



Influence of shearing rate on the residual strength characteristic of three landslides soils in loess area

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Abstract

In order to investigate the effect of the shearing rate on the residual shear strength of slip zone soils, a series ring shear tests were carried out on slip zone soils from three landslides in loess area at the two shearing rates (0.1mm/min and 1 mm/min). The slip zone soil specimens used in present study were from the northwest of China. Results indicated that the shear displacement to achieve the residual stage for specimens with higher shearing rate is greater than that of the lower rate. Relationship between the residual friction coefficients and normal stress shows that the residual friction coefficients for all specimens under the lower normal stress were greater than that under the higher normal stress at two shearing rate. Furthermore, the difference in the residual friction angle ϕ_r at the two shearing rates, $\phi_r(1) - \phi_r(0.1)$, under each normal stress level were either positive or negative values, with the maximum absolute value of $\phi_r(1) - \phi_r(0.1)$ reach up to 2.218° . However, the difference $\phi_r(1) - \phi_r(0.1)$ under all normal stresses was negative, which indicates that the residual shear parameters reduced with the increasing of the shearing rate in loess area.

Keywords: Residual shear strength; Ring shear test; Shearing rate; Normal stress; Slip zone soils



1. Introduction

Residual strength of soil is of great significance for evaluating the reactivating potential of the slope, in which consists of pre-existing sliding surface. Residual strength of a landslide soil is defined as the minimum constant value of strength along the slip plane, in which the soil particles are reoriented and subjected to sufficiently large displacements in relatively low shearing rate (Skempton, 1985).

Numerical studies have been done to assess the residual strength through the laboratory tests using ring shear tests and reversal direct shear tests (Chen and Liu, 2013; Vithana et al., 2012). It is a generally accepted fact that the measurement of the residual strength is most preferred done with a ring shear test since it allows the soil specimen be sheared at unlimited displacement which can simulate the field conditions more accurately (Lupini et al., 1981; Tiwari and Marui, 2005; Bhat, 2013; Sassa et al., 2004). Until now, several relationships between the residual strength and soil index parameters have been reported in the literature with a wide range of soil by using various kinds of ring shear apparatus (Hoyos et al., 2014; Jiang et al., 2016; Kimura et al., 2015; Li et al., 2013; Skempton, 1964). Furthermore, many studies have shown that the shearing rate may or may not affect the minimum value of soil strength at residual states (Suzuki et al., 2007; Grelle and Guadagno, 2010; Gonghui et al., 2010; Bhat, 2013; Tika and Hutchinson, 1999; Lemos, 1985; Morgenstern and Hungr, 1984; Tika, 1999).

From the high shearing rate aspect in the geotechnical literatures, Morgenstern and Hungr (1984) carried out ring shear tests on two types of coarse sand in high velocity and found that the frictional behavior was not affected by either the velocity or the normal stress. However, there were many researchers asserted that the effect of shearing rate on the shear behavior of soil cannot be ignored. For example, Skempton (1985), Tika and Hutchinson (1999) found that the faster shearing rate above 100 mm/min may bring about great qualitative changes in the residual behavior. Moreover, Tika et al. (1996) conducted fast ring shear tests on a wide range of natural soils and concluded that there are three types of rate effects on the residual strength, namely, a positive rate effect (the residual strength of soil at fast rate is higher than that of the slow rates), a neutral rate effect (the residual strength of soil is independent of the shearing rate) and a negative rate effect (the residual strength of the soil at higher speed is lower than that of the lower speed). Recently, Gratchev Ivan and Sassa (2015) reported that the residual strength of the clay decrease with the shear rate increase from 0.2 to 5 mm/s.

On the other hand, in the slow shearing rate range, Skempton (1985) reported that variation in the value of the residual



friction angle for shearing rates in a range of 0.05 to 0.35 mm/min was less than a 5% and concluded that the impact of shearing rate on the residual strength of clay is almost negligible within slow rate displacement. Similarly, Bhat (2013) concluded that there is hardly increase in residual strength of kaolin clay with the shearing rate ranging from 0.233 mm/min to 0.586 mm/min. Furthermore, Yokota et al. (1995) showed that residual strength is not affected by shearing rate lower than 1.01 mm/min in ring-shear tests. Except the above studies, other similar results were also found in clays that the residual strength is independent of the shearing rate (Chen and Liu, 2013; Tiwari and Marui, 2001). However, Suzuki et al. (2001) has reported the shearing rate ranging from 0.02 to 2.0 mm/min significantly affected the residual strength of kaolin clay and mud stone. Moreover, Gonghui et al. (2010) also has reported that the residual shear strength of the weathered serpentinite is positively dependent on the shear rate in the slow rate.

On general, the effect of the shearing rate on the residual strength of the soil has not been sufficiently studied in high and slow shearing rate range. Furthermore, except for a few studies, researchers have not widely reported the impact of the shearing rate on the residual strength of loess soil in relatively lower shearing rate range from 0.1mm/min to 1 mm/min. However, it should be noted that the residual strength parameters obtained from using different shearing rate may be adopted to provide a guide for designing some precision engineering which require high accuracy of the design parameters, thus, the effect of the shearing rate on the residual strength of soils should be fully understood to determine the parameters with high reliability. In addition, residual strength of soil plays a key role in assessing the stability analysis and evaluating the reactivation potential of landslides which consists of pre-existing slip plane surface. Therefore, accurate determination of the residual strength parameters and their dependence on the shearing rate may affect the stability evaluation of landslides. Thus, it is necessary to study the residual strength variation of loess in rate of shearing in order to have a good understanding of the suitable approach for the residual strength measurement.

In this backdrop, the present study investigated the effect of the shearing rate on the residual strength of soil samples obtained from three landslides in loessic- developed areas at two different shearing rates (0.1mm/min and 1 mm/min) by using a ring shear apparatus. The main objective of this study was to examine the change in the residual strength parameters of loess at different shearing rates and their relationship with the normal stress in naturally drained ring shear tests.



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85 **2. Geological setting of landslide sites**

86 Soil samples from three reactivated landslides in the northwest of China were selected for this study. Soil samples used
 87 for the ring shear tests and index measuring tests predominantly consist of loess deposits and were collected in a disturbed
 88 condition. For convenience, the names of landslide sites were abbreviated into Djg, Ydg, and Dbz. Figure 1 shows the study
 89 sites and some views of the landslides.

90 **Dingjiagou landslide**

91 The Djg landslide, located at the mouth of Dingjia Gully in Yan'an of China, is geologically composed of upper loess
 92 and lower sand shale in the Yan-chang formation. The dustpan-shaped landslide is inclined to the east, with its inclination
 93 75.85°. The landslide is 350 m in width, 180 m in length, 70 m in elevation. The average thickness of slip mass is around 20
 94 m, and the volume of landslide totaled approximately $105 \times 10^4 \text{ m}^3$. The slip mass is mainly constituted by loess, whereas the
 95 sliding bed consists of sand shale in Yan-chang formation. The thickness of the sliding zone varied from 30 to 50 cm. The
 96 front lateral region of the main slide section of the Djg landslide, where the sampling was performed, was found to be silty
 97 clay.

98 **Yandonggou landslide**

99 The Ydg landslide is located in the Qiaogou town of Yan'an in Shaan xi province of China. The top and the toe altitude of
 100 the landslide are about 1165 m and 1110 m above the sea level, with the height difference between the toe and the top of
 101 landslide about 55 m. The slides have well-developed boundaries with the main sliding direction of 240° and slope angle of
 102 30°. From the landslides profile, the sliding masses from top to bottom were classified by Q₃ loess, Q₂ loess and clay soil,
 103 respectively. Multiple landslide activities had occurred in this site, and the soil samples used in this study were collected
 104 from Q₂ loess stratum within the slide ranged from 4.5 to 18 m in height.

105 **Dabuzi landslide**

106 The Dbz landslide is located in the middle part of Shaanxi province (about 108°51'36" east longitude and 34°28'48" north
 107 latitude), China, which is a semi-arid zone dominated by loessic geology. In this region, the investigated site is classified as a
 108 typical loess tableland with quaternary stratum. The sedimentary losses in this area are grey yellow, and the exposure stratum



in this area has been divided into two stratigraphic units, namely, the upper late Pleistocene (Q_3) loess and the lower mid-Pleistocene (Q_2) loess, of which the Q_3 loess is younger. The Q_3 loess is closest to the surface and is up to approximately 12 m thick, while the thickness of Q_2 loess may reach an upper limit of about 50 m (Leng et al., 2018). The loess in this area have well-developed vertical joints (Sun et al., 2009). The travel distance and the maximum width of the slip mass are roughly estimated to be 121.55 m and 133.46 m, respectively. The armchair-shaped landslide shows an apparent sliding plane, with an area of approximately 15,660 m² and about 66.25 m maximum difference in elevation. The main direction of this landslide is approximately 355°. The exposed slip zone in the side scarp of the landslide, where the sampling was done, was found to be entirely in the Q_2 loess stratum of the Dbz landslide site. The thickness of narrow-band slip zone loess is less than 1.0 cm, inclined at around 65° to the horizontal direction. Since the band of slip zone is thin, mix soils which consist of the slip zone soils, the very thin upper and lower parts of the loess are mingled together and served as the representative samples of the slip zone loess in this site.

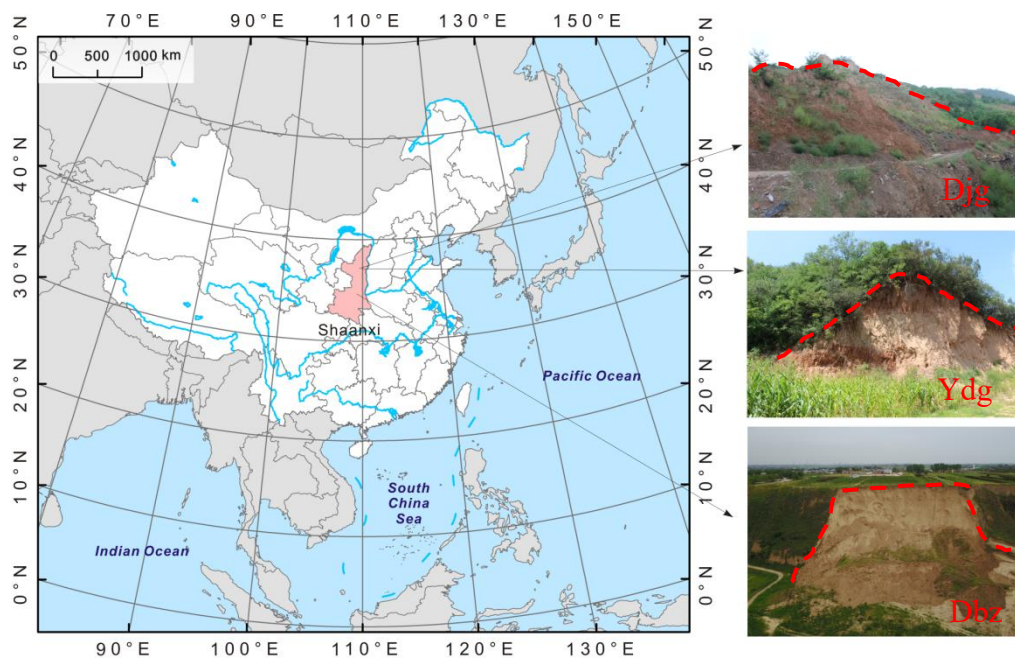


Figure 1. Location of study sites and some views of landslides

3. Experimental scheme

3.1. Testing sample



124 The fact that the residual shear strength is independent of the stress history was reported by many researchers (Bishop et
 125 al., 1971; Vithana et al., 2012; Stark Timothy et al., 2005). Thus, disturbed loess samples from each of the three landslides
 126 weighing about 25 kg were collected from the slip surface soil of each slide and used to determine residual shear strength.

127 The soil samples were air-dried and then crushed with a mortar and pestle, and subsequently processing it through 0.5 mm
 128 sieve. Distilled water was added to the soil samples until desired density and water content were obtained. The physical
 129 parameters such as natural moisture content, specific gravity, bulk density, plastic limit, and liquid limit were determined in
 130 accordance with the Chinese National Standards (CNS) GB/T50123-1999 (standards for soil test methods) (SAC, 1999),
 131 but clay size was defined to be less than $2\mu\text{m}$ followed ASTM, D422 (ASTM, 2007). Each soil sample was separated into
 132 clay (sub 0.002 mm), silt (0.002-0.075 mm), and sand (0.075-0.5 mm) fractions. The physical indexes of the soil are listed in
 133 Table 1.

134 The grain size distribution of soil was measured using a laser particle size analyzer Bettersize 2000 (Dandong Bettersize
 135 Instruments Corporation, Dandong, China). The sieved soil samples were used to determine particle size distribution. In this
 136 study, soil samples were treated with sodium hexaphosphate, serving as a dispersant, to disaggregate the bond between the
 137 particles. The results show that the clay fraction in Djg landslide soil (24%) is more than two times than that from Ydg (9%)
 138 and Dbz (9.1%). Furthermore, the particle size analyses illustrates that the percentage of silt-sized soil in three landslides
 139 ranged from 75.66% to 87.4%. In addition, Ydg landslide soil consists of the greatest percentage of the sand fraction which
 140 reaches up to 10.55%.

141 In present study, a total of twenty four specimens were tested at two shearing rate (0.1mm/min and 1 mm/min) and under
 142 normal stresses ranged from 100kN/m^2 to 400kN/m^2 in a ring shear apparatus.

143 **Table 1** Physical parameters of slip-zone loess

<i>sites</i>	ρ_d	W	ρ	G_s	W_L	W_p	Grain size fractions (%)			
							<0.002m	0.002-0.005m	0.005-0.075	0.075-0.5m
Djg	1.74	19.	2.0	2.6	3	20	24	11.48	64.18	0.34
Ydg	1.47	18	1.7	2.7	3	19	9	5.28	75.17	10.55



Dbz	1.48	16	1.7	2.7	3	21	9.1	6.4	81	3.5
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Notes: ρ_d = dry density; w=moisture water content; ρ = bulk density; G_s = specific gravity; W_L =liquid limit; W_p = plastic limit

3.2. Testing apparatus

The advantage of a ring shear test apparatus to measure residual shear strength including its ability to allow unidirectional shearing of a soil specimen (Bishop et al., 1971;Tika, 1999;Suzuki et al., 2007;Bromhead, 1979). Thus, a ring shear apparatus was used in this study.

An advanced ring shearing apparatus (SRS-150) manufactured by GCTS (Arizona, USA) was adopted in ring shear tests and the photos of apparatus were shown in Fig.2, which consists mainly of a shear box with an outer diameter of 150 mm, an inter diameter of 100 mm and the maximal sample height of 250 mm. The shearing box consists of the upper shearing box and the lower shearing box. In the shearing process, the upper shearing box keeps still while the lower one rotates. The apparatus, which provides effective specimen area of 98 cm², is capable of shearing the specimen for large displacement in single direction. The annular specimen is confined by inside and outside metal rings. Moreover, the specimen is confined by bottom annular porous plates and top annular porous plates in which have sharp-edged radial metal fins which protrude vertically into the top and bottom of the specimen at the shearing process. The normal stress, shearing strength and shearing displacement can be monitored while shearing by computer. The measurement features of the ring shear apparatus employed in this study are described as follows: shearing rate range from 0.001 to 360 degrees per minute, 10 kN axial load capacity, 300 N.m continuous torque capacity, maximum normal stress of 1000 kN/m².

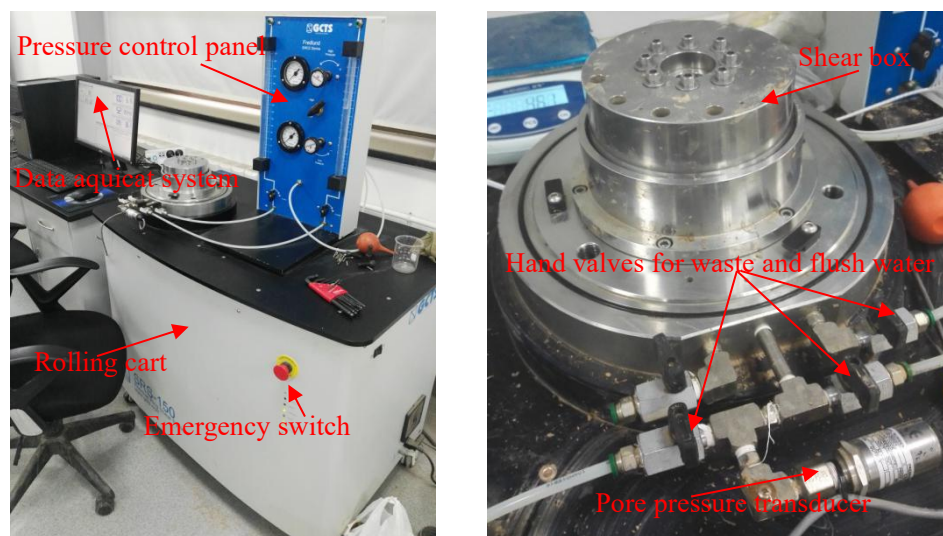


Figure 2. Ring shear apparatus (SRS-150)

3.3. Testing procedure

In present study, reconstituted samples of the sub 0.5 mm soil fraction were used in the testing as it was reported that the residual strength of the soil was unaffected by its initial structure (Bishop et al., 1971; Vithana et al., 2012). Specimens were first prepared by adding distilled water to the air-dried soil until the saturated moisture contents of the three landslide soils were obtained. Then, specimens were kept in a sealed container for at least one week to fully hydrate. Specimens are then reconstituted in the ring-shaped chamber of the apparatus by compaction. In order to make the sample uniform while packing, the sample was placed in three layers, and each layer was tamped under a vertical stress which is lower than the given normal stress to achieve the design height. The final height of the specimen in the ring shear apparatus after tamp varied but was typically about 20 mm (to achieve a specific bulk density). The specimen was then consolidated under a specific effective normal stress in a range of 100kN/m² to 400kN/m² until required consolidation was achieved. Then, the consolidated specimen is subjected to shearing under constant normal stress by rotating the lower half of the shear box attached to a gear, while the upper half remains still. In ring shear tests, the normal stress at the shearing was the same as at consolidation stage.

In this study, ring shear tests were performed in a single stage under drained condition and the samples were subjected to shear until the residual state was achieved. Drain condition of the shearing process is provided by two porous stones



attached on the top and the bottom platen of the specimen container. As for soil specimens with low permeability, the rate of excess pore pressure generation in the shear box may exceeded that of pore-pressure dissipation, this type of condition is identified as naturally drained condition in previous studies(Okada et al., 2004). Furthermore, Tiwari (2000) asserted that it was acceptable to use a shearing rate below 1.1 mm/min to simulate the field naturally drained condition. Thus, shearing rates of 0.1 mm/min and 1 mm/min were used in this study to simulate the naturally drained condition of the slip zone soils.

4. Results and discussions

Twenty four specimens were tested to investigate the shear characteristics of the slip-zone soils in the ring shear apparatus. Tests results are shown in this section.

4.1. Shear behavior

Figures 3a, 4a and 5a show the typical shear characteristics of the slip-zone soils (shearing rate 0.1 mm/min and 1 mm/min) obtained from three different locations, where, the shear stress is plotted against the shear displacement at the normal stress ranged from 100kN/m² to 400kN/m². It is a widely accepted fact that normal stress has effect on the shear behavior of the soil (Kimura et al., 2015;Stark Timothy et al., 2005;Stark et al., 2005;Eid, 2014), thus, the shear behavior of samples at the peak and residual stages, where, the determined peak friction coefficient as well as residual friction coefficient are plotted in Figure 3b, 4b, and 5b against the corresponding effective normal stresses. The friction coefficient is defined as the shear stress divided by the effective normal stress.

Figures 3a, 4a and 5a demonstrate that shear stress increases dramatically within small shear displacement and then reduces with shearing displacement, until residual conditions were achieved at large displacements. Furthermore, it is clear that peak strength as well as residual strength of the samples with high shearing rate is almost smaller than that of the samples with low rate. In Figures 3a, 4a and 5a, a clear drop can be seen, at any normal stress, for specimens obtained from all sites. It is obvious that Djg specimens showed greater peak-post drop than that of Ydg and Dbz specimens. According to the conclusion that the residual stage is attained if a constant shear stress is measured for more than half an hour (Bromhead, 1992), it can be seen that the shear displacement to achieve the residual stage for specimens with higher shearing rate is greater than that of the lower rate. For example, the minimum shear displacements for attaining residual condition for Djg specimens with low and high shearing rate were about 360mm and 650mm, respectively. Under the shearing rate of

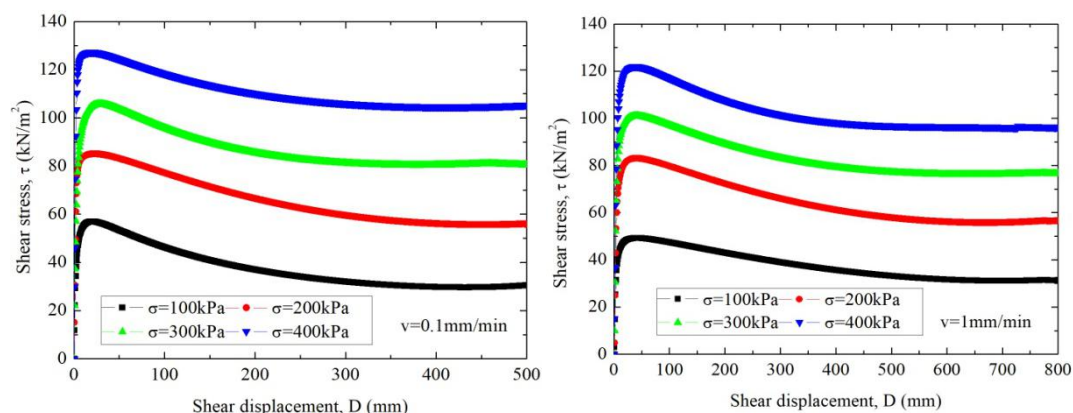


0.1mm/min and 1mm/min, Ydg specimens need approximately 80mm and 1,400 mm displacement to achieve residual stage. However, Dbz specimens require about 40mm and 60mm displacement to reach residual condition for low and high shearing rate, respectively.

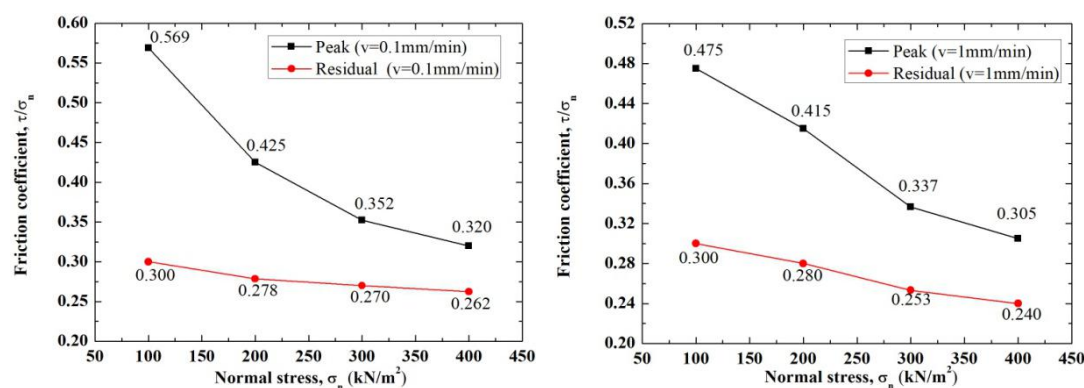
4.2. Effect of normal stress on the friction coefficients

It can be seen from the Figures 3b, 4b and 5b that the friction coefficients (peak and residual) are higher at lower effective normal stress levels. For example, the peak and residual friction coefficient of Djg landslide soils at the shearing rate of 0.1mm/min reduced from 0.569 to 0.32 and from 0.3 to 0.262, respectively. Similarly, results obtained from other two landslides soils also showed that the friction coefficients decrease nonlinearly with the normal stresses. Furthermore, specimens with shearing rate of 0.1mm/min attained greater friction coefficients than that with shearing rate of 1mm/min.

In order to get an insight into the effects of the normal stress on the slip zone shear strength, the shear behavior of the soil sheared at the normal stress of 100kN/m² and 400kN/m² were selected for analysis. At the normal stress of 100kN/m², Djg samples showed about 47.3% and 36.8% decrease in the friction coefficient from the peak friction coefficient at the shearing rate of 0.1 and 1 mm/min, respectively, which is greater than in the Ydg (about 9.8% and 10.3%) and Dbz (about 2.4% and 3.2%) samples. In Figures 3b, 4b and 5b, on average, it is obvious that the decrease of the friction coefficient from the peak strength to the residual strength in the Djg sample was almost 18.1% and 21.3% for the sample consolidated at normal stress of 400kN/m² under the shearing rate of 0.1mm/min and 1mm/min (Figure 3b), While the friction coefficient reduction in Ydg sample with low and high shearing rate were only about 4.1% and 4.8% (Figure 4b). And the friction coefficient reduction in Dbz samples with low and high rate were only approximately 5.6% and 6.0% (Figure 5b) from the peak strength, respectively. Based on the conclusion that the post-peak drop in strength of soil is only due to particle reorientation after the peak strength (Mesri and Shahien, 2003;Skepmton, 1964), the results in this study demonstrated that the Djg landslide soil existed the greater particle reorientation compared with that of other two landslide soils.

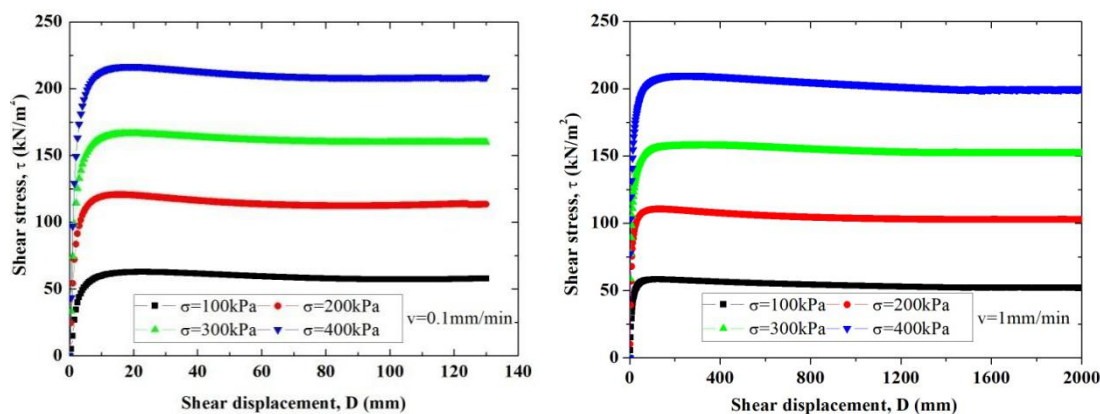


(a) Relationship between shear stress and shear displacement

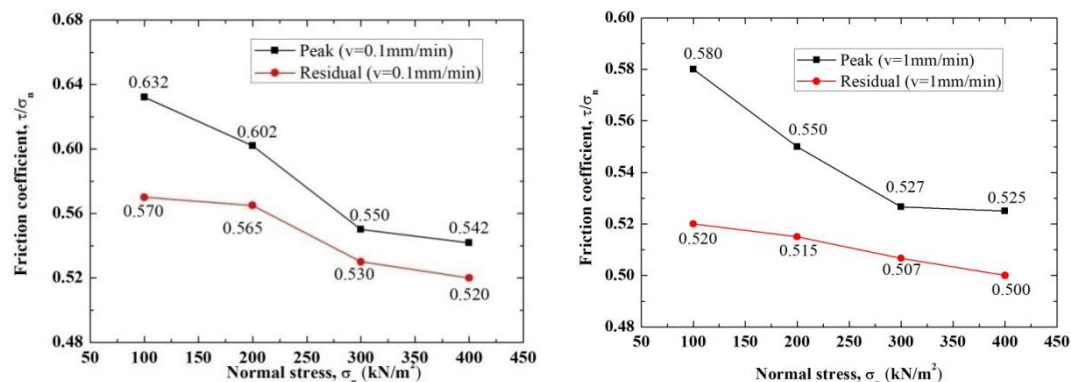


(b) Relationship between friction coefficient and normal stress

Figure 3. Shear behavior characteristics of Djg soil samples

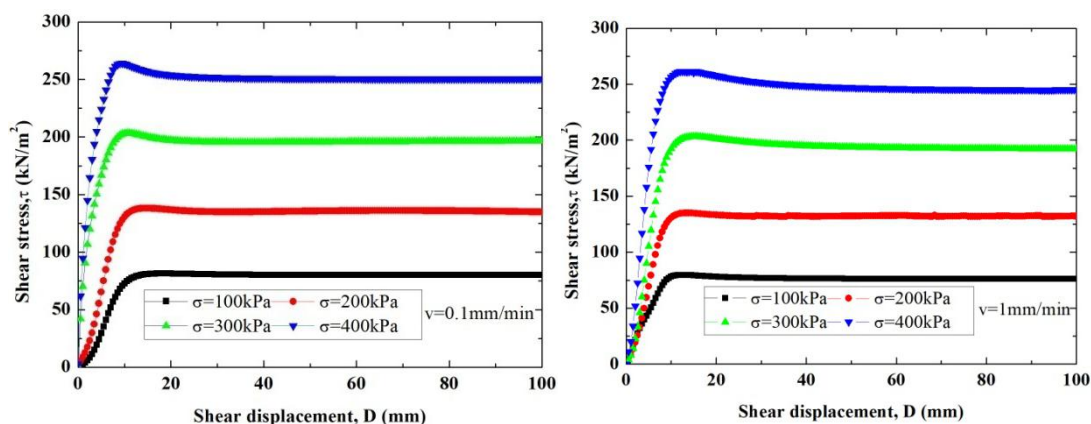


(a) Relationship between shear stress and shear displacement

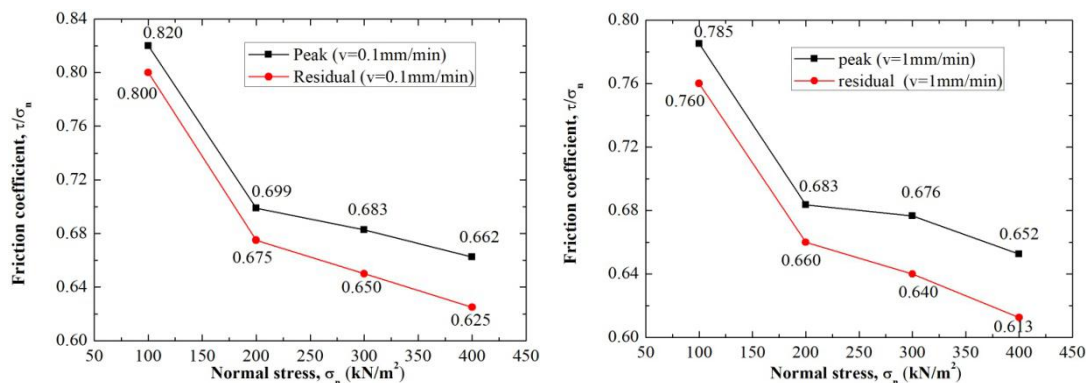


(b) Relationship between friction coefficient and normal stress

Figure 4. Shear behavior characteristics of Ydg soil samples



(a) Relationship between shear stress and shear displacement



(b) Relationship between friction coefficient and normal stress

Figure 5. Shear behavior characteristics of the Dbz soil samples



4.3. Effects of shearing rate on residual strength parameter

For the representative samples described above, Figures 6, 7 and 8 show the relationships between the residual friction coefficient and the normal stress, and the residual strength parameters. The residual friction coefficient is plotted against the normal stress. The residual friction coefficient is defined as the residual shear strength divided by normal stress. It is widely recognized that the shear strength parameters including cohesion and friction angle (Terzaghi, 1951; Stark et al., 2005). However, according to the previous studies, the residual angle of soils varies depended on the soil properties as well as the magnitude of normal stress provided the residual cohesion of soil is zero (Skepmton, 1964; Bishop, 1971; Kimura et al., 2014). Thus, in this study, the residual frictions are calculated by Coulomb's law assumed the residual cohesion is zero. The residual strength parameters were defined as $\phi_r(0.1)$ and $\phi_r(1)$ at the low shearing rate and high shearing rate, respectively. And the difference between the residual friction angles at two shearing rate was defined as $\phi_r(1) - \phi_r(0.1)$. Comparatively, the residual friction coefficient was defined as $\tau_r/\sigma_n(0.1)$ at the low shearing rate and $\tau_r/\sigma_n(1)$ at the high shearing rate, respectively. Furthermore, the difference between the residual friction coefficients was defined as $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$. Table 2 summarized the residual parameters of the landslide soils.

Figure 6 shows that the residual friction coefficients were relatively low in Djg samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ at the normal stress of 100 kN/m^2 to 400 kN/m^2 ranged from 0.3 to 0.262 and from 0.3 to 0.24, respectively. The difference between the friction coefficients, $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$, at each normal stress level are varied in a range of -0.022 to $+0.002$. For the difference between the residual friction angles, $\phi_r(1) - \phi_r(0.1)$, ranged from -1.212° to $+0.079^\circ$ (Table 2). For normal stress above 200 kN/m^2 , the coefficient $\tau_r/\sigma_n(0.1)$ was found to be greater in the magnitude than the coefficient $\tau_r/\sigma_n(1)$.

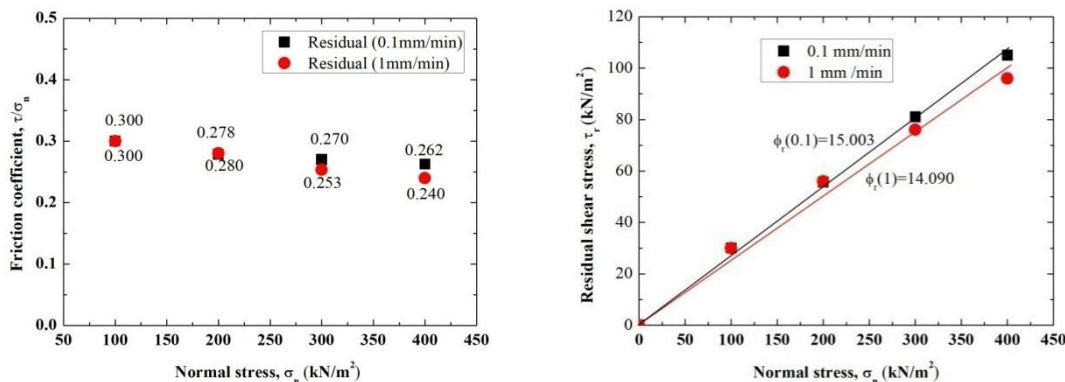




Figure 6. Relationships between residual shear stress and normal stress, and residual strength parameter for Djg soil sample

Figure 7 gives the relationship between the residual coefficient and normal stress, and residual shear strength parameter for Ydg samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ under the normal stress of 100kN/m² to 400kN/m² ranged from 0.57 to 0.52 and from 0.52 to 0.50, respectively. Furthermore, the difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at each normal stress was from -0.05 to -0.02. As for the difference between the residual friction angles, $\phi_r(1) - \phi_r(0.1)$, was in a range of -2.218° to -0.909°. In case of Ydg soil sample, there was insignificant reduction in residual friction coefficients with the increasing of shearing rate for all normal stresses.

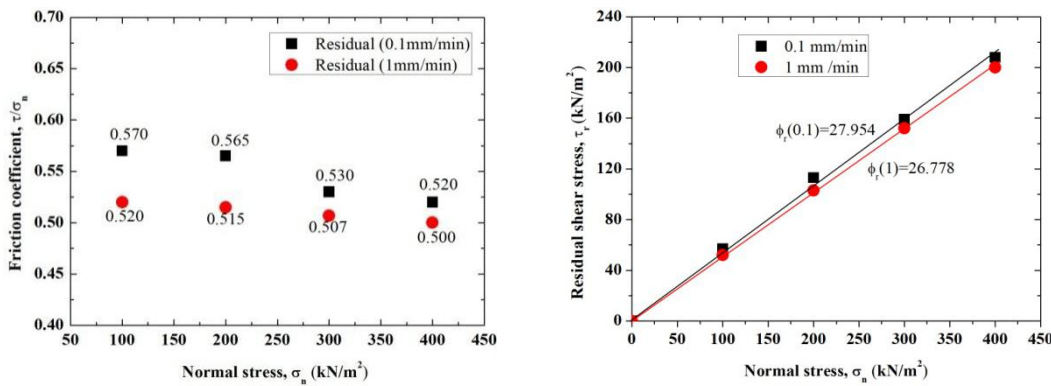


Figure 7. Relationships between residual shear stress and normal stress, and residual strength parameter for Ydg soil sample

Figure 8 presents the results of the Dbz samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ at the normal stress of 100kN/m² to 400kN/m² ranged from 0.8 to 0.625 and from 0.76 to 0.613, respectively. The difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at each normal stress was from -0.04 to -0.01. The difference $\phi_r(1) - \phi_r(0.1)$ was from -1.425° to -0.405°. For Dbz samples, there was somewhat decrease of the residual friction coefficients with the increasing of the shearing rate for all normal stress levels. It is noted that the maximum difference was found at the lowest normal stress of 100kN/m².

Table 2 summarizes residual strength parameters including $\phi_r(0.1)$ and $\phi_r(1)$ of all specimens obtained from the ring shear tests in this study. As for the Djg samples, the residual strength parameter $\phi_r(0.1)$ and $\phi_r(1)$ for all normal stress were



found to be 15.003° and 14.09° , respectively. However, the residual friction angles $\phi_r(0.1)$ and $\phi_r(1)$ of the Ydg samples were obtained to be 27.954° and 26.778° , respectively. In the case of Dbz sample, the friction angles $\phi_r(0.1)$ and $\phi_r(1)$ were high, 32.822° and 32.293° , respectively. The residual friction angles $\phi_r(0.1)$ and $\phi_r(1)$ under all normal stresses were from 15.003° to 32.822° and from 14.09° to 32.293° , respectively.

Due to the influence of the shearing rate, the difference $\phi_r(1) - \phi_r(0.1)$ in the Djg, Ydg and Dbz samples, were -0.913° , -1.176° and -0.529° , respectively. Wang (2014) and Fan et al. (2017) asserted that the residual shear strength of remoulded loess hardly affected by shearing rate below 5mm/min. However, the results in this study shown that $\phi_r(1) - \phi_r(0.1)$ under all normal stress levels were negative for slip zone loess. Moreover, the maximum value of the difference $\phi_r(1) - \phi_r(0.1)$ even reached about 1.176° .

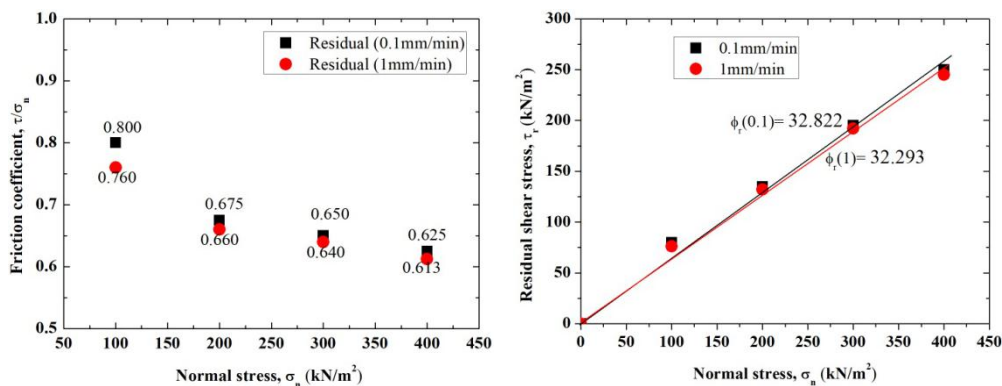


Figure 8. Relationships between residual shear stress and normal stress, and residual strength parameter for Dbz soil sample

Table 2 Residual shear strength parameter of landslide soils

No	Sample	Normal stress(kN/m ²)	Residual strength parameter				Difference in parameter	
			0.1 mm/min $\phi_r(0.1)$ ($c_{r(0.1)}=0$) (Degrees)	1 mm/min $\phi_r(1)$ ($c_{r(1)}=0$) (Degrees)			$\phi_r(1)-\phi_r(0.1)$ (Degrees)	
			Under each σ_n	Under all σ_n	Under each σ_n	Under all σ_n	Under each σ_n	Under all σ_n



1	Djg	100	16.699	15.003	16.699	14.090	0	-0.913
		200	15.563		15.642		0.079	
		300	15.110		14.216		-0.894	
		400	14.708		13.496		-1.212	
2	Ydg	100	29.683	27.954	27.474	26.778	-2.209	-1.176
		200	29.466		27.248		-2.218	
		300	27.923		26.870		-1.053	
		400	27.474		26.565		-0.909	
3	Dbz	100	38.660	32.822	37.235	32.293	-1.425	-0.529
		200	34.019		33.425		-0.594	
		300	33.024		32.619		-0.405	
		400	32.005		31.487		-0.518	

5. Influence of the shearing rate on the residual friction angles according to soil properties

Figure 9 depicts the relationships between residual friction angles as well as the difference in the residual friction angles and soil properties including liquid limit (LL), plasticity index (Ip) and clay fraction (CF) at two shearing rates. The residual friction angles at two shearing rates decreased nonlinearly with the increasing of the LL. As for the relationship between the ϕ_r and Ip, the ϕ_r under the low and high shearing rates decreases from about 32° to 15° with increasing the plasticity index from 11 to 16. With increasing of CF from 9% to 24%, the residual friction angles under low and high shearing rates were found to decrease. Interestingly, for Dbz and Ydg soils of which have similar percentage of clay fraction, the residual friction angles at both shearing rates varied. However, in the relationships between the difference in the residual friction angles and the soil properties, no clear correlations were found.

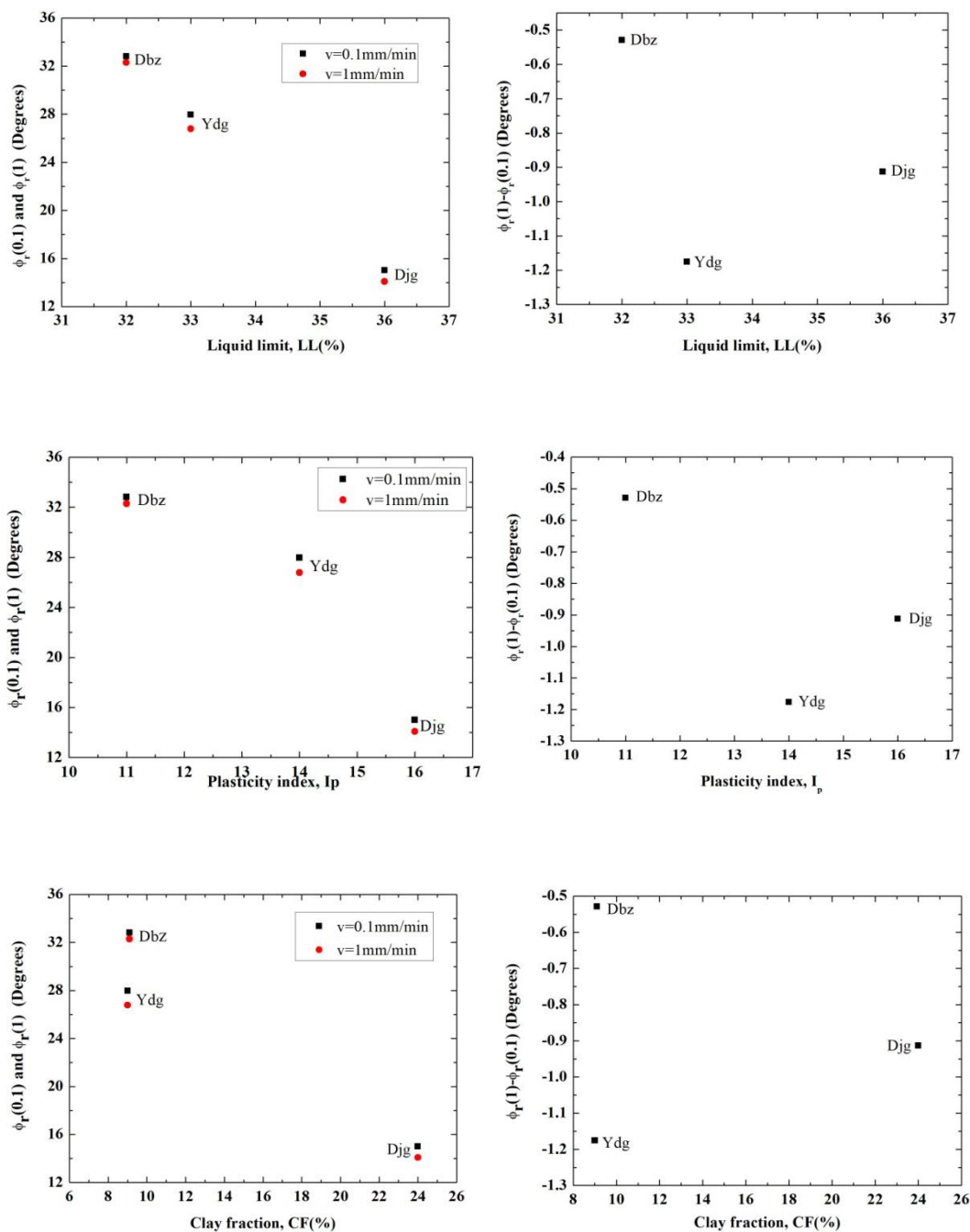


Figure 9. Relationships between residual shear parameter, the difference in residual shear parameter and the soil properties at two shearing rates



Conclusions

The shearing rate of slip zone soil of landslide may be changed after the occurrence of the first landslide activity, thus, the residual shear strength of the slip zone soil could be changed accompanying this process. As for some precision engineering which require high accuracy of the design parameters, the shearing rate effect on the residual shear strength should be fully investigated. In this study, at the shearing rate of 0.1 mm/min and 1 mm/min, a series of ring shear tests under normal stress ranged from 100kN/m² to 400kN/m² were performed on reconstituted slip zone soil samples obtained from three landslides in loess area. The main results can be summarized in the following points:

- (i) The shear displacement to achieve the residual stage for specimens with higher shearing rate is greater than that of the lower rate.
- (ii) As for slip zone soils in this study, specimens with lower shearing rate attained greater friction coefficients than that with higher shearing rate.
- (iii) At the two shearing rate (0.1 mm/min and 1 mm/min), the residual friction coefficient under the lower normal stress was higher than that under the higher normal stress in all samples. In addition, there was a nonlinearly decrease trend of the residual friction with the normal stress.
- (iv) For slip zone soils in this study, the difference at the two shearing rate, $\phi_r(1) - \phi_r(0.1)$, under each normal stress level were either negative or positive. However, under all normal stress, the difference at the two shearing rate $\phi_r(1) - \phi_r(0.1)$ was found to be positive.
- (v) The residual friction angles reduce with the increasing of shearing rate. Furthermore, the maximum magnitude of the difference between the residual friction angle $\phi_r(1)$ and $\phi_r(0.1)$ was even obtained to be approximately 1.176° in loess area.
- (vi) The relationships between the ϕ_r under two shearing rates and soil properties including liquid limit and plasticity index, demonstrated that the ϕ_r at both shearing rates decrease gradually with the increasing of LL and Ip. However, no clear correlations between the difference in the ϕ_r at low and high shearing rates and the soil properties were found.



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