

Response to Referee #2

Palacio Cordoba, J., Mergili, M., and Aristizábal, E.: Probabilistic landslide susceptibility analysis in tropical mountainous terrain using the physically based r.slope.stability model, Nat. Hazards Earth Syst. Sci. Discuss., <https://doi.org/10.5194/nhess-2019-223>, in review, 2019

We are grateful for the constructive comments on our paper and have included the suggestions in the revised manuscript. Please find below our responses in detail. Changes in the manuscript are highlighted in blue letter.

- in the world? in South America?

South America – corrected.

- I don't understand the relation between the colombian andes and the world. I think the sentence has to be revise because for the reader the relation is not evident. How many landslides or proportion of landslides are triggered by rainfall in Colombian Andes ? I think the percentage mentionned by Petley has to be put above in the text.

The paragraph was rewritten for clarification (L37-40):

“Petley (2008) mentioned that in 2007, 89.6% of the fatalities due to landslides worldwide were related to rainfall-triggered landslides. In tropical environments and complex terrain such as the Colombian Andes, a high percentage of landslides are triggered by heavy or prolonged rainfall (Van Westen and Terlien, 1996; Terlien, 1998).”

- I suggest: Therefore there is a strong societal and economic request to include landslide susceptibility etc....

Thank You - Corrected according to the suggestion. (L53-54):

- It will be interesting to have information related to the landslide in South America and to know the "position" of Columbia in term of number of landslide, number of fatalities etc.... in this paragraph

There are not many studies with a specific rank, however considering the recommendation of the reviewer additional to the rank of Mergili et al (2015) a new rank according to Petley and Sepulveda (2015) was incorporated (L60-65):

“... where Colombia is the country with the second highest number of fatal landslides in Latin America and Caribbean, only Brazil shows a slightly higher number of fatal landslides (119) (Sepúlveda and Petley, 2015). Mergili et al. (2015) expose that Colombia, according to EM-DAT, presents a victim/event ratio of 77.34 (3171/41) and is only exceeded by Peru and Ecuador in this respect. However, the EM-DAT database only considers events with high numbers of fatalities: actually, the real number of fatal landslides is much higher (Aristizábal and Gómez, 2007; Mergili et al., 2015).”

- You have to improve this part because since SAFELAND project, cascini 2008 Corominas et al 2014 and van Westen et al 2007 it is possible to divide the method by qualitative and quantitative method. I think you have to be in the recent scientific paradigm to describe the different method and approach for LSA and LHA. Data driven are separate from Physically in the group Quantitative methods (see Corominas et al 2014) and the different works of the JTC 1.

The LSA classification was adjusted according to Corominas et al. (2014) (L66-69):

“... qualitative or by quantitative methods (Corominas et al., 2014). Qualitative methods correspond to knowledge-driven approaches based entirely on the judgment of experts using geomorphological criteria in the field (van Westen et al., 2000). Quantitative approaches are subdivided into data-driven methods and physically-based models. Statistical or data-driven methods ...”

- it is very nice to cite ourself but Mergili never introduce the concept of FoS, I prefer you put Duncan for instance. I think the text of Mergili et al 2014 is cited after in the text.

The Mergili et al. (2014) and Ahmed et al. (2012) citations were eliminated, we think that the reference to Lam and Fredlund (1993) is enough.

- idem of the comments above.

The Mergili citation was eliminated.

- Mergili is not the first to mention this problem. I think there is in the literature some key paper mentioning this difference, I suggest to change that.

Thanks for the observation, more references have been included. However the authors think it is difficult to establish who was the first. The references only try to provide some books and papers where this issue is discussed in detail ([L80](#)):

“... (Bishop, 1954; Carson & Kirkby, 1972; Crozier, 1986; Duncan & Wright, 2005) ...”

- I think the transition between the two paragraphs is not very well structured. The first sentence will be put above in the text. Later you can speak about Columbia.

Thanks for the recommendation, the paragraph was adjusted. ([L105-112](#))

- I have a problem, this is a scientific paper and I don't see the scientific problematic. The introduction is well written, the problematic misses. Where are the scientific questions? I think you have to include a little paragraph between to explain the problematic and why you want to use R slope stability and what do you expect with it.

Yes, this is an important point which we have not explained in sufficient clarity in the original manuscript. We have extended the explanation accordingly ([L105-120](#)):

“In terms of scale or spatial resolution, physically-based models are suggested to be applied to finer-scale study areas, whereas data-driven approaches are recommended for broader-scale landslide susceptibility analyses (van Westen et al., 2006; Corominas et al., 2014). In fact, the implementation of data-driven methods or physically-based models for the incorporation of landslide hazard mapping into land use planning has been regulated in several countries. In Colombia, according to the Decree 1807/2014, the implementation of deterministic or probabilistic physically-based methods is obligatory in urban and urban expansion areas, whereas statistical and knowledge-driven models are only permitted for rural areas.

In order to meet this requirement, we have to obtain a more detailed knowledge on the suitability of the available physically-based landslide susceptibility models, with the final goal to use model ensembles in order to obtain a broader, more robust picture of the landslide susceptibility conditions. We think that r.slope.stability could be a candidate to enrich such an ensemble, considering shallow planar and deep-seated ellipsoidal failure surfaces. Consequently, the scientific aims of the present study are (i) to evaluate the suitability of r.slope.stability for physically-based landslide susceptibility mapping in tropical mountainous terrain, and (ii) to identify its fit with other potentially suitable models, helping to learn about strengths, limitations, and uncertainties.

Following these aims, ...”

- Maybe this is the point for a problematic, why you want to compare results with another results obtained with another tools? I think this is one of the key points of your paper and you have to develop it in the introduction

See the response to the previous comment, particularly ([L113-116](#)):

“In order to meet this requirement, we have to obtain a more detailed knowledge on the suitability of the available physically-based landslide susceptibility models, with the final goal to use model

ensembles in order to obtain a broader, more robust picture of the landslide susceptibility conditions. We think that r.slope.stability could be a candidate to enrich such an ensemble, ...”

and (L118-119):

“... (ii) to identify its fit with other potentially suitable models, helping to learn about strengths, limitations, and uncertainties.”

- is it possible to differentiate the type of landslides in the figure ?

This is a good point, such a differentiation would be useful, indeed. However, it is unfortunately not possible, as the landslide database did not provide this information: it is just described in a general way that rotational and translational landslides were found (Hermelin et al., 1992; Garcia, 1995). (L159-160):

“However, there is no distribution map which permits the differentiation between translational and rotational landslides”

- you speak about rotational and translational landslide above, have you got this typology ?for the area?

See above comment.

- i think a desription of instabilities is a prerequisite for this type of study. The reader want to know which type of phenomena wll be modelled and the number etc etc.... I suggest to add a paragraph about labdslides type, number, a a bried description of them with the main predisposing factors

Yes, we agree. The description provided in the first paragraph of this section was rather short in the initial manuscript, so we have complemented it according to the suggestion of the reviewer (L168-174):

“A total of 699 landslides were identified and mapped in the La Arenosa catchment. According to Velásquez and Mejía (1991), most of the landslides started as shallow translational slides and transformed into debris flows, from very to extremely rapid with high water content (Fig. 1). The landslide bodies were small with respect to the flow length, and the slip surfaces were parallel to the slope surface. Field studies carried out by The majority of the landslides initiated within residual soils in hollows and open slopes with a slope steepness ranging from 35° to 42° (Velásquez and Mejía, 1991). The same authors described that the depth of failure surface was less than 3 m and corresponded to the contact of the residual soil with the saprolite.”

- the sentence has to rewritten. for instance "The tool is able to take into account planar failure with an infinite slope stability module and another type of failure with a slip surface module . "

Thanks for this recommendation, the formulation was adjusted accordingly. (L184-185):

“The tool is able to take into account planar failure with an infinite slope stability module and a slip surface module”

- I suggest to inverse the two paragraph about the modules

1. planar module

2. slope surface module

in order to be consistent with the sentence above and the order

Thanks for the recommendation, the order was adjusted accordingly.

- I suggest to split the paragraph of the too; in two part

1. description of the oS modules with slope stability model and infinite slope stability model

2. probabilistic model

the paragraph will be clearer

Thank You also for this recommendation, we have split the paragraph accordingly

- I understand the concept and what you mean, but the sentence is not very clear. I suggest to rewrite it.

Yes, we have rewritten this sentence as follows (L213-214):

“The slip surface model includes a probabilistic component where the dimensions of the ellipsoids are randomly varied for the computation of FoS.”

-please explain why you use this approach. Justify it please. Is there a strong relationship between slope and thickness of surficial formations and weathered layers? If this is the case, you can refer to the paper written by Thiery et al., 2019 in engineering geology, in addition to Aritzabal 2013, the authors refer to the relation between slope degrees and different weathered horizons in crystalline rocks.

This is an important point. We had only provided the Aristizábal (2013) reference, because in that work they established the relationship between the slope and the soil weathering profile according to field work in this catchment. However according to the suggestion of the reviewer the references to Thiery et al. (2015) was included as well as a reference to Catani et al. (2010).

- for comparison of what? of another results with another tool ? if this is the case please explain or better put it in the introduction as it mentioned in my antecedent comments

The sentence was rewritten for clarification (L247-248):

“The r.slope.stability model is applied with the probabilistic approach and deterministic approach for comparison with other models tested in the catchment.”

- in R slope stability you have different shape of probability function. I think it will be good to explain that in the paragraph dedicated to the model.

Yes – Section 3 (L213-214):

“Thereby, rectangular, normal, log-normal, or exponential probability density functions are employed.”

- which method ?

The paragraph was partly rewritten. The maximum width and length of the ellipsoid is determined according to what is seen in the field according to the inventory. The number of ellipsoids to be tested (1000) was the one where the results did not vary too much when the number was further increased and where data processing does not require a high-performance computer.

The text in the revised manuscript read as follows (L255-259):

“... 1000 simulated surfaces touching each cell. Preliminary tests have indicated that this value represents a good compromise between computational efficiency and accuracy of the results. Each ellipsoidal slip surface is defined by the coordinates of the centre and the variable lengths of the three half-axes, the aspect, the inclination and the offset centre above the terrain, ...”

- why you use this values? they correspond to the size of landslides (min, max?) If you put a paragraph about the slope instabilities and description of them, these values could be justified. I encourage you to write this paragraph as suggested before

The assumed dimensions (and also the depths) of the simulated landslides have been defined according to the landslide inventory (we have added a more detailed description of the landslides to Section 2 – see above comment).

- justify the value. The max depth of translational landslide is 2.8 m ? see the comment just before about size and depth justification.

See response to the comment above.

-evaluation or validation? I think you validate the models with the inventory. The evaluation will be made with the different previous maps and the expert judgement. Please revise this paragraph.

We are aware that there is no general agreement on the terminology with regard to “*evaluation*” and “*validation*”. We think that, as the reference information (landslide inventory) is also uncertain, the term “*validation*” would be too strong, and rather prefer to keep “*evaluation*”.

- put the descriptions in a table. These indices are very well known by the scientific community working on landslides with different spatial susceptibility and hazard approaches

We agree that those are very well-known indices. However, we still think that particularly for this reason, an additional table is not justified. So, we rather prefer to keep the description of these parameters as it is.

- "to evaluate the performance of the tool". Moreover for me the problem of this type of procedure (just statistical comparison) is that the judgement of the expert is neglected. Results of these models have to be kept as 'indicators'. You are the expert of this area, normally you are able (and this is the duty of the scientist) to have a criticism about results and the different location of high, or another susceptibility classes. For instance sometimes the models give good ROC curve and good landslide recognition, but some areas are not taken into account or inversely you can obtain huge areas with very high susceptibility but in reality (real world) it is impossible (because you don't take into account some parameters) and with your data you oversimplify the real world.

At the beginning of the paragraph, we have added “... *to evaluate the consistency between the model results.*” We cannot evaluate model performance by the comparison with the results of other models, because they might be as true or wrong as the *r.slope.stability* results.

We fully agree that it is necessary to somehow combine expert opinion and simulation. The present paper focuses very much on simulation-based approaches of slope stability analysis. Evaluation of the model builds on a landslide inventory, which has been mapped by experts through subjective judgement, based on the available evidence. Consequently, there is – in our opinion – a strong expert-based component in the evaluation of the model results, even though the evaluation itself is a quantitative one. However, expert judgements are also necessary for converting the model results into hazard classes and associated recommendations. This is made clear in the final part of the revised conclusions (L451 -462; see also response to the last comment).

- what is the goal of this comparison/evaluation of *r slope stability* results? what do you want to explain with this comparison? You have to add some explanations about the goal, the objective of this comparison

In general, using different models for one and the same area and landslide event increases the robustness of the results, as each model is based on different assumptions and different ways of implementation of the equations. Similar results yielded by different models indicate some degree of plausibility, whereas different results help to identify possible shortcomings, limitations, or even bugs. (L113-116):

*“In order to meet this requirement, we have to obtain a more detailed knowledge on the suitability of the available physically-based landslide susceptibility models, with the final goal to use model ensembles in order to obtain a broader, more robust picture of the landslide susceptibility conditions. We think that *r.slope.stability* could be a candidate to enrich such an ensemble, ...”*

and (L118-119):

“... (ii) to identify its fit with other potentially suitable models, helping to learn about strengths, limitations, and uncertainties.”

- which type of landslide you try to model with infinite slope model ? all ?

The inventory available to us does not allow for the distinction between shallow translational and deep-seated rotational landslides, but we know that both shallow and deep-seated landslides occurred in the 1990 event (see comment above). Therefore, we have used both models in parallel, in order to cover both types of landslides, expecting that the “truth” is somewhere in between the two results. A major advantage of r.slope.stability is that it includes both the infinite slope stability model and the sliding surface model. A possible future strategy could be to combine both models in one single probabilistic simulation.

- you have two type of landslide? you compare tyoe by type as suggested by Thiery et al., 2007 , Corminas et al., 2014 ? This is important because each landslide type has its own prediposing factors diferent from one to another one. I am dubitative, are the mudslide has the same prediposing factors as the soil slip ? The reader is not able to know because no paragraph about landslides is included in the paper.

We have tried to clarify this issue in Section 2 by describing the landslides in more detail (see comment above).

- why you used this thresholds for the FoS classes ? please justify beacuse in some studies authors divide it in 4 classes (0-1/1-1.2/1.2-1.5/>1.5)

Yes, we are aware about this issue – to our knowledge, there is no generally accepted “standard” classification. We just consider the chosen classification useful for our purpose of comparison, without claiming its general validity.

- I repeat but why calculate two maps ? please justify in the text if you have two type of landslides (translational shallow and deeper landslides with rotational or another slip surface)

Please see above comment.

- and in an expert point of view..... the different probability can be divide following the different thesholds defining by the JTC 1 works, they give somme classes or by old paper written by Kirsten HA. (1983) Significance of the probability of failure in slope engineering. Civ Eng 25:17–27

The paper of Kirsten (1983) was checked in detailed and the text was modified, however the thresholds were not modified considering that they were used just to compare between different models, they are not proposed for further use. According to Kirsten (1983), different thresholds should be used for specific scenarios and according to the local conditions. The following sentence was added (L92-94):

“Although FoS represents a quantitative – and seemingly objective – approach to evaluate slope stability, it has to be used carefully. Authors such a Kirsten (1983) mention that different values of FoS could be obtained from slopes with equal probability of failure. However, several software packages use this concept ...”

- i don't understand (once again) you compare all landslide type with each final map. in this sense you compare maybe shallow landslides with map simulated for deeper landslides. This is not correct.

This issue is related to the missing distinction of shallow and deep-seated landslides in the inventory. However, we know that both shallow and deep-seated landslides occurred in the 1990 event (see comment above). Therefore, we have evaluated both models in parallel with the entire inventory, in order to cover both types of landslides, expecting that the “truth” is somewhere in between the two simulation and evaluation results. A major advantage of r.slope.stability is that it includes both the infinite slope stability model and the sliding surface model. A possible future strategy could be to combine both models in one single probabilistic simulation.

- ok and what is the signification really ? it is not very clear for me

The following explanation was added (L362 –369):

“However, both deterministic analyses of r.slope.stability show the highest false alarm rates. The major reasons for this result – and for the similar patterns yielded with SHIA_Landslide – are most likely an underestimate of the geotechnical stability of the material, the neglect of the effects of vegetation, and/or the overestimate of water saturation. Geotechnical testing is often performed on material which is disturbed in the one or the other way, so that the resulting parameters do not necessarily represent the natural conditions over larger areas. Roots could lead to some degree of stabilization, and it could also be the vegetation retaining sufficient water to avoid full saturation of the soil throughout the catchment.”

- Never you mention that before an dthe fact that you try to replicate this event !!! Please revise the text before to explain this point !!!!

We have mentioned this aspect in the last paragraph of the introduction ([L121-123](#)):

“This model is evaluated using a landslide inventory prepared after a major and destructive rainfall-triggered multi-landslide event in the La Arenosa catchment on 21 September 1990.”

- yes but this is not the only tool which can used probabailistic analysis TRIGRS now made the same type of calculation, ALICE by the BRGM can made the same calculation. I think you must mention them in your discussion? some references are available on the web.

Yes, this is a good point. We have added the following sentence ([L383-385](#)):

“Besides r.slope.stability, there are several physically-based models which provide a probabilistic module, such as TRIGRS (Baum et al., 2008) and ALICE (Thiery et al., 2017). Much more work has to be done in this direction.”

- or a ground water level because inmany cases this is not the rainfall introduced in the model but the considered GWL see and mention Fressard et al 2016 please, they explain that

Yes, this is an important aspect. We have replaced “rainfall event” by “hydro-meteorological event” and added the following text to the description of the study area: ([L135-138](#)):

“Most of the landslides in the catchment are triggered by such events, even though it is often not the rainfall itself, but the response of the groundwater level – also depending on its initial state due to antecedent rainfall conditions – which is decisive for the occurrence of landslides (Fressard et al., 2016).”

- slope conditions of what ? I think you want to write the slope instabilities or slope stability conditions
Thanks also for this suggestion, we have changed the text to “slope stability conditions”.

- and the type of failure, but you never mention that. please you have to improve your text to mention this point !!!

It was included in Section 2 (please see response to above comment).

- Why ? maybe during this event only translationnal and shallow landslides occurred ?

Thank You for this (and also the next two) comments. This issue was not clearly explained in the original manuscript, so that we have now tried to clarify it ([L405-408](#)):

“... which is confirmed by the reference information: Even though Hermelin et al. (1992) and Garcia (1995) mentioned that both shallow and deep-seated landslides were observed after the studied rainfall event, Velásquez and Mejía (1991) underlined the dominance of shallow landslides.”

- because you have only translationnal landslide ?!!!!?

Please see response to the above comment.

- ok but in your area is it only this type of landslides ? if yes why you calculate another type of slope failure ?

Please see response to the above comment.

- YES !!! but in your case ?

We have tried to explain this aspect in a clearer way in the revised manuscript (L415-419):

“... or – if both mechanisms are relevant – evaluate the model results with different sub-datasets of the inventory. In the present study, due to the lacking distinction of landslide mechanisms in the inventory at the level of individual landslides, we applied both the infinite slope stability model and the slip surface model in order to cover a broad range of mechanisms. The result of the evaluation against the entire landslide inventory was in line with the generally reported dominance of shallow translational landslides.”

- For more readability, it will be better to put the same number of classes for each map (3 or 4)

The map derived with Shalstab was reduced to 3 classes, considering a single class for “unstable” and “unconditionally unstable” (Fig. 6D).

- I am disappointed because you don't speak about the real validity of the final map and if the different areas simulated in different susceptibility classes correspond to slope considered by expert as susceptible or not. For you if the landslides are well recognized by the model(s) the maps are ok. This is not the case everywhere. The different maps could be used for future developments (you explain that in your introduction) but now which map could you use to produce a regulatory map ? it will be a very interesting discussion and one paragraph about that will improve the paper and give your point of view about that. You can find one interesting example Thiery et al. (2017) in GEOSCIENCES about the use of this type of simulation by PBM for future developments. I encourage you to consult the paper.

Thanks for the recommendation – indeed, this is a very important aspect. We see a major role of expert opinion in the conversion of slope failure probabilities to hazard classes and recommendations for action, something which is a highly non-trivial task. We conclude the revised manuscript with the following piece of text, also mentioning another important aspect (L451 -462):

“In Colombia, hazard mapping is mandatory for use in land planning in urban areas. Results like the slope failure probability (Fig. 6F and Fig. 6G) are considered suitable for this purpose. However, they only represent one step on a long way:

- (i) Much of the landslide risk in the Arenosa catchment is not so much related to the failure of unstable slopes, but rather to the downslope propagation of the mobilized material as debris flows (Velásquez and Mejía, 1991). Therefore, coupling of the slope stability model results with mass flow simulation tools is absolutely necessary (Mergili et al., 2017; Bout et al., 2018).*
- (ii) Slope failure probabilities – or impact indicator indices (Mergili et al., 2017) – neither take into account the dimension of time nor the fine-scale patterns of the geotechnical characteristics of the catchment. Consequently, those results have to be interpreted in a relative, qualitative rather than an absolute, quantitative sense. The combination of these maps with expert judgement is therefore essential to define suitable thresholds separating the study area into different zones with their individual hazard levels and recommendations for action.”*

Probabilistic landslide susceptibility analysis in tropical mountainous terrain using the physically based r.slope.stability model

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15 **Abstract.** Landslides triggered by rainfall are very common phenomena in complex tropical environments such as the Colombian Andes, one of the regions of South America most affected by landslides every year. Currently in Colombia, physically based methods for landslide hazard mapping are mandatory for land use planning in urban areas. In this work, we perform probabilistic analyses with r.slope.stability, a spatially distributed, physically based model for landslide susceptibility analysis, available as an open-source tool coupled to GRASS GIS. This model
20 considers alternatively the infinite slope stability model or the 2.5D geometry of shallow planar and deep-seated landslides with ellipsoidal or truncated failure surfaces. We test the model in the La Arenosa catchment, northern Colombian Andes. The results are compared to those yielded with the corresponding deterministic analyses and with other physically based models applied in the same catchment. Finally, the model results are evaluated against a landslide inventory using a confusion matrix and Receiver Operating Characteristic (ROC) analysis. The model
25 performs reasonably well, the infinite slope stability model showing a better performance. The outcomes are, however, rather conservative, pointing to possible challenges with regard to the geotechnical and geo-hydraulic parameterization. The results also highlight the importance to perform probabilistic instead of – or in addition to – deterministic slope stability analyses.

1. Introduction

30 Landslides cause substantial human and economic losses every year (Kjekstad and Highland, 2009; Petley, 2012; Schuster and Highland, 2001). According to Dilley et al. (2005), the worldwide area exposed to landslides is around 3.7 million km², where 66 million people live in the 820,000 km² identified as the high-risk zone. Although economic losses tend to concentrate in industrialized and developed countries, the numbers of human fatalities and affected persons are highest in densely populated, less developed countries (Petley, 2012; Sepúlveda and Petley,
35 2015).

Landslides triggered by earthquakes and rainfall are a frequent phenomenon in mountainous terrain (Keefer et al., 1987; Van Westen et al., 2008; Varnes, 1978). Petley (2008) mentioned that in 2007, 89.6% of the fatalities due to landslides worldwide were related to rainfall-triggered landslides. In tropical environments and complex terrain such as the Colombian Andes, a high percentage of landslides are triggered by heavy or prolonged rainfall (Van Westen and Terlien, 1996; Terlien, 1998).

Colombia, located in the north-western corner of South America, exhibits complex geographical and hydro-climatological features arising from its tectonic setting and equatorial location. The mountainous configuration of Colombia is the result of the Caribbean Plate moving south-westward relative to the South American Plate and the eastward subduction of the Nazca Plate beneath the northern Andes along the western margin of Colombia (Kellogg et al., 1995; Taboada et al., 2000; Trenkamp et al., 2002). Related to the hydro-climatological conditions and because of its equatorial location, rainfall in Colombia is highly intermittent in space and time, influenced by the atmospheric circulation patterns over the neighboring tropical Pacific Ocean and the Caribbean Sea and the combined hydro-climatic and ecological dynamics of the Amazon and Orinoco basins (Poveda et al., 2007).

Shallow landslides triggered by rainfall are very common phenomena in tropical environments such as the Colombian Andes, where hillslopes are characterized by deep weathering profiles and are subjected to periods of intense tropical rainfall (Aristizábal, 2013). In Colombia, landslide-prone regions are densely populated. As a consequence, hundreds of fatalities are associated with landslides triggered by rainfall every year (Sanchez and Aristizábal, 2018). Therefore, there is a strong social and economic need to include landslide susceptibility and hazard zoning in land use planning, to reduce landslide fatalities and economic losses.

According to the Emergency Events Database (2011), Colombia is one of the South American countries with most landslides. In the period between 1901 and 2017, 45 landslide disasters were registered with 3,619 fatalities, 78,395 people affected, and economic losses of 2.4 million USD. In the Global Landslide Catalog, Colombia has 87 entries with a total of 464 deaths (Kirschbaum et al., 2015) from 2007 to 2013. The Latin America and Caribbean landslide database has compiled a record of 110 fatal landslides in Colombia between the years 2004 and 2013 with a total of 880 deaths, where Colombia is the country with the second highest number of fatal landslides in Latin America and Caribbean, only Brazil shows a slightly higher number of fatal landslides (119) (Sepúlveda and Petley, 2015). Mergili et al. (2015) expose that Colombia, according to EM-DAT, presents a victim/event ratio of 77.34 (3171/41) and is only exceeded by Peru and Ecuador in this respect. However, the EM-DAT database only considers events with high numbers of fatalities: actually, the real number of fatal landslides is much higher (Aristizábal and Gómez, 2007; Mergili et al., 2015).

Landslide susceptibility assessment can be determined by qualitative or by quantitative methods (Corominas et al., 2014). Qualitative methods correspond to knowledge-driven approaches based entirely on the judgment of experts using geomorphological criteria in the field (van Westen et al., 2000). Quantitative approaches are subdivided into data-driven methods and physically-based models. Statistical or data-driven methods evaluate the relationship between landslides and causative factors to predict the landslide spatial probability (Carrara, 1983; Gorsevski et al., 2000; Lee, 2005; Lee and Pradhan, 2007; Süzen and Doyuran, 2004). Physically based models for landslide susceptibility and hazard assessment of detailed areas include the interaction between hydrology, topography, soil

properties, and in some cases, vegetation in order to understand and predict the location and timing of landslide occurrence. Such models generally compute slope stability, using the Factor of Safety (*FoS*). *FoS* is given by the dimensionless ratio between the resisting forces and the driving forces (Lam and Fredlund, 1993). Most of the physically based models available in the literature build on the limit equilibrium concept and the assumption of a planar slope of infinite length with a potential failure surface parallel to the topographic surface (Chen and Chameau, 1983; Lam and Fredlund, 1993). However, the infinite slope stability approach is proposed only for shallow, planar sliding surfaces in friction-dominated soils and fails to capture the complexity of deep-seated landslide phenomena (Bishop, 1954; Carson & Kirkby, 1972; Crozier, 1986; Duncan & Wright, 2005).

Limit equilibrium models have been extended to three-dimensional (3D) failure surfaces: geometric shapes such as spheres or ellipsoids represent non-planar slip surfaces in a much better way and are important to consider in areas of complex lithological conditions or for soils with high cohesion values. The first 3D slope stability model was presented by Baligh and Azzouz (1975). Later, Chen and Chameau (1983) developed a method to analyse cohesive and frictional slopes with different pore water conditions. Dennhardt and Forster (1985) proposed a method using an ellipsoidal slip surface. Kalatehjari and Ali (2013) carried out a review of different 3D analysis models in which they exposed the fact that many of the methods considered the slope and slip surface as symmetrical shapes in order to determinate the static condition of equilibrium. Hovland (1997) presented a method for cohesive and frictional soils based on the Fellenius method (Fellenius, 1927): in this approach, the forces that act between columns are disregarded and *FoS* is determined by normal and shear forces that act at the bases of the columns (Lam and Fredlund, 1993).

Although *FoS* represents a quantitative – and seemingly objective – approach to evaluate slope stability, it has to be used carefully. Authors such as Kirsten (1983) mention that different values of *FoS* could be obtained from slopes with equal probability of failure. However, several software packages use this concept for 3D slope stability analysis, e.g. STAB3D (Baligh and Azzouz, 1975), 3D-PCSTABL (Thomaz, 1986), CLARA (O Hungr, 1988), and TSLOPE3 (Pyke, 1991). Most of these models include some limitations reducing the accuracy of *FoS* obtained (Stark, 2003). One of the most important limitations is that they were designed to analyses individual landslides or slopes; they are not appropriate for regional or catchment-scale slope stability analyses (Mergili et al., 2014b). A few 3D (or, strictly speaking, 2.5D) slope stability models in Geographic Information Systems (GIS) have been used for landslide susceptibility mapping (Carrara and Pike, 2008; Qiu et al., 2007; Xie et al., 2003). Recently, the r.slope.stability model, a C- and Python-based raster module in the open-source software GRASS GIS (GRASS Team, 2019) has been proposed. The r.slope.stability model considers the 2.5D geometry of the sliding surface for analyzing a number of randomly selected potential sliding surfaces that are ellipsoidal or truncated in shape (Mergili et al., 2014a, b), and also offers an implementation of the infinite slope stability model.

In terms of scale or spatial resolution, physically-based models are suggested to be applied to finer-scale study areas and their results are strongly influenced by the level of detail in the input data (Zizioli et al., 2013), whereas data-driven approaches are recommended for broader-scale landslide susceptibility analyses (van Westen et al., 2006; Corominas et al., 2014). In fact, the implementation of data-driven methods or physically-based models for the incorporation of landslide hazard mapping into land use planning has been regulated in several countries. In

110 Colombia, according to the Decree 1807/2014, the implementation of deterministic or probabilistic physically-based methods is obligatory in urban and urban expansion areas, whereas statistical and knowledge-driven models are only permitted for rural areas.

In order to meet this requirement, we have to obtain a more detailed knowledge on the suitability of the available physically-based landslide susceptibility models, with the final goal to use model ensembles in order to obtain a
115 broader, more robust picture of the landslide susceptibility conditions. We think that r.slope.stability could be a candidate to enrich such an ensemble, considering shallow planar and deep-seated ellipsoidal failure surfaces. Consequently, the scientific aims of the present study are (i) to evaluate the suitability of r.slope.stability for physically-based landslide susceptibility mapping in tropical mountainous terrain, and (ii) to identify its fit with other potentially suitable models, helping to learn about strengths, limitations, and uncertainties.

120 Following these aims, we present a probabilistic analysis of slope stability in GIS for modelling landslide susceptibility in a tropical and mountainous environment using the r.slope.stability model. This model is evaluated using a landslide inventory prepared after a major and destructive rainfall-triggered multi-landslide event in the La Arenosa catchment on 21 September 1990. A quantitative performance evaluation of the model by ROC analysis is carried out. The results are compared with those obtained through the corresponding deterministic analyses with
125 r.slope.stability, with SHALSTAB (Dietrich and Montgomery, 1998) and with SHIA_Landslide (Aristizábal et al., 2016), which represents a new model developed for tropical mountainous terrain.

2. Study area

The La Arenosa catchment, with an area of 9.9 km², is located on the north-western side of the Colombian Andes, at 1000–1900 m above sea level (Velásquez and Mejía, 1991; Aristizábal et al., 2016). The climate is tropical
130 humid with a mean annual temperature of 23°C and a mean annual precipitation of 4,300 mm. However, precipitation is highly variable between the different seasons, and between different years. The annual cycle of precipitation shows a bimodal period of rainfall (inter-annual scale) with rainfall peaks in the months of April (450 mm) and October (600 mm) (IDEAM, 2010). Rainfall often occurs in the afternoon and at night in the form of heavy rainstorms or cloudbursts of short duration (Aristizábal et al., 2015; Garcia, 1995). Most of the landslides in
135 the catchment are triggered by such events, even though it is often not the rainfall itself, but the response of the groundwater level – also depending on its initial state due to antecedent rainfall conditions – which is decisive for the occurrence of landslides (Fressard et al., 2016).

Although the natural vegetation of the La Arenosa catchment would correspond to very humid premontane forest, all the primary forest has been removed, and the lands are exclusively dedicated to agricultural use. In the highest
140 and steepest parts of the basin, the predominance of coffee crops, sugar cane, pastures and very small areas of secondary forest is maintained. This situation is considered a conditional cause of landslide occurrence.

Residual soils which have evolved from granodiorite rocks covered by slope deposits and fluvio-torrential deposits are characteristic for the area. Slope deposits cover approx. 15% of the catchment. Strong in situ weathering occurs due to chemical decomposition in the humid tropical climate (Velásquez and Mejía, 1991). Indicators of rapid,
145 extensive and progressive spheroidal decomposition of the granite are observed down to an average depth of 30 m.

The saprolite is fairly well graded. Its texture is described as sandy silt or silty sand with some gravel and a small fraction of clay. Relict joints in the parent rock are preserved in the saprolite. They facilitate preferential flow and therefore strongly influence the observed hydraulic conductivity of the surrounding soil matrix (INTEGRAL, 1990; Aristizábal et al., 2015). The matrix-supported slope deposits are formed by boulders of granite, residual soils, and vegetation debris (Aristizábal et al., 2015). Slope deposits generally accumulate at foot slopes or in gullies. Those usually poorly consolidated deposits are the consequence of past landslides. Their content in cobbles and boulders is high, and natural soil pipes are common (Velásquez and Mejía, 1991).

On 21 September 1990, the La Arenosa catchment was strongly affected by a rainfall event of high intensity and short duration. In less than 3 hours, 208 mm of precipitation, with a maximum intensity of $90 \text{ mm}\cdot\text{h}^{-1}$, was recorded within the study area, triggering approximately 800 landslides. Based on the intensity-duration-frequency (IDF) curve, a return period of 200 years was estimated for this event (Velásquez and Mejía, 1991).

The strong rainfall in the catchment, imposed upon a general saturation of the soils in a rugged topography, triggered a series of almost simultaneous landslides of rotational or translational type in the catchment (Hermelin et al., 1992; Garcia, 1995). However, there is no distribution map which permits the differentiation between translational and rotational landslides. The population was strongly affected: 20 fatalities were counted and 260 people had to be evacuated, 27 houses were destroyed and 30 damaged, and so were several bridges (Hermelin et al., 1992; Aristizábal et al., 2016). Estimates arrived at more than 6 million US\$ of total loss.

INTEGRAL (1990) with Velásquez and Mejía (1991) analysed a set of aerial images and conducted a detailed field survey to produce a detailed landslide inventory for the event. However, it was not feasible to generate a complete inventory of landslides for the entire catchment as aerial photographs and topographic maps were not available for an area of approx. 2 km^2 . Only the area covered by the landslide inventory was considered for this study, corresponding to an area of 7.6 km^2 .

A total of 699 landslides were identified and mapped in the La Arenosa catchment. According to Velásquez and Mejía (1991), most of the landslides started as shallow translational slides and transformed into debris flows, from very to extremely rapid with high water content (Fig. 1). The landslide bodies were small with respect to the flow length, and the slip surfaces were parallel to the slope surface. The majority of the landslides initiated within residual soils in hollows and open slopes with a slope steepness ranging from 35° to 42° (Velásquez and Mejía, 1991). The same authors described that the depth of failure surface was less than 3 m and corresponded to the contact of the residual soil with the saprolite.

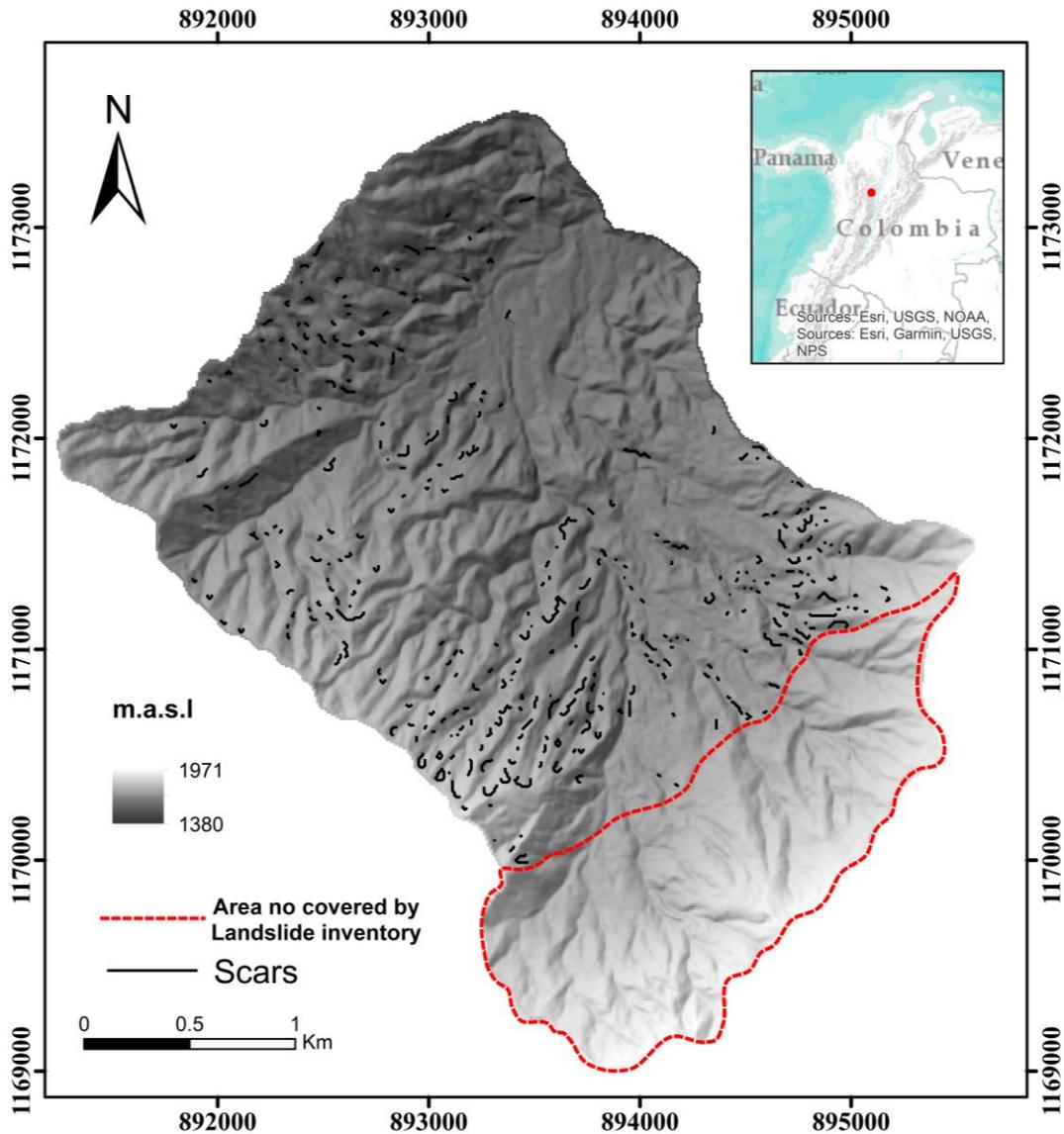


Figure 1. Landslides scars inventory database according to landslide inventory. The area with a red line does not have landslide inventory. The plot of a nonlinear surface. Source: Adapted from (INTEGRAL, 1990; Velásquez and Mejía, 1991)

180

3. The r.slope.stability model

The r.slope.stability model is a GIS-based, free, and open-source slope stability modelling software (www.slopestability.org) developed by Mergili et al. (2014a, b) as a C- and Python- based raster module of the GRASS GIS software package (GRASS Team, 2019). The tool is able to take into account planar failure with an infinite slope stability module and a slip surface module.

185

For shallow and planar landslides, r.slope.stability includes a classic infinite slope stability approach. For the infinite slope stability analysis, S acts parallel to the shear plane and seepage is considered parallel to the slope. For ellipsoid-shaped slip surfaces, in contrast, S is generally not parallel to the shear plane of the columns, even if it is parallel to the slope (Mergili et al., 2014b). The infinite slope stability model is run independently from the ellipsoidal failure surface analysis: FoS_{inf} for each raster cell is calculated according to Eq. (2).

$$FoS_{inf} = \frac{c' \cdot A + G' \cos \beta \cdot \tan \varphi'}{G' \cos \beta + S}, \quad (2)$$

where β is the slope angle of the slip surface (corresponding to the inclination of the terrain).

The slip surface model considers the 2.5D geometry of the sliding surface and evaluates FoS or the probability of slope failure (P_f) for many randomly selected potential ellipsoidal or truncated slip surfaces (Fig. 2). Each raster cell can be affected by various slip surfaces and is characterized by a unique value of FoS or P_f for each raster cell of the study area. Thereby, the lowest value of FoS or the highest value of P_f out of the values for all the sliding surfaces touching the cell is considered relevant. The model permits the users to impose restrictions with respect to the width, length, and depth of the ellipsoids (Mergili et al., 2014a, b).

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The slip surface model used in r.slope.stability represents a revision and extension of the 2.5D sliding surface model of Hovland (1997) (Xie et al., 2003). The calculation of FoS is based on the basic principle of equilibrium (Eq. 1).

$$FoS = \frac{\sum (c' \cdot A + (G' \cos \beta_c + N_s) \tan \varphi') \cos \beta_m}{\sum (G' \sin \beta_m + T_s) \cos \beta_m}, \quad (1)$$

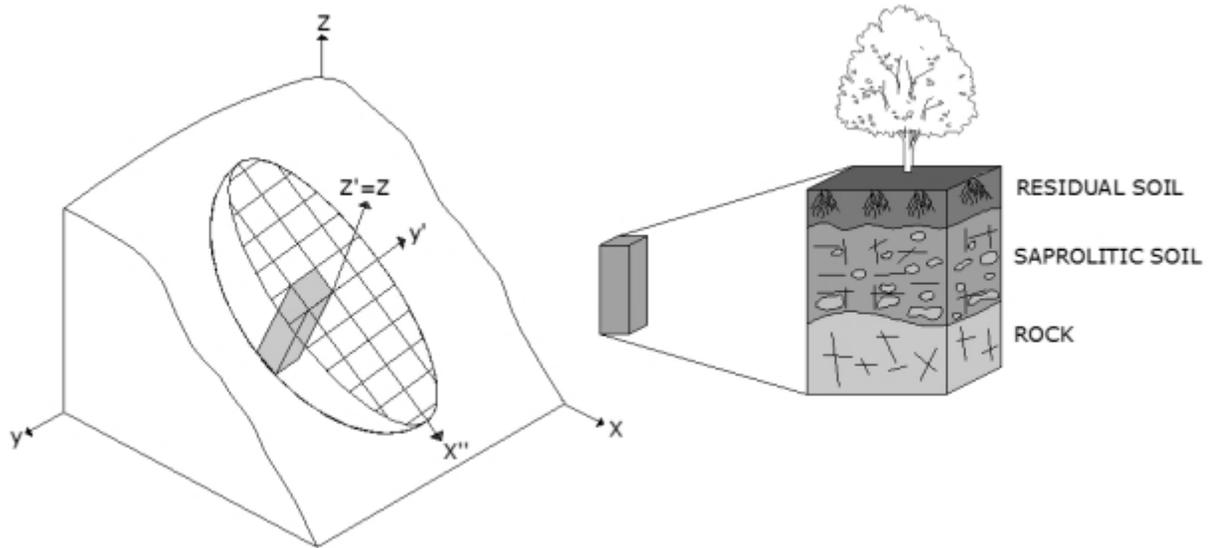
where c' is the effective cohesion (N m^{-2}), G' is the weight of the moist soil (N), β_c is the inclination of the slip surface, φ' is the effective internal friction angle, β_m is the apparent dip of the sliding surface in direction of the aspect α , N_s and T_s (N) are the contributions of the seepage force to the normal force and the shear force, respectively, and A (m^2) is the slip surface area assigned to each column. Inter-column forces and external forces (e.g seismic loading) are neglected (Mergili et al., 2014a, b). The slip surface model is further based on the model of King (1989), in which the direction of the seepage force (S) corresponds with the direction of the hydraulic gradient, approximated by slope and aspect of the groundwater table.

210

Both the infinite slope stability model and the slip surface model can be used for a probabilistic analysis, applying a range of geotechnical parameters (c' , φ') and the depth of the failure surface. **Thereby, rectangular, normal, log-normal, or exponential probability density functions are employed.** The result is a probability of failure (P_f), representing the fraction of tested parameter combinations yielding $FoS < 1$ (Mergili et al., 2014a).

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Note that, in the present work, with probabilistic model, we always refer to the random variation of the geotechnical parameters. **The slip surface model includes a probabilistic component where the dimensions of the ellipsoids are randomly varied for the computation of FoS .**



220 **Figure 2.** The typical ellipsoid used, slip surface with a model column in r.slope.stability, typical weathering profile of tropical environments and complex terrains. Source: Adapted from Aristizábal et al. (2016) and Qiu et al. (2007).

4. Data and procedure

225 The input of the r.slope.stability model consists of a Digital Terrain Model (DTM), spatial datasets of the mechanic and hydraulic characteristics of the study area, and finally the restraints imposed to the model by the user, based on the knowledge of the study area. A DTM with a spatial resolution of 10 x 10 m was provided by the Instituto Geografico Agustín Codazzi. A soil thickness map was built using interpolation, employing an empirical relationship between soil thickness and slope angle in the study area (Aristizábal, 2013; Catani et al., 2010; Thiery et al., 2019). The computed residual soil depth ranges from 1 m to 2.8 m.

230 The La Arenosa catchment is basically composed of two soil types, alluvial and residual, with properties strongly related to the parental material (IGAC, 2007). Alluvial soils cover 6.7% of the total area; they correspond to quaternary deposits composed of alluvial sediments of moderate depth, limited by the presence of fragments of rock and gravel. Residual soils cover 93.3% of the total area; they are derived from igneous rocks such as granites and quartz-diorites. The residual soils are medium to fine textured, well-drained, in some cases characterized by gravel or stones in the profile. These soils have deep weathering profiles depending on parent rock lithology and local conditions, partly reaching down to a depth of 100 m (Aristizábal et al., 2005; Suarez, 1998). The geotechnical parameters were obtained based on studies and laboratory tests carried out in La Arenosa by Velásquez and Mejía (1991). For the residual soils, cohesion values range between 5 kPa and 12.5 kPa, whereas the internal friction angle of the soil ranges from 16° to 24°. The dry unit weight ranges from 14.3 to 14.9 kN.m⁻³. No geotechnical laboratory tests are available for the alluvial deposits; however, they show very gentle slopes generally not prone to landsliding; the cohesion and friction angle values were therefore assumed based on literature values (Ameratunga et al., 2016; Aristizábal, 2013; Aristizábal et al., 2015; Geotechdata, 2013).

Table 1. Geotechnical parameters of La Arenosa catchment from laboratory tests. γ_d = specific weight dry of soil, c' =effective cohesion, φ' = effective angle of internal friction and θ_s =saturated water content minimum and maximum values for c' and φ' presented

	Alluvial soil	Residual soil
θ_s (%)	25	40
γ_d (kN/m ³)	17	14.9
c' (kN/m ²)	1 (0.6-1.4)	5 (3-7)
φ' (°)	34 (30-38)	24 (21-27)
Depth (m)	2.5-2.8	1.2-2.8

245 Adapted from (Aristizábal, 2013; Aristizábal et al., 2016)

The r.slope.stability model is applied with the probabilistic approach and deterministic approach for comparison with other models tested in the catchment. Both are used in combination with the infinite slope stability model and the slip surface model, resulting in a total of four model runs. Rectangular probability density functions for c' , φ' , and the soil depth d (m) are considered for the probabilistic analysis. The rectangular distribution is suitable for representing random variables which have known upper and lower bounds and an equal likelihood of occurring anywhere between these bounds (Fenton and Griffiths, 2008). For each parameter (c' , φ' , and d) we use a sample size of ten values, which are randomly selected from the ranges in Table 1.

250 With the infinite slope stability model, FoS for each raster cell is calculated with respect to the bottom of the soil. The analysis with truncated ellipsoidal failure surfaces is performed in a procedure with 1000 simulated surfaces touching each cell. Preliminary tests have indicated that this value represents a good compromise between computational efficiency and accuracy of the results. Each ellipsoidal slip surface is defined by the coordinates of the centre and the variable lengths of the three half-axes, the aspect, the inclination and the offset centre above the terrain, the failure surfaces include widths between 10 and 100 m and lengths between 10 and 200 m, with a maximum truncated depth of 2.8 m.

260 Quantitative evaluation of the empirical adequacy of the r.slope.stability model results are accomplished through a confusion matrix and an ROC analysis for the continuous output. Each raster cell of both the observed data (inventory) and the predicted data (model result) is assigned to one of two classes: stable (no landslide mapped or $FoS \geq 1$, respectively) or unstable (landslide mapped or $FoS < 1$, respectively) (Fawcett, 2005). The overlay of the classes is known as confusion matrix, where each raster cell is assigned to one of four classes: true positive (observed unstable cell is predicted as unstable), true negative (observed stable cell is predicted as stable), false positive (observed stable cell is predicted as unstable) and false negative (observed unstable cell is predicted as stable).

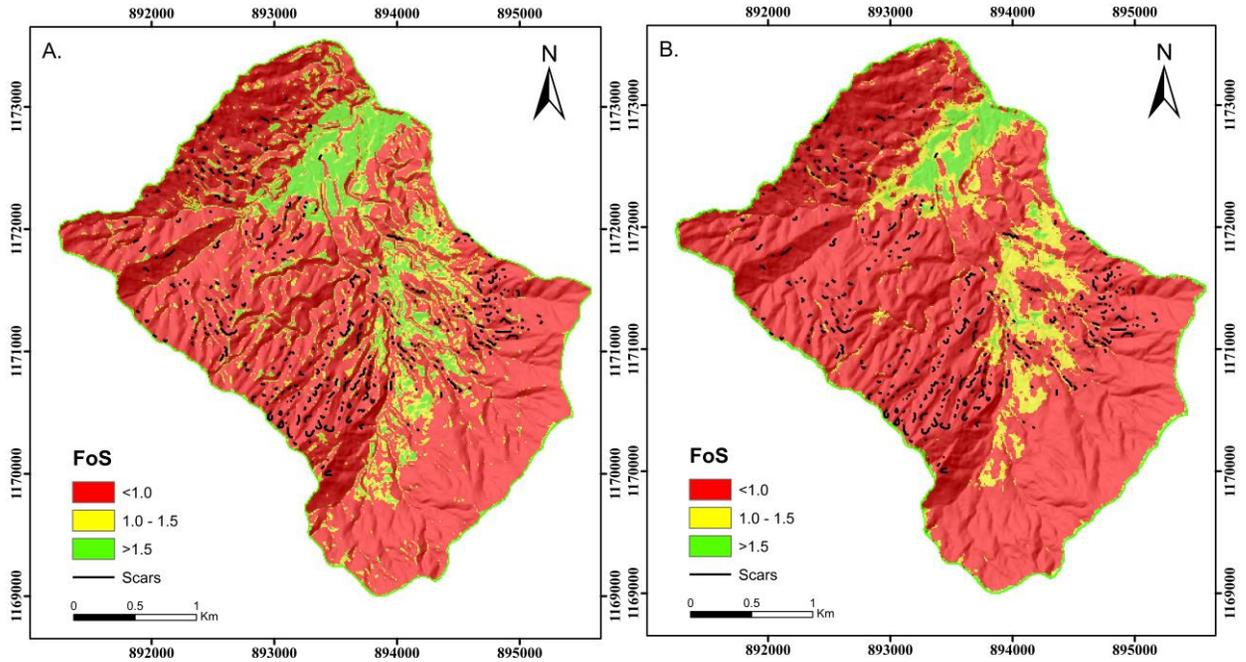
270 The statistical indexes measuring the performance (Table 4) are the true positive rate or hit rate (TPr) is defined as the ratio between the true positives and the observed positives. The true negative rate or specificity (TNr) is the ratio between the true negatives and the observed negatives. The false positive rate or false alarm rate (FPr) is

defined as the ratio between the false positives and the observed negatives, and the positive predictive value, also called the precision, is the ratio between the true positives and the total predicted positives (Aristizábal et al., 2015, 2016). Evaluation only considers the area covered by the landslide inventory. The ROC analysis plots TPr against FPr for various threshold levels of P_f .

r.slope.stability is compared with the SHALSTAB and SHIA_Landslide models to evaluate the consistency of the results. The SHALSTAB model, developed by Montgomery and Dietrich (1994), applies a topographic index to estimate the saturation of the soil as a function of rainfall infiltration. This procedure builds on the assumption that surface topography can be used as a main indicator of landslide susceptibility (Aristizábal et al., 2015). The model employs the hydrological model TOPOG which uses steady-state rainfall and an infinite slope approach for the geotechnical component (O'Loughlin, 1986). SHIA_Landslide is a physically based and conceptual model, developed by Aristizábal et al. (2016), for computing positive pore pressure changes as well as the resulting changes in FoS due to rainfall infiltration, coupling a distributed hydrological model with a classical infinite slope stability model.

5. Results

The results using the deterministic analysis with the infinite slope stability model in r.slope.stability are shown in Figure 3A, whereas the results obtained with the slip surface model are shown in Figure 3B, both of them in terms of FoS . Table 2 shows the confusion matrix calculated by comparing the deterministic analysis results with the scars in the landslide inventory map. For the infinite slope stability model, unstable conditions with $FoS < 1$ are shown for 79.2% of the catchment area, whereas only 10.5% show acceptably stable conditions with FoS values > 1.5 ; these areas correspond to the lower parts of the catchment formed by alluvial sediments with very gentle slopes. With regard to the slip surface model, 84% of the catchment area show $FoS < 1$, and only 5.8% show acceptable stability conditions with $FoS > 1.5$.



295

Figure 3. Landslide susceptibility classified according to *FoS*. Deterministic analyses of *r.slope.stability* (infinite slope stability model) (A) and *r.slope.stability* (ellipsoid-based model) (B).

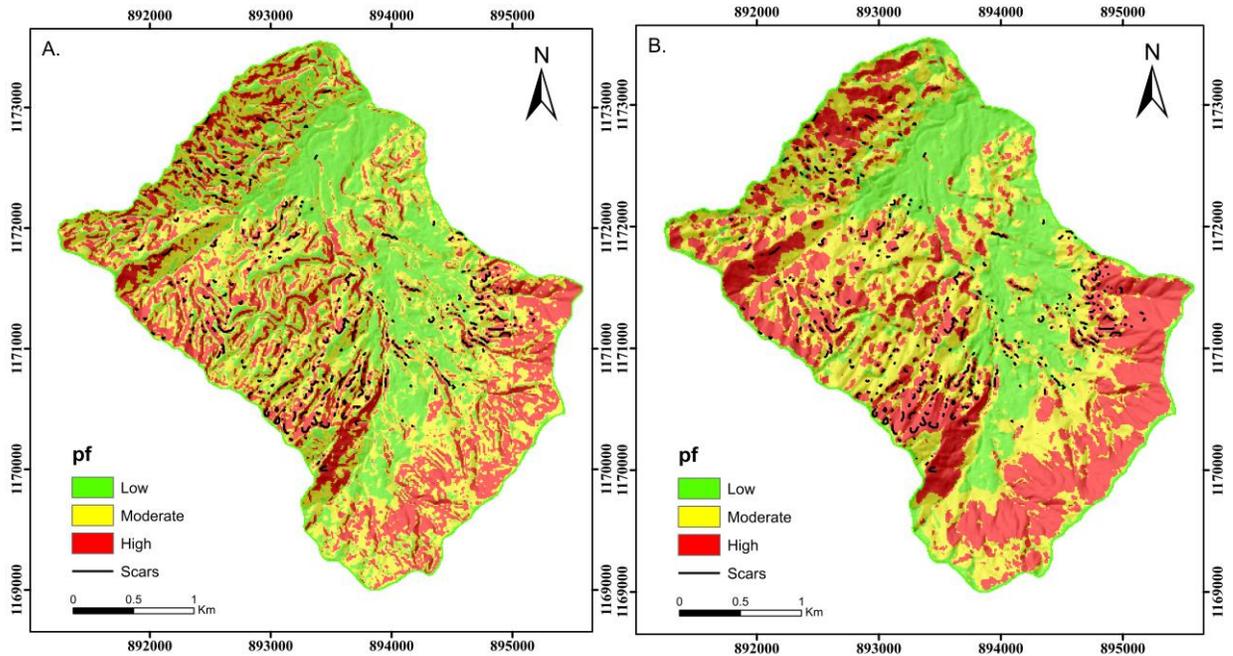
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Figure 4 illustrates the results of the probabilistic component of *r.slope.stability* used with the infinite slope stability model (Fig. 4A) and with the slip surface model (Fig. 4B) in terms of P_f . This probability is computed as the proportion of parameter combinations predicting $FoS < 1.0$ at a specific raster cell. Table 3 shows the confusion matrix calculated by comparing the probabilistic analysis results with the scarps in the landslide inventory map. To define the critical P_f thresholds, the distance to perfect classification parameter (r) proposed by Medina-cetina and Cepeda (2010) is used:

305

$$r = \sqrt{(FPr)^2 + (1 - TPr)^2} \quad (3)$$

The threshold values yielding the lowest value of r , indicating the best model performance, are used to discriminate between predicted positive and predicted negative cells. We are fully aware that this is a purely statistical optimization approach, not necessarily meaningful from a geotechnical point of view – an issue that will be further elaborated in the Discussion.



310

Figure 4. Probability of failure for r.slope.stability (infinite slope stability model) (A) and r.slope.stability (ellipsoid-based model) (B).

315

For the infinite slope stability model a minimum value of $r = 0.31$ is obtained for $P_f = 0.96$, whereas for the slip surface model a minimum value of $r = 0.46$ is obtained for $P_f = 0.99$. For the infinite slope stability model, unstable conditions with $P_f > 0.96$ are shown just for 30.5% of the catchment area, whereas 69.5% show stable conditions ($FoS \geq 1.0$). 36.4% of the catchment display values of $P_f > 0.99$ according to the slip surface model, whereas 63.6% display stable conditions. Unstable hillslopes, according to this criterion, are located mostly in the southern portion of the catchment, where no landslide inventory is available.

320

Table 2. The confusion matrix for the deterministic analysis

Classifier	r.slope.stability (ellipsoid-based model)			
	Pixel	Area (m ²)	Total percentage	Partial percentage
Observed landslide areas				
TP	2089	208900	2.7%	95%
FN	100	10000	0.1%	5%
Observed non-landslide areas				
TN	17314	1731400	22.7%	23%
FP	56838	5683800	74.5%	77%

Classifier	r.slope.stability (Infinite slope stability model)			
	Pixel	Area (m ²)	Total percentage	Partial percentage
Observed landslide areas				
TP	2154	215400	2.8%	98%
FN	35	3500	0.05%	2%
Observed non-landslide areas				
TN	17326	1732600	22.7%	23%
FP	56826	5682600	74.4%	77%

325 According to the confusion matrix, the deterministic analysis of r.slope.stability correctly predicts 98% and 95% of
the observed landslide areas with the infinite slope stability model and the slip surface model, respectively. The
other 2% and 5% are predicted as stable but did experience landslides according to the inventory. However, for the
observed non-landslide areas, only 23% are correctly predicted as stable by the infinite slope model and slip surface
model, whereas the other 77%, which are predicted as unstable, did not fail according to the inventory. The
330 deterministic model is more efficient in correctly classifying slopes where landslides occurred and less efficient at
classifying slopes on which landslides did not occur. With the thresholds of r applied in this study, the confusion
matrix for the probabilistic analysis shows a correct prediction of 83% and 65% of the observed landslide areas
with the infinite slope stability model and the slip surface model, respectively. The other 17% and 35% are
erroneously predicted as stable according to the inventory. For the observed non-landslide areas 74% and 70% are
335 correctly predicted as stable by the infinite slope model and slip surface model, respectively, whereas only 26% and
30% are predicted as unstable, but did not fail.

Table 3. The confusion matrix for the probabilistic analysis

Classifier	r.slope.stability (ellipsoid-based model)			
	Pixel	Area (m ²)	Total percentage	Partial percentage
Observed landslide areas				
TP	1416	141600	1.9%	65%
FN	773	77300	1.0%	35%
Observed non-landslide areas				
TN	51623	5162300	67.6%	70%
FP	22529	2252900	29.5%	30%

Classifier	r.slope.stability (Infinite slope stability model)			
	Pixel	Area (m ²)	Total percentage	Partial percentage
Observed landslide areas				
TP	1812	181200	2.4%	83%

FN	377	37700	0.5%	17%
Observed non-landslide areas				
TN	54890	5489000	71.9%	74%
FP	19262	1926200	25.2%	26%

340 The area under the ROC curve (AUC) values indicate a good ability of the probabilistic results to distinguish between susceptible and less susceptible areas (Fig. 5). The infinite slope stability model yields areas under the ROC curve of 0.82 and 0.83 for the deterministic and probabilistic analyses, respectively, whereas the slip surface model performs worse, but still fair (0.73 and 0.71).

345 **Table 4.** Statistical indexes measuring the performance of r.slope.stability and the other models

	Hit rate (0-100)	False alarm rate (0-100)	Specificity (0-100)	Precision
r.slope.stability (P_f ellipsoidal based model)	65	30	70	0.06
r.slope.stability (P_f infinite slope stability model)	83	26	74	0.08
r.slope.stability (FoS infinite slope stability model)	98	77	23	0.04
r.slope.stability (FoS ellipsoid-based model)	95	77	23	0.04
SHIA_Landslide	77	24	76	0.07
SHALSTAB	91	42	58	0.05

Adapted from (Aristizábal et al., 2015; Aristizábal et al., 2017)

350 Table 4 summarizes the statistical indices measuring the performance and prediction of the r.slope.stability model compared to SHALSTAB and SHIA_Landslide for the La Arenosa catchment. Figure 5 focuses on the performance of the area under curve ROC for the r.slope.stability model compared to SHIA_Landslide in the La Arenosa catchment, the better yield corresponds to the infinite slope stability model.

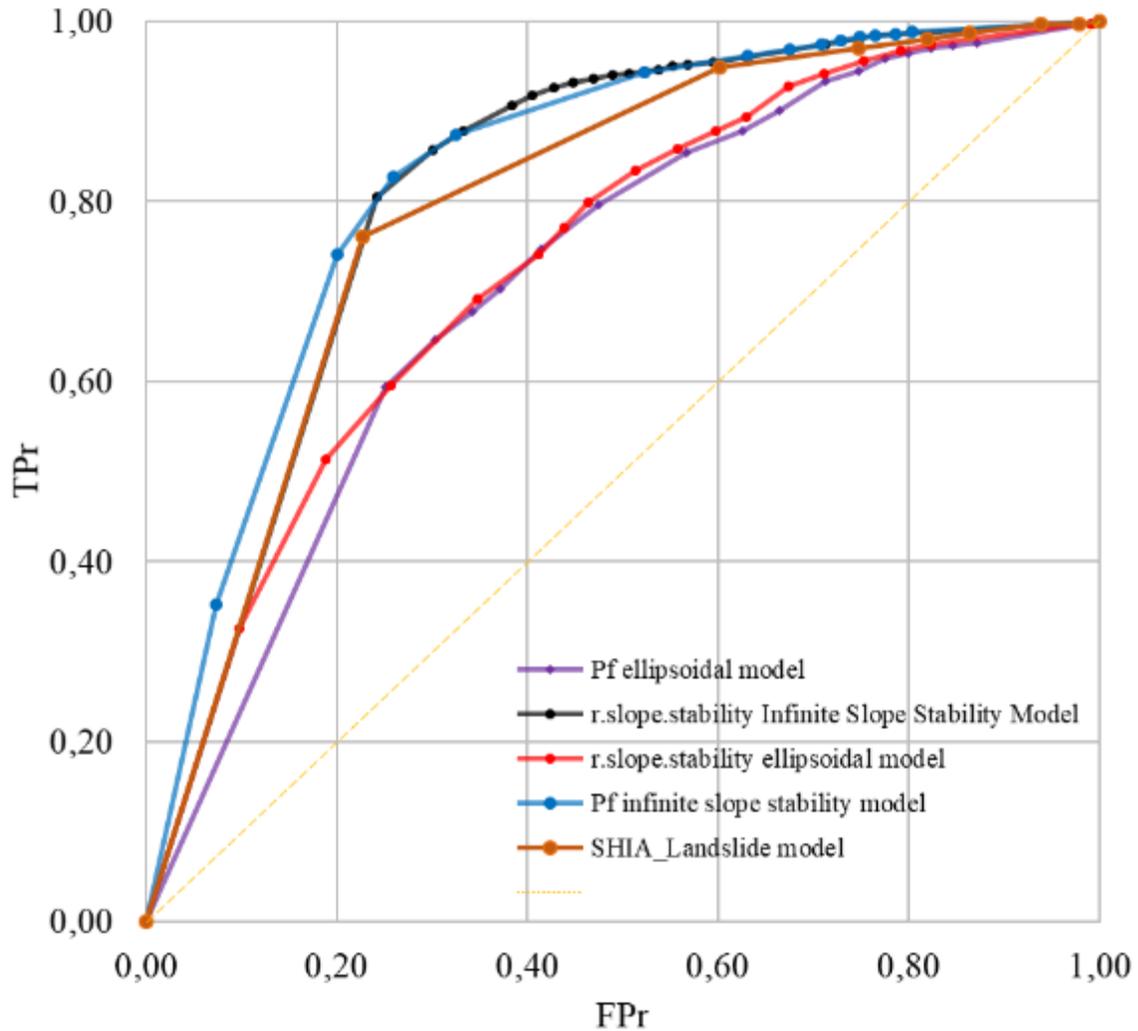


Figure 5. ROC curve models, r.slope.stability (infinite slope stability model) deterministic curve (0.83) and P_f (0.82) for r.slope.stability (ellipsoid-based model) deterministic curve (0.73) and P_f (0.71), SHIA_Landslide (0.77).

Figure 6 compares the results of r.slope.stability with SHALSTAB and SHIA_Landslide for a specific area of the catchment. SHALSTAB shows more areas classified as unconditionally unstable and unstable displaying similarities to the deterministic analysis of r.slope.stability, whereas SHIA_Landslide and the probabilistic analysis of r.slope.stability tend to show fewer areas with $FoS < 1$ or with high values of P_f . In summary, the deterministic analysis of the r.slope.stability model shows considerably higher hit rates, compared to the probabilistic analysis of r.slope.stability, SHALSTAB and SHIA_Landslide. However, both deterministic analyses of r.slope.stability show the highest false alarm rates. The major reasons for this result – and for the similar patterns yielded with SHIA_Landslide – are most likely an underestimate of the geotechnical stability of the material, the neglect of the effects of vegetation, and/or the overestimate of water saturation. Geotechnical testing is often performed on material which is disturbed in the one or the other way, so that the resulting parameters do not necessarily represent

the natural conditions over larger areas. Roots could lead to some degree of stabilization, and it could also be the vegetation retaining sufficient water to avoid full saturation of the soil throughout the catchment.

370 6. Discussion

In this work, the r.slope.stability tool was applied to the La Arenosa catchment, where the models SHALSTAB and SHIA_Landslide had been tested before (Aristizábal et al., 2015, 2016). An ideal model performance simultaneously maximizes *TPr* and minimizes *FPr*. In the La Arenosa catchment, the failure area associated with the rainfall event under investigation corresponds just to 2.2% of the whole catchment. Considering this situation, SHALSTAB tends to predict more unstable areas for this specific rainstorm, increasing the prediction capacity of the model but at the same time increasing the false positive error. Related to this aspect, Cervi et al., (2010) expose the limitations of SHALSTAB due to the hydrological assumptions (flow in steady-state conditions). By contrast, the SHIA_Landslide model shows a strong capacity for prediction and a very low *FPr*. In the case of r.slope.stability, the probabilistic analysis with the infinite slope stability model replicates the September 21, 1990 event with a very good reliability, in a much more successful way compared to SHALSTAB, and to the deterministic r.slope.stability approach; and a similar performance of SHIA_Landslide with a higher hit rate and a slightly higher value of *FPr*.

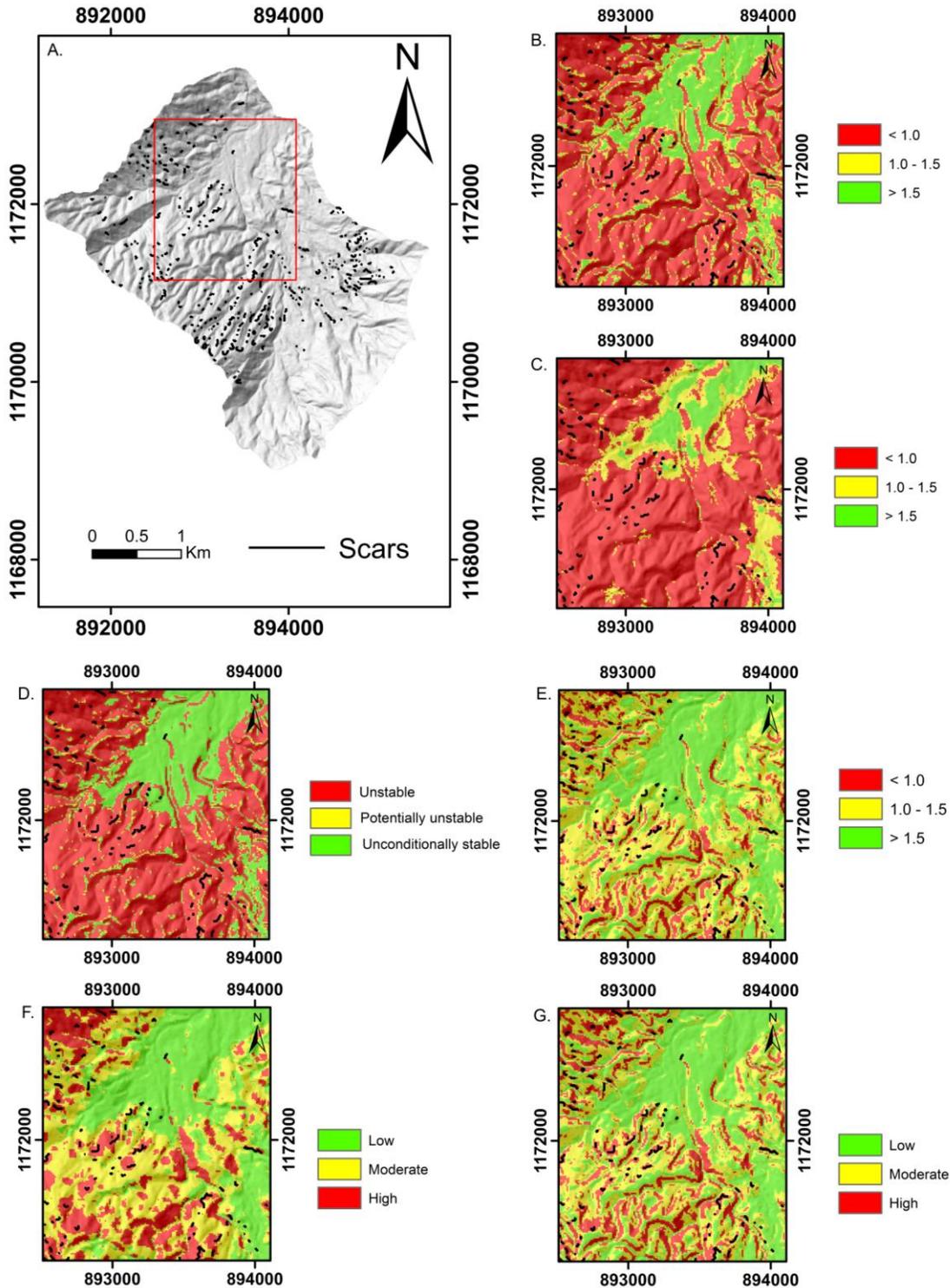
An important advantage of r.slope.stability compared to SHALSTAB and SHIA_Landslide is the possibility of carrying out a probabilistic analysis in terms of considering ranges of the key model parameters. Besides r.slope.stability, there are several physically-based models which provide a probabilistic module, such as TRIGRS (Baum et al., 2008) and ALICE (Thiery et al., 2017). Much more work has to be done in this direction. Measuring geotechnical and hydraulic parameters for large areas is difficult, time-consuming, and expensive, and there is an inherent variability in parameters associated with lithology and soil formation processes (Canli et al., 2017). Similarly, soil thickness shows a very high variability and uncertainty. This means there is a lot of uncertainty related to the horizontal and vertical natural variations of soil hydraulic and geotechnical parameters (Christian and Baecher, 2001). In general, soil properties show a pseudo-random pattern rather than a constant value. Additionally, landslides are more complex than their representation in the physically-based models adopted, and the geometrical and mechanical parameters that control slope stability are not known with sufficient accuracy (Griffiths et al., 2012; Guzzetti, 2016).

The conventional deterministic approach neglects uncertainties in the slope stability analysis. Although the *FoS* computation is more likely to identify areas prone to slope failure during a given hydro-meteorological event rather than to predict the exact locations of specific landslides (Baum et al., 2010), *FoS* is often not a reliable indicator of the slope stability conditions because it is – in terms of interpretation – a binary value derived from several uncertain parameters (Chowdhury, 2009). Thus, considering that physically based models are very sensitive to soil properties and soil depth, the probability distribution of failure is a much better indicator of the slope stability conditions.

Probabilistic analyses permit the inclusion of natural soil variations in the analysis, but also the mechanism of failure is fundamental for obtaining adequate results. In the case of the September 21, 1990 rainstorm, the infinite

slope stability analysis of the r.slope.stability model using planar shallow failure surfaces shows a much better performance. The results obtained show that the failure geometry is most appropriately approximated by shallow planar surfaces, which is confirmed by the reference information: Even though Hermelin et al. (1992) and Garcia (1995) mentioned that both shallow and deep-seated landslides were observed after the studied rainfall event, Velásquez and Mejía (1991) underlined the dominance of shallow landslides. This also agrees with the type of landslides often triggered by short, heavy rainfall events causing a rapid increase in pore pressure (Crosta, 1998). Such landslides are characterized by small and shallow, slope-parallel failure planes (depth of 0.3–2 m) (Anderson and Sitar, 1995). The displaced material, by means of processes of static liquefaction and rapid reduction of shear strength in undrained conditions, develops into flows that spread downward (Anderson and Sitar, 1995; Wang and Sassa, 2003). This type of landslide is described by Hungr et al. (2014) as debris avalanche. In cases where failures occur on slip surfaces with curved shapes, the assumption of ellipsoid-shaped slip surfaces would be expected to provide better results. This indicates the necessity of evaluating a priori the mechanism of failure and then employing the most appropriate model, or – if both mechanisms are relevant – evaluate the model results with different sub-datasets of the inventory. In the present study, due to the lacking distinction of landslide mechanisms in the inventory at the level of individual landslides, we applied both the infinite slope stability model and the slip surface model in order to cover a broad range of mechanisms. The result of the evaluation against the entire landslide inventory was in line with the generally reported dominance of shallow translational landslides.

The current version of r.slope.stability does not permit variation of the saturation level, meaning that the analysis has to be carried out in either dry or saturated conditions. In the r.slope.stability model, including the role of the infiltration process under saturated conditions in particular, could strongly improve the model performance in terms of being more effective in considering local hydrological conditions which govern slope instability processes (Mergili et al., 2014a, b), given that the required data are available.



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Figure 6. Comparison of results. (A) Analysis area (B) FoS obtained with r.slope.stability (infinite slope stability model) (C) FoS obtained with r.slope.stability (slip surface model) (D) Analysis with SHALSTAB (Aristizábal et al., 2015) (E) FoS obtained with SHIA_Landslide (Aristizábal et al., 2016) (F) P_f obtained with r.slope.stability (ellipsoid-based model) (G) P_f obtained with r.slope.stability (infinite slope stability model).

430 7. Conclusions

In this work, we have presented the results of r.slope.stability for the La Arenosa catchment in the Colombian Andes. r.slope.stability is a 2.5D slope stability model capable of dealing with shallow and deep-seated landslides triggered by rainfall. The model was evaluated with a set of observed landslides triggered by the 21 September 1990 rainstorm, for which different slope stability models had been previously applied. Considering the
435 probabilistic analyses performed, r.slope.stability shows a high hit rate, suggesting an acceptable prediction capacity for failure areas (83%–65%) for the infinite slope stability model and the slip surface model, respectively. The false alarm rate is relatively low for the infinite slope stability model (26%) and for the slip surface model (30%). The areas under the ROC curve yielded by the probabilistic approach are 0.83 for the infinite slope stability model and 0.71 for the slip surface model. These results clearly suggest a higher performance for the assumption of
440 shallow, planar failure surface (infinite slope stability) model than for the deep-seated slip surface model, a finding which is in line with the physical characteristics of the observed landslides. Despite the generally good model performance, the results were far too conservative, compared to the observations, meaning that either (1) the assumption of the saturation patterns was inappropriate; or (2) the geotechnical parameters fed into the model are not representative for the study area. The same challenges were identified for the SHALSTAB model. Future
445 studies shall further elaborate on this issue.

Compared to many other models, r.slope.stability has the advantage that it supports the derivation of a failure probability in terms of considering ranges of the key model parameters, instead of fixed values. Since in any landslide susceptibility analysis it is necessary to consider that soil parameters and their spatial variability are highly uncertain, the computation of failure probabilities in addition to *FoS* is highly recommended. In Colombia,
450 hazard mapping is mandatory for use land planning in urban areas. Results like the slope failure probability (Fig. 6F and Fig. 6G) are considered suitable for this purpose. However, they only represent one step on a long way:

- (i) Much of the landslide risk in the Arenosa catchment is not so much related to the failure of unstable slopes, but rather to the downslope propagation of the mobilized material as debris flows (Velásquez & Mejía, 1991). Therefore, coupling of the slope stability model results with mass flow simulation tools is absolutely
455 necessary (Mergili et al., 2017; Bout et al., 2018).
- (ii) Slope failure probabilities – or impact indicator indices (Mergili et al., 2017) – neither take into account the dimension of time nor the fine-scale patterns of the geotechnical characteristics of the catchment. Consequently, those results have to be interpreted in a relative, qualitative rather than an absolute, quantitative sense. The combination of these maps with expert judgement is therefore essential to define suitable
460 thresholds separating the study area into different zones with their individual hazard levels and recommendations for action.

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