Towards a comprehensive view of dust event from multiple satellite and ground measurements: exemplified by the East Asia May 2017 dust storm

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Abstract. One or several aspects of the source, distribution, transport, optical properties of airborne dust have been characterized using different types of satellite and ground measurements each with unique advantages. In this study, a dust event occurred over the East Asia area in May 2017 was exemplified to demonstrate how all the above mentioned aspects of a dust event can be pictured by combining the advantages of different satellite and ground measurements. The used data included the Himawari-8 satellite Advanced Himawari Imager (AHI) true-colour images, the Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO) Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) aerosol vertical profiles, the Aura satellite Ozone Monitoring Instrument (OMI) aerosol index images, and the ground based Aerosol Robotic Network (AERONET) aerosol properties and the ground station particulate matter (PM) measurements. From the multi-satellite/sensor (AHI, CALIOP and OMI) time series observations, the dust storm was found to originate from the Gobi Desert on the morning of 3 May 2017 and transport northeastward to the Bering Sea, eastward to the Korean Peninsula and Japan, and southward to south-central China. The air quality in China deteriorated drastically: the PM_{10} (PM<10 μm in aerodynamic diameter) concentrations measured at some air quality stations located in northern China reached 4333 μg/m^3. At the AOE_Baotou, Beijing, Xuzhou-CUMT, and Ussuriysk AERONET

sites, the maximum aerosol optical depth values reached 2.96, 2.13, 2.87, and 0.65 and the extinction Ångström exponent dropped to 0.023, 0.068, 0.03, and 0.097, respectively. The dust storm also induced unusual aerosol spectral single-scattering albedo and volume size distribution.

1. Introduction

Large amounts of dust particles are emitted from the deserts in western/northern China and southern Mongolia every year, especially in spring (Shao et al., 2011). The annual dust emissions of eastern Asia reach approximately 25% of the total global dust emissions (Ginoux et al., 2004). These massive emissions produce significant influences on the Earth's radiation balance, climate, ambient air quality and human health (Goudie, 2009;Shao et al., 2011;Rodríguez et al., 2012). Dust aerosols exert both direct and indirect effects on the climate system. Dust can directly scatter and absorb solar radiation over ultraviolet, visible, and infrared wavelengths, resulting in positive or negative forcing (Rosenfeld et al., 2001;Tegen, 2003). Dust is also involved in cloud formation and precipitation processes and can alter the albedo of snow and ice surfaces, thereby causing indirect effects on the Earth's energy budget (Rodríguez et al., 2012;Rosenfeld et al., 2001;Bangert et al., 2012).

Due to the long-distance transport of dust plumes (Zhu et al., 2007), dust particles can alter the atmospheric conditions in regional even global scale(Goudie, 2009). The dust aerosols from the Taklimakan and Gobi Deserts can travel thousands of miles, thereby affecting large areas of China (Wang et al., 2013;Lee et al., 2010;Chen et al., 2015;Tan et al., 2012), South Korean and Japan (Mikami et al., 2006), and even the Northern Pacific Ocean and North America (Fairlie et al., 2007;Creamean and Prather, 2013;Guo et al., 2017). Dust storms can cause poor air quality and low visibility and have severe effects on the human health and environment (Goudie, 2009;Lee et al., 2010). Desert dust is the main contributor to aerosol loading and particulate matter (PM) mass concentrations in China during the spring season (Wang et al., 2013). During heavy dust outbreaks, PM_{10} (PM less than 10 μm in aerodynamic diameter) mass concentrations can even reach 20 exceedances of the internationally recommended limit value in northern China. Moreover, dust particles can interact with anthropogenic pollution and smoke, causing air conditions with greater complexity (Dall'Osto et al., 2010).

Many literatures have studied desert dust from different perspectives using different satellite data, ground-based observations and model simulations (Badarinath et al., 2010;Wang et al., 2013;Teixeira et al., 2016). For example, some analysed the dust chemical composition and dust radiative effects, i.e., dust optical and micro-physical properties (Alam et al., 2014;Basha et al., 2015;Srivastava et al., 2014). The other focused on the long-distance transport of dust plumes using satellite observations and/or model simulations (Huang et al., 2008;Guo et al., 2017;Athanasopoulou et al., 2016). However, few studies have been carried out to fully examine the source, distribution, transport, optical properties of the dust storm. This is possibly because each observation system can only characterize one or several aspects of them.

This study tried to picture a comprehensive view of dust event using different satellite and ground measurements with a recent heavy dust storm over northern China and southern Mongolia from 3 to 8 May 2017 as an example. The dust storm spread with wind across south eastern Russia and even reached the Bering Sea on 7-8 May 2017. Satellite time series observations (the Himawari-8 satellite Advanced Himawari Imager (AHI) true-colour images and the Ozone Monitoring Instrument (OMI) aerosol index (AI) images) were used to capture the dust transport. The OMI AI was also used to provide information about the absorbing aerosol distribution. The Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO) data were used to monitor the dust aerosol type and vertical distribution. The Air Resources Laboratory's HYbrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model was used to generate back trajectories to identify the dust sources. Ground-based measurements form both Aerosol Robotic Network (AERONET) and air quality stations were used to analyse the variations in aerosol properties caused by the dust storm. The connections and correspondences among different observations are briefly analysed.

2. Data and methods

2.1 General description of the study area

Fig. 1 shows the topography of the study area. The deserts in China and Mongolia, where an abundance of dust events occur, constitute the second-largest dust source in the world. During the spring, the Gobi region is affected by the Mongolian cyclones, which is the main factor to the severe Asian dust storms (Shao et al., 2011).

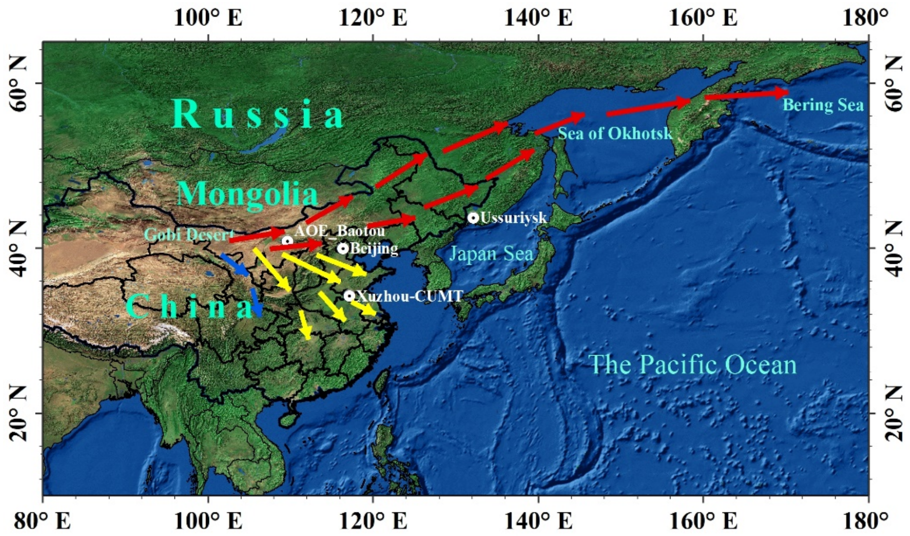

Fig. 1. Study area. The white circles with black point represent four AERONET stations and the arrows show the dust transport directions.

2.2 AHI/Himawari-8

The Himawari-8 (H8) satellite was launched on 7 October 2014 by the Japan Meteorological Agency (JMA). It started operation on 7 July 2015. The Advanced Himawari Imager (AHI) on board H8 can provide multi-spectral observations with a high spatial resolution and high temporal frequency. It has 16 channels with a spatial resolution of 0.5-2 km. The AHI level 2 calibrated data provided by JMA have a spatial coverage of 120° by 120° centred at 0° N, 140° E, and the observation area includes most of eastern Asia, Australia and the Pacific Ocean. In addition, the AHI provides full-disk observations every 10 minutes. AHI level 2 calibrated data provided by the JMA and downloaded from the Japan Aerospace Exploration Agency (JAXA) Earth Observation Research Center (EORC) were used (http://www.eorc.jaxa.jp/ptree/terms.html).

2.3 OMI/Aura

The OMI sensor aboard the Aura satellite measures the Earth in the ultraviolet (UV) and visible spectra (270-550 nm) with a wide swath. The OMI provides a parameter called the UV aerosol index (UV-AI), which is a qualitative parameter that detects UV-absorbing aerosols. The UV-AI is sensitive to absorbing aerosols, including mineral dust, black carbon, and biomass burning aerosols (Eck et al., 2001). Therefore, the UV-AI can be used to identify aerosol types through positive values for dust and biomass burning particles and near-zero or small positive values for clouds and weakly absorbing aerosols (Torres et al., 2007). In addition, the UV-AI can be obtained under both cloudy and cloudless conditions. The surface reflectance also has no impact on the UV-AI, which makes it capable of detecting absorption by aerosols over highly reflective surfaces (Torres et al., 2007). Since this dust event occurred in May, a high UV-AI can be a good indicator of high dust aerosol loading when combined with CALIPSO observations, as Aura and CALIPSO have similar equatorial crossing times. Here, level 3 OMI UV-AI data were used with a 0.125° x 0.125° spatial resolution.

2.4 CALIOP/CALIPSO

The Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) instrument on board the CALIPSO satellite provides vertical profiles of the elastic backscatter at two wavelengths (532 nm and 1064 nm) during both day and night. The CALIOP payload also provides linear depolarization at 532 nm that can be used to identify dust aerosols since dust aerosols have a high linear depolarization ratio due to their non-sphericity. Aerosol types are also provided in the CALIPSO aerosol product. The CALIPSO algorithm defines six aerosol types, including smoke, dust, polluted dust, clean continental, polluted continental, and marine (Omar et al., 2009;Omar et al., 2013). It has been evaluated that the CALIPSO aerosol classification works well in most cases (Wu et al., 2014). It should be noted that the accuracy of aerosol detection is decreased over highly reflected land surfaces such as deserts and snow-covered regions. Here, CALIPSO level 2 vertical feature mask (VFM) aerosol layer products were used to provide independent information about dust aerosols, especially for the night-time, as the signal-to-noise ratio during the night-time is better than that during the daytime for CALIPSO (Liu et al., 2009). The VFM products have a vertical resolution of 30 m below 8.2 km, 60 m for 8.2-20.2 km, and 180 m for 20.2-30.1 km (Winker et al., 2007). The dust vertical distribution and dust layer height, were analysed using CALIPSO VFM data.

2.5 AERONET data

The AERONET is a ground-based remote sensing aerosol network (Holben et al., 1998) that provides spectral aerosol optical depth (AOD) and inversion products derived from direct and diffuse radiation measurements by Cimel sun/sky-radiometers (Dubovik et al., 2006). The inversion products include both, microphysical parameters (e.g., the size distribution and complex refractive index) and radiative properties (e.g., the single-scattering albedo and phase function) (Dubovik et al., 2006).

In this study, Level 1.5 cloud screened data including both sun direct data (Version 2 and Version 3) and

Inversion data (Version 2) from four AERONET sites in the study area were used to analyse the temporal variations in aerosol properties, including the AOD, the extinction Ångström exponent (EAE), volume size distribution (VSD), and single-scattering albedo (SSA). Fig. 1 shows the locations of the four sites (white circles), namely, AOE_Baotou, Beijing, Xuzhou-CUMT, and Ussuriysk.

2.6 PM measurements

There are thousands of air quality stations over China that can provide hourly PM measurements during both, day-time and night-time. In addition, the measurements are free from the influences of clouds, making it a perfect complement to AERONET observations and satellite observations. Ground-based measurements of the PM mass concentration over the Chinese mainland were collected to evaluate the dust-affected areas and to further analyse the transport of the dust plume. Furthermore, the temporal variations in the PM concentrations at 14 typical stations were analysed in detail to examine the propagation of dust particles in different directions. Detailed information about these 14 air quality stations is given in Table 1.

Table 1. The cities and locations of the 14 air quality stations to examine the propagation of dust particles

City (Site)	Longitude (°)	Latitude (°)	City (Site)	Longitude (°)	Latitude (°)
Bayannao'er (BYN)	107.5936	40.916	Shanghai (SHS)	121.536	31.2659
Changsha (CSS)	112.9958	28.3586	Taiyuan (TYS)	112.5583	37.7394
Chengde (CDS)	117.9664	40.9161	Tianshui (TSS)	105.7281	34.5814
Guangyuan (GYS)	105.8153	32.4246	Tongliao (TLS)	122.2603	43.6267
Heihe (HHS)	127.4961	50.2486	Weihai (WHS)	122.0508	37.5325
Huhhot (HHT)	111.7277	40.8062	Zhengzhou (ZZS)	113.6113	34.9162
Jiangchang (JCS)	102.1878	38.5247	Zhongwei (ZWS)	105.18	37.0172

2.7 HYSPLIT trajectories and meteorological data

The HYSPLIT model developed by NOAA's Air Resources Laboratory was employed (Draxler and Rolph, 2013). It is widely used for computing air mass forward/backward trajectories to analyse the transport of air/pollution parcels. The start/end point as well as the time of the HYSPLIT computation can be customized. Here, HYSPLIT was used to generate air mass backward trajectories to trace the air movement. The data from Global Data Assimilation System (GDAS) with a spatial resolution of 0.5 degree were used as meteorology input. The backward trajectories ending at 9 selected points were calculated for determining the dust source.

The meteorological data including wind vectors and geopotential height (GH) from ECMWF ERA Interim reanalysis dataset were also used (http://apps.ecmwf.int/datasets/data/interim-full-daily/levtype=pl/). The distribution of the wind direction, wind speed and GH with a spatial resolution of 1 degree during 3-8 May 2017 were analysed to understand the origin and transport of the dust storm.

3. Results

3.1 Source and transport of the dust event

Fig. 2 shows the spatial distribution of the UV-AI over East Asia from 3 to 8 May 2017 obtained from the OMI-Aura observations. High AI values (>2.0) can be observed over northern China, especially over Inner Mongolia on 3 May, north eastern China on 4-5 May, and south eastern Russia on 5-6 May. The OMI-AI time series revealed one of the long-distance transport path of the strong absorbing aerosols that originated from the Gobi Desert, i.e., moving towards the east and then northeast (hereafter referred to as northeast direction for simplicity). This can be explained by the strong west and southwest wind evident in Fig. 3, which showing the spatial distribution of the wind vectors and geopotential height field at 500-pha level at 06:00 UTC during 3-8 May. The dust storm was initially developed over western Inner Mongolia (~ 40° N, 100° E) on 3 May 2017 (see Fig. 2a) and then swept through the North China plain and reached north eastern China (~50° N, 125°E) on 4 May 2017 due to a strong west wind (Fig. 2b and Fig. 3). On 5 May, the dust plume was transported to the western Sea of Okhotsk (~56°N, 140°E). For the next two days, the elevated dust plume travelled across the Sea of Okhotsk and finally reached the Bering Sea (see Fig. 2e-f). Furthermore, there is a small portion of the high AI values in the Japan Sea on 7 May (Fig. 2e) indicating that there is a second dust transport path of all the way east and the Korean Peninsula and Japan were affected. This is because the wind field diverged to two directions at North China plain, i.e., towards northeast and towards east (Fig. 3).

To confirm that the high AI values were caused by dust aerosols, CALIPSO observations that passed through the high AI value regions during the night-time were employed to provide aerosol type and vertical distribution information. Fig. 4 depicts the vertical distributions of the aerosol types for six overpass trajectories shown in Fig.2 (the deep blue lines). The dust aerosol (yellow) in Fig. 4 corresponded well to the high AI value region in Fig. 2. Furthermore, the dust layer was thick and distributed from the surface to a 10 km height. Moreover, part of the aerosol layer was marked as the polluted dust subtype over the Central China on 4 May and over the northern China on 5 May. This may be explained by the mixture of dust and anthropogenic pollution.

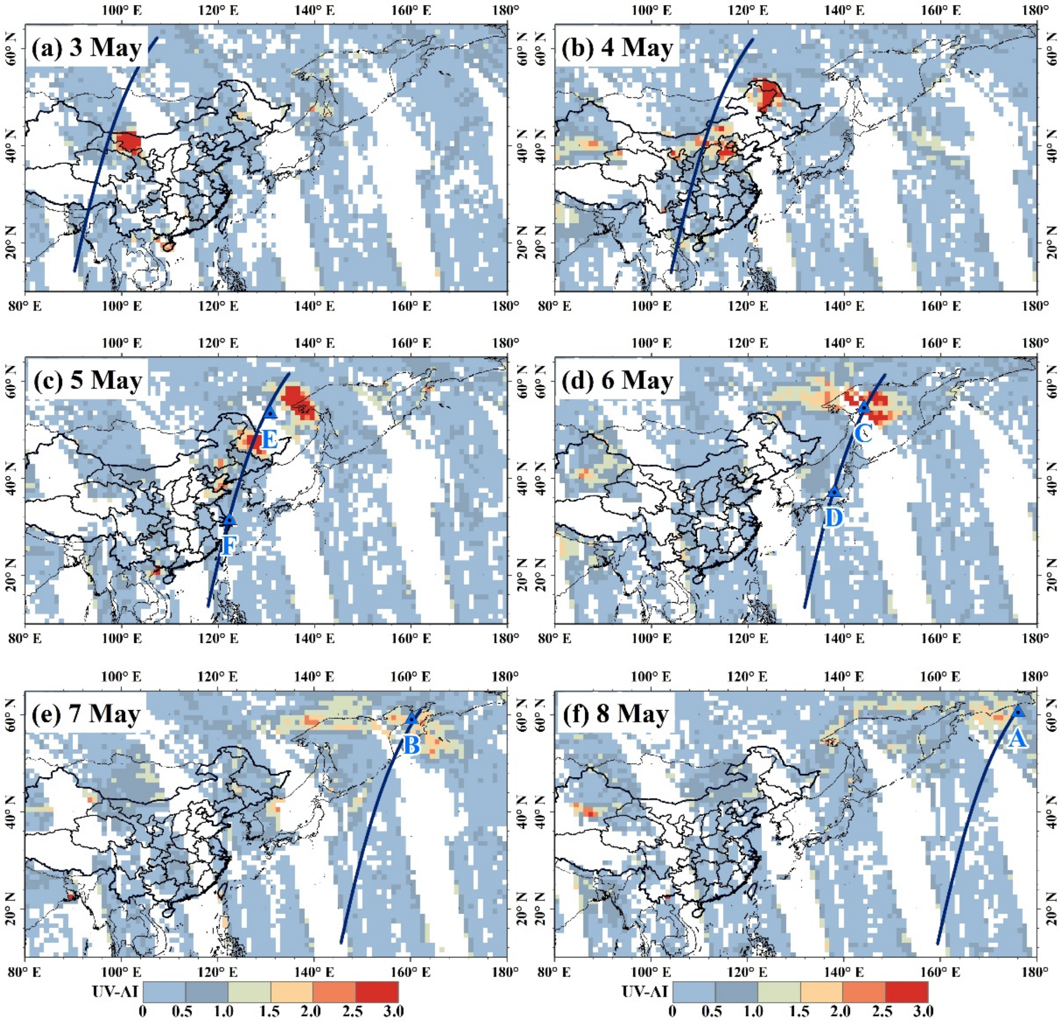

Fig. 2. a-f. Spatial distributions of the OMI UV AI from 3 to 8 May 2017. The deep blue lines are the overpass trajectories of the used CALIPSO observations on that day. The blue triangles are the end points of the HYSPLIT computation.

Fig. 5 shows the backward trajectories at different sources (the blue triangles in Fig. 2) during 5-8 May 2017. The trajectories are computed at three different altitudes (1000 m, 2000 m, and 3000 m). The HYSPLIT backward trajectory analysis revealed that the air masses that reached the Bering Sea (Fig. 5a), the Kamchatka Peninsula (Fig. 5b), the Sea of Okhotsk (Fig. 5c), and the Japan Sea (Fig. 5d) originated from the Gobi Desert. This result is consistent with that from the OMI-AI and CALIPSO aerosol type information.

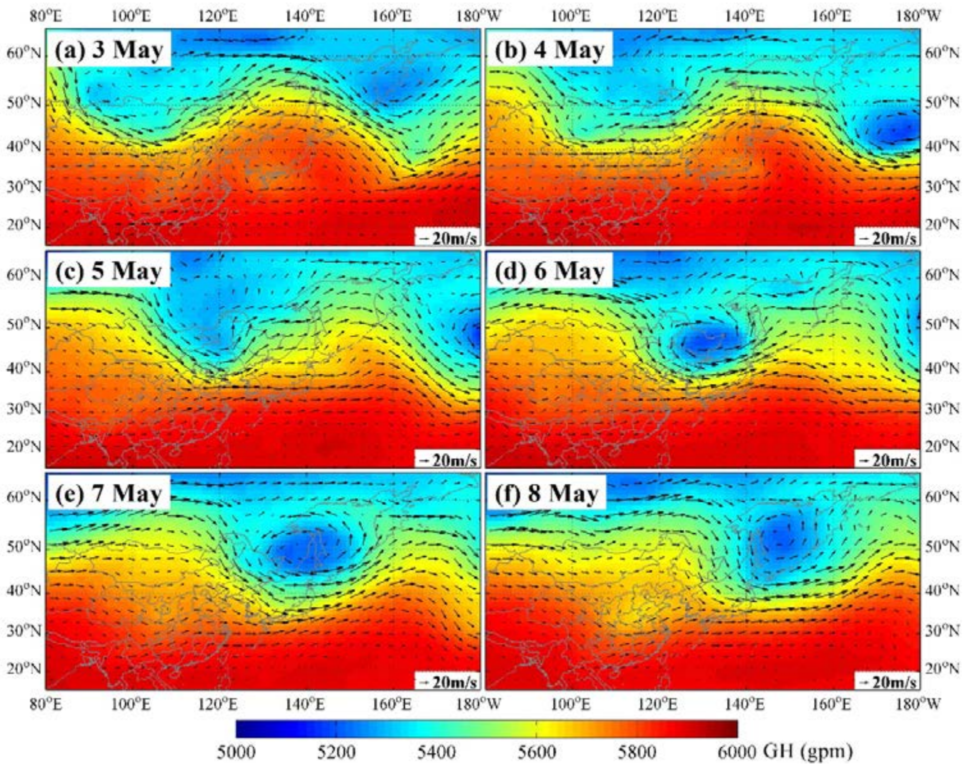

Fig. 3. a-f. Spatial distribution of wind vectors and geopotential height (GH) at 500-hPa level at 06:00 UTC during 3-8 May. The GH and wind vectors were derived from ECMWF ERA Interim reanalysis dataset.

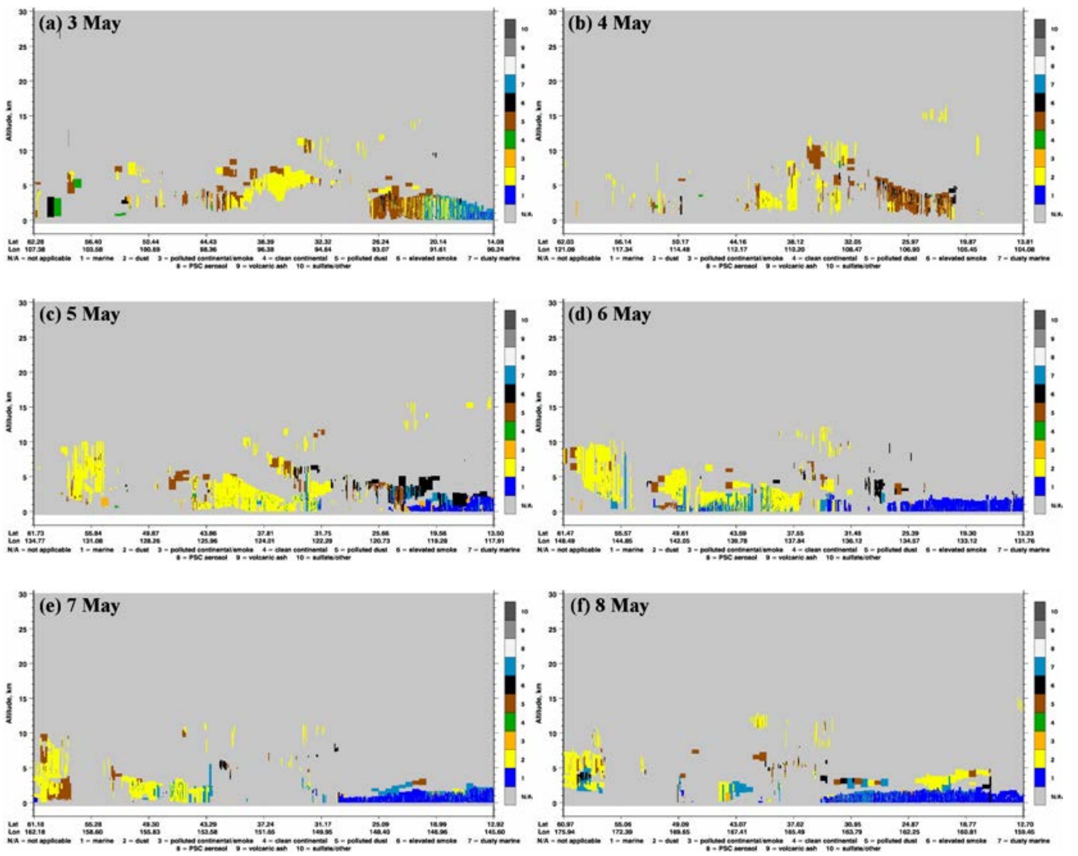

Fig. 4. a-f. CALIPSO aerosol subtypes on 3-8 May 2017, corresponding to the overpass trajectories shown in Fig. 2. The dust aerosol is shown in yellow.

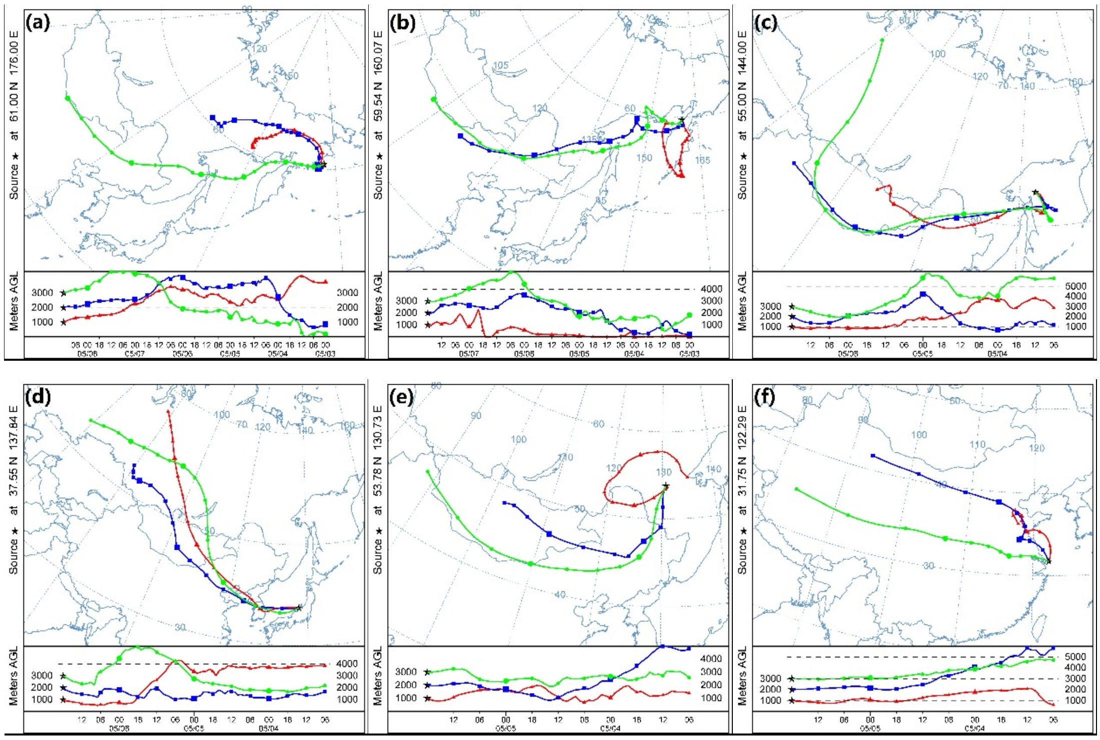

Fig. 5. Backward trajectories derived from the HYSPLIT model at different locations during 5-8 May 2017: (a) 132-h back trajectories ending at the Bering Sea on 8 May, (b) 108-h back trajectories ending at the Kamchatka Peninsula on 7 May, (c-d) 84-h back trajectories ending at the Sea of Okhotsk and the Sea of Japan on 6 May, respectively, and (e-f) 60-h back trajectories ending at southeastern Russia and the Yangtze River estuary region on 5 May, respectively.

For a better view of the dust plume transport, the high-temporal-resolution observations from the Himawari-8/AHI true-colour composite images on 3 May (Fig. 6) and 4 May (Fig. 7) were shown. The results in Fig. 6 suggest that the strong dust storm was originated from the western part of the Gobi Desert and was formed by several distinct dust clusters (Fig. 6b). In the morning of 3 May, only a small area was covered by a dust plume in the Gobi Desert (Fig. 6a), as the dust storm continuously increased and quickly moved. Part of the dust plume over south western Inner Mongolia moved along the edge of the Qinghai-Tibet Plateau and then finally reached the northern Sichuan basin (Fig. 6c), revealing the third path of the dust transport. This path of the dust transport is not revealed in the OMI AI time series maps possibly because the dust in this path is not very severe. On the other hand, thick dust plume travelled along the China-Mongolia border with continually increasing dust intensity and moved quickly towards the northeast and east. In addition, in the late afternoon of 4 May 2017 (Fig. 7), another thick dust plume was found that originated from northern Inner Mongolia (Fig. 7c) that was quickly transported eastward due to strong westerly winds. High-frequency observations from the AHI presented more information about this severe dust storm, revealing multi-plumes propagation and several different transport directions, including south eastward, eastward and north eastward. The longest-distance transport occurred in the north eastward direction, as OMI-AI and CALIPSO-VFM illustrated in the previous section, and finally arrived at the Bering Sea.

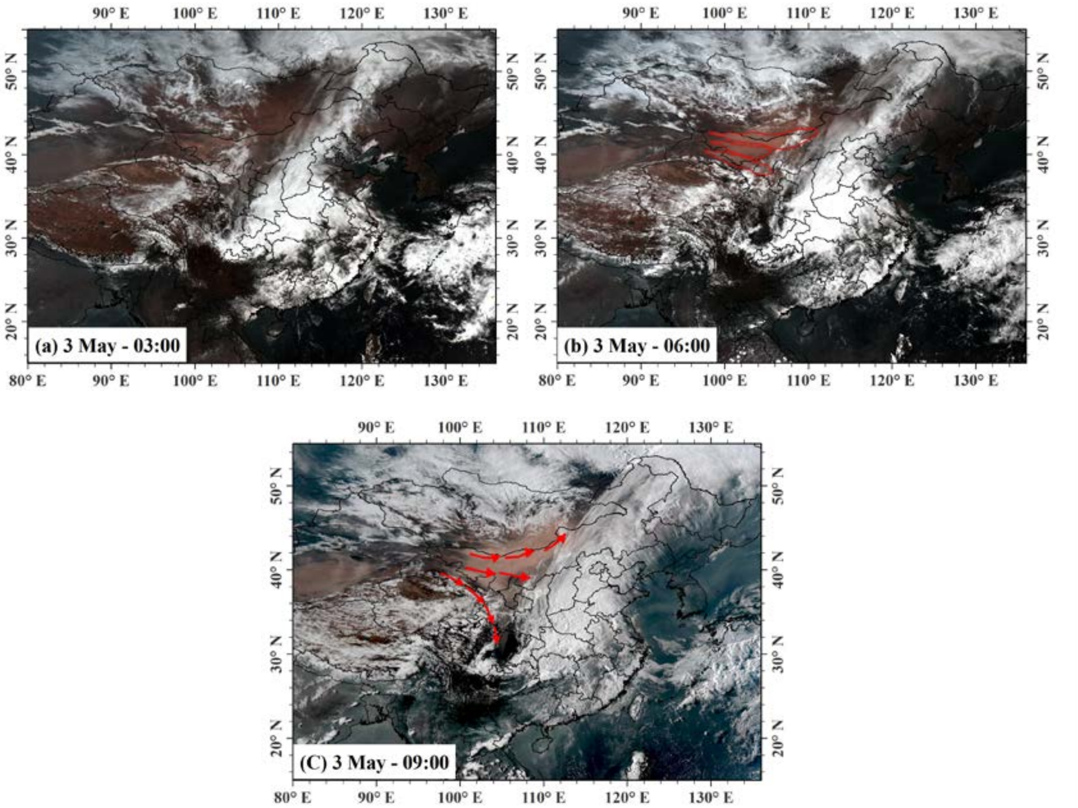

Fig. 6. True-colour composite images of mainland China (a-c) from AHI data over a 3-h interval on 3 May 2017. (a) 03:00, (b) 06:00, (c) 09:00. The red polygons in (b) are dust clusters. The arrows in (c) represents the dust transport directions.

3.2 PM characterization in China during the dust event

In this section, the temporal variations in the $PM_{2.5}$ (PM<2.5 μm in aerodynamic diameter) and PM_{10} mass concentrations over mainland China were analysed and the third path of the dust transport, i.e., towards southeast, is obvious. Fig. 8 depicts the hourly PM_{10} concentration distribution over mainland China over a 12-h interval from 06:00 a.m. on 5 May to 06:00 p.m. on 7 May (Beijing time) using a total of 1350 stations, the PM values are real-time measurements per hour. The PM concentration value less than 200 were shown in grey. Interestingly, south eastward transport was revealed through the intensive PM concentration measurements, which was almost missed by most of the satellite observations because the central and eastern China were covered by cloud during 5-7 May. The high PM_{10} concentration was mostly distributed over 35-40°N at 06:00 on 5 May (Fig. 8a). After 12 hours, the dust plume moved to Shandong Peninsula (eastern China close to the Yellow Sea) and further affected Central China on 6 May (Figs. 8 c-d). On May 7 the dust events were found in most stations of eastern and central China (Figs. 8 e-f). The southward propagation of the dust plume caused a high PM_{10} concentration (>500) in south-central China (~28° N, 118° E) as well as the eastern coastal areas including the Shandong Peninsula (eastern China close to the Yellow Sea), and the Yangtze River Delta.

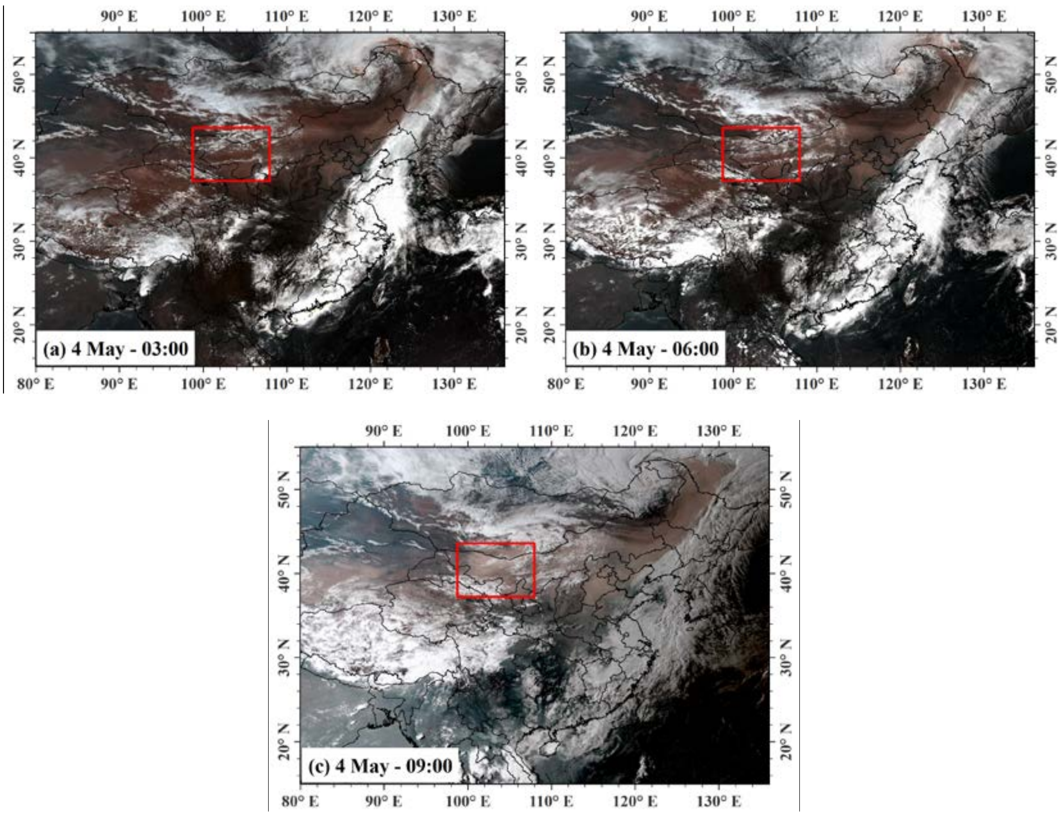

Fig. 7. True-colour composite images of mainland China (a-c) from AHI data over a 3-h interval from 03:00 UTC to 09:00 UTC on 4 May 2017. (a) 03:00, (b) 06:00, (c) 09:00. The red rectangles marked the area where another dust plumes originated from in the afternoon of 4 May.

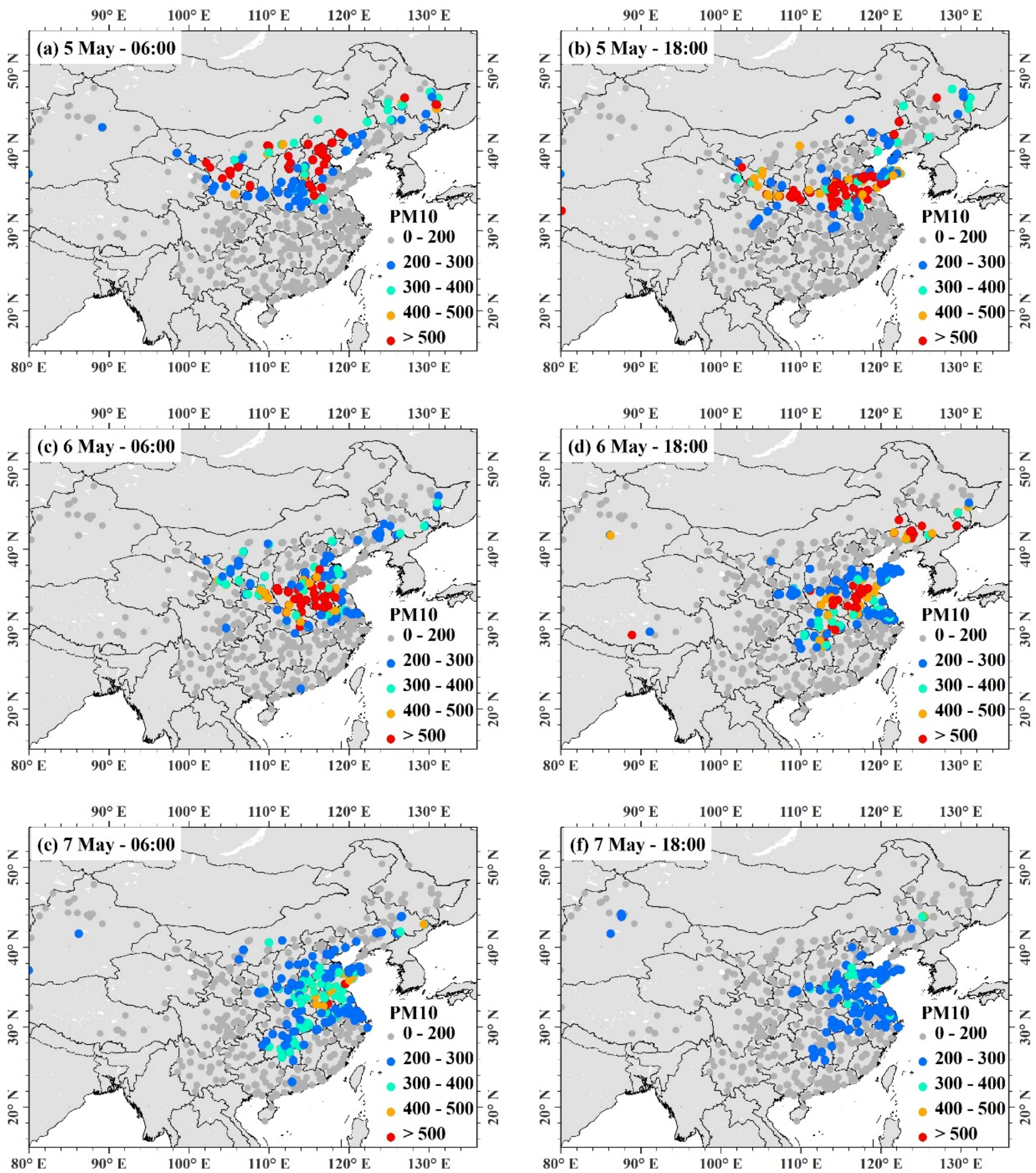

Fig. 8. PM_{10} mass concentrations measured by ground-based air quality stations in mainland China over a 12-h interval from 06:00 on May 5 to 18:00 on May 7.

To obtain better insight into the dust evolution, measurements from 14 typical air quality stations (the colour circles in Fig. 9) situated within the source and transport areas of the dust were analysed in detail. As the PM concentration was measured during both the day-time and night-time, the data can provide much more information about this continuous dust plume. Fig. 10 shows the PM temporal variations along three different dust transport directions during 2 to 7 May, including the northeastward propagation (a), southward propagation (b) and southeastward propagation (c). It is clearly observed that both the $PM_{2.5}$ and the PM_{10} increased dramatically, and the PM_{10} showed much larger increments than the $PM_{2.5}$ during this dust event from all three figures.

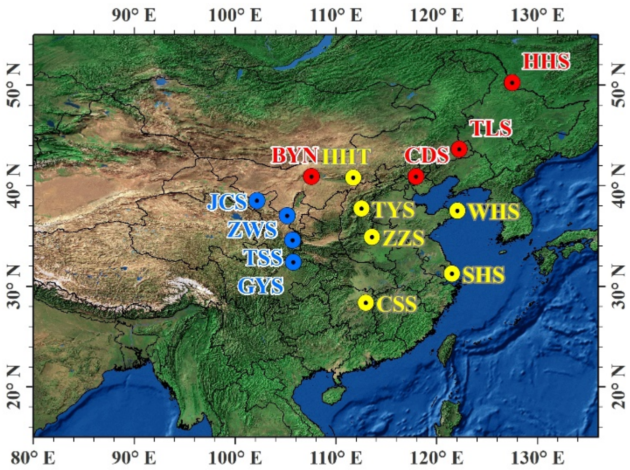

Fig. 9. The locations of the 14 air quality stations. Different color represent different dust transport directions names by sequencing site names in the time order when the dust passes: the northeast direction of BYN-CDS-TLS-HHS, and the two south directions of JCS-ZWS-TSS-GYS and HHT-TYS-ZZS-CSS. Note the east direction travelling to the Japan Sea is not shown as most of its path is over sea without air quality stations.

As Fig. 10a illustrates, the sharp increase in the PM mass concentration was first observed at BYN station on the morning of 3 May, followed by the CDS (23:00 UTC on 3 May) and TLS (8:00 UTC on 4 May), and reached the northeastern-most city, namely, Heihe (HHS) (06:00 UTC on 4 May). The maximum value of PM_{10} concentration at BYN reached 4333 μg/m^3 on 4 May. High PM_{10} concentrations occurred successively at those sites. These drastic changes in the PM_{10} are in agreement with the dust movements revealed from the satellite observations.

PM measurements at 4 stations distributed along the eastern edge of the Qinghai-Tibet Plateau, including JCS, ZWS, TSS and GYS, are shown in Fig. 10b. Within one day, the dust plume was transported across Gansu and reached GYS, which is located in the Sichuan Basin. This transport was also revealed by the high-frequency AHI observations (Fig. 6c), although it is not as noticeable.

Fig. 10c displays the PM concentration variations over the cities located in Central China, including Taiyuan (TYS), Zhengzhou (ZZS), and Changsha (CSS). In addition, high PM concentrations were observed in the coastal areas of eastern China, as shown in Fig. 11. Note that the increases of PM_{10} are much larger than the increments of $PM_{2.5}$ in these stations, suggesting that the dust particles were transported to southern and eastern China.

To confirm this southward propagation of dust, the backward trajectories ending at GYS, CSS, and SHS were analysed by HYSPLIT, as shown in Fig. 12. The trajectories are computed at three different altitudes

(500 m, 1000 m, and 1500 m). The northwestern air masses at all three locations originated from sources in the Gobi Desert. Thus, dust could be the main reason for the sudden increase in the PM concentrations.

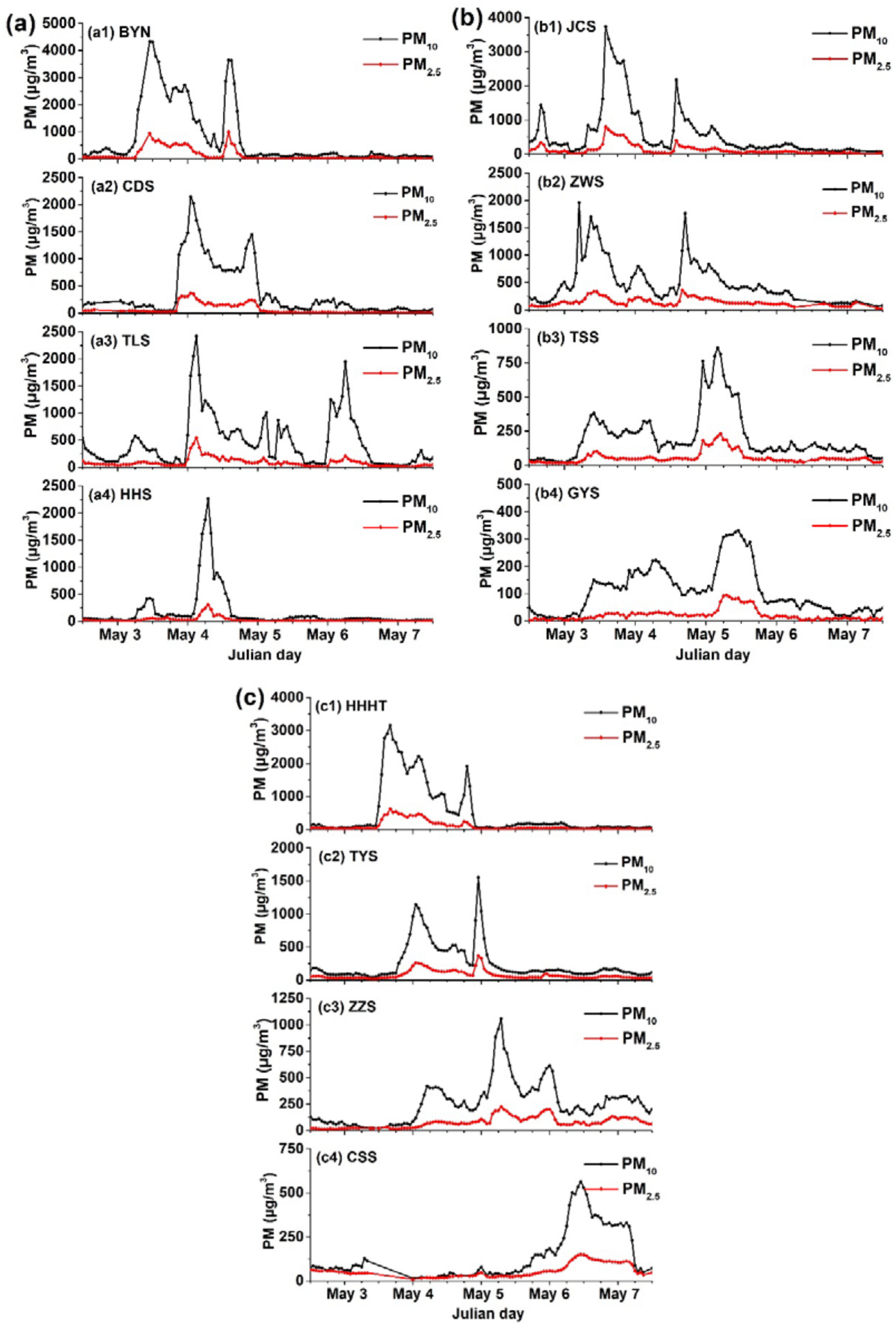

Fig. 10. Time series of the $PM_{2.5}$ (red curves) and PM_{10} concentrations (black curves) during 2-7 May at 14 air quality stations in three directions: (a) northeastward propagation, including (a1) BYN, (a2) CDS, (a3) TLS and (a4) HHS, (b) southward propagation, including (b1) JCS, (b2) BYS, (b3) TSS and (b4) GYS, and (c) southeastward propagation, including (c1) HHT, (c2) TYS, (c3) ZZS and (c4) CSS.

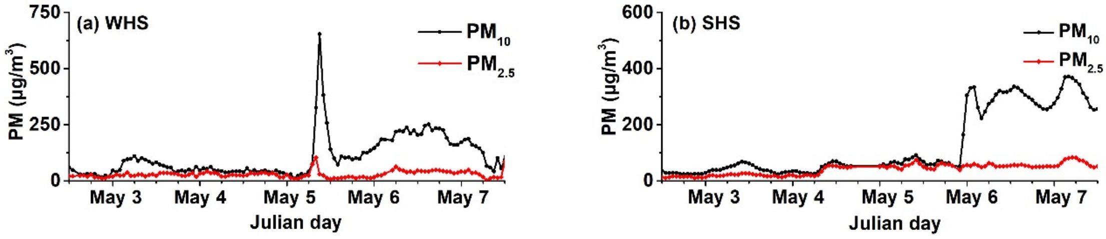

Fig. 11. Time series of the PM$_{2.5}$ (red curves) and PM$_{10}$ concentrations (black curves) during 2-7 May at (a) WHS and (b) SHS.

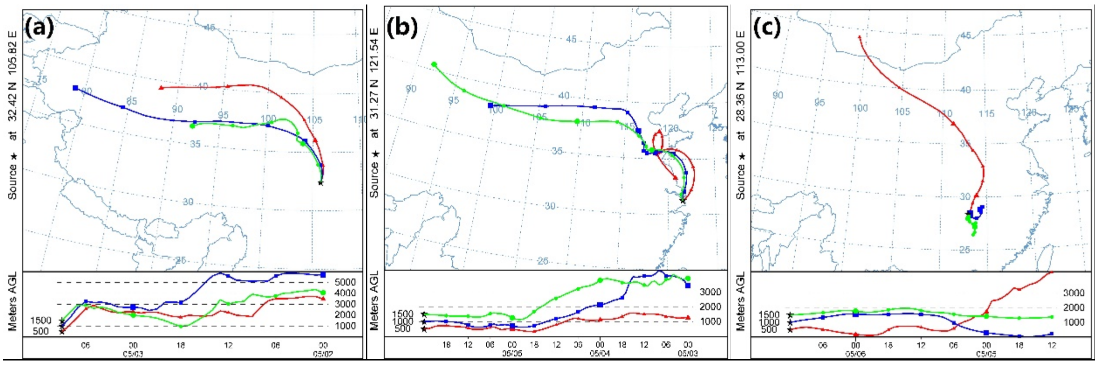

Fig. 12. Backward trajectories derived from the HYSPLIT model at different altitude levels (500 m, 1000 m, and 1500 m) at (a) GYS on 3 May, (b) SHS on 6 May, and (c) CSS on 6 May 2017.

3.3 Aerosol property variations during the dust event

In order to understand the effects of dust storm on aerosol properties, the changes in the aerosol properties at four typical AERONET stations located in the study area were investigated. The longitudes of AOE_Baotou, Beijing, and Ussuriysk increase from west to east, and the latitudes of AOE_Baotou, Beijing, and Xuzhou-CUMT decrease from north to south (Fig. 1). Several key aerosol properties, including the AOD, EAE, SSA, and aerosol VSD, were analysed in detail.

The temporal variations in the daily AOD and EAE at the four AERONET sites during dusty and non-dusty days are plotted in Fig. 13. The maximum AODs at 440 nm caused by the dust storm were 2.96, 2.13, 2.87, and 0.65 at AOE_Baotou, Beijing, Xuzhou-CUMT, and Ussuriysk, respectively. The maximum AOD at Baotou (the westernmost station) was recorded on 2 May 2017 and became lower afterwards. Then, the dust storm moved eastward, and the highest AOD value of 2.13 was observed over Beijing on 4 May 2017. As the dust storm travelled northeastward, the Ussuriysk, located in southern Russia, was affected with a slight increase in the AOD (from ~0.25 to ~0.65) and a sharp decrease in the EAE (from ~1 to ~0.1) on 5 May 2017. Xuzhou-CUMT, which is located in southern Central China, was also severely affected by the strong dust on 4-5 May. The maximum AODs occurred at different times at the four sites due to the movement of the dust storm. In addition, there are obvious negative correlations between the AOD and EAE during the dust event. The dust storm brought numerous large particles, causing the low EAE and high extinction properties.

The SSA is strongly related to absorption/scattering characteristics. Fig. 14 shows the variability of the spectral SSA before, during and after this dust event, and the monthly average as a benchmark. The SSA at longer wavelengths (e.g., 675, 870, and 1020 nm) at AOE_Baotou varied from ~0.8 to ~0.9 during non-dusty days (1 May and 2 May), and the monthly average of SSA_{675nm} (SSA at 675 nm) was approximately 0.9. In contrast the SSA_{675nm} increased to 0.97-0.98 during dust days (2 and 4 May). In addition, the spectral behaviour of the SSA showed significant differences. The SSA increased with the wavelength on 2 May and 4 May. Especially on 4 May, the SSA largely increased from 440 nm to 675 nm (from 0.93 to 0.98), and the dSSA ($dSSA=SSA_{870nm}-SSA_{440nm}$) also increased to 0.07. According to Dubovik et al. (2002), mineral dust aerosols tend toward a dSSA value greater than 0.05. In contrast, the monthly average of the spectral SSA as well as the spectral SSA during non-dusty days obviously decreased with the increase in wavelength. The high SSA and increasing spectral behaviour indicates that aerosol particles are dominated by large particles with strong scattering. However, it was noticed that the SSA_{440nm} on 2 May was high. This could be explained by the mixture of dust aerosols with large amounts of anthropogenic aerosols from industrial emissions, which are more absorbent.

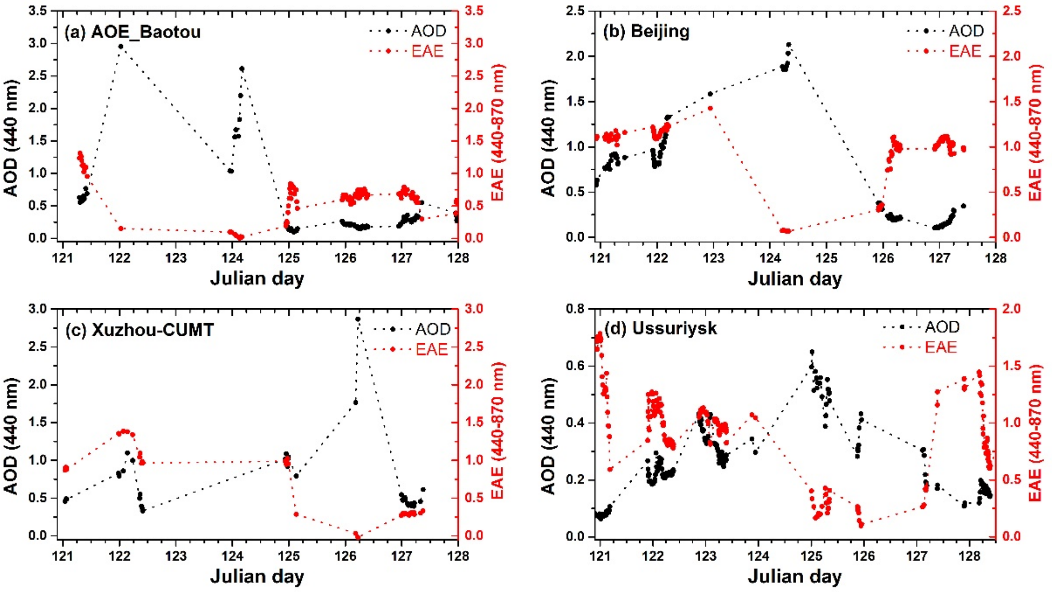

Fig. 13. Variations in the AOD (440 nm) and Ångström exponent (440-870 nm) at (a) AOE_Baotou, (b) Beijing, (c) Xuzhou-CUMT, and (d) Ussuriysk during 1-8 May 2017.

Similar properties can be observed over Beijing, as the dusts over both Baotou and Beijing have similar sources. However, there are still a few differences. The monthly average of the spectral SSA in Beijing was lower than that in AOE_Baotou, and an opposite spectral dependence was observed between these two sites. Baotou was affected by a greater quantity of industry emissions than Beijing, as it is a heavy industry city.

The VSD variation showed a more obvious distinction between dusty and non-dusty days. As Fig. 14 illustrates, the particle volumes of fine-mode aerosols are comparable with those of coarse-mode aerosols in Beijing and Baotou during non-dusty days. The strong dust storm caused a dramatic increase in coarse-mode particles. The volume median radius also showed differences between dusty and non-dusty days. The VSD peaks occurred at radii of ~2 μm with peak values of 1.05 and 1.8 on 4 May at Baotou and Beijing, respectively. Meanwhile, no significant variation was observed for fine-mode particles. It is observed that the volumes of both fine- and coarse-mode particles were large at AOE_Baotou on 2 May due to the combination of fine-mode aerosols with dust particles. This also explains the spectral SSA behaviour on that day.

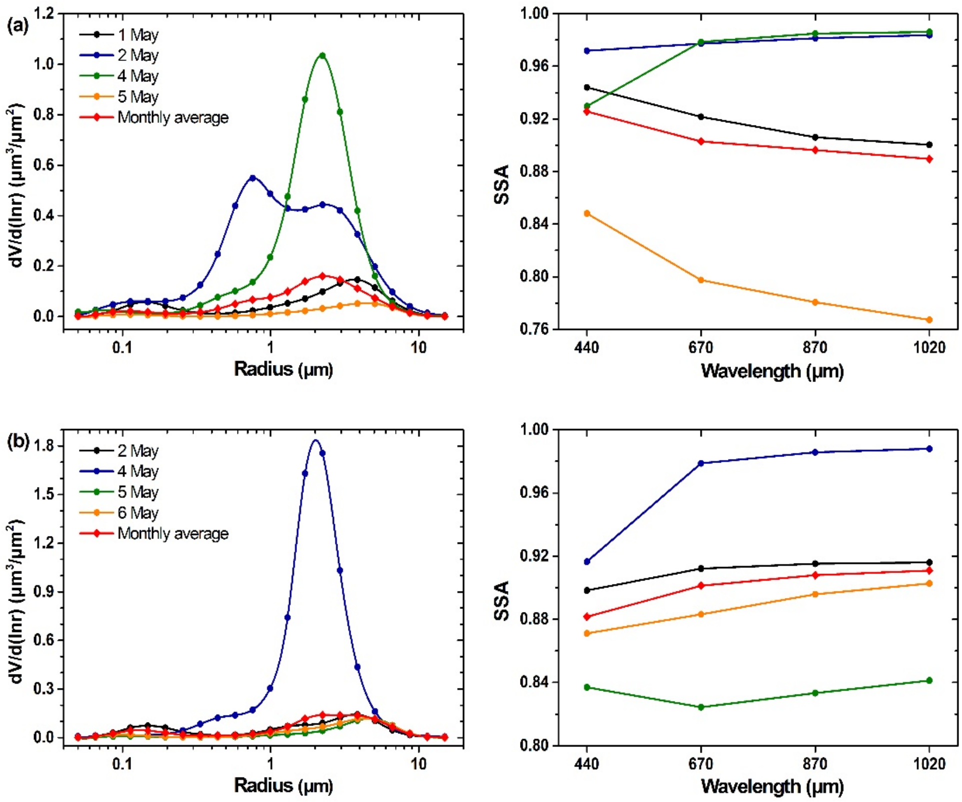

Fig. 14. Variations in the daily aerosol volume size distribution and spectral SSA during the dust event at (a) AOE_Baotou and (b) Beijing. Different colours represent different days, and the red curves represent the average VSD and SSA in May 2017. There was no VSD and SSA inversion product for Xuzhou-CUMT and Ussuriysk sites during May 3-8 2017.

4. Conclusions

In this study, we described a strong dust storm that occurred in northern China and Mongolia in early May 2017. The source and transport were investigated using multi-satellite data (including OMI, CALIPSO, and AHI), ground-based measurements (including PM measurements and AERONET observations), and HYSPLIT model computations. Benefiting from the high frequency of geostationary satellite observations, the rapid spatial-temporal variations in the dust plume were captured, including the continuous dust storms originating from the Gobi Desert region and different transport directions over China region. The OMI-AI and CALIPSO observations during the night-time provided more comprehensive information with larger coverage for the large-scale transport and vertical distribution of the dust plume. Intensive measurements (in both time and space) of the PM mass concentration revealed additional details when the region was covered by thick clouds and CALIPSO covered limited observation areas. The backward trajectories computed from the HYSPLIT model also confirmed the dust transport directions. From the combined observations, this severe dust storm was found to originate from the Gobi Desert, and travel to three different directions affecting large areas of China, including northern China, southeast China, and even Central China due to the strong winds. In addition, southern and eastern Russia and the Bering Sea were influenced by the long-distance transport of the strong dust plume. The aerosol properties (EAE, SSA, and VSD) have changed greatly during the dusty days as numerous large particles contributed to strong scattering and extinction. Overall, the combined observations of satellite and ground-based data contributed to the comprehensive monitoring of the source and long-distance transport of the dust storms, providing complete information on the spatial-temporal distribution.

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