

37 In this context early warning systems (EWS) play a key role. UNISDR (2009) defines
38 early warning systems as “the set of capacities needed to generate and disseminate timely
39 and meaningful warning information to enable individuals, communities and
40 organizations threatened by a hazard to prepare and to act appropriately and in sufficient
41 time to reduce the possibility of harm or loss”. A complete EWS is divided into four steps:
42 (1) risk knowledge, (2) monitoring, forecasting and warning, (3) communication of an
43 early warning system and (4) response capability (UN, 2006). The first two steps are
44 related to the field of physical sciences while the two last steps are associated to social
45 science aspects. There are several works related to the impact of the early warning system
46 in the prevention of floods. Baudoin et al. (2014) and UNISDR (2015) show some
47 interesting examples on how early warning systems can save lives and reduce the damage
48 to the people. Borga et al. (2011) developed an early warning system methodology for
49 flash floods in Europe through the HYDRATE project. The authors enhanced the
50 capability of flash flood forecasting in ungauged basins by exploiting the extended
51 availability of flash flood data and the improved process understanding. Alfieri et al.
52 (2012) analysed several early warning systems applied to detect surface water flooding,
53 flash floods, debris flows, land-slides induced by extreme rainfall events, river and coastal
54 floods. The authors proposed several tasks to palliate the main drawbacks of some of
55 these systems. Also, Hossain et al. (2014) developed a system to measure the water depth
56 of the river at the “Valley of Death” and Cools et al. (2012) developed an early warning
57 system to detect flash floods in the Sinai Peninsula, both based on a satellite-based
58 forecast system. In Europe a very interesting example of an early warning system is the
59 EWS applied to the region of Flanders (Schelfaut et al., 2012 and CIW, 2013). In this
60 work, the different steps are analysed under the FREEMAN project (Flood REsilience
61 Enhancement and MANagement). The European Flood Awareness System (EFAS) is
62 also another example of an EWS developed to the sponsorship of the European
63 Commission. This system provides daily streamflow forecast for Europe starting from up
64 to 10-days weather forecast (medium-term forecast). More details of this model can be
65 shown in Thielen et al. (2009), Pappenberger et al. (2011), Cloke et al. (2013) and Alfieri
66 et al. (2014). Using this model Dottori et al. (2017) develop a methodology to adapt EFAS
67 to real time forecasting. Demerit et al. (2013) analyse the problems derived from the use
68 of the early warning system to medium and long-term flood forecast, mainly the
69 dissemination of the information to people potentially affected by these events. They
70 reveal that flood forecasters usually wait the confirmation from local institutions

71 (Hydrologic Confederations...) instead of acting following the information provided by
72 the early warning systems. These local systems are focused in short-term forecast (0 to
73 48 h) that are more suitable to evacuation than fore damage mitigation. Some examples
74 of these short-term local systems focused on river floods are: the River Forecast Centers
75 (<https://water.weather.gov/ahps/rfc/rfc.php>) in the United States of America or “Sistema
76 de Ayuda a la Decisión”
77 (<http://www.chebro.es/contenido.visualizar.do?idContenido=12789&idMenu=2902>)
78 developed by the Hydrographic Confederation of the Ebro river (Spain). In Europe the
79 meteoalarm (http://www.meteoalarm.eu/?lang=en_UK) provides advice on exceptional
80 weather events including floods with a temporal window of 48 h. There are mainly two
81 kind of floods derived from precipitation events: flash-floods and river-floods. On the one
82 hand, flash-floods are characterised by a delay time, from the peak precipitation time to
83 the peak of flood, from 3 to 6 hours. These floods are usually registered in dry climate
84 and rocky terrain areas due to the lack of vegetation to infiltrate the precipitation into the
85 ground. These kind of floods have associated a very high level of risk due their velocity
86 of propagation. On the other hand, river-floods are generally registered in larger rivers in
87 areas with a wet climate and the delay time is greater than 6 hours. The consequences
88 associated to the latter ones can be also dramatic to the people and their properties. This
89 make necessary to develop an EWS to improve the security of the areas exposed to these
90 events. The area of study analysed in this work is mainly affected by river-floods.

91 In this paper, a flood early warning system based on precipitation forecast is presented.
92 The system, which is being developed in collaboration with the Hydrographic
93 Confederation of Miño-Sil River, consists of three steps: i) precipitation forecast; ii) use
94 of a hydrologic model to predict extreme flows; iii) use of a hydraulic model that is
95 applied at certain areas only under extreme flows. Starting from 1-day, 2-day and 3-day
96 precipitation forecast windows provided by the Regional Meteorological Office
97 (MeteoGalicia), the outflows associated to the catchment of the Miño River (NW Spain)
98 were obtained using the HEC-HMS model (U.S. Army Corps of Engineers, 2018). This
99 model was calibrated for the area of study by means of series of historical flood events
100 detected over the last decade. The numerical model Iber (Bladé et al., 2014) was used to
101 obtain water depth and velocity under extreme flow conditions for some risk areas where
102 previous events have caused damages or material loses. Both models (i.e., HEC-HMS
103 and Iber) are freely available software so the system can be applied at any location without

104 costs derived from the licences of commercial codes. The main contribution of the EWS
105 presented in this work respect to the systems shown in the bibliography is that all the
106 components are freely available and easily adaptable to different areas of the world.
107 The paper, which aims to describe the steps followed to develop the EWS, is organised
108 as follows. First, a description of the area of study (the upper reach of Miño River and the
109 city of Lugo, NW Spain) is shown. Then the methodology to obtain the weather forecast,
110 the computation of the run-off and the hydraulic processes are briefly presented. Also the
111 communication among all the models (Precipitation Forecast - Run-Off - Hydraulic
112 processes) is explained. Next, the results of the precipitation and outflow forecast of a
113 series of historical flood events are presented along with a statistic analysis of their
114 accuracy. Finally, the numerical water depth obtained for a particular flood event at the
115 city of Lugo is shown and compared with field data measured during the event.

116 **2. Study area**

117 The area of study is located in north-western Spain (Figure 1). It corresponds to the upper
118 reach of the Miño River. This sub-catchment area is about 2200 km² and the elevation
119 ranges from 360 to 980 m.a.s.l. The average annual precipitation ranges from 144 to 1300
120 mm year⁻¹. Miño River presents an annual hydrologic cycle characterised by a pluvial
121 regime, presenting maximum river discharges during winter months descending then to
122 reach its minimum values during summer (Fernández-Nóvoa et al., 2017). Specifically,
123 considering the period under study at Lugo station, Miño River reaches maximum flows
124 of 114 and 128 m³s⁻¹ in January and February and minimum ones of 7 and 8 m³s⁻¹ in
125 August and September, respectively.

126 Figure 1 (upper-left panel) shows the catchment of the upper reach of the Miño River,
127 which is divided into three main sub-basins according to their topographic characteristics.
128 Seven rain gauges operated by MeteoGalicia are located in the entire sub-catchment.
129 Table 1 shows the location and the elevation of each of the rain gauges located in the
130 upper reach of the Miño River. The outlet of this catchment is located in the city of Lugo
131 (Figure 1, lower panel). This area is usually flooded during the events of extreme
132 precipitations in the upper reach of the Miño River. The absence of dams in the catchment
133 to regulate the flow also affects the high frequency of these events.

134 **3. Methodology**

135 In this work, an automatic EWS is proposed. This system is composed of several elements
136 as shown in Figure 2. All these components are orchestrated by a Python script that is the
137 responsible of gather and transform the data properly in order to feed the models used in
138 the system. First of all, the rainfall forecast performed with the Weather Research and
139 Forecasting model is provided by the weather agency (MeteoGalicia). Details are
140 provided in next section. Forecasted data are automatically downloaded and the rainfall
141 relative to each sub-basin is extracted to feed the hydrological model HEC-HMS. When
142 the catchment outflow obtained with HEC-HMS surpasses the 90th percentile of historical
143 data, it is considered as a possible extreme event and the following steps will be applied.
144 In that case, this outflow will be used as inlet condition for the hydraulic simulation using
145 the model Iber to provide flood maps with water depths and velocities at certain risk areas
146 (the city of Lugo in this particular case). Data provided by Iber are processed for hazard
147 evaluation. At this stage the system checks if there is a risk condition in the areas
148 accessible by pedestrians. These areas are user defined and can be changed depending on
149 seasonal events. In order to emit a warning alert, the criteria of Cox et al. (2010) are used
150 to define safety limits for children since they are the most vulnerable population group.
151 Following this criterion, a warning will be emitted if there is a zone where any of the
152 following thresholds are surpassed: the water depth (h) is higher than 0.5 m, the
153 magnitude of water velocity (v) is higher than 0.2 ms^{-1} or the product ($h \cdot v$) excess $0.4 \text{ m}^2\text{s}^{-1}$.
154 This warning is sent in form of report to a decision maker, so an expert can validate the
155 resulting data and discard false positives.
156 The details of the components of the EWS, the data sources, and the calibration processes
157 are described in the following sections.

158 **3.1 Precipitation data**

159 **3.1.1 Forecasted precipitation data**

160 Forecasted precipitation data were obtained from the Regional Meteorological Office
161 (MeteoGalicia, <http://www.meteogalicia.gal/>). MeteoGalicia publishes weather forecast
162 results based on the Weather Research and Forecasting (WRF) Model (Skamarock et al.,
163 2005) (<https://www.mmm.ucar.edu/weather-research-and-forecasting-model>). The WRF
164 model is a numerical weather prediction system at regional mesoscale designed mainly
165 for forecasting applications. WRF is run operationally since 2008 providing daily data
166 until the end of 2012 (00 UTC) and twice a day (00 UTC and 12 UTC) from then on, with

167 a 72 hour forecast window, a temporal resolution of 1 hour and maximum spatial
168 resolution of 4 km (Sousa et al., 2013). Data provided by MeteoGalicia are freely
169 available at its THREDDS (Thematic Realtime Environmental Distributed Data Service)
170 server, also maintaining an historical archive of past forecast since 2008. The model
171 outputs provide several variables related to weather. In the case of this study, precipitation
172 information was automatically obtained for the areas under interest at the 00 UTC of each
173 day during the period 2008-2018.

174 **3.1.2 Measured precipitation data**

175 Real precipitation data at hourly scale were obtained from the rain gauges managed by
176 MeteoGalicia, which is responsible of their maintenance and data quality control. Data
177 from these rain gauges was used to assess the performance of the MeteoGalicia Weather
178 Forecast to predict extreme rain events. The mentioned rain gauges are pictured in Figure
179 1 and their location and elevation is detailed in Table 1.

180 **3.2 River discharge data**

181 Daily discharge data of the Miño river were provided by the corresponding river Basin
182 Authority (Confederación Hidrográfica del Miño-Sil, <https://www.chminosil.es>). In this
183 case of study, Miño flow data at Lugo station covering the period 2008-2018 were
184 selected. River data were used to calibrate and validate the hydrologic model system used
185 during the development of this study.

186 **3.3 HEC-HMS & Iber+**

187 Here the hydrological and hydraulic models used in the study will be briefly described
188 along with the methods to analyse their accuracy.

189 The semi-distributed model HEC-HMS (Feldman, 2000 and U.S. Army Corps of
190 Engineers, 2018) was used to analyse the rain-runoff processes and the numerical model
191 Iber (Bladé et al., 2014) was used to compute the hydraulic processes.

192 The HEC-HMS is a model developed by the US Army Corps of Engineers that is applied
193 to simulate continuous hydrological processes. The HEC-HMS model can be used to
194 analyse various hydrological aspects, such as flooding events, reservoir capacity,
195 stormwater warnings, and stream restoration (U.S. Army Corps of Engineers 2008).
196 HEC-HMS is divided into four components: (i) an analytical model: calculation of direct

197 runoff and channel routing; (ii) a basin model: representation of hydrological elements in
 198 a watershed; (iii) a system to manage input data and store data; (iv) a post-processing tool
 199 to report and illustrate simulation results. Two main processes were taken into account in
 200 the methodology developed in this case of study: loss (infiltration) and transform
 201 methods. In the first case, the Soil Conservation Service (SCS) curve number was
 202 selected. This method implements the curve number methodology for incremental losses,
 203 since it was designed to calculate the infiltration during periods of heavy rainfall, and
 204 therefore is well suited to this type of studies. Respect to the transform process, based on
 205 the way of convert the excess precipitation as runoff, the SCS unit hydrograph method
 206 was also selected for the reasons mentioned above. More information about the loss and
 207 transform methods used in this work are detailed in NRCS (2007). By last, the
 208 Muskingum-Cunge Routing method was selected for runoff propagation because it
 209 provides a good approach in basins with similar slopes. This method takes into account
 210 the conservation of mass as well as the diffusion representation of the conservation of
 211 momentum (U.S. Army Corps of Engineers 2008). Other parameters like the baseflow
 212 were not considered because suppose less than 3% of the peak flow for this kind of events
 213 and can be neglected.

214 Taylor diagrams (Taylor, 2001) were used to compute the accuracy of the results obtained
 215 with HEC-HMS by means of the normalised standard deviation (Eq. 1), normalised
 216 centred root-mean square difference (Eq. 2) and correlation (Eq. 3).

217

$$218 \quad \sigma_{n,A} = \frac{\sqrt{\frac{\sum_{i=1}^N (A_i - \bar{A})^2}{N}}}{\sigma_B} \quad (1)$$

219

$$220 \quad E_{n,A} = \frac{\sqrt{\frac{\sum_{i=1}^N [(A_i - \bar{A}) - (B_i - \bar{B})]^2}{N}}}{\sigma_B} \quad (2)$$

221

$$222 \quad R_A = \frac{\sum_{i=1}^N [(A_i - \bar{A})(B_i - \bar{B})]}{N \sigma_A \sigma_B} \quad (3)$$

223

224 where A is a numerical variable and B a reference variable. The subscript n refers to the
 225 normalised parameter, subscript i refers to the different samples, N is the number of
 226 samples, barred variables refer to mean values and σ is the standard deviation.

227 The hydraulic simulations were carried out using the numerical model Iber (Bladé et al.
 228 2014). Iber is a numerical code that solves the 2D (Two-Dimensional) Shallow Water
 229 Equations by means of finite volume schemes (FVS). The software package is formed by
 230 three elements: pre-processing tool, numerical model and post-processing tool. The first
 231 and the last modules are based in the software GID (GID, 2018). It provides a user
 232 friendly graphical interface (GUI) to create the case and edit the parameters that define
 233 the problem to solve. It also provides tools to analyse the results of the numerical
 234 simulations. The pre-processing and post-processing tools were used only during the
 235 modelling and testing of the study area. However, the automatic EWS runs the model in
 236 batch mode without user interaction. Iber was recently improved in terms of efficiency
 237 becoming Iber+ (García-Feal et al. 2018). This new parallel implementation of the Iber
 238 model takes advantage of GPU (Graphics Processing Unit) computing using the Nvidia
 239 CUDA (NVIDIA CO., 2019) platform. Using this technology, the new implementation is
 240 able to run up to 100 times faster. This fact makes Iber+ especially suitable for the
 241 implementation of an EWS where the response times can be crucial to issue an early alert.
 242 The accuracy of the water depth results computed with Iber+ at five control points was
 243 assessed by means of the *bias* and the *RMSE* (Root Mean Square Error) for the extreme
 244 event recorded on January 2013

$$245 \quad RMSE = \sqrt{\frac{\sum_{i=1}^N (A_i - B_i)^2}{N}} \quad (4)$$

$$246 \quad bias = \frac{\sum_{i=1}^N (A_i - B_i)}{N} \quad (5)$$

247 where A is numerical value, B the measured value and N the number of control points.

248 **4. Results and discussion**

249 **4.1 Accuracy of MeteoGalicia Precipitation Forecast**

250 The capability of MeteoGalicia Weather Forecast system to predict rain events was
 251 evaluated by means of the comparison with real precipitation data provided by the rain
 252 gauges in the area of study. For that purpose, the predicted (numerical) precipitation was
 253 obtained at the closest grid points to the location of the rain gauges. The correlation
 254 between predicted and measured precipitation was calculated for each rain gauge during
 255 the available period (2008-2018). For this calculation, Spearman rank correlation was
 256 used due to its robustness to deviations from linearity, as well as its strength to the

257 influence of outliers. This procedure was carried out for 3 forecast windows (1-24 h, 25-
258 48 h and 49-72 h; 1-day, 2-day and 3-day forecast from now on) to determine the accuracy
259 of the forecast at different temporal scales. The comparison is carried for an aggregation
260 time of 24 h, which matches the recording frequency of rain data provided by
261 MeteoGalicia and is compatible with the kind of flood events (mainly river floods) of the
262 area.

263 The values of the correlation and the normalised standard deviation for each rain gauge
264 are shown in tables 2 and 3. Table 2 shows the analysis for the complete series and table
265 3 shows the results considering only rainy events (precipitation above the 75 percentile).
266 In general, considering the complete series, precipitation prediction offers a good
267 representation of the registered values and the variability of precipitation. In fact,
268 correlations above 0.8 were obtained for the first two windows (1-day and 2-day forecast),
269 although with a higher correlation for the first one. The correlation is slightly lower for
270 the 3-day forecast, although it is still close to 0.8. When only rainy events are considered
271 mean correlation values are slightly lower than considering the complete series, although
272 showing a good representation of the registered data. It is specially remarkable the high
273 correlation showed under 1 day forecast window with a mean value above 0.7 (Table 3).
274 Respect to the normalised standard deviation, most of cases in both series are similar to
275 1, which shows a good agreement between forecast and real precipitation. Therefore, it
276 can be concluded that the precipitation forecast provided by MeteoGalicia offers results
277 very close to the real rain events for the entire time series of precipitation data (2008-
278 2018). This shows the accuracy of MeteoGalicia models to forecast precipitation events
279 up to three days in advance.

280 **4.2 Calibration and validation of hydrological processes using HEC-HMS**

281 A set of 15 extreme flood events registered during the period 2008-2018 were used to
282 calibrate and validate the rain-runoff model HEC-HMS (Table 4) by comparing the
283 outflows measured at the gauge station located at Lugo with the flows obtained with
284 HEC-HMS using the 1-day forecast of precipitation. Forecasted rain data were considered
285 because they are used to feed the model in its forecast version. In situ data would be only
286 valid for hindcast purposes. Calibration was carried out using the specific calibration tools
287 implemented in HEC-HMS (Feldman, 2000) in order to choose two independent
288 parameters, the curve number (CN) and lag time (L_g), for each sub-basin. The values of
289 CN and L_g were computed using particle swarm algorithm (Kennedy and Eberhart, 1995,

290 Pedersen, 2010 and Mezura-Montes and Coello, 2011) to minimise the error between the
291 measured streamflow and the numerical one. No empirical formulas were used for CN
292 and L_g due the uncertainty associated to their definition (Fang et al., 2008; Upegui and
293 Gutierrez, 2011; Grimaldi et al., 2012). Eleven flood events were used for calibration
294 purposes and the rest of cases were used to validate the model. Table 5 shows the values
295 of the CN and L_g for each sub-basin obtained for each event used in the calibration step.
296 The mean values of CN and L_g of each sub-basin were used to validate the model in four
297 flood events (01/2013, 01/2014, 02/2016 and 03/2018) by means of a Taylor diagram
298 (Figure 3).

299 The values of normalised standard deviation (σ_n) range from to 0.8 to 1.2, the values of
300 the root mean squared difference (RMSD) range from 0.3 to 0.6 and the correlation of the
301 numerical results range from to 0.85 to 0.95. The values of σ_n means that the variability
302 of the numerical results are quite similar to the variability of the reference time series
303 (difference less than the 20%) and the values of E_n can be considered as good values
304 according to Moriasi et al. (2007). These values of σ_n , E_n and correlation show that the
305 mean values of CN and L_g obtained in the calibration step characterise the behaviour of
306 the basin with a high accuracy.

307 Figure 4 compares the numerical and measured streamflow for the event that happened
308 in January 2013 using the three forecast windows. The left panel shows that time series
309 of the flows predicted by the model are similar to those measured at the gauge station.
310 The right panel is the Taylor diagram corresponding to the three forecast windows. The
311 standard deviation is observed to range from 0.8 to 1.2 for the three forecasts. RMSD
312 values for 1-day and 2-day forecasts are around 0.3, being around 0.6 for the 3-day
313 forecast. Finally, the correlation coefficient for 1-day and 2-day forecasts are close to
314 0.95, being around 0.85 for the 3-day forecast.

315 **4.3 Case of study**

316 Once the predicted water flow showed to reproduce the real events with a high accuracy
317 ($E_n \sim 0.8$, $\sigma_n \sim 0.3$ and $R \sim 0.95$), the water depth and velocity during the flood event that
318 affected Lugo on 20th January, 2013 were computed using the numerical code Iber+
319 (Garcia-Feal et al., 2018). Figure 5 shows the numerical domain at Lugo, where seven
320 land uses were defined to model the characteristics of the terrain. The Manning's
321 coefficient associated to each land use are shown in Table 6. Figure 5 also shows the

322 location of the inlet and outlet boundary conditions. The initial water depth was obtained
323 from data provided by the gauge station located at Lugo. The inlet condition was defined
324 by means of the input hydrograph (Critical/Subcritical) and the outlet condition was
325 defined using a supercritical/critical outflow. Turbulence was not taken into account as
326 suggested by (SNCZI, 2011) and according with similar works (Ercicum et al., 2010; Liu
327 et al. 2013; Segura-Beltrán et al., 2016).

328 The topography of the area of study was obtained from raster files freely downloaded
329 from the Instituto Geográfico Nacional website (<https://www.ign.es/web/ign/portal>). The
330 computational domain was discretised using a mesh with near 200,000 unstructured
331 triangular elements, with an average area of 2 m².

332 Five control points were defined at the area of study (see Figure 6) to analyse the accuracy
333 of the numerical results. Points from 1 to 4 are located in places next to the riverbank
334 usually frequented by pedestrians while the last one is located in the riverbed. Therefore,
335 the first four points are of special interest to issue an alert.

336 Figure 7 shows the values of the water depth obtained in the numerical simulations along
337 with the water depth obtained at the control points during the flood event. These field
338 values were obtained from photographs provided by volunteers and local media and taken
339 within the interval 12:00 – 16:00 on January 20th. The numerical water depth is expressed
340 in terms of a mean value and a range that corresponds to 3 times the standard deviation
341 of the values within that interval. Visually, the numerical results are quite similar to the
342 field data when considering the 1-day forecast, especially if one considers that the
343 accumulation of the small inaccuracies of the three models involved can give rise to
344 biases. The values are slightly less accurate when considering the 2-day forecast and
345 worse for the 3-day forecast due to lower accuracy in rainfall forecast. Finally, it must be
346 mentioned that the depicted values do not correspond to the peak flow that took place on
347 21th January, 2013 (at approx. 4:00).

348 Apart from the visual comparison, the accuracy of the model to calculate water elevation
349 was analysed in terms of two estimators (*RMSE* and *bias*) computed using the three
350 forecast windows. The minimum values of *RMSE* and *bias* are obtained with the 24h
351 forecast window (21 cm and 0 cm, respectively). The *RMSE* is satisfactory when
352 compared with the mean upward displacement of water during the event, which is about
353 2.5 m. In addition, the *bias* is null, showing that the model (in average) neither
354 overestimate nor underestimate real water elevation. The accuracy decreases with the
355 forecast window, although results are still good for a 2-day forecast (*RMSE*= 28 cm and

356 bias =4 cm). Finally, the accuracy is acceptable for a 3-day forecast (RMSE= 41 cm and
357 bias =-35 cm), although with limitations in terms of bias, since the model clearly tends to
358 underestimate field measurements. In summary, the agreement between measured and
359 computed values indicates that the system can be used to issue alert up to 3 days in
360 advance.

361 Figure 8 shows the maximum water depth and maximum velocity obtained for 1-day
362 forecast. Hazard maps (Figure 9) can be computed from these data according to the
363 criterion of Cox et al. (2010). Several recreation areas near the riverbanks show to have
364 surpassed the aforementioned hazard threshold. Therefore, decision-makers can use the
365 map to restrict activities in these areas, in order to mitigate the consequences of floods.

366 **5. Conclusions**

367 In this paper an Early Warning System for flood prediction using precipitation forecast
368 was presented. This system starts automatically using rain forecast data retrieved from
369 Regional Meteorological Office (MeteoGalicia) and concatenates two freely available
370 software packages (HEC-HMS and Iber+). The upper reach of the Miño River (NW
371 Spain) and, in particular, the city of Lugo were used as a benchmark.

372 A Python script was developed to deal with all the components involved in the system
373 without user interaction. First, the precipitation forecast provided by MeteoGalicia is
374 automatically obtained for the area of study. Second, rain forecast is provided to HEC-
375 HMS as an input to compute the streamflow in the catchment area. When the streamflow
376 obtained with HEC-HMS surpasses the 90th of the historical percentile at some
377 previously selected risk area (the city of Lugo in this particular case), the possibility of
378 an extreme event is detected and that streamflow is automatically defined as an inlet
379 condition for Iber+. Finally, data obtained from Iber+ are processed for risk assessment
380 and, if applicable, decision makers are reported.

381 The accuracy of the different models was assessed to analyse the capability of the system
382 to provide reliable results. First, the accuracy of the precipitation forecast provided by
383 MeteoGalicia was analysed for the period 2008-2018 showing that the 1-day forecast is
384 slightly more accurate than the 2-day forecast, being the 3-day slightly worse, although
385 the three forecast windows showed a reasonable agreement with field data. As a second
386 step, the accuracy of HEC-HMS to reproduce extreme flows was assessed by means
387 fifteen flood events recorded over for the period 2008-2018. Taylor diagrams were used

388 to compute the accuracy of the numerical streamflow compared with field data obtained
389 at the control station located near Lugo. Once again, results were satisfactory for the three
390 forecast windows, especially for the 1-day and 2-day forecast. Finally, a historical flood
391 event recorded in January, 2013 was used to assess the accuracy of Iber+ to reproduce
392 real water elevation at 5 control points located at the riverbank and riverbed. Both the
393 *RMSE* and the *bias* between the measured and computed elevations were satisfactory,
394 especially for the 1-day forecast.

395 The system needs less than 1 hour to run the models for a 3-day forecast horizon. While
396 data can be downloaded in a few seconds and the hydrologic model can be run in less
397 than a minute, no matter the extent of the area, the real bottleneck in the system is the
398 hydraulic model. Fortunately, the execution time does not necessarily increase with the
399 number of risk areas since different areas can be run concurrently when the available
400 hardware resources allow it. Taking into account that meteorological data are available
401 every day at 5:00 a.m. the system can provide an alert report to decision makers before
402 6:00 a.m. Additional improvements can be applied without additional cost in term of
403 runtime. For example, an ensemble approach can be applied when rain forecasts from
404 different sources are used as an input condition for HEC-HMS, in such a way that Iber+
405 is only executed when at least one of the hydrological realizations indicates a possible
406 extreme event.

407 Additional research is still needed to cover the entire Miño river basin, where other
408 problems may arise from the presence of dams. The system, when fully developed, can
409 even help to manage dams intelligently, maximizing energy production and dampening
410 floods at the same time.

411 The Early Warning System can be easily adapted for any area of the world since the
412 required input data can be obtained freely from public institutions and the models to
413 compute the hydrological and the hydraulic processes (HEC-HMS and Iber+,
414 respectively) are both freely available. Therefore, the EWS is especially interesting for
415 developing countries where the acquisition of commercial software is not sustainable.

416

417

418

419 *Code and data availability.* Freely available data and software (HEC-HMS and Iber+)
420 were used for this work. The detailed processing flowchart is shown in Fig. 2 (Section 3
421 – Methodology).

422

423 *Author contributions:* JGC, OGF and DFN designed the research, conducted the analysis
424 and wrote the paper, ; JMDA and MGG supervised the research and revised the paper.

425

426 *Competing interests.* The authors declare that they have no conflict of interest.

427

428

429 **Acknowledgements**

430 This work was partially supported by Water JPI-WaterWorks Programme under project
431 Improving Drought and Flood Early Warning, Forecasting and Mitigation
432 (IMDROFLOOD, code: PCIN-2015-243), and by Xunta de Galicia under Project
433 ED431C 2017/64-GRC “Programa de Consolidación e Estructuración de Unidades de
434 Investigación Competitivas (Grupos de Referencia Competitiva). We especially thank
435 Carlos Ruiz del Portal Florido, Head of the Hydrological Planning Office, Hydrographic
436 Confederation of Miño-Sil River for helpful discussions and for providing access to real
437 data within the context of INTERREG-POCTEP Programme project RISC_ML (Code:
438 0034_RISC_ML_6_E).

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441 **References**

- 442 Alfieri, L., Salamon, P., Pappenberger, F., Wetterhall, F., and Thielen, J.: Operational
443 early warning systems for water-related hazards in Europe, *Environ. Sci. Policy*,
444 21, 35-49, <https://doi.org/10.1016/j.envsci.2012.01.008>, 2012.
- 445 Alfieri, L., Pappenberger, F., Wetterhall, F., Haiden, T., Richardson, D., and Salamon,
446 P.: Evaluation of ensemble streamflow predictions in Europe, *J. Hydrol.*, 517,
447 913–922, <https://doi.org/10.1016/j.hydrol.2014.06.035>, 2014.
- 448 Alfieri, L., Bisselink, B., Dottori, F., Naumann, G., de Roo, A., Salamon, P., Wyser, K.,
449 and Feyen, L.: Global projections of river flood risk in a warmer world, *Earth's*
450 *Future*, 5 (2), 171-182., <https://doi.org/10.1002/2016EF000485>, 2017.
- 451 Baudoin, M., Henly-Shepard, S., Fernando, N., Sitati, A., and Zommers, Z.: Early
452 warning systems and livelihood resilience: exploring opportunities for community
453 participation, UNU-EHS Working Paper Series, No.1, United Nations University
454 Institute of Environment and Human Security (UNU-EHS), Bonn, 2014.
- 455 Bladé, E., Cea, L., Corestein, G., Escolano, E., Puertas, J., Vázquez-Cendón, E., Dolz, J.,
456 and Coll, A.: Iber -River modelling simulation tool [Iber: herramienta de
457 simulación numérica del flujo en ríos], *Revista Internacional de Metodos*
458 *Numericos para Calculo y Diseno en Ingenieria*, 30 (1), 1-10.
459 <https://doi.org/10.1016/j.rimni.2012.07.004>, 2014.
- 460 Borga, M., Anagnostou, E. N., Blöschl, G., and Creutin, J. D.: Flash flood forecasting,
461 warning and risk management: The HYDRATE project, *Environ. Sci. Policy*, 14
462 (7), 834-844, <https://doi.org/10.1016/j.envsci.2011.05.017>, 2011.
- 463 CIW, Rapport Globale Evaluatie Overstromingen, Report in Dutch [Evaluation Report
464 for Floods] Committee for Integrated Water Resources Management (CIW),
465 Flemish Authority, Brussels, Belgium, 2013.
- 466 Cloke, H., Pappenberger, F., Thielen, J., and Thiemiig, V.: Operational European flood
467 forecasting, in: *Environmental Modelling: Finding Simplicity in Complexity*, 2nd
468 ed., Wainwright, J. and Mulligan, M. (Eds.), John Wiley and Sons, Ltd,
469 Chichester, UK. <https://doi.org/10.1002/9781118351475.ch25>, 2013.
- 470 Cools, J., Vanderkimpen, P., El Afandi, G., Abdelkhalek, A., Fockede, S., El Sammany,
471 M., Abdallah, G., El Bihery, M., Bauwens, W., and Huygens, M.: An early
472 warning system for flash floods in hyper-arid Egypt, *Nat. Hazard Earth Syst. Sci.*,
473 12 (2), 443-457, <https://doi.org/10.5194/nhess-12-443-2012>, 2012.

474 Cox, R. J., Shand, T. D., and Blacka, M.J.: Australian Rainfall and Runoff revision project
475 10: appropriate safety criteria for people, *Water Res.*, 978, 085825-9454, 2010.

476 Dankers, R., and Feyen, L.: Climate change impact on flood hazard in Europe: An
477 assessment based on high-resolution climate simulations”. *J. Geophys. Res-*
478 *Atmos.*, 113 (19), <https://doi.org/10.1029/2007JD009719>, 2008.

479 Demeritt, D., Nobert, S., Cloke, H. L., and Pappenberger, F.: The European Flood Alert
480 System and the communication, perception, and use of ensemble predictions for
481 operational flood risk management, *Hydrol. Process.*, 27 (1), 147-157,
482 <https://doi.org/10.1002/hyp.9419>, 2013.

483 Dottori, F., Kalas, M., Salamon, P., Bianchi, A., Alfieri, L., and Feyen, L.: An operational
484 procedure for rapid flood risk assessment in Europe, *Nat. Hazards Earth Syst. Sci.*,
485 17 (7), 1111-1126. <https://doi.org/10.5194/nhess-17-1111-2017>, 2017.

486 Erpicum, S., Dewals, B., Archambeau, P., Detrembleur, S., Piroton, M.: Detailed
487 inundation modelling using high resolution DEMs, *Eng. Appl. Comp. Fluid.*, 4(2),
488 106-208, <https://doi.org/10.1080/19942060.2010.11015310>, 2010.

489 Feldman, A.D.: Hydrologic Modeling System HEC-HMS, Technical Reference Manual,
490 p. 157. Institute for Water Resources Davis, USA, 2000.

491 Fang, X., Thompson, D.B., Cleveland, T.G., Pradhan, P., Malla, R.: Time of
492 concentration estimated using watershed parameters determined by automated
493 and manual methods, 134. 202-211.
494 [https://doi.org/10.5194/10.1061/\(ASCE\)0733-9437\(2008\)134:2\(202\)](https://doi.org/10.5194/10.1061/(ASCE)0733-9437(2008)134:2(202)).

495 Fernández-Nóvoa, D., deCastro, M., Des, M., Costoya, X., Mendes, R., Gómez-Gesteira,
496 M.: Characterization of Iberian turbid plumes by means of synoptic patterns
497 obtained through MODIS imagery, *J. Sea Res.* 126, 12-25.
498 <https://doi.org/10.1016/j.seares.2017.06.013>.

499 García-Feal, O., González-Cao, J., Gómez-Gesteira, M., Cea, L., Domínguez, J.M.,
500 Formella, A.: An accelerated tool for flood modelling based on Iber, *Water*, 10
501 (10), art. no. 1459, <https://doi.org/10.3390/w10101459>, 2018.

502 *GID Reference Manual*. <https://www.gidhome.com/>, 2018.

503 Grimaldi, S., Petroselli, A., Tauro, F., Porfiri, M.: Time of concentration: a paradox in
504 modern hydrology, *Hydrol. Sci. J.* 57, 217-228.
505 <https://doi.org/10.1080/02626667.2011.644244>.

506 Hossain, F., Siddique-E-Akbor, A. H. M., Yigzaw, W., Shah-Newaz, S., Hossain, M.,
507 Mazumder, L. C., Ahmed, T., Shum, C. K., Lee, H., Biancamaria, S., Turk, F. J.,

508 and Limaye, A.: Crossing the "valley of Death": Lessons learned from
509 implementing an operational satellite-based flood forecasting system, *B.*
510 *Am.Meteorol. Soc.*, 95 (8), 1201-1207, [https://doi.org/10.1175/BAMS-D-13-](https://doi.org/10.1175/BAMS-D-13-00176.1)
511 00176.1, 2014.

512 Jonkman, S. N.: Global perspectives on loss of human life caused by floods, *Nat. Hazards*,
513 34 (2), 151-175, <https://doi.org/10.1007/s11069-004-8891-3>, 2005.

514 Kennedy, J., Eberhart, R. Particle swarm optimization. *IEEE International Conference on*
515 *Neural Networks - Conference Proceedings*, 4, pp. 1942-1948, 1995.

516 Liu, Y., Zhou, J., Song, L., Zou, Q., Liao, L., Wang, W.: Numerical modelling of free-
517 surface shallow flows over irregular topography with complex geometry, *Appl.*
518 *Math. Model.*, 37(23), 9482-9498, <https://doi.org/10.1016/j.apm.2013.05.001>,
519 2013.

520 Mezura-Montes, E., Coello Coello, C.A. Constraint-handling in nature-inspired
521 numerical optimization: Past, present and future. *Swarm and Evolutionary*
522 *Computation*, 1 (4), pp. 173-194. 2011.

523 Moriasi, D.N., Arnold, J.G., Van Liew, M.W., Bingner, R.L., Harmel, R.D., Veith, T.L.:
524 Model evaluation guidelines for systematic quantification of accuracy in
525 watershed simulations. *Transactions of the ASABE*, 50 (3), 885-900.
526 [https://www2.scopus.com/inward/record.uri?eid=2-s2.0-](https://www2.scopus.com/inward/record.uri?eid=2-s2.0-34447500396&partnerID=40&md5=50b5724614f28257edef46d43db96018)
527 34447500396&partnerID=40&md5=50b5724614f28257edef46d43db96018.

528 Natural Resources Conservation Service (NCRS), 2007. *National Engineering*
529 *Handbook: Chapter 16 Hydrographs*. Washington, DC.

530 Noji, E. K.: The public health consequences of disasters, *Prehospital and disaster*
531 *medicine : the official journal of the National Association of EMS Physicians and*
532 *the World Association for Emergency and Disaster Medicine in association with*
533 *the Acute Care Foundation*, 15 (4), pp. 147-157, 2000.

534 NVIDIA Corporation. *CUDA C Programming Guide*. Available online:
535 https://docs.nvidia.com/cuda/pdf/CUDA_C_Programming_Guide.pdf

536 Pappenberger, F., Thielen, J., and Del Medico, M.: The impact of weather forecast
537 improvements on large scale hydrology: analysing a decade of forecasts of the
538 European Flood Alert System, *Hydrol. Process.*, 25, 1091–1113,
539 <https://doi.org/10.1002/hyp.7772>, 2011.

540 Pedersen, M.E.H. Good parameters for particle swarm optimization. *Tech. Rep. HL1001*.
541 Hvas Lab Copenhagen, Denmark. 2010

542 Schelfaut, K., Pannemans, B., van der Craats, I., Krywkow, J., Mysiak, J., and Cools, J.:
543 Bringing flood resilience into practice—the FREEMAN project, *Environ. Sci.*
544 *Policy*, 14 (7), 825-833, <https://doi.org/10.1016/j.envsci.2011.02.009>, 2011.

545 Segura-Beltrán, F., Sanchis-Ibor, C., Morales-Hernández, M., González-Sanchís, M.,
546 Bussi, G., Ortiz, E.: Using post-flood surveys and geomorphologic mapping to
547 evaluate hydrological and hydraulic models: The flash flood of the Girona River
548 (Spain in 2007, *J. Hydrol.*, 541, 310-329,
549 <https://doi.org/10.1016/j.jhydrol.2016.04.039>, 2011.

550 Skamarock, W. C. , Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Wang, W., and
551 Powers, J. G.: A Description of the Advanced Research WRF Version 2. Tech.
552 Rep., National Center for Atmospheric Research, 2005.

553 SNCZI, Guía Metodológica para el Desarrollo del Sistema Nacional de Cartografía de
554 Zonas Inundables, Ministerio de Medio Ambiente y Medio Rural y Marino,
555 Gobierno de España, 2011.

556 Sousa, M.C., Alvarez, I., Vaz, N., Gomez-Gesteira, M., and Dias, J. M.: Assessment of
557 wind pattern accuracy from the QuikSCAT satellite and the WRF model along the
558 Galician coast (Northwest Iberian Peninsula), *Mon. Weather Rev.*, 141 (2), 742-
559 753, <https://doi.org/10.1175/MWR-D-11-00361.1>, 2013.

560 Taylor, K.E.: Summarizing multiple aspects of model performance in a single diagram,
561 *J. Geophys. Res.*, 106, 7183-7192, 2001

562 Thielen, J., Bartholmes, J., Ramos, M. H., and De Roo, A.: The European flood alert
563 system-part 1: Concept and development, *Hydro. Earth Syst. Sc.*, 13 (2), 125-140,
564 <https://doi.org/10.5194/hess-13-125-2009>, 2009.

565 UNISDR. UNISDR terminology on disaster risk reduction. United Nations Office for
566 Disaster Risk Reduction, 2009.

567 UNISDR. Making development sustainable: the future of disaster risk management.
568 Global Assessment Report on Disaster Risk Reduction, United Nations Office for
569 Disaster Risk Reduction (UNISDR), Geneva, Switzerland, 2015.

570 UN, Global survey of early warning systems, A report prepared at the request of the
571 Secretary-General of the United Nations, 2006.

572 Upegui, J.J.V., Gutiérrez, A.B. Estimation of the time of concentration and the lag time
573 at san Luis creek basin, Manizales [Estimación del tiempo de concentración y
574 tiempo de rezago en la cuenca experimental urbana de la quebrada san luis,
575 manizales]. *DYNA (Colombia)*, 78 (165), pp. 58-71, 2011.

576 U.S. Army Corps of Engineers. Hydrologic Modeling System (HEC-HMS) Applications
577 Guide: Version 3.1.0. Davis: Institute for Water Resources, Hydrologic
578 Engineering Center. 2008.

579 U.S. Army Corps of Engineers, Hydrologic Modeling System (HEC-HMS) User's
580 Manual: Version 4.3. Institute for Water Resources Davis: Hydrologic
581 Engineering Center, 2018.

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584 **Figure Captions**

585

586 **Figure 1.** Area of study. In the upper right panel, the location of the entire catchment of
587 the shared Portuguese-Spanish river (shaded area) in the Iberian Peninsula and the
588 riverbed of the Miño river (blue line) are shown. The rain gauges (rg1, ...,rg7) located in
589 the catchment and the sub-basins (Sb₁, Sb₂ and Sb₃) of the domain (upper left panel), as
590 well as the area of study in Lugo (lower panel) are also shown. (PNOA courtesy of ©
591 Instituto Geográfico Nacional).

592 **Figure 2.** Flowchart of the proposed EWS.

593 **Figure 3.** Time series of the registered streamflow (dashed line) and numerical
594 streamflow (orange line) of the validation events: a) 01/2013, b) 01/2014, c) 02/2016 and
595 d) 03/2018). Taylor diagram of the validation cases are also shown.

596 **Figure 4.** Time series of the outflow at the control point obtained in the gauge station
597 (dashed line) and calculated using the three forecast windows (left panel) and Taylor
598 diagram for the same cases (right panel).

599 **Figure 5.** Numerical domain at Lugo. The land uses and the location of the boundary
600 conditions (red lines) are also shown. (PNOA courtesy of © Instituto Geográfico
601 Nacional).

602 **Figure 6.** Location of the five control points at the area of study in Lugo. (PNOA courtesy
603 of © Instituto Geográfico Nacional).

604 **Figure 7.** Comparison between water depth (h in meters) between the numerical model
605 (.) and the field data (x) for the three forecast windows 1-day (left), 2-day (middle) and
606 3-day (right). The range of the numerical values correspond to 3 times the standard
607 deviation of the elevations obtained from 12:00 to 16:00 on January 20th, 2013.

608 **Figure 8.** Maximum water depth (upper panel) and maximum velocity (lower panel)
609 obtained with Iber+ for the 1-day precipitation forecast. (PNOA courtesy of © Instituto
610 Geográfico Nacional).

611 **Figure 9.** Areas where hazard criterion is surpassed. (PNOA courtesy of © Instituto
612 Geográfico Nacional).

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615 **Table 1.** Location and elevation of the rain gauges located in the area of study (The
616 system of reference for latitude and longitude is the EPSG: 4326).

617 **Table 2.** Values of the correlation (Spearman's r) and normalised standard deviation (σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
618 σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
619 considering the complete time series of precipitation. The averaged values for each
620 precipitation forecast are also shown.

621 **Table 3.** Values of the correlation (Spearman's r) and normalised standard deviation (σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
622 σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
623 considering only rainy events (above the 75th percentile). The averaged values for each
624 precipitation forecast are also shown.

625 **Table 4.** Main characteristics of the analysed flood events

626 **Table 5.** Curve number (CN) and lag time (L_g) values for each sub-basin for different
627 flood events. The mean value and the standard deviation are provided in lower rows.

628 **Table 6.** Manning's coefficients of the numerical domain.

629 **Table 1.** Location and elevation of the rain gauges located in the area of study (The
630 system of reference for latitude and longitude is the EPSG: 4326).
631

Rain gauge id.	Name	Latitude	Longitude	Elevation (m.a.s.l.)
rg ₁	Labrada	43.4054	-7.50205	662
rg ₂	Lanzós	43.3746	-7.64468	470
rg ₃	Guitiriz-Mirador	43.2266	-7.78307	684
rg ₄	Sanbreixo	43.1457	-7.79112	496
rg ₅	Castro de Rei Lea	43.1559	-7.48588	428
rg ₆	Pol	43.1626	-7.28258	647
rg ₇	Corno do Boi	43.0374	-7.89265	731

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634 **Table 2.** Values of the correlation (Spearman's r) and normalised standard deviation (σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
635 considering the complete time series of precipitation. The averaged values for each
636 precipitation forecast are also shown.
637
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Rain gauge	Forecast window (h)					
	1-24		25-48		49-72	
	r	σ_n	r	σ_n	r	σ_n
rg ₁	0.84	0.80	0.82	0.81	0.77	0.80
rg ₂	0.84	1.09	0.82	1.07	0.79	1.07
rg ₃	0.83	1.00	0.81	0.96	0.77	0.99
rg ₄	0.81	0.97	0.79	0.96	0.75	0.98
rg ₅	0.81	1.13	0.80	1.10	0.76	1.12
rg ₆	0.84	1.16	0.83	1.07	0.79	1.07
rg ₇	0.83	1.05	0.81	1.06	0.77	1.10
Mean value	0.83	1.03	0.81	1.00	0.77	1.02

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641 **Table 3.** Values of the correlation (Spearman's r) and normalised standard deviation (σ_n) of the precipitation forecast using the measured data as reference at each rain gauge,
642 considering only rainy events (above the 75th percentile). The averaged values for each
643 precipitation forecast are also shown.
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Rain gauge	Forecast window (h)					
	1-24		25-48		49-72	
	r	σ_n	r	σ_n	r	σ_n
rg ₁	0.66	0.72	0.61	0.70	0.53	0.72
rg ₂	0.71	1.00	0.63	0.98	0.56	0.99
rg ₃	0.70	0.98	0.61	0.93	0.59	0.98
rg ₄	0.73	0.93	0.65	0.90	0.60	0.93
rg ₅	0.68	1.02	0.63	1.01	0.54	1.04
rg ₆	0.69	1.14	0.65	0.98	0.56	1.00
rg ₇	0.74	1.03	0.68	1.02	0.63	1.10
Mean value	0.70	0.97	0.64	0.93	0.57	0.97

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650 **Table 4.** Main characteristics of the analysed flood events

651

Date of the flood event	Duration (days)	Initial flow (m^3s^{-1})	Initial depth (m)
28/12/09	4	52	1.3
17/11/10	5	116	1.7
17/01/13	10	164	1.9
11/03/13	5	179	2.0
05/11/13	7	234	2.3
14/01/13	10	165	1.9
28/01/14	15	202	2.1
01/03/14	4	134	1.8
30/01/15	3	184	2.0
01/03/15	3	134	1.8
10/02/16	7	216	2.1
26/02/16	3	137	1.8
05/03/16	4	175	2.0
10/03/18	6	154	1.9
30/03/18	4	201	2.1

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654 **Table 5.** Curve number (CN) and lag time (L_g) values for each sub-basin for different
655 flood events. The mean value and the standard deviation are provided in lower rows.

656

Date of the flood event	Sb ₁		Sb ₂		Sb ₃	
	CN	L_g (min)	CN	L_g (min)	CN	L_g (min)
12/09	92	1154	97	2700	98	2770
11/10	80	1140	84	2702	80	2781
03/13	79	1157	96	2701	99	2774
11/13	80	1148	86	2685	83	2778
01/14	78	1155	96	2700	98	2767
03/14	81	1153	88	2706	92	2764
01/15	96	1153	99	2701	99	2773
03/15	81	1151	91	2700	98	2771
02/16	81	1155	88	2700	98	2767
03/16	82	1153	80	2711	84	2764
03/18	80	1152	82	2691	93	2769
Mean	85	1152	90	2700	93	2771
σ	6	4	6	7	7	5

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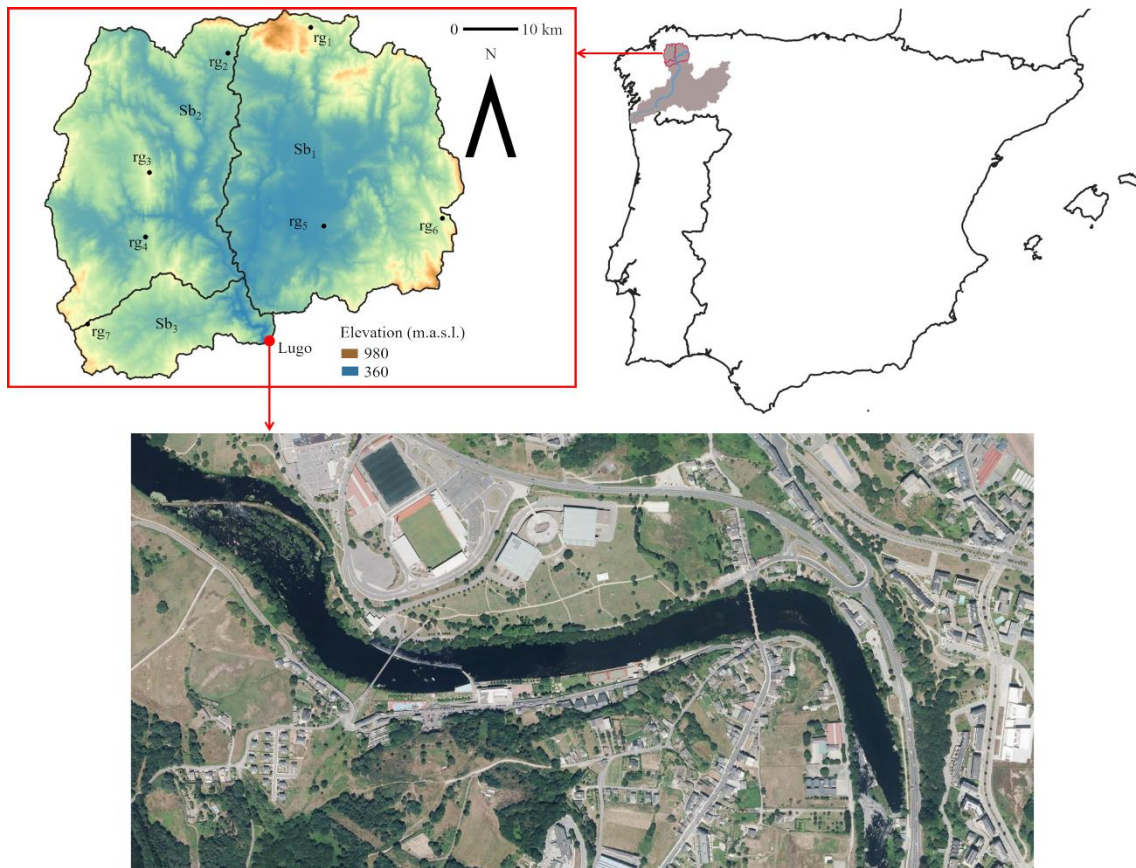
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659 **Table 6.** Manning's coefficients of the numerical domain.

Land's uses	Manning's coefficient ($\text{s m}^{-1/3}$)
River	0.025
Brush	0.050
Trees	0.120
Sparse vegetation	0.080
Infrastructure	0.020
Industrial	0.100
Residential	0.150

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663 **Figure 1.** Area of study. In the upper right panel, the location of the entire catchment of
 664 the shared Portuguese-Spanish river (shaded area) in the Iberian Peninsula and the
 665 riverbed of the Miño river (blue line) are shown. The rain gauges (rg1, ...,rg7) located in
 666 the catchment and the sub-basins (Sb₁, Sb₂ and Sb₃) of the domain (upper left panel), as
 667 well as the area of study in Lugo (lower panel) are also shown. (PNOA courtesy of ©
 668 Instituto Geográfico Nacional).

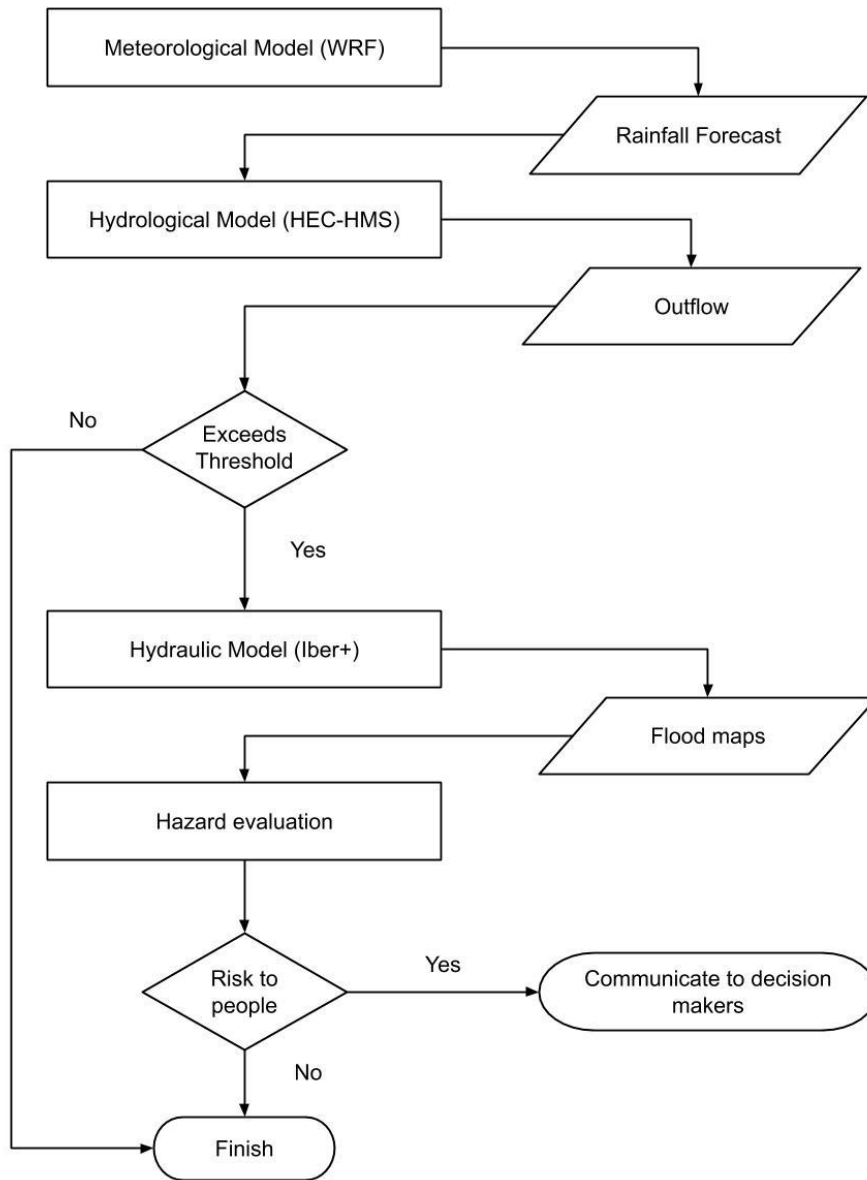
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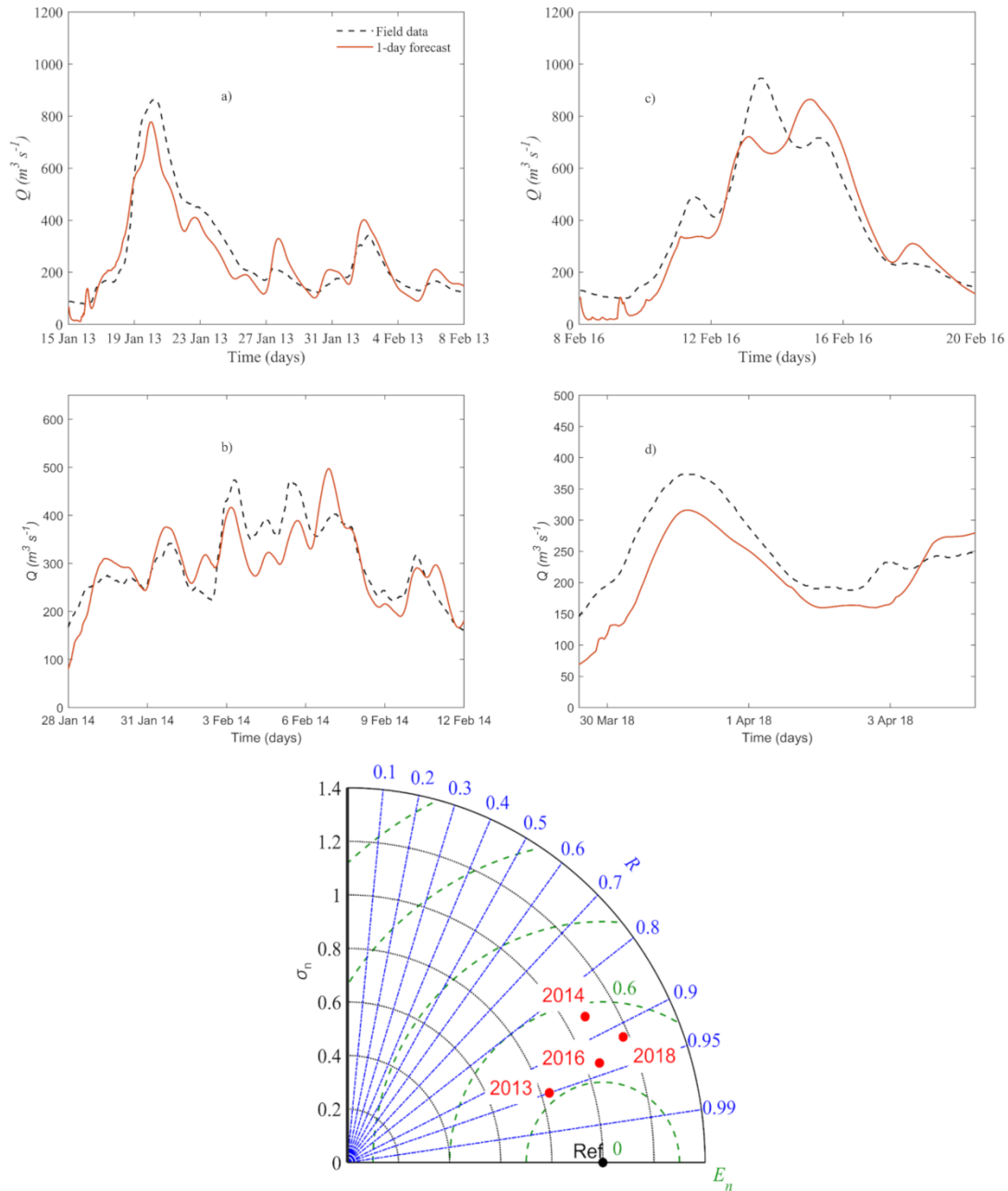
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675 **Figure 2.** Flowchart of the proposed EWS.

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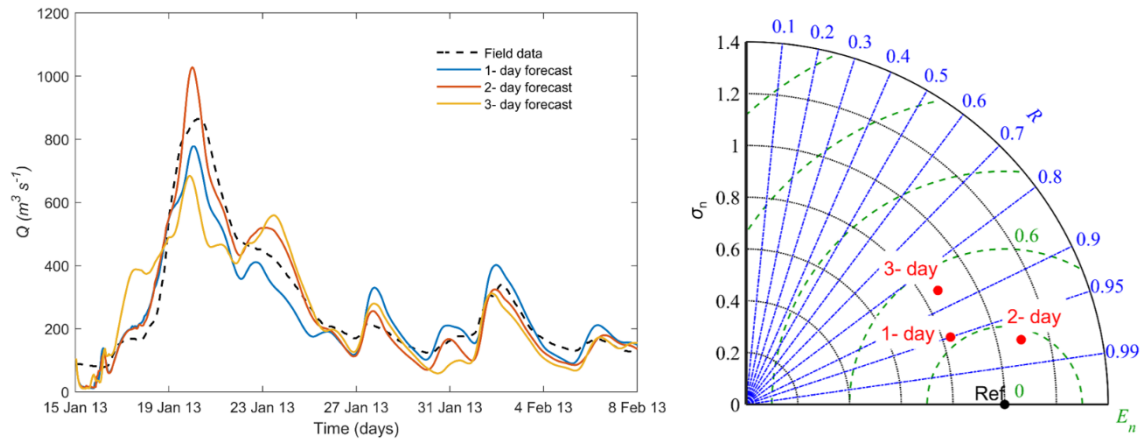
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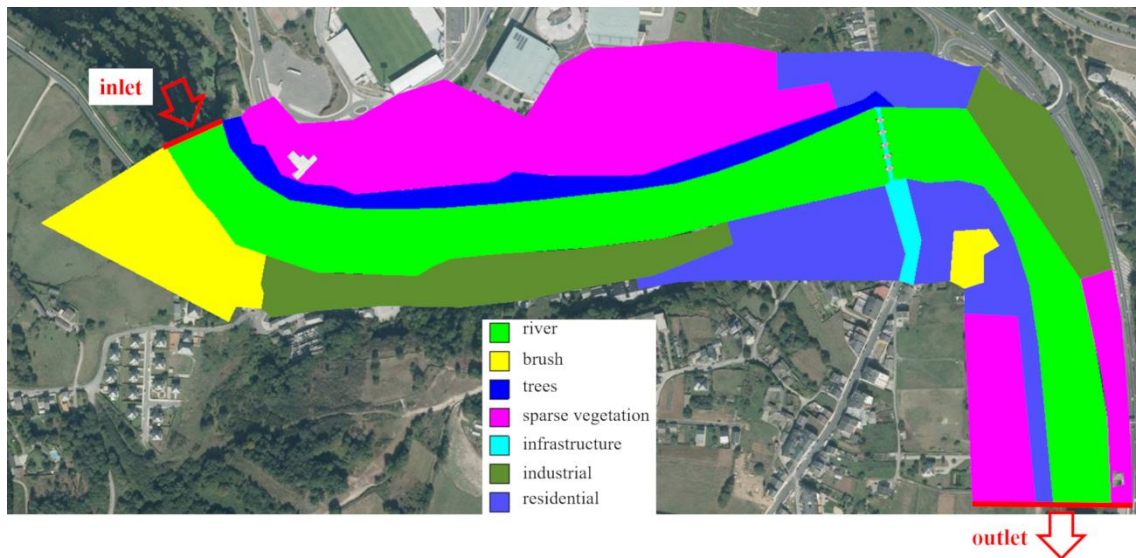
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 681 streamflow (orange line) of the validation events: a) 01/2013, b) 01/2014, c) 02/2016 and
 682 d) 03/2018). Taylor diagram of the validation cases is also shown.

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Figure 4. Time series of the outflow at the control point obtained in the gauge station (dashed line) and calculated using the three forecast windows (left panel) and Taylor diagram for the same cases (right panel).



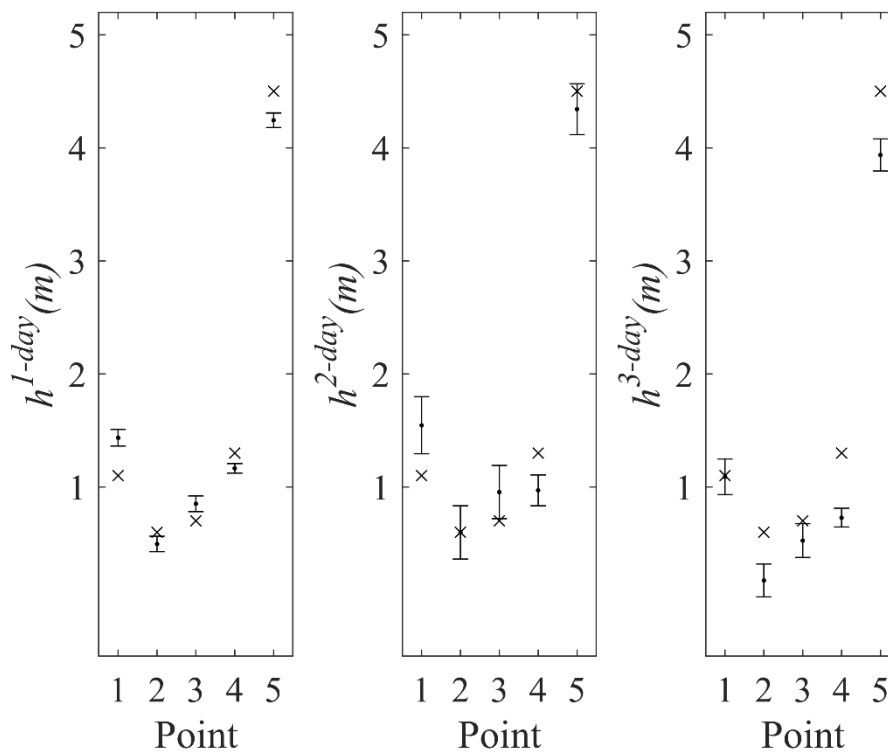
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Figure 5. Numerical domain at Lugo. The land uses and the location of the boundary conditions (red lines) are also shown. (PNOA courtesy of © Instituto Geográfico Nacional).



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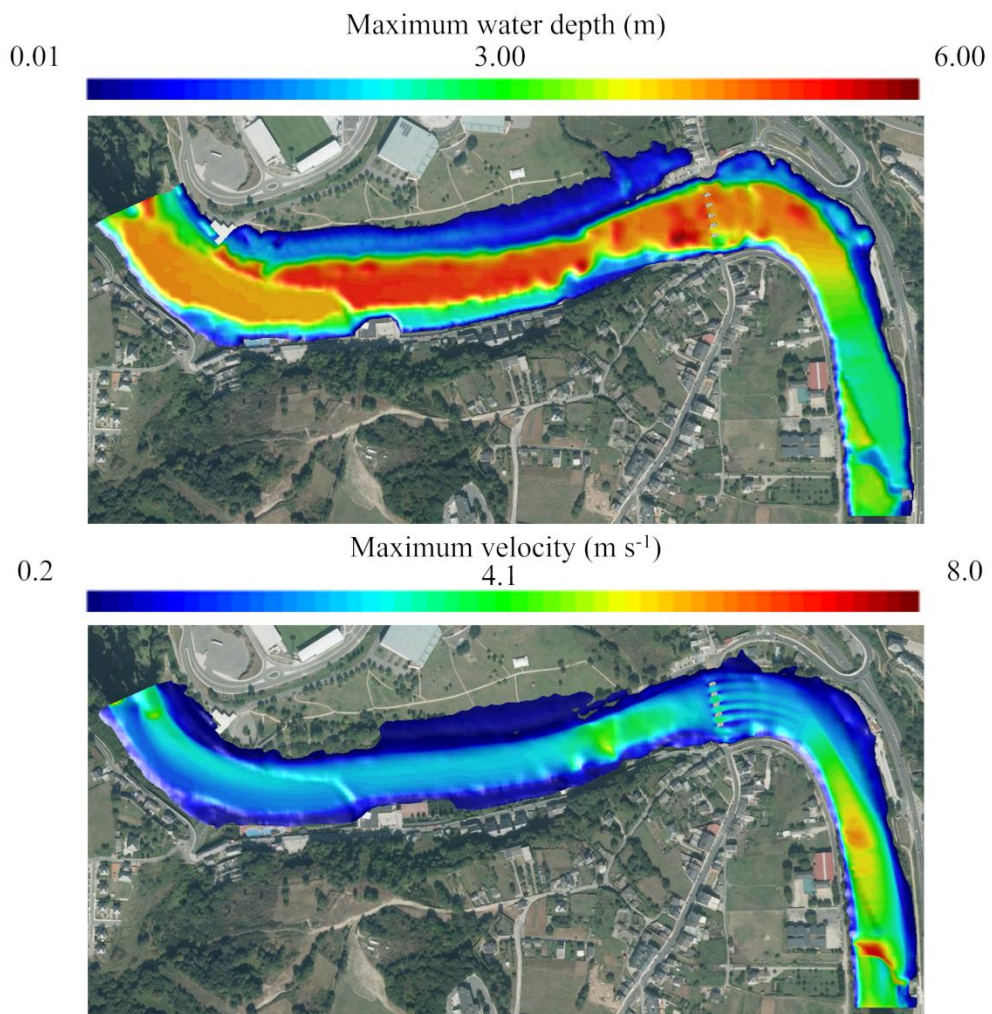
Figure 6. Location of the five control points at the area of study in Lugo. (PNOA courtesy of © Instituto Geográfico Nacional).



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715 **Figure 8.** Maximum water depth (upper panel) and maximum velocity (lower panel)
716 obtained with Iber+ for the 1-day precipitation forecast. (PNOA courtesy of © Instituto
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724 **Figure 9.** Areas where hazard criterion is surpassed. (PNOA courtesy of © Instituto

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