

Response to Reviewers of the manuscript

"Global detection of rainfall triggered landslide clusters"

by Benz and Blum, submitted to *Natural Hazards and Earth System Sciences*.

Manuscript Number: nhess-2018-391

Revision due before: 29 May 2019

Reviewer #1 comments:

Comment I:

The submitted paper deals with the identification of clusters in the Global Landslide Catalogue (GLC) in relation to individual rainfall events extracted from CHIRPS precipitation data. It is therefore also related to the investigation of biasing effects in small-scale (global) landslide catalogues that are not compiled through landslides inventorying but rather based on media or governmental reports. In this respect, the paper is timely, interesting and well suited to NHESS. The presentation is clearly structured and the language of the article is fluent, the Figures are of good quality. The clustering algorithm proposed by the authors to relate reported landslides to individual rainfall events is interesting and may be worth to be published. However, the validity of the constraints of the clustering algorithm is not discussed in detail by the authors. In this respect, a sensitivity analysis is missing in such that different time span thresholds in precipitation may be tested in order to investigate the behavior of the identified clusters. Also, the effect of different values for the spearman coefficient to discriminate clusters may be investigated.

Reply: Thank you very much for the kind words and constructive comments.

We agree and therefore included a sensitivity analysis, which investigates the different time span thresholds in precipitation and the effects of different spearman coefficients. The following discussion is now added:

“The threshold value of the spearman correlation coefficient was determined by testing the robustness of the identified clusters for different threshold values between zero and one (Fig. S2). Our results indicate that mean duration, area, and number of landslides per cluster are comparably robust to changes of the spearman correlation coefficient. In contrast maximum duration, area and number of landslides per cluster change drastically for different threshold values. From a correlation coefficient threshold of 0.35 to 0.7, maximum number of landslide events per cluster decreases from close to 500 to slightly above 100, maximum duration decreases from more than 80 days to approximately 25, and area decreases from 60,000,000 km² (approximately 1/3 of the planet’s surface area) to 200,000 km². For threshold values greater 0.7, only minor changes are observed. Hence, the latter was set as the correlation threshold value for this study (Fig. S2).

Additionally, we tested the robustness of the method to the time period of precipitation for which the correlation coefficient was determined (Fig. S3). It appears that the number of days is much less influential than the set correlation coefficient threshold (Fig. S2). Again, maximum number of landslides, area, and duration are impacted most, however remain stable for time period longer than 30 days prior to the second event.”

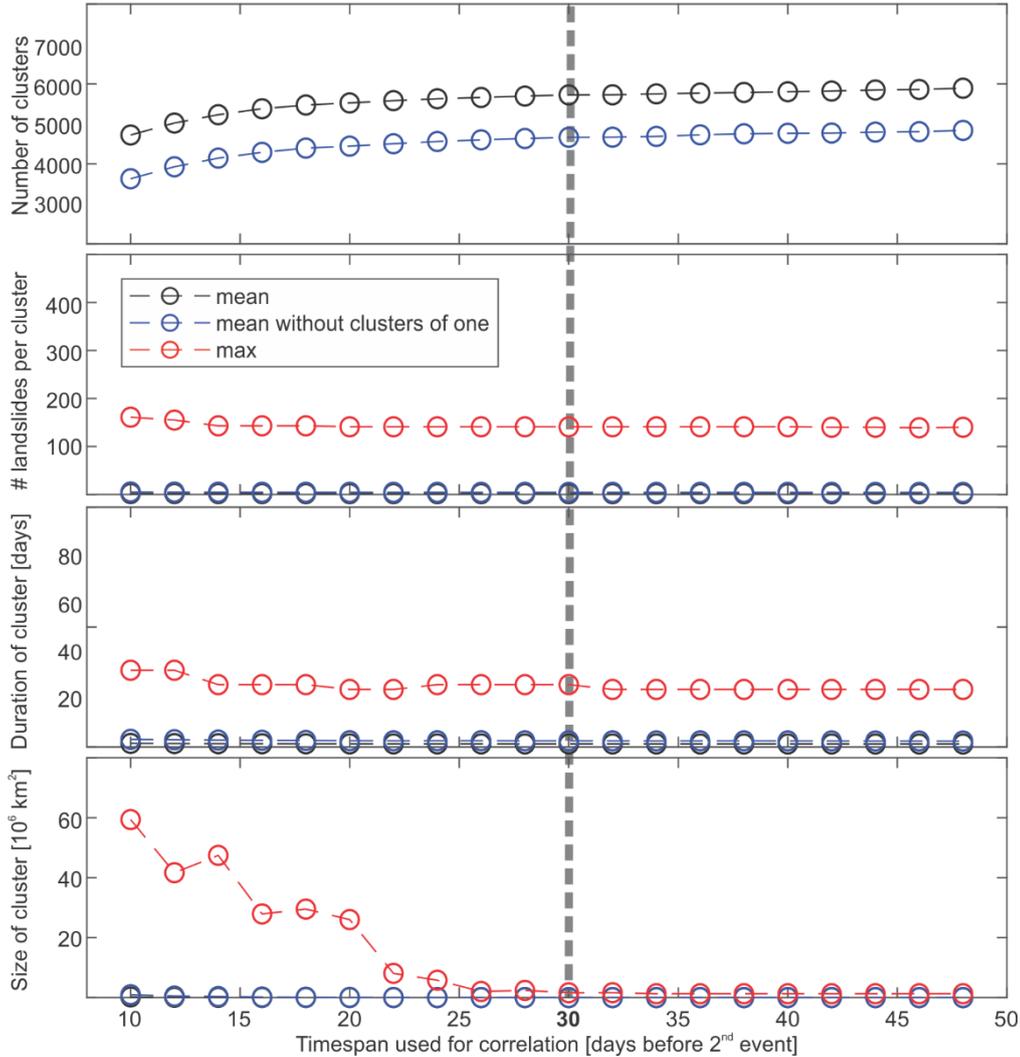


Figure S3. Impact of the chosen threshold for the timespan for which the Spearman correlation coefficient is determined on the total number of clusters in the global landslide catalog, on the average number of landslides per cluster, on the average duration of landslide clusters, and on the average area of landslide clusters. The correlation coefficient threshold was set to 0.7 for this analysis. In this study a threshold of 30 days was chosen, as from this point onwards number of clusters, maximums size of, duration of and landslides per cluster, becomes stable.

Comment II:

Another issue is that the only constraint applied for landslide clustering is precipitation. Environmental information like e.g. climatic setting, subsoil lithology or relief parameters are

not introduced into the clustering algorithm and the effect of introducing those is not investigated or discussed.

Reply: Correct, the introduced algorithm focus solely on rainfall, detecting clusters of landslides triggered by the same rainfall event. While different lithology and relief parameters impact the rainfall intensity-duration threshold, two landslide events located in different areas with different thresholds might still be triggered by the same rainfall event. We therefore decided to not include any of these parameters in the algorithm. However, an additional sentence was added to the chapter describing the algorithm that discusses this issue:

“The introduced algorithm is independent of subsoil topography and relief parameters. While these impact the precipitation intensity-duration threshold that is commonly expected to trigger landslides, locations with different thresholds might still experience landslides triggered by the same rainfall event.”

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Reviewer #2 comments:

Comment I:

The paper proposes an algorithm to detect and group clusters of landslide events that occurred or were triggered by the same rainfall event. The algorithm is then applied to the Global Landslide Catalogue (GLC). The paper has a good structure, even if, some improvements are needed to increase its quality and clarity. The research topic, from my point of view, is useful and of interest. In the following my revision.

The abstract should contain a description of the main aim and the innovative features of the research proposed. In my opinion, the abstract focuses too much on the results. The rows 14-20 should be summarized. It is not useful, and probably counterproductive, going too much into detail in the abstract. Then, I would suggest the authors to describe the aim of the algorithm and the innovative features of the re-search, focusing on how the paper is pushing a step forward in this topic.

Reply: We agree. Hence the abstract was rewritten accordingly:

“An increasing awareness of the cost of landslides on the global economy and of the associated loss of human life, has led to the development of various global landslide databases. However, these databases typically report landslide events instead of individual landslides, i.e. a group of landslides with a common trigger and reported by media, citizens and/or government officials as a single unit. The latter results in significant cataloging and reporting biases. To counteract this biases, this study aims to identify clusters of landslide events that were triggered by the same rainfall event. An algorithm is developed that finds a series of landslide events that a) is continuous with no more than two days between individual events, and b) precipitation at the location of an individual event correlates with precipitation of at least one other event. The developed algorithm is applied to the Global Landslide Catalog (GLC) maintained by NASA. The results show that more than 40 % of all landslide events are connected to at least one other event, and that 14 % of all studied landslide events are actually part of a landslide cluster consisting of at least 10 events and up to 108 events in one day. Duration of the detected clusters also varies greatly from 1 to 24 days. Our study intends to enhance our understanding of landslide clustering and thus will assist in the development of improved, internationally streamlined mitigation strategies for rainfall related landslide clusters.”

Comment II:

Concerning section 2.3 "Method", I would suggest going more deeply into the explanation of the algorithm. A flowchart can be useful to fully describe the processes behind it.

Reply: We agree. Hence, an additional and more detailed flowchart was created. As Fig. 2 already provides a simplified schematic drawing, we decided to place the flowchart in the supplementary material.

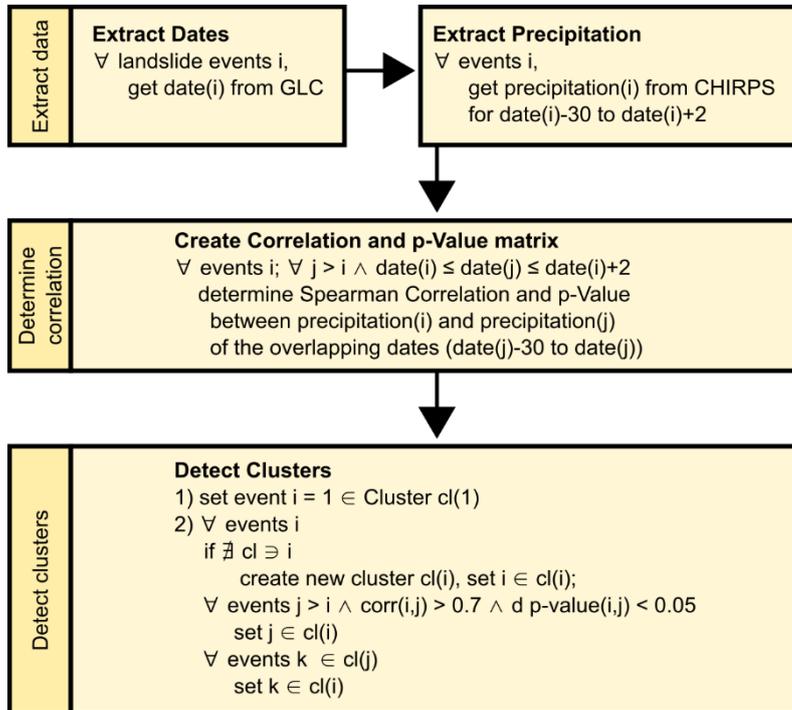


Figure S1. Flowchart of the algorithm to detect clusters within the global landslide catalog (GLC). Symbols included: \forall - for all; \wedge - logical conjunction; \in - element of; \nexists - there does not exist; \ni - contains.

Comment III:

More-over, the description of the two conditions for gathering landslide events into the same rainfall events should be described more in detail, in particular the condition (II). The choice of those values, for the spearman correlation and the p-value, should be fully described, also commenting about the limitations connected to those choices (as done for condition (I)).

Reply: We agree. Hence, following additional feedback from reviewer I, the sensitivity analysis was severely extended and the choices set for condition (II) are discussed in more detail. Additionally, a more inclusive discussion of limitations is now included:

“The threshold value of the spearman correlation coefficient was determined by testing the robustness of the identified clusters for different threshold values between zero and one (Fig. S2). Our results indicate that mean duration, area, and number of landslides per cluster are comparably robust to changes of the spearman correlation coefficient. In contrast maximum duration, area and number of landslides per cluster change drastically for different threshold values. From a correlation coefficient

threshold of 0.35 to 0.7, maximum number of landslide events per cluster decreases from close to 500 to slightly above 100, maximum duration decreases from more than 80 days to approximately 25, and area decreases from 60,000,000 km² (approximately 1/3 of the planet's surface area) to 200,000 km². For threshold values greater 0.7, only minor changes are observed. Hence, the latter was set as the correlation threshold value for this study (Fig. S2).

Additionally, we tested the robustness of the method to the time period of precipitation for which the correlation coefficient was determined (Fig. S3). It appears that the number of days is much less influential than the set correlation coefficient threshold (Fig. S2). Again, maximum number of landslides, area, and duration are impacted most, however remain stable for time period longer than 30 days prior to the second event.

It is important to note that the introduced method does not limit the spatial extent of the found landslide clusters. While this ensures that previously undetected, large-scale connections between individual landslide events are found, it is also susceptible to link landslides occurring in different parts of the world, where rainfall coincidentally correlates. Hence, when applying the method to another dataset, the robustness of the threshold values for correlation coefficient and time analyzed needs to be rechecked.“

Comment IV:

Then, I would also suggest moving in this section the method used to define the rainfall events, using section 2.1, and 2.2 only to describe the dataset available.

Reply: We agree. Section “2.3 Method” was renamed to “2.3. Detection of Landslide Clusters and a new section “2.4 Rainfall Analysis” was created. 2.4 now contains all of the information previously in 2.2 describing how precipitation before landslide events was analyzed in order to better understand its impact:

“2.4 Rainfall Analysis

In order to compare rainfall during a landslide event to overall rainfall at the location, the 95th percentile of precipitation excluding non-rainy days was determined for 10 years prior to the event. This comparison was also previously used by Kirschbaum et al (2015) to identify rainfall triggered landslide events. However, in their case, rainfall data from the Tropical Rainfall Measuring Mission (TRMM) was used for the time period 2000–2013 independent of the date of the landslide event. Due to its higher spatial resolution CHIRPS data was used here instead.

In addition to the 95th percentile of rainfall, the global rainfall threshold by Guzzetti et al. (2008) was also utilized to determine the likelihood of the individual landslide events being triggered by rainfall. In their study 2626 rainfall events that have resulted in shallow landslides and debris flows were analyzed in order to determine

the following global rainfall intensity–duration threshold [<http://rainfallthresholds.irpi.cnr.it>]:

$$I = 2.2 \cdot D^{-0.44} \quad (1)$$

Here the threshold intensity (I) was determined for each 24 hours starting with a duration (D) of 12 hours. This results in an average precipitation of 0.73 mm/h for $D = 12$ h, 0.45 mm/h for $D = 36$ h, and 0.35 mm/h for $D = 60$ h. The rainfall threshold was then compared to the cumulative mean precipitation of the rainfall event preceding each landslide event.”

Comment V:

Please consider using more tables in Section 3 to summarize and better describe the results obtained. Currently the text may result a bit confusing.

Reply: We agree. Thus, we added an additional row to Table 1 giving information on the global dataset and created Tables S2 and S3, providing information about trigger and size for each region. Because this information is also available in Figure 4, we decided to place these tables in the supplementary material.

Table 1. Regional statistics for all landslide clusters (LC) with at least ten landslide events (LE).

Region	# LC	# LE	LE per LC	Average duration of LCs	LEs per day	Average area of LCs [km ²]	Percentage of LE in a LC ≥ 10 LE
Global	50	1,209	24.2	7	3.5	35,441	13
West Coast, North America	29	829	28.6	9	3.3	52,970	31
Central and Eastern USA	8	107	13.4	6	2.4	23,357	12
South and Central America	5	168	33.6	3	11.2	1,320	18
Himalaya	4	48	12.0	5	2.3	5,476	3
South-East Asia	4	57	14.3	5	3.2	5,143	4

Table S2. Size of the landslides of clusters with at least 10 events for the different regions.

Region	Small	Medium	Large	Very large	Unknown
West Coast, North America	26%	29%	1%	-	43%
Central and Eastern USA	62%	38%	-	-	-
South and Central America	4%	88%	5%	3%	-
Himalaya	6%	88%	6%	-	-
South-East Asia	4%	88%	9%	-	-

Table S3. Apparent trigger of the landslides of clusters with at least 10 events for the different regions.

<i>Region</i>	<i>Downpour</i>	<i>Rain</i>	<i>Continuous rain</i>	<i>Snow melt</i>	<i>Flooding</i>	<i>Tropical cycloon</i>	<i>Con-struction</i>	<i>Unknown</i>
<i>West Coast, North America</i>	27%	19%	3%	2%	1%	-	-	47%
<i>Central and Eastern USA</i>	64%	15%	9%	3%	-	-	1%	6%
<i>South and Central America</i>	85%	-	8%	-	-	7%	-	-
<i>Himalaya</i>	40%	4%	48%	-	2%	-	-	6%
<i>South-East Asia</i>	16%	19%	-	-	-	65%	-	-

Comment VI:

Finally, I would suggest to clearly split the discussion from the results. Please consider creating Section 4. "Discussion. In this way the authors' comments are highlighted and easy to be understood for a reader.

Reply: We would like to keep our manuscript concise, thus we would like to keep our current structure.

Specific comments

Comment VII:

Avoid in the text: ">", "<", etc...

Reply: All Instances of ">" and "<" have been replaced with "more than"/"greater" and "less than" respectively.

Comment VIII:

Fig.1 difficult to distinguish among dots. I would suggest enlarging this figure.

Reply: We agree, the Figure was updated and now shows a heat map instead of individual dots.

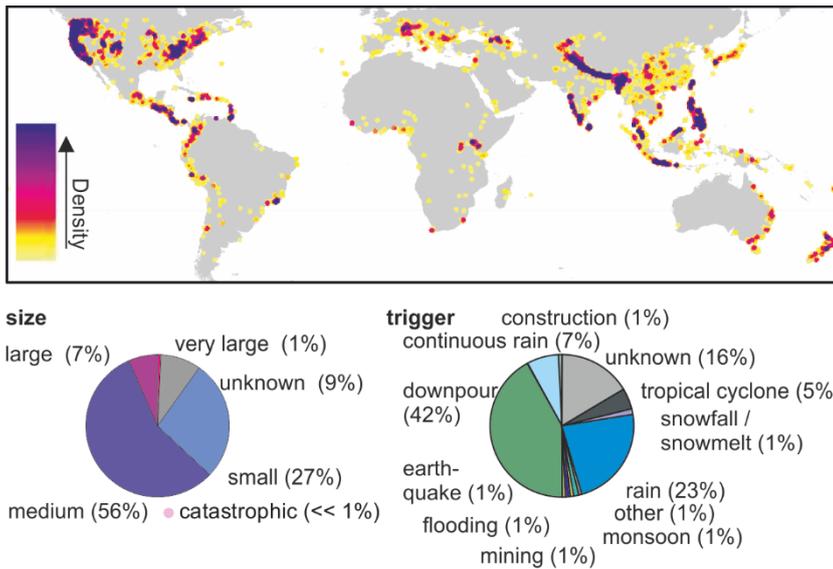


Figure 1. Heat map of all landslide events analyzed in this study and their size and apparent trigger. Overall a total of 9279 events were tested for clustering.

Comment IX:

Please check the numbers for tables and figures (i.e., lines 122, 124: Fig. S1 ??; line 169: Table S1??)

Reply: With Table S1 and Fig. S1 (and following Figures S2 – S10) we meant to refer to tables in the supplementary material. This information is now given when first mentioning Figure S1 and Table S1:

“A schematic drawing of this algorithm is provided in Fig. 2, and a more detailed flowchart in Fig. S1 in the supplementary material.”

“Table S1 in the supplementary material gives more detail of the 50 clusters with at least 10 events.”

1 Global detection of rainfall triggered landslide clusters

2 Susanne A. Benz^{1,2}, Philipp Blum¹

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5 *Correspondence to:* Philipp Blum (blum@kit.edu) and Susanne A. Benz (sabenz@ucsd.edu)

6 Abstract

7 An increasing awareness of the cost of landslides on the global economy and of the associated loss
8 of human life, has led to the development of various global landslide databases. However, these
9 databases typically report landslide events instead of individual landslides, i.e. a group of landslides
10 with a common trigger and reported by media, citizens and/or government officials as a single unit.
11 The latter results in significant cataloging and reporting biases. To counteract this biases, this study
12 aims to identify clusters of landslide events that were triggered by the same rainfall event. An
13 algorithm is developed that finds a series of landslide events that a) is continuous with no more
14 than two days between individual events, and b) precipitation at the location of an individual event
15 correlates with precipitation of at least one other event. ~~Here, t~~The developed algorithm is applied
16 to the Global Landslide Catalog (GLC) maintained by NASA. The results show that more than 40 %
17 of all landslide events are connected to at least one other event, and that 14 % of all studied
18 landslide events are actually part of a landslide cluster consisting of at least 10 events and up to
19 108 events in one day. Duration of the detected clusters also varies greatly from 1 to 24
20 days. ~~However, in a more regional analysis this number ranges from 30 % for the West Coast of~~
21 ~~North America to 3 % in the Himalaya Region. The cluster with most landslide events in a day is~~
22 ~~located in Rio de Janeiro, Brazil, with 108 events on 6th April 2010. In contrast, the longest running~~

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23 ~~cluster was observed on the West Coast of North America with 132 events occurring in an area of~~
24 ~~over 120,000 km² during 24 days in December 2015.~~ Our study intends to enhance our
25 understanding of landslide clustering and thus will assist in the development of improved,
26 internationally streamlined mitigation strategies for rainfall related landslide clusters.

27 **Keywords:** Landslide events; Database; Extreme weather; Rainfall induced; Early warning
28 systems;

29 1. Introduction

30 The fatal and catastrophic nature of landslides has led to the development and maintenance of
31 various global databases, such as the NASA Global Landslide Catalogue (GLC; e.g. Kirschbaum
32 et al. 2015) and recently the Global Fatal Landslide Database (GFLD) by Froude & Petley (2018).
33 Typically, these databases have a distinct focus. For example, the Global Landslide Catalogue
34 (GLC) operated by NASA focuses on rainfall triggered landslides (Kirschbaum et al., 2010, 2015),
35 whereas the Global Fatal Landslide Database records fatal landslides (Froude and Petley, 2018;
36 Petley, 2012). Through these databases we are able to provide first estimates on the number of
37 recorded fatalities, which were ~~>~~more than 55,000 between 2004 and 2016 (Froude and Petley,
38 2018) and map near real-time risk for landslides almost on a global scale (Kirschbaum and Stanley,
39 2018). Still, while they play a key role in understanding the effects of landslides on our society, it
40 is important to note that they are primarily based on news and government reports. These databases
41 therefore do not count landslides, but landslide events, which contain either a single or a multitude
42 of landslides within an area that are assumed to be triggered by the same event (Malamud et al.,
43 2004). The exact number of slope failures in each event is often unknown and depends on the
44 quality of the reporting. For some databases this number is included in a parameter of intensity or
45 size of each event. Typically, for large databases however, this is merely qualitative and describes

46 not only the number of individual landslides, but also an impact such as economic or human losses.
47 This classification is commonly based on press releases and is therefore heavily biased on the news
48 outlet reporting each event (e.g. Carrara et al., 2003).

49 Landslides triggered by catastrophic events, such as earthquakes or major storms, are often counted
50 as one event containing thousands of individual landslides (Kirschbaum et al., 2015). In contrast,
51 landslides caused by non-catastrophic events such as reasonable rainfall, are commonly counted as
52 individual events, disregarding their shared trigger. Consequently, the overall extent of clustering
53 in landslides is often unknown. But only if we better understand the extent of clustering between
54 individual landslide events, will we be able to understand the patterns they occur in and have the
55 chance to utilize these patterns to improve our forecast models (e.g. Martelloni et al., 2012).

56 Until now, few studies have focused on rainfall triggered landslide clusters and rather on temporal
57 clusters over a long time period within a confined region (e.g. Samia et al., 2017; Witt et al., 2010).
58 Biasutti et al. (2016) investigated the spatiotemporal clustering due to rainfall events for three
59 selected urban areas of the US West Coast: Seattle, San Francisco and Los Angeles. Over the nine
60 year study period, they found approximately 20 days within each city with multiple (up to eight)
61 landslide events. Additionally, they could identify close to 40 landslide events that were followed
62 by another event within the next week. However, with a focus on only selected study areas, they
63 did not show the overall extend of these clusters.

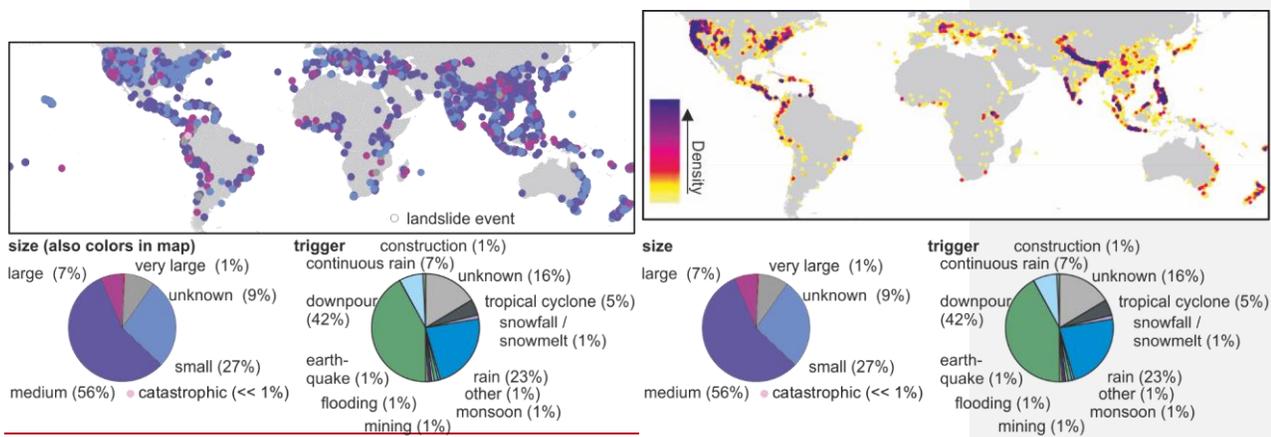
64 The objective of this study is therefore to develop an algorithm, which is able to identify such
65 clusters on a global scale. By applying the algorithm to the Global Landslide Catalog (GLC) the
66 overall degree of clustering in the database is shown, and spatial patterns of clusters with at least
67 10 landslide events are described. Additionally, landslide events and rainfall patterns of the most
68 intense and longest clusters are comprehensively discussed. In contrast to previous studies, such as

69 by Biasutti et al. (2016), clusters here are not constricted by a maximum spatial extent, instead they
 70 are grouped by analyzing and comparing rainfall prior to the event at the event locations.

71 **2. Material and Method**

72 **2.1 Landslide Data**

73 All landslide events within this study are part of the Global Landslide Catalog (GLC) operated by
 74 NASA and introduced in Kirschbaum et al. (2010, 2015). Data within the catalogue is based on
 75 online news articles that are found through search engine options such as Google Alerts. In the
 76 presented study, only events with a location accuracy ≤ 25 km are considered. As the rainfall data
 77 used is only available within $\pm 50^\circ$ Latitude, landslide events outside of this range are not
 78 considered. Overall, a total of 9279 landslide events, ranging from 1988 to 2018 are analyzed (Fig.
 79 1). However, only 45 of these events occurred before 2007, when the GLC was established.



80
 81 **Figure 1.** HeatM map of all landslide events analyzed in this study and their size (also color in map) and apparent
 82 trigger. Overall a total of 9279 events were tested for clustering.

83 For each event, the GLC provides a landslide type, e.g. land- or mudslide, and a landslide trigger,
 84 e.g. rainfall, downpour, earthquakes or construction work. Detailed descriptions on these

85 classifications can be found in Kirschbaum et al. (2010, 2015). Furthermore, within the GLC the
86 intensity, impact, and number of landslides per event is expressed in a variable called “size”. While
87 events classified as small in the database are only a single landslide, medium or larger landslide
88 events may consist of multiple landslides within an unspecified range. About 64 % of the studied
89 events are classified as medium or larger in size. However, a precise count of the number of
90 landslides contained within these events does not exist in this database nor in any other of the global
91 scale databases currently available. Within the GLC most of the small events that contain only a
92 single landslide, are located within the United States (Fig. 1).

93 2.2 Rainfall Data

94 For the rainfall analysis, the Climate Hazards Group InfraRed Precipitation with Station data
95 (CHIRPS) (Climate Hazards Group, 2015) is used, which has a resolution of $0.05^\circ \times 0.05^\circ$ and
96 daily time steps. For each landslide event location, precipitation data were downloaded for 10 years
97 ~~30 days~~ preceding the event and up to two days after the event using Google Earth Engine (Gorelick
98 et al., 2017).

99 ~~In order to compare rainfall during the event to overall rainfall at the location, the 95th percentile~~
100 ~~of precipitation excluding non-rainy days was determined for 10 years prior to the event. This~~
101 ~~comparison was also previously used by Kirschbaum et al (2015) to identify rainfall-triggered~~
102 ~~landslide events. However, in their case, rainfall data from the Tropical Rainfall Measuring Mission~~
103 ~~(TRMM) was used for the time period 2000–2013 independent of the date of the landslide event.~~
104 ~~Due to the higher spatial resolution CHIRPS data was used here instead.~~

105 ~~Here the threshold intensity (I) was determined for each 24 hours starting with a duration (D) of~~
106 ~~12 hours. This results in an average precipitation of 0.73 mm/h for $D = 12$ h, 0.45 mm/h for $D =$~~

~~26 and 0.25 mm for D. On their full threshold, we compared the cumulative precipitation of their full preceding block level~~

108 The main objective of this study is to identify clusters of landslide events that occurred during, and
109 are likely triggered by the same rainfall event. To determine if two events, A and B, occurred during
110 the same rainfall event, two conditions have to be fulfilled: (I) A and B occurred within three days
111 of each other, and (II) spearman correlation between daily precipitation at A and at B is >greater
112 0.7 and has a p-value <less than 0.05 for the 30 days preceding the later of the two events. Other
113 landslide events that fulfill these conditions with either A or B, are considered to be part of the
114 cluster. A schematic drawing of this algorithm is provided given in Fig. 2, and a more detailed
115 flowchart in Fig. S1 in the supplementary material. The threshold value of three days maximum
116 between two events was used following Biasutti et al. (2016), who found it unlikely that landslide
117 events occurring more than three days apart, occurred during the same rainfall event. However, it
118 is important to note that their study was set in three metropolitan areas on the West Coast of the
119 USA and might not be applicable everywhere.

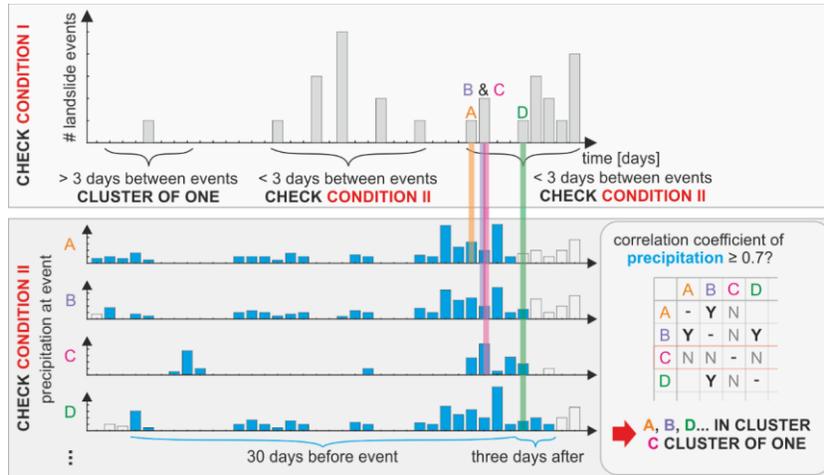
120 The threshold value of the spearman correlation coefficient was determined by testing the
121 robustness of the identified clusters for different threshold values between zero and one (Fig. S2+).
122 Our R results indicate that mean duration, area and number of landslide per cluster are comparably
123 robust to changes of the spearman correlation coefficient. In contrast maximum duration, area, and
124 number of landslides per cluster change drastically for different threshold values. From a
125 correlation coefficient threshold of 0.35 to 0.7, maximum number of landslide events per cluster
126 decreases from close to 500 to slightly above 100, maximum duration decreases from more than
127 80 days to approximately 25, and area decreases from 60,000,000 km² (approximately 1/3 of the
128 planet's surface area) to 200,000 km². For threshold values ~~It was set to be greater~~ 0.7, as from here
129 on numbers of landslides per cluster, duration of clusters, and area of clusters are stable for their

130 ~~mean and maximum values minor little changes are observed. Hence, Thus, the latter value was set as the correlation threshold value for this~~
131 ~~study (Fig. S+S2).~~

132 Additionally, we tested the robustness of the method to the time period of precipitation for which
133 the correlation coefficient was determined (Fig. S3). It appears that the number of days is much
134 less influential than the set correlation coefficient threshold (Fig. S2). Again, maximum number of
135 landslides, area, and duration are impacted most, however remain stable for time period longer than
136 30 days prior to the second event., however

137 It is important to note that the introduced method does not limit the spatial extent of the found
138 landslide clusters. While this ensures that previously undetected, large-scale connections between
139 individual landslide events are found, it is also susceptible to link landslides occurring in different
140 parts of the world, where rainfall coincidentally correlates. Hence, when applying the method to
141 another dataset, the robustness of the threshold values for correlation coefficient and time analyzed
142 needs to be rechecked.;

143 The introduced algorithm is independent of subsoil topography and relief parameters. While these
144 impact the precipitation intensity-duration threshold that is commonly expected to trigger
145 landslides, locations with different thresholds might still experience landslides triggered by the
146 same rainfall event.



147
148 **Figure 2.** Schematic drawing of the algorithm used to identify, if two landslide events within the Global Landslide
149 Catalog (GLC) occurred during the same rainfall event and hence belong to the same cluster. For condition II only
150 events occurring within three days of each other are compared.

151 2.4 Rainfall Analysis

152 In order to compare rainfall during the a landslide event to overall rainfall at the location, the 95th
153 percentile of precipitation excluding non-rainy days was determined for 10 years prior to the event.
154 This comparison was also previously used by Kirschbaum et al (2015) to identify rainfall triggered
155 landslide events. However, in their case, rainfall data from the Tropical Rainfall Measuring Mission
156 (TRMM) was used for the time period 2000–2013 independent of the date of the landslide event.
157 Due to ~~the~~ its higher spatial resolution CHIRPS data was used here instead.

158 In addition to the 95th percentile of rainfall, the global rainfall threshold by Guzzetti et al. (2008)
159 was also utilized to determine the likelihood of the individual landslide events being triggered by
160 rainfall. In their study 2626 rainfall events that have resulted in shallow landslides and debris flows

161 were analyzed in order to determine the following global rainfall intensity–duration threshold

162 [<http://rainfallthresholds.irpi.cnr.it>]:

$$I = 2.2 \cdot D^{-0.44} \quad (1)$$

163 Here the threshold intensity (I) was determined for each 24 hours starting with a duration (D) of
164 12 hours. This results in an average precipitation of 0.73 mm/h for $D = 12$ h, 0.45 mm/h for $D =$
165 36 h, and 0.35 mm/h for $D = 60$ h. The rainfall threshold was then compared to the cumulative
166 mean precipitation of the rainfall event preceding each landslide event.

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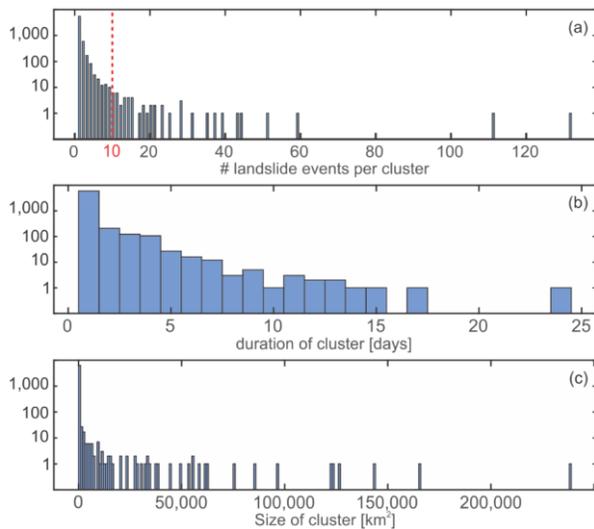
167 **3. Results and Discussion**

168 **3.1 Clustering Characteristics**

169 The presented algorithm divided the 9279 landslide events of the Global Landslide Catalog (GLC)
170 into 6474 clusters of events connected through precipitation. However, 85 % of these clusters
171 consist of only a single landslide event, containing in total 59 % of all recorded landslide events.
172 This implies that a large number of landslide events are in fact isolated events with no association
173 to other events. Nevertheless, 67 % of these ‘single landslide event’-clusters are categorized as
174 medium or larger and might contain more than one landslide (in comparison 58 % of the landslide
175 events in clusters \geq one landslide event are categorized as medium or larger). Hence, the number
176 of isolated landslides is likely to be significantly smaller than the number of isolated landslide
177 events.

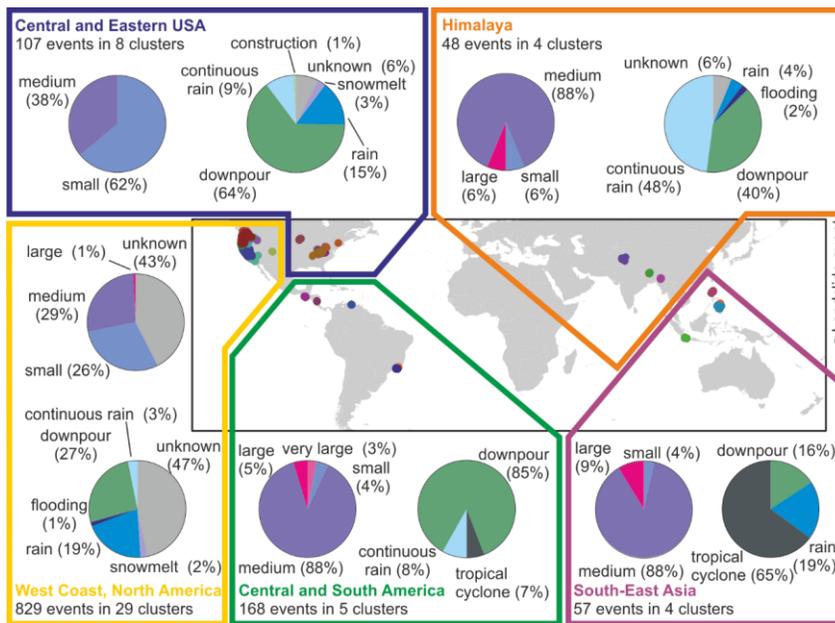
178 In the Global Landslide Catalog (GLC) only 3 % of the analyzed landslide events are linked to
179 triggers unrelated to rainfall such as construction, volcanos or earthquakes. This number is reduced
180 to 1.5 % for landslides in a cluster of more than one event. Due to the low number of events in this
181 category, future research is necessary to test and thoroughly validate these findings as well as to

182 assess possible reasons and implications of this phenomenon. For now, we assume that this is
183 mainly caused by biased reporting and cataloging of landslide events, where events linked to larger
184 disasters such as earthquakes, might be reported as one large landslide event, whereas landslides
185 linked to rainfall, might be individually reported. Similar observations were previously made by
186 Kirschbaum et al. (2015) for events in the GLC that are linked to major storms. An example of this
187 is the catastrophic magnitude 7.8 Gorkha earthquake in Nepal in 2015. While more than 25,000
188 landslides occurred during the earthquake and its aftershock sequence (e.g. Roback et al., 2018),
189 they are only reported as 13 landslide events in the excerpt from the GLC analyzed here. In it, they
190 are described as ranging in size from small to large and their trigger is given as “unknown”,
191 “earthquake” and in one case “snowmelt”. Our algorithm sorts these events into eight clusters of
192 up to three events.



193
194 **Figure 3.** Histogram of the number of events per cluster, duration of clusters and area of the convex hull of each cluster.
195 Clusters with only a single landslide event were appointed an area of zero. Within this study, all clusters with at least
196 10 landslide events were analyzed more closely.

197 Figure 3 provides histograms of the landslide events per cluster, duration of clusters and area
 198 covered by clusters (convex hull) in a logarithmic scale. As expected, for all three aspects frequency
 199 reduces drastically for larger numbers. In the following section all 50 clusters with at least 10 events
 200 (marked in red in Figure 3) are evaluated more closely.



201
 202 **Figure 4.** Location of all landslide events within clusters ≥ 10 events (different colors indicate different clusters).
 203 Overall, clusters in five distinct regions could be identified in the GLC (see Table S1 for more detail). Size and trigger
 204 (GLC categorization) of the associated landslide events are also shown (also see Tables S2 and S3).

205 **3.2 Clusters with more than ten Landslide Events**

206 **3.2.1 Global Analysis**

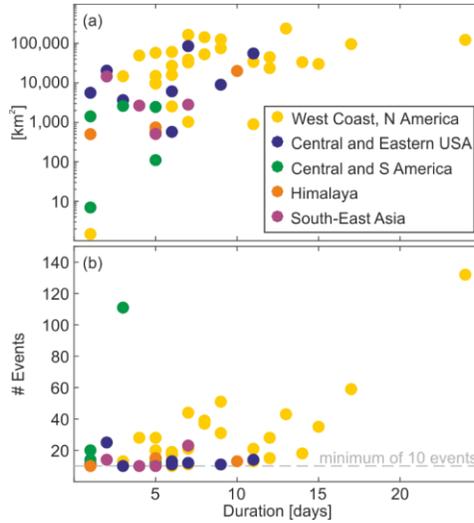
207 Table S1 [in the supplementary material](#) gives more detail of the 50 clusters with at least 10 events.
 208 In total 13 % of all landslide events are associated with one of these clusters ([Table 1](#)). As the
 209 database is most likely incomplete, the true number is expected to be higher. Overall the algorithm

210 detects clusters in five distinct regions: (1) West Coast of North America, (2) Central and Eastern
211 USA, (3) Central and Southern America, (4) Himalaya Region and (5) South-East Asia (Fig. 4).
212 However, close to three quarters of all clusters ≥ 10 events are found within the USA mostly due
213 to a bias in the GLC database (Kirschbaum et al., 2015) (Fig. 1). This is also shown in the size of
214 recorded landslide events (Fig. 4 and Table S2).

215 In North America events are often classified as small in size, while clusters in the other regions
216 contain mainly medium events. This might be due to English speaking media, on which the GLC
217 is based, only picking up on large international events that consist of multiple landslides within an
218 area and smaller ones are under or not reported at all.

219 The median clusters with at least 10 events last six days, consist of 15 events, and span over an
220 area of 15,000 km² (Fig. 5). As expected, there is a positive correlation between cluster duration
221 and area (spearman correlation coefficient of 0.70, p-value: 0.001). However, this cannot be
222 observed for cluster duration and number of landslide events within the cluster (spearman
223 correlation coefficient of 0.44, p-value: 0.001). When comparing the different regions, clusters
224 located on the West Coast of North America are on average the longest and cover the largest area.
225 In contrast, events in South America are shortest and smallest, nevertheless they have the highest
226 number of events and clusters per day (Table 1).

227 On a global scale, no significant trend over time can be observed and clusters with ≥ 10 events
228 occur around the year (Fig. S42). Similarly, the total number of reported landslide shows no
229 significant increase in the GLC (Kirschbaum et al., 2015) as well as in other global databases such
230 as the Global Fatal Landslide Database (Froude and Petley, 2018). More regional observations
231 show seasonal variation and are described more closely in the following chapters. However, for
232 three out of the five regions, there are only five clusters or even less.



233
 234 **Figure 5.** Link between the duration of the individual clusters ≥ 10 events and a) the covered area and b) the number
 235 of landslide events per cluster. The color of the scatter plots indicates the region, in which each cluster occurred.

236 **Table 1.** Regional statistics for all landslide clusters (LC) with at least ten landslide events (LE).

Region	# LC	# LE	LE per LC	Average duration of LCs	LEs per day	Average area of LCs [km ²]	Percentage of LE in a LC ≥ 10 LE
Global	50	1,209	24.2	7	3.5	35,441	13
West Coast, North America	29	829	28.6	9	3.3	52,970	31
Central and Eastern USA	8	107	13.4	6	2.4	23,357	12
South and Central America	5	168	33.6	3	11.2	1,320	18
Himalaya	4	48	12.0	5	2.3	5,476	3
South-East Asia	4	57	14.3	5	3.2	5,143	4

237
 238 **3.2.2 West Coast, North America**

239 Landslides in the west of North America have been intensively investigated, mainly in the form of
 240 case studies that discuss landslides along the Pacific coast in the states of California (Collins and
 241 Sitar, 2008; Wieczorek, 1988), Oregon (Benda, 1990; Miller and Burnett, 2008) and Washington

242 (LaHusen et al., 2016; Perkins et al., 2017). This region is also one of the few, where the clustering
243 of rainfall triggered landslide events was previously investigated, showing qualitatively that there
244 are many instances in which landslides occur on consecutive days (Biasutti et al., 2016).

245 About 31 % of all landslide events recorded in this area belong to a cluster of at least ten events.
246 This is the highest number compared to the other regions of the world (Table 1). However, this
247 effect might be amplified by the high number of reported landslides. The large number of events
248 and clusters is mainly due to geologic, topographic, climatic conditions and construction practices.
249 For example in Oregon, steep slopes and heavy rainfalls are as well as poor construction practices
250 result in high economic losses (Wang et al., 2002). Burns et al. (2017) estimated an average annual
251 loss of \$15.4 million due to landslides in Oregon alone. In years with heavy storms such as 1996,
252 this can accumulate to more than \$100 million (Wang et al., 2002).

253 The observed clusters in this area are among the longest and have the largest areas of all regions
254 (Table 1). While the size of landslide events (as given by the GLC) in the west of North America
255 are small compared to most other regions, there is also a considerable amount of events, where the
256 size is unknown (43 %, Fig. 4, [Table S2](#)). While about half of the landslide events within clusters
257 ≥ 10 events are classified as “trigger unknown” (47 %), landslide events with a known cause are
258 mainly triggered by downpour (27 %) or rain (19 %) (Fig. 4, [Table S3](#)). However, when looking at
259 satellite based rainfall data preceding the clusters, rainfall cannot always be identified as a trigger
260 (Fig. [S53](#)). While it generally exceeds the global rainfall threshold (Guzzetti et al., 2008), the 95th
261 percentile of precipitation on rainy days is not reached for the majority of the clusters. Although,
262 several studies linked landslides within California to earthquakes (e.g. Harp and Jibson, 1996;
263 Keefer, 2000), they occurred before 2007 and are not registered in the GLC.

264 While there appears to be no significant change in the number of clusters over time (Fig. [S4](#)

265 2), most clusters occur during the rainy season (November to March), when most landslide events
266 occur. Within the west of North America this time period is therefore often referred to as the
267 “landslide season” (e.g. Mirus et al., 2018). Only one cluster in this region appears in June (Cluster
268 ID 21, Table S1). However, the center of this cluster is located more inland (in San Miguel County,
269 Colorado) and is also the shortest cluster (only one day) within the region as well as the most local
270 of all clusters in this study, covering only 1 km². While this cluster is triggered by downpour
271 according to the GLC, this is not apparent from satellite derived precipitation (Fig. S53). The small
272 size of the cluster might be the reason, why low-resolution satellite derived precipitation does not
273 record any anomalies here.

274 3.2.3 Central and Eastern USA

275 While most of the clusters with ≥ 10 landslides events of this region, are located in the Appalachian
276 Plateau (Ohio, West Virginia and Kentucky), one cluster can be found in Minnesota (ID 34 in Table
277 S1 and Fig. S64). While it is considerably smaller (580 km² compared to ~~>~~more than 9,000 km²),
278 it is comparable to the Appalachians cluster in its number of landslide events and duration. The
279 Appalachian Plateau is well known for its landslides and the annual direct cost in Kentucky exceeds
280 \$10 million (Crawford and Bryson, 2017).

281 Like the landslide clusters observed in the west of North America, clusters here consist mainly of
282 small landslides, which is most likely linked to the news alerts on which the GLC is based.
283 Checking sources in the GLC, they are mainly reported within smaller, more local news outlets
284 compared to landslide events outside of the US. To our knowledge the individual events grouped
285 by our algorithm into clusters have never been linked before. Clusters in this region occur
286 predominantly in spring (February to June), when rainfall is highest, slightly later than events on
287 the West Coast (Fig. S42). According to GLC they are predominantly triggered by downpours

288 (64 %, Fig. 4, [Table S3](#)). However, extreme rainfall is not always visible in satellite derived
289 precipitation (Fig. [S64](#)). For most clusters, it is below the 95th percentile, but above the global
290 threshold. It is worth noting that one cluster located in West Virginia (Cluster ID 35) shows no
291 rainfall on the satellite before day three of the cluster. Following the GLC, early landslide events
292 within this cluster are linked to snowmelt.

293 **3.2.4 Central and South America**

294 In contrast to the clusters in North America, more than 95 % of landslide events within clusters of
295 this region are medium in size or larger and might consist of several landslides themselves (Fig.
296 4). Thus, the number of landslides per cluster and per day is likely to be significantly higher than
297 the number of events per cluster and per day. Still, clusters in this area are on average only two and
298 a half days long, covering an area of slightly over 1,500 km² and they are the smallest and shortest
299 of all regions (Fig. 5, Table 1). It is important to note that this region covers the largest area reaching
300 from Rio de Janeiro in Brazil to Guatemala in Central America. From the few clusters we could
301 identify, it appears that there are dissimilarities between the clusters in Central America and South
302 America. The two clusters in Nicaragua (ID 42) and Guatemala (ID 39) are triggered by continuous
303 rain and a tropical cyclone, respectively. In contrast, all events located in South America (IDs 38,
304 40, and 41) are all triggered by downpour (Table S1 and Fig. [S75](#)).

305 **3.2.5 Himalaya**

306 Like in South America, most landslide events (94 %) associated with clusters with ≥ 10 events in
307 the Himalaya region are categorized as medium and larger. Thus, the number of landslides per
308 cluster is again expected to be significantly higher than the number of landslide events per cluster.
309 However, there may be differences between regions. Event ID 44, located in India and Pakistan
310 around Jammu and Kashmir, is classified as medium to small, much longer (10 days) and covers

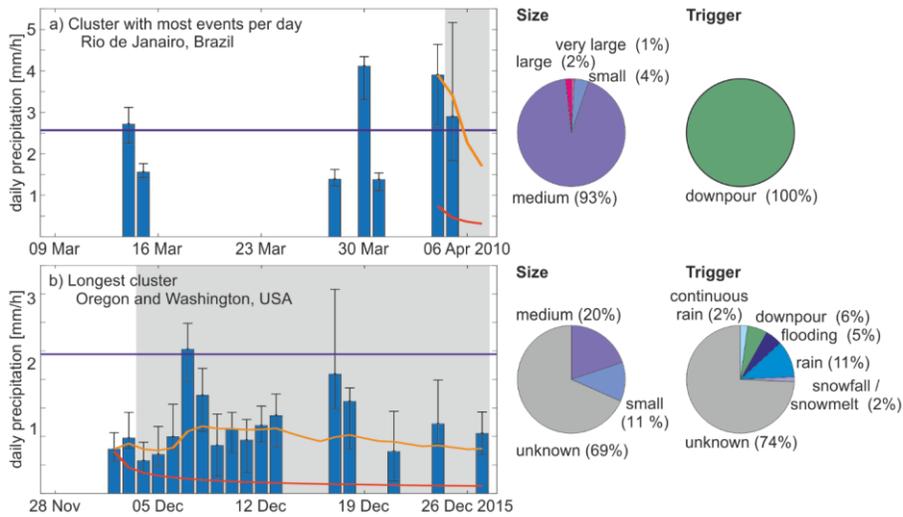
311 an area more than 10 times larger than the other clusters. All of them are classified as medium or
312 large and are located in the East of India with some events in Nepal (Table S1). In both regions,
313 clusters are triggered by continuous rain or downpour. For all clusters satellite based rainfall data
314 exceeds the global threshold, and in most cases the 95th percentile of rainfall on rainy days (Fig.
315 S86). It is important to note that while earthquake triggered landslides are common in the region
316 (e.g. Parkash, 2013; Roback et al., 2018), the presented algorithm is by design only able to pick up
317 clusters that are linked by rainfall.

318 3.2.6 South-East Asia

319 As only four clusters are identified in this region, a detailed analysis is impossible. Again, 96 % of
320 the events associated are categorized as medium or larger and the main triggers are tropical
321 cyclones (Cluster IDs 47 and 48), downpour (Cluster ID 49), and rain (ID 50) (Table S1). Here,
322 satellite based rainfall data before clusters is both above the global rainfall threshold and in most
323 cases above the 95th percentile (Fig. S79). While only one of the four clusters (ID 50) is recorded
324 outside of the Philippines (in Indonesia), there is no apparent difference between both countries
325 (Table 1).

326 3.3 Most Intense Cluster

327 The cluster with the most events in one day, i.e. most intense cluster, happened in Rio de Janeiro,
328 Brazil, as well as neighboring cities Niteroi and Sao Goncalo in 2010. In an area of approximately
329 2,800 km², 111 landslide events were recorded within only three days, however predominantly on
330 6th April 2010 (Table S1, ID 38). This is almost four times as many landslide events in a single day
331 than the second most intense clusters (IDs 1 and 3) located in Washington and Oregon, USA. Both
332 recorded 29 events in one day.

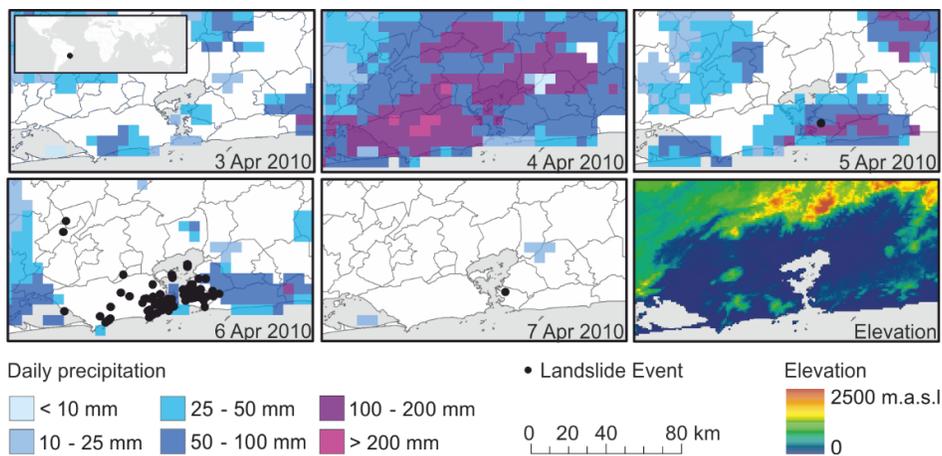


333
 334 **Figure 6.** Daily precipitation for 30 days preceding the last landslide event of the cluster with the size of the associated
 335 landslide events and their trigger according to the GLC. Shown is the median precipitation for all landslide locations
 336 with the inner quartiles as an error bar. The 95th percentile of daily rainfall (rainy days only) in the ten years preceding
 337 the event is given in blue, in red the global rainfall threshold ID (Guzzetti et al., 2008) and in orange the cumulative
 338 mean for the rainfall event preceding the cluster. a) Cluster with the most events per day (ID 43), and b) longest running
 339 cluster (ID 22).

340 Most of the 111 events associated with the cluster in Rio de Janeiro were recorded as medium in
 341 size, all of which were triggered by downpour (Fig. 6a). This is confirmed by satellite derived
 342 precipitation. Heavy rainfalls (Figs. 6a, 7) occurred on the 4th and 5th of April of up to 210 mm per
 343 day. In comparison, the 95th percentile in the 10 years preceding this cluster is on average only 62
 344 mm per day (rainfall for each individual location shown in Fig. S108). While the rainfall covered
 345 a large area, landslide events were primarily reported for steep slopes just outside the densely
 346 populated city center. Due to its location close to, and inside the urban area of Rio de Janeiro, the
 347 cluster caused approximately 200 fatalities according to CNN news reports
 348 (<http://www.cnn.com/2010/WORLD/americas/04/12/brazil.flooding.mudslides/>).

349 The location in the city might also be the reason for the large number of events being reported, as
350 we can expect more individual landslides being reported here compared to the countryside.

351 While studies not based on English speaking news alerts report a large number of landslides within
352 and around Rio de Janeiro (Calvello et al., 2015; Sandholz et al., 2018), only nine additional
353 landslide events inside the area of this cluster were reported in the GLC between 2009 and 2018.
354 Additionally, just northwest of the cluster another cluster occurred in January 2011 (ID 41 in Table
355 S1, Fig. S75). Although, this cluster only counts 20 individual landslide events within the GLC, it
356 is being reported as thousands individual landslides (Coelho Netto et al., 2013).

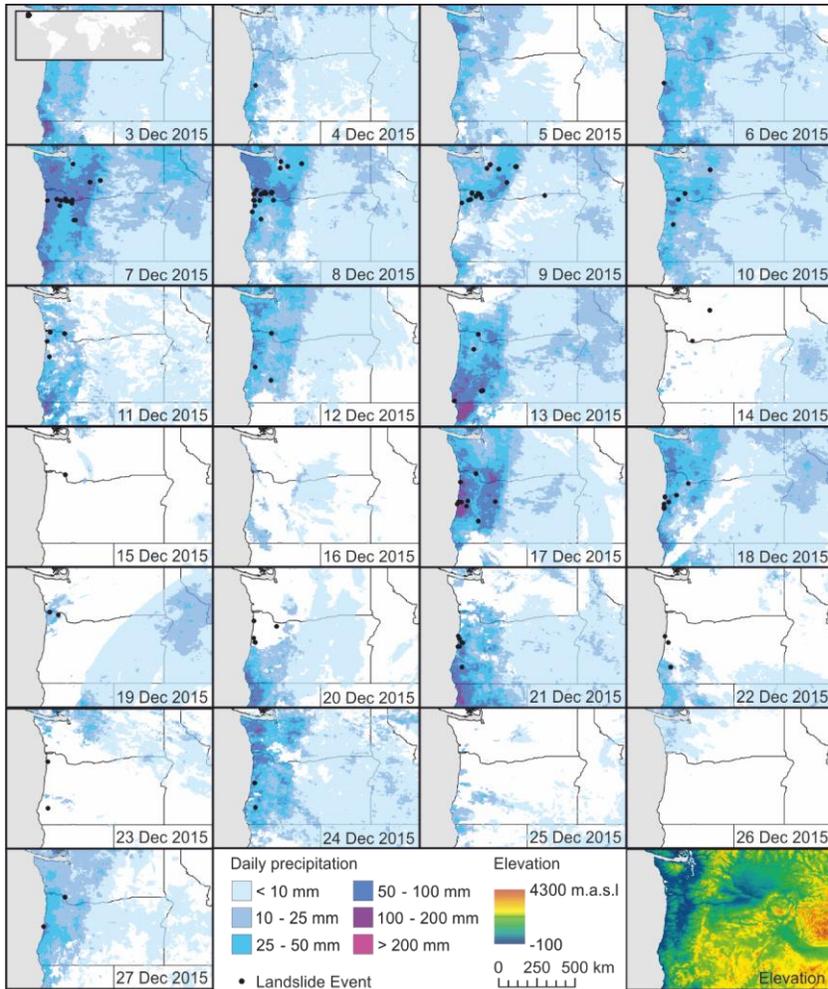


357
358 **Figure 7.** Location of the events in the cluster with the most events per day located in Rio de Janeiro, Brazil. Also
359 shown are daily precipitation and elevation. Elevation data is taken from the US Geological Survey (GTOPO30).

360 3.4 Longest Cluster

361 The longest running cluster identified in this study occurred in Oregon and Washington, USA from
362 4th to 27th December 2015 for a total of 24 days with 132 landslide events (Cluster ID 18, Table
363 S1). The second longest cluster lasted 17 days over January and February in 2012 and was also

364 located in Oregon and Washington, USA (Cluster ID 7). Overall, most events within the longest
365 cluster are unknown in size (69 %) and trigger (74 %) (Fig. 6b). However, inspecting satellite based
366 rainfall data, continuous rainfall appears to be the main trigger (Fig. 6b, Fig. 8 and Fig. [S9-S11](#) for
367 rainfall at the individual event locations). While daily rainfall is mainly below the 95th percentile,
368 cumulative mean rainfall is continuously above the global rainfall threshold. Although, heavy
369 rainfall is common in this area during winter times, for this cluster it lasted longer than usual and
370 was followed by shorter rain events in short successions (Fig. 8). Thus, the series of landslides did
371 not halt resulting in the longest cluster in the GLC. Following the information on sources within
372 the GLC, it appears that local media reported about the individual landslide events, but did not
373 detect on the extreme length of the continuous series of landslide events at this point in time (e.g.
374 <https://kval.com/news/local/landslide-blocks-i-5-in-sw-washington>;
375 [https://q13fox.com/2015/12/09/landslide-above-puget-sound-damages-several-homes-at-least-](https://q13fox.com/2015/12/09/landslide-above-puget-sound-damages-several-homes-at-least-one-vehicle/)
376 [one-vehicle/](#)). As landslide events are such a common occurrence in this region, and due to the
377 large area covered by this cluster, there is currently little to no emphasis on the longevity of this
378 specific series of landslide events in media and scientific studies.



379
 380 **Figure 8.** Location and time series of the longest cluster, located mainly in Oregon, USA. Also shown are daily rainfall
 381 and elevation. Elevation data is available from the US Geological Survey (GTOPO30).

382 **4. Conclusion**

383 In this study an algorithm is presented that detects clusters of landslide events that occur during,
 384 and are likely triggered by the same rainfall events. Here this algorithm is applied to the Global

385 Landslide Catalog (GLC), where it detects that more than 40 % of all recorded events can be linked
386 to at least one other event. The global analysis shows that 14 % of all landslide events are part of a
387 cluster ≥ 10 events. However, this percentage varies dramatically by the region, ranging from 30 %
388 on the West Coast of North America to 3 % in the Himalayas. Part of this is caused by sampling
389 and reporting bias. As the GLC is based on English speaking media, events in the USA are reported
390 and cataloged in much greater detail than events abroad. Nevertheless, within the GLC we could
391 detect clusters ≥ 10 landslide events in five distinct regions: (1) West Coast of North America, (2)
392 Central and Eastern USA, (3) Central and Southern America, (4) Himalaya Region, and (5) South-
393 East Asia. In South America, the studied clusters are the shortest, but contain the most events per
394 day. However, this is mainly due to a cluster in Rio de Janeiro, where 108 of events were recorded
395 on 6th April 2010. As most of these events are classified as medium or larger, the absolute number
396 of landslides is expected to be significantly higher. In contrast, the longest and largest clusters are
397 observed on the West Coast of North America. On average clusters here last nine days and cover
398 an area of more than 50,000 km². The steep slopes and continuous rainfalls present in the area
399 combined with the above average reporting of landslide events, makes a more detailed analysis of
400 rainfall related landslide clusters possible. The longest of all detected clusters ≥ 10 landslide events
401 is also located in this region: In December 2015, 132 landslide events were recorded over a time
402 period of 24 days spanning more than 120 thousand km², which were all triggered by the same
403 rainfall event. Detection of large scale clusters such as this one can not only help to improve our
404 understanding of the link between individual events, but also be used in our mitigation strategies.
405 Only once we improve our understanding of the relation between individual landslide events, we
406 will be able to predict their behavior and forecast their economic losses and fatalities. While our
407 study does not replace case specific and small scale studies, as well as the identification of threshold
408 values, it can provide an improved understanding for managing landslide mitigations on a larger

409 scale. Within the area covered by individual clusters the same mitigation strategies, including early
410 warning systems (EWS) based on weather forecast simulations, can be developed and validated.
411 For future research we recommend to use the presented algorithm not only for the correlation with
412 precipitation data, but also to include the geometry of atmospheric rivers during cluster detection.
413 Finally, the algorithm could be applied to more regional and other global landslide databases
414 thereby improving our understanding on the spatial and temporal occurrence of landslide clusters.

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