

1 **Contrasting large fire activity in the French Mediterranean**

2
3 Anne Ganteaume¹, Renaud Barbero¹

4 ¹ Irstea, Mediterranean Ecosystems and Risks,, Aix-en-Provence, France

5
6 *Correspondence to:* Anne Ganteaume (anne.ganteaume@irstea.fr)

7 **Abstract**

8 In the French Mediterranean, large fires have significant socio-economic and environmental impacts.
9 We used a long-term geo-referenced fire time series (1958-2017) to analyse both spatial and temporal
10 distributions of large fires (LF; ≥ 100 ha). The region was impacted in some locations up to 6 times by
11 recurrent LF and 21% of the total area burned by LF occurred on a surface that previously burned in
12 the past, with potential impact on forest resilience. We found contrasting patterns between the East
13 and the West of the study area, the former experiencing fewer LF but of a larger extent compared to
14 the latter, with an average time of occurrence between LF exceeding 4,000 ha < 7 years mostly in the
15 eastern coastal area and > 50 years in the West. This longitudinal gradient in LF return level contrasts
16 with what we would expect from mean fire weather conditions strongly decreasing eastwards during
17 the fire season but is consistent with larger fuel cover in the East, highlighting the strong role of fuel
18 continuity in fire spread. Additionally, our analysis confirms the sharp decrease in both LF frequency
19 and burned area in the early 1990s, due to the efficiency of fire suppression and prevention reinforced
20 at that time, thereby weakening the functional climate-fire relationship across the region.

32 **1 Introduction**

33 It is now unanimously agreed that large fires have most significant socio-economic and environmental
34 impacts, threatening or damaging infrastructures, ecosystems, and even costing human life, especially
35 in the expanding wildland-urban interfaces (WUI) (Blanchi et al., 2014; Syphard and Keeley 2015;
36 Radeloff et al., 2018). However, the definitions of what can be considered as a large fire are numerous
37 (Shvidenko and Nilsson, 2000; Stocks et al., 2002; Barbero et al., 2014a, Stavros et al., 2014; Nagy et
38 al., 2018; Tedim et al., 2018), the cutoff being arbitrary or statistically assessed. Usually, large fires
39 represent only a small proportion of the total number of fires but they typically account for the bulk of
40 burned area in many regions throughout the world (Stocks et al., 2002; San Miguel-Ayanz et al., 2013;
41 Stavros et al., 2014, Barbero et al., 2014a, 2014b, Ganteaume and Guerra, 2018) and determine in fact
42 the trend and interannual variability in the total burned area.

43 Large fires and fire severity have increased over the past several decades across parts of the
44 globe (Pausas and Fernández-Muñoz, 2012; Dennison et al., 2014; Stephens et al., 2014), these
45 changes being attributed to a combination of climate change (Westerling et al., 2006; Bradstock et al.,
46 2009; Flannigan et al., 2009; Barbero et al., 2015; Abatzoglou and Williams, 2016) and past fire
47 suppression (McKenzie et al., 2004; Littell et al., 2009; Miller et al., 2009). However, these upward
48 trends are not universal and some landscapes, mostly in southern Europe, have not experienced such
49 increases in large fires and even showed a decline since the 1990s (San Miguel-Ayanz et al., 2013;
50 Ruffault and Mouillot, 2015; Ganteaume and Guerra, 2018), albeit conflicting signals were found
51 across parts of Portugal and Spain (Turco et al., 2016). This overall fire reduction has been attributed
52 to an increased effort in fire management after the large fires in the 1980s (Turco et al., 2016;
53 Fréjaville and Curt, 2017).

54 In Mediterranean systems, bottom-up drivers are generally thought to play a strong role in fire
55 activity. Indeed, ignitions are mainly due to human activities (negligence or arson) as seen in
56 California (Syphard and Keeley, 2015; Kolden and Abatzoglou, 2018) or in southeastern France
57 (Ganteaume and Jappiot, 2013) where very few fires are started by lightning strikes (Ganteaume et al.,
58 2013). Likewise, fuel structure and composition control fire spread and, therefore, the location of the
59 largest fires (Duane et al., 2015; Fernandes et al., 2016). The fuel structure is also subject to human
60 activities (Moreira et al., 2011), with agricultural land abandonment or systematic fire suppression
61 leading to the build-up of large amount of fuels (Pausas and Fernández-Muñoz, 2012). Additionally,
62 top-down drivers including fire weather conditions, can help define areas where large fires are most
63 likely to occur (Moritz et al., 2012; Ruffault et al., 2016) but also provide windows of opportunity for
64 fire spread. Large fires in Mediterranean climate ecosystems are often enabled by episodes of severe
65 fire weather of varying duration that can be generated by dry and hot winds as seen in California
66 (Abatzoglou et al., 2013; Kolden and Abatzoglou, 2018) or by cold but dry wind as seen in
67 southeastern France (Ruffault et al., 2016). Collectively, climatic factors alongside ignition sources,

68 fuels, but also suppression forces are thought to influence fire spread. It is noteworthy that changes in
69 fire suppression policy over the last few decades mentioned above have induced sharp decreases in
70 fires, partially modifying the functional relationships linking fire to climate (Fréjaville and Curt, 2017;
71 Syphard et al., 2017), and thus, decreasing fire activity independently of the climate forcing
72 (Hawbaker et al., 2013; Syphard et al., 2007).

73 We focused here on the French Mediterranean, the most fire-prone region of France, where the
74 largest fire on record reached 11,580 ha despite a highly fragmented landscape. This is also a highly
75 populated area characterized by an extensive WUI and high network density which are highly
76 impacted by fire ignitions especially in the western part (Ganteaume and Long-Fournel 2015) with the
77 potential for several consecutive reburns. The region includes plant communities well adapted to
78 Mediterranean climate conditions that confer on this area a high fire risk but an increase in fire
79 recurrence and a shortening of the period between fires were shown to impact vegetation structure,
80 especially with the decrease in mature tree cover (Ganteaume et al. 2009), including the loss of
81 resilience of *Pinus halepensis* stands (Eugenio et al. 2006). It is thus of interest to quantify reburns
82 across the region given their detrimental impacts on ecosystems.

83 Additionally, little attention has been devoted to understanding the spatial distribution of large
84 fires along a longitudinal transect. From a bottom-up perspective, fire prone areas along the
85 Mediterranean coast have been extensively built up in the western part of the region, thereby reducing
86 the availability of fuel while increasing the probability of human-started fires (Ganteaume et al.,
87 2013). From a top-down perspective, climatological annual precipitation is increasing eastwards,
88 gradually lowering the weather-induced fire danger. How these two factors, namely fuel continuity
89 and fire weather, modulate the occurrence of large fires is still unclear.

90 Previous works in the French Mediterranean were based on gridded fire data commencing
91 from the mid-1970s (e.g., Ruffault et al., 2016; Fréjaville and Curt, 2017; Ganteaume and Guerra,
92 2018; Lahaye et al., 2018). Here, we used for the first time longer time-series of georeferenced fires
93 extending back to 1958 and sought to examine both spatial and temporal distributions of large fires
94 (>100 ha) across the French Mediterranean. More specifically, this paper has a three-fold objective.
95 First, we sought to identify the locations associated with large fire recurrence and quantify the spatial
96 extent of the region with reburns. Second, we sought to establish the mean fire extent and the fire
97 return level along a longitudinal transect spanning the French Mediterranean and identify the possible
98 role of climate conditions and fuel continuity in shaping this longitudinal gradient. This exploratory
99 analysis may provide some insights on a fire aspect that was overlooked in previous studies. Finally,
100 building on previous research, we sought to re-estimate trends in large fires across the region taking
101 advantage of a fire record spanning almost six decades.

102

103

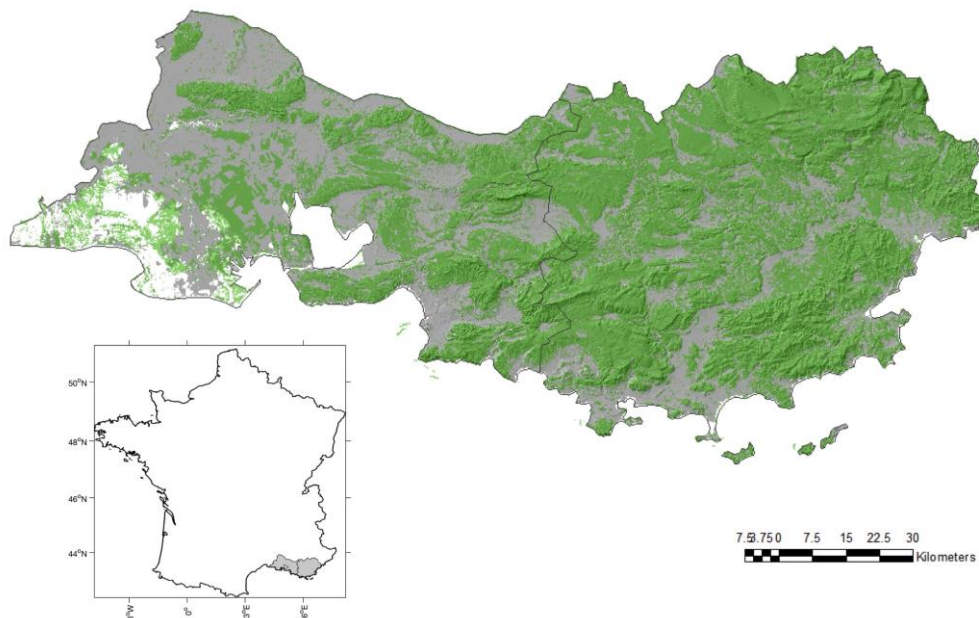
104 2 Material and Methods

105 2.1 Study Area

106 The study area (total surface area of 11 157 km²) is one of the most fire-prone region of SE France in
107 terms of fire frequency (i.e. number of fires) and burned area (Ganteaume and Jappiot, 2013;
108 Ganteaume and Guerra, 2018). The western part is characterized by an extensive WUI where the
109 ignitions are the most frequent (47% of the total ignitions occurred in the WUI) (Ganteaume and
110 Long-Fournel, 2015). Most large fires occur in summer but their cause is often unknown and when it
111 is known, these large fires are mainly due to arson (Ganteaume and Guerra, 2018).

112 The two parts of the study area (Fig. 1), located on a West-East gradient of the Mediterranean,
113 share most climate characteristics albeit the amount of annual precipitation increases eastwards
114 (Ruffault et al., 2017). These areas also differ in the structure of landscapes; forested massifs are larger
115 in the eastern zone while the proportion of WUI and the urbanization are higher in the western area
116 (respectively, 15% vs 7%, Ganteaume unpublished data, and 394 vs 174 inhabitants km⁻²,
117 <https://www.geoportail.gouv.fr>), as well as in the main flammable fuel types, due to the nature of the
118 bedrock (acidic soils being mainly located in the East as opposed to limestone-derived soils in the
119 West). All these differences are hypothesized to affect fire spread and ultimately, the distribution of
120 large fires.

121



122

123 Figure 1: Map of the study area. Forested systems in green were extracted from the “BD Forêt 2014”
124 of the National Geographic Institute (<https://www.geoportail.gouv.fr>).

125

126

2.2 Fire Data

127 Large fires in the French Mediterranean have already been studied in previous works using shorter
128 time series based on the gridded regional fire database Prométhée that recorded fires since 1973
129 (Fréjaville and Curt, 2015; Ruffault and Mouillot, 2017; Ruffault et al., 2018). However, this gridded
130 data provides neither the fire perimeter needed to assess reburns nor the temporal length needed to
131 assess return periods in large fires. Here, we used the georeferenced fire perimeter database compiled
132 by the Office National des Forêts (ONF) and Directions Départementales des Territoires et de la Mer
133 (DDTM Bouches du Rhône and Var) available from 1961 to 2017 in the western part and from 1958
134 to 2016 in the eastern part of the study area. Fire perimeters were derived from aerial photography and
135 remote sensing (the latter since 2016) and confirmed by ground truth targeting mostly fires larger than
136 10 ha in the earliest period. Approximate perimeters of older fire events (i.e., before 1990) have been
137 corrected using aerial photos and Landsat satellite images when available (i.e. a more accurate
138 delineation of fire perimeters adjustment were performed) (Faivre, 2011).

139 We focused on large fires ≥ 100 ha (hereafter LF), representing only 28% of the total number
140 of fires ≥ 1 ha (N=1277) but accounting for 94% of the total burned area. This detection threshold is
141 within the range of thresholds used in previous works in the French Mediterranean ranging from 30 ha
142 (Ruffault and Mouillot, 2017) to 250 ha (Ruffault et al., 2017).

143

144

2.3 Climate and Land Cover Data

145 We computed the daily Fire Weather Index (FWI) from the Canadian Forest Fire Weather Index
146 system using daily surface meteorological variables at a 8-km spatial resolution from the quality-
147 controlled SAFRAN dataset providing maximum temperature, minimum relative humidity,
148 precipitation and wind speed over France from 1959-2017 (Vidal et al., 2009, 2010, 2012). The FWI
149 computation usually requires noon observations. However, given that SAFRAN is a daily
150 meteorological database, we calculated FWI using maximum temperature and minimum relative
151 humidity as surrogates of noon observations following prior analyses (e.g., Jolly et al. 2015;
152 Abatzoglou et al., 2018). Although the FWI was empirically calibrated for estimating whether
153 atmospheric conditions and fuel moisture content are prone to wildfire development in Canada
154 (VanWagner, 1987), the FWI has already proven useful to track large fire in Mediterranean regions
155 (Dimitrakopoulos et al., 2011) including the French Mediterranean (Barbero et al., 2019). Grid cells of
156 the FWI lying within the study area were first averaged across the June-September season and then
157 averaged across all latitudes spanning the region of interest to form a longitudinal cross-section of
158 mean summer FWI conditions.

159

160

We extracted fuel cover data from the “BD Forêt 2014” of the National Geographic Institute
(<https://www.geoportail.gouv.fr>) and regrided the data onto 8-km spatial grid. The percentage of land

161 area covered by forest types was computed across all latitudes spanning the region of interest to form a
162 longitudinal cross-section as described above.

163

164 **2.4 Spatial Analyses**

165 Based on a sequence of 58 layers of annual large fire scars covering the 1958–2017 period, the
166 following fire attributes were extracted: (i) fire frequency or the number of fires that occurred on a
167 same location over the period studied and (ii) time since the last fire.

168

169 **2.5 Temporal Analyses**

170 Monotonic trends in LF frequency and in burned area due to LF were assessed using the non-
171 parametric Mann-Kendall test (Kendall, 1975) and a change point detection test (Standard Normal
172 Homogeneity Test (SNHT); Alexandersson and Moberg, 1997) was used to identify potential abrupt
173 changes in the time series.

174 We estimated annual maximum burned area (AMBA) return levels in the eastern and western
175 part of the study area using the so-called block (here 1-year) maxima approach. We extracted the
176 AMBA in both areas and selected the type of distribution that best fitted both series using the Akaike
177 Information Criteria (AIC). In both areas, the gamma distribution was found to best describe the
178 AMBA series. Using this distribution, the inverse cumulative distribution was calculated allowing the
179 determination of the theoretical quantiles from which we derived the return levels (AMBA) associated
180 to different return periods ranging from 5 to 100 years. Asymmetric confidence intervals were
181 calculated using a resampling approach. This approach consists in creating new sub-samples from the
182 original sample (75% of the original sample are extracted at random) using a bootstrapping process
183 with replacement and then estimating a return level for each of the resampled data (N=1000). The
184 resulting empirical distribution can then be used to derive the 95% confidence intervals from the
185 resulting collection of estimates.

186

187

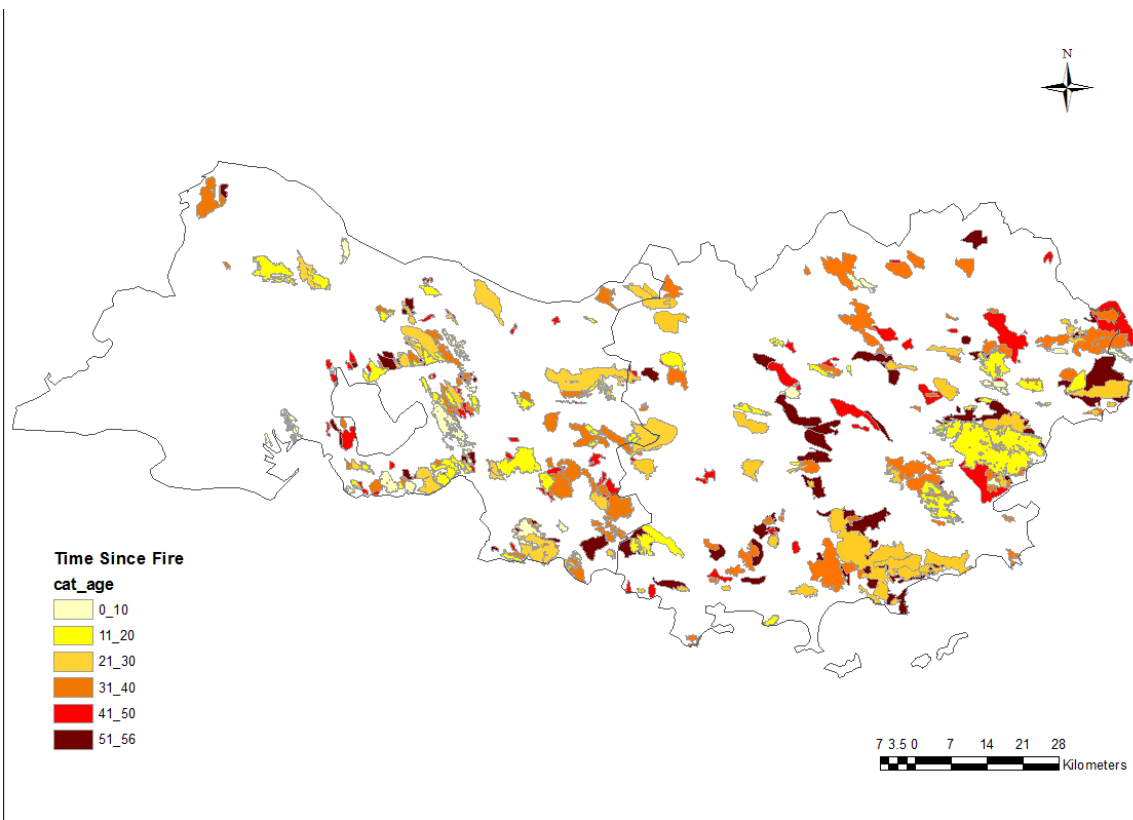
188 **3 Results**

189 **3.1 Spatial distribution of large fires and reburned areas**

190 In total, 353 LF were recorded in the region between 1958 and 2017 (194 in the western part and 159
191 in the eastern part) with, however, a higher burned area in the East nearly doubling the area burned in
192 the West (respectively, 199,404 and 112,043 ha representing 3379.7 and 2000.8 ha burned per year;
193 Tab. 1). LF were responsible for most of the total burned area in the East (97%) as well as in the West
194 (87%), which supports the relevance of the fire-size threshold selected (100 ha).

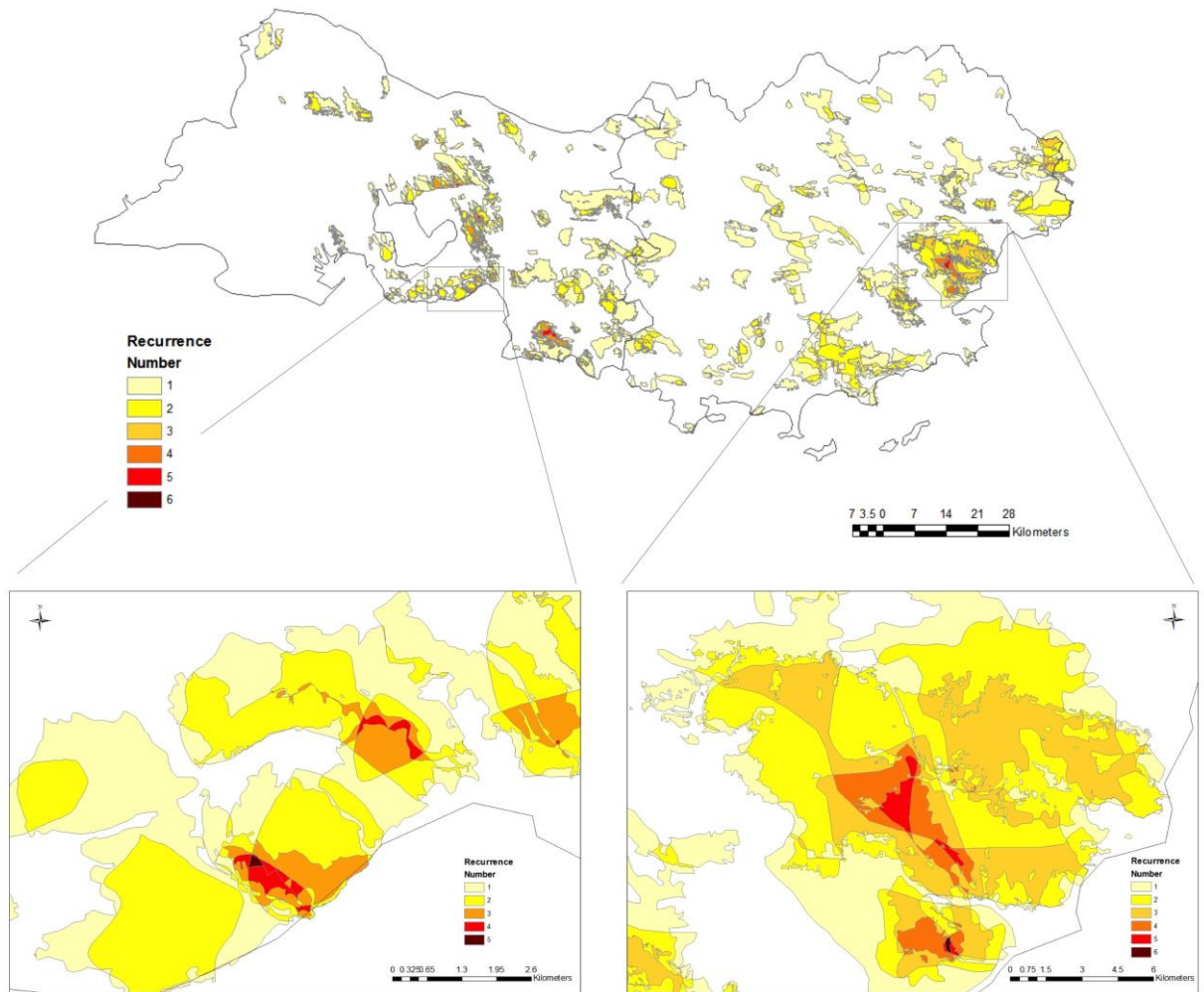
195 Regarding the LF age distribution (Fig. 2), the most frequent LF belonged to the 31-40 year-
196 class resulting in the most LF-prone decade. In the East, recent LF were mainly located on the coast
197 while the age distribution was more homogeneous in the western part. Notice that most LF growths
198 were in the main wind direction blowing from Northwest. A total surface area of 312,447 ha was
199 burned during the period studied of which 21% occurred on a surface that already burned in the past
200 (Fig. 3), due to multiple overlaps in burned areas by recurrent fires (i.e. LF occurrence on the same
201 surface). LF reburns occurred up to 6 times in the East but represented only a small part of the
202 recurrence (0.3%; Tab. 2). One to two reburns were the most frequent patterns in the western part of
203 the study area (39.4 and 39.9% of the recurrence, respectively; Tab. 2) while in the East, most reburns
204 occurred only once (46.3%). The surface impacted by only one LF represented 74.5% and 71.2% of
205 the total area burned by LF in the West and the East, respectively (Tab. 2).

206



207

208 Figure 2: Time since the last LF (cat_age in years).



209

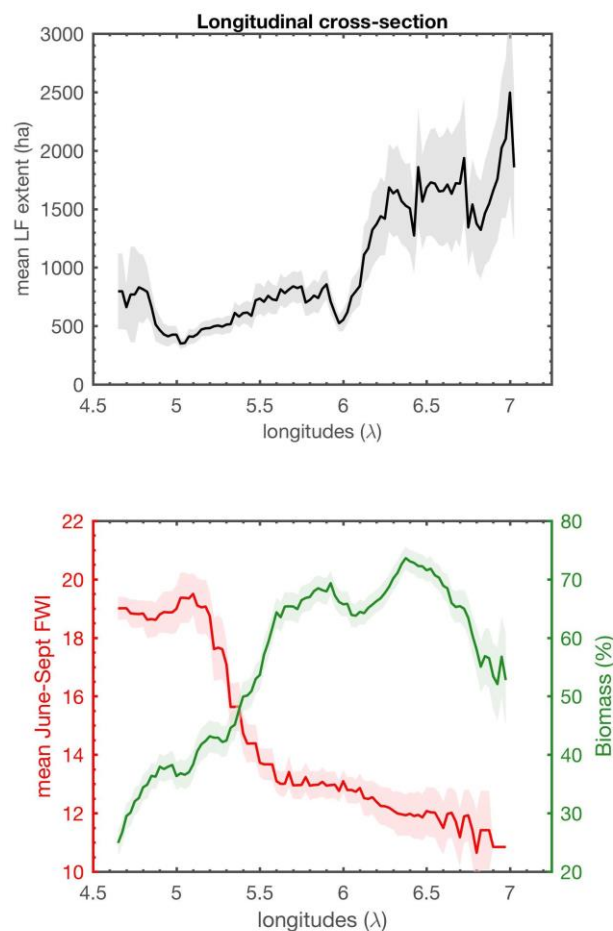
210 Figure 3: Fire recurrence on the 1961-2017 and 1958-2016 period in the western and eastern part,

211 respectively.

212

3.2 Longitudinal contrast in large fire extent

213 The mean LF extent varied along a longitudinal gradient, increasing from the West to the East
214 (Fig. 4 top). This signal contrasts with the mean summer FWI gradient decreasing towards the East but
215 is consistent with the sharp increase in biomass towards the East (Fig. 4 bottom). This suggests that LF
216 spread is not limited by climate conditions across the region but strongly fuel-limited in the West, due
217 to landscape fragmentation and the high proportion of WUI. Indeed, the landscape has undergone
218 substantial transformation with time in the western part contributing to reduce fuel cover and thereby
219 mean LF extent. This highlights the role of fuel continuity on fire spread as shown in previous
220 research (Hargrove et al., 2000; Finney et al., 2007) and the need to include fuel cover in future
221 projections of fire activity based on fire weather indices only.



222

223 Figure 4: Top) Longitudinal cross-section of mean LF extent computed over 30-km sliding windows.
224 The 95% confidence intervals were estimated using a bootstrapping approach. Bottom) Same as top
225 panel but for mean June-September FWI (in red) and the percent of biomass (in green).

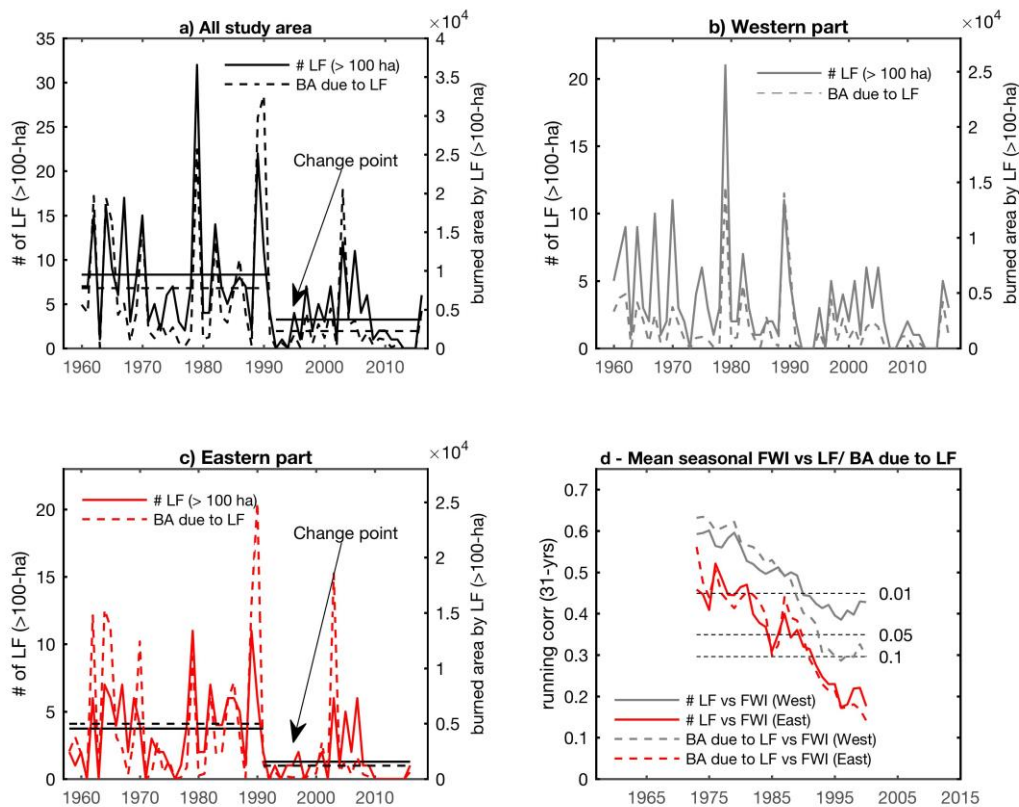
226

227

3.3 Long-term trends in large fires

228 A significant decline in annual LF frequency alongside area burned by LF was found across the region
 229 according to a Man-Kendall test (Fig. 5). This overall decline is consistent with a significant change
 230 point in both LF metrics in 1991 as shown in previous findings (Fox et al., 2015; Ruffault and
 231 Mouillot, 2015). This signal was especially evident in the eastern part (Fig. 5c) while neither a change
 232 point nor a significant trend ($p > 0.05$) were detected in the western part for both LF metrics (Fig. 5b).
 233 We then examined how interannual correlations between mean June-September FWI and LF activity
 234 have changed over time across both regions (Fig. 5d). Higher correlations prevailed in the western part
 235 throughout the period but the relationships strongly weakened with time in both regions in agreement
 236 with previous findings (Ruffault and Mouillot, 2015), passing below significance levels across recent
 237 years.

238

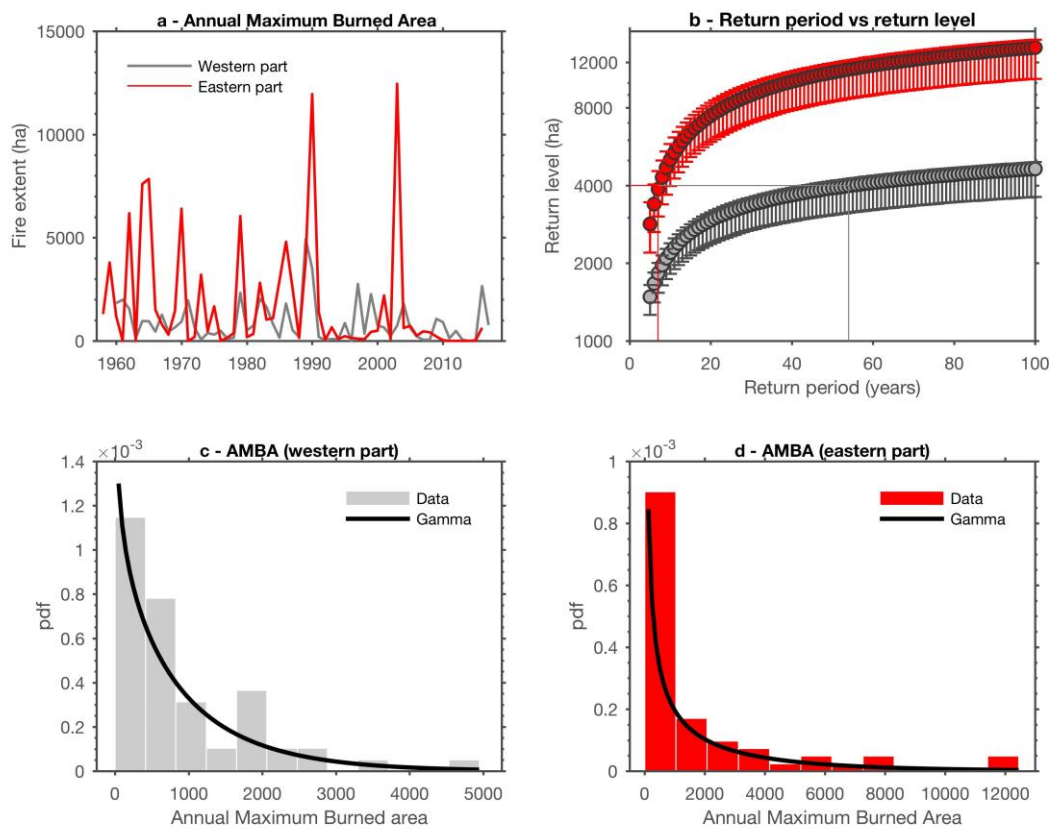


239 Figure 5: a) Annual number of LF (in black) and area burned by LF (in red) across the region.
 240 Significant change points at the 5% confidence level according to a Standard Normal Homogeneity
 241 Test (SNHT) in both metrics are indicated. Horizontal solid lines indicate the overall mean observed
 242 before and after the change point. b) Same as a) but for the western part. c) Same as a) but for the
 243 eastern part. d) Sliding correlations on 31-year windows between mean June-September FWI and
 244 annual LF frequency (solid lines) and annual burned area due to LF (dashed lines) in the western
 245

246 (gray) in eastern part (red). The horizontal dashed lines indicate different significance levels of the
 247 Pearson correlations. Correlations are indicated for the middle of the sliding windows.

248

249 Figure 6 shows the AMBA in each part of the study area (panel a) as well as the Gamma
 250 distribution models that were found as the best fit to the data (panels c,d). Estimates of AMBA return
 251 intervals show that a LF >4000 ha occurs on average every 7 years in the eastern part and every 55
 252 years in the western part (Figure 6b), supporting results of Figure 4 indicating an overall increase in
 253 LF extent eastwards.



254

255 Figure 6: a) Time series of the annual maximum burned area (AMBA) in the western part (in gray)
 256 and in the eastern part (in red). b) Return levels in AMBA in the western part (in gray) and in the
 257 eastern part (in red) for different return periods ranging from 5 to 100 years. The 95% confidence
 258 intervals were estimated using a bootstrapping approach. c) Distribution of AMBA in the western part
 259 (bars) with the gamma distribution (black line) that was found to best describe the data. d) Same as c
 260 but for the eastern part.

261

262

263 4 Discussion

264 Improving our understanding of large fire activity is of utmost importance to fire prevention and
265 management to mitigate their impacts. Here, we presented a comprehensive analysis of spatial and
266 temporal patterns of LF in the French Mediterranean. To our knowledge, the fire database compiled
267 and analysed in this framework provides for the first time a detailed description of LF recorded on
268 geo-referenced long time series.

269

270 4.1 Spatial distribution of large fires and reburned areas

271 In total, 21% of the burned area occurred on a surface that already burned in the past due to multiple
272 overlaps in burned areas by recurrent fires (up to 6 times in the East). These areas of higher recurrence
273 could induce a loss of resilience of the forest types such as *Pinus halepensis* stands with an increase in
274 the number of fires and/or a decrease in the time-since-fire (Eugenio et al., 2006). Results showed that
275 there was a strong spatial variation of LF according to the time-since-fire, with clusters of recent LF
276 along the coast (where the recurrence was the highest) and more ancient LF in the central and northern
277 part of the eastern area where the tourist pressure is lower. In contrast, LF were homogeneously
278 distributed in the West, regardless of their age and most reburns corresponded to WUI areas.

279 We found that the return level was higher in the eastern part of the study area although LF
280 were more frequent in the West. These contrasted regional return levels may provide critical and
281 useful information for risk assessment and local decision-making. Indeed, LF >4000 ha may occur
282 within seven years in the East against 55 years in the West. In other words, LF are less probable in the
283 east where fire ignitions are more limited but when an ignition does occur, the fire is likely to spread
284 over larger areas. This longitudinal gradient is likely due to the variation in landscape fragmentation.
285 Indeed, the western area presents a mosaic of wildlands interspersed with agricultural areas and WUI,
286 LF being thereby concentrated in natural spaces less extended than in the eastern part where large
287 forested massifs mostly located on the coast allowed fire spread. By contrast, LF were more frequent
288 in the West where population density, the proportion of WUI, and infrastructures (railroads and roads)
289 are the highest, as shown in previous works (Keane et al., 2008; La Puma, 2012; Alexandre et al.,
290 2016; Nagy et al., 2018). Fox et al. (2015) showed that, in an area located East of our study area,
291 neither WUI characteristics (despite the 60% increase between 1964 and 2009 in this area) nor fire
292 weather were major drivers of fire frequency and burned area, the climate control becoming less
293 important as the fire regime shifted to more frequent human-started fires (Zumbrunnen et al., 2009).

294 Some recent studies across Euro-Mediterranean countries emphasized that large fire
295 preferentially occurred under specific synoptic patterns associated with high temperature (Pereira et
296 al., 2005; Trigo et al., 2013; Hernandez et al., 2015). In southern France, large fires were also
297 facilitated by wind events blowing from Northwest (Ruffault and Mouillot, 2015, 2017). The shapes of

298 LF which were more elongated in the wind direction in the western part support the results of Ruffault
299 et al. (2018) pinpointing that the main wind-driven large fires that had occurred in 2016 were located
300 in the western part while the main heat-driven large fires that occurred in 2003 were located in the
301 East of the area. Taking into account other metrics describing the LF patch complexity (e.g. azimuthal
302 angle or shape index) as in Laurent et al. (2018) could allow deriving additional information on the
303 role of wind on their geometry or on the fraction of LF driven by wind.

304

305 **4.2 Long-term trends in large fires**

306 The overall reduction in both LF frequency and burned area observed over the last 6 decades is in
307 agreement with previous works that highlighted a decrease in fire activity across parts of southern
308 Europe in response to an increased effort in fire suppression (Turco et al., 2016), taking place in early
309 1990s in the French Mediterranean (Ruffault and Mouillot, 2015, Fox et al., 2015; Curt and Fréjaville,
310 2017). Indeed, the region was highly impacted by fires during the 1970-1990 period and developed a
311 thorough fire suppression and prevention system in the beginning of the 1990s, allocating more means
312 for fire management that allowed faster reactivity in case of fire start (the strategy became
313 extinguishing the fires at their initial stage by massive attack to prevent their spread). The decrease in
314 both LF frequency and burned area since 1991, especially evident in the eastern part of the region, is
315 likely due to this change in firefighting policy and fire prevention regulations (fire suppression could
316 be more intense in the East as fires were historically larger in that region).

317 Climate projections suggest that atmospheric conditions conducive to large fire will increase in the
318 future. Indeed, the warming and drying trends projected in southern Europe are expected to facilitate
319 fire spread (Turco et al., 2018), at least where fuel and ignitions are not limiting. This trend towards
320 more extreme fire weather conditions is likely to overcome prevention efforts in the French
321 Mediterranean (Lahaye et al., 2018), a region where expanding forests (Abadie et al. 2017) are
322 increasing fuel loading and may offer opportunities for future fire spread.

323

324

325 **5 Conclusions**

326 This work, based on long-term geo-referenced fire time series (1958-2017) analysed both spatial
327 and temporal variations of LF throughout one of the most impacted areas of the French Mediterranean.
328 On the whole, 21% of the total area burned by LF occurred on a surface that already burned in the
329 past, the region being impacted in some locations up to 6 times by recurrent LF (coastal areas of the
330 eastern part of the study area). LF were less frequent in the eastern part but larger than LF occurring in
331 the West mostly in WUI. This longitudinal gradient in LF extent, featuring a shorter time of

332 occurrence between LF in the East with respect to the West, contrasts with what we would expect
333 from mean fire weather conditions strongly decreasing eastwards but is consistent with larger fuel
334 cover in the East. Indeed, fuel continuity in the East allows fire to grow large and to reach on average
335 4,000 ha every 7 years, a spatial extent in burned area observed only every 50 years in the West.

336 An abrupt decline in LF was evident across the eastern part in the early 1990s, mostly due to a
337 change in fire management policy thereby contributing to the weakening of the climate-fire
338 relationship. However, despite large means allocated to fire suppression, large fire outbreak is still
339 possible in the French Mediterranean (such as in 2003 or 2016), as specific weather conditions can
340 overwhelm fire suppression efforts (Fernandes et al., 2016; Lahaye et al., 2018). A better knowledge
341 of LF drivers is necessary to strengthen fire prevention by providing valuable information on priority
342 areas where LF are more likely to occur.

343

344 *Acknowledgements.* The authors wish to thank Adeline Bellet and Denis Morge, for preprocessing the
345 data with ArcGis. The authors also sincerely thank Aimee Mac Cormack for English revision. The
346 authors also thank two anonymous reviewers for their constructive comments and suggestions that
347 helped improved the quality of the manuscript.

348

349

350 **References**

351 Abadie, J., Dupouey, J.-L., Avon, C., Rochel, X., Tatoni, T. and Bergès, L., Forest recovery since
352 1860 in a Mediterranean region: drivers and implications for land use and land cover spatial
353 distribution. *Landscape Ecology*, <https://doi.org/10.1007/s10980-017-0601-0>, 2017.

354 Abatzoglou, J.T., Williams, A.P., Boschetti, L., Zubkova, M., Kolden, C.A.: Global patterns of
355 interannual climate-fire relationships, *Global Change Biology*, 24 (11), 5164-5175, 2018.

356 Abatzoglou, J.T. and Williams, A.P.: Impact of anthropogenic climate change on wildfire across
357 western US forests, *Proc Natl Acad Sci USA*, 113, 11770–11775, 2016.

358 Abatzoglou, J.T., Barbero, R. and Nauslar, N.J.: Diagnosing Santa Ana winds in southern California
359 with synoptic-scale analysis, *Weather and Forecasting*, 28, 704-710, doi:10.1175/WAF-D-13-00002.1,
360 2013.

361 Alexandersson, H. and Moberg, A.: Homogenization of Swedish temperature data. Part I:
362 Homogeneity test for linear trends, *International Journal of Climatology*, 17(1), 25-34, 1997.

363 Alexandre, P.M., Stewart, S.I., Keuler, N.S., Clayton, M.K., Mockrin, M.H., Bar-Massada, A.,
364 Syphard, A.D. and Radeloff, V.C.: Factors related to building loss due to wildfires in the conterminous
365 United States, *Ecol. Appl.*, 26, 2323–2338, 2016.

366 Barbero, R., Curt, T., Ganteaume, A., Maillé, E., Jappiot, M., and Bellet, A.: Simulating the effects of
367 weather and climate on large wildfires in France, *Nat. Hazards Earth Syst. Sci. Discuss.*,
368 <https://doi.org/10.5194/nhess-2018-283>, in review, 2018.

369 Barbero, R., Abatzoglou, J.T., Larkin, N.K., Kolden, C.A. and Stocks, B.: Climate change presents
370 increased potential for very LF in the contiguous United States, *Int. J. Wildland Fire*, 24, 892–899.
371 2015.

372 Barbero, R., Abatzoglou, J.T., Kolden, C., Hegewisch, K., Larkin, N.K. and Podschwit, H.: Multi-
373 scalar influence of weather and climate on very large-fires in the eastern United States, *Int. J.*
374 *Climatol.* doi:10.1002/joc.4090, 2014a.

375 Barbero, R., Abatzoglou, J.T., Steel, E.A. and Larkin, N.K.: Modeling very large-fire occurrences over
376 the continental United States from weather and climate forcing, *Environ. Res. Lett.*, 9, 124009, 2014b.

377 Blanchi, R., Leonard, J., Haynes, K., Opie, K., James, M. and Dimer de Oliveira, F.: Environmental
378 circumstances surrounding bushfire fatalities in Australia 1901–2011, *Environmental Science and*
379 *Policy*, 37, 192-203, 2014.

380 Bradstock, R.A., Cohn, J.S., Gill, A.M., Bedward, M. and Lucas, C.: Prediction of the probability of
381 LF in the Sydney region of south-eastern Australia using fire weather, *Int. J. Wildland Fire*, 18, 932-
382 943, 2009.

383 Curt, T. and Fréjaville, T.: Wildfire Policy in Mediterranean France: How Far is it Efficient and
384 Sustainable? Risk analysis, DOI: 10.1111/risa.12855, 2017.

385 Dennison, P.E., Brewer, S.C., Arnold, J.D. and Moritz, M.A.: Large wildfire trends in the western
386 United States, 1984–2011, *Geophys. Res. Lett.*, 41, 2928–293, 2014.

387 Dimitrakopoulos, A.P., Bemmerzouk, A.M. and Mitsopoulos, I.D.: Evaluation of the Canadian fire
388 weather index system in an eastern Mediterranean environment, *Meteorological Applications*, 18(1),
389 83–93, <https://doi.org/10.1002/met.214>, 2011.

390 Duane, A., Piqué, M., Castellnou, M. and Brotons, L.: Predictive modelling of fire occurrences
391 from different fire spread patterns in Mediterranean landscapes, *Int. J. Wildland Fire*, 24, 407–418,
392 doi:10.1071/WF14040, 2015.

393 Eugenio, M., Verkaik, I., Lloret, F. and Espelta, J.M.: Recruitment and growth decline in *Pinus*
394 *halepensis* populations after recurrent wildfires in Catalonia (NE Iberian Peninsula), *Forest Ecology*
395 *and Management*, 231(1-3), 47-54, 2006.

396 Faivre, N.: Which pyrodiversity for what biodiversity? A multi-scale comparative study of two
397 Mediterranean ecosystems. Ph.D. Dissertation, Aix-Marseille University, France, 2011.

398 Fernandes, P.M., Pacheco, A.P., Almeida, R. and Claro, J.: The role of fire suppression force in
399 limiting the spread of extremely large forest fires in Portugal, *European Journal of Forest Research*
400 135, 253–262, doi:10.1007/S10342-015-0933-8, 2016.

401 Flannigan, M.D., Krawchuk, M.A., de Groot, W.J., Wotton, B.M., and Gowman, L.M.: Implications of
402 changing climate for global wildland fire, *Int. J. Wildland Fire*, 18, 483–507, 2009.

403 Fox, D.M., Martin, N., Carrega, P., Andrieu, J., Adnès, C., Emsellem, K., Ganga, O., Moebius, F.,
404 Tortorollo, N. and Fox, E.A.: Increases in fire risk due to warmer summer temperatures and wildland
405 urban interface changes do not necessarily lead to more fires, *Appl. Geogr.*, 56, 1–12,
406 <https://doi.org/10.1016/j.apgeog.2014.10.001>, 2015.

407 Fréjaville, T. and Curt, T.: Spatiotemporal patterns of changes in fire regime and climate: defining the
408 pyroclimates of south-eastern France (Mediterranean Basin), *Clim. Change*, 129, 239–51, 2015.

409 Fréjaville, T. and Curt, T.: Seasonal changes in the human alteration of fire regimes beyond the
410 climate forcing, *Environ. Res. Lett.* 12, 035006. doi:10.1088/1748-9326/AA5D23, 2017.

411 Ganteaume, A. and Jappiot, M.: What causes LF in Southern France. *Forest Ecology and*
412 *Management*, 294, 76-85, DOI 10.1016/j.foreco.2012.06.055, 2013.

413 Ganteaume, A. and Long-Fournel, M.: Driving factors of fire density can spatially vary at the local
414 scale in SE France, *Int. J. Wildland Fire*, 24(5), 650-664, 2015.

415 Ganteaume, A. and Guerra, F.: Explaining the spatio-seasonal variation of fires by their causes: The
416 case of southeastern France, *Appl. Geography*, 90, 69-81, 2018.

417 Ganteaume, A., Jappiot, M., Lampin-Maillet, C., Curt, T. and Borgniet, L.: Fuel characterization and
418 effects of wildfire on limestone soils in southeastern France, *For. Ecol. Manag.*, 258S, S15-S23, 2009.

419 Ganteaume, A., Camia A., Jappiot M., San Miguel-Ayanz J., Long-Fournel, M. and Lampin C.: A
420 Review of the Main Driving Factors of Forest Fire Ignition Over Europe, *Environ. Manag.*, 51 (3),
421 651-662, 2013.

422 Hawbaker, T.J., Radeloff, V.C., Stewart, S.I., Hammer, R.B., Keuler, N.S. and Clayton, M.K.: Human
423 and biophysical influences on fire occurrence in the United States, *Ecol. Appl.*, 23, 565–82, 2013.

424 Hernandez, C., Drobinski, P. and Turquety, S.: How much does weather control fire size and intensity
425 in the Mediterranean region? *Ann. Geophys.*, 33, 931–939, 2015.

426 Jolly, W.M., Cochrane, M.A., Freeborn, P.H., Holden, Z.A., Brown, T.J., Williamson, G.J., and
427 Bowman, D.M.J.S.: Climate-induced variations in global wildfire danger from 1979 to 2013, *Nature*
428 *Communications*, 6, 7537, 2015.

429 Keane, R.E., Agee, J.K., Fulé, P., Keeley, J.E., Key, C., Kitchen, S.G., Miller, R. and Schulte, L.A.:
430 Ecological effects of large-fires on US landscapes: benefit or catastrophe? *Int. J. Wildland Fire*, 17,
431 696–712, 2008.

432 Kendall, M.G.: *Rank Correlation Methods*, 4th ed., Charles Griffin, London, 1975.

433 Kolden, C.A. and Abatzoglou, J.T.: Spatial Distribution of Wildfires Ignited under Katabatic versus
434 Non-Katabatic Winds in Mediterranean Southern California USA, *Fire*, 1, 19, 2018.

435 Lahaye, S., Curt, T., Fréjaville, T., Sharples, J., Paradis, L. and Hély, C.: What are the drivers of
436 dangerous fires in Mediterranean France? *Int. J. Wildland Fire*, 27, 155–163, 2018.

437 La Puma, I.P.: *Fire in the pines: A landscape perspective of human-induced ecological change in the*
438 *pinelands of New Jersey*, Ph.D. Dissertation, Rutgers University, New Brunswick, NJ, USA, 2012.

439 Littell, J.S., McKenzie, D., Peterson, D.L. and Westerling, A.L.: Climate and wildfire area burned in
440 western US ecoprovinces 1916–2003, *Ecological Applications*, 19, 1003–1021, 2009.

441 McKenzie, D., Gedalof, Z., Peterson, D.L., and Mote, P.: Climatic change, wildfire, and conservation,
442 *Conservation Biology*, 18, 890–902, 2004.

443 Miller, J.D., Safford, H.D., Crimmins, M. and Thode, A.E.: Quantitative evidence for increasing forest
444 fire severity in the Sierra Nevada and southern Cascade Mountains, California and Nevada, USA,
445 *Ecosystems*, 12, 16–32, 2009.

446 Moreira, F., Viedma, O., Arianoutsou, M., Curt, T., Koutsias, N., Rigolot, E., Barbati, A., Corona, P.,
447 Vaz, P., Xanthopoulos, G., Mouillot, F., and Bilgili, E.: Landscape–wildfire interactions in southern
448 Europe: implications for landscape management, *J. Environ. Manage.*, 92, 2389–2402, 2011.

449 Moreno, M.V., Conedera, M., Chuvieco, E. and Pezzatti, G.B.: Fire regime changes and major driving
450 forces in Spain from 1968 to 2010, *Environ. Sci. Policy*, 37, 11–22, 2014.

451 Moritz, M.A.: Analyzing extreme disturbance events: fire in Los Padres National Forest, *Ecological*
452 *Applications* 7, 1252–1262, doi:10.1890/1051-0761(1997)007[1252:AEDEFI]2.0.CO;2, 1997.

453 Moritz, M.A., Parisien, M-A., Batllori, E., Krawchuk, M.A., Van Dorn, J., Ganz, D.J. and Hayhoe, K.:
454 Climate change and disruptions to global fire activity, *Ecosphere*, 3, 49, 2012.

455 Mouillot, F. and Field, C.B.: Fire history and the global carbon budget: a 1 degrees x 1 degrees fire
456 history reconstruction for the 20th century, *Global Change Biology*, 11, 398–420, 2005.

457 Nagy, R.C., Fusco, E., Bradley, B., Abatzoglou, J.T. and Balch, J.: Human-Related Ignitions Increase
458 the Number of Large Wildfires across U.S. Ecoregions, *Fire*, 1, 4, 2018.

459 Pausas, J.G. and Fernandez-Munoz, S.: Fire regime changes in the Western Mediterranean Basin:
460 From fuel-limited to drought-driven fire regime, *Clim. Chang.*, 110, 215–226, 2012.

461 Pereira, M.G., Trigo, R.M., da Camara, C.C., Pereira, J. and Leite, S.M.: Synoptic patterns associated
462 with large summer forest fires in Portugal, *Agric. For. Meteorol.*, 129, 11–25, 2005.

463 Pezzatti, G.B., Zumbunnen, T., Bürgi, M., Ambrosetti, P. and Conedera, M.: Fire regime shifts as a
464 consequence of fire policy and socio-economic development: an analysis based on the change point
465 approach, *For. Policy Econ.*, 29, 7–18, 2013.

466 Radeloff, VC, Helmers, D.P., Kramer, H.A., Mockrin, M.H., Alexandre, P.M., Bar-Massada, A.,
467 Butsic, V., Hawbaker, T.J., Martinuzzi, S., Syphard, A.D. and Stewart, S.I., Rapid growth of the US
468 wildland-urban interface raises wildfire risk, *PNAS*, <https://doi.org/10.1073/pnas.1718850115>, 2018.

469 Ruffault, J. and Mouillot, F.: How a new fire-suppression policy can abruptly reshape the fire-weather
470 relationship, *Ecosphere*, 6, 199, <https://doi.org/10.1890/ES15-00182.1>, 2015.

471 Ruffault, J. and Mouillot, F.: Contribution of human and biophysical factors to the spatial distribution
472 of forest fire ignitions and large wildfires in a French Mediterranean region, *Int. J. Wildl. Fire*, 26,
473 498–508, <https://doi.org/10.1071/WF16181>, 2017.

474 Ruffault, J., Moron, V., Trigo, R. M. and Curt, T.: Objective identification of multiple large fire
475 climatologies: an application to a Mediterranean ecosystem, *Environ. Res. Lett.*, 11, 75006,
476 <https://doi.org/10.1088/1748-9326/11/7/075006>, 2016.

477 Ruffault, J., Moron, V., Trigo, R. M. and Curt, T.: Daily synoptic conditions associated with large fire
478 occurrence in Mediterranean France: evidence for a wind-driven fire regime, *Int. J. Climatol.*, 37, 524–
479 533, <https://doi.org/10.1002/joc.4680>, 2017.

480 Ruffault, J., Curt, T., Martin-StPaul, N.K., Moron, V. and Trigo, R.M.: Extreme wildfire events are
481 linked to global-change-type droughts in the northern Mediterranean, *Nat. Hazards Earth Syst. Sci.*,
482 18, 847–856, <https://doi.org/10.5194/nhess-18-847-2018>, 2018.

483 San-Miguel-Ayanz, J., Moreno, J.M. and Camia, A.: Analysis of LF in European Mediterranean
484 landscapes: Lessons learned and perspectives, *Forest Ecology and Management*, 294, 11-22, 2013.

485 Shvidenko, A.Z. and Nilsson, S.: Extent, distribution, and ecological role of fire in Russian forests, In
486 ‘Climate Change, and Carbon Cycling in the Boreal Forest’, *Ecological Studies*, 138 (Eds ES
487 Kasischke, BJ Stocks), pp. 132–150 (Springer-Verlag: Berlin), 2000.

488 Stephens, S.L., Burrows, N., Buyantuyev, A., Gray, R.W., Keane, R.E., Kubian, R., Liu, S., Seijo, F.,
489 Shu, L. and Tolhurst, K.G.: Temperate and boreal forest mega-fires: Characteristics and challenges,
490 *Front. Ecol. Environ.*, 12, 115–122, 2014.

491 Stocks, B.J., Mason, J.A., Todd, J.B., Bosch, E.M., Wotton, B.M., Amiro, B.D., Flannigan, M.D.,
492 Hirsch, K.G., Logan, K.A., Martell, D.L. and Skinner, W.R.: Large forest fires in Canada, 1959–1997,

493 Journal of Geophysical Research–Atmospheres, 107, FFR 5-1–FFR 5-12, doi:10.1029/2001JD000484,
494 2002.

495 Syphard, A.D., Clarke, K.C. and Franklin, J.: Simulating fire frequency and urban growth in southern
496 California coastal shrublands, *Landscape Ecology*, 22, 431–445, doi:10.1007/S10980-006-9025-Y,
497 2007.

498 Syphard, A.D. and Keeley, J.E.: Location, timing and extent of wildfire vary by cause of ignition, *Int.*
499 *J. Wildland Fire*, 24(1), 27-36, doi.org/10.1071/WF14024, 2015.

500 Syphard, A.D., Keeley, J.E., Pfaff, A.H. and Ferschweiler, K.: Human presence diminishes the
501 importance of climate in driving fire activity across the United States, *Proc. Natl. Acad. Sci. USA*,
502 114, 13750, 2017.

503 Tedim, F., Leone, L., Amraoui, M., Bouillon, C., Coughlan, M.R., Delogu, G.M., Fernandes, P.M.,
504 Ferreira, C., McCaffrey, S., McGee, T.K., Parente, J., Paton, D., Pereira, M.G., Ribeiro, L.M., Viegas,
505 D.X. and Xanthopoulos, G.: Defining Extreme Wildfire Events: Difficulties, Challenges, and Impacts,
506 *Fire*, 1, 9, doi:10.3390/fire1010009, 2018.

507 Trigo, R.M., Sousa, P.M., Pereira, M.G., Rasilla, D. and Gouveia, C.M.: Modelling wildfire activity in
508 Iberia with different atmospheric circulation weather types, *Int. J. Climatol.*, doi: 10.1002/joc.3749,
509 2013.

510 Turco, M., Rosa-Cánovas, J.J., Bedia, J., Jerez, S., Montávez, J.P., Llasat, M.C., Provenzale, A.:
511 Exacerbated fires in Mediterranean Europe due to anthropogenic warming projected with
512 nonstationary climate-fire models. *Nature Communications*, 9:3821, doi:10.1038/s41467-018-06358-
513 z, 2018.

514 Turco, M., Bedia, J., Di Liberto, F., Fiorucci, P., von Hardenberg, J., Koutsias, N., Llasat, M.C.,
515 Xystrakis, F. and Provenzale A.: Decreasing Fires in Mediterranean Europe, *PLoS one*, 11, e0150663,
516 2016.

517 Turco, M., Llasat, M.C., Tudela, A., Castro, X. and Provenzale, A.: Brief communication Decreasing
518 fires in a Mediterranean region (1970-2010, NE Spain), *Nat. Hazards Earth Syst. Sci.*, 13(3), 649–652,
519 doi: 10.5194/nhess-13-649-2013, 2013.

520 Van Wagner, C.E.: Development and structure of the Canadian forest fire weather index system,
521 *Forestry*, <https://doi.org/19927>, 1987.

522 Vidal, J.-P., Martin, E., Franchistéguy, L., Habets, F., Soubeyroux, J.-M., Blanchard, M. and Baillon,
523 M.: Multilevel and multiscale drought reanalysis over France with the Safran-Isba-Modcou
524 hydrometeorological suite, *Hydrology and Earth System Sciences Discussions*, 6(5), 6455–6501,
525 <https://doi.org/10.5194/hessd-6-6455>, 2009.

526 Vidal, J.-P., Martin, E., Franchistéguy, L., Baillon, M. and Soubeyroux, J.-M. : A 50-year high-
527 resolution atmospheric reanalysis over France with the Safran system, *International Journal of*
528 *Climatology*, 30(11), 1627–1644, <https://doi.org/10.1002/joc.2003>, 2010.

529 Vidal, J.-P., Martin, E., Kitova, N., Najac, J. and Soubeyroux, J.-M.: Evolution of spatio-temporal
530 drought characteristics: Validation, projections and effect of adaptation scenarios, *Hydrology and*
531 *Earth System Sciences*, 16(8), 2935–2955, <https://doi.org/10.5194/hess-16-2935-2012>, 2012.

532 Westerling, A.L., Hidalgo, H.G., Cayan, D.R. and Swetnam, T.W.: Warming and earlier spring
533 increase Western US forest wildfire activity, *Science*, 313, 940–943, 2006.

534 Zumbrunnen, T., Bugmann, H., Conedera, M. and Bürgi, M.: Linking forest fire regimes and
535 climate—a historical analysis in a dry inner Alpine valley *Ecosystems*, 12, 73–86, 2009.

536

537

TABLES

538 Table 1: Statistics on fires (≥ 1 ha) and LF (≥ 100 ha)

Study area	Total number of fires	Total burned area (ha)	Number of LF:	%	Area burned by LF (ha)	%	Record length (years)
West	975	128 196	194	20	112 043	87	56
East	302	204 535	159	52	199 404	97	58
Total	1277	332 731	353	28	312 447	94	

539

540 Table 2: Percentages of burned area (relative to the total burned area) affected by recurrent LF and
 541 percentages of recurrence relative to the LF frequency (when number=1, LF is considered as non-
 542 recurrent).

Number of LF on same location	Western part		Eastern part	
	Area burned by recurrent LF	Frequency	Area burned by recurrent LF	Frequency
1	74.5%	39.4%	71.2%	46.3%
2	20.3%	39.9%	22.3%	34.7%
3	4.5%	16.6%	5.5%	13.1%
4	0.7%	3.9%	0.8%	4.1%
5	0.005%	0.2%	0.2%	1.5%
6	-	-	0.008%	0.3%

543

544