



Tectonic Origin Tsunami Scenario Database for the Marmara Region

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Abstract. This study presents the first tsunami scenario database in Marmara Sea, Turkey referring to 30 different earthquake scenarios obtained with the combinations of 32 possible fault segments. The fault mechanisms in Marmara Sea have been studied in detail within FP-7 MARSite project, which were derived from various databases and literature review. Tsunami simulations have been performed according to these defined 30 earthquake scenarios by tsunami numerical code NAMI DANCE (NAMIDANCE, 2011) which solves Nonlinear Shallow Water Equations (NLSWE) using leap-frog scheme. For each earthquake scenario, tsunami hydrodynamic parameters, mainly maximum water surface elevations, arrival time of first wave and maximum wave, and water level fluctuations were calculated at 1333 synthetic gauge points meticulously selected along the coasts of Marmara Sea. The overall simulation results indicate that maximum expected wave heights due to these earthquake scenarios are between 1m and 2m and even more than 2m at some locations along Marmara coasts, such as Kadikoy, Halic and Silivri coasts in Istanbul and Bayramdere and Kursunlu districts along the coasts of Bursa province. The estimated maximum water levels at Bostanci, Pendik and Buyukada coasts in Istanbul, Cinarcik and Bandirma towns and at the entrance of Izmit Bay would reach up to 2m. Tekirdag coasts and Buyuk Cekmece and Bakirkoy coasts in Istanbul and Yalova coasts would experience maximum tsunami wave amplitudes around 1.5m. The waves reach up to 1m at Izmit and Gemlik Bays, Erdek Peninsula and Marmara Island. The overwiew the results reveal that higher historical tsunami wave heights observed in Marmara Sea cannot be explained by only earthquake-generated tsunamis. Therefore, there is strong agreement on considering submarine landslides as the primary tsunami hazard component in the Marmara Sea as experienced during history and expected in the future.

1 Introduction

Marmara Region is located at the northwest part of Turkey, being one of the most important settlements throughout history as the passageway between two continents, Europe and Asia. The region literally acts as the heart of economy in the country hosting mainly the megacity Istant vith a population of higher 15 n 15 million and huge capacity of trading, various industrial facilities, ports, airports as well as other densely populated cities.

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Marmara Sea located in the middle of this region is seismic very active area having various faults of different types with great possibility of triggering big earthqu larger than 7.0 (Ambraseys, 2002). There are are number of references in literature proving the generation of historical tsunamis in the Marmara Sea caused by these earthquakes (Mihailovic 1927; Gundogdu 1986; Oztin and Bayülke 1991; Oztin 1994; Ambraseys and Finkel 1987, 1995; Altınok and Ersoy 2000; Arel and Kiper 2000; Altınok et al. 2001, 2003, 2011; Yalciner et al. 2001a, b, c; Ambraseys 2002; Cetin et al. 2004; Rothaus et al. 2004; Tinti et al. 2006; Meral Ozel et al. 2011). The most significant ones among those the Istanbul earthquakes on 10 September 1509, on 22 May 1766, and on 10 July 1894; the Sarkoy-Murefte earthquake on 9 August 1912; and the 17 August 1999 Kocaeli earthquake (Necmioğlu, 2016). During the 1509 earthquake, with a magnitute lose to 8.0, the sea flooded the shores along Istanbul coasts, waves crashed against city walls and around 4000-5000 people died in the city (Ambraseys and Finkel 1995). Tsunami waves with probably more than 6.0 m height overtopped the city walls and caused flooding (Oztin and Bayülke 1991). 1766 Istanbul earthquake, on the other hand, triggered tsunami with considerab of Gemlik Bay in Eastern Marmara (Ambraseys and Finkel 1995; Altınok et al. 2003). There are mainly two motivations of list study: the first one is investigating the nature of historical tsunamis in Marmara Sea, namely whether they are generated solely due to those significant earthquakes or not. The second aim is directly correlated with the operations of Regional Earthquake and Tsunami Monitoring Center (RETMC) in Kandilli Observatory and Earthquake Research Institute (KOERI) acting as 24/7 National Warning Center (NTWC-TR), which provides monitoring and warning services to the Eastern Mediterranean, Aegean, and Black Seas since 1 July 2012 under the framework of the Intergovernmental Coordination Group for the Tsunami Early Warning and Mitigation System in the North-eastern Atlantic, the Mediterranean and connected seas (ICG/NEAMTWS) initiative and has been accredited during 13th Session of ICG in Bucharest, Romania, 26–28 September 2016 as a Tsunami Service Provider (TSP-TR). Due to short arrival times of first waves in ∇ mara coasts, having prepared tsunami scenarios covering various possible earthquakes is quite vital.

On account of this purpose and necessity, we have identified a comprehensive set of earthquake scenarios for the whole Marmara Region obtained by the compilation of historical records, past studies in literature and empirical results. These scenarios constitute the basis for tsunami numerical modeling conducted to obtain tsunami scenario database in Marmara Sea.

32 different fault segments are finally determined in that manner of which methodology is de probed in the following section.

2. Methodology

2.1 Identification of Earthquake Scenarios

The main structural element controlling the morphological and structural features in Marmara Sea region is the northern strand of the North Anatolian Fault zone, which idered as a principal e-slip displacement zone (Alpar and Yaltırak, 2002). The North Anatolian Fault (NAF) is a major right-lateral, strike-slip fault that extends more than 1200 km from eastern Turkey to the north Aegean Sea (Sengör et al., 2005). It accommodates the relative right-lateral motion between the Anatolian region and Eurasia at a geodetic rate of ~25 mm/yr (Meade et al., 2002; Reilinger et al., 2006). Along its westernmost segment, the



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fault bifurcates into northern and southern branches, the northern branch following Izmit Bay and entering the Sea of Marmara southeast of Istanbul. By far the majority of long-term fault slip occurs on the northern fault branch following the northwest striking Princes' Islands Fault (PIF) and joining the east-west striking Central Marmara Fault (CMF) immediately south of Istanbul [Le Pichon et al., 2001; Armijo et al., 2005]. After traversing much of the Sea of Marmara, the CMF merges with the Ganos Fault, exiting the Sea along the Ganos Peninsula. Ergintav et al. (2014) concluded that the Princes' Islands segment is most likely to generate the next M > 7 earthquake along the Sea of Marmara segment of the NAF. Armijo et al. (2005) stated that a zone of maximum loading with at least 4–5 m of slip deficit encompassing the strike-slip gment 70 km long between the Cinarcik and Central Basins would alone be capable of generating a lage magnitude earthquake in the order of Mw 7.2. Hergert et al. (2011) argues in the that the Main Marmara Fault can be interpreted as a through-going fault that slips almost purely in a strike-slip sense, but they also point out that, not contradictory to the previous statement, there is significant dipslip motion at some sections of the Main Marmara Fault. The South Marmara Fault lies between the highly active northern branch and the weakly active (but still capable of generating magnitude 7 earthquakes) southern branch (Pichor 11., 2014).

The geometry of the possible tsunamigenic faults in the Marmara Sea has been derived from the GIS Database of the fault parameters in the Marmara region produced within MARsite - New Directions in Seismic Hazard assessment through Focused Earth Observation in the Marmara Supersite (FP7 Project-ENV.2012 6.4-2, Grant 308417). Fault parameters, on the other hand, were subjectively assigned through an extensive review of the literature (Alpar and Yaltırak, 2002; Altınok and Alpar, 2006; Armijo et al., 2005; Ergintav et al., 2014; Gasperini et al., 2011; Hebert et al., 2005; Hergert et al., 2011; Hergert and Heidbach, 2010; Imren et al., 2001; Kaneko, 2009; Le Pichon et al., 2001; Le Pichon et al., 2003; Le Pichon et al., 2014; Oglesby and Mai, 2012; Sengor et al., 2014; Tinti et al., 2006; Utkucu et al., 2009). Based on the databases and literature review, faults segments were simplified in order to be able to use them as input for tsunami modelling, where each segment correspond to a rectangular area with an associated hypothetical uniform slip (Figure 1). All parameters required for the identification of the segments, such as geographical coordinates for the start- and end-points of the segments, hypocentre, type of fault, strike, dip, rake, length and width of the segment, focal depth (where the top of the fault has been set to 0.5 km depth) and corresponding displacements according to empirical relations provided by Leonard (2010) and Wells and Coppersmith (1994) are presented in Table 1. This was followed by the definition of different hypothetical rupture scenarios reaching a total number of 30 scenarios as described in Table 2. In each scenario, special attention was given that the total earthquake moment derived from the summation of the earthquake moments, which are assigned to each segment, has been preserved by each scenario obtained through the combination of several segments. Slip values have been assigned using the same logic but in an arbitrary manner without any prior assumption, so that heterogeneous earthquake rupture scenarios can be represented. Figure 2 shows the map of all earthquake scenarios collectively including the fault segments (green lines) assigned accordingly.





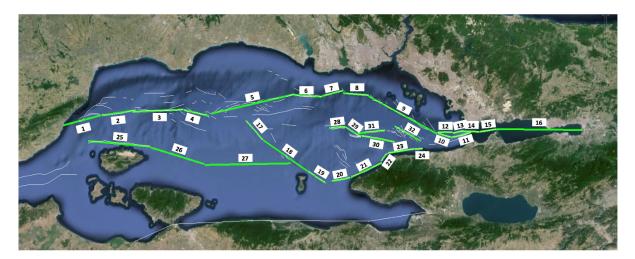


Figure 1: Simplified fault segments in Marmara identified for tsunami modelling. Segments correspond to a rectangular area with an associated uniform slip.

5 2.2 Tsunami Numerical Modeling

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Based on 30 earthquake scenarios defined in Section 2.1, tsunami numerical modeling has been performed using the numerical code NAMI DANCE (NAMIDANCE, 2011), which solves Nonlinear Shallow Water Equations (NLSWE) with leap-frog scheme both in Cartesian and Spherical coordinate system. NAMIDANCE was tested, validated and verified against analytical solutions, laboratory measurements and field observations in several scientific articles (NTHMP, 2017; Lynett et al., 2017; Velioğlu, 2017).

Tsunami numerical modelling is performed using 90m grid sized bathymetry - topography data as a single study domain, which includes multi-beam bathymetric measurements and 900m grid sized GEBCO data in the sea and 30m grid sized ASTER data on land. Besides, coastline and coastal defence structures i.e. breakwaters, groins alarge docks in the ports were also digitized in GIS environment and added to bathymetry - topography data for increasing the resolution and precision in coastal zones.

The initial sea surface at the time of fault rupture for each segment has been calculated using Okada (1985) formula. In each scenario, it was assumed that all designated fault segments are ruptured at the same ance. NAMIDANCE calculates the sea surface after the rupture of each segment and combines them in order to output the final sea surface that was used as the tsunami source of each scenario. For instance for Earthquake Scenario #1 (SN01), segment-1, segment-2, segment-3 and segment-4 are the fault components designated for this scenario. The sea surface for each fault segment was calculated using Okada (1985) formula as if they are rupturing at the same instance. NAMIDANCE then outputs a final sea surface as the combination of these four segments as an overall tsunami source for SN01. We applied same edure for all earthquake scenarios accordingly.





The synthetic gauge points along the coasts of Marmara Sea were selected with very sensitive analysis so as to locate them in shallow zone at water depths less than 20m. We considered the locations of industrial facilities, residential areas, harbors, marinas, factories and six Tsunami Forecast Points (TFPs) while selecting those gauge points (TFPs are located at Marmara Eregli, Haydarpasa, Yalova, Mudanya, Erdek and Degirmencik, where the arrival time of first wave and tsunami alert level are calculated and included in national tsunami alert messages disseminated from Regional Tsunami and Earthquake Monitoring Center in KRDAE; see Figure 3). The total number of gauge points is 1333 located at the water depths of less than 5m. The northern part of the area has much more important locations then the southern part, therefore gauge points in the northern part is denser than the southern part.

The maximum and minimum water surface elevations of tsunami sources calculated are given in Table 3. As seen from the table, the initial sea surface disturbances for all scenarios are less than 1m. The highest sea surface was calculated for SN23, which includes the rupture of segments 17, 18 and 19 located at the center of Marmara Sea raking W-SE direction.

Tsunami simulations were conducted for each scenario during 2 hours the corresponding tsunami sources. Tsunami hydrodynamic parameters such as maximum and minimum wave amplitudes, arrival times of first and maximum wave, flow depths and current velocities were calculated throughout Marmara basin and at 1333 synthetic gauge points. We discuss the simulation results in the following section in an integrated manner instead of giving the details of each scenario results in the main text of this paper. However, the distributions of maximum wave amplitudes and arrival time of first wave plotted for each scenario are provided in Supplementary Material. These plots directly reflect the hydrodynamic parameters calculated at each synthetic gauge point according to the defined color scale. The maximum wave amplitudes less than 25cm (accepted as threshold for critical water level that has potential to fall a person do were colored with green as a representative of relatively safer coastal zones of Marmara Sea. Besides, Supplementary Material also includes water level fluctuations calculated for each scenario at six TFPs (Haydarpasa, Yalova, Mudanya, Erdek, Degirmencik and M. Eregli) and also 20 other locations that are selected out of 1333 synthetic gauge points considering the locations of industrial facilities, harbors, marinas, refineries and shipyards. These fluctuation plots also indicate the arrival time of first wave at each location.

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Particular Par			Moment	n-cm)	8,913E+25	7,079E+25	1,413E+26	1E+26	7,943E+26	2,512E+25	1,778E+25	2,512E+25	3,981E+26	2,512E+25	2,512E+25	1,259E+25	6,31E+24	2,512E+25	1,778E+25	1,995E+26	5,012E+25	1,995E+26	3,548E+25	1,259E+25	3,548E+25	4,467E+24	2,512E+25	4,467E+24	1,995E+26	1,413E+26	2,818E+26	2,512E+25	3,548E+25	3,548E+25	5,012E+25
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Name		L20	ΜW		-	-										_	-	-		9'9	-	-	-	-		-		-	9'9	-	-		-		6,1
Name			0	E)	9'0	9′0	2'0	5′0	6′0	0,4	4′0	0,4	2′0	6,0	0,4	6,0	0,2	9′0	0,3	8′0	0,4	9′0	0,4	6'0	0,4	0,2	0,4	0,2	2'0	2'0	6′0	6'0	6'0	0,4	0,4
Santa Fault Secone Fault State			Focal Depth	(km)	6'1	6′2	6′2	6′2	7,9	5,4	5,4	5,4	5,4	6,7	6,7	6,7	6,7	6,7	6,7	6,7	5,4	5,4	5,4	5,4	5,4	5,4	5,4	5,4	6'2	6'2	6'2	5,4	5,4	5,4	5.4
Santa Fault Secone Fault State			egment Area	(km2)	247,5	225,0	330,0	165,0	547,5	115,0	95,0	110,0	335,0	8,89	112,5	75,0	43,8	100,0	87,5	450,0	110,0	225,0	0'06	0,08	145,0	40,0	115,0	35,0	397,5	360,0	540,0	75,0	80,0	145,0	120.0
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Table 1: Parametric information on the segments.





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	Davg (m)		1,28	2,43	2,12	2,36	2,24	4,46	3,06	3,23	2,91	2,70	2,39	0,65	2,19	1,90	1,69	1,59	3,66	2,67	2,21	1,86	0,62	1,06	2,06	0,61	1,54	1,38	1,48	66'0	1,24	2,58
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5	Mo	(dyn/cm)	4,01173E+26	1,1955E+27	1,26352E+27	1,66163E+27	1,71187E+27	7,94328E+26	8,62349E+26	1,26046E+27	1,31069E+27	1,32226E+27	1,52178E+27	6,80205E+25	4,66128E+26	5,16365E+26	5,27928E+26	7,27454E+26	3,98107E+26	4,48345E+26	4,60E+26	6,59434E+26	6,18005E+25	2,61327E+26	2,85126E+26	8,21231E+25	1,99526E+26	3,4078E+26	6,22618E+26	9,60815E+25	1,10719E+26	1,41254E+26
1	Davg (m)		85'0	99'0	0,61	0,62	09'0	0,79	0,65	0,65	0,61	65'0	0,61	0,41	0,53	0,49	0,47	0,54	0,65	0,53	0,50	0,58	0,34	0,54	0,55	0,42	0,77	69'0	0,74	0,44	0,40	0,46
91	Mw		8'9	6'9	0'2	2,0	2,0	6,7	8'9	6'9	6'9	6'9	0'2	6,4	9'9	6,7	6,7	6'9	6,5	9'9	9'9	8'9	6,3	6,7	6,5	6,4	9'9	8,9	6,9	6,4	6,3	6,2
12010	Mo	(dyn/cm)	1,81514E+26	3,22767E+26	3,65729E+26	4,36523E+26	4,55422E+26	1,41254E+26	1,84215E+26	2,5501E+26	2,73908E+26	1508,8 2,88586E+26	3,88586E+26	4,29613E+25	1,13756E+26	1,32655E+26	1,47332E+26	2,47332E+26	7,07946E+25	8,96934E+25	1,04E+26	2,04371E+26	3,35766E+25	1,33577E+26	7,52972E+25	5,72152E+25	1E+26	1,70795E+26	3,12048E+26	4,29439E+25	3,56078E+25	2,51189E+25
Fault	Area	(km2)	5'296	1515	1835	2170	2351,3	547,5	867,5	1202,5	1383,8	8'8051	1958,8	320	655	836,25	961,25	1411,3	335	516,25	641,25	1091,3	306,25	756,25	425	415	397,5	757,5	1297,5	300	275	168,75
	32	168,8					2			,	-	-	,			8	5	1		21	9	,	(6)	7	П	П			-			2,58
Ì	31	145,0 120,0																													1,24	
	30																													1,232		
	53	80,0																												0,75	1,24	
	28	75,0																												0,75	1,24	
	27	0 540,0																	_										5 1,19		_	
-	56	360,0							_	Н							_	_	_			Н	Н		Н		4	7 1,73	9 2,215	_	_	_
-	1 25	35,0 397,5		Н	L				_	Н									_			Н	Н		Н	0,	1,54	1,07	1,19		_	_
	23 24			Н	H	L				Н						L						Н	Н		Н	0,70 07,0		L				_
ł	22 2	40,0 115,0	_	Н						Н							_	_	_			Н	Н		Н	0,70					_	
ł	21 2	145,0 40	_	Н					_	Н	_	_				_	_	_	_			Н	Н		Н	0,70 0,		H	_	_	_	
ł	20	80,0		Н		Н			_	Н						Н	_	_	_			Н	Н		Н	0,70		Н	_	_	_	
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a	18	225,0		П																			П		2,58	П						
SEGMENTS INCLUDED	17	110,0																				П	П		1,48							
MENTS	16	450,0											2,99					2,379				2,32		1,329								
SEG	15	87,5										1,75	1,12				0,79	0,59			1,00	0,67	0,35	0,67								
	14	100,0										1,75	1,12				0,79	0,59			1,00	0,67	0,931 0,931	0,67								
	13	43,8										1,75	1,12				0,79	0,59			1,00	0,67		0,67								
	12	5 75,0										1,75	1,12			-	0,79	0,59			1,00	0,67	0,35	0,67								
-	11	112,5					5 1,45				3 1,63					2 1,18				1,44												
-	10	,0 68,8				34	3,36 1,45			3	1,63	71	88		35	2,85 1,42	8	62	9	3,34 1,44	1	24	Н									
	8	95,0 110,0 335,			72	3,5	1,45 3,3		1,75	1,00 4,03	1,63 1,63 3,643	15 3,3	1,12 1,12 1,12 2,988	25	3,285	1,42 2,8	0,79 3,38	9 2,379	3,66	3,3	3,31	2,324	Н						_	_	_	_
ł	7	0,0		Н	1,84 1,84	55 1,5	1,45 1,4		1,75 1,7	1,00 1,0	63 1,6	75 1,7	12 1,1	181 0,52	1,04 1,04	1,18 1,4	7,0 67,0	65'0 65'0	_			Н	Н	Н	Н							
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ł	4	65,0 54	1,06	1,40 1,40 4,249	1,84 1,84 3,708 1,84	1,55 1,55 1,55 3,534 1,55 1,55 1,55 3,534	1,45 1,45 1,45 1,45 3,36 1,45	4	3,	4	3,	3,	2,	\vdash		H	\vdash	\vdash	Н			Н	Н	\vdash	Н	\vdash		\vdash	\vdash	\vdash		_
1	6	330,0	1,06 1,06 1,06	1,40	1,84	1,55	1,45			H						H	\vdash	Н	Н			H	H		Н	Н						
Ì	2	225,0	1,06	1,40	1,84	1,55	1,45			П						П						П	П		П							
	1	247,5	1,91	1,40	1,84	1,55	1,45															П	П									
SCENARIO		Segment Area 247,5 225,0 330,0 165,0 547,5 115,0	SN01	SN02		SN04	SN05	90NS	SN07	SN08	60NS	SN10	SN11	SN12	SN13	SN14	SN15	SN16	SN17	SN18	SN19	SN20	SN21	SN22	SN23	SN24	SN25	SN26	SN27	SN28	SN29	SN30

Table 2: Displacements assigned to fault segments for 30 scenarios to be used in tsunami modelling





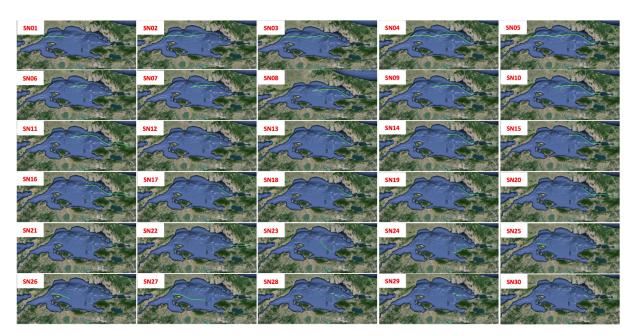


Figure 2: Map representation of the scenarios considered in this study.



5 Figure 3: The locations of Tsunami Forecast Points (white dots and white-highlighted names written in red) and other important coastal districts (red stars and white names) where estimated tsunami effect is significant.





NAME OF SCENARIO	MIN. WATER SURFACE ELEVATION (m)	MAX. WATER SURFACE ELEVATION (m)	SEGMENTS
SN01	-0.08	0.07	1-2-3-4
SN02	-0.54	0.33	1-2-3-4-5
SN03	-0.47	0.29	1-2-3-4-5-6-7-8
SN04	-0.45	0.28	1-2-3-4-5-6-7-8-9
SN05	-0.43	0.26	1-2-3-4-5-6-7-8-9- 10-11
SN06	-0.58	0.33	5
SN07	-0.50	0.29	5-6-7-8
SN08	-0.53	0.31	5-6-7-8-9
SN09	-0.48	0.28	5-6-7-8-9-10-11
SN10	-0.44	0.26	5-6-7-8-9-12-13-14- 15
SN11	-0.39	0.23	5-6-7-8-9-12-13-14- 15-16
SN12	-0.03	0.05	6-7-8
SN13	-0.39	0.25	6-7-8-9
SN14	-0.33	0.22	6-7-8-9-10-11
SN15	-0.41	0.26	6-7-8-9-12-13-14-15
SN16	-0.29	0.18	6-7-8-9-12-13-14- 15-16
SN17	-0.48	0.28	9
SN18	-0.44	0.25	9-10-11
SN19	-0.44	0.25	9-12-13-14-15
SN20	-0.31	0.18	9-12-13-14-15-16
SN21	-0.033	0.05	12-13-14-15
SN22	-0.04	0.06	12-13-14-15-16
SN23	-0.61	0.39	17-18-19
SN24	-0.03	0.04	20-21-22-23-24
SN25	-0.07	0.10	25
SN26	-0.08	0.09	25-26
SN27	-0.06	0.10	25-26-27
SN28	-0.05	0.07	28-29-30
SN29	-0.09	0.23	28-29-31
SN30	-0.44	0.25	32

Table 3: Maximum and minimum wate rface elevations of tsunami sources calculated for each earthquake scenario with corresponding fault segments

5 2.3 Summary of Results

As described in previous section, the simulation results are presented here as an integrated distribution of maximum wave amplitudes for overall tsunami scenario database in Marmara Sea. The maximum wave amplitudes were calculated at each synthetic gauge point for 30 earthquake scenarios. The calculated results of all scenarios at each gauge point were sorted from





larger to smaller and the highest value was stored as the representative maximum wave amplitude at this gauge point. After defining all maximum wave amplitude values at 1333 synthetic gauge points, their integrated distribution was plotted for entire Marmara Sea (Figure 4). As described in previous section, the coastal zones with green color represents relatively safer locations according to the earthquake scenarios used in this database.

Following the same procedure, the integrated distribution of arrival time of maximum waves (that is exactly the time of occurrence of maximum wave amplitude at each gauge point in entire earthquake scenarios) was plotted in Figure 5. The results show that the arrival of maximum waves is expected at Prince Islands, Yalova coasts, some parts of Kadikoy and Silivri coasts within 5 minutes as right ally short time for evacuation (refer to Figure 3 for the locations of affected districts).

Due to the evaluation of the modeling results for all earthquake scenarios, maximum wave amplitudes would be estimated between 1m and 2m and even more than 2m at some locations along Marmara coasts. The results show that the maximum wave amplitudes for Kadikoy, Halic and Silivri coasts in Istanbul and Bayramdere and Kursunlu districts along the coasts of Bursa province would be more than 2m. The estimated maximum water levels at Bostanci, Pendik, Buyukada coasts in Istanbul; and Cinarcik, Bandirma and at the entrance of Izmit Bay would reach up to 2m. Tekirdag coasts especially M. Eregli, B. Cekmece and Bakirkoy coasts in Istanbul and Yalova coasts would experience maximum tsunami wave amplitudes around 1.5m. The waves reach up to 1m at Izmit and Gemlik Bays, Erdek Peninsula and Marmara Island. A summary of the simulation results for all earthquake scenarios are given in Table 4 including the calculated maximum wave amplitudes for corresponding earthquake scenario and the names of most affected coastal regions.



Figure 4: The integrated distribution of maximum wave amplitudes calculated at each gauge points for all earthquake scenarios





Figure 5: The integrated distribution of arrival times of maximum wave amplitudes calculated at each gauge points for all earthquake scenarios

Maximum wave amplitude calculated	Corresponding scenario	Name of most affected coastal region
2m – 2.2m	SN06	Kursunlu, Bayramdere, Eastern Silivri coasts, Kadikoy, Halic
1.75m - 2m	SN02, SN03, SN07, SN08, SN09, SN23	Bostanci, Pendik, Cinarcik, Bandirma and Buyukada
1.5m – 1.75m	SN04, SN05, SN10	Western Silivri coasts, Kadikoy
1.25m – 1.5m	SN11, SN13, SN15, SN17, SN18, SN19, SN30	Prince Islands, Tekirdag, M. Eregli, Bakirkoy, B. Cekmece,
1m - 1.25m	SN14	Yalova, Western Silivri coasts, Tuzla, Pendik
0.75m - 1m	SN16, SN20, SN29	Izmit and Gemlik Bays, Erdek Peninsula, Marmara Island
Negligible effect	SN01, SN12, SN21, SN22, SN24, SN25, SN26, SN27, SN28	Gelibolu, Lapseki, Degirmencik; Marmara, Avsa and Pasalimani Islands; innermost locations of Izmit and Gemlik Bays, northern Bosphorus, west of Erdek Peninsula, Sarkoy coasts

Table 4: The calculated maximum wave amplitudes with corresponding scenarios and the names of most affected coastal regions in Marmara Sea (the locations of affected coastal regions are shown on the map in Figure 3).

5





As of doing deterministic tsunami hazard assessme ne maximum wave height calculated within entire tsunami scenario database in this study does not exceed 2.2m. However, we should note that this value would be naturally higher in terms of having probabilistic tsunami hazard assessment, i.e. as Hancilar (2016) provides inundation maps resulting from probabilistic tsunami hazard analysis for a 10% probability of exceedance in 50 yr including the building numbers and types, lifeline systems and demographic data in Istanbul. Hancilar (2016) also highlights that the residential buildings at risk are mainly located in Kadikoy, Tuzla, Bakirkoy and Prince Islands where our study points out significant wave heights as well.

3. Conclusion

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This study, as the most comprehensive study so far to investigate the tsunami hazard due to earthquakes in the Marmara region, provides comprehensive tsunami scenario database including the estimated arrival times of first and maximum waves and maximum wave heights along Marmara coasts as well as integrated coastal impact maps. The water level fluctuations at TFPs and critical important locations in the Marmara Sea are also presented for each earthquake scenario. The results show that higher historical tsunami wave heights observed in Marmara Sea cannot be explained by only earthquake generated tsunamis and submarine landslides should be considered as the primary tsunami hazard component in the Marmara Sea. Depending on the landslide volume fundamentally, as proved by several previous studies, per sible tsunamis from submarine landslides in the Marmara Sea could be significantly higher than those from earthquakes and waveforms from all the coasts around the Marmara Sea indicate that other residential areas might have had a high risk of tsunami hazards from submarine landslides, which can generate higher tsunami amplitudes and shorter arrival times, compared to Istanbul (Latcharote et al., 2016). To address this issue, Necmioglu (2016) proposed a tsunami warning system in the Marmara region coupled with the existing earthquake early warning system, which could work without waiting for any focal mechanism parameter determination that may lead to underestimate tsunami hazards in the case of a strike-slip fault earthquake, due to the fact that submarine landslides could generate large tsunamis in the Marmara Sea. That coupling as well as the use of tsunami scenario database would enable RTEMC-KOERI to have most efficient operations in tsunami early warning in Marmara region without having any gaps in the system.

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