



1 **Water-level attenuation in broad-scale assessments of exposure to coastal flooding: a
2 sensitivity analysis**

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15

16 **Abstract**

17 This study explores the uncertainty introduced in global assessments of coastal flood exposure and risk
18 when not accounting for water level attenuation due to land-surface characteristics. We implement a
19 range of plausible water-level attenuation values for characteristic land-cover classes in the flood
20 module of the Dynamic and Integrated Vulnerability Assessment (DIVA) modelling framework and
21 assess the sensitivity of flood exposure and flood risk indicators to differences in attenuation rates.
22 Results show a reduction of up to 47% in area exposure and even larger reductions in population
23 exposure and expected flood damages when considering water level attenuation. The reductions vary
24 by country, reflecting the differences in the physical characteristics of the floodplain as well as in the
25 spatial distribution of people and assets in coastal regions. We find that uncertainties related to not
26 accounting for water attenuation in global assessments of flood risk are of similar magnitude to the
27 uncertainties related to the amount of SLR expected over the 21st century. Despite using simplified
28 assumptions to account for the process of water level attenuation, which depends on numerous
29 factors and their complex interactions, our results strongly suggest that an improved understanding
30 and representation of the temporal and spatial variation of water levels across floodplains is essential
31 for future impact modelling.

32

33 **1. Introduction**

34

35 Increased flooding due to sea-level rise (SLR) is a major natural hazard that coastal regions will face in
36 the 21st century, with potentially high socio-economic impacts (Kron, 2013; Wong et al., 2014). Broad-
37 scale (i.e. continental to global) assessments of coastal flood exposure and risk are therefore required
38 to inform mitigation targets and adaptation decisions (Ward et al., 2013a), related financial needs and
39 loss and damage estimates. Towards these ends, a number of recent studies have assessed the



40 exposure of area, population and assets to coastal flooding at national to global scales (Nicholls, 2004;
41 Brown et al. 2013; Jongman et al., 2012a; Ward et al., 2013b; Arkema et al., 2013; Muis et al., 2017) as
42 well as flood risk (Hinkel et al. 2014; Vousdoukas et al., 2018a).
43 Although methods for broad-scale coastal-flood exposure and risk assessment vary between studies,
44 flood extent and water depth have commonly been assessed based on spatial analysis, assuming that
45 all areas with an elevation below a certain water level that are hydrologically connected to the sea are
46 flooded (the “bathtub” method) (Poulter and Halpin, 2008; Licher et al., 2010). Notable exceptions
47 are the studies of Dasgupta et al. (2009), who used a simple approach to account for wave height
48 attenuation with distance from the coast, and Vousdoukas et al. (2018b) who, for the Iberian
49 peninsula, adopted a modified version of the bathtub approach that also considers water volume. The
50 use of simplified methods for assessing flooding is primarily related to difficulties of using
51 hydrodynamic methods at broad scale, namely the limited availability and large volume of the
52 necessary high-resolution input data; and the prohibitive computational costs, which render
53 hydrodynamic modelling applications impractical at global scales (Ramirez et al., 2016). Therefore,
54 global applications have utilised elevation data with a spatial resolution of 1 km and a vertical
55 resolution of 1m (Mondal and Tatem, 2012; Jongman et al., 2012b; Ward et al., 2014), with only a few
56 recent studies employing higher spatial resolution (90m) datasets (e.g. Hinkel et al., 2014; Vousdoukas
57 et al., 2018a; see also de Moel et al., 2015).
58 Hydrodynamic models are normally used only for local-scale applications. This is because they require
59 detailed data on parameters such as coastal topography/bathymetry and land use in order to
60 represent local-scale processes and to account for hydraulic properties. A range of simpler inundation
61 models that partly account for hydraulic processes at intermediate scales using medium resolution
62 elevation data (<100m²) have also been applied at subnational scales (e.g., Bates et al., 2010; Wadey
63 et al., 2012; Lewis et al. 2015; Ramirez et al., 2016), and these models are beginning to inform analysis
64 at broader scales (e.g. Vousdoukas et al., 2016; 2018a). There is also a developing literature on
65 hydrodynamic modelling of water level attenuation over coastal wetlands at the landscape scale (<1
66 km) for saltmarshes (Loder et al., 2009; Wamsley et al., 2009, 2010; Barbier et al., 2013; Smith et al.,
67 2016) and mangrove forests (McIvor et al., 2012; Zhang et al., 2012; Liu et al., 2013). However, the
68 incorporation of the above processes in global models is still very limited.
69 Not accounting for hydrodynamic processes in global models can, however, lead to overestimation of
70 flood extent and water depth. Hydrodynamic models capture processes that are not included in global
71 models, e.g. the effects of surface roughness (both natural and anthropogenic) and channel network
72 density and connectivity (and its effect on landscape continuity) on the timing, duration and routing of
73 floodwaters. For example, inundation extent has been shown in some cases to significantly decrease



74 in urban and residential areas when the built environment is represented in numerical simulations (e.g.
75 tsunami inundation: Kaiser et al., 2011; storm surge inundation: Brown et al., 2007; Orton et al., 2015).
76 To our knowledge, there is no study that has explored the uncertainty introduced into global models
77 by not accounting for water level attenuation due to hydrodynamic processes related to surface
78 roughness. This paper aims to address this gap. We derive a range of plausible water-level attenuation
79 values from existing literature and implement them in the flood module of the Dynamic Interactive
80 Vulnerability Assessment (DIVA) modelling framework (Hinkel et al., 2014). Next, we assess the
81 sensitivity of flood exposure and flood risk indicators to plausible changes in water-level attenuation
82 values under a range of different SLR scenarios. Finally, we compare the uncertainty due to water level
83 attenuation rates with the uncertainty range associated with expected SLR during the 21st Century.

84

85 **2. Methods and Data**

86

87 *2.1 The Dynamic Interactive Vulnerability Assessment (DIVA) modelling framework*

88 DIVA is an integrated, global modelling framework for assessing the biophysical and socio-economic
89 consequences of SLR, and associated extreme water levels, under different physical and socio-
90 economic scenarios and considering various adaptation strategies (Hinkel and Klein, 2009). DIVA has
91 been widely used for global and continental scale assessments of SLR impacts, vulnerability and
92 adaptation (e.g., McLeod et al., 2010; Hinkel et al. 2010; Brown et al. 2013; Hinkel et al., 2013; Hinkel
93 et al., 2014; Spencer et al., 2016; Schuerch et al., 2018). It is underpinned by a global coastal database
94 which divides the world's coastline (excluding Antarctica) into 12,148 coastal segments (Vafeidis et al.,
95 2008). Each segment contains approximately 100 elements of data concerning the physical, ecological
96 and socio-economic characteristics of the coast. Here we focus on the impacts of increased exposure
97 to coastal flooding and potential damages of extreme sea level events (due to the combination of
98 storm surges and astronomical high tides). We used the flood module of DIVA (for details see Hinkel
99 et al., 2014) to estimate potential coastal flood damage, SLR impacts and associated costs.

100 We specifically considered the following five indicators, which progressively include additional
101 components of flood risk:

- 102 1. Area below the 1-in-100 year flood event (km²), an estimate based on elevation data and
103 information on water levels for a single hazard event (i.e. the height of the 1-in-100 year sea
104 flood);



105 2. People living in the 1-in-100 year floodplain, a calculation based on spatial data on elevation
106 and population as well as on information for a single hazard event (i.e. the height of the 1-in-
107 100 year sea flood);
108 3. Assets in the 1-in-100 year floodplain (US \$), a calculation that uses data on elevation,
109 population, Gross Domestic Product (GDP) and information for a single hazard event (i.e. the
110 height of the 1-in-100 year sea flood);
111 4. Expected value of the number of people flooded per year (hereafter, people flooded), a
112 calculation based on elevation and population data and the probability distribution of the
113 hazard (i.e. sea flood heights and their probability of occurrence); and
114 5. Expected value of annual damages to assets (hereafter, flood damage) (US \$), a calculation
115 based on elevation, population and GDP data and the probability distribution of the hazard
116 (i.e. sea flood heights and their probability of occurrence).
117 For each coastline segment, a cumulative exposure function for area and population that gives the
118 areal extent (hydrologically connected to the sea) and number of people below a given elevation was
119 constructed. Damages to assets were assessed using a depth-damage function with a declining slope,
120 with 50% of the assets being destroyed at a water depth of one metre (Messner et al., 2007).
121
122 *2.2 Coastal Elevation and Rate of Water level Attenuation*
123 To simulate the effect of different values of attenuation at the broad scale, we implemented a stylised
124 elevation profile to represent the process of water level attenuation. We assumed that water levels
125 decrease at a constant slope (α) with increasing distance from the coastline. Location-specific coastal
126 profiles for every coastline segment were based on floodplain areas contained within the DIVA
127 database. The database reports total land area within different elevation increments (<1.5m, 1.5-2.5m,
128 2.5-3.5m, 3.5-4.5m, 4.5-5.5m, 5.5-8.5m, 8.5-12.5m, 12.5-16.5m) for each coastal segment. The
129 elevation dataset that was used for estimating floodplain areas and developing the segment elevation
130 profiles is the Shuttle Radar Terrain Mission (SRTM) Digital Elevation Database (Jarvis et al., 2008)
131 which has a vertical resolution of 1m and a spatial resolution of 3 arc seconds (~90m at the equator).
132 We approximated the average coastal profile for every segment by assuming that elevation
133 continuously increases with distance from the shore. Starting with the lowest elevation increment, the
134 floodplain areas of all elevation increments were cumulatively summed to retrieve the total area below
135 a certain elevation. The total areas were then divided by the segment length to derive the inundation
136 length of the respective floodplain (dx_i). To evaluate the representativeness of the assumption of
137 continuously increasing elevation with increasing distance from the shore, we used the original SRTM
138 dataset and calculated the Euclidian distance of each cell to the nearest coastline for every pixel. Mean
139 distances from the coast were calculated for each of the floodplain areas of each segment.
140 Subsequently, we compared these mean distances with the respective average floodplain elevation



141 for each DIVA coastline segment to analyse the validity of the “continuous-increase” assumption. This
142 comparison revealed that 55% of the DIVA coastline segments show either a continuous increase or
143 no change in the mean distance along the elevation profile (Figure 1a), suggesting that elevation does
144 not decrease with distance from the coast. Comparing all elevation increments of all segments (i.e.
145 pairwise comparison of the mean distances of consecutive elevation increments in a segment), there
146 was an increase, or no change, in the mean distance from the coastline in 88% of cases. Only 12% of
147 cases showed a decrease (Figure 1b). This result indicates that the stylised continuous profile (Figure
148 1a) can be regarded as representative of global coastal topography (see also Schuerch et al., 2018).

149

150

151

(a)

152



(b)



153

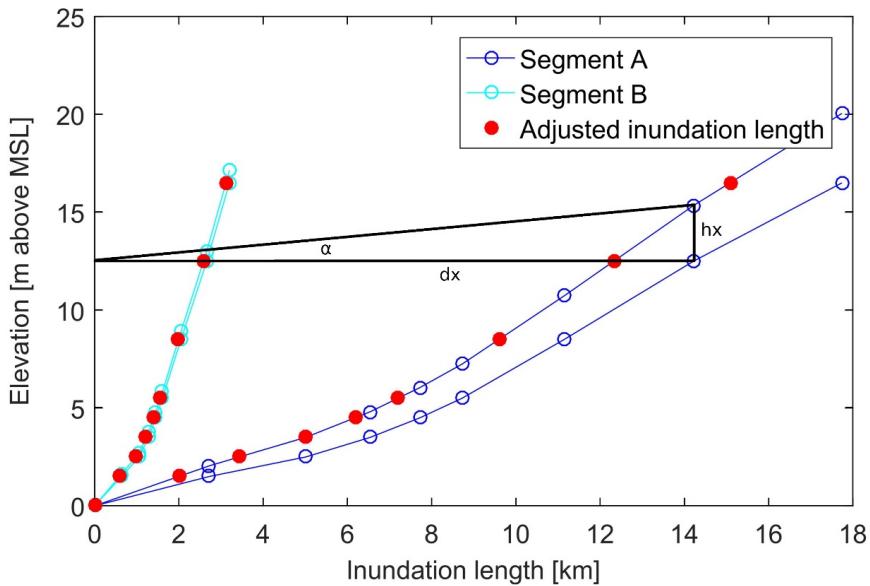
154 **Figure 1:** Stylised coastal profile with (a) continuous and (b) discontinuous increase in elevation with distance
155 from the shore.

156 We then adjusted the coastal profile using a range of possible attenuation rates that represent
157 different water surface slopes. Depending on the applied value for water level attenuation, the slope
158 (α) of the inundating water surface was employed to modify (incline) the coastal profile. Based on this
159 slope, the coastal profile is thereby elevated by the amount of the water level reduction (hx_i) computed
160 at a distance dx_i (Fig. 1):

161 $hx_i = \tan(\alpha) * dx_i$ (equation 1)

162 In this way the original floodplain areas and inundation depths are reduced in order to account for the
163 reduced (i) inundation length (dx) and (ii) inundation depth (hx) (see Fig. 2).

164



165

166 **Figure 2:** The stylised coastal profile, based on the floodplain areas in the DIVA database (lower line), for two
 167 characteristic coastline segments (A with a flat and B with a steep profile). Water level attenuation is accounted
 168 for by inclining the coastal profile according to equation 1 (upper line). Red dots on the adjusted coastal profile
 169 indicate the inundation length in the case of a water level with a constant slope of α , which represents the
 170 attenuation rate and for an incident water level equal to the corresponding increment height.

171

172 For the sensitivity analysis we used a range of attenuation rates that embraces the values reported in
 173 the literature (Table 1), where water level under storm conditions has been shown to decrease with
 174 distance from the coast. For reviewing the literature we employed the ISI Web of Knowledge and based
 175 our search on the keywords “surge”, “attenuation”, “water-level”. We selected studies that directly
 176 reported values of water level reduction with distance and did not include studies focussing on wave
 177 attenuation. We must note that the aim was not to conduct a systematic literature review but rather
 178 to identify a characteristic range of values that could support the sensitivity analysis. The identified
 179 studies all relate to coastal wetland environments. Although there are published studies of localised
 180 water level dynamics from flow-form interactions in urban and other settings, we have not come
 181 across similar landscape-scale assessments for other land use types. Therefore we broadened this
 182 review, where reported attenuation values were up to 70cm/km, by directly contacting scientists and
 183 data analysts with experience in field or modelling studies. Following their expert advice, we extended
 184 our analysis to include attenuation rates of up to 100 cm/km as an upper limit.

185



Event type	Landcover type	Location	Rate of water-level reduction	Method	Source
Storm surge	Bare land and Marsh	Modelled platform +0.5 m above sea level	10 cm / km (no vegetation, no channels) 26 cm / km (100% vegetation cover, no channels) 8 cm / km (100% vegetation cover, channel network)	Numerical modelling	Temmerman et al., 2012
Hurricane Isaac (2012)	Marsh	Louisiana	Up to 70cm/km water level reduction in presence of vegetation; 37 % reduction of total inundation volume	Numerical modelling	Hu et al., 2015
Hurricanes	Marsh	Multiple	1 m per 14.5 km 6.9 cm/km (range from 1m per 5km to 1m per 60km 20 - 1.7 cm/km)	Field Study	Corps of Engineers (1963) – In Wamsley et al., 2010
Hurricane Andrew (1992)	Marsh	Louisiana	1m per 20km-23.5km 5 - 4.3 cm/km	Field Study	Lovelace 1994
Hurricane Rita (2005)		Louisiana	1m per 4km to 1m per 25km 25 – 4 cm/km	Field Study	McGee et al. 2006 in Wamsley et al., 2010
Hurricanes Wilma (2005) and Charley (2004)	Mangroves Marsh	Florida	9.4 - 4.2 -cm/km	Field Study	Krauss et al., 2009
Hurricanes	Mangroves	Louisiana	23.3 – 1.7 cm/km	Field Studies	McIvor et al., 2012 (from various studies)
Hurricane Wilma (2005)	Mangroves	South Florida	Up to 50 cm/km (6-10 cm per km in the absence of mangroves)	Field study & modelling	Zhang et al., 2012
Hurricanes	Mangroves	South Florida	7.7 - 5.0 cm/km	Modelling	Liu et al., 2013

186 **Table 1:** Water level reduction rates, for different types of landcover, as reported in the literature.

187

188 We further constrained the sensitivity analysis by adjusting the range water attenuation rate for each
 189 segment based on the predominant land use type covering the area of every elevation increment. For
 190 estimating the predominant land use we employed the GlobCover Land Cover V2.3 dataset, a global
 191 land cover dataset with a resolution of 10 arc second (~300 meter at the equator). It is based on
 192 the ENVISAT satellite mission's MERIS sensor (Medium Resolution Image Spectrometer) covering the
 193 period between January and December 2009 and includes 22 land cover classes. As the available
 194 information on water attenuation rates by land use type is limited, we reclassified the data to seven



195 classes (forest, urban, cropland, grassland, mangroves, saltmarshes and Unknown) and assigned
196 maximum attenuation rates to each class (Table 2). For the model runs we used the five attenuation
197 categories (no, low, medium, high and maximum attenuation) corresponding to 0, 25%, 50%, 75% and
198 100% of the maximum values found in the literature / from expert review, for each class. These rates
199 were then used to incline the water surface in order to represent a constant water level attenuation
200 and the associated reduction in water levels (α) across the floodplain for each coastline segment.

201

Land Use Class	Maximum Attenuation (cm/km)
FOREST (1)	50
URBAN (2)	100
CROPLAND (3)	40
GRASSLAND (4)	25
MANGROVES (5)	50
SALTMARSHES (6)	25
UNKNOWN (0)	25

202 Table 2: Maximum attenuation rates per land use class used in the sensitivity analysis

203 *2.4 Sea-Level Rise and Socio-Economic Scenarios*

204 For global SLR in 2100 from a 1985 – 2005 baseline we used three scenarios: the 5% quantile of the
205 low Representative Concentration Pathway (RCP) 2.6; the median of the medium scenario RCP 4.5;
206 and the 95% quartile of the high scenario RCP 8.5. These scenarios are represented by SLR estimates
207 of 29, 50 and 110 cm (by 2100), respectively and were developed in the Inter Sectoral Model
208 Intercomparison Project Fast Track (for full details see Hinkel et al., 2014). Following Menendez and
209 Woodworth (2010), once mean sea level had been determined, future extreme water levels were
210 obtained by displacing upwards extreme water levels for different return periods (as included in the
211 DIVA database) with the rising sea level.

212 We used a single shared socio-economic pathway (SSP), namely SSP2, to represent changes in coastal
213 population and assets. SSP2 reflects a world with medium assumptions between the other four SSPs,
214 in terms of resource intensity and fuel dependency as well as GDP and population development (O’Neil
215 et al., 2014). Finally, we ran the DIVA model using a no-dike scenario, where no defence measures for
216 preventing coastal flooding are present.

217

218 **3. Results**

219 We present results for the different classes of attenuation rates, across the five indicators that
220 progressively include additional components of flood risk:



221

222 *3.1 Reduction of current flood exposure and risk*

223 Table 3 shows the results from the five categories of attenuation rates and both the absolute and
224 percentage reductions in the values of the five indicators against this baseline.

	Water Level Attenuation Category				
	NO	LOW (% decrease)	MEDIUM (% decrease)	HIGH (% decrease)	FULL (% decrease)
Area below the 1-in-100 year flood [km²]	859,059	654,474 (24%)	572,546 (33%)	519,871 (40%)	481,379 (44%)
People below the 1-in-100 year flood [million]	196	133 (32%)	115 (41%)	104 (47%)	97 (51%)
Assets below the 1-in-100 year flood [billion US\$]	52,663	34,814 (34%)	29,486 (44%)	26,467 (50%)	24,567 (53%)
People flooded [million/yr]	4429	3086 (30%)	2673 (40%)	2341 (47%)	2156 (51%)
Flood damages to assets for the 1-in-100 year flood [billion US\$/yr]	3,577	2,361 (34%)	1,972 (45%)	1,782 (50%)	1,660 (54%)

225 **Table 3:** Reduction, relative to the bathtub method, of five indicators of global exposure and risk, for different
226 water-level attenuation rates. Values are for a medium SLR scenario (median of the medium scenario RCP 4.5;
227 50 cm by 2100)

228

229 Our results show that accounting for water-level attenuation in the assessment of flooding results in
230 large differences in the values of the five indicators. For example, the area exposed to the 1 in 100-
231 year flood in 2015 reduces by up to 44% with the application of attenuation rates. The low attenuation
232 category results in an area reduction of 24% while the use of medium attenuation rates results in a
233 reduction of 33% (see Table 3). Interestingly, the number of people in the 1 in 100-year floodplain
234 reduces to 97 million when considering high attenuation. This is a reduction of 51%, which is similar to
235 the respective reduction in assets (53%) but higher than the reduction in area (44%) exposure. This
236 result reflects the high population density near the coast that has been reported in previous studies
237 (e.g. Neumann et al., 2015). Flood damages from the 1-in-100 year event are reduced in similar
238 proportion, totalling a reduction of more than 1.9 trillion US\$ (54%) globally, when considering
239 maximum attenuation rates.

240 The reduction in impacts is not uniform across the globe and varies considerably between different
241 countries. Some examples are given in Table 4, where accounting for water level attenuation reduces
242 area exposure by up to 73% in China, 40% in Bangladesh and 50% in the USA. At the same time, the



243 reduction in annual flood costs follows a different trend, with exposed assets reducing by up to 71% in
 244 China, 42% in Bangladesh and 25% in the USA, reflecting differences in the physical characteristics of
 245 the floodplain as well as in the spatial distribution of people and assets in the coastal regions of these
 246 countries.

247

Water Level Attenuation Rate	0 cm/km	10 cm/km (reduction)	20 cm/km (reduction)	50 cm/km (reduction)	100 cm/km (reduction)
Area below 1-in-100 year flood					
(km ²)					
Bangladesh	5831.4	4642.7	4204.7	3864.1	3528.3
		(20%)	(28%)	(34%)	(40%)
China	85864.8	43770.7	32588.3	27018.2	23413.4
		(49%)	(63%)	(69%)	(73%)
USA	69924.1	54244.5	43528.3	39346.5	35386.5
		(24%)	(39%)	(44%)	(50%)
Assets below 1-in-100 year flood					
(billion \$US)					
Bangladesh	67.2	54.7	48.8	43.6	39.4
		(19%)	(27%)	(35%)	(42%)
China	5137.8	2319.6	1724.2	1432.4	1258.1
		(55%)	(66%)	(72%)	(76%)
USA	533.7	431.4	388.2	360.7	341.9
		(19%)	(27%)	(33%)	(36%)

248

249 **Table 4:** Absolute and relative reduction of the 1-in-100-year floodplain area and associated exposed assets when
 250 applying different water-level attenuation rates for Bangladesh, China and USA. Values assume a medium SLR
 251 scenario (median of the medium scenario RCP 4.5; 50 cm in 2100).

252

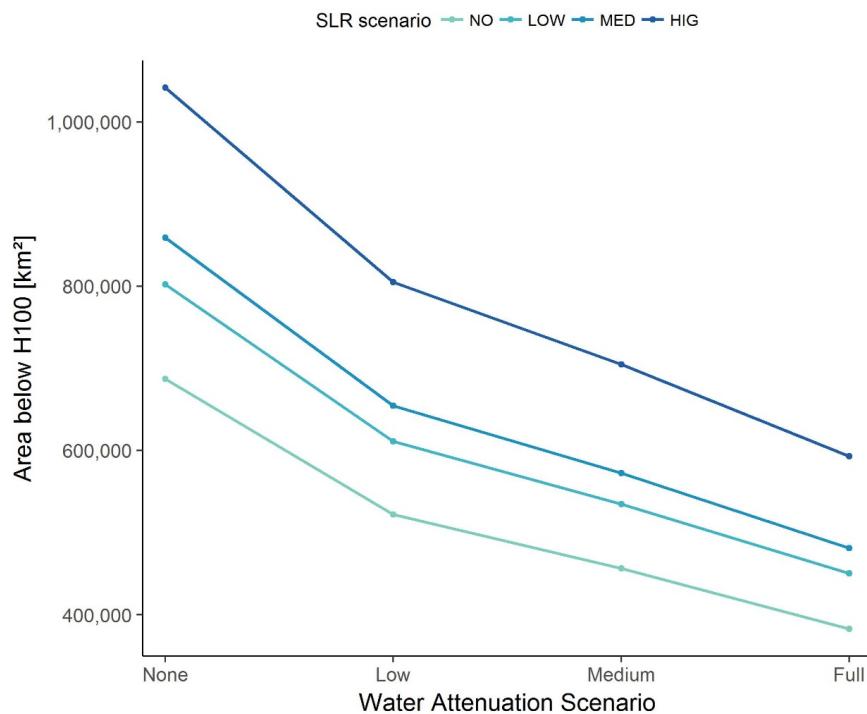
253 3.2 Comparison of attenuation rate uncertainty with sea-level rise uncertainty

254

255 Figure 3 illustrates the area of land located below the 1-in-100 year storm surge level (H100), plotted
 256 against the different attenuation rates for water level change. The inclusion of water-level attenuation
 257 in the assessment of flooding results in large reduction in the extent of the 100-year floodplain in 2100
 258 (Figure 3) under all SLR scenarios. Even the use of low attenuation of water levels results in a reduction
 259 of 230,000 km² of area exposed to the 1-in-100-year flood under the no-SLR scenario. This increases
 260 to 350,000 km² under the high SLR scenario. For the medium SLR scenario (median of the medium



261 scenario RCP 4.5; 50 cm by 2100), this reduction amounts to 31% and 40% of the total exposed area
262 at medium and full water level attenuation respectively. The relative reduction is larger (up to 60%)
263 for the high SLR scenario compared to the medium-, low- and no-SLR scenarios. Importantly, the
264 overall difference in the extent of the area of the 100-year floodplain between the no- and high-SLR
265 scenarios is of a similar order of magnitude to the difference in area extent between the no and low
266 water level attenuation rates, under any scenario. This indicates that when assessing area exposure
267 accounting for even relatively moderate rates of water level attenuation can be of similar importance
268 to the differences that result from different scenarios of SLR. This analysis, therefore, strongly suggests
269 that uncertainties related to the omission of this factor in global assessments of flood risk are of similar
270 magnitude to the uncertainties related to the magnitude of SLR expected over the 21st century.



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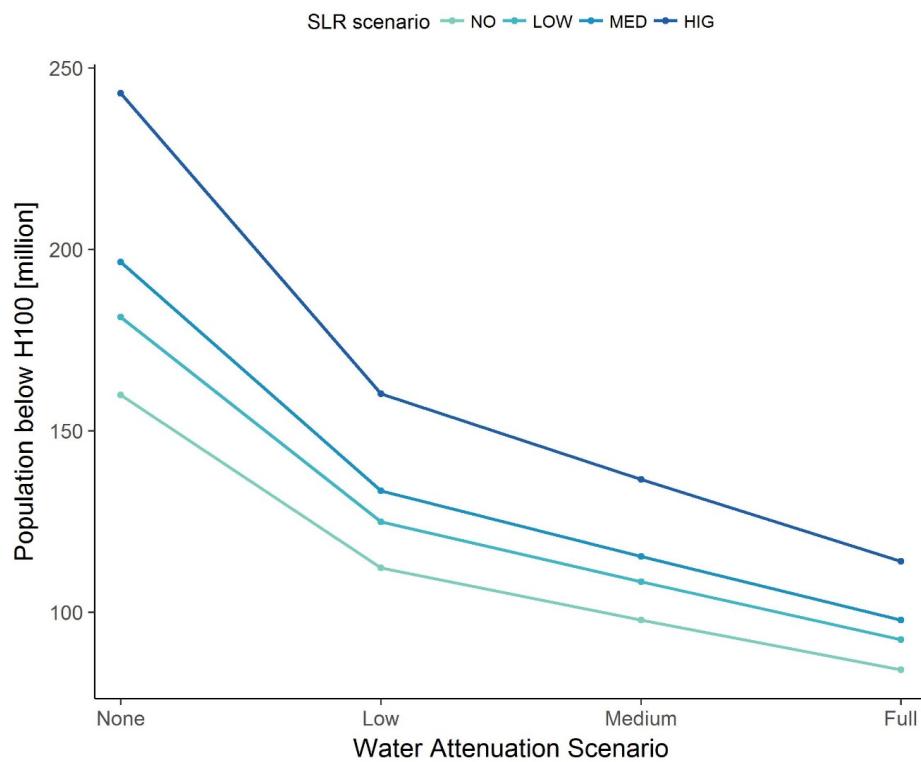
272 **Figure 3:** Global total extent of the 100-year floodplain, for different water level attenuation rates and SLR
273 scenarios.

274

275 Similar patterns can be observed for the exposure of population to the 1-in-100-year flood (Figure 4).
276 Low attenuation (Table 1), leads to a reduction of more than 30% in the exposure of population in
277 2100, under the high SLR scenario, bringing the number of people at risk in the 100-year floodplain



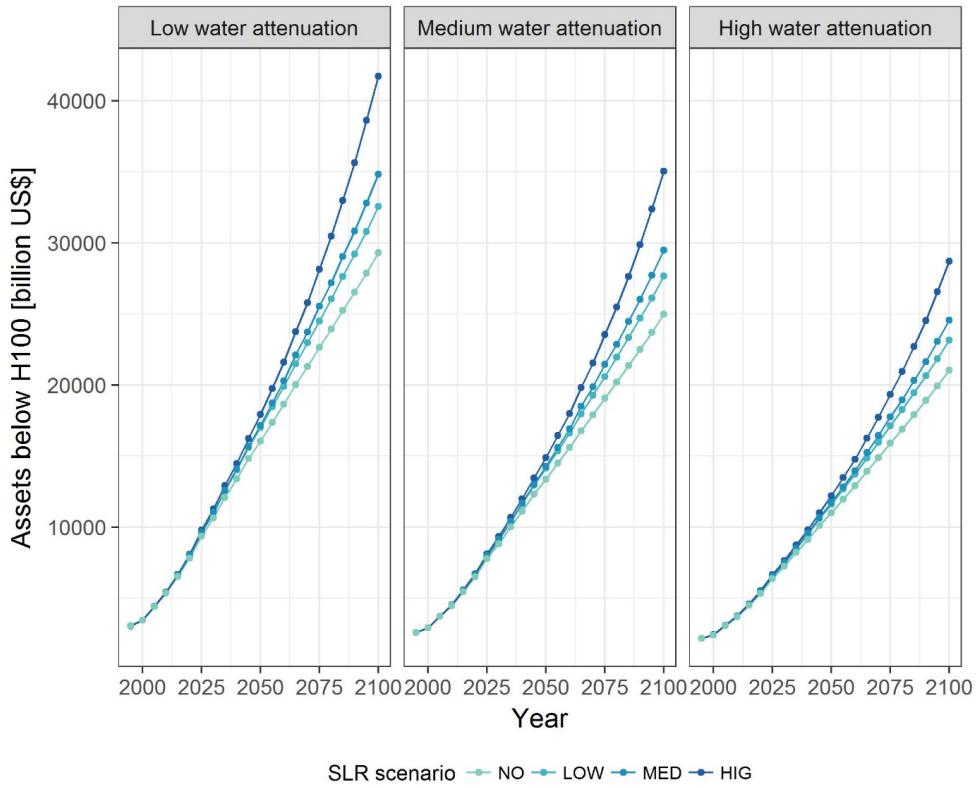
278 down by approximately 75 million. Moreover, medium attenuation leads to a reduction in flood
279 exposure by 100 million people, making population exposure lower than the exposure under no SLR
280 when attenuation is not considered. Again, this result suggests that accounting for water level
281 attenuation may be equally important to accounting for SLR uncertainty when assessing the exposure
282 of people to coastal flooding due to SLR.



283

284 **Figure 4:** Global estimates of population in the 100-year floodplain for different water-level reduction rates
285 (Table 1) and SLR scenarios.

286 The value of assets exposed to the 1-in-100-year flood, under all scenarios, are also reduced
287 substantially when accounting for water level attenuation (Figure 5). Considering low attenuation rates
288 results in a decrease in the exposure of assets of approximately 34% in 2100, for a medium SLR
289 scenario. A reduction of 50% in assets' exposure occur when high attenuation is used. Furthermore,
290 the use of a relatively moderate attenuation rate shifts the impact on assets' exposure by
291 approximately 30 years, under all SLR scenarios.

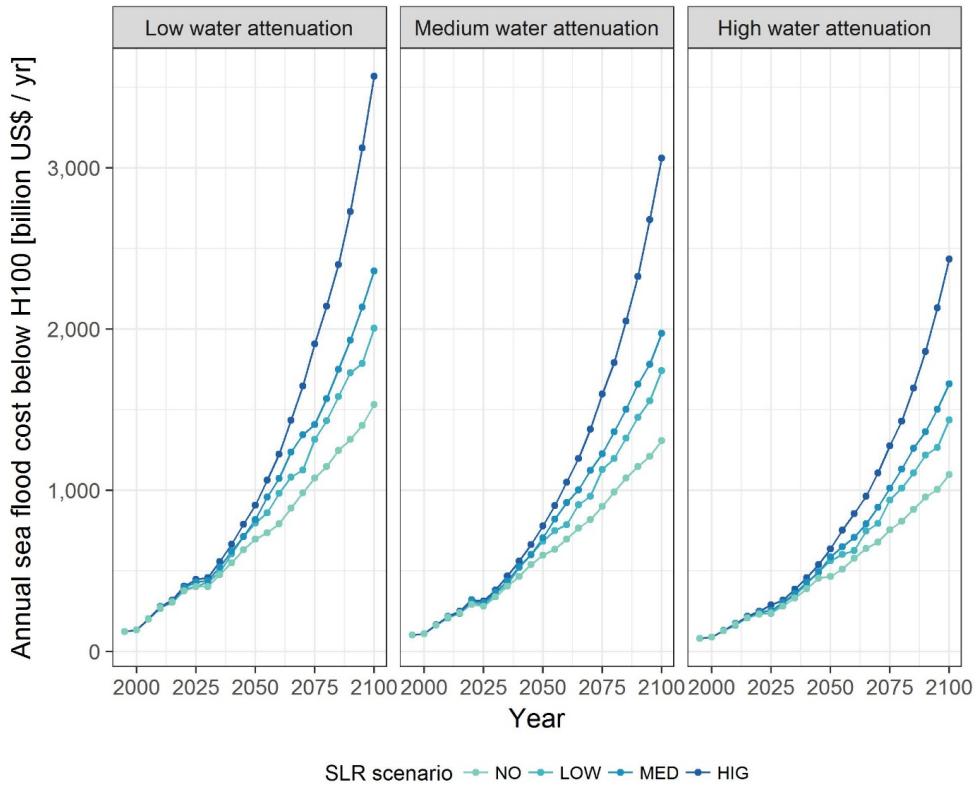


292

293 **Figure 5:** Temporal evolution of the amount of assets that are located in the 100-year floodplain for different
294 water-level reduction rates (Table 1) and SLR scenarios.

295

296 Damages also reduce considerably with the introduction of water level attenuation rates (Figure 6).
297 For example, the use of a low attenuation rate results in a 34% reduction in damages to assets in 2100
298 from the 1-in-100 year flood. The larger decrease in damages due to water level attenuation compared
299 to population and area exposure is due to the fact that, besides the decrease of the flood area extent,
300 water level attenuation leads to an additional reduction of flood depth with distance from the coast.
301 As water depth is an important parameter for calculating damages to assets (Thieken et al., 2005;
302 Penning-Rowsell et al., 2013), depth reduction further reduces the potential damages of assets due to
303 flooding.



304
305 Figure 6: Comparison of temporal evolution of sea-flood damage estimates for attenuation rates of 0,
306 10 and 50 cm/km, for different SLR scenarios.

307

308 **4. Discussion and Conclusions**

309 This study highlights the importance of accounting for the effects of hydrodynamic processes when
310 assessing the impacts of coastal flooding at national to global scales. In particular, water level
311 attenuation from the interaction of extreme inundation events with vegetated surfaces can lead to
312 considerably lower estimates of exposure of land area and population to coastal flooding.
313 Furthermore, this effect can lead to large reductions in potential damages, as lower water depths
314 combined with smaller flood extents give significantly lower flood-damage costs. The reduction in
315 exposure and risk is very pronounced, even when considering low water level attenuation rates.

316 Accounting for water level attenuation appears to be as important in assessing impacts as accounting
317 for uncertainties related to the total magnitude of SLR. In many of the cases explored, the difference
318 in impacts between no- and high-SLR scenarios is similar to the difference in impacts between no- and



319 low attenuation rates of up to 12.5 cm/km (excluding urban land use). This finding is of particular
320 relevance in environments where the floodplain substantially extends inland, such as in many of the
321 world's deltas and coastal plains.

322 It is widely acknowledged that the use of simplified methods, such as the bathtub method, can provide
323 useful first-order estimates of global impacts of SLR and associated flooding (Lichter et al., 2010; Hinkel
324 et al., 2014), although an overestimation of flood extent and depth with the use of the bathtub method
325 should be generally anticipated (Vousdoukas et al., 2016). Further, we must note that the reduction
326 that we observe with the use of water level attenuation rates does not necessarily reflect actual
327 impacts. These are likely to depend on additional factors which are usually not considered in global
328 assessments. For example, damage to assets in our analysis is based solely on water depth; factors
329 such as high local flow velocities from channelized flow, storm wave impacts, inundation by saline
330 water and sedimentation from flood waters are not taken into account. Such contributory factors can
331 lead to an increased cost of damages and thus counteract the lower impacts predicted from the use of
332 a water level attenuation term alone. Furthermore, the analysis reported here is predicated on the
333 assumption of a continuous increase in elevation with increasing distance from the shore. This study
334 shows that whilst this assumption is valid for the majority of coastal segments, there are segments
335 where this assumption does not hold true. In these cases model outputs may poorly describe flood
336 areas, flooded population numbers and asset damages and incorrectly predict the effect of changes in
337 the rate of water level attenuation. Nevertheless, and despite these caveats, our results emphasise
338 the importance of accounting for uncertainties in impact assessments stemming from the lack of
339 consideration of water level attenuation over coastal plains.

340 Our approach means to provide an illustration of the potential effects of water level attenuation, as
341 this process is not constant throughout the floodplain and depends on numerous parameters beyond
342 the type of the surface cover. These factors include storm duration, wind direction, water depth and
343 vegetation traits (Resio and Westerink, 2008; Smith et al., 2016; Stark et al., 2016). Furthermore,
344 applying a constant slope to account for water level attenuation is a strong simplification, since this
345 will vary between different storm events, but also under the influence of SLR. Nevertheless, given the
346 very high sensitivity of the outputs to even small changes in water level reduction rates; and the
347 obvious lack of sufficient data on the actual effect of different types of surface on attenuating water
348 levels during surges, we suggest that future work needs to focus on quantifying the water level
349 attenuation terms for different land uses. Thus, for example, both Brown et al. (2007), in the case of
350 modelled flooding following storm surge-induced sea defence failure, and Kaiser et al. (2011), in the
351 case of modelled tsunami wave impacts, have shown that disregarding buildings and associated
352 infrastructure (roads, gardens, ditches) when assessing inundation can lead to a large overestimation



353 of the extent of flooding. Furthermore, given the large range of uncertainty with respect to the actual
354 values of water level reduction associated with just one surface cover, wetland habitat (Table 1), future
355 impact modelling needs to focus on a better understanding of the temporal and spatial variation of
356 water levels across floodplains that show a wide variety of land use types and human occupancy,
357 including densely urbanised regions (e.g. Lewis et al., 2013; Blumberg et al., 2015).

358 Given that coastal wetlands can efficiently attenuate surge water levels, the results of this study give
359 a first estimate of how much of an impact reduction may result from the implementation of large-
360 scale, ecosystem-based flood risk reduction management schemes (e.g. Temmerman et al., 2013). In
361 addition, achieving lower water levels through the establishment of coastal wetlands not only reduces
362 impacts but may also affect the timing of potential adaptation tipping points by extending the
363 anticipated lifetime of adaptation measures. This would allow the development of alternative
364 adaptation pathways, a sequential series of linked adaptation options triggered by changes in external
365 conditions (Barbier, 2015), for coastal regions.

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368 **References**

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