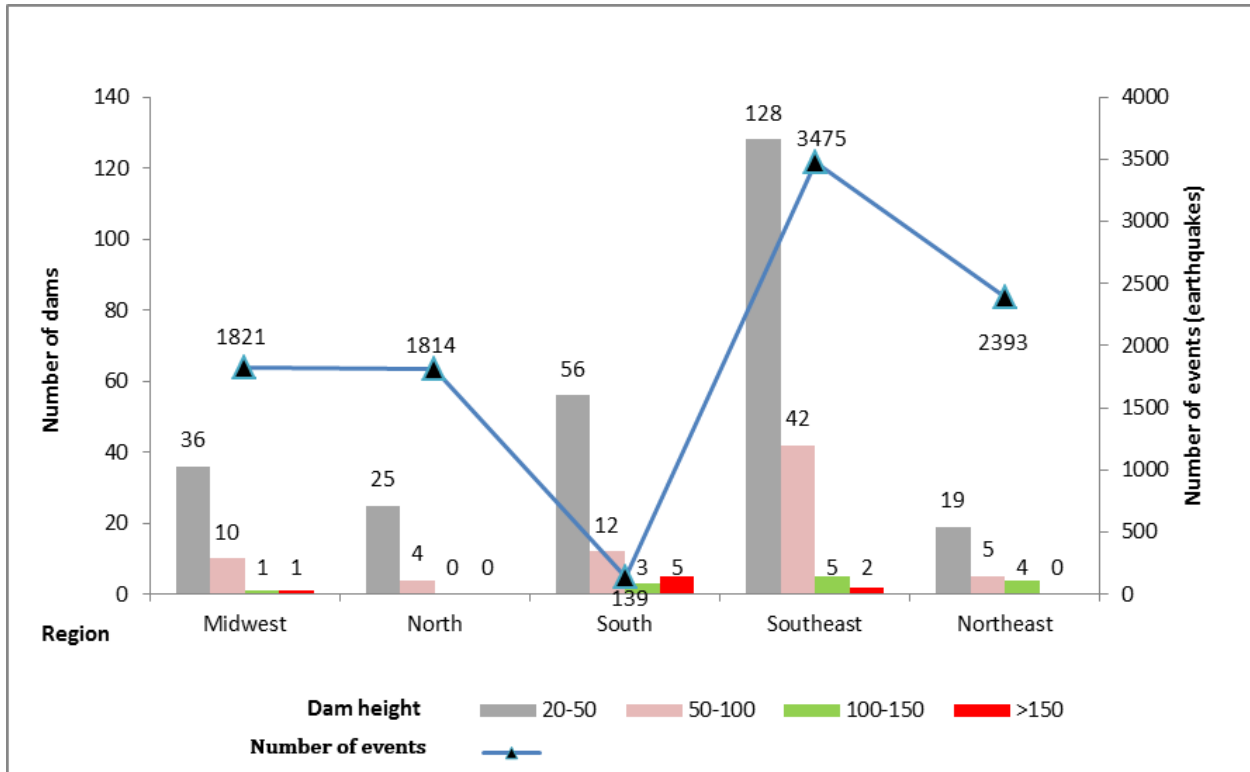
630 (



631 Fig 6 – Map of Brazil showing natural earthquakes (white circles, with magnitude) and RTS in
632 Brazil (red circles, with magnitude, numbered as stated by Table 2). Data from the bulletin of the
633 IAG-USP and SISBRA-UnB.

634

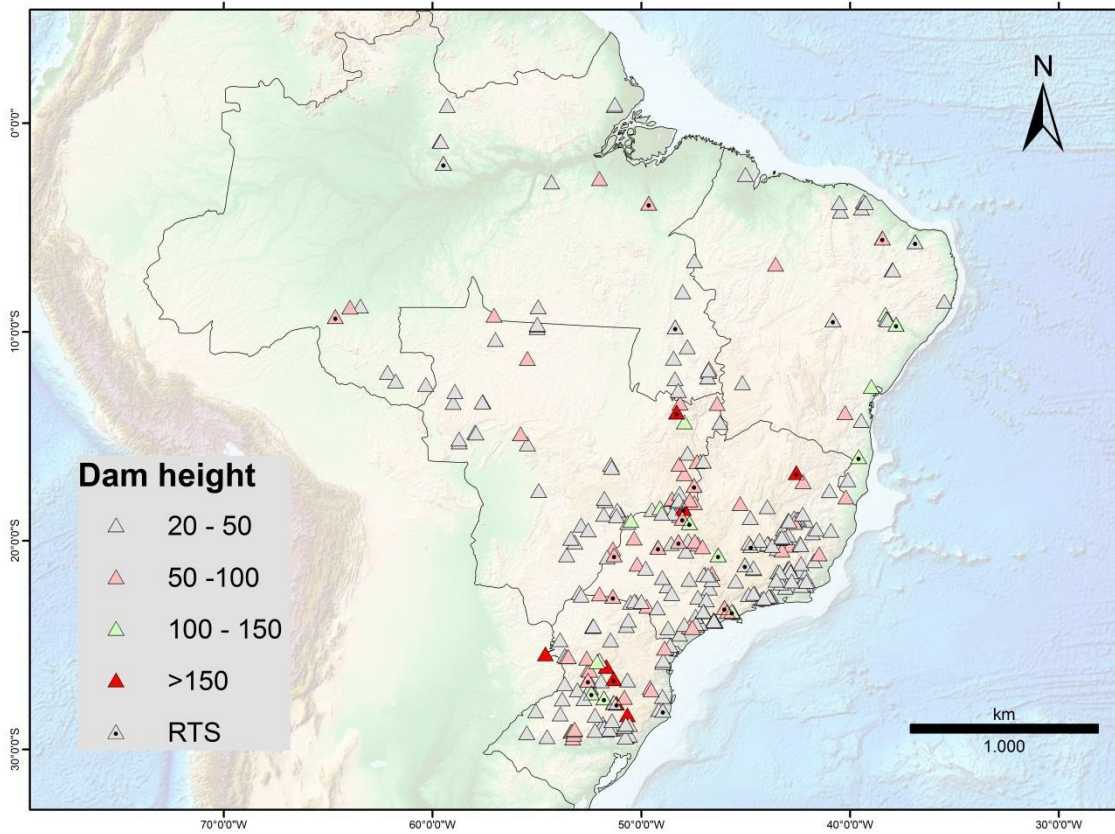
635



636

637 Fig 7- Graph showing the earthquakes, dams, and regions of the country. The southeastern region
638 concentrates the highest and the most dams in the country.

639



640

641 Fig 8- Map showing the location and classification by the dam height. Data by Brazilian
642 Committee on Dam.

643

644

645

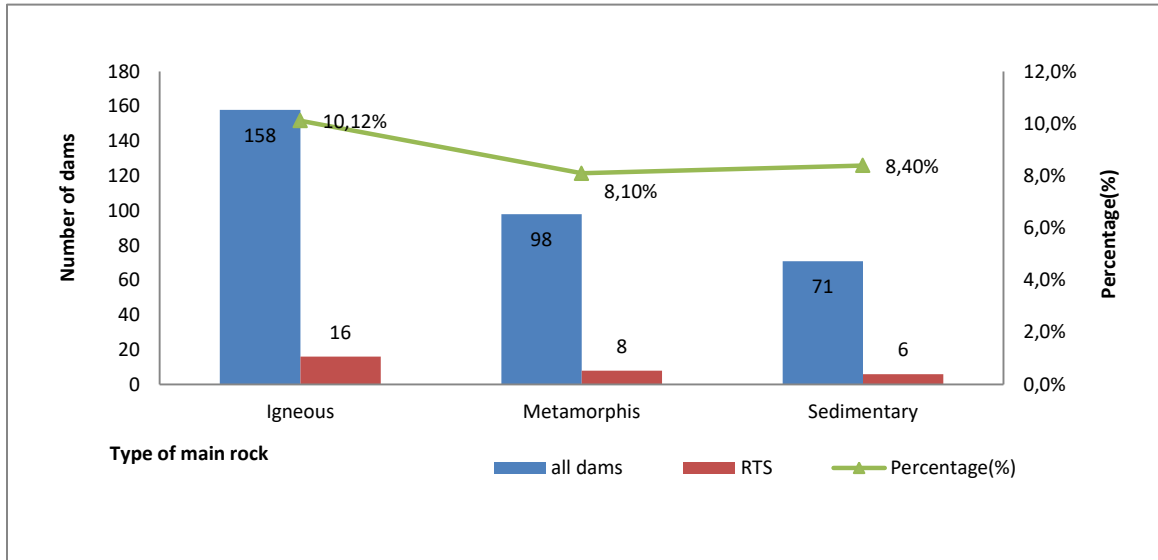
646

647

648

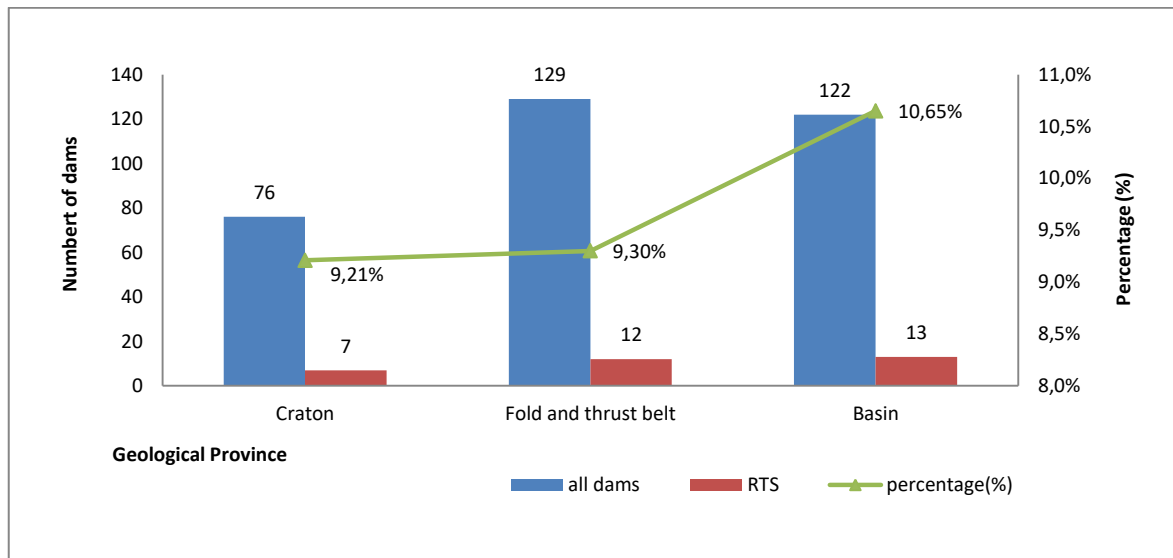
649 a)

650



651 b)

652

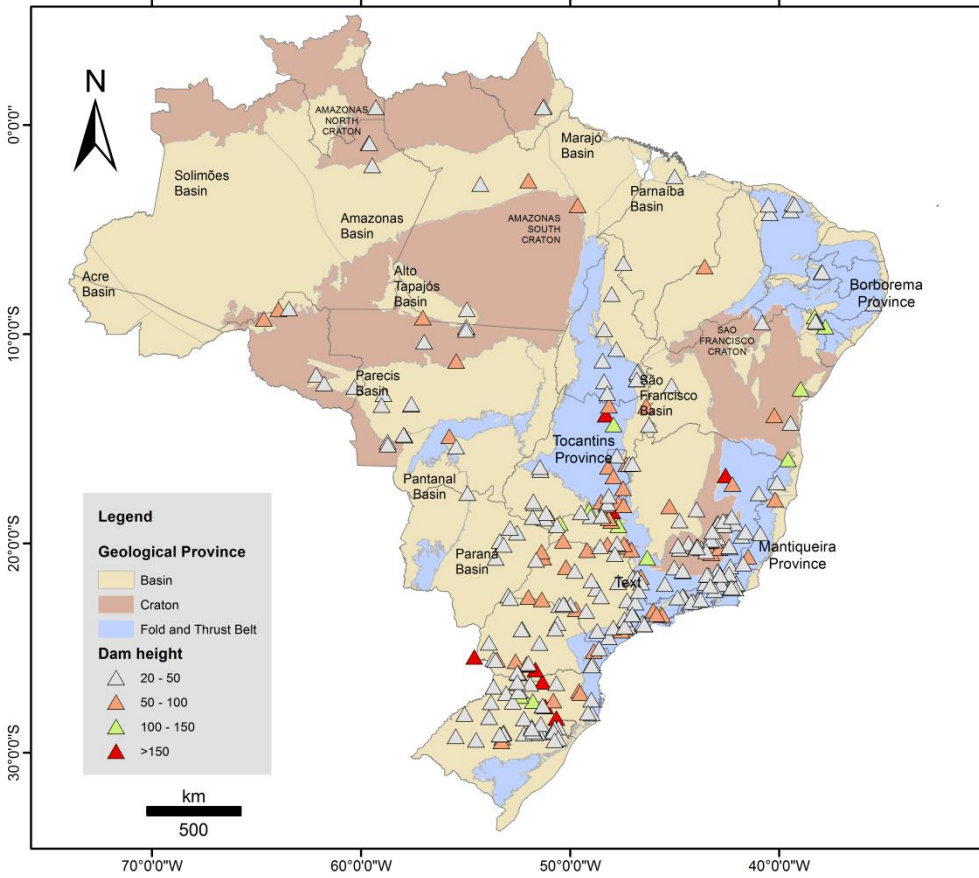


653 Fig 9- a) Percentage of cases of Reservoir-triggered Seismicity in Brazil as stated by main rock

654 types (sedimentary, metamorphic and igneous) in the dam area. b) classification as stated by the

655 main geological provinces.

656



657 Fig 10- Map of Brazil with 348 dams with a height of 20m or more (data from the Brazilian
658 Committee on Dams-2018). The colors refer to the main geological provinces (data from CPRM-
659 Mineral Resources Research Company).

660

661

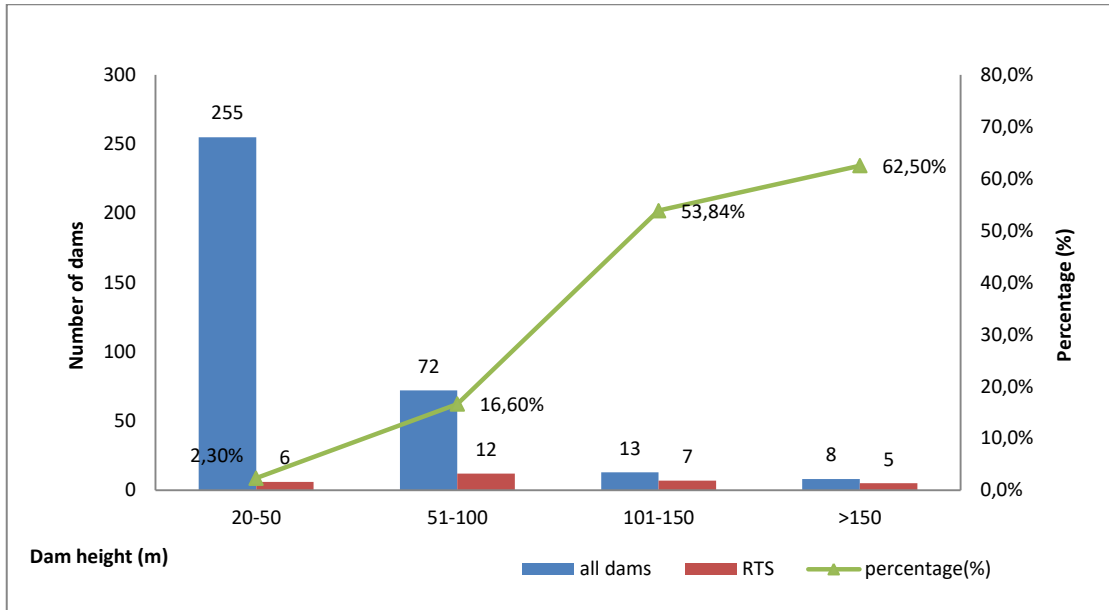
662

663

664

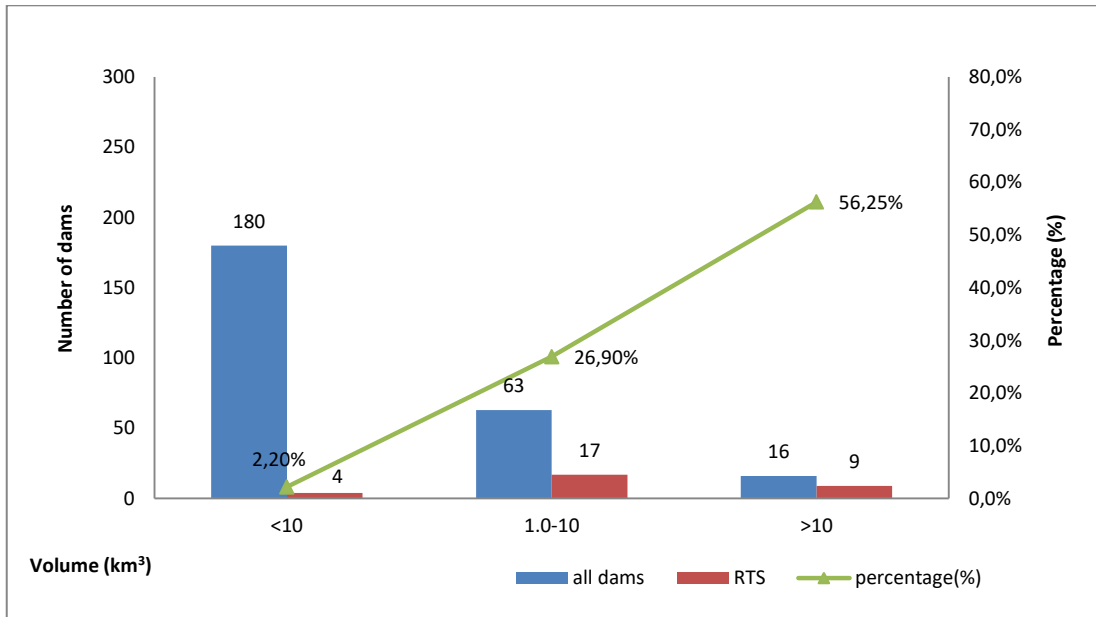
665 a)

666



667 b)

668

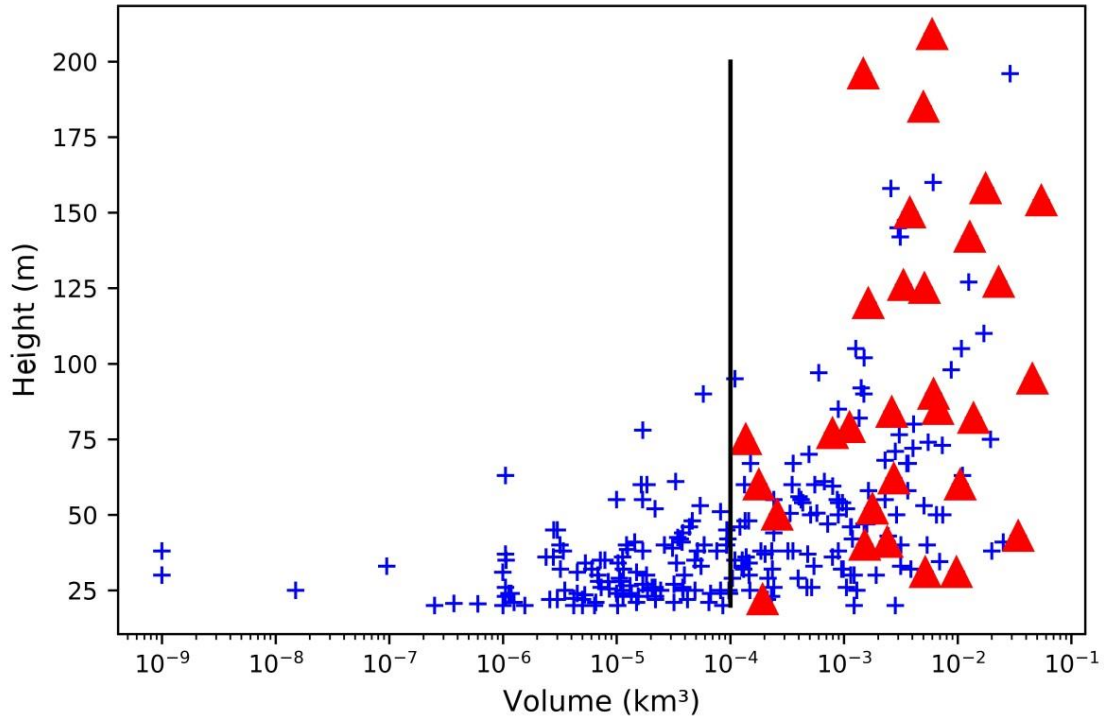


669 Fig 11- Percentage of cases of Reservoir-triggered Seismicity as stated by (a) dam height and (b)

670 reservoir volume. 54% of dams taller than 100 m trigger earthquakes and 32% of reservoirs

671 larger than $1 \times 10^{-3} \text{ km}^3$ trigger earthquakes.

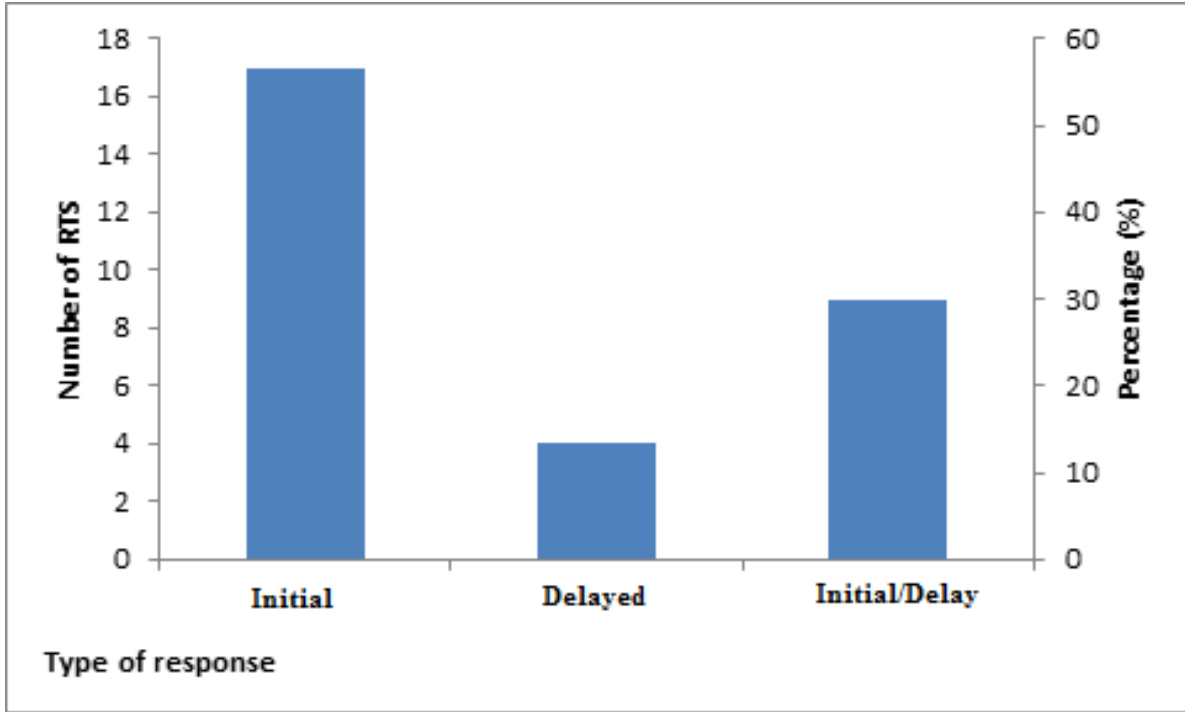
672



673 Fig 12- Graph of reservoir volume and dam height for all dams in Brazil. The triangles indicate

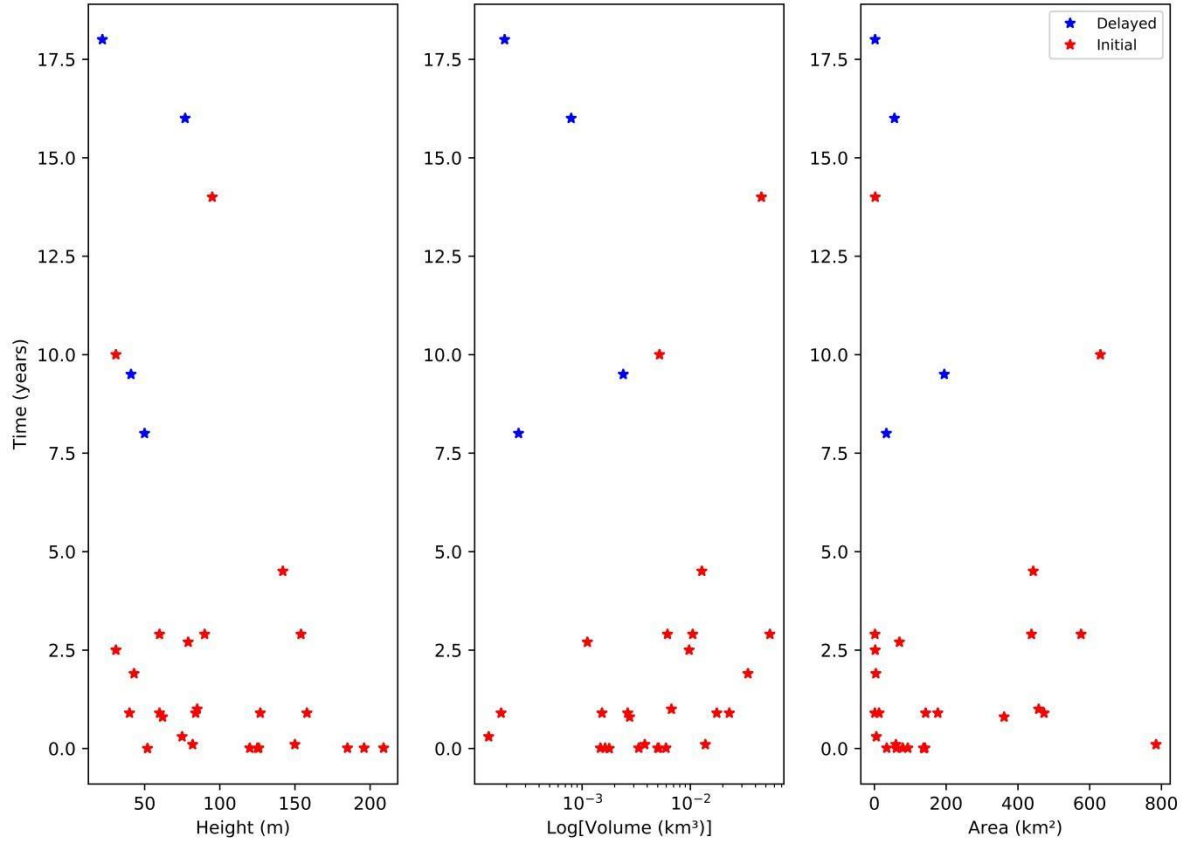
674 the RTS cases and the crosses, other reservoirs. The black bar is the limit of RTS cases.

675



676 Fig 13- Graph of the type of response for RTS cases.

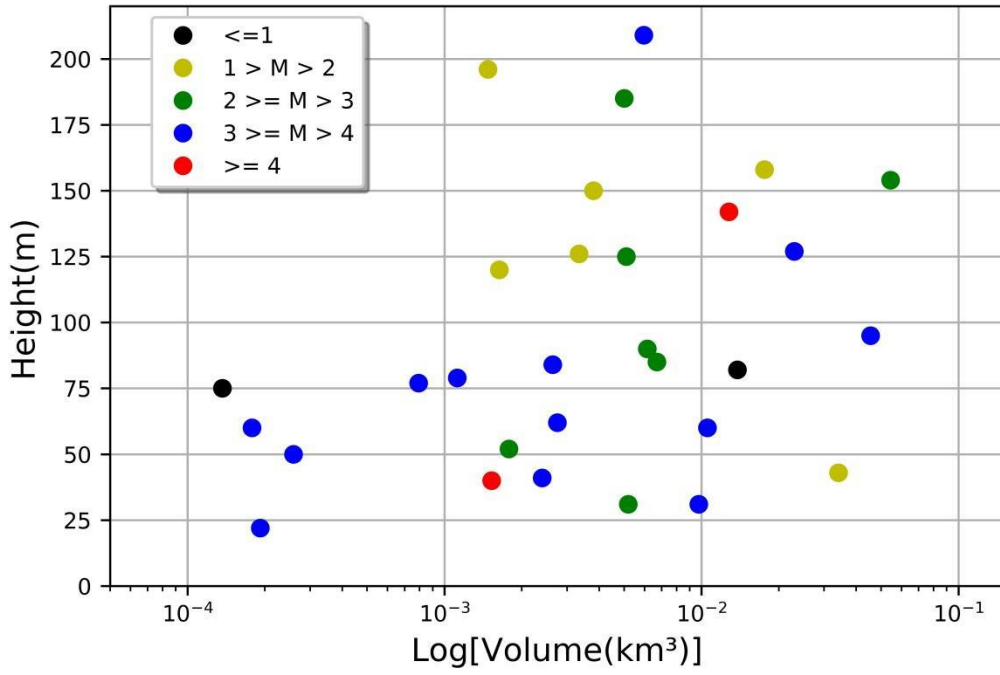
677



678 Fig 14- Graph of delay time/response versus dam height, volume, and area.

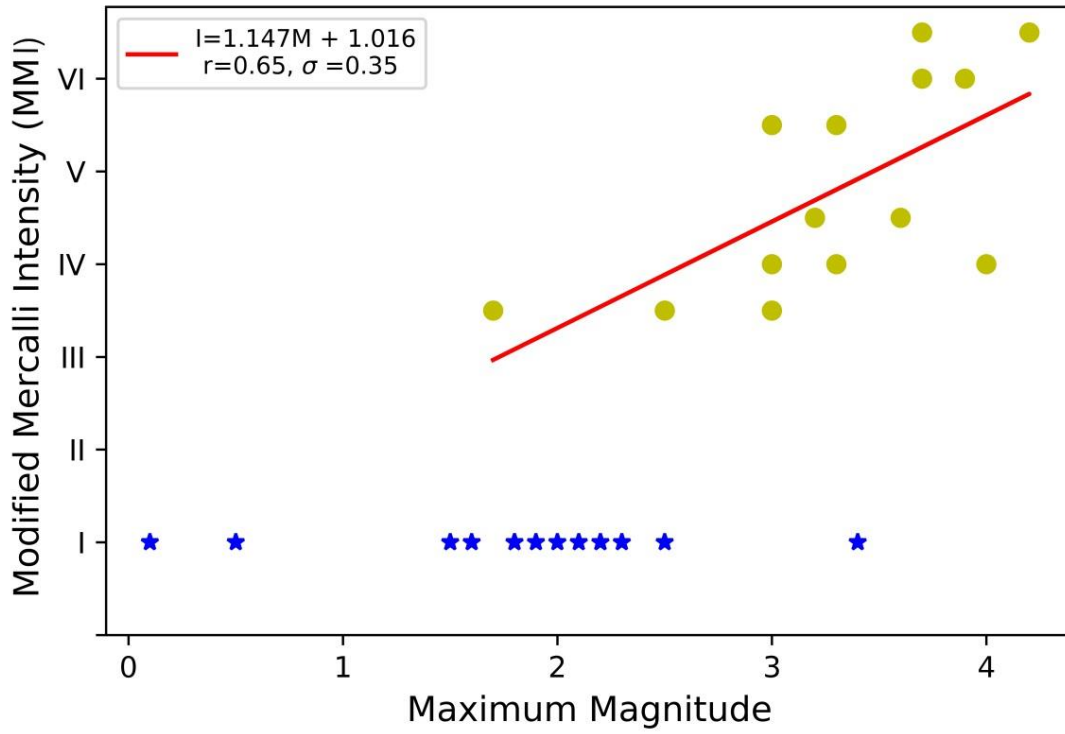
679

680



681 Fig 15- Distribution of reservoir volume and dam height versus the Reservoir-triggered

682 Seismicity maximum magnitude cases.



684 Fig 16- Graph showing maximum magnitude and intensity. The linear adjustment (bar) was

685 performed only with data represented by circles. The blue stars indicate cases of intensity I.