



1     **Forecasting post-failure landslide mobility using a SPH model**  
2             **and data from ring shear strength tests: A case study**

3                     Miao Yu<sup>1</sup>, Yu Huang<sup>2,3\*</sup>, Wenbin Deng<sup>2</sup>, Hualin Cheng<sup>2</sup>

4     <sup>1</sup> Faculty of Engineering, China University of Geosciences, Wuhan, Hubei 430074,  
5     China

6     <sup>2</sup> Department of Geotechnical Engineering, College of Civil Engineering, Tongji  
7     University, Shanghai 200092, China

8     <sup>3</sup> Key Laboratory of Geotechnical and Underground Engineering of the Ministry of  
9     Education, Tongji University, Shanghai 200092, China

10    \* Corresponding author: Yu Huang (Tel.: +86-21-6598-2384; Fax: +86-21-6598-5210.  
11    E-mail: [yhuang@tongji.edu.cn](mailto:yhuang@tongji.edu.cn))

12    **Abstract** Flowlike landslides, such as flowslides and debris avalanches, have caused  
13    serious infrastructure damage and casualties for centuries. Effective numerical  
14    simulation incorporating accurate soil mechanical parameters is essential for  
15    predicting post-failure landslide mobility. In this study, smoothed particle  
16    hydrodynamics (SPH) incorporating soil ring shear test results was used to forecast  
17    the long-runout mobility for a landslide on an unstable slope in China. First, a series  
18    of ring shear tests under different axial stresses and shear velocities were conducted to  
19    evaluate the residual shear strength of slip zones after extensive shear deformation.  
20    Based on the ring shear test results, SPH modeling was conducted to predict the  
21    post-failure mobility of a previously identified unstable slope. The results indicate that



22 the landslide would cut a fire road on the slope after 12 s and cover an expressway at  
23 the foot of that slope after 36 s. In the model, the landslide would finally stop sliding  
24 about 25 m beyond the foot of the slope after 120 s. This study shows that combining  
25 the SPH method with ring shear test results to forecast landslide mobility can provide  
26 basic information for landside disaster mitigation.

27 **Keywords:** Landslide hazard; Post-failure mobility; Ring shear tests; Smoothed  
28 particle hydrodynamics (SPH); Residual strength

## 29 **1. Introduction**

30 Flowlike landslides triggered by intense earthquakes or rainfall, such as debris and  
31 rock avalanches, have caused serious infrastructure damage and casualties for  
32 centuries (Okada et al. 2007; Wang et al. 2005). This kind of landslide is commonly  
33 high-speed and has a long runout distance. For example, a large landslide in southern  
34 Italy in February, 2010, had a runout distance of 1.2 km and necessitated the  
35 evacuation of nearly 2,300 people. This landslide was triggered by heavy and  
36 prolonged rainfall between August 2009 and February 2010 (Gattinoni et al. 2012).  
37 The 2009 Shiaolin landslide in Taiwan, induced by a cumulative rainfall of nearly  
38 1700 mm from Typhoon Morakot, buried Shiaolin Village and resulted in more than  
39 400 people dead and missing (Tsou et al. 2011). Numerical simulations that  
40 incorporate accurate soil mechanical parameters are a powerful tool for simulating  
41 landslide runout distances; these simulations can provide fundamental reference  
42 information for landside disaster mitigation (Yerro et al. 2016; Žic et al. 2015).



43           The main numerical methods for simulating landslides are the discrete element  
44    methods and the continuum methods (Lu et al. 2014; Wu et al. 2017). Using a discrete  
45    element method, such as the distinct element method (DEM) or discontinuous  
46    deformation analysis (DDA), the nonphysical parameters cannot be determined  
47    exactly (Huang et al. 2014). However, continuum methods based on grids, like the  
48    finite element method (FEM) and the finite difference method (FDM), have the  
49    shortcomings of grid distortion and low accuracy for the numerical analysis of a  
50    landslide with a long runout. Recently, a new numerical method has been used to  
51    overcome these limitations, namely the smoothed particle hydrodynamics method  
52    (SPH) (Bui et al. 2008). This method is in the framework of continuum methods. SPH  
53    is a pure Lagrangian, meshless hydrodynamics method and it is capable of simulating  
54    flow deformation, free surfaces, and deformation boundaries (Liu and Liu 2003).  
55    Several studies have demonstrated the efficiency of the SPH method for the large  
56    deformation analysis post landslide. Huang et al (2014) provided a general view of  
57    SPH applications for solving large deformation and failure problems such as dam  
58    breaks, slope failure, and soil liquefaction flow. Pastor et al (2009) applied a  
59    depth-integrated, coupled SPH model successfully to simulate catastrophic flow-like  
60    landslides that occurred in southern Italy in 1998. Cascini et al (2014) proposed a  
61    SPH model to represent two actual flow-type events accurately. Cuomo et al (2016)  
62    used SPH to simulate flow-like landslides (debris flows and debris avalanches) and  
63    discussed the influence of bed entrainment on landslide propagation. Hu et al (2015)



64 conducted two- and three-dimensional SPH numerical simulation of flow-like  
65 landslides triggered by the 2008 Wenchuan earthquake in China and proposed that the  
66 SPH method is well-suited for modeling free surfaces, moving interfaces, and  
67 extensive deformation.

68 Study into the residual shear strength property of slip zones under large shear  
69 deformation is essential to landslide long-runout mechanism explanation (Tika and  
70 Hutchinson 1999; Wen et al. 2007). Because the physical sample displacement using  
71 conventional laboratory shear tests, like direct shear tests and triaxial shear tests, is  
72 limited to about 10 mm (Okada et al. 2007; Casagli et al. 2006; Van Asch et al 2007),  
73 the shear behavior for large shear displacements cannot be assessed by these methods  
74 (Dai et al. 2016). Ring shear tests, which can impart extremely large shear strains, may  
75 be the ideal laboratory tool for extensive shear deformation testing (Okada et al. 2007;  
76 ASTM Standard D7608-10, 2010). Several studies have applied ring shear tests to  
77 study the residual shear strength of soils (Fukuoka et al, 2007; Hoyos et al. 2014; Li et  
78 al. 2013; Wang et al. 2005). For example, Fukuoka et al (2007) applied a newly  
79 developed ring shear test to study shear zone development during large displacements.  
80 That study pointed out that a ring shear test is the most appropriate test for studying  
81 long-travel landslides. Kimura et al (2014) studied the effect of the shearing rate on  
82 the residual strength of landslide soils using ring shear tests. Zhang et al (2011) used  
83 ring shear tests to study the transform mechanism of the slide-debris flow under large



84 deformation. Li et al (2017) explored the residual strength of silty sand under different  
85 degrees of over consolidations and different shear rates using ring shear tests.

86 This study presents an effective numerical simulation method, namely SPH, that  
87 incorporates accurate soil mechanical parameters derived from ring shear tests. The  
88 aim is to predict the downslope flow after slope failure of a previously identified  
89 unstable slope and thereby provide basic information for landside disaster mitigation.  
90 First, this paper describes the geomorphological and geological setting, hydrogeology  
91 and rainfall, and triggering factors of the landslide examined for this case study.  
92 These descriptions are based on detailed fieldwork. Next, a series of ring shear tests  
93 under several different normal stresses and shear rates were performed to identify the  
94 shear strength of the landslide soil. Finally, a SPH-based numerical simulation of the  
95 landslide was run to predict the extent of the landslide and track the slide velocity at  
96 different times.

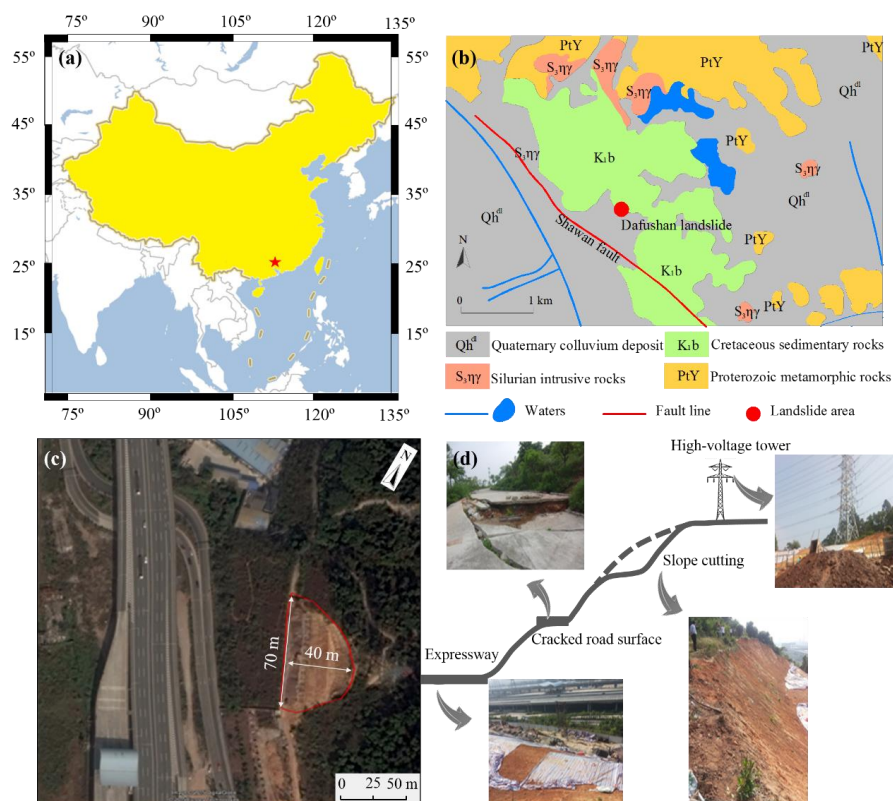
## 97 **2. A case study – the Dafushan landslide**

### 98 *2.1 Geomorphological and geological setting*

99 The Dafushan landslide, located in the Panyu District, Guangzhou City, South China,  
100 was selected for this case study (Fig. 1(a)). The slope is primarily composed of  
101 Cretaceous silty mudstone, conglomerate, and sandstone overlain by Quaternary silty  
102 clay (Yu et al. 2017) (Fig. 1(b)). The landslide is creeping from the northeast to the  
103 southwest covering an area of about 70 m × 40 m (Fig. 1(c)). The height difference  
104 between the toe and the crown is approximately 20 m with an average gradient of 25°.



105 The Dongxin expressway and a 50 t, high-voltage power line tower are located at the  
106 toe and top of the slope, respectively. In addition, there is a fire response service road  
107 that runs along the slope that is affected by the slide (Fig. 1(d)).



108  
109 **Fig. 1.** Overview of the Dafushan landslide. (a) Landslide location; (b)  
110 Geomorphologic and geologic map of the landslide area; (c) Aerial view of the  
111 unstable (potential landslide area (the area inside the red lines) (image from Google  
112 Earth®); (d) Engineering activities on the slope (reprinted from Yu et al. (2017) with  
113 permission of Springer).

114 *2.2 Landslide triggering factors*



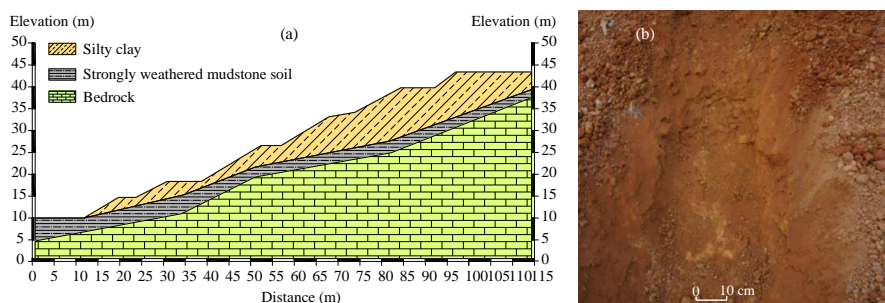
115 The ground was first found to be unstable in May 2013. This instability was  
116 manifested mainly by cracks in the ground surface and cracks in the  
117 round-the-mountain road. The road was built for fire response services in May 2011.  
118 The relevant departments repaired the damaged road immediately to guarantee the  
119 normal operation of the road. However, addition evidence of instability was found in  
120 the middle of August 2013 after a period of intense rainfall. The road was damaged  
121 again and the trees up the hill began to tilt. Based on preliminary field investigation,  
122 the main factors that triggered the landslide were deduced.

123 *(1) Hydrogeology and rainfall*

124 Rainfall is the main supply source of groundwater in the study area. The average  
125 annual rainfall is 1635.6 mm. Most of the rain falls between April and September; this  
126 rainfall accounts for 81% of the yearly precipitation. In the rainy season, the  
127 groundwater level rise significantly and reduces the shear strength of the soil.  
128 Combined with the rainfall flushing effect on the slope surface, the stability of the  
129 slope is decreased significantly.

130 *(2) Mechanical properties of landslide soil*

131 The shallow part of the landslide is mainly composed of silty clay (Fig. 2) and a  
132 strongly weathered mudstone soil with a low shear strength. These materials soften  
133 and disintegrate when wet, thus the slope is stable in the dry season but shows signs  
134 of instability in the rainy season.



135

136 **Fig. 2.** Geology and soil at the Dafushan landslide. (a) Longitudinal geologic section  
137 of the unstable slope shown in Fig. 1(c). (b) Photograph of the silty clay landslide soil.

### 138 (3) Human engineering activities

139 Human engineering activities impaired the natural stability of the slope. Two  
140 examples: a) to build the fire service road, a cut was made in the slope; b) the heavy  
141 high-voltage power line tower increases the downward pressure on the slope (Fig.  
142 1(d)).

### 143 3. Ring shear tests

144 A GCTS Residual Ring Shear Testing System (model SRS-150) produced by  
145 Geotechnical Consulting and Testing Systems (GCTS) in 2012 in the USA was used  
146 for the ring shear tests conducted for this study (Fig. 3). The SRS 150 is a fully  
147 automated electro-pneumatic and servo-controlled testing system used for  
148 determining the residual strength of continuously sheared soil. Shear torques of up to  
149 820 Nm can be applied, consolidation stress can be up to 1000 kPa, and unlimited  
150 angular rotation is allowed (Dai et al. 2016; Hoyos et al. 2014). The unit is capable of  
151 applying shearing rates of 0.001 to 360 degrees per minute continuously with



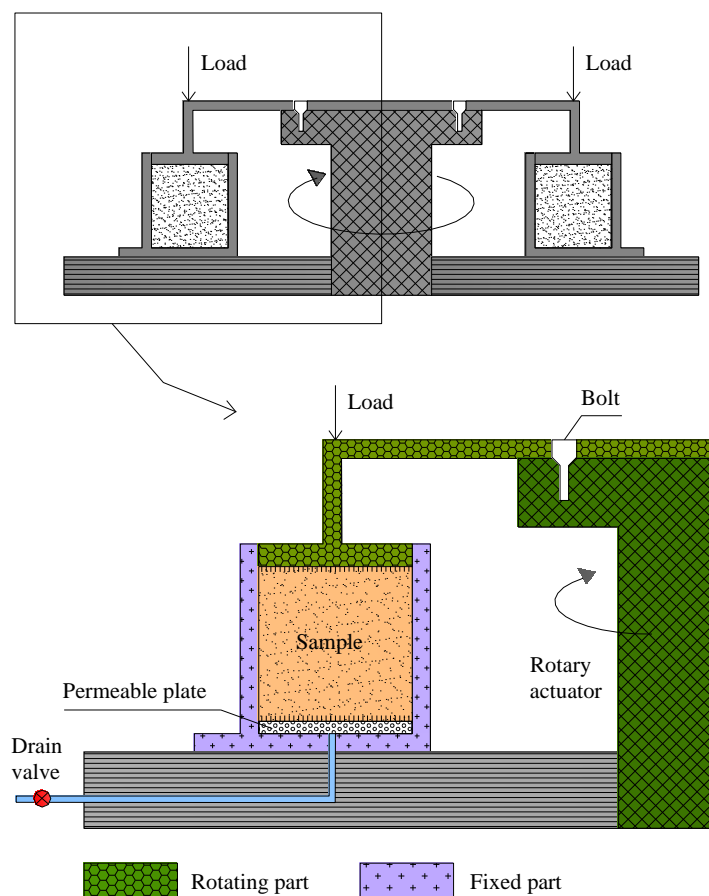


152 zero-backlash for replication of true in-situ strain rates during failure. (Hoyos et al.  
153 2011).



154  
155 **Fig. 3.** Photograph of the GCTS SR-150 Residual Ring Shear testing device and  
156 image of the GCTS software interface.

157 A schematic illustration of a sample in the apparatus is shown in Fig. 4. For  
158 testing granular materials, the device accepts ring-shaped samples with a 150 mm  
159 outer diameter and a 100 mm inner diameter. The sample is sheared by rotating the  
160 upper half of the testing unit and keeping the lower half motionless. Two types of  
161 shearing modes, either a shear speed control mode or a shear torque control mode, can  
162 be chosen.



163

164 **Fig. 4.** Schematic cross sections of ring shear apparatus shown in Figure 3.

165 *3.1 Sample preparation and test procedures*

166 The samples studied were samples of the silty clay soil from the Dafushan landslide

167 shown in Fig. 2(b). The soil's physical properties are listed in Table 1.

168 **Table 1** Physical properties of a soil from the Dafushan landslide.

Density	Dry density	Water	Liquid	Plastic	Plastic	Liquidity
$\rho$ (g/cm <sup>3</sup> )	$\rho_d$ (g/cm <sup>3</sup> )	content $\omega$ (%)	limit $\omega_L$	limit $\omega_P$	index $I_P$	index $I_L$



			(%)	(%)		
1.77	1.43	21.4	29.8	17.5	12.3	0.32

169 A series of ring shear tests were performed to determine the physical properties  
 170 of the landslide soil after it had been extensively sheared. The saturated soil sample  
 171 was first consolidated under a normal stress and then it was sheared to a residual state  
 172 under naturally drained conditions using the shear speed control mode of the ring  
 173 shear test system. For these tests, normal stresses of 50, 100, 200, 300, and 400 kPa  
 174 were used to consolidate the soil samples and different shear rates (1, 5, 10, 20 °/min)  
 175 were employed. Test parameters are listed in Table 2.

176 **Table 2** Consolidation stresses, shearing rates, and saturations for soil specimens subjected to  
 177 laboratory ring shear tests.

Test	Normal stress $\sigma$ (kPa)	Shear rate $\alpha$ (°/min)	Saturation (%)
1-1	50	5	100
1-2	100	5	100
1-3	200	5	100
1-4	300	5	100
1-5	400	5	100
2-1	200	1	100
2-2	200	5	100
2-3	200	10	100

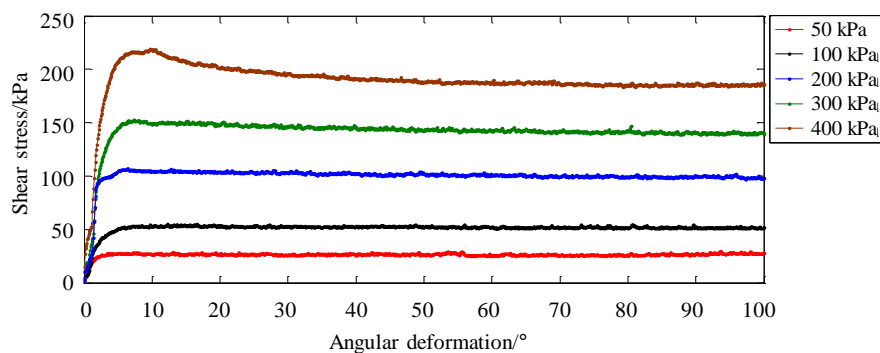


2-4	200	20	100
3-1	50	5	0
4-2	100	5	0
3-3	200	5	0
3-4	300	5	0
3-5	400	5	0

178 *3.2 Test results and discussion*

179 (1) Axial stress

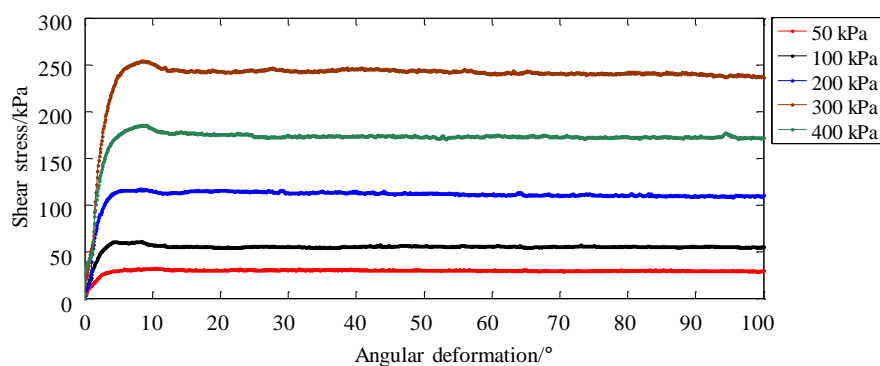
180 Figure 5 shows the relationships between shear stress and angular displacement  
181 under a shear rate of 5 °/min and axial stresses of 50, 100, 200, 300, and 400 kPa. At  
182 the same shear rate, shear strength increases with increasing axial stress. In the initial  
183 shear stages, shear stresses increase rapidly along with shear displacement and reach a  
184 peak shear strength. The greater the axial stresses, the larger the shear displacement at  
185 peak shear strength. When the axial stress is low (e.g., 50 kPa and 100 kPa), the shear  
186 stresses do not change after peak shear strength is reached. When the axial stress is  
187 high (e.g., 200 kPa, 300 kPa, or 400 kPa), the shear stresses decrease after peak shear  
188 strength but eventually stabilize. This stable strength is the residual shear strength and  
189 is the result of strain softening.



190

191

(a) Saturated soil



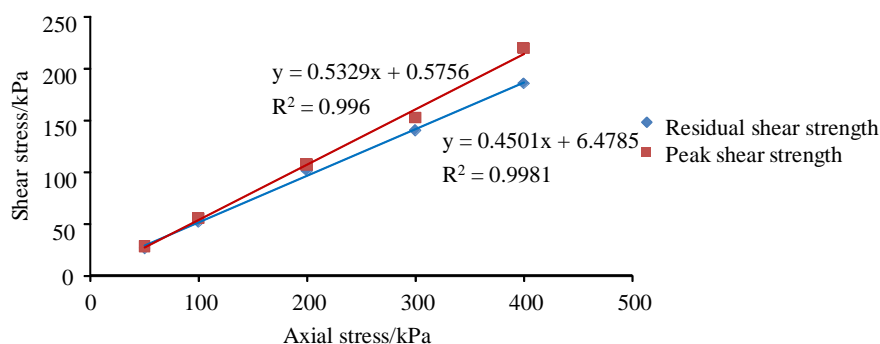
192

193

(b) Dry soil

194 **Fig. 5.** Shear stress–angular displacement curves for the landslide soil at a shear rate  
195 of 5°/min and different axial stresses for (a) saturated soil and (b) dry soil.

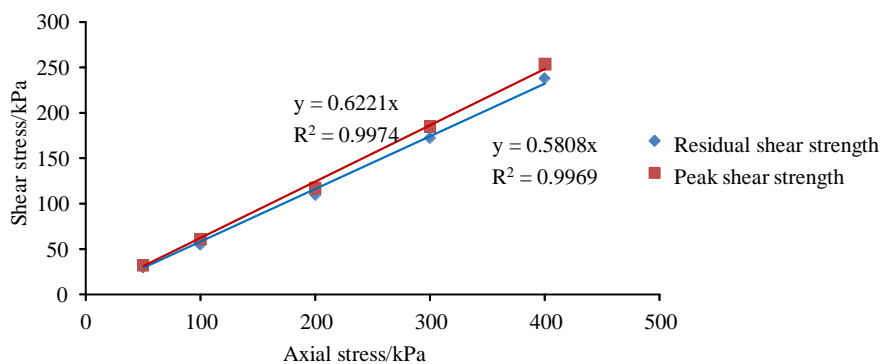
196 The residual strength envelope of the soil can be illustrated by plotting the shear  
197 stress against axial stress, as shown in Fig. 6.



198

199

(a) Saturated soil



200

201

(b) Dry soil

202 **Fig. 6.** Residual strength envelopes for the landslide soils; (a) saturated soils, (b) dry

203

soils.

204

Based on Coulomb's equation, the peak and residual shear strengths of the

205

landslide soil were obtained and are listed in Table 3. Because the main trigger for the

206

Dafushan landslide was heavy rain, the residual strength of saturated soil is used for

207

the numerical simulation presented in Section 4 of this paper.

208

**Table 3** Cohesion and internal friction for landslide soils at peak and residual shear

209

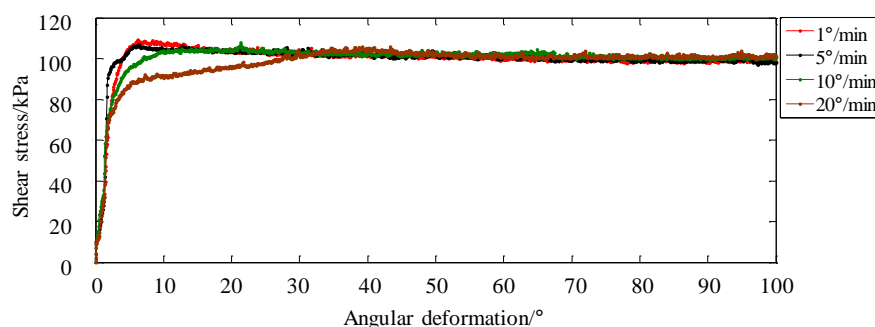
strengths calculated from the Coulomb (Mohr-Coulomb) equation.



Soil	Peak shear strength		Residual shear strength	
	Cohesion $c_r$ /kPa	Internal friction angle $\varphi_r$ /°	Cohesion $c_r$ /kPa	Internal friction angle $\varphi_r$ /°
Saturated soil	0.58	28.05	6.48	24.23
Dry soil	0	31.89	0	30.15

210 (2) Shear rate

211 Figure 7 shows the relationships between shear stress and angular deformation  
 212 under a normal stress of 200 kPa at shear rates of 1, 5, 10, and 20 °/min. As the shear  
 213 rate increases, the residual shear strengths increase slightly but the peak shear  
 214 strengths show the opposite reaction. However, the angular displacements at peak  
 215 shear strength increase significantly, as shown in Table 4.



216  
 217 **Fig. 7.** Shear stress–angular displacement curves for saturated landslide soil under  
 218 200 kPa axial stress.

219 **Table 4** Differences in shear strengths and angular displacements for saturated  
 220 landslide soil at different shearing rates.



Shearing rate (°/min)	Peak shear strength (kPa)	Residual shear strength (kPa)	Difference between peak and residual shear strength (kPa)	Angular displacement at peak shear strength (°)
1	109.10	99.35	9.75	6.264
5	107.00	99.52	7.48	6.444
10	105.00	100.55	4.45	16.992
20	105.80	100.99	4.81	39.168

221 To analyze the relationship between the residual shear strength of the saturated  
 222 soil and the shear strain rate, the residual shear stress–shear strain rate curve can be  
 223 drawn (Fig. 8). The formula for calculating the shear strain rate is:

$$224 \quad \dot{\gamma} = \frac{R\omega}{H} \quad (1)$$

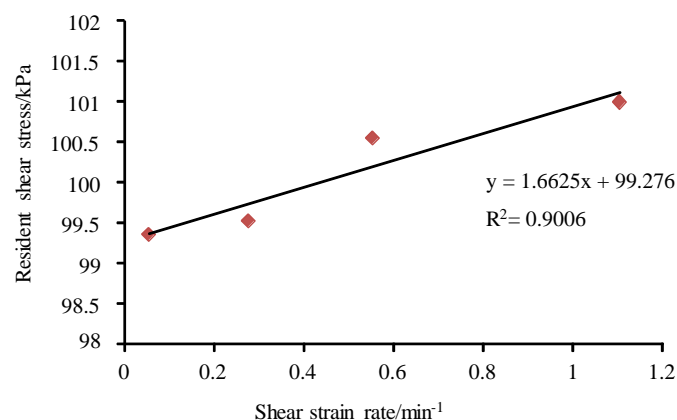
225 where  $\dot{\gamma}$  is the shear strain rate,  $R$  is the average radius of the sample,  $\omega$  is the  
 226 angular velocity,  $H$  is the sample height.

227 As shown in Fig. 8, the residual shear strength of the saturated soil as determined  
 228 by these experiments increases linearly with shear strain rate. This result agrees with  
 229 the results reported by Li et al. (2013) and Dai et al. (2016). This relationship is  
 230 similar to the behavior of a viscous fluid and can be expressed by Eq. (2):

$$231 \quad \tau = \eta\dot{\gamma} + f(\sigma) \quad (2)$$

232 where  $\tau$  is shear stress,  $\eta$  is the coefficient of viscosity. The intercept  $f(\sigma)$   
 233 represents the shear stress when the shear strain rate equals 0.





234

235 **Fig. 8.** Residual shear stress–shear strain rate curves for the saturated landslide soil.

#### 236 **4. SPH-based numerical simulation for landslides**

##### 237 *4.1 Calculation principles and SPH process methods*

###### 238 *(1) Basic SPH concepts*

239 Smoothed particle hydrodynamics is a mesh-free and fully Lagrangian method based  
240 on fluid dynamics. In Lagrangian models, the coordinates move with the medium  
241 being modeled. The continuous medium is discretized into a series of arbitrarily  
242 distributed discrete elements (called particles) and field variables (like energy,  
243 velocity, density, or any other variable) for each particle can be calculated in the form  
244 of SPH (Dao et al. 2013; Huang and Dai 2014).

245 The SPH method is built on interpolation theory with two essential  
246 approximations. These approximations are smoothing and the particle (Huang et al.  
247 2014). The smoothing approximation, also known as kernel approximation, describes  
248 a function in a continuous form as an integral representation. The particle  
249 approximation means that the value of a function for a particle can be determined by



250 the average value of all the particles in the support domain. The smoothing and the  
 251 particle approximations can be expressed, respectively, by the following two  
 252 equations:

$$253 \quad \langle f(x) \rangle = \int_{\Omega} f(x') W(x-x', h) dx' \quad (3)$$

$$254 \quad \langle f(x) \rangle = \sum_{j=1}^N m_j \frac{f_j(x)}{\rho_j} W(x-x', h) \quad (4)$$

255 where the angle brackets represent a kernel approximation,  $x$  is the location  
 256 vector of the particle,  $x'$  denotes neighboring particle in the support area,  $W$  is the  
 257 smoothing function,  $h$  stands for the smoothing length,  $\Omega$  stands for the volume of the  
 258 integral that contains  $x$ ,  $m$  is the mass, and  $\rho$  is the density,  $N$  is the total number of  
 259 particles.

### 260 (2) Governing equations

261 The Navier–Stokes equations in a computational fluid dynamics framework are used  
 262 as governing equations in this study. The equations of continuity and motion in the  
 263 SPH version can be expressed as:

$$264 \quad \frac{d\rho_i}{dt} = \sum_{j=1}^N m_j (u_i^\beta - u_j^\beta) \frac{\partial W_{ij}}{\partial x_i^\beta} \quad (5)$$

$$265 \quad \frac{du_i^\alpha}{dt} = \sum_{j=1}^N m_j \left[ \frac{\sigma_i^{\alpha\beta}}{(\rho_i)^2} + \frac{\sigma_j^{\alpha\beta}}{(\rho_j)^2} \right] \frac{\partial W_{ij}}{\partial x_j^\beta} + F_i \quad (6)$$

266 where  $W_{ij}$  represents the smoothing function of particle  $I$  calculated at particle  $j$ ,  $t$   
 267 is time,  $u$  denotes the velocity vector,  $\sigma$  is the stress tensor,  $F$  represents the vector of  
 268 external force, and  $\alpha$  and  $\beta$  are the coordinate directions.

### 269 (3) Model for a landslide simulation



270 The Bingham model has been proved as one of the most effective models for runout  
271 simulation of flowlike landslides (Marr et al. 2002; Moriguchi et al. 2009). In this  
272 paper, the Bingham flow model is also adopted as the constitutive model for the  
273 Dafushan landslide in this study. The relationship between shear stress and strain rate  
274 can be written as:

$$275 \quad \tau = \eta \dot{\gamma} + \tau_y. \quad (7)$$

276 Equation (8) can be modified by combining it with the Mohr-Coulomb yield  
277 criterion to yield (Moriguchi et al. 2009):

$$278 \quad \tau = \eta \dot{\gamma} + \sigma \tan \varphi + c \quad (8)$$

279 where  $\tau$  denotes the shear stress,  $\eta$  and  $\tau_y$  represent the Bingham yield viscosity  
280 and stress, respectively,  $\dot{\gamma}$  is the shear strain rate,  $\sigma$  is the pressure,  $\varphi$  is the friction  
281 angle, and  $c$  is the cohesion.

282 For this study, the concept of equivalent viscosity was adopted to better integrate  
283 the Bingham model into the SPH framework. The equivalent viscosity can be  
284 expressed as:

$$285 \quad \eta' = \eta + \tau_y / \dot{\gamma}. \quad (9)$$

286 The maximum value was defined by Uzuoka et al. (1998) as:

$$287 \quad \eta' = \eta_0 + \frac{\tau_y}{\dot{\gamma}} \quad \text{when } \eta' < \eta_{\max} \quad (10)$$

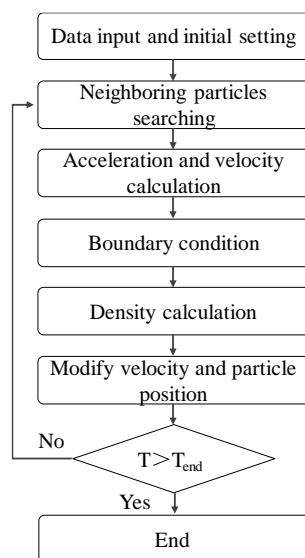
$$288 \quad \eta' = \eta_{\max} \quad \text{when } \eta' > \eta_{\max} \quad (11)$$

289 where  $\eta_{\max}$  is the maximum value of  $\eta'$ .

290 *(4) Procedure for the numerical simulation*



291 A flow chart for the SPH numerical simulation is shown as Fig. 9. Details about how  
292 the calculations are carried out can be found in Huang et al. 2014. The accuracy of  
293 SPH program in landslide modelling was also fully validated in Huang et al. 2014.



294

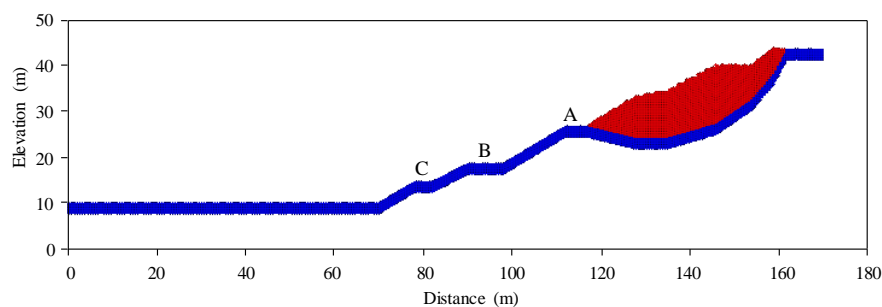
295 **Fig. 9.** Flow chart for the SPH numerical simulation used in this study.

#### 296 4.2 Dafushan landslide SPH simulation and results

297 Based on a terrain model derived from an unmanned aerial vehicle and  
298 structure-from-motion (Yu et al. 2017), an SPH simulation of the failure process of  
299 the Dafushan landslide was conducted. This simulation was used to assess the  
300 landslide's effects when failure occur. The numerical model is calculated on the basis  
301 of a total of 3,242 particles, 1,537 particles for the slide mass and 1,705 for the fixed  
302 boundary. Figure 10 is a longitudinal section of the model slide with the particles in  
303 the slide mass shown in red, the boundary particles shown in blue. The diameter of



304 each particle is 0.5 m. The soil particles in the model can be deformed in both the  
 305 vertical and horizontal directions under gravitational force in the vertical direction.



306  
 307 **Fig. 10.** Longitudinal section of the SPH numerical model of the Dafushan landslide.  
 308 The particles representing the slide mass are shown in red, the particles representing  
 309 the fixed boundary are shown in blue.

310 Table 5 lists the parameters used in the SPH simulation of the landslide. The  
 311 shear strength parameters listed in Table 5,  $c$  and  $\phi$ , are the values calculated from the  
 312 ring shear tests (Table 3).

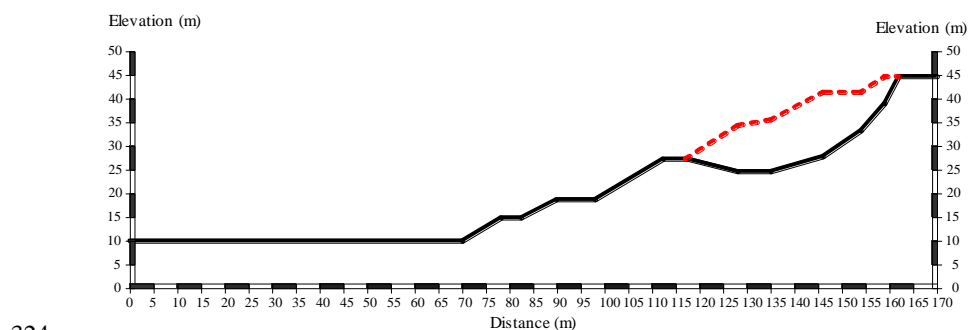
313 **Table 5** Parameters used in the SPH simulation of the Dafushan landslide.

Density $\rho$ (kg/m <sup>3</sup> )	1770
Residual cohesion $c$ (kPa)	6.48
Residual internal friction Angle $\phi$ (°)	24.23
Acceleration of gravity $g$ (m/s <sup>2</sup> )	9.80
Unit time step $\Delta t$ (s)	0.003
Time step ( $n$ )	40000

314 Figures 11(a)-11(g) show the flow process of Dafushan landslide predicted by  
 315 the SPH simulation. In Fig. 11, the solid black line represents the bed on which the



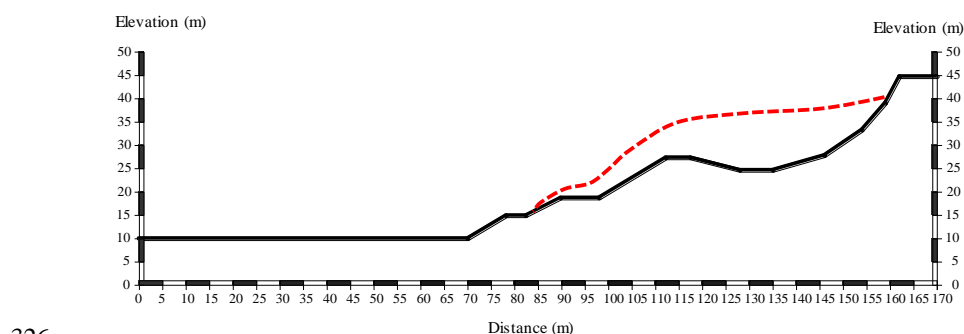
316 mass slides, the red dashed line represents the SPH-modeled ground surface. At time  $t$   
317  $= 0$ , this red line is the ground surface before slide failure. For times after  $t = 0$ , it is  
318 the top surface of the flowing mass of soil that constitutes the moving landslide mass  
319 as predicted by the SPH simulation results. In the model, the time the failed Dafushan  
320 landslide lasts, from initiation to the whole landslide mass coming to rest, is 120 s.  
321 The model predicts that the landslide would cut the fire road at  $t = 12$  s and cover the  
322 expressway at  $t = 36$  s. When the landslide stops sliding at 120 s, slide material would  
323 cover about a 25 m wide swath of ground beyond the foot of the topographic slope.



324

325

(a)  $t = 0$  s



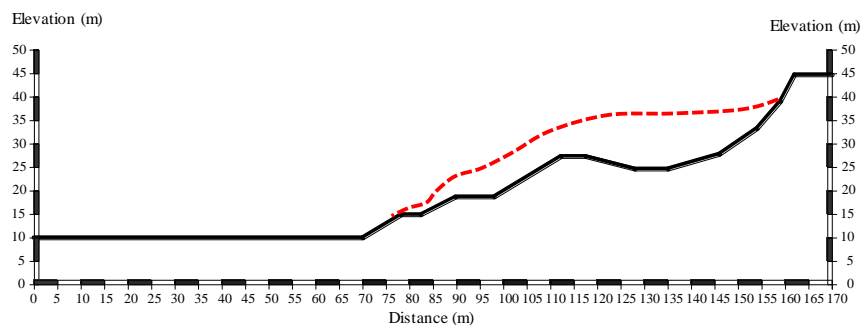
326

327

(b)  $t = 12$  s



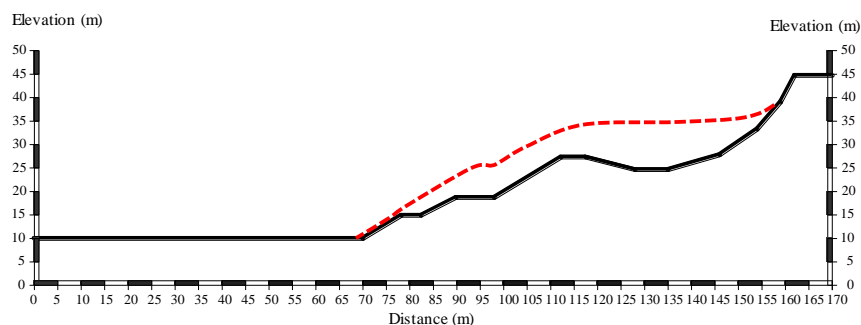
328



329

(c)  $t = 24$  s

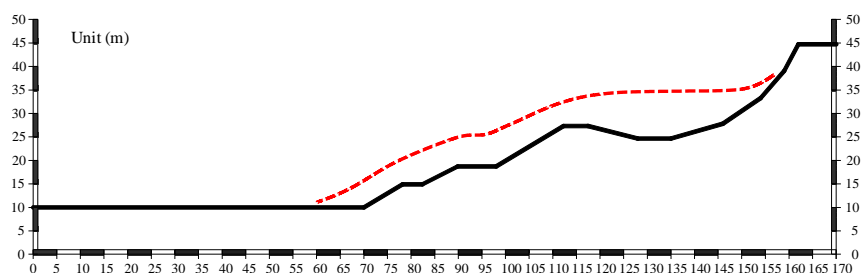
330



331

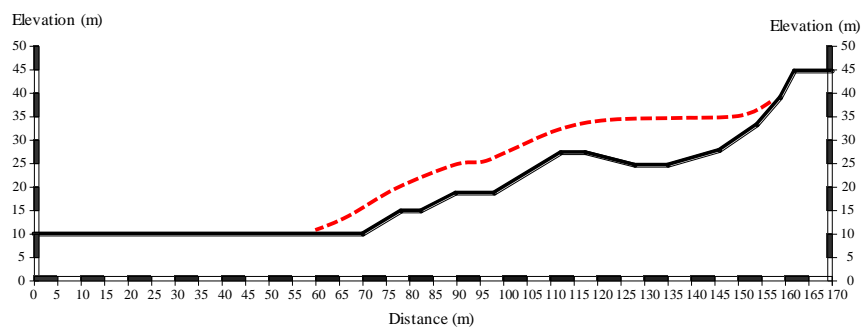
(d)  $t = 36$  s

332





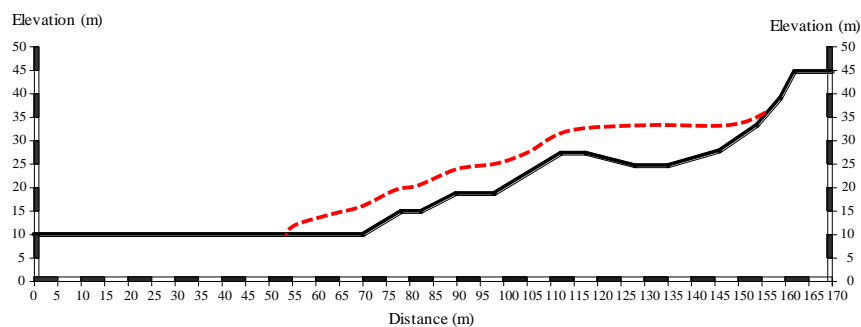
333



334

(e)  $t = 60$  s

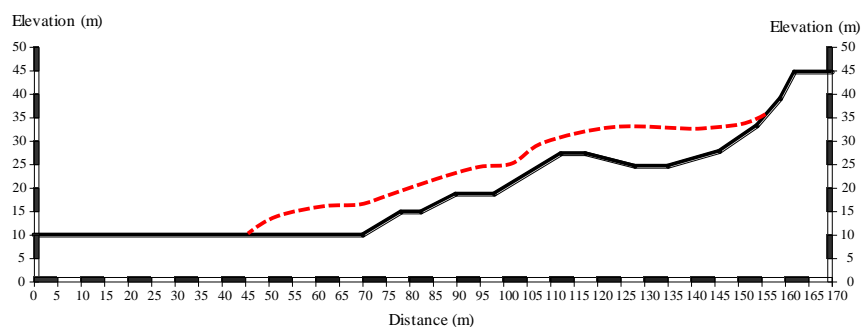
335



336

(f)  $t = 84$  s

337



338

(g)  $t = 120$  s

339 **Fig. 11.** Longitudinal profiles showing the results of the SPH forecasting model. The

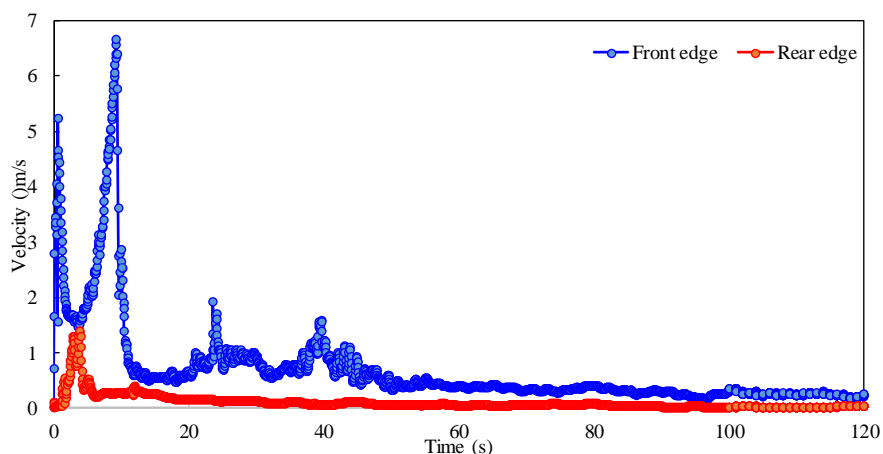
340 panels represent the outline of the Dafushan landslide from the time the slide is





341 initiated at  $t = 0$  s (panel a) through the slide finally coming to rest at  $t = 120$  s (panel  
342 g).

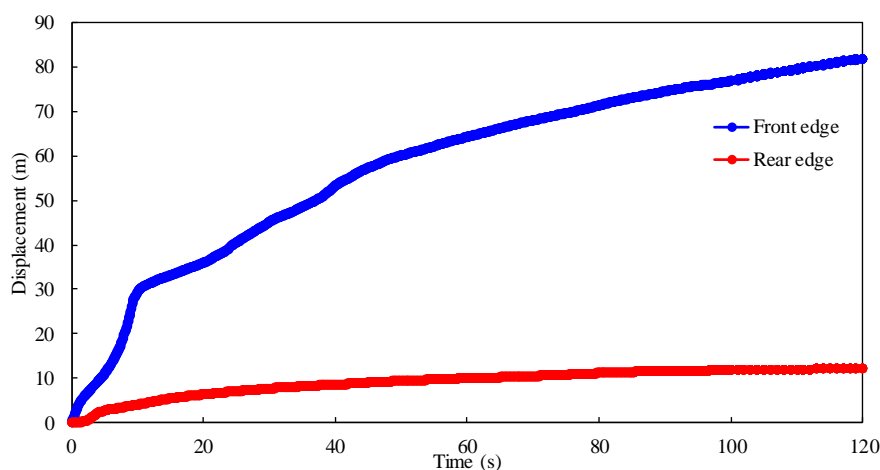
343 Because this SPH simulation is a Lagrangian method, it can track the velocity  
344 and displacement of each particle accurately. The velocity and displacement curves  
345 for the front and rear edges of simulated landslide are shown in Figs. 12 and 13. As  
346 shown in Fig. 12, the velocity of the front edge increases rapidly after slope failure  
347 begins and reaches three velocity peaks as the slide passes the three steps labeled A, B,  
348 and C shown in Fig. 10. The speed of the front and the times after initiation that it  
349 reaches these three steps are 5.23 m/s at 0.6 s at step A, 6.66 m/s at 9.3 s at step B,  
350 and 1.92 m/s at 23.6 s at step C. Unlike the front edge of the landslide, the velocity of  
351 the landslide's rear edge shows only a single peak. The maximum speed is 1.40 m/s;  
352 this appears 3.8 s after the slide is initiated.



353  
354 **Fig. 12.** Velocity curve of the front and rear edge of Dafushan landslide as predicted  
355 by the SPH model.



356 According to the Fig. 13, the maximum flow distances of the front and rear edge  
357 are up to 82 m and 12.3 m, respectively. The front edge of the slide will destroy the  
358 fire road about 10–12 s after the slide starts and reach the highway at  $t = 36$  s.  
359 Thereafter, the velocity gradually approaches zero as the flow distance increases. The  
360 maximum distance the landslide flows is approximately 82 m, and the speed of the  
361 flow can be divided into three stages. The flow is fastest from 0–10 s, slower from  
362 10–45 s, and relatively slow from 45–120 s. However, once signs of failure are  
363 observed at the Dafushan landslide site, evacuation of personnel and vehicles within  
364 about 25 m of the slope should begin immediately.



365

366 **Fig. 13.** Displacement curve of the front and rear edge of Dafushan landslide as  
367 predicted by the SPH model.

## 368 5. Conclusions

369 In this study, the SPH method incorporating soil mechanical parameters derived from  
370 ring shear tests is used to predict the flow of a potential landslide that could develop



371 on an unstable slope in Guangzhou City, China. This study provides basic information  
372 for landside disaster mitigation. The conclusions are:

373 (1) Under the same shear rate, soil shear strength increases with increasing axial  
374 stress. For the conditions used in this study, under high axial stress ( $> 200$  kPa) the  
375 soil exhibits strain softening.

376 (2) During ring shear tests, as the shear rate increases, the residual shear  
377 strengths increase slightly but the peak shear strengths decrease as the angular  
378 displacements at peak shear strength increase significantly.

379 (3) A SPH-based numerical simulation of the potential Dafushan landslide  
380 conducted to predict the scope of the landslide and track the slide velocity at different  
381 times shows that the landslide would cut the fire road at  $t = 12$  s and cover the  
382 expressway at  $t = 36$  s. And once signs of failure are observed at the Dafushan  
383 landslide site, evacuation of personnel and vehicles within about 25 m of the slope  
384 should begin immediately.

### 385 **Acknowledgements**

386 This work was supported by the National Science Fund for Distinguished Young  
387 Scholars of China (Grant No. 41625011), the Fundamental Research Funds for  
388 National University, China University of Geosciences (Wuhan) (Grant No.  
389 CUGL170806), and the National Key Technologies R&D Program of China (Grant  
390 No. 2012BAJ11B04).

### 391 **References**



- 392 [1] ASTM Standard D7608-10 (2010) Standard Test Method for Torsional Ring  
393 Shear Test to Determine Drained Fully Softened Shear Strength and Nonlinear  
394 Strength Envelope of Cohesive Soils (Using Normally Consolidated Specimen)  
395 for Slopes with No Preexisting Shear Surfaces. ASTM International, West  
396 Conshonocken, PA
- 397 [2] Bui HH, Fukagawa R, Sako K, Ohno S (2008) Lagrangian meshfree particles  
398 method (SPH) for large deformation and failure flows of geomaterial using  
399 elastic–plastic soil constitutive model. *International Journal for Numerical and*  
400 *Analytical Methods in Geomechanics* 32(12): 1537–1570
- 401 [3] Casagli N, Dapporto S, Ibsen M L, Tofani V, Vannocci P (2006) Analysis of the  
402 landslide triggering mechanism during the storm of 20th–21st November 2000, in  
403 Northern Tuscany. *Landslides* 3(1): 13–21
- 404 [4] Cascini L, Cuomo S, Pastor M, Sorbino G, Piciullo L (2014) SPH run-out  
405 modelling of channelised landslides of the flow type. *Geomorphology* 214(2):  
406 502–513
- 407 [5] Cuomo S, Pastor M, Capobianco V, Cascini L (2016) Modelling the space–time  
408 evolution of bed entrainment for flow-like landslides. *Engineering Geology* 212:  
409 10–20
- 410 [6] Dai, Z., Huang, Y., Deng, W., Jiang, F., & Wang, D. (2016) Constitutive flow  
411 behavior of a municipal solid waste simulant at post-failure: experimental and  
412 numerical investigations. *Environmental Earth Sciences* 75(11): 1–9



- 413 [7] Dao MH, Xu H, Chan ES, Tkalich P (2013) Modelling of tsunami-like wave  
414 run-up, breaking and impact on a vertical wall by sph method. *Natural Hazards &*  
415 *Earth System Sciences* 13(12): 3457–3467
- 416 [8] Fukuoka H, Sassa K, Wang G (2007) Influence of shear speed and normal stress  
417 on the shear behavior and shear zone structure of granular materials in naturally  
418 drained ring shear tests. *Landslides* 4(1): 63–74
- 419 [9] Gattinoni P, Scesi L, Arieni L, Canavesi M (2012) The February 2010 large  
420 landslide at Maierato, Vibo Valentia, Southern Italy. *Landslides* 9(2): 255–261
- 421 [10] Hoyos LR, Velosa CL, Puppala AJ (2011) A servo/suction-controlled ring shear  
422 apparatus for unsaturated soils: Development, performance, and preliminary  
423 results. *Geotechnical Testing Journal* 34(5): 413–423
- 424 [11] Hoyos LR, Velosa CL, Puppala AJ (2014) Residual shear strength of unsaturated  
425 soils via suction-controlled ring shear testing. *Engineering Geology* 172: 1–11
- 426 [12] Hu M, Liu MB, Xie MW, Liu GR (2015) Three-dimensional run-out analysis and  
427 prediction of flow-like landslides using smoothed particle hydrodynamics.  
428 *Environmental Earth Sciences* 73(4):1629–1640
- 429 [13] Huang Y, Dai ZL (2014) Large deformation and failure simulations for  
430 geo-disasters using smoothed particle hydrodynamics method. *Engineering*  
431 *Geology* 168: 86–97
- 432 [14] Huang Y, Dai ZL, Zhang WJ (2014) Geo-disaster modeling and analysis: an  
433 SPH-based approach. Springer, Heidelberg, pp 184–185



- 434 [15]Kimura S, Nakamura S, Vithana SB, Sakai K (2014) Shearing rate effect on  
435 residual strength of landslide soils in the slow rate range. *Landslides* 11(6):  
436 969–979
- 437 [16]Li DY, Yin KL, Glade T, Leo C (2017) Effect of over-consolidation and shear  
438 rate on the residual strength of soils of silty sand in the Three Gorges Reservoir.  
439 *Scientific Reports* 7(1): 5503
- 440 [17]Li YR, Wen BP, Aydin A, Lu NP (2013) Ring shear tests on slip zone soils of  
441 three giant landslides in the Three Gorges Project area. *Engineering Geology*  
442 154(2): 106–115
- 443 [18]Liu GR, Liu MB (2003) Smoothed particle hydrodynamics: a meshfree particle  
444 method. World Scientific Press, Singapore
- 445 [19]Lu CY, Tang CL, Chan YC, Hu JC, Chi CC (2014) Forecasting landslide hazard  
446 by the 3D discrete element method: A case study of the unstable slope in the  
447 Lushan hot spring district, central Taiwan. *Engineering Geology* 183: 14–30
- 448 [20]Marr J G, Elverhøi A, Harbitz C, Imran J, Harff P (2002) Numerical simulation  
449 of mud-rich subaqueous debris flows on the glacially active margins of the  
450 Svalbard–Barents Sea. *Marine Geology* 188(3–4): 351–364
- 451 [21]Moriguchi S, Borja R I, Yashima A, Sawada K (2009) Estimating the impact  
452 force generated by granular flow on a rigid obstruction. *Acta Geotechnica* 4(1):  
453 57–71



- 454 [22]Okada Y, Ochiai H, Okamoto T, Sassa K, Fukuoka H, Igwe, O (2007) A complex  
455 earth slide–earth flow induction by the heavy rainfall in July 2006, Okaya City,  
456 Nagano Prefecture, Japan. *Landslides* 4(2): 197–203
- 457 [23]Pastor M, Haddad B, Sorbino G, Cuomo S, Drempetic V (2009) A  
458 depth-integrated, coupled SPH model for flow-like landslides and related  
459 phenomena. *International Journal for Numerical and Analytical Methods in*  
460 *Geomechanics* 33(2): 143–172
- 461 [24]Tika TE, Hutchinson JN (1999) Ring shear tests on soil from the Vaiont landslide  
462 slip surface. *Geotechnique* 49(1): 59–74.
- 463 [25]Tsou CY, Feng ZY, Chigira M (2011) Catastrophic landslide induced by typhoon  
464 Morakot, Shiaolin, Taiwan. *Geomorphology* 127(3): 166–178
- 465 [26]Uzuoka R, Yashima A, Kawakami T, Konrod JM (1998) Fluid dynamics based  
466 prediction of liquefaction induced lateral spreading. *Computers and Geotechnics*  
467 22(3–4): 234–282
- 468 [27]Van Asch TW, Van Beek LPH, Bogaard TA (2007) Problems in predicting the  
469 mobility of slow-moving landslides. *Engineering geology* 91(1): 46–55
- 470 [28]Wang G, Suemine A, Furuya G, Kaibori M, Sassa K (2005) Rainstorm-induced  
471 landslides at Kisawa village, Tokushima Prefecture, Japan, August 2004.  
472 *Landslides* 2(3): 235–242



- 473 [29]Wen BP, Aydin A, Duzgoren-Aydin NS, Li YR, Chen HY, Xiao SD (2007)  
474 Residual strength of slip zones of large landslides in the Three Gorges area,  
475 China. *Engineering Geology* 93(3): 82–98
- 476 [30]Wu JH, Lin WK, Hu HT (2017) Assessing the impacts of a large slope failure using  
477 3DEC: The Chiu-fen-erh-shan residual slope. *Computers & Geotechnics* 88:32–45.
- 478 [31]Yerro A, Alonso EE, Pinyol NM (2016) Run-out of landslides in brittle soils.  
479 *Computers & Geotechnics* 80:427–439
- 480 [32]Yu M, Huang Y, Zhou JM, Mao LY (2017) Modeling of landslide topography  
481 based on micro-unmanned aerial vehicle photography and structure-from-motion.  
482 *Environmental Earth Sciences* 76(15): 520
- 483 [33]Zhang M, Yin Y, Hu R, WS, Zhang Y (2011) Ring shear test for transform  
484 mechanism of slide–debris flow. *Engineering Geology* 118(3): 55–62
- 485 [34]Žic E, Arbanas Ž, Bićanić N, Ožanić N (2015) A model of mudflow propagation  
486 downstream from the grohovo landslide near the city of rijeka (croatia). *Natural*  
487 *Hazards & Earth System Sciences* 15(2): 293–313