

Responses to the Comments by Editors and Reviewers

Dear Professor Töpfer:

Upon your recommendation, we have carefully revised Paper nhes-2019-156 entitled “Shear rate effect on the residual strength characteristics of saturated loess” after considering all the comments made by the reviewers.

We thank Professor Töpfer, Editor and two anonymous reviewers for their constructive comments. The manuscript has been significantly improved by incorporating their suggestions. The following are our point-to-point responses to their comments.

Responses to the Comments from Editor and reviewer 1

Comments from the editors and reviewers:

-Reviewer #1

This paper provides interesting data on soils obtained from three landslides and could be of interest to many readers. However, I feel some additional work is required prior to it being suitable for publication in the journal.

Reply: Thank you for your encouraging comments on our work.

1. Some more details need to be provided on the soils (Djg, Ydg and Dbz), e.g. particle size distribution curves.

Reply: Implemented. See Figure 2 of the revised manuscript.

2. Line 91, “their relationships” could be changed to “the relationship between the residual strength parameters”.

Reply: Implemented. See lines 98-99 of the revised manuscript.

3. Line 146, “crushed”, does this affect results?

Reply: Implemented. The procedure “crushed” would not affect results. The purpose of crushing with a mortar and pestle is to disintegrate aggregate. Crushing samples has been found suitable to determine the residual strength of the remoulded soils (Stark et al., 2005). This should be done with care so as not to destroy silty-dominated loess. See details in lines 154-159 of the revised manuscript.

4. Line 169, to keep consistency with the text body, change “moisture water content” to “moisture content”, please revise it.

Reply: Implemented. See line 184 of the revised manuscript.

5. Lines 172-173, there are 2 main types discussed in the literature, the Bromhead device and the IC/NGI device, which one is this? Please point out it in the paper.

Reply: Implemented. SRS-150 used in this study is a type of Bromhead ring shear apparatus. See lines 197-198 in the revised manuscript.

6. Line 198, please give more detail about compaction.

Reply: Implemented. See lines 229-232 of the revised manuscript.

7. It seems that you do not need to mention the shearing process in lines 203-204 again since you have mentioned the procedure in Lines 176-177.

Reply: Implemented. We have deleted the contents in original lines 203-204 according to the review's suggestion.

8. Line 207, "the sampling rate was increased to 1 min", please check the sampling rate unit.

Reply: Implemented.

9. Line 209, in my opinion, "the samples were subjected to shear" could be changed to "the samples were subjected to shearing".

Reply: Implemented. See line 241 of the revised manuscript.

10. Lines 209-210, how do you define the residual state was achieved?

Reply: Implemented. In this study, following the Bromhead (1992), the residual state was defined when a constant shear stress is obtained for more than half an hour. See lines 242-243 in the revised manuscript.

11. Lines 238-239, The authors should clearly define what are low and high shear rate.

Reply: Implemented. See lines 272-274 of the revised manuscript.

12. Lines 375-376, the authors do not need to write Liquid limit (LL) again since you have mentioned that in lines 372-373, just use LL in lines 375-376.

Reply: Implemented. See line 453 of the revised manuscript.

13. Line 400, change "Figs. 7, 8 and 9" to "Figs. 7-9", please revise it.

Reply: Implemented. See line 403 of the revised manuscript.

14. In Table 1, units missing on the header. Feel PSD curve is necessary. Please revise it.

Reply: Implemented. See Table 1 in the revised manuscript.

With regards to PSD curve, we have added the PSD curves in the revised manuscript, see Figure 2 in the revised manuscript.

15. The use of the English language needs some work. I really recommend the authors to send the manuscript to be reviewed thoroughly by a native English speaker.

Reply: Implemented. The revised manuscript has been reviewed thoroughly by a native English speaker to improve the grammar and readability.

-Reviewer #2

This paper deals with the effect of the shear rate on the residual shear strength of loess from three landslides by using a ring shear apparatus. Overall, this is an interesting manuscript because the topic can be considered of large significance for international researches in the field; however, this manuscript needs some important improvements to get it into a position to be acceptable for publication. Thus, a major revision is recommended. My critical review is summarized in the following sentences:

Reply: Thank you for your encouraging comments on our work.

1. The title could be more informative although it is pertinent and understandable

Reply: Implemented. The title has been changed as “Shear rate effect on the residual strength characteristics of saturated loess in naturally drained ring shear tests” according to the review’s suggestion. See title in the revised manuscript.

2. The abstract should be more precise and clear, although the most important results have been mentioned (Please, find the file attached for details).

Reply: Implemented. See lines 29-35 of the revised manuscript.

3. Authors should better emphasize the aim, importance and results of this study, and why it should be considered as relevant to be published in an international journal.

Reply: Implemented. See lines 70-89 of the revised manuscript.

4. The introduction provides relevant background information. Important scientific publications, on which the work is based, are cited but some recent original papers are not considered.

Reply: Implemented. We have cited some recent original papers, see lines 45-46, 73-75, 85 of the revised manuscript.

5. Geological setting and sampling sites if, on the one hand, require a brief description, on the other hand, should contain all the useful information for the purpose of the work.

Reply: Implemented. We have cited relevant references to describe the geological setting and sampling sites, see lines 110-111, 126, 132-133 of the revised manuscript.

6. Congruent bibliographic references are missing. Please, find the file attached for details

Reply: Implemented. See references in the revised manuscript.

7. Description of the materials used (grain size distribution, percentage and mineralogy of the clay fraction, plasticity of fine) is very important and can explain some discrepancies between different interpretations. Please, find the file attached for details

Reply: Implemented. See lines 71-76 of the revised manuscript.

8. Description of the method used in this study should be detailed and complete.

Reply: Implemented. See lines 218-220, 223-225, 229-232, 241-243 of the revised manuscript.

9. What does it mean for low or high shear rate and low or high effective normal stress?
Please, find the file attached for details

Reply: Implemented. See lines 272-274, 313-315 of the revised manuscript.

10. Results and discussion may be combined into a single section to avoid repetitions in the discussion, which would thus be more interesting and complete, also with references to earlier or contemporary studies relevant to the topic.

Reply: Implemented. See lines 253, 263-264, 307-308, 341-348, 354, 401-423, 440-442 of the revised manuscript.

11. Discussion of the results should include aspects related to dilatancy and **critical state**

Reply: With regards to the dilatancy effect of the samples, we have added the relevant content in the manuscript, see lines 283-289 of the revised manuscript.

With regards to the critical state of the loess, we did not measure normal displacement of loess samples, therefore the critical state of samples was not discussed in the study.

12. Conclusions summarize the main findings of the experimental research but could describe their significance or implication, in light of what was already known about the subject of the study, and present fresh insights or possible new ways of approaching research questions

Reply: Implemented. See lines 488-496, 506-511 of the revised manuscript.

13. Text, tables citations and references should be formatted according to the journal's instructions

Reply: Implemented. See Tables and references in the revised manuscript.

14. A thorough revision of the text with the help of a native English speaker is suggested.

Reply: Implemented. The manuscript has been revised with the help of a native English speaker.

1 Shear rate effect on the residual strength characteristics of saturated loess in naturally
2 drained ring shear tests

3

4 Baoqin Lian^{a,b}, Jianbing Peng^{a*}, Qiangbing Huang^a

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6 ^aCollege of Geological Engineering and Surveying, Chang'an University, Key
7 Laboratory of Western China Mineral Resources and Geological Engineering, Xi'an
8 710054, China

9

10 ^bDepartment of Geology & Geophysics, Texas A&M University, College Station, TX
11 77843-3115, United States

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13 *Corresponding author: Jianbing Peng (dicexyl@gmail.com)

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19 **Abstract**

20 Residual shear strength of soils is an important soil parameter for assessing the
21 stability of landslides. To investigate the effect of the shear rate on the residual shear
22 strength of loessic soils, a series of **naturally drained** ring shear tests were carried out
23 on loess from three landslides at two shear rates (0.1 mm/min and 1 mm/min).
24 **Experimental** results showed that the shear displacement to achieve the residual stage
25 for specimens with higher shear rate was greater than that of the lower rate; both the
26 peak and residual friction coefficient became smaller with increase of shear rate for
27 each sample; at two shear rates, the residual friction coefficients for all specimens
28 under the lower normal stress were greater than that under the higher normal stress.
29 **Moreover, specimens with almost the same low fraction of clay (CF) showed similar**
30 **shear rate effect on the residual friction coefficient with normal stress increasing,**
31 **whereas specimen with high CF (24%) showed the contrast tendency, indicating that**
32 **such effect is closely associated with CF.** The tests results revealed that the difference
33 in the residual friction angle ϕ_r at the two shear rates, $\phi_r(1) - \phi_r(0.1)$, **under each**
34 **normal stress level were either positive or negative values of which the maximum**
35 **magnitude is about 0.8° .** However, the difference $\phi_r(1) - \phi_r(0.1)$ **determined** under all
36 normal stress levels was negative, which indicates that the residual shear parameters
37 reduced with the increasing of the shear rate in loess area. Such negative shear rate
38 effect on loess could be attributed to a greater ability of clay particles in specimen to
39 restore broken bonds at low shear rates.

40 **Keywords:** Loess; Residual shear strength; Ring shear test; Shear rate; Residual shear
41 parameter

42

43 **1. Introduction**

44 Residual shear strength of soil is of great significance for evaluating the stability
45 for the slip surface of first-time landslides as well as reactivated landslides (Bishop et
46 al., 1971; Mesri and Shahien, 2003; Tiwari and Latha, 2019; Li et al., 2017). The
47 residual strength of soils is defined as the minimum constant value of strength along
48 the slip plane, in which the soil particles are reoriented and subjected to sufficiently
49 large displacements in relatively low shear rate (Skempton, 1985).

50 Numerical studies have been done to assess the residual strength through the
51 laboratory tests using ring shear tests and reversal direct shear tests (Vithana et al.,
52 2012; Summa et al., 2018; Moeyersons et al., 2008; Chen and Liu, 2013; Summa et al.,
53 2010). It is a generally accepted fact that the measurement of the residual strength is
54 most preferred done with a ring shear test since it allows the soil specimen be sheared
55 at unlimited displacement which can simulate the field conditions more accurately
56 (Sassa et al., 2004; Tiwari and Marui, 2005; Lupini et al., 1981; Bhat, 2013). Until now,
57 great efforts have been paid to the study of the shear rate effect on the minimum value
58 of clay or sand strength at residual states (Li et al., 2017; Tika and Hutchinson,
59 1999; Suzuki et al., 2007; Grelle and Guadagno, 2010; Lemos, 1985; Tika,
60 1999; Morgenstern and Hungr, 1984). As a result, the residual strength of clay or sand
61 under the effect of shear rate has been made relatively clear. However, compared with
62 the results of tests on clay or sand, understanding of the shear characteristics of silty
63 soil, such as loess, is not yet complete. As pointed out by Ding (2016), some drained
64 ring shear tests have concluded that the increase in shear rate causes the residual

65 strength of loess to increase. On the contrary, Kimura et al. (2014) reported that the
66 residual strength of Malan loess decreases with the increase of shear rate. Furthermore,
67 Wang et al. (2015) found that the effect of shear rate on residual strength of loess is
68 closely associated with the normal stress levels, and the change in residual strength of
69 loess samples under high normal stress levels is small in ring shear tests.

70 Therefore, some inconsistent or even opposite results have been reported in the
71 ring shear tests on loess above, which maybe attributed to the differences in the grain
72 size distribution and mineral composition of the different material tested in previous
73 studies (Ajmera et al., 2012). Particularly, this discrepancy maybe due to the
74 difference in quantity and mineralogy of clay fraction (Nakamura et al., 2010;Li et al.,
75 2013). Therefore, the previous studies on the effect of shear rate on residual strength
76 of loess implied that there is still a lack of experimental data on this topic. From the
77 above investigations, it can be concluded that the effect of the shear rate on the
78 residual strength of the loess is not fully understood and needs further scrutiny.
79 Moreover, it should be noted that the residual strength parameters (friction angle)
80 obtained from using different shear rates may be adopted to provide a guide for
81 designing some precision engineering which require high accuracy of the design
82 parameters, thus, the effect of the shear rate on the residual strength of soils should be
83 fully investigated to determine the parameters with high reliability. In addition,
84 residual strength parameters of soil play a key role in assessing the stability analysis
85 of landslides (Xu et al., 2018;Wesley, 2018). Therefore, accurate determination of the
86 residual strength parameters and their dependence on the shear rate may affect the

87 stability evaluation of landslides. Thus, it is necessary to study the change of residual
88 strength of loess with shear rate in order to have a good understanding of the suitable
89 approach for the residual strength parameters measurement.

90 In this backdrop, to clarify the residual shear characteristics of loess under the
91 effect of the shear rate, a series of naturally drained ring shear tests were conducted on
92 loess obtained from three landslides **on the Loess Plateau in China** at two shear rates
93 (0.1 mm/min and 1 mm/min). The residual shear characteristics of loess at the
94 residual state was examined. Considering that shear strength of loess reduces with
95 moisture content (Picarelli, 2010;Zhang et al., 2009;Dijkstra et al., 1994), ring shear
96 tests were conducted on saturated loess samples corresponding to the worst condition
97 in field engineering. Furthermore, this study investigated the change in the residual
98 strength parameters of loess at different shear rates and **the relationship between the**
99 **residual strength parameters** with the normal stress in naturally drained ring shear
100 tests as well.

101

102 **2. Geological setting of landslide sites**

103 Soil samples from three landslides in the northwest of China were selected in this
104 study. Soil samples used for the ring shear tests and index measuring tests
105 predominantly consist of loess deposits and were collected in a disturbed condition.
106 For convenience, the names of landslide sites were abbreviated into Djg, Ydg, and
107 Dbz. Fig. 1 shows the study sites and some views of the landslides.

108 **Dingjiagou landslide (Djg)**

109 The Djg landslide, located at the mouth of Dingjia Gully in Yan'an of China, is

110 geologically composed of upper loess and lower sand shale in the **Triassic Yanchang**
111 formation (**She, 2015**). The dustpan-shaped landslide is inclined to the east, with its
112 inclination 75.85° . The landslide is 350 m in width, 180 m in length, 70 m in elevation.
113 The average thickness of slip mass is around 20 m, and the volume of landslide
114 totaled approximately $105 \times 10^4 \text{ m}^3$. The slip mass is mainly constituted by loess,
115 whereas the sliding bed consists of sand shale in Yan-chang formation. The thickness
116 of the sliding zone varied from 30 to 50 cm. The front lateral region of the main slide
117 section of the Dlg landslide, where the sampling was performed, was found to be silty
118 clay.

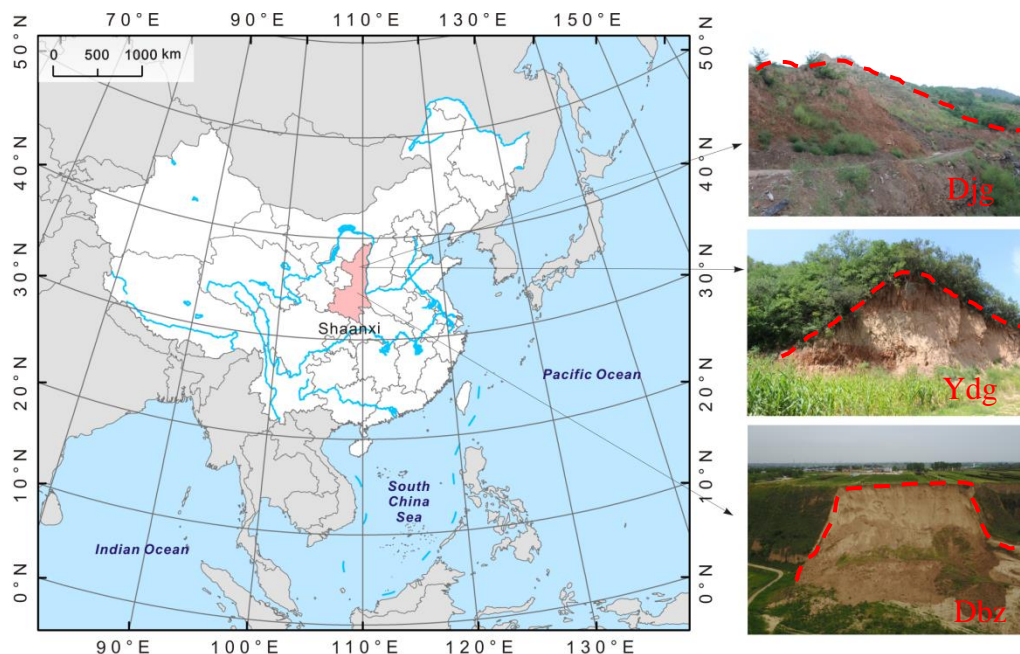
119 **Yandonggou landslide (Ydg)**

120 The Ydg landslide, located in the Qiaogou town of Yan'an in Shaan xi province of
121 China. The top and the toe altitude of the landslide are about 1165 m and 1110 m
122 above the sea level, with the height difference between the toe and the top of landslide
123 about 55 m. The slides have well-developed boundaries with the main sliding
124 direction of 240° and slope angle of 30° . From the landslides profile, the sliding
125 masses from top to bottom were classified by late Pleistocene (Q_3) loess, Lishi (Q_2)
126 loess and clay soil, respectively (**Zhang et al., 2006**). Multiple landslides had occurred
127 in this site, and the soil samples used in this study were collected from Q_2 loess
128 stratum within the slide ranged from 4.5 m to 18 m in height.

129 **Dabuzi landslide (Dbz)**

130 The Dbz landslide located in the middle part of Shaanxi province (about E
131 $108^\circ 51' 36''$ and N $34^\circ 28' 48''$), China, which is a semi-arid zone dominated by loessic
132 geology (**Yan et al., 2015**). In this region, the investigated site is classified as a typical

133 loess tableland with Quaternary stratum (Ma et al., 2019). The sedimentary losses in
134 this area are grey yellow, and the exposure stratum in this area has been divided into
135 two stratigraphic units, namely, the upper Malan (Q₃) loess and the lower Lishi (Q₂)
136 loess, of which the Q₃ loess is younger. The Q₃ loess is closest to the surface and is up
137 to approximately 12 m thick, while the thickness of Q₂ loess may reach an upper limit
138 of about 50 m (Leng et al., 2018). The loess in this area have well-developed vertical
139 joints (Sun et al., 2009). The travel distance and the maximum width of the slip mass
140 are roughly estimated to be 122 m and 133 m, respectively. The armchair-shaped
141 landslide shows an apparent sliding plane, with an area of approximately 15,660 m²
142 and about 66.25 m maximum difference in elevation. The main direction of this
143 landslide is approximately 355°. The exposed side scarp of the landslide, where the
144 sampling was done, was found to be entirely in the Q₂ loess stratum.



145
146 Fig. 1. Location of study sites and some views of landslides

147 Notes: Red dashed lines in the Fig. 1 represent landslide boundary.

148 3. Experimental scheme

149 **3.1. Testing sample**

150 The fact that the residual shear strength is independent of the stress history has
151 been reported by many researchers (Bishop et al., 1971; Stark et al., 2005; Vithana et
152 al., 2012). Thus, disturbed loess samples from each landslide weighing about 25 kg
153 were collected to investigate the residual shear strength.

154 The soil samples were air-dried, and then crushed with a mortar and pestle as it
155 has been reported that crushing samples were suitable to determine the residual
156 strength of the remoulded soils (Stark et al., 2005). It was found that small lumps may
157 exist in air-dried samples, which may be too big for the cell, so lumps were crushed in
158 order to make sample uniform. This should be done with care so as not to destroy
159 silty-dominated loess. After that, soil samples were processed through 0.5 mm sieve.
160 Distilled water was then added to the soil samples until saturated water content were
161 obtained. The physical parameters such as natural moisture content (*in-situ* moisture
162 content), specific gravity, bulk density, plastic limit, and liquid limit were determined
163 in accordance with the Chinese National Standards
164 (CNS) GB/T 50123-1999 (standards for soil test methods) (SAC, 1999), but clay size
165 was defined to be less than 2 μm followed ASTM, D 422 (ASTM, 2007). Each soil
166 sample was separated into clay (sub 0.002 mm), silt (0.002-0.075 mm), and sand
167 (0.075-0.5 mm) fractions. The physical indexes of the soil are listed in Table 1.

168 The grain size distribution of soil was measured using a laser particle size
169 analyzer Bettersize 2000 (Dandong Bettersize Instruments Corporation, Dandong,
170 China). The sieved soil samples were used to determine particle size distribution. In

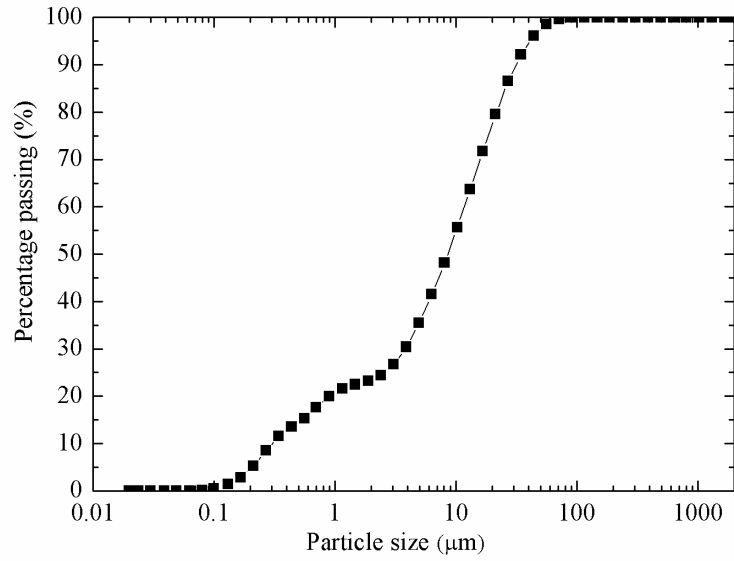
171 this study, soil samples were treated with sodium hexaphosphate, serving as a
 172 dispersant, to disaggregate the bond between the particles. **The particle size**
 173 **distribution curves of soils were shown in Fig. 2.** The results show that the clay
 174 fraction in Djg landslide soil (24%) is more than two times than that from Ydg (9%)
 175 and Dbz (9.1%). Furthermore, the particle size analysis illustrated that the percentage
 176 of silt-sized soil in three landslides ranged from 75.66% to 87.4%. In addition, Ydg
 177 landslide soil consists of the greatest percentage of the sand fraction which reaches up
 178 to 10.55% (Table 2 and Fig. 2).

179
 180
 181
 182

183 **Table 1.** Physical parameters of slip-zone loess.

<i>sites</i>	ρ_d	<i>W</i>	ρ	G_s	W_L	W_p	Grain size fractions (%)			
							<0.002mm	0.002-0.005mm	0.005-0.075mm	0.075-0.5mm
Djg	1.74	19.5	2.08	2.65	36	20	24	11.48	64.18	0.34
Ydg	1.47	18	1.74	2.71	33	19	9	5.28	75.17	10.55
Dbz	1.48	16	1.72	2.70	32	21	9.1	6.4	81	3.5

184 Notes: ρ_d = dry density (g/cm³); w=**moisture content** (%); ρ = bulk density (g/cm³); G_s
 185 = specific gravity; W_L =liquid limit (%); W_p = plastic limit (%).

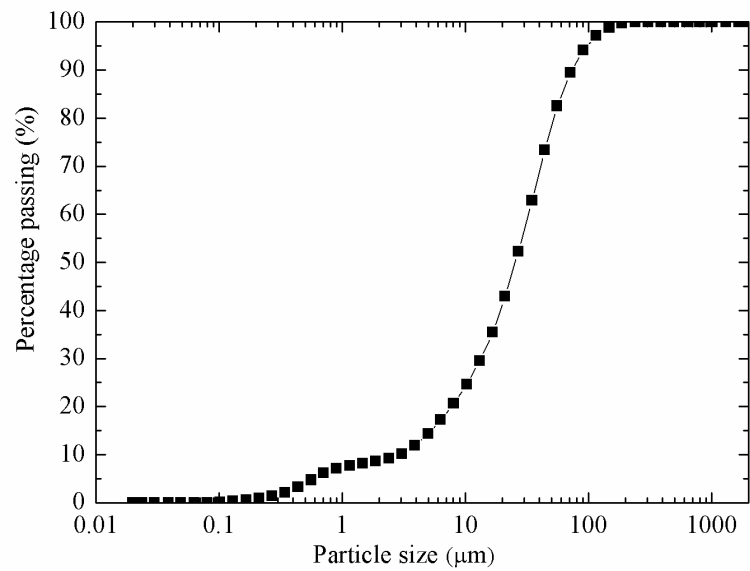


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(a) Particle size distribution curve of soil obtained from DJG

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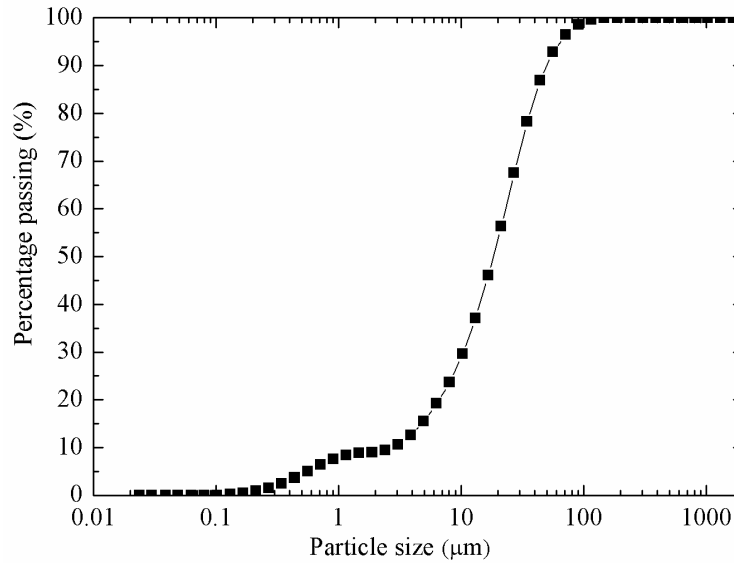


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(b) Particle size distribution curve of soil obtained from YDG

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(c) Particle size distribution curve of soil obtained from DBZ

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194 Fig. 2. Particle size distribution curves.

195

196 3.2. Testing apparatus

197 An advanced ring shearing apparatus (SRS-150), the Bromhead-type ring shear

198 apparatus, manufactured by GCTS (Arizona, USA) was adopted in ring shear tests

199 and the photos of apparatus were shown in Fig. 3, which consists mainly of a shear

200 box with an outer diameter of 150 mm, an inter diameter of 100 mm and the maximal

201 sample height of 250 mm. The shearing box consists of the upper shear box and the

202 lower shear box. In the shearing process, the upper shear box keeps still while the

203 lower one rotates. The apparatus, which provides an effective specimen area of 98

204 cm², is capable of shearing the specimen for large displacements. The annular

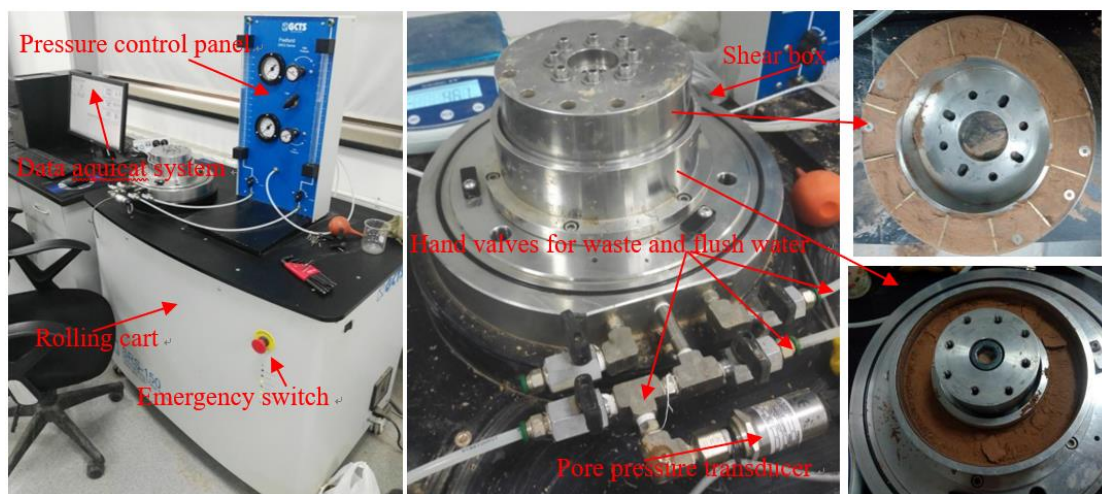
205 specimen is confined by inside and outside metal rings. Moreover, the specimen is

206 confined by bottom annular porous plates and top annular porous plates in which have

207 sharp-edged radial metal fins which protrude vertically into the top and bottom of the

208 specimen at the shearing process. Two annual porous plates were used to provide

209 drainage condition in the test following previous research (Stark and Vettel, 1992).
210 The normal stress, shear strength and shear displacement can be monitored by
211 computer in shearing process. The measurement features of the ring shear apparatus
212 employed in this study are described as follows: shear rate range from 0.001 degrees
213 to 360 degrees per minute, 10 kN axial load capacity, 300 Nm continuous torque
214 capacity, maximum normal stress of 1000 kN/m².



215
216 Fig. 3. Ring shear apparatus (SRS-150)

217 3.3. Testing procedure

218 This study was comprised of 3 groups of test results, in which 24 remolded
219 saturated loess samples were sheared with normal stress ranging from 100 to 400
220 kN/m² under two shear rates. In present study, reconstituted samples of the sub 0.5
221 mm soil fractions were prepared for the shear tests as it was reported that the residual
222 strength of the soil was unaffected by its initial structure (Vithana et al., 2012; Bishop
223 et al., 1971). Consolidated drained (CD) tests with single-stage shear was performed.
224 Here, the single-stage shear means shearing the sample under effective pressure or
225 stress conditions after the consolidation of the sample. Specimens were first prepared
226 by adding distilled water to the air-dried soil until the saturated moisture contents

227 were obtained. Then, specimens were kept in a sealed container for at least one week
228 to fully hydrate. Afterwards, specimens are reconstituted in the ring-shaped chamber
229 of the apparatus by compaction. During the compaction process, samples were
230 divided into equal five parts and each part was poured into the shear box and
231 compacted. Samples with a height of 2.5 cm in this study were prepared in five layers
232 of equal height to achieve the required density. The specimen was then consolidated
233 under a specific effective normal stress in a range of 100 kN/m² to 400 kN/m² until
234 consolidation was achieved. In this study, consolidation was completed when the
235 consolidation deformation was smaller than 0.01 mm within 24 hr (Kramer et al.,
236 1999;Shinohara and Golman, 2002). In ring shear tests, the normal stress at the
237 shearing was the same as at consolidation stage. Shear strength of loess specimen was
238 recorded at intervals of 1s before the peak shear strength, after the peak, the sampling
239 rate was increased to 1 min.

240 In this study, ring shear tests were performed in a single stage under naturally
241 drained condition and the samples were subjected to shearing until the residual state
242 was achieved. Following the Bromhead (1992), the residual state was defined when a
243 constant shear stress is obtained for more than half an hour. Drained condition of the
244 shearing process is provided by two porous stones attached on the top and the bottom
245 platen of the specimen container. As for soil specimens with low permeability, the
246 rate of excess pore pressure generation in the shear box may exceeded that of
247 pore-pressure dissipation, this type of condition is identified as naturally drained
248 condition in previous studies (Okada et al., 2004). Furthermore, Tiwari (2000)

249 asserted that it was acceptable to use a shear rate below 1.1 mm/min to simulate the
250 field naturally drained condition. Thus, shear rates of 0.1 mm/min and 1 mm/min
251 were used in this study to simulate the naturally drained condition of the slip zone
252 soils.

253 **4. Results and discussion**

254 Twenty -four specimens were tested to investigate the residual shear
255 characteristics of the saturated loess in the ring shear apparatus. Residual shear
256 strength of loess was determined following the research conducted by Bromhead
257 (1992) who pointed out that the residual stage is attained if a constant shear stress is
258 measured for more than half an hour. Tests results are shown in this section.

259 **4.1. Shear behavior**

260 Figs. 4(a)- 6(a) show the typical shear characteristics of the loess (shear rate of 0.1
261 mm/min and 1 mm/min) obtained from three different locations, where, the shear
262 stress is plotted against the shear displacement. It is a widely accepted fact that
263 normal stress has effect on the shear behavior of the soil (Wang et al., 2019;Eid,
264 2014;Kimura et al., 2015;Stark et al., 2005;Eid et al., 2019), thus, the shear behavior
265 of samples at the peak and residual stages, where, the determined peak friction
266 coefficient as well as residual friction coefficient are plotted in Figs. 4(b)-6(b) against
267 the corresponding effective normal stresses as well. The friction coefficient is defined
268 as the shear stress divided by the effective normal stress.

269 Figs. 4(a)-6(a) demonstrate that shear stress increases dramatically within small
270 shear displacement and then reduces with shear displacement, until residual

271 conditions were achieved at large displacements. Furthermore, it is obvious that the
272 peak strength and the residual strength of samples with high shear rate (shear rate
273 equal to 1 mm/min) are almost smaller than that of the samples with low shear rate
274 (shear rate equal to 0.1 mm/min). It can be found that shear displacement to achieve
275 the residual stage for specimens with high shear rate is greater than that of the low
276 rate. For example, the minimum shear displacements for attaining residual condition
277 for Djg specimens with low and high shear rate were about 360 mm and 650 mm,
278 respectively. Under the shear rate of 0.1 mm/min and 1 mm/min, Ydg specimens need
279 approximately 80 mm and 1,400 mm displacement to achieve residual stage. However,
280 Dbz specimens require about 40 mm and 60 mm displacement to reach residual
281 condition for low and high shear rate, respectively.

282 In Figs. 4(a)- 6(a), a clear drop can be seen, at any normal stress, for specimens
283 obtained from all sites. During shearing, as reported by Terzaghi et al. (1996), strain
284 softening exhibits a dilative behavior for soils. It is seen that the shear behavior is
285 non-linear against the shear displacement. The loess in Djg, Ydg and Dbz exhibited
286 the typical shear stress and shear strain relationship, i.e., the strain softening behavior
287 for a given shear rate (Figs. 4(a)-6(a)). As seen in Figs. 4(a)-6(a), the lower shear rate
288 results in a more obvious dilation effect during the shearing process with a specific
289 normal stress. It is obvious that Djg specimens showed greater peak-post drop than
290 that of Ydg and Dbz specimens. For example, at the normal stress of 100 kN/m², Djg
291 samples show approximately 47.3% and 36.8% decrease from the peak friction
292 coefficient to the residual friction coefficient at low and high shear rates (Fig. 4(b)),

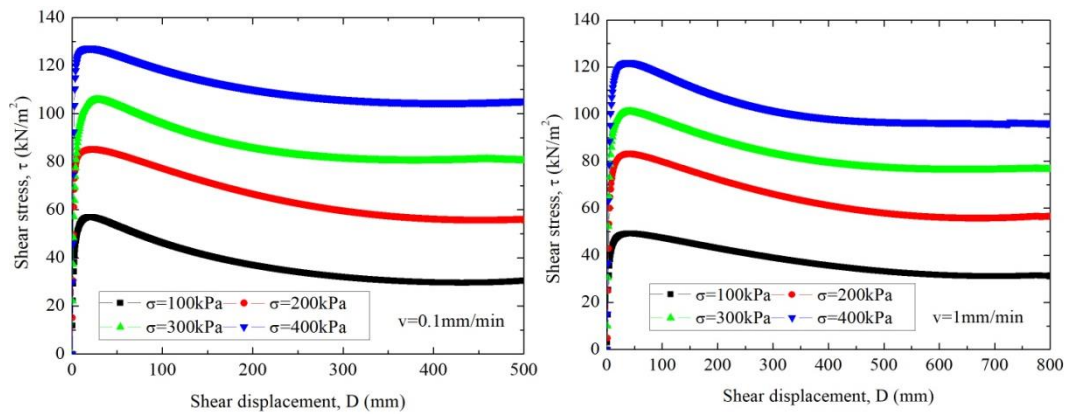
293 respectively, which is greater than in the Ydg samples (about 9.8% and 10.3% in Fig.
294 5(b)) and Dbz samples (about 2.4% and 3.2% in Fig. 6(b)). In Djg samples, an
295 obvious slickenside was observed on the shear surface (Fig. 7). This phenomenon
296 indicates a high degree of reorientation of platy clay minerals parallel to the direction
297 of shearing. In Figs. 4(b)- 6(b), on average, it was found that the decrease in the
298 friction coefficient from the peak strength in the Djg sample is almost 18.1% and
299 21.3% for the sample consolidated at normal stress of 400 kN/m² under the low and
300 high shear rate (Fig. 4(b)), while such reduction in friction coefficient in Ydg sample
301 are only about 4.1% and 4.8% (Fig. 5(b)). Furthermore, under the low and high shear
302 rate, the friction coefficient reduction in Dbz samples are only approximately 5.6%
303 and 6.0% (Fig. 6(b)). Skempton (1985) reported that the strength of soils falls to the
304 residual value in ring shear tests, owing to reorientation of platy clay minerals parallel
305 to the direction of shearing. Based on the conclusion that the post-peak drop in
306 strength of normally consolidated soil is only due to particle reorientation after the
307 peak strength (Skepmtton, 1964;Mesri and Shahien, 2003;Habibbeygi and Nikraz,
308 2018), the results demonstrated that the Djg landslide soil existed the greater particle
309 reorientation compared with that of other two landslide soils.

310

311 **4.2. Effect of normal stress on the friction coefficients**

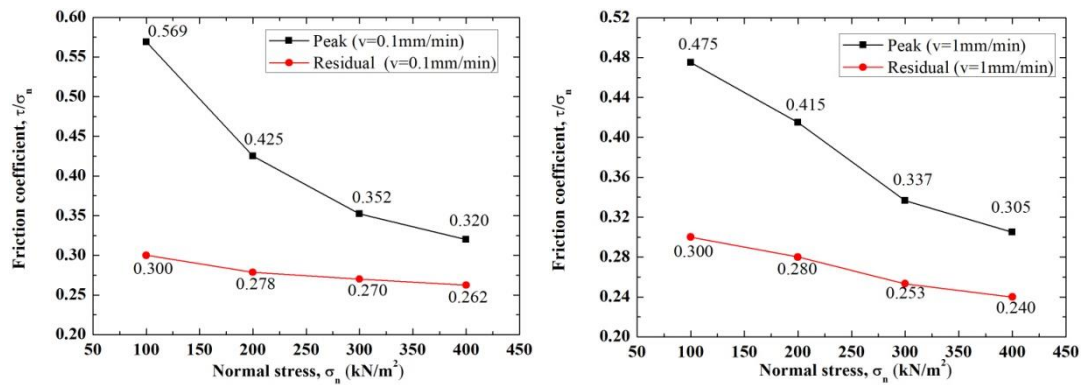
312 It can be seen from the Figs. 4(b)-6(b) that the friction coefficients (peak and
313 residual) are higher at low effective normal stress levels (effective normal stress equal
314 or less than 100 kN/m²) compared with that at high normal stress (effective normal

315 stress between 200 and 400 kN/m²). For example, with normal stress increasing from
 316 100 kN/m² to 400 kN/m², the peak and residual friction coefficient of Djg landslide
 317 soils at the shear rate of 0.1 mm/min reduce from 0.569 to 0.32 and from 0.3 to 0.262
 318 (Fig. 4(b)), respectively. Similarly, results obtained from other two landslides loess
 319 also show that the friction coefficients decrease nonlinearly with normal stresses (Figs.
 320 5(b) and 6(b)). Furthermore, specimens with shear rate of 0.1 mm/min attained greater
 321 friction coefficients than that with shear rate of 1 mm/min (Figs. 4(b)-6(b)).



322

323 (a) Relationship between shear stress and shear displacement

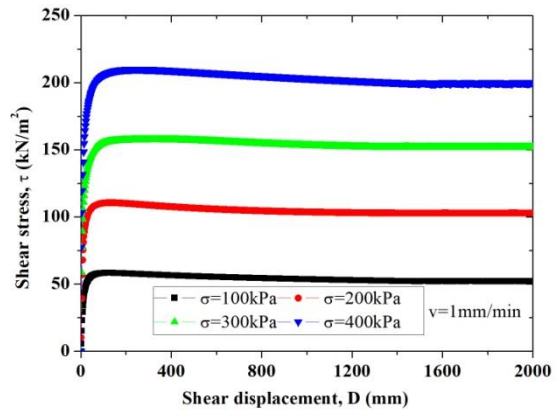
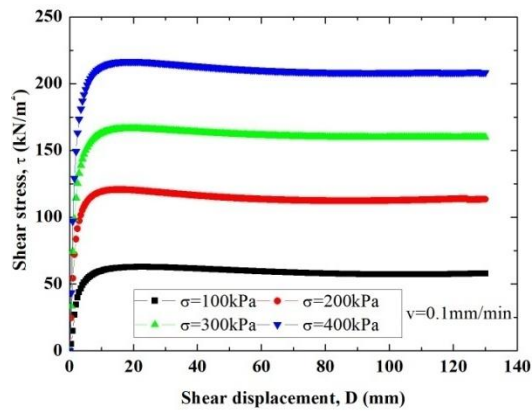


324

325 (b) Relationship between friction coefficient and normal stress

326 Fig. 4. Shear behavior characteristics of Djg soil samples

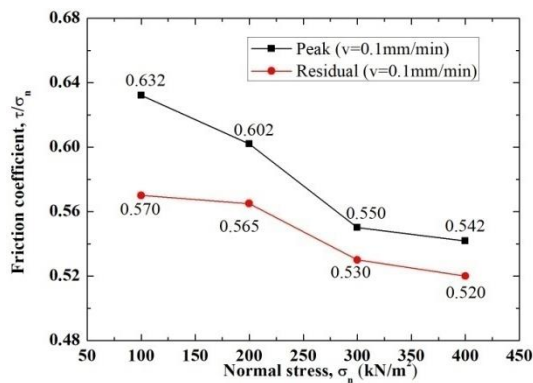
327



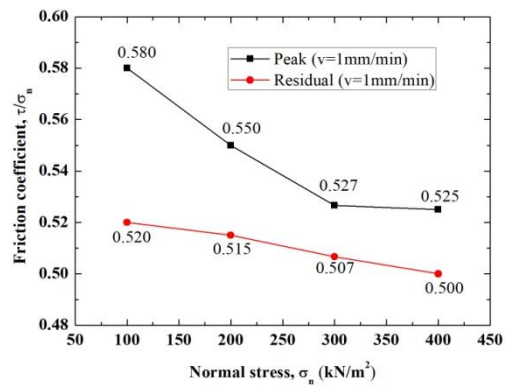
328

329

(a) Relationship between shear stress and shear displacement



330

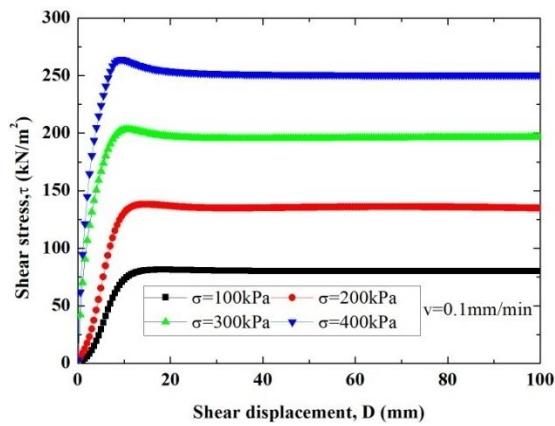


331

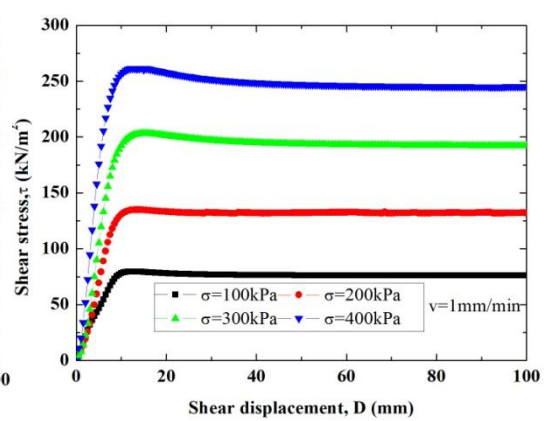
(b) Relationship between friction coefficient and normal stress

332

Fig. 5. Shear behavior characteristics of Ydg soil samples

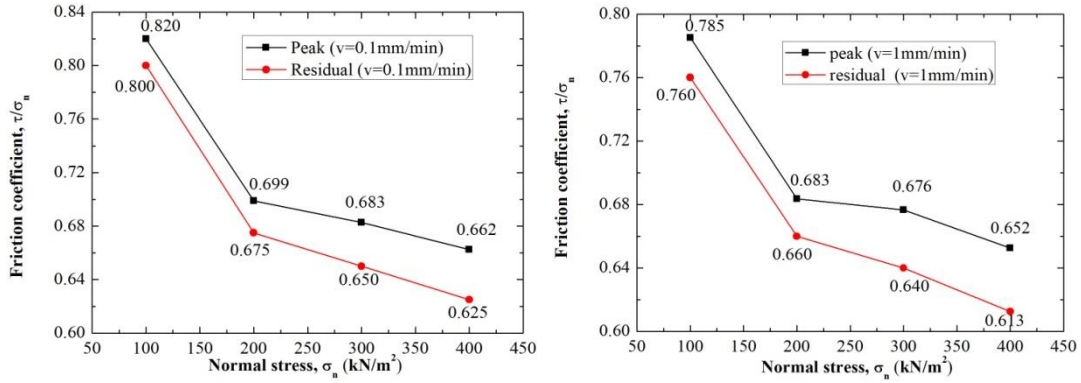


333



334

(a) Relationship between shear stress and shear displacement



(b) Relationship between friction coefficient and normal stress

Fig. 6. Shear behavior characteristics of the Dbz soil samples

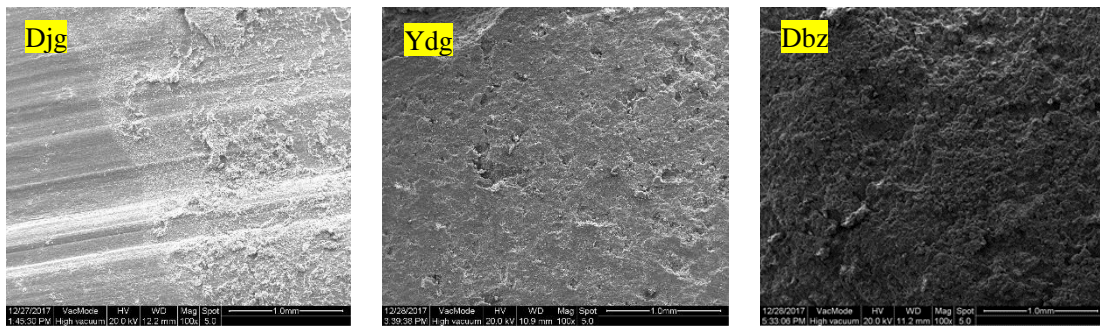


Fig. 7. SEM photographs of the shear surface of loess samples (100 magnification)

4.3. Effects of shear rate on residual strength parameter

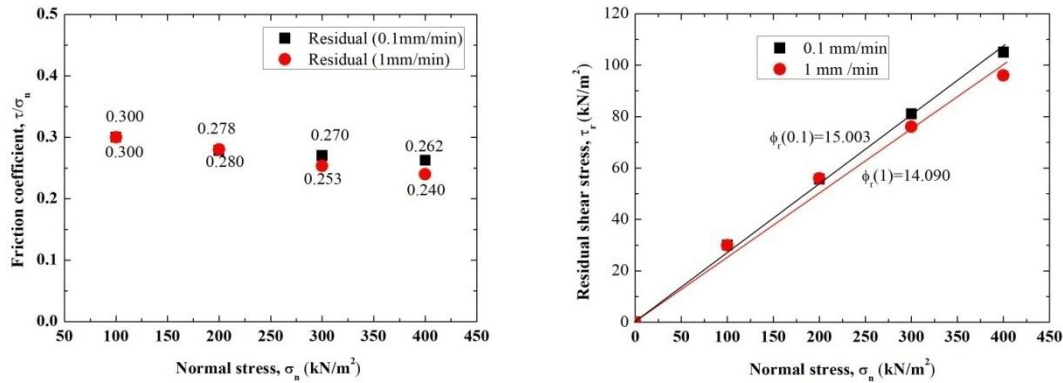
Following the previous study reported by (Eid et al., 2019; Terzaghi et al., 1996), the maximum value during shear process can be the peak shear stress, whereas the minimum value can be the minimum shear stress. Correspondingly, the maximum value can be referred to as the peak shear strength, whereas the minimum value can be referred to as the residual shear strength that resulted from particle rearrangements after a large shear displacement. Furthermore, the peak and residual strength parameters are determined by using Mohr–Coulomb failure criterion (Terzaghi et al., 1996). In this study, the residual strength parameters were analyzed and discussed.

For the samples described above, Figs. 8-10 show the relationships between the residual friction coefficient and the normal stress, and the residual strength parameters.

351 The residual friction coefficient is plotted against the normal stress. The residual
352 friction coefficient is defined as the residual shear strength divided by normal stress. It
353 has been recognized that the shear strength parameters including cohesion and friction
354 angle (Terzaghi, 1951; Stark Timothy et al., 2005; Pakbaz et al., 2018). However,
355 according to the previous studies, the residual angle of soils varies depended on the
356 soil properties as well as the magnitude of normal stress provided the residual
357 cohesion of soil is zero (Kimura et al., 2014; Skempton, 1964). Thus, in this study, the
358 residual frictions are calculated by Coulomb's law assumed the residual cohesion is
359 zero following the previous studies (Skempton, 1985). The residual strength
360 parameters were defined as $\phi_r(0.1)$ and $\phi_r(1)$ at the low shear rate and high shear rate,
361 respectively. And the difference between the residual friction angles at two shear rates
362 was defined as $\phi_r(1) - \phi_r(0.1)$. Comparatively, the residual friction coefficient was
363 defined as $\tau_r/\sigma_n(0.1)$ at the low shear rate and $\tau_r/\sigma_n(1)$ at the high shear rate,
364 respectively. Furthermore, the difference between the residual friction coefficients
365 was defined as $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$. Table 2 summarized the residual shear
366 parameters of the landslide soils.

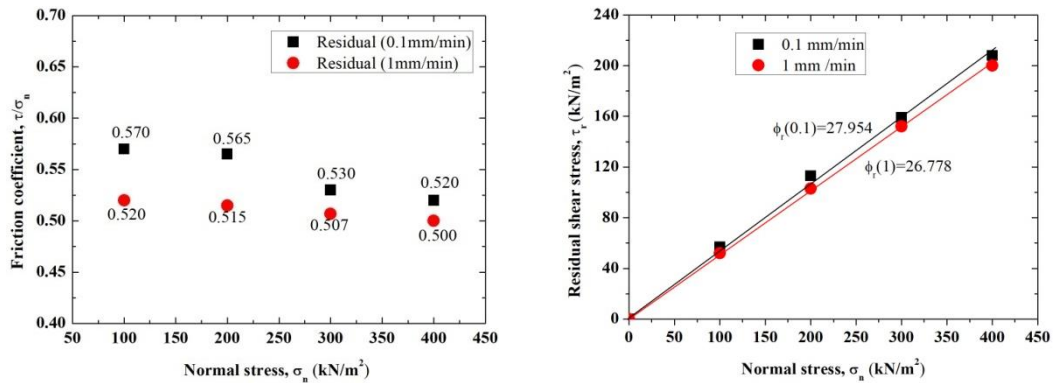
367 Fig. 8 shows that the residual friction coefficients are relatively low in Djg
368 samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ at the normal stress of 100 kN/m²
369 to 400 kN/m² ranged from 0.3 to 0.262 and from 0.3 to 0.24, respectively. The
370 difference between the friction coefficients, $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$, at each normal
371 stress level are varied in a range of -0.022 to +0.002. For the difference between the
372 residual friction angles, $\phi_r(1) - \phi_r(0.1)$, ranged from -1.212° to +0.079° (Table 2). For

373 normal stress above 200 kN/m², the residual friction coefficient $\tau_r/\sigma_n(0.1)$ was found
 374 to be greater than the residual friction coefficient $\tau_r/\sigma_n(1)$. For this sample, residual
 375 friction coefficients show a slight decrease with the shear rate for normal stress above
 376 200 kN/m².



377
 378 Fig. 8. Relationships between residual shear stress and normal stress, and residual
 379 strength parameter for Djg soil sample

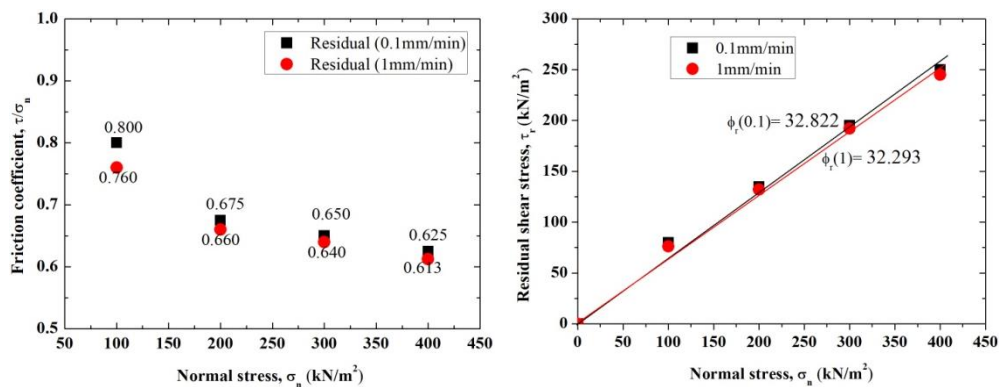
380 Fig. 9 gives the results of the Ydg samples. The coefficients $\tau_r/\sigma_n(0.1)$ and τ_r/σ_n
 381 (1) under the normal stress of 100 kN/m² to 400 kN/m² ranged from 0.57 to 0.52 and
 382 from 0.52 to 0.50, respectively. Furthermore, the difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at
 383 each normal stress was from -0.05 to -0.02. As for the difference between the residual
 384 friction angles, $\phi_r(1) - \phi_r(0.1)$, was in a range of -2.218° to -0.909°. In case of Ydg
 385 soil sample, the residual friction coefficients decreased with increase of shear rate for
 386 all normal stress levels.



387

388 Fig. 9. Relationships between residual shear stress and normal stress, and residual
 389 strength parameter for Ydg soil samples

390 Fig. 10 presents the results of the Dbz samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$
 391 $\tau_r/\sigma_n(1)$ at the normal stress of 100 kN/m² to 400 kN/m² ranged from 0.8 to 0.625 and
 392 from 0.76 to 0.613, respectively. The difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at each normal
 393 stress was from -0.04 to -0.01. The difference $\phi_r(1) - \phi_r(0.1)$ was from -1.425° to
 394 -0.405°. For Dbz samples, there was somewhat decrease tendency of the residual
 395 friction coefficients with the increasing of the shear rate for all normal stress levels. It
 396 is noted that the maximum difference was found at the lowest normal stress of 100
 397 kN/m².



398
 399 Fig. 10. Relationships between residual shear stress and normal stress, and residual
 400 strength parameter for Dbz soil sample

401 From the experimental results on the three selected landslides, it was found that
 402 there is a negative relationship between residual friction coefficients and shear rates
 403 for all samples (Figs. 8-10). Such a negative effect of shear rate (higher residual
 404 friction coefficients at lower rates) has been reported in the literature for fine-grained
 405 soils (Gratchev Ivan and Sassa, 2015;Tika et al., 1996). This effect may be closely
 406 associated with ability of clay particles in specimen to restore broken bonds at

407 different shear rates. Previous studies (Osipov et al., 1984; Perret et al., 1996).
408 concluded that with higher shear rates, the breakdown of the bonds between clay
409 particles or flocs exceeds the restoration bond, leading to reduction in residual friction
410 coefficients. In contrast, the bonds between particles are rebuilt quickly and the
411 recovery rate can catch up the breakdown rate at lower shear rates. Therefore, the
412 weaker bonding between particles could explain the strength drop with the increasing
413 of the shear rate in this study.

414 As for Ydg and Dbz specimen, it is found that the shear rate effect on the friction
415 coefficient can be seen to decrease with normal stress (Figs. 9 -10). By contrast, there
416 is an increasing tendency in the influence of shear rate on the friction coefficient with
417 normal stress in Djg specimen (Fig. 8). Gibo et al. (1987) reported that the residual
418 friction angle of soils was controlled by the effective normal stress as well as by the
419 CF. Interestingly, Ydg (with CF 9%) and Dbz (with CF 9.1%) specimens with almost
420 the same fraction of clay showed similar shear rate effect on the residual friction
421 coefficient with normal stress increasing, however, Djg (with 24% CF) showed the
422 contrast tendency of shear rate effect on residual friction coefficient with normal
423 stress, indicating that such effect is closely associated with CF.

424 Table 2 summarizes residual strength parameters including $\phi_r(0.1)$ and $\phi_r(1)$ of
425 all specimens obtained from the ring shear tests in this study. As for the Djg samples,
426 the residual strength parameter $\phi_r(0.1)$ and $\phi_r(1)$ for all normal stress were found to
427 be 15.003° and 14.09° (Fig. 8), respectively. However, the residual friction angles ϕ_r
428 (0.1) and $\phi_r(1)$ of the Ydg samples were obtained to be 27.954° and 26.778° (Fig. 9),

429 respectively. In the case of Dbz sample, the friction angles $\phi_r(0.1)$ and $\phi_r(1)$ were
 430 high, 32.822° and 32.293° (Fig. 10), respectively. The residual friction angles $\phi_r(0.1)$
 431 and $\phi_r(1)$ under all normal stresses were from 15.003° to 32.822° and from 14.09° to
 432 32.293° , respectively.

433 Due to the influence of the shear rate, the difference $\phi_r(1) - \phi_r(0.1)$, at each
 434 normal stress level varies in different locations, while the value of $\phi_r(1) - \phi_r(0.1)$
 435 under all normal stress for the Djg, Ydg and Dbz samples were -0.913° , -1.176° and
 436 -0.529° , respectively (Table 2). Wang (2014) and Fan et al. (2017) asserted that the
 437 residual shear strength of remolded loess hardly affected by shear rate below 5
 438 mm/min. However, the results in this study shown that $\phi_r(1) - \phi_r(0.1)$ under all
 439 normal stress levels were negative for loess. Moreover, the absolute value of $\phi_r(1)-$
 440 $\phi_r(0.1)$ in Ydg samples even reached up to 1.176° . **Therefore, the ring shear test**
 441 **results provides a basis for some general comments on the use of tests results with**
 442 **different shear rates, partially deepening some aspects deriving from previous studies.**

443

444 Table 2. Residual shear strength parameter of landslide soils.

No	Sample	Normal stress(kN/m ²)	Residual strength parameter			Difference in parameter	
			0.1 mm/min $\phi_r(0.1)$ ($c_r(0.1)=0$) (Degrees)	1 mm/min $\phi_r(1)$ ($c_r(1)=0$) (Degrees)		$\phi_r(1) - \phi_r(0.1)$ (Degrees)	
			Under each σ_n	Under each σ_n	Under each σ_n	Under each σ_n	Under all σ_n
		100	16.699	16.699		0	
1	Djg	200	15.563	15.642		0.079	
		300	15.110	14.216	15.003	-0.894	-0.913
		400	14.708	13.496	14.090	-1.212	

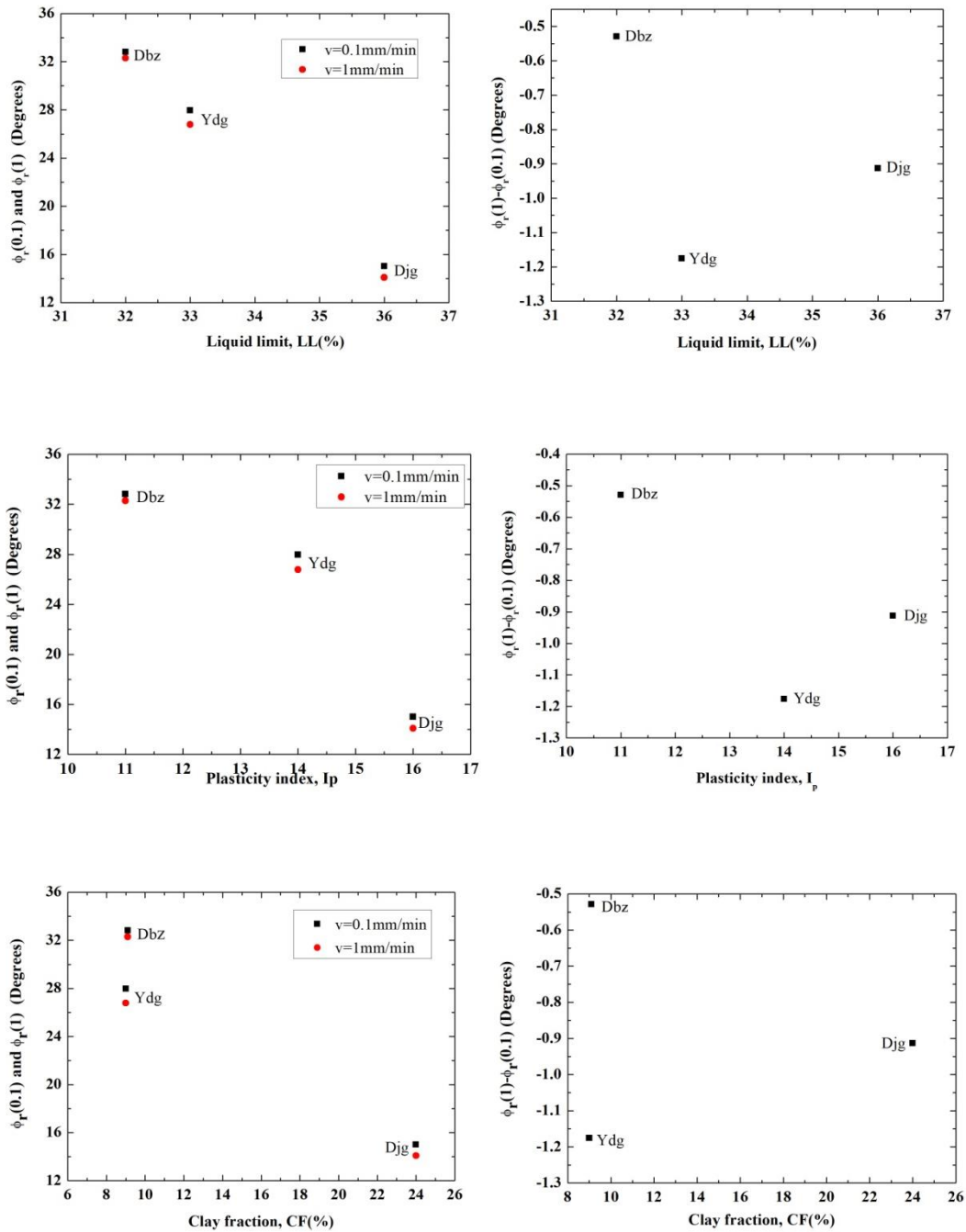
		100	29.683		27.474		-2.209
2	Ydg	200	29.466	27.954	27.248	26.778	-2.218
		300	27.923		26.870		-1.053
		400	27.474		26.565		-0.909
		100	38.660		37.235		-1.425
3	Dbz	200	34.019	32.822	33.425	32.293	-0.594
		300	33.024		32.619		-0.405
		400	32.005		31.487		-0.518

445

446 **4.4. Influence of the shear rate on the residual friction angles according to soil** 447 **properties**

448 It has been recognized that residual shear strength of soils is closely related with
449 soil properties, such as particle size distribution (PSD), liquid limit (LL), plasticity
450 index (Ip) and clay fraction (CF) (Terzaghi et al., 1996; Sayyah et al., 2016; Xu et al.,
451 2018; Eid et al., 2016). Fig. 11 depicts the relationships between residual friction
452 angles as well as the difference in the residual friction angles and soil properties
453 including LL, plasticity index (Ip) and clay fraction (CF) at two shear rates. The
454 residual friction angles at two shear rates decreased nonlinearly with the increasing of
455 the LL. As for the relationship between the ϕ_r and Ip, the ϕ_r under the low and high
456 shear rates decreases from about 32° to 15° with increasing the Ip from 11 to 16. These
457 findings agree well with the early studies (Wesley, 2003; Tiwari et al., 2005). With
458 increasing of CF from 9% to 24%, the residual friction angles under low and high
459 shear rates were found to decrease (Fig. 11). These observations are consistent with
460 previous studies (Lupini et al., 1981; Gibo et al., 1987). Interestingly, for Dbz and Ydg

461 soils of which have similar percentage of clay fraction, the residual friction angles at
 462 both shear rates varied. However, in the relationships between the difference in the
 463 residual friction angles and the soil properties, no clear correlations were found.



464

465

466

467

468

469 Fig. 11. Relationships between residual shear parameter, the difference in residual
 470 shear parameter and the soil properties at two shear rates

471

472 **5. Conclusion**

473 A series of ring shear tests were conducted on loess obtained from three landslides
474 to study the residual shear characteristics of saturated loess. Based on the test results,
475 the effect of the shear rate on the residual shear characteristics of loess in naturally
476 drained condition was examined. The following conclusions can be drawn:

- 477 1. Ring shear test revealed that (i) shear displacement to achieve the residual stage
478 with high shear rate is greater than that of the low shear rate; (ii) Both the peak
479 and residual friction coefficient became smaller with increase of shear rate for
480 each sample;(iii) The greater difference between the peak and the residual friction
481 coefficient in loess samples could be attributed to relatively well-developed
482 slickenside on the shear surface.
- 483 2. At the two shear rates, there was a nonlinearly decrease trend of the residual
484 friction coefficient with the normal stress in all loess samples. The difference
485 between the friction coefficients, $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ was found to decrease
486 with normal stress in Ydg and Dbz specimens while increase with normal stress in
487 Djj specimens, indicating that CF may be closely associated with shear rate effect
488 on residual friction coefficient with normal stress. **Therefore, as for Ydg and Dbz**
489 **with relatively low fraction of CF, there is an increase effect of shear rate on**
490 **residual friction coefficient with decreasing of normal stress. Thus, for the**
491 **application of measured residual friction coefficient for stability analysis of**
492 **shallow landslides with lower overburden pressure, it is significant for us to use a**
493 **low shear rate in ring shear tests to measure residual shear strength parameters. On**

494 other hand, for D_{jg} with high CF, it is more reliable to use a low shear rate in ring
495 shear tests to determine residual friction coefficient for stability analysis of deep
496 landslides with high overburden pressure.

- 497 3. The difference at the two shear rates, $\phi_r(1) - \phi_r(0.1)$, under each normal stress
498 level were either negative or positive. However, under all normal stress, the
499 difference at the two shear rates $\phi_r(1) - \phi_r(0.1)$ was found to be negative. Such
500 negative shear rate effect on loess could be attributed to greater ability of clay
501 particles in specimen to restore broken bonds at low shear rates.
- 502 4. The relationships between the ϕ_r under two shear rates and soil properties (LL, I_p),
503 demonstrated that the ϕ_r at both shear rates decreased gradually with the
504 increasing of LL and I_p . However, no clear correlations between the difference in
505 the ϕ_r at low and high shear rates and the soil properties were found.

506 A first attempt was made in this work to describe some shear rate effect on the
507 residual characteristics of the saturated loess. The obtained experimental results do
508 suggest that the residual shear behavior of saturated loess, can be affected, to a certain
509 extent, by the shear rate. However, a more quantitative evaluation of such effects, and
510 a deeper understanding on the underlying processes must be achieved in order to
511 assess their role in the initiation and mobility of loess landslides.

512

513

514 **Code availability:** Code can be made available by the authors upon request.

515 **Data availability:** Data can be made available by the authors upon request.

516 **Author contributions:** BL,JP and QH conceived and designed the method; BL
517 produced the results, and wrote the original manuscript under the supervision of JP.
518 JP and QH writing-review and editing.

519 **Competing interest:** The authors declare that they have no conflicts of interest.

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