

Dear Dr. Didenkulova,

We thank you for handling the review of our paper “Tsunami run-up estimation based on a hybrid numerical flume and a parameterization of real topobathymetric profiles”. Please, find in this document, our responses to reviewers in which we have included detailed replies to their comments.

Considering the topics that this reviewers have come up we have re-edited the manuscript, which is included at the end of this document. The changes are highlighted in green for reviewers1’s comments and in yellow for reviewers2’s comments.

Íñigo Aniel-Quiroga

Corresponding author

### Anonymous Referee #1

The manuscript proposes a methodology to estimate tsunami runup by mixing up a classical Tsunami code (COMCOT), for the first stages, and an averaged NavierStokes model for the runup process. In my opinion, the manuscript exhibits a well and organized work and I suggest that it should be published after minor revision regarding specific points that should be clarified, because they affect in the understanding and make the manuscript not fully reproducible.

REPLY: We thank reviewer 1 for the thorough review of our manuscript and the positive comments regarding its organization.

We have added their ideas and modifications to the paper (green color in revised version), increasing its overall significance.

Major comments:

1) Time computation is regularly mentioned, however there is no solid numbers. For example, how long it takes a regular tsunami running? How long it takes obtain the

Final run-up estimation with the presented methodology?

REPLY: Thanks a lot for coming up this topic. Computational time information is basic to have a more global approach of the process of construction of the database, and to understand how it makes easier and faster the run-up calculation.

On hazard assessments, in particular of large areas, the computational time become a key element on the methodology to apply. For example, in the case of a simulation of an event that travels through the ocean basin and then floods a local area, it can require several levels of nested grids to simulate the tsunami including a high resolution grid for the local area. The computational time to conduct this simulation depends on many aspects but it can take 10 to 16 hours in a common computer.

In the methodology presented in this paper, the interpolation itself takes just some seconds. If a numerical simulation with SWE model is carried out to obtain the wave conditions to use them as input for the IH-TRUST, then just a single grid for the whole ocean basin is necessary, what could take around one hour, depending again on the simulation domain.

Finally, the simulations that are of the database required a long time for calculation depending mainly on the size of the VOF domain. Typical times range from 2 hours to 16 hours.

This data has been added to the new version of the manuscript in the conclusions section (page 41, lines 6-8)

2) It is also not mentioned, but I guess authors have assumed an instantaneous tsunami generation. This have to be very clear. In general, there is lack of details on the tsunami modeling. Domain size, computation time, CFL condition (depending on your chosen grid size), etc. You should, at least, comment some lines due to the fact that time characteristics of the seismic source can enhance the tsunami amplification. This becomes important in huge and rare events as The 1960 Chilean Earthquake and 2004 Sumatra Earthquake, where the source time function is not well resolved (specially in the Chilean event). Besides of all the earthquake parameters, there is the slip distribution. It is demonstrated that the runup can be amplified up to six times (Geist (2002), Ruiz et al. (2015)). So, the kind of seismic sources should be clearly defined.

REPLY: As the reviewer correctly addresses, the tsunami generation follows some hypothesis or simplifications. Specifically, an instantaneous generation and a regular and constant slip

distribution were assumed. In this way, it is interesting to highlight, as the reviewer remarks, that when historic or past events are being simulated a proper source could be evaluated and used to determine the  $H$  and  $T$  of the tsunami wave to be used as an accurate input for the methodology. On the other hand, potential events that are part of tsunami hazard assessments commonly use idealized sources, that can be used to evaluate  $H$  and  $T$ . Anyway, the tsunamigenic sources used for elaborating the database were idealized parameterizations that were transformed into initial water surface displacement by means of Okada model.

Regarding the grid size for COMCOT simulations, it was set to  $\Delta x=500\text{m}$ . Depending on the maximum depth of the grid the necessary time step to satisfy Courant Condition was calculated and used, based on the restriction of the model for the condition  $Cr=0.5$ :

$$\frac{c \cdot \Delta t}{\Delta x} < C_r, \text{ where } c = \sqrt{g \cdot h}$$

Clarifying lines have been added to the manuscript (page 9 and page 14)

3) The approximation of the topo-bathymetric profiles are fitted from the GEBCO data, but no resolution is mentioned. The authors fixed the profiles with four (4) segments: a constant depth (1) connected to two lines offshore (2) and another line inland (1). The first and natural question is why to set 4 segments?? Is it because the 5-space of parameters is already big enough? Another issue, is that the trench morphology is not captured, or at least, not showed in the manuscript. This is because in subduction zones, before the ocean becomes "constant", there is a huge depression, especially in The Marianas trench, where the water column is higher and faster.

REPLY:

Regarding the GEBCO bathymetry, the topobathymetric profiles that were used in the elaboration of the database were obtained from GEBCO, using the resolution provided by this database, 30" (around 900 m in the Equator).

Regarding the chosen geometry, we analyzed worldwide profiles trying to find a parameterization that covered two main aspects. First, and mainly, that they could represent appropriately most of the profiles and second, that the selected parameterization allowed managing the database to be created. The technique for classification (Maximum dissimilitude) and interpolation (radial basis functions) are specially designed to work on high dimensional domains (i.e. Camus et al., 2011), therefore the number of segments is not an issue. Nevertheless, a run-up calculation requires the parameterization of a profile as input, therefore that parametrization must be functional. After considering other options, like adding a new segment, we considered that our parameterization achieved this equilibrium between representability and functionality.

Regarding the subduction trench, its applicability, as it can be seen in the applications cases is limited to those generation areas that are deep enough to be part of a profile included in the database. In this sense, the system works quite well if this is the case, as it was observed in the examples of Chile. However there is a limitation, well noted by the reviewer: the profile parameterization falls out of the application range (see page 34 line 13). The result of assuming that the sea bed is constant seawards the generation area gives a good approximation of the "trench problem".

Some lines explaining this aspects have been included in the new version of the manuscript (page 7, 5, 35.) and in the conclusions section

4) Authors "cheats" the tsunami interaction of the reflected wave by assuming a constant and flat region with open boundary condition. However, would not this add some kind of artifacts to the model? Test regarding this issue should be do it.

REPLY:

One of the biggest issues to perform the coupling between models was to obtain a clean input wave for forcing the VOF domain. The problem arises because the wavelength of tsunami waves is, in some cases, longer than VOF model domain. This implies that before a

tsunami wave passes completely through the boundary between models, the wave front reaches the coast, is reflected and return to the initial boundary aliasing the wave amplitude.

Several tests have been run to assure that this artifice of assuming plane beach and open boundary do not affect the numerical simulation. This artifice avoids the interaction of the tsunami wave arriving to the coast and the reflected wave, “cleaning” optimally the signal of the tsunami waves that were included in the development of the database, as explained in figure 6.

5) Authors make use of analytical solution of Synolakis (1987), however, I’m not convinced that is the good one here. There are analytical solutions in piecewise bathymetries (e.g. Kanoglu & Synolakis (1998), Fuentes et al. (2015), Riquelme et al. (2015)). Actually, in figure (13) the results do not agree with those analytical solution which state that offshore slope closest to the coast controls the runup process.

We really appreciate this comment. We have used Synolakis as an example of comparison because although it was created for Solitary waves it has been commonly used on tsunami risk assessments, despite its application is not appropriate as highlighted by Madsen (2008). In order to improve this, we have added a new column to the comparison tables with the value of the run-up but calculated with Madsen and Shaffer (2010) for single waves, as also suggested by Reviewer 2. In addition we have included the given references in order to have a more complete view of the existing solutions,

In the case of solitary waves it was shown by Kanoglu and Synolakis(1998) that the slope closest to the coast controls the run-up. Our results show that, actually, the slope closest to the coast is very important in the final value of the tsunami run-up. However, according to our results, in the case of tsunami waves the influence of the other slopes, especially the one of the next segment seawards should not be neglected, remarking again the difference in the wavelength of solitary waves and tsunamis, what leads to a different behavior, a different time while the wave is affected by each segment of the bathymetry, affecting the reflection, shoaling etc.

This aspect, also noted in Naik & Behera (2016) using numerical models, has been also added to the manuscript, in the page 26 and in the discussion section.

6) It is not mentioned the criterion to trace the profiles. Perpendicular to the shore? Paralel to the wave travel??

REPLY: Orientation of profile is a key piece on the runup calculation using this method. As the database were constructed using a numerical flume in which wave direction of propagation and profile coincide, this is of course the best configuration to trace profiles. Nevertheless, in real scenarios is not so simple to define this direction.

This aspect has been added in the new version of the manuscript (page 18)

7) The methodology is compared with numerical models and retrieves same estimations. The fact that inaccessible high-resolution data can be overcome should be more highlighted. Again, I don't think Synolakis's formula is comparable here, since it uses a Solitary wave as initial condition, and there are analytical solutions that can handle with arbitrary shapes (Madsen & Schaffer,(2010) , Fuentes (2017) ).

REPLY: The fact that inaccessible High Resolution data can be overcome is now highlighted appropriately (Page26) and it has been included as well in conclusions section (page 41, line 16).

AS explained in the reply to the comment 5, we have included the results of the application of Madsen formula, highlighting the fact, remarked by the reviewer 1, that Synolakis formula, although widely applied, was created for Solitary waves.

Specific comments: Thanks a lot for this specific comments. We have added all of them to new version of the manuscript.

- First line of intro: Add the 2010 Chile tsunami.
- Page 9, 5: It seems that "H" is unnecessary here. Also, it should be clarified that period relation is valid in the linear regime.
- Please add geographic axis to the map plots.
- Page 36, 5: Authors say "the results are accurate". Please, add a percentage based on the results.

References:

- Kanoglu, U. & Synolakis, C.E., (1998). Long wave runup on piecewise linear topographies, J. Fluid Mech. (JFM), 374, 1–28.
- Fuentes, M., J. Ruiz, and S. Riquelme (2015), The run-up on a multilinear sloping beach model, Geophys. J. Int., 201(2), 915–928.
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Interactive comment on Nat. Hazards Earth Syst. Sci. Discuss., <https://doi.org/10.5194/nhess2017-445>, 2017.

## Anonymous Referee #2

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General comments: This paper presents a method for quickly assessing tsunami runup for different tsunami wave shapes and bathymetries. The method is built on a hybrid approach, where the Non-Linear Shallow water model COMCOT is used for the deep-water propagation and a RANS models is used for the near shore processes. The results from this hybrid model, then enters an interpolation model, which can be used to assess run-up. The approach is novel and innovative, and I especially like adding a fast interpolation model. I have however, a few major concerns. The actual implementation of the hybrid model is not well described and I think the coupling between the two models can pose big problems. Further, the hybrid model is not validated on its own in a controlled environment.

REPLY: We thank Reviewer 2 for their careful review and for the positive comments regarding the developed interpolation modeling. We have considered their modifications and suggestions, what has increased its overall significance. Changes included following Reviewer 2 comments are in yellow color in the revised version. We respond in detail to their comments below.

We specially thank their comments towards explaining more deeply the hybrid model implementation. Initially we included some more details regarding this specific part of the work, that concentrated our efforts, like a specific validation. However, in the end, we decided to skip some of this data to keep the focus of the paper on the run-up calculation. Following Reviewer 2 comments we have reconsidered this aspect, and we have added more information to make it easier all the process to follow. This incorporation provides a substantially improved approach of the hybrid model.

## Major comments:

1) Parts of implementation and usage of the hybrid model is not well described.

a. What are typical grid sizes in the RANS model? Are these sufficient to handle the processes, which NLSW models cannot handle? Like wave breaking.

## REPLY:

The design of the domain of the RANs model (IH2VOF) followed 2 criteria.

First, there is a limitation of the model that do not allow grid more than 5499 cells in X dimension ( $n_x < 5499$ ), and the ratio between dimensions must be constant ( $r = \Delta x / \Delta z = \text{constant}$ ). In this case, and due to the different scale necessities on each dimension, the applied ratio was  $r = 5/1$ . Therefore, the maximum length covered with RANS model was  $L_x = n_x \cdot \Delta x$ .

And second, to control and avoid false wave breaking, the Z dimension of the RANS model grid must be discretized in a minimum number of cells, satisfying the expression:

$$\Delta z = \left\lceil \frac{K \cdot H_{COMCOT}}{10 \cdot 0.05} \right\rceil \cdot 0.05$$

Where K is a safety margin of the model=1.08 and  $\Delta z$  is defined in the range ( $0.05 < \Delta z < 1$ ). The effect of the down-open brackets is “rounding to the lowest integer”.

To sum up, in this sense, the model itself limits the length  $L_x$  and the grid size  $\Delta x$ . First,  $\Delta z$  is calculated with  $H_{comcot}$ , then  $\Delta x$  is obtained  $\Delta x = \Delta z \cdot r$ , and finally,  $L_x = n_x \cdot \Delta x$ . This approach results in values of  $L_x$ , depending on the Tsunami wave height of the COMCOT model. For  $H_{comcot}=0.5$ , then  $L_x=1400$  m. For  $H_{comcot}=4.5$ m, then  $L_x=12400$ m.

Following this process, the grid size is enough to handle the processes that the LSWE model cannot.

**b. What are typical  $L_x$  lengths?**

$L_x = n_x \cdot \Delta x$ , and the value of  $\Delta x$  depends on  $H_{comcot}$ . In the same way as it has been described, the RANS model limits in terms of number of cells drive the generation of the grid, and also the value of  $L_x$ , that is calculated following two rules:

“Maximize the area” means to use all the available cells to cover the maximum length.

And, since the number of cells is limited, we did not want to lose cells onshore far away from the flooded area. Thus, we pre-calculate a rough value of the run-up using COMCOT and approach a first value of the run-up.

Taking this into account, typical values of the RANS model length are from 500 to 25000 m.

**c. How are the boundary conditions for the turbulence mode?**

REPLY: In the case of turbulent flows, IH2VOF numerical model uses smooth wall, log-law distribution for the mean tangential velocity.

This aspect has been included in page 4 of the new version of the manuscript.

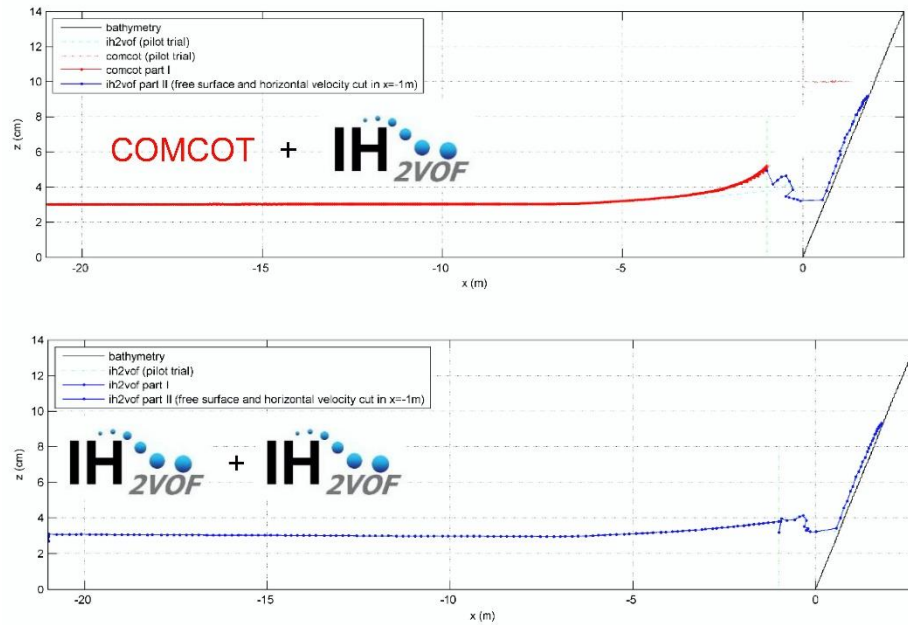
d. It is unclear how  $x_{cut}$  is determined. In the paper two criteria is given. One is to maximize the area of the IH2VOF domain and the other is to ensure that flooding does not exceed the inland end. Regarding the first criteria, letting IH2VOF cover the entire numerical flume would achieve that, but that is clearly not what is being done. Regarding the latter, I cannot see how the end position of the IH2VOF domain influences the position of  $x_{cut}$ .

REPLY: AS it has been described, the limit of the RANS model in terms of number of cells drives the generation of the grid. The position where the models are coupled,  $x_{cut}$ , is given then by the value of  $L_x$ . As commented previously,  $L_x$  is set trying to maximize the cells that are effectively used in the simulation. This important aspect of the hybrid model was not included in the first submitted version of the manuscript, but following reviewer 2 comments we have added it. Regarding specifically the first criteria, since the flume is non-scaled, it was not possible to cover the whole domain with RANS model due to computational restrictions, i.e., we cannot calculate the generation-propagation and inundation areas without assuming other limitations of scale. Moreover, offshore generation and propagation is well solved by LSWE model, where non-linearities are not relevant in the calculation.

e. One of the advantages of the hybrid model is that the RANS model can handle processes that the simpler COMCOT model cannot. One of such processes is the wave splitting into an undular bore, which can happen when the wave travels long in shallow water and this has been witnessed in many real life tsunamis. To be able to capture this effect  $x_{cut}$  needs to be positioned sufficiently off shore. How is this ensured?

REPLY: Due to the exposed characteristics of the model it is not possible to extend the RANS model grid seawards more than what this limitation allows. Nevertheless, to avoid loss of processes on coupling, several coupling tests were performed. These test were conducted reproducing the flume on scale, so a target simulation were performed using only the IH2VOF model. Lately, simulations of coupling between IH2VOF + IH2VOF and COMCOT + IH2VOF were performed and compared to single IH2VOF simulations. In both cases IH2VOF cases coupling methods reproduce evolution of single IH2VOF simulations adequately. In the attached figure one of the conducted tests is given. In this figure it can be observed that, when the last part of the flume is calculated by means of IH2VOF model,

the comparison between the different combinations of models is accurate in terms of run-up.



As a consequence, the part of the processes that are not covered by LSWE model are incorporated in the last part of the flume, where the RANS model works.

- f. To avoid reflection, the numerical flume of COMCOT is altered, to properly access the incoming wave. I have a problem with this approach. In reality, especially in cases with steep slopes, there will be significant reflection from the beach which will and should affect the incoming wave. This effect cannot be captured with the current approach. Further it is also unclear what would happen when the reflected wave from the IH2VOF domain meets the hard boundary between the two models. Will this cause additional reflections in IH2VOF domain?

REPLY: The artifice that has been applied to avoid reflection effects on the input wave focus precisely on avoiding that the reflected wave affects the tsunami wave between borders incident one. This unaltered wave is used to force the IH2VOF domain, in which simulation reflection on the beach is of course observed and considered for the runup calculations. Figure 6 of the manuscript tries to explain how this artifice corrects the possible reflection effect. Fig 6a shows the reflected wave and Fig 6b shows how the artifice works and the reflective effect is almost imperceptible. This figure quote has been improved in order to make clear the artifice intention.

- g. The calculations of  $L$  does not match Fig. 4. E.g.  $L_i$  is given as  $L_i = 1/50 \tan(\beta_0)$ . This will result in a very low  $L_i$ . Further  $L_{off}$  is given as  $L_f + x_2$ , but according to Fig. 4 it should be  $L_{off} = L_f + x_2 + x_1$ . Finally, there is no need for  $\delta X$  in the calculations of  $L_f$  as it is present both in the denominator and the numerator.

REPLY: Regarding the calculation of  $L_i$ , the expression in the manuscript is incorrect, as describe by the reviewer. It should be:

$$L_i = \frac{50}{\tan \beta_0}$$

Regarding Calculation of  $L_{off}$ , again, we appreciate reviewer correction. It should be:

$$L_{off} = L_f + x_2 + x_1$$

Finally, regarding the  $L_f$  calculation we used  $\lfloor x \rfloor$  “ceiling brackets” which is the largest integer less than or equal to  $x$ , commonly used in mathematics and computer science. Then, using  $\Delta x$  both in the denominator and numerator, allows us rounding values to the order of  $\Delta x$ .



h. How is the run-up height determined in the IH2VOF model? In a VOF computation, the interface can span across several cells.

REPLY: The numerical model calculates for the last flooded cell the ratio that is actually flooded and provides the run-up in accordance.

Indeed, due to the fact that the RANS model does not calculate directly the free surface but it tracks the changes in cell density, there are mainly 2 ways to tackle this calculation:

- Assessing the iso-surface, determining the contour where VOF function is 0.5
- Calculating the water contained, accumulated in a column of the grid. In order to avoid diffusion, the quantity of water is added.

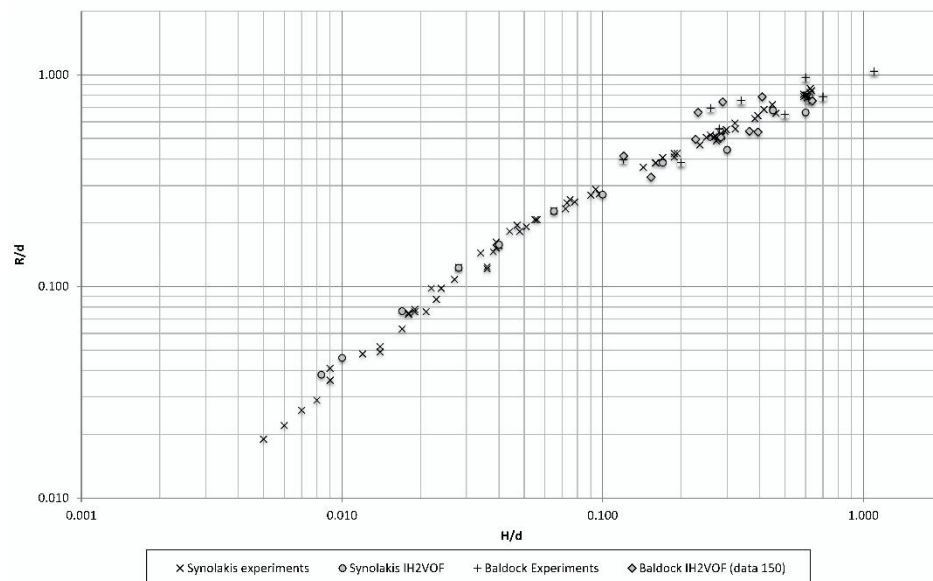
In this case, after several “trial and error” tests, the second method has been applied, although in the end the difference were not serious. None of the methods is perfect but both of them provided a good approach.

This aspect is included in page 4 of the new version of the manuscript.

**Responses to comments regarding models coupling have been included in the paper section regarding the characteristics of the numerical flume.**

2) The first validation case is performed by comparing the interpolation model to the hybrid model. This is an important and satisfying validation case, but I am lacking validation of the actual hybrid model. How will the model perform using the approach outlined in the paper e.g. in cases with both breaking and non-breaking waves running up a constant slope.

REPLY: Both, COMCOT and IH2VOF models are models that have already been successfully validated in the past. The validation of the numerical flume and the coupling of the models was made by comparing its results with those conducted by Synolakis (1987) and Baldock (2009). The scenarios of these experiments were calculated using COMCOT, IH2VOF, COMCOT+IH2VOF (the complete numerical flume) and the results were compared to the results of the physical experiments. The attached figure shows the run-up obtained in this comparison and how the results of both series of experiments fit adequately with the numerical flume results. This validation allowed us to continue with the database elaboration, and the database itself was then validated by comparing the results with both numerical models and field work data.



IH2VOF model is, in the hybrid model, the responsible of incorporating non-linearities and breaking effects. In this sense, apart from the validation as part of the numerical flume, it has also been validated and applied in many studies e.g.

The shown figure, together with an explanation of the validation has been added to the manuscript, at the end of section 2 (pages 13).

3) The performance of the iterative solver is compared to the Synolakis formula. I do not believe this is a fair comparison, as it is created for the run-up of a solitary wave, which as highlighted by the author does represent a real geophysical tsunami event. A more relevant comparison could be to the analytical model proposed by Madsen and Schäeffer (2010), as also highlighted by the authors in the introduction. (The Synolakis formula require a proper reference).

REPLY: We appreciate this comment. It was also highlighted by Reviewer 1. We have added the results of applying Madsen and Shaffer (2010) to the table where the result of applying the run-up database interpolation are given. AS explained to Reviewer 1 we think that this complementary validation definitely improves the significance of the paper. In addition a proper reference to Synolakis formula has been included. We have decided to leave the Synolakis formula application results as well due to its common use in literature.

4) The periods are estimated as the time between the first two zero crossings for positive heights. Does this mean that the model cannot differentiate between tsunamis having only positive surface displacement and e.g. a leading depression?

REPLY: This criteria (period=time between 2 positive zeros), is the one that IHTRUST, the interpolation tool and system, uses to automatically calculate the period of a time series. However the system itself allows to manually edit the period: It shows the time series and the part of it that is going to be considered in the interpolation. If other period, like in the case of a leading depression is to be used, it can be manually corrected. However, several tests were performed, and we did not found serious differences on simulations led by crest or trough regarding the run-up results. Therefore, the tool was scripted following the explained criteria, although as explained it allows its modification.

5) With this approach of estimating period and wave height, I see a potential problem in the case where the leading wave is not the largest. Can you please elaborate on this?

REPLY: This is a real limitation that concerned us during the process of scripting the IHTRUST. It was the main reason why we added the “manual way” to include the height and period. This allows to use an ad hoc input data on the interpolation process. Although, obviously it cannot take into account several waves of the tsunami, it assures to use the proper part of the time series as input.

We have reinforced the IHTRUST explanation in page 20

6) One of the main points off this work is to be able to quickly access tsunami runup without doing long complicated simulations. Therefore for this work to fulfill this, it would be beneficial is the TRD database was made available to engineers. Are there any plans regarding this?

REPLY: So far, we are still making the most out of the tsunami run-up database. We are specifically working on a better definition of the influence of each parameter on the final value of the run-up. Once all the analyses are finished, we are planning to release the data, by itself or on a new issue.

Smaller comments: Thanks a lot for this smaller comments. We have included them to the new version of the manuscript.

1) Page 1, line 8: It is stated that Run-up is accurately calculated by means of numerical models. This is a rather strong statement. I would prefer it rewritten as: can be accurately calculated

2) Page 1, line 14. The models here, and several other places are referred to as schemes. They are however not schemes, but models. Please change the formulations.

REPLY: We have changed it.

- 3) Fig. 5. It is stated that only the COMCOT model is used with the altered domain. If this is indeed the case, then please remove the IH2VOF domain from the figure, as it is causing confusion.
- 4) Fig. 6 units and legends are missing on the colorbars. Please add these.
- 5) Page 14, line 2. It is stated that  $d^2 - d^1$  was always shorter than 2200 m and  $x^1$  shorter than 210 km. How does this correspond with table A1 where  $d_2 - d_1$  is always larger than 2200 m? [It should say larger.](#)
- 6) Fig 8. Many of the axis are missing units. Please add these.
- 7) Fig 10. Please add missing units to the axis.
- 8) Page 19, line 6. It is stated that T corresponds to the time between the first two zero crossings for positive heights. However in Fig 11. It looks as if the second zero crossing has not occurred within the shaded area? [The IHTRUST tool calculates the best fitting to this criteria. In this case the red shaded area approximates the period. The time series is built with discrete points and the system catches the closest one.](#)
- 9) Page 24. It is described how low values of  $\tan(\beta_2)$  gives lower run-up height due to friction. From Fig 13, it can however also be seen that the run-up heights reduce with large values of  $\tan(\beta_2)$ . Please elaborate on this.
- 10) It is unclear exactly what Fig 15 is describing. Please rewrite the description for clarity and add units to the axis.

[REPLY: We have modified this figure in order to make it clear that it is just a scheme of the main "regimes" that have been found. It is the figure 17 in the new version of the manuscript](#)

# Tsunami run-up estimation based on a hybrid numerical flume and a parameterization of real topobathymetric profiles

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**Abstract.** Tsunami run-up is a key value to determine when calculating and assessing the tsunami hazard in a tsunami-prone area. Run-up can be accurately calculated by means of numerical models, but these models require high-resolution topobathymetric data, which are not always available, and long computational times. These drawbacks restrict the application of these models to the assessment of small areas. As an alternative method, to address large areas, empirical formulae are commonly applied to estimate run-up. These formulae are based on numerical or physical experiments on idealized geometries. In this paper, a new methodology is presented to calculate tsunami hazard at large scales. This methodology determines the tsunami flooding by using a coupled model that combines a nonlinear shallow water model (2D-H) and a volume-of-fluid model (RANS 2D-V) and applies the optimal numerical models in each phase of the tsunami generation-propagation-inundation process. The hybrid model has been widely applied to build a tsunami run-up database (TRD). The aim of this database is to form an interpolation domain with which to estimate the tsunami run-up of new scenarios without running a numerical simulation. The TRD was generated by simulating the propagation of parameterized tsunami waves on real non-scaled profiles. A database and hybrid numerical model were validated using real and synthetic scenarios. The new methodology provides feasible estimations of the tsunami run-up; engineers and scientists can use this methodology to address tsunami hazard at large scales.

## 1. Introduction

Recent tragic tsunami events, like those that occurred in the Indian Ocean in 2004, in Chile in 2010, and in Japan in 2011 have exposed the need for further work to develop and apply tsunami risk reduction measures. The adequate evaluation of tsunami hazard in tsunami-prone areas is the first step in a proper risk evaluation (UNESCO-IOC, 2009). Determination of the tsunami hazard focuses on the estimation of the area that would be flooded during a tsunami and on the calculation of the variables or parameters that define the phenomenon in that area, e.g., wave amplitude, current depth, tsunami travel time, etc. Among these parameters, maximum run-up provides the elevation to which water from a tsunami wave will rise during its flooding process. Therefore, run-up is a key parameter that must be adequately determined when assessing the inundation of affected areas.

When tsunami hazard is addressed at a local scale (tens of kilometers or one coastal city), the optimal methodology to calculate the flooding and run-up is typically the application of validated deterministic numerical models (Álvarez-Gómez et al., 2013; Titov et al., 2011; Wang, 2009). These models allow reproduction of the 3 main tsunami processes: generation, propagation and inundation. To address these processes and to properly estimate the flooded area, high-resolution topography-bathymetry data of the study area are required, as well as the focal parameters that define the tsunamigenic mechanism. Nevertheless, the application of tsunami numerical models has some limitations and uncertainties (Park et al., 2015; Selva et al., 2016). First, their use requires a high computational cost and expert modelers. Second, the necessary high-resolution data to properly study the hazard in local areas are not always available. In addition, the correct definition of the tsunamigenic mechanisms, e.g., the parameters of the focal mechanism, contains uncertainties in itself. Finally, even though models are evolving to reduce uncertainties, there is still ongoing work on several aspects, such as wave transformation near the coast, interaction of waves with coastal structures, and accurate incorporation of bottom friction.

On the other hand, in large-scale studies (hundreds of kilometers or the coast of a whole country), the drawbacks of numerical models are more evident, and the lack of continuous high-resolution topobathymetry and the elevated computational cost foster the use of other approaches. An alternative methodology to estimate the tsunami run-up and, consequently, the flooded area, includes the application of run-up analytical or empirical formulae. In these cases, numerical models, despite the lower resolution of bathymetry, adequately calculate the tsunami wave characteristics offshore and can then be used as input for the formulae. Afterwards, by applying this method to several topobathymetric profiles along the coast, the total flooded area due to tsunami action can be estimated.

The calculation and analysis of run-up was initially approached by Carrier and Greenspan (1958). They found the exact solution for the nonlinear shallow water equations for a sloping beach with non-breaking regular waves. Keller and Keller (1964) derived an analytical solution for linear shallow water waves at a constant depth moving up a constant slope beach. This geometry has become the canonical problem. Synolakis (1987) extended Carrier and Greenspan's result to this problem by joining Carrier and Greenspan's and Keller and Keller's solutions to provide a closed-form solution for solitary wave run-up. Synolakis' results are remarkable, as solitary waves have been widely used to model tsunamis, numerically and physically. Li and Raichlen (2001) revisited Synolakis's results to determine the importance of a higher order correction to the analytical approach. Later, Madsen et al. (2008) demonstrated that solitary waves do not represent the large scale of a tsunami, and Chan and Liu (2012) confirmed this affirmation. Madsen and Schäffer (2010) found closed-form solutions for the run-up of waves of several shapes; their solutions included other parameters, such as the period, achieving more realistic results. Finally, (Fuentes et al., 2015) studied the run-up on multilinear sloping beaches.

In addition, run-up has been commonly linked with the Iribarren number (Iribarren and Nogales, 1949), also called the surf similarity parameter (Battjes, 1974). Hunt (1959) joined this parameter with the non-dimensional run-up of regular waves.

Kobayashi and Karjadi (1994) combined physical and numerical simulations to derive an equation to calculate run-up, using the ratio between the run-up and the wave amplitude and its relationship with the surf-similarity parameter. Fuhrman and Madsen (2008) demonstrated that the relationship between surf-similarity and solitary waves was similar to the that between surf-similarity and period waves.

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More recently, several authors have focused their work on calculating tsunami run-up by developing new models with other approaches. Sepúlveda and Liu (2016) presented expressions for the calculation of the run-up based on the parameters that defined the focal mechanism of the tsunamigenic seism. Riquelme et al. (2015) derived simple solutions to estimate run-up on nearfield tsunamis by extending Synolakis solution (Synolakis, 1987) and Park et al. (2015) defined the run-up for compound slopes, based on the work of Madsen and Schäffer (2010) and numerical simulations of tsunami waves on two-slope topobathymetric profiles.

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However, the application of these equations and formulae is not always evident, and each approach considers different inputs. Moreover, the parameterization presented by Carrier and Greenspan (1958), extended by Synolakis (1987) and modified by Park et al. (2015), is based on theoretical bathymetric profiles. It does not explicitly consider real profiles or the geometry of the whole area, from the tsunami generation zone to the flooded area. Furthermore, the numerical models that do consider the natural geometry of the bathymetric profiles adequately predict propagation, but they cannot accurately solve the flooding calculation, in addition to the other exposed drawbacks.

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20 Complementing these methodologies, this work presents an alternative methodology to calculate tsunami flooding at large scales and is focused on assessing the run-up. The methodology is then applied to further develop a database from which the tsunami run-up of new scenarios can be interpolated.

The main component of the methodology is a numerical flume where the simulations are run. This flume was developed by combining a nonlinear shallow-water-equations model and a Navier-Stokes volume-of-fluid model to create a hybrid model that applies the optimal numerical model in each area of the flume. Time series of tsunami waves and topobathymetric profiles are used as input to calculate the run-up.

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This hybrid model has been applied to further develop a database from which the run-up of new tsunami scenarios can be interpolated. This database contains an adequate representation of natural bathymetric profiles worldwide and the variability in tsunami wave shapes, allowing calculation of the tsunami run-up of new scenarios by interpolation without running a numerical simulation.

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The aim of this methodology is to help specialists to further develop tsunami hazard maps at large scales, where the application of numerical models is not computationally affordable and high-resolution data are not available. This method can be used to quickly estimate the run-up in tsunami-prone areas or accurately estimate the flooded area for new tsunami scenarios.

5 The paper is structured as follows: Section 2 describes the developed methodology, including the parameterization of realistic bathymetric profiles and tsunami wave shapes and the construction of the numerical flume. In section 3, the application of the methodology to calculate the tsunami run-up database is discussed, together with a sensitivity analysis of the influence of each parameter on the final value of the run-up. Section 4 includes details of the tool that has been developed in order to use the database to calculate new tsunami event run-ups. Section 5 presents the validation of the methodology with real and numerical  
10 scenarios. Finally, section 6 discusses the conclusions drawn from this work.

## 2. Tsunami run-up hybrid model methodology

The run-up calculation methodology presented in this paper consists of the numerical simulation of tsunami waves along real non-scaled bathymetric profiles that were previously parameterized.

To carry out these simulations, a numerical flume was designed. This flume is formed by the coupling of two numerical  
15 models.

The Cornell Multi-grid Coupled Tsunami Model (COMCOT, (Wang, 2009)) solves the nonlinear shallow water equations (NLSWE) using a leap-frog finite differences scheme on a 2D horizontal domain. In addition, the IH2VOF model solves volume-averaged Reynolds-averaged Navier-Stokes (VARANS) equations based on the decomposition of the velocity and pressure fields into mean and turbulent components using a  $\kappa$ - $\epsilon$  turbulent model on a 2D vertical domain, using a log-law  
20 distribution for the mean tangential velocity (Lara et al., 2006). The former model is prepared to simulate the stages of tsunami propagation; meanwhile, the latter model is specially designed to simulate the coastal processes and wave transformations present when the waves reach the coastal areas. IH2VOF model calculates the run-up by evaluating the water accumulated in each column of the grid, tracking the changes on each cell density.

25 In the flume, the strengths of both models are used to design a numerical space where tsunami waves are propagated, using COMCOT from the deep ocean (~4 km depth) to the coast, where the capabilities of the IH2VOF model are applied to calculate the flooding. As a result, a hybrid model that adequately solves the tsunami processes in both deep and shallow waters was achieved.

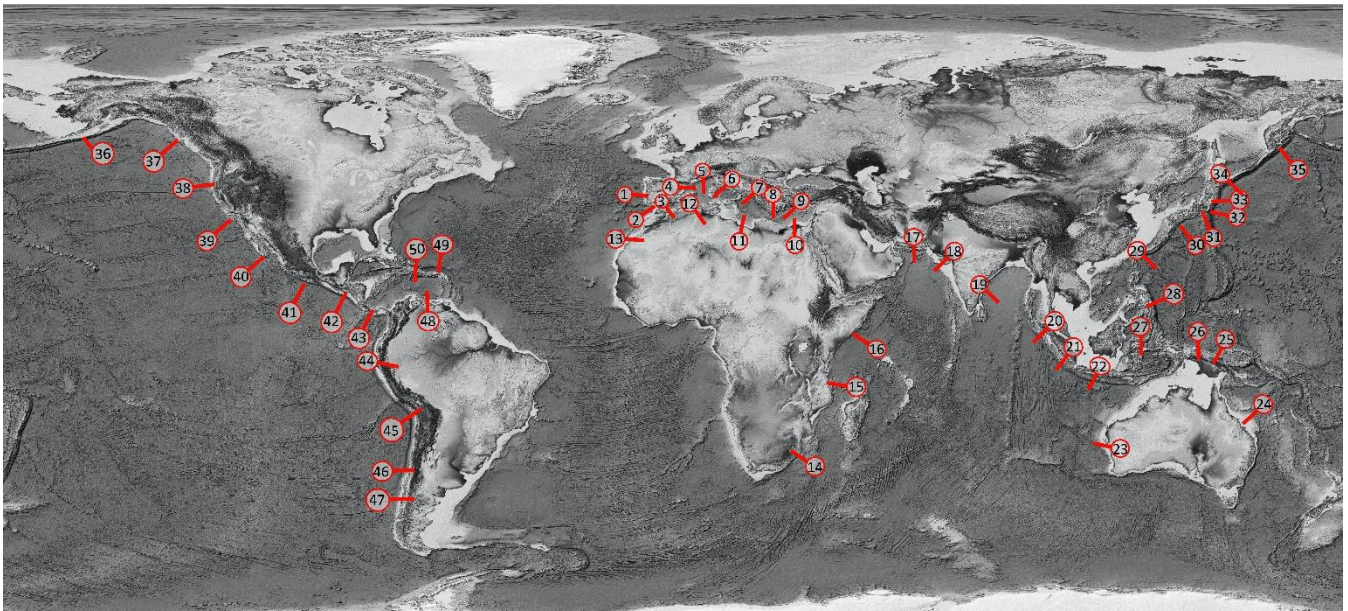
30 Parameterized profiles and a tsunami wave time series dataset are used as input for the numerical flume. These inputs, the most relevant aspects of the numerical flume geometry, and the coupling of the models are described below.



## 2.1 Bathymetric profile characterization

Worldwide bathymetric profiles were analyzed, with a focus on finding a parameterization that properly represents natural shapes.

- 5 To cover the existing variability in the world bathymetry, a representative sample of 50 averaged profiles was obtained from tsunami-prone coastal areas and basins, namely, the Pacific Ocean, Indian Ocean, Mediterranean Sea and Caribbean Sea (Fig. 1). Topographic and bathymetric information was obtained from the **General Bathymetric Chart of the Oceans (GEBCO, International Hydrographic Organization, 2014, with a cell size  $\Delta x=30''$ )**, The European Marine Observation and Data Network (Bathymetry Consortium EMODnet, 2016) and the local bathymetry data that was available. The shape of these profiles was
- 10 analyzed to perform an adequate parameterization.

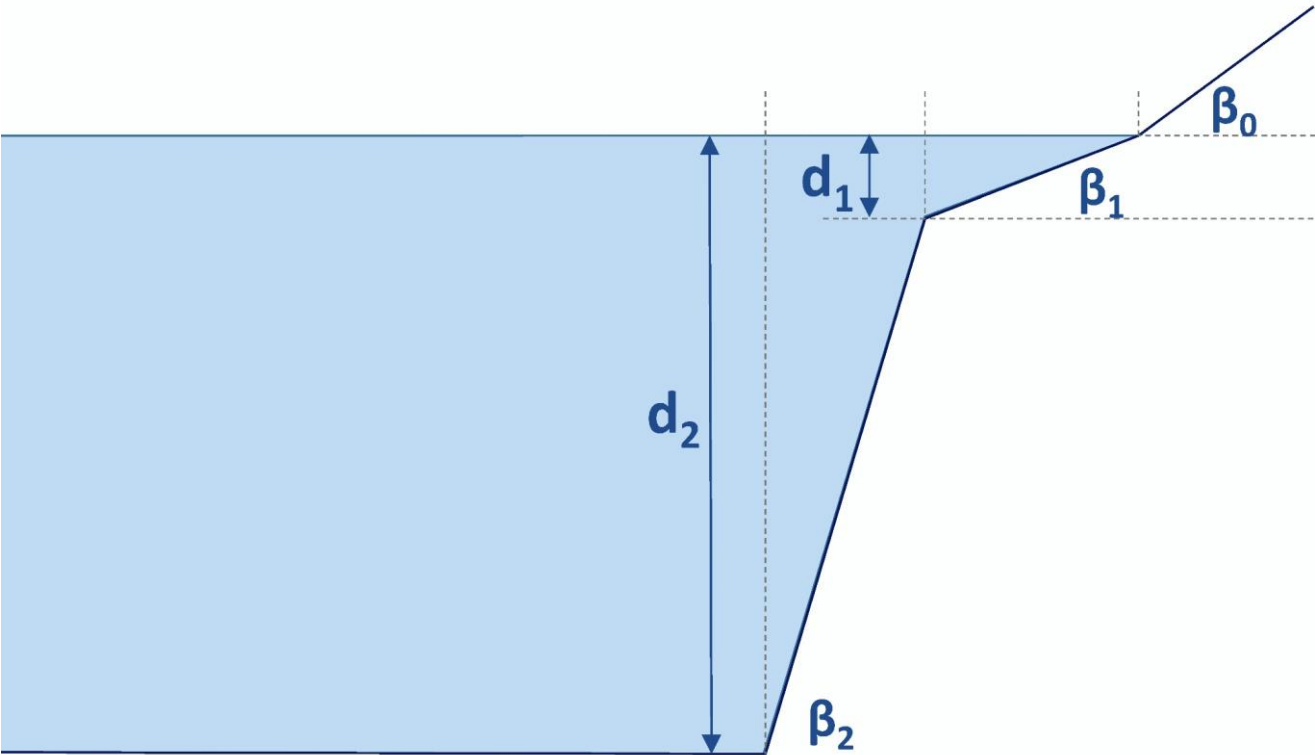


**Fig. 1. Distribution of sample profiles**

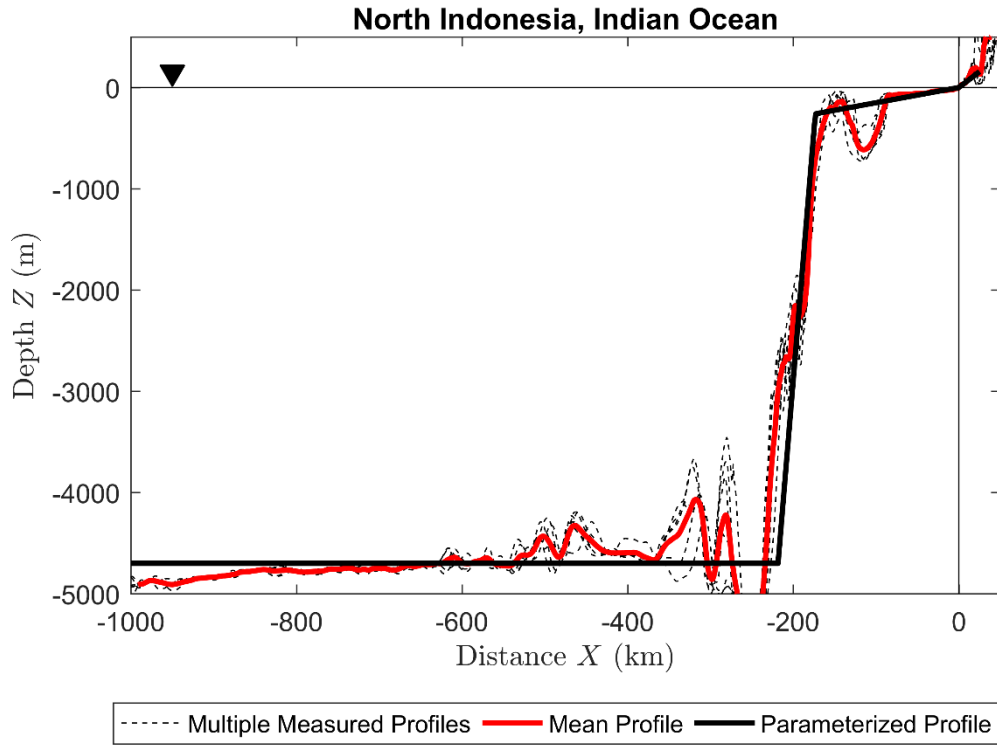
- 15 The propagation of a tsunami can affect thousands of kilometers; thus, the profiles must extend under both deep water and shallow water to capture tsunami generation to flooding. Considering this requirement, profiles were defined from inland (50 m height) to the deep ocean (~4000 m depth). To avoid singularities, each defined profile is the average profile of a 10-km-wide coastal segment. **Based on the bathymetric shapes observed in this selection, a representative and functional parameterization the profiles was tackled, using five parameters: three slopes ( $\tan \beta_1$ ,  $\tan \beta_2$ , and  $\tan \beta_3$ ) and two depths ( $d_1$  and  $d_2$ ). Fig. 2 shows the five-parameter geometry.**
- 20



As an example, in Fig. 3, a selected profile from the Indonesian coast is shown, as well as its parameterized profile that was created by applying the five-parameter geometry. The parameters for each considered profile were fitted by using a least-squares method.



5    **Fig. 2.** Scheme of the parameterized profiles, based on real profiles analyses. The profiles are defined by 3 angles ( $\tan \beta_1$ ,  $\tan \beta_2$ , and  $\tan \beta_3$ ) and 2 depths ( $d_1$  and  $d_2$ )



**Fig. 3. Sample of measured topobathymetric profiles on the Indonesian coast, as well as the mean and parameterized profiles.**

**GEBCO was used as source for the topobathymetric data (Cell size  $\Delta x=30''$ )**

The maximum and minimum values of the 5 parameters are shown in Table 1. These values cover a wide range of the profiles that can be found in nature. Despite not including all the existing geometries, the maximum and minimum values certainly provide enough information to characterize the topobathymetric profiles.

**Table 1. Maximum and minimum values of the profile parameters**

Parameter	Min	Max
$d_1$	20 m	1100 m
$d_2$	2200 m	6000 m
$\tan\beta_0$	$5.0e^{-4}$	$1.5e^{-1}$
$\tan\beta_1$	$5.0e^{-4}$	$2.5e^{-2}$
$\tan\beta_2$	$1.0e^{-2}$	$2.0e^{-1}$

## 2.2 Initial tsunami wave characterization

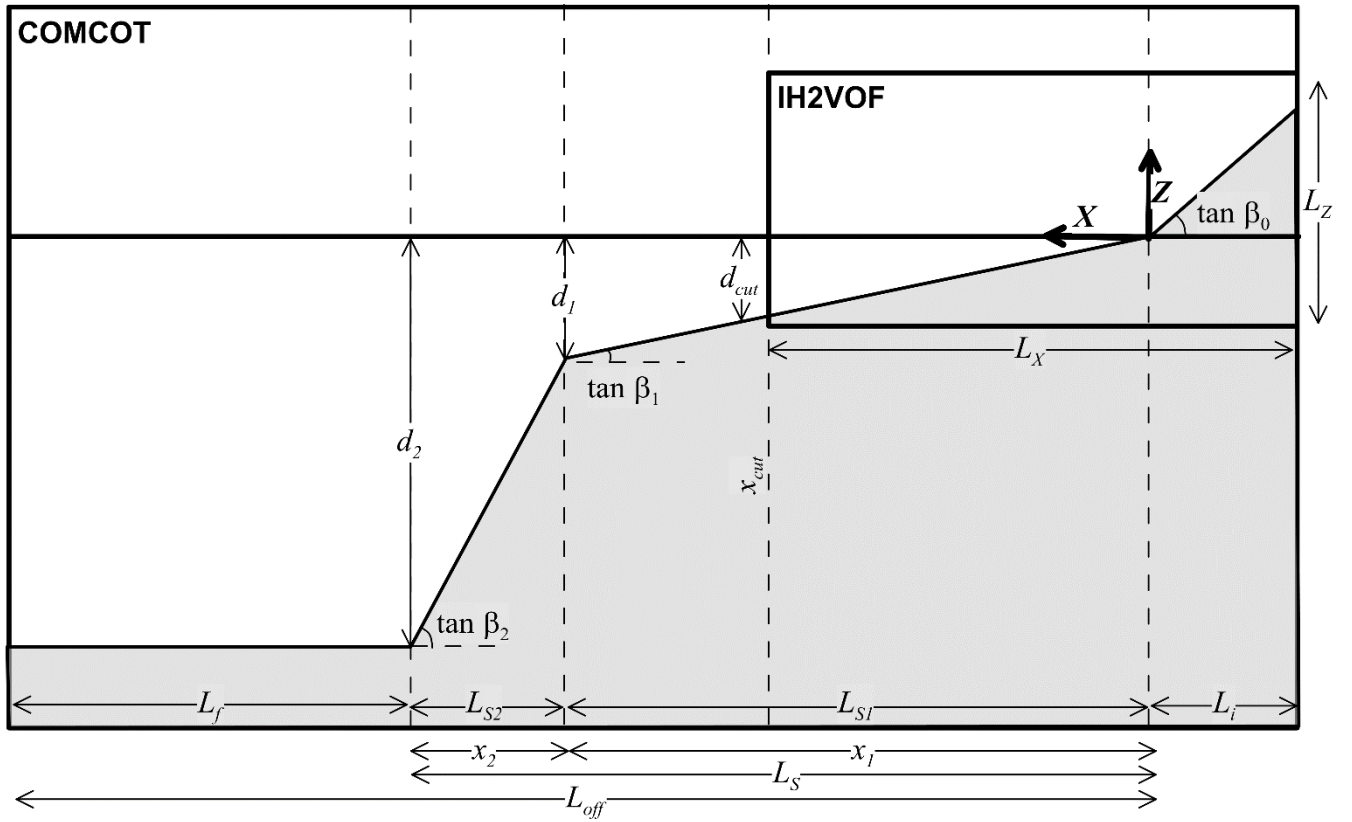
The numerical flume described in detail in the next subsection requires not only the topo-bathymetric profile characterization but also the characteristics of the tsunami waves as input. To use these data as input for the hybrid model, a time series of the offshore wave amplitude must be provided. These time series could be obtained from either records of real tsunamis, e.g., from DART buoys, or from the results of numerical model tsunami propagation. In this case, COMCOT (Wang, 2009) was adopted. This model calculates all stages of tsunami modeling (generation, propagation and coastal flooding).

The generation of the tsunamis in COMCOT is approached via elastic finite fault plane theory, using the so-called Okada model (Okada, 1985). This model assumes an idealized rectangular fault plane as a representation of two colliding tectonic plates. The Okada model requires 7 focal mechanism parameters as input to calculate the initial deformation of the water

10 surface due to the earthquake. These parameters are the focal depth ( $h_{focal}$ ), rupture length ( $L$ ) and width ( $W$ ) of the fault plane, dislocation ( $D$ ), strike direction ( $\theta$ ), dip angle ( $\delta$ ) and slip (rake) angle ( $\lambda$ ). An instantaneous tsunami generation with a constant distribution of the slip is assumed. A simulation of the numerical model provides the wave amplitude time series to be used as input for the hybrid model.

## 2.3 Numerical flume geometry

15 The dimensions of the numerical flume vary with the profile characteristics, adapting the domain for each simulation. The geometry of the flume is shown in Fig. 4. The total length  $L$  of the flume is split in two components:  $L_{off}$  is the submerged part of the profile and  $L_i$  is the inland part of the profile.



**Fig. 4. Numerical flume geometry, including the 5 parameters that define each profile ( $\tan \beta_1$ ,  $\tan \beta_2$ ,  $\tan \beta_3$ ,  $d_1$  and  $d_2$ ) and the general location of  $x_{cut}$ , where numerical models are coupled**

- 5  $L$  is determined for each simulation according to the profile parameters ( $\tan \beta_1$ ,  $\tan \beta_2$ ,  $\tan \beta_3$ ,  $d_1$  and  $d_2$ ), and the tsunami wave length is  $\lambda = T \cdot \sqrt{g \cdot h}$ , where  $T$  is wave period and  $g$  is the gravitational acceleration:

$$L = L_i + L_{off} \quad (1)$$

$$L_i = \frac{50}{\tan \beta_0} \quad (2)$$

$$L_{off} = L_f + x_2 + x_1 \quad (3)$$

$$L_f = \left\lceil \frac{1.2 \cdot \lambda}{10 \cdot \Delta x} \right\rceil \cdot \Delta x \quad (4)$$

$$x_2 = \frac{d_1}{\tan \beta_1} + \frac{d_2 - d_1}{\tan \beta_2} \quad (5)$$

where  $\Delta x$  is the resolution (cell size) of the simulation with the COMCOT numerical model, as described in detail in the next section. In Equation 4,  $\Delta x$  is in both denominator and numerator to round  $L_f$  to the order of  $\Delta x$ , by means of “ceiling brackets”.

- 5 The IH2VOF domain is located in the shallowest part of the profile, with a sufficient area of the inland domain to obtain an accurate measurement of run-up and an area as long as possible for the wave propagation. The length covered with RANS model was between 500 m and 25000 m.

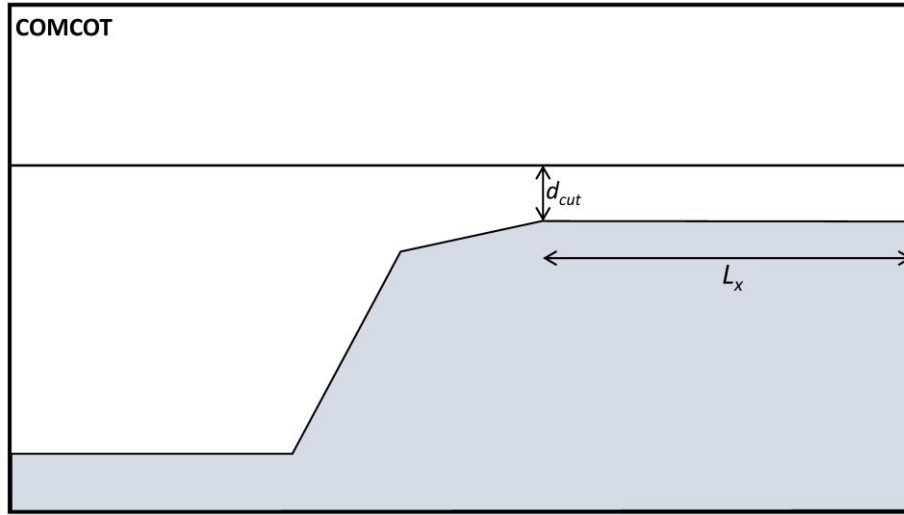
## 2.4 Numerical models coupling

- The coupling of the numerical models was focused on accurately locating the border position between the models,  $x_{\text{cut}}$  (see Fig. 4). This location is optimized in the domain of the IH2VOF model for every tsunami scenario, since that area is the most computationally demanding. Two criteria are followed for this optimization: 1) maximize the area of the IH2VOF domain and 2) simultaneously ensure that the flooding does not exceed the inland end of the IH2VOF domain.

- To achieve this optimization, it is necessary to know a rough value of the run-up in advance in order to fit the inland part of the grid. In this sense, the more accurate the rough estimate of the run-up is, the fewer inland cells are wasted (without flooding), meaning that the IH2VOF performance is optimized. Clearly, the complete run-up must be fully covered by this model domain, meaning that the vertical length of the onshore grid  $L_z$  must be adequate. To determine this horizontal length in advance, each simulation is precalculated with only COMCOT to obtain an approximation of the run-up of the considered tsunami scenario.

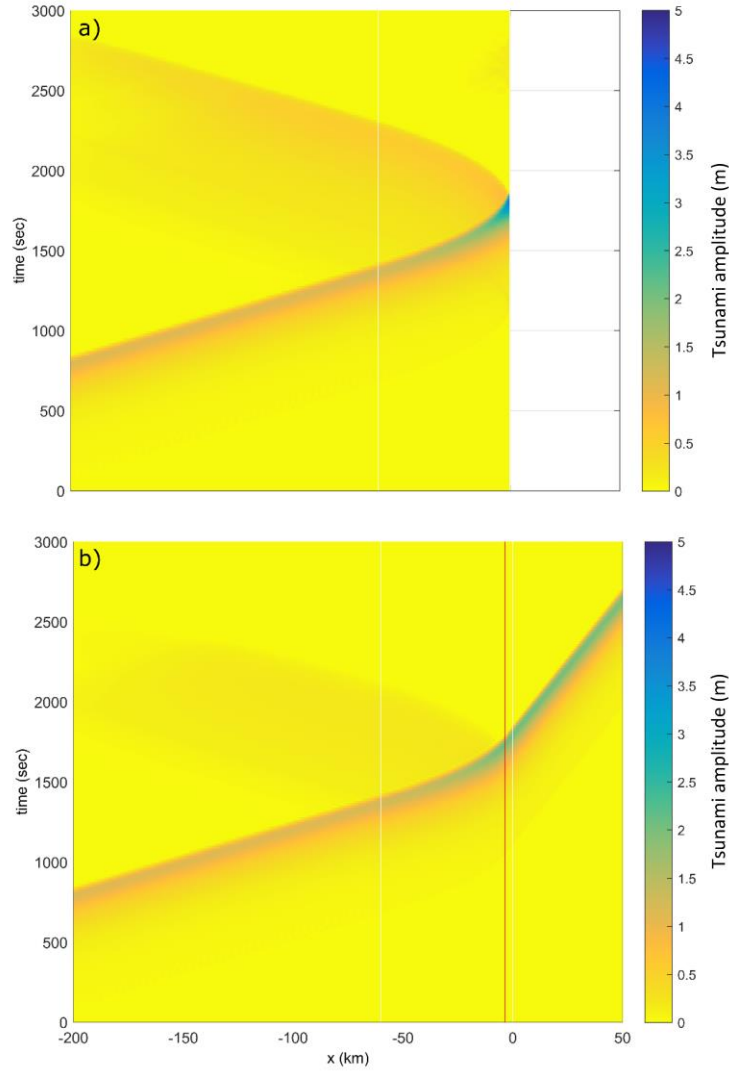
- The transference of data between models occurs at  $x_{\text{cut}}$ . IH2VOF requires as input a time series of sea surface deformation and a velocity profile; hence, these data are obtained as an output of the COMCOT model. Nevertheless, in most cases, the tsunami wave length  $\lambda$  is considerably longer than the length of the IH2VOF grid  $L_x$  ( $\lambda > L_x$ ). Therefore, before the entire wave has passed  $x_{\text{cut}}$ , the reflected wave has already reached back to that point; as a consequence, the amplitude tsunami wave series from COMCOT used in IH2VOF along  $x_{\text{cut}}$  would be "contaminated" with the reflected wave.

- To avoid this situation, a second simulation with only COMCOT is performed for the considered scenario. In this simulation, the topobathymetric profile is the same, but the slope is set to 0 from  $x_{\text{cut}}$ , and the right inshore boundary is left open (see Fig. 5). This approach minimizes the influence of the reflection, allowing the input data that COMCOT transfers to IH2VOF to be accurately obtained at the  $x_{\text{cut}}$  position. Thus, the unaltered wave is used to force the IH2VOF domain, in which simulation reflection on the beach is correctly observed and considered for the run-up calculations.



**Fig. 5. Numerical flume geometry with a modified profile to avoid the reflection phenomena on the  $x_{\text{cut}}$  position**

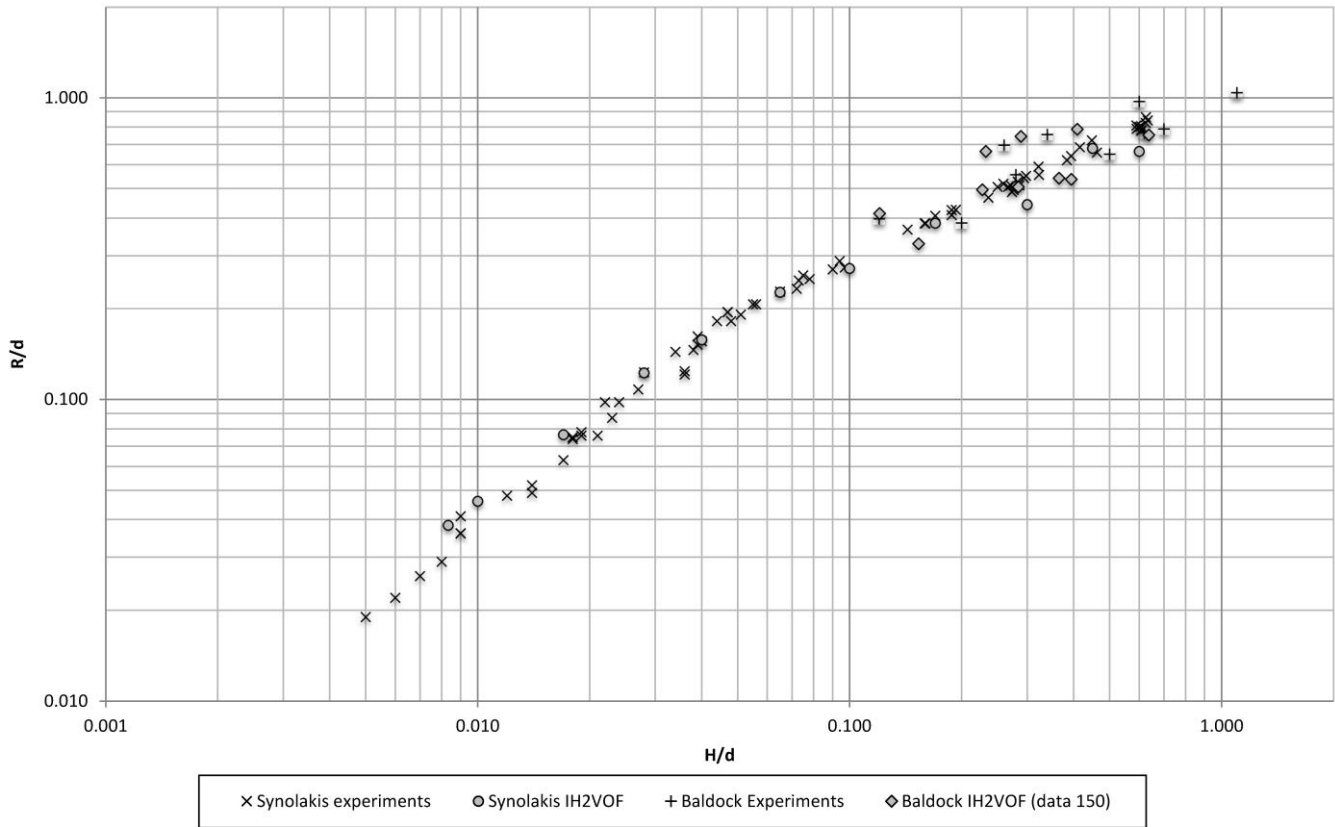
To attest the effectiveness of this approach, an example of the wave height propagation of a tsunami wave with  $H=5.6$  m for  $T=40$  minutes in the numerical flume is shown in Fig. 6. Fig. 6a shows the propagation of the wave height in the unaltered flume with the reflection effects, and Fig. 6b shows the same propagation in the modified flume, in which the reflection effects are minimized. In the plots in Fig. 6, the x-axis is along the length of the flume, the y-axis is the time of the simulation, and the  $x_{\text{cut}}$  position is marked as a red line. The example wave enters the flume after 10 minutes (600 seconds) of simulation and propagates towards the coast (zero on the x-axis), reaching the  $x_{\text{cut}}$  position after 30 minutes (1800 s) of simulation. At the  $x_{\text{cut}}$  position, a reflected wave on the order of 1 m height is reduced by 95%, making it possible to obtain the boundary conditions for the IH2VOF simulation.



**Fig. 6. Propagation of a tsunami wave in the numerical flume on the distance-time plane.  $x_{\text{cut}}$  position is indicated on the plane with a red line. a) In the initial domain, a reflected wave is observed, aliasing the boundary condition for the IH2VOF domain. b) The modified domain, in which a constant infinite depth is modeled after the  $x_{\text{cut}}$  position. In this case, the reflection of the wave is eliminated, allowing the boundary conditions for the IH2VOF model to be obtained**

vi) COMCOT simulation to obtain the boundary conditions (input) for the IH2VOF simulation; vii) IH2VOF simulation and viii) run-up determination in the IH2VOF domain.

The validation of the numerical flume and the coupling of the models was made by comparing its results with the results of the physical experiments conducted by (Synolakis, 1987) and (Baldock et al., 2009). The scenarios defined on these experiments were calculated using the complete numerical flume and the results were compared to the results of the physical experiments. Fig. 7 shows the run-up obtained in this comparison and how the results of both series of experiments fit adequately with the numerical flume results.



**Fig. 7. Validation of the numerical flume by comparing the results of its application with the results of the physical experiments of (Synolakis, 1987) and (Baldock et al., 2009)**

To assess the tsunami hazard in a tsunami-prone area, this coupled model can be applied to several profiles all along the studied coastal area. The methodology provides the run-up at each of the profiles, allowing the flooded area to be estimated as an envelope of the run-up limits.



### 3. Application example: further development of a tsunami run-up database

The presented hybrid model was created with the aim of applying it to generate a tsunami run-up database (TRD) from a combination of bathymetric profiles and tsunami waves. The objective of this database is to create an interpolation space that allows the instant evaluation of the tsunami run-up, without needing to run numerical simulations.

- 5 This database contains the run-up of tsunami scenarios that are combinations of parameterized tsunami waves and parameterized bathymetric profiles. These scenarios have been simulated within the described numerical flume.

The following section is focused on explaining the details of the development of the TRD, the selection of the bathymetric profiles and tsunami waves, and the simulations run with the hybrid numerical model. Finally, an interpolation tool, which was developed ad hoc to apply the database to the instantaneous estimation of the run-up, is presented.

#### 10 3.1 TRD bathymetric profiles

The parameterization of the 50 bathymetric profiles that were selected worldwide (see Fig. 1 and Table 1) were added to the TRD. The 5 parameters of each of these 50 measured profiles were obtained by means of a least-squares fitting method.

To increase the number of cases included in the TRD, more realistic profiles were added as combinations of the 5 parameters.

- 15 To generate these profiles, the ranges of the values of each parameter over the 50-profile sample set were analyzed, with a focus on identifying trends or rules that characterize their variability. The new profiles follow these trends, avoiding the inclusion of unrealistic combinations of parameters (e.g.,  $(d_2 - d_1)$  was always larger than 2200 m and  $x_l$  was always shorter than 210 km). By using these realistic combinations of parameters, the TRD was expanded to 5000 profiles.

- 20 Finally, from those 5000 profiles, a selection of 49 profiles was made by means of the maximum dissimilarity algorithm (Camus et al., 2011). These 49 profiles assure a maximum variability in the profiles to further develop the TRD. The parameters of the 49 chosen profiles are given in Table A1 of Annex A.

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#### 3.2 TRD tsunami wave parameterization

The tsunami waves for the TRD were obtained by means of simulations of realistic scenarios using the COMCOT model, applying a grid size of  $\Delta x=500$  m and satisfying the CFL condition given by the model ( $\frac{C\Delta t}{\Delta x} < 0.5$ , where  $\Delta t$  is the time step of the simulation). Based on these simulations, the tsunami waves were characterized by 2 parameters: tsunami wave height ( $H$ ) and period ( $T$ ) at the depth  $d_2$ .

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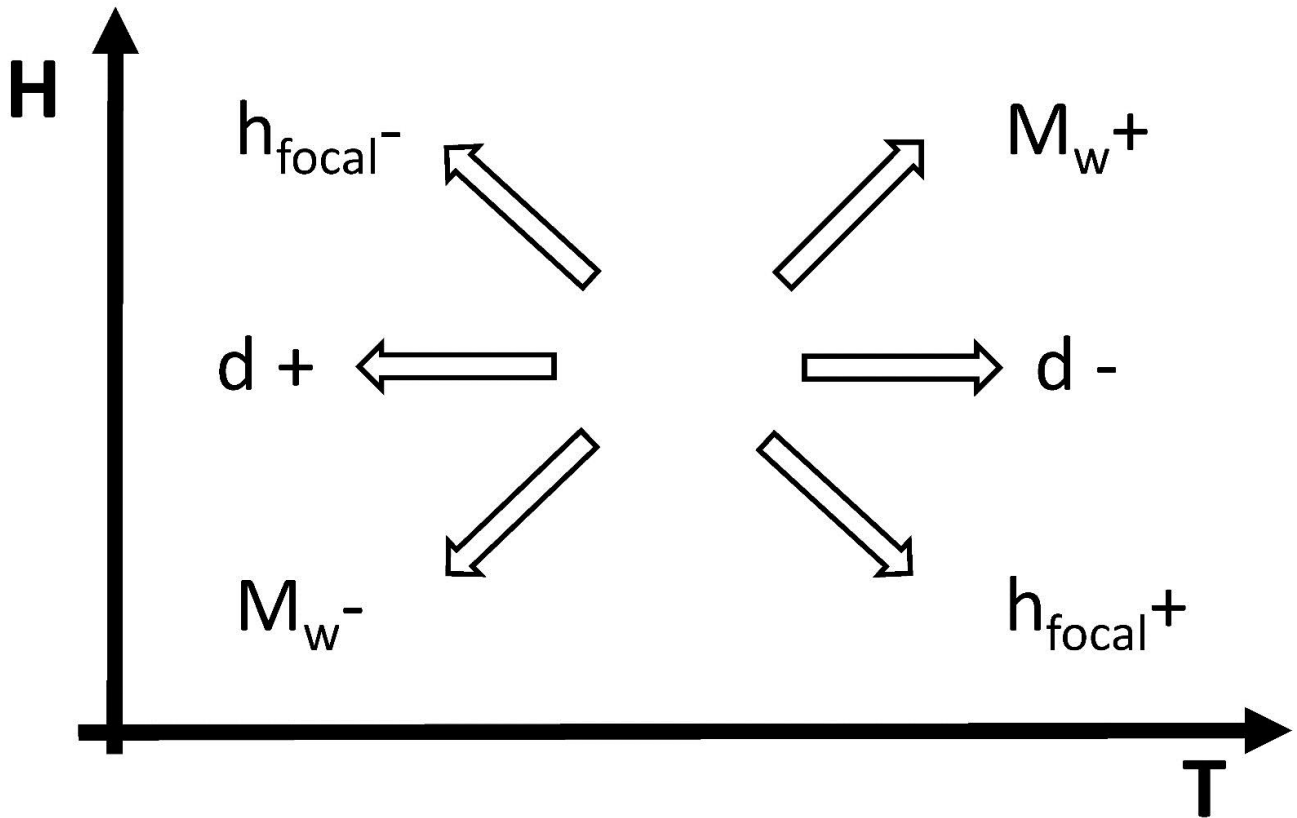
To generate the tsunami wave shapes with COMCOT, an infinite horizontal domain with a constant water depth was used. In the analyses of the seven focal mechanism parameters (see 2.2), some simplifications were assumed. First, to generate higher tsunami waves, the three angles were fixed to the combination that provides the maximum tsunami height, i.e.,  $\theta=0^\circ$ ,  $\delta=90^\circ$  and  $\lambda=90^\circ$ . Second,  $D$  was defined in Table 2, and a width of  $W = (M_0/6.25\mu\gamma)^{1/3}$  is obtained by assuming a rectangular fault with proportion  $L/W = 2.5$  and using the  $M_0$  formulation (Table 2). Finally, Kanamori and Anderson (1977) provided a relationship between the seismic moment and earthquake magnitude:  $M_w = 2/3 \cdot \log(M_0) - 6.07$ .

**Table 2. Formulations used in the definition of the parameters of the tsunami waves included in the addition to the database**

Formulation	Definitions	Source
$D=\gamma L$	$\gamma=6.5\times 10^{-5}$	Scholz and Harris (2002)
$M_0=\mu SD$	$S=LW$ ; $\mu=2.5\times 10^{11}$ Aki (1972)	Steketee (1958); Burridge, R. and Knopoff (1967)

10 Taking these parameters into account, the tsunami waves can be obtained in terms of earthquake magnitude ( $M_w$ ), focal depth ( $h_{focal}$ ) and water depth ( $d$ ). Therefore, the influence of these three parameters on the tsunami wave height and period was explored and depicted in a general scheme in Fig. 8. As it could be intuitively expected, the higher the magnitude of the earthquake is, the higher the tsunami wave height; however, the deeper the focal depth is, the lower the tsunami wave height. On the other hand, regarding the tsunami wave period, the period increases with the earthquake magnitude or focal depth. The

15 water depth in the rupture area affects only the tsunami period, which increases when this depth decreases.



**Fig. 8. General scheme of the tsunami wave height ( $H$ ) and period ( $T$ ) behavior in relation to the water depth ( $d$ ) in the generation area and earthquake focal depth ( $h_{focal}$ ) and magnitude ( $M_w$ )**

- 5 Following this characterization, a set of tsunami waves was selected, covering period values from 5 to 40 minutes and wave height values in the source area (depth  $d_2$ ) from 0.2 to 2.0 m. In generation areas, tsunami waves are commonly within these ranges (Papadopoulos, 2016). The considered waves are peak or positive waves, meaning that the wave height was considered from the sea level. The tsunami wave heights (from A to K) and periods are given in Table 3. The interpolation is limited to the  $H$  and  $T$  ranges included in the database. The incorporation of new waves will complement the existing wave database and
- 10 increase the range of application of the methodology.

**Table 3. Tsunami wave height and period at the source area for the current events included in the TRD**

Name	Height (m)	Period (mins)
A	1.6	35
B	0.5	35
C	0.5	8

<b>D</b>	0.2	5
<b>E</b>	1.5	15
<b>F</b>	0.5	15
<b>G</b>	1.5	25
<b>H</b>	0.5	25
<b>I</b>	1.0	35
<b>J</b>	1.0	15
<b>K</b>	1.0	25

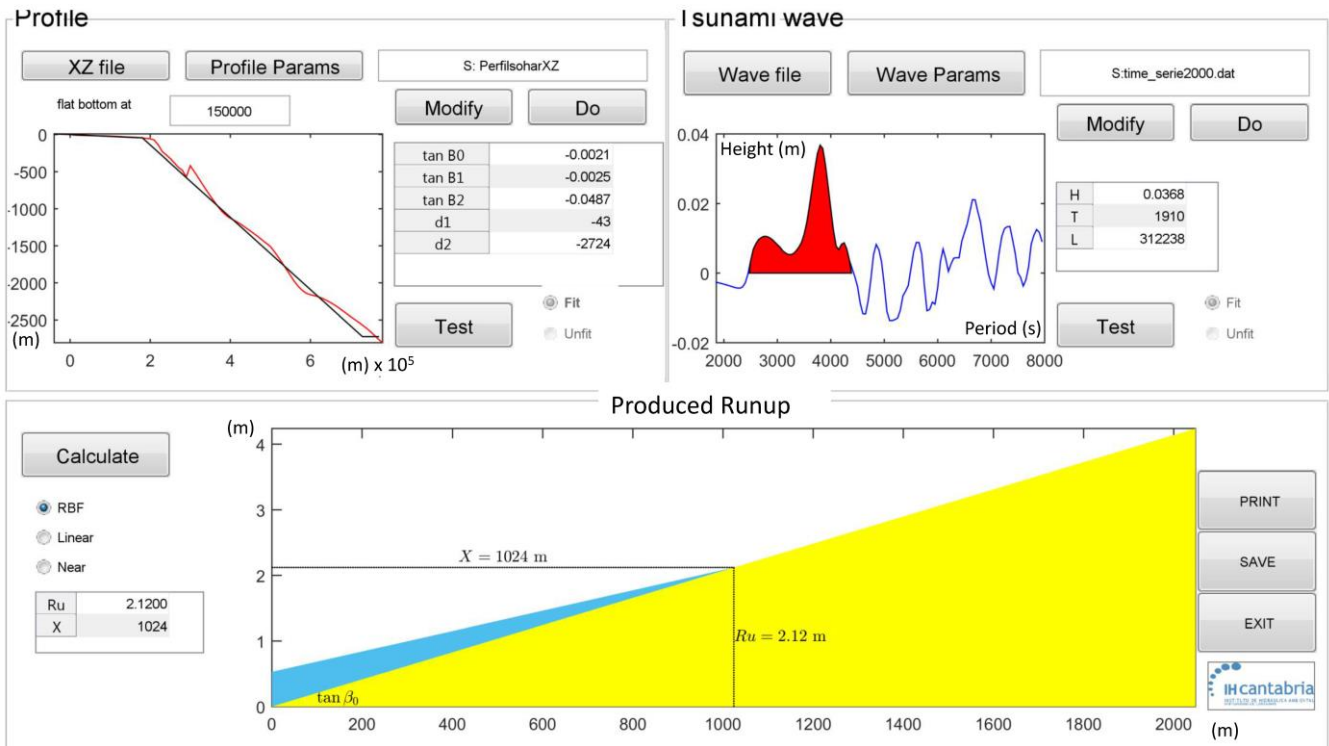
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#### 4. Run-up estimation by interpolation of the TRD

The procedure explained at the end of section 2.4 was followed to calculate the run-up of the combination of the tsunami waves and the 49 topobathymetric profiles. Therefore, the TRD is increased by 539 scenarios, provided by 7 parameters ( $\tan\beta_0$ ,  $\tan\beta_1$ ,  $\tan\beta_2$ ,  $d_1$ ,  $d_2$ ,  $H$  and  $T$ ). These simulated scenarios constitute the 7-dimension interpolation domain in which new run-up calculations are carried out. The interpolation procedure and the result of its application are described next. A tool to calculate the run-up by interpolation was developed. This tool allows the simultaneous analysis of the influence of each parameter on the final value of the run-up.

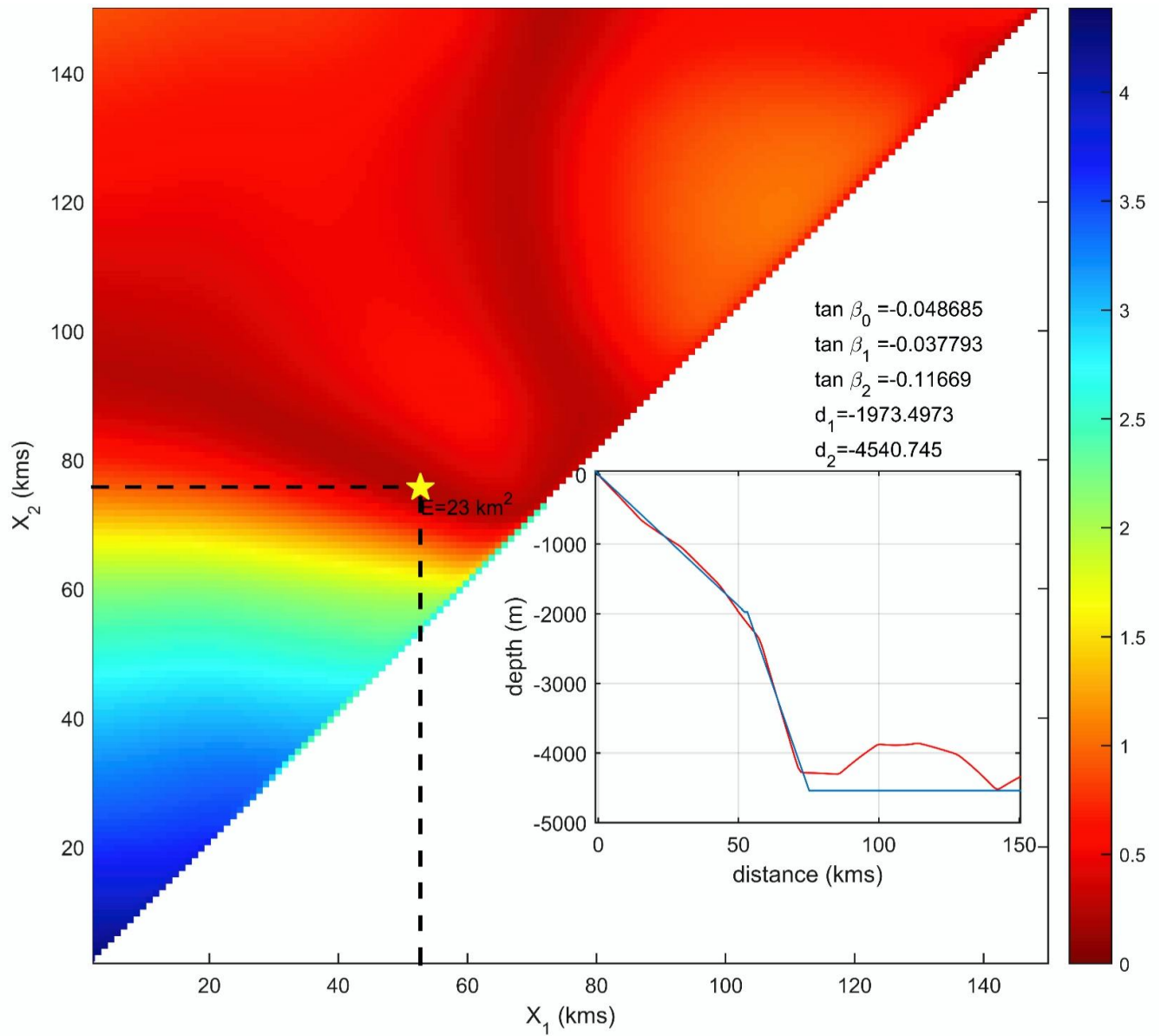
##### 4.1 The interpolation tool (IH-TRUST)

To interpolate the value of the run-up for new scenarios, a numerical tool was programmed. This tool, called *Instituto Hidráulica-Tsunami Run-Up Simulation Tool* (IH-TRUST), processes the profile and wave data to calculate the run-up, and it performs an interpolation of the 7 parameters considered in the TRD. IH-TRUST consists of three modules, or elements (Fig. 9).



**Fig. 9. IH-TRUST interface, showing its 3 elements: the bathymetric profile, the tsunami wave parameterization and the run-up calculation**

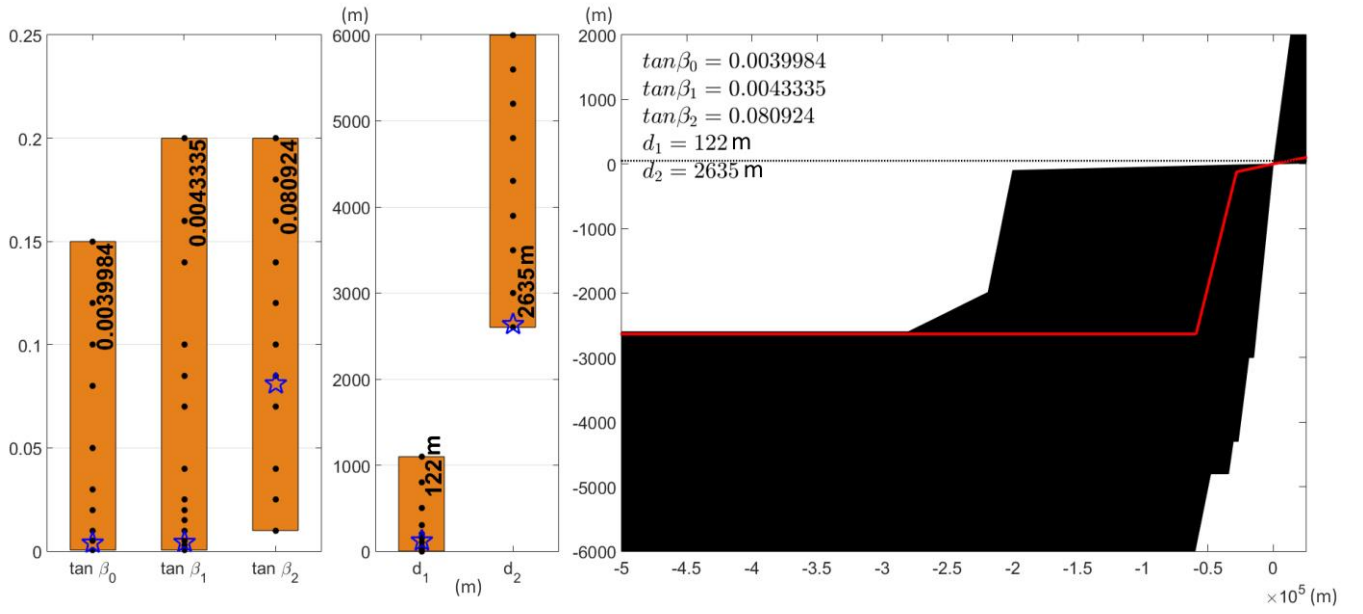
- 5 In the first element, the tool calculates the parameterization of the real topobathymetric profile into 5 parameters:  $\tan\beta_0$ ,  $\tan\beta_1$ ,  $\tan\beta_2$ ,  $d_1$ , and  $d_2$ . **The real profile ideal trace should be parallel to the potential tsunami wave travel.** The parameterization is approached by means of a least-squares fitting. An example of the fitting of a profile is shown in Fig. 10, in which the main plot shows the quadratic error in terms of distances from the coast to  $d_1$  ( $X_1$ ) and to  $d_2$  ( $X_2$ ), and the star indicates the minimum error, which is consequently the position of the best set of five parameters. The subplot shows the original profile and the parameterized profile. Afterwards, the tool verifies that the parameterized profile is included in the ranges of the parameter values contained in the interpolation space.
- 10



**Fig. 10. Parameterization of the bathymetric profiles.** The figure maps the  $E_{rms}$  values as a function of the distance from the coast to  $d_1$  ( $X_1$ ) and the distance from the coast to  $d_2$  ( $X_2$ ) obtained during the process of finding the best parameters for a profile.

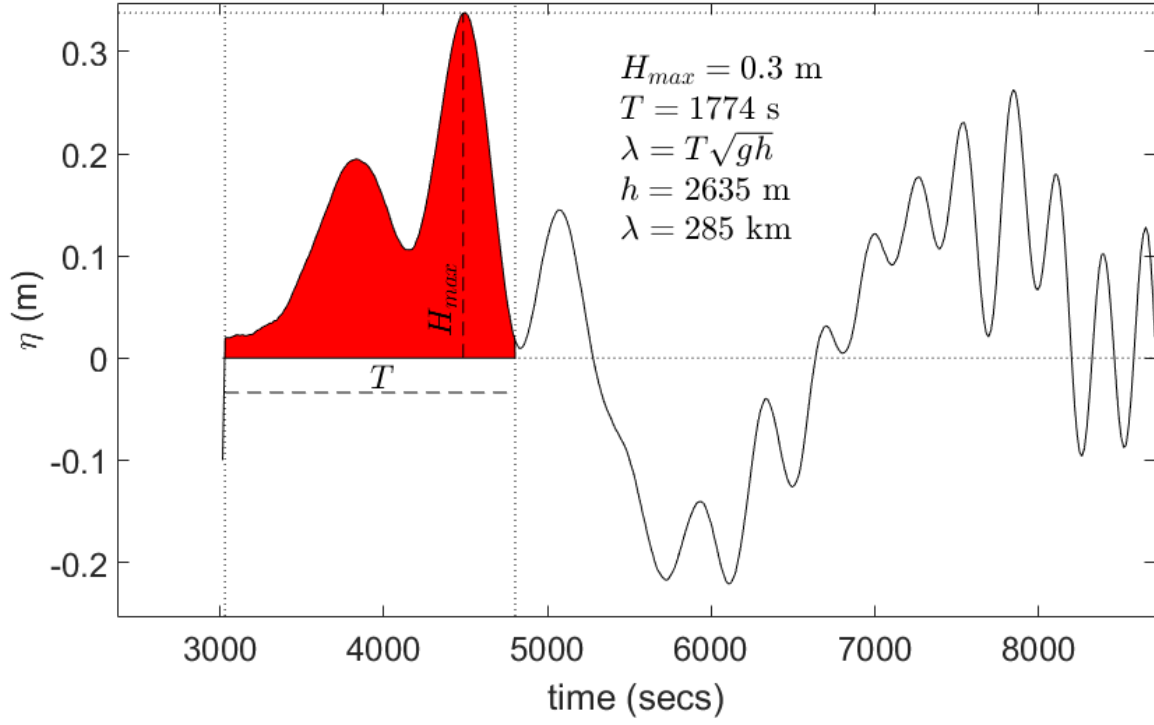
In the subplot, the original profile and the parameterization are shown

- 5 Fig. 11 shows a representation of all the profile domains in black and the introduced profile in red. A set of bars indicates the acceptable values for each parameter, and a star marks the position of each parameter for the new profile.



**Fig. 11. Fitting of a topobathymetric profile in the TRD. Each parameter value (left) and parameterized profile shape (red) for the profiles are included in the TRD (black)**

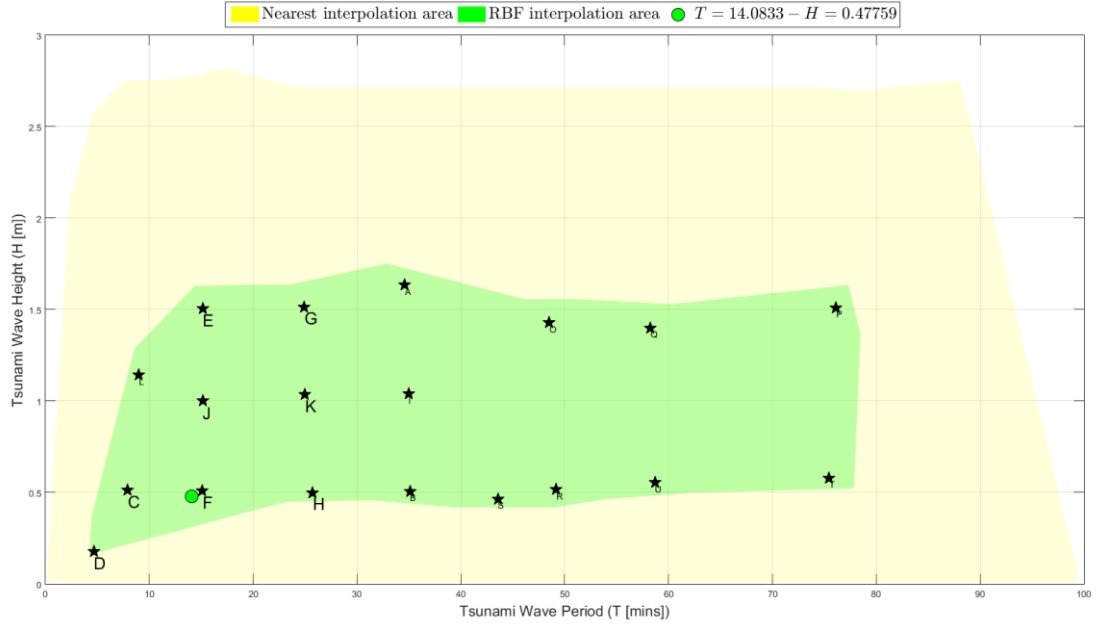
- 5 In the second element, IH-TRUST calculates the values of the  $H$  and  $T$  of the tsunami wave to be assessed at depth  $d_2$ . The tool reads a time series containing the tsunami shape and calculates  $H$  and  $T$ .  $T$  corresponds to the time between the first two zero crossings for positive heights, and  $H$  is the maximum wave height observed within period  $T$ , but the tool allows to manually set the height and period within the time series, if desired (Fig. 12). This is especially useful when the wave shape is not standard, the wave has a leading depression or, simply, it is not the largest wave.



**Fig. 12. Tsunami wave profile parameterization in the IH-TRUST tool, including the part of the tsunami time series considered (in red) to calculate the period  $T$  and the height  $H$**

After the wave parameters are calculated, IH-TRUST checks if the tsunami wave fits in the interpolation domain of the database. Fig. 13 shows the tsunami waves included on the database, the area where the interpolation is valid and the position of the tsunami wave that is being studied.





**Fig. 13. Tsunami wave height and period cases included in the database. The point corresponding to any new wave should fall in the green shadow in order for it to be able to be interpolated with the generated TRD**

- 5 Finally, in the third element of IH-TRUST, the results of the calculation of the run-up  $Ru$  is given based on the profile parameters and the tsunami wave. The interpolation (Fig. 9) is calculated by means of the RBF (Camus et al., 2011) and linear and nearest interpolation methods. In addition, the horizontal flooding distance  $X$  is calculated using the inland slope. The tool uses an RBF interpolation by default, but the nearest or linear methods are also available, since they are useful to calculate events that plot closer to the boundary of the valid interpolation area.

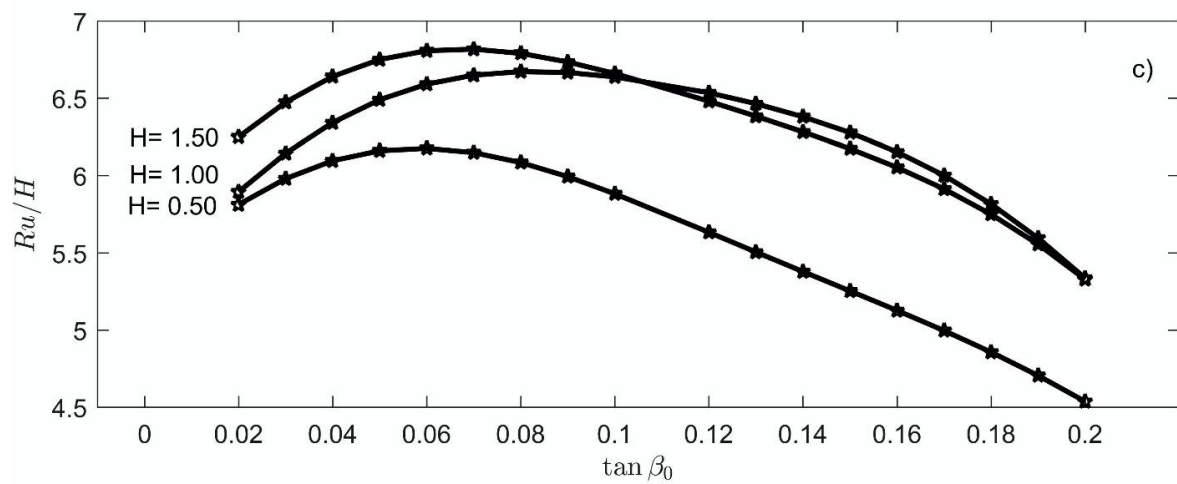
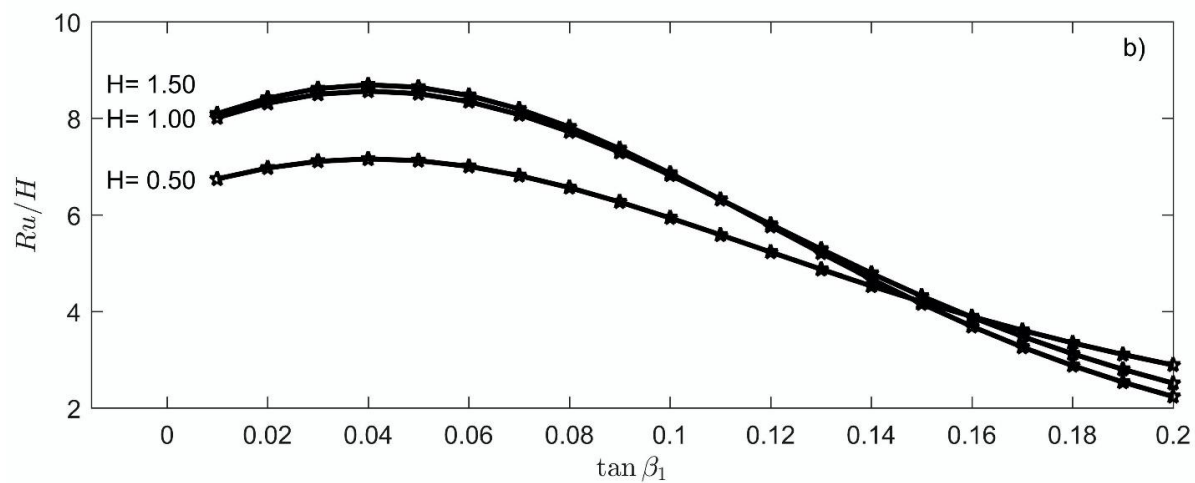
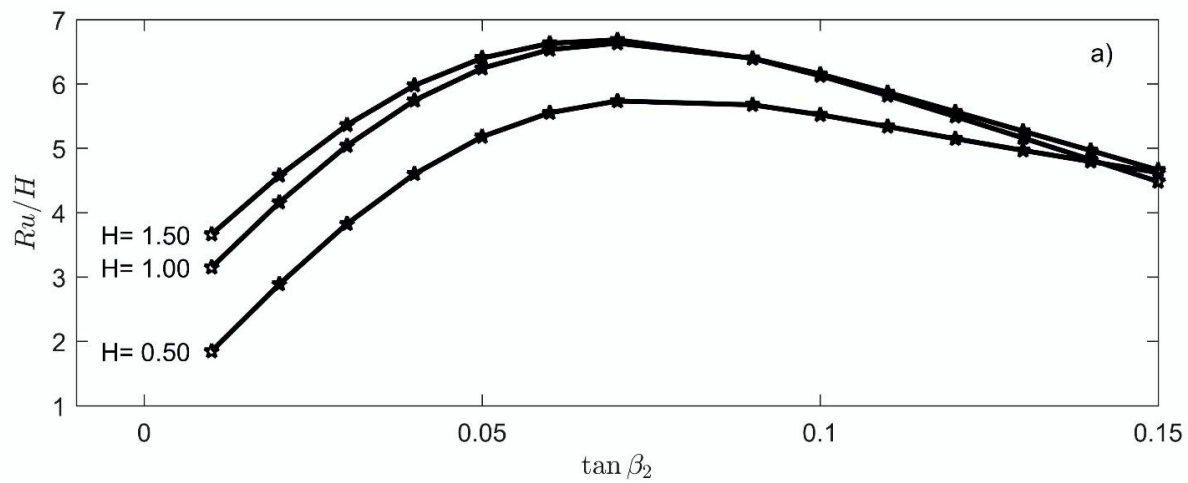
10

## 4.2 Influence of the profile parameters on the tsunami run-up

- The TRD and IH-TRUST were used to explore and analyze the influence of each parameter of the profile on the final value of the run-up. This analysis was approached by evaluating scenarios in the TRD. Although it is out of the scope of this paper, to understand the influence of each parameter of the bathymetric profile, several tests were conducted with a mean profile
- 15 ( $\tan\beta_{0m}=0.080$ ,  $\tan\beta_{1m}=0.09$ ,  $\tan\beta_{2m}=0.110$ ,  $d_{1m}=500$ , and  $d_{2m}=4350$ ) and by varying only one of the 5 parameters that define the profile at a time; additionally, several values of  $H$  and  $T$  inside the boundaries of the domain were considered.

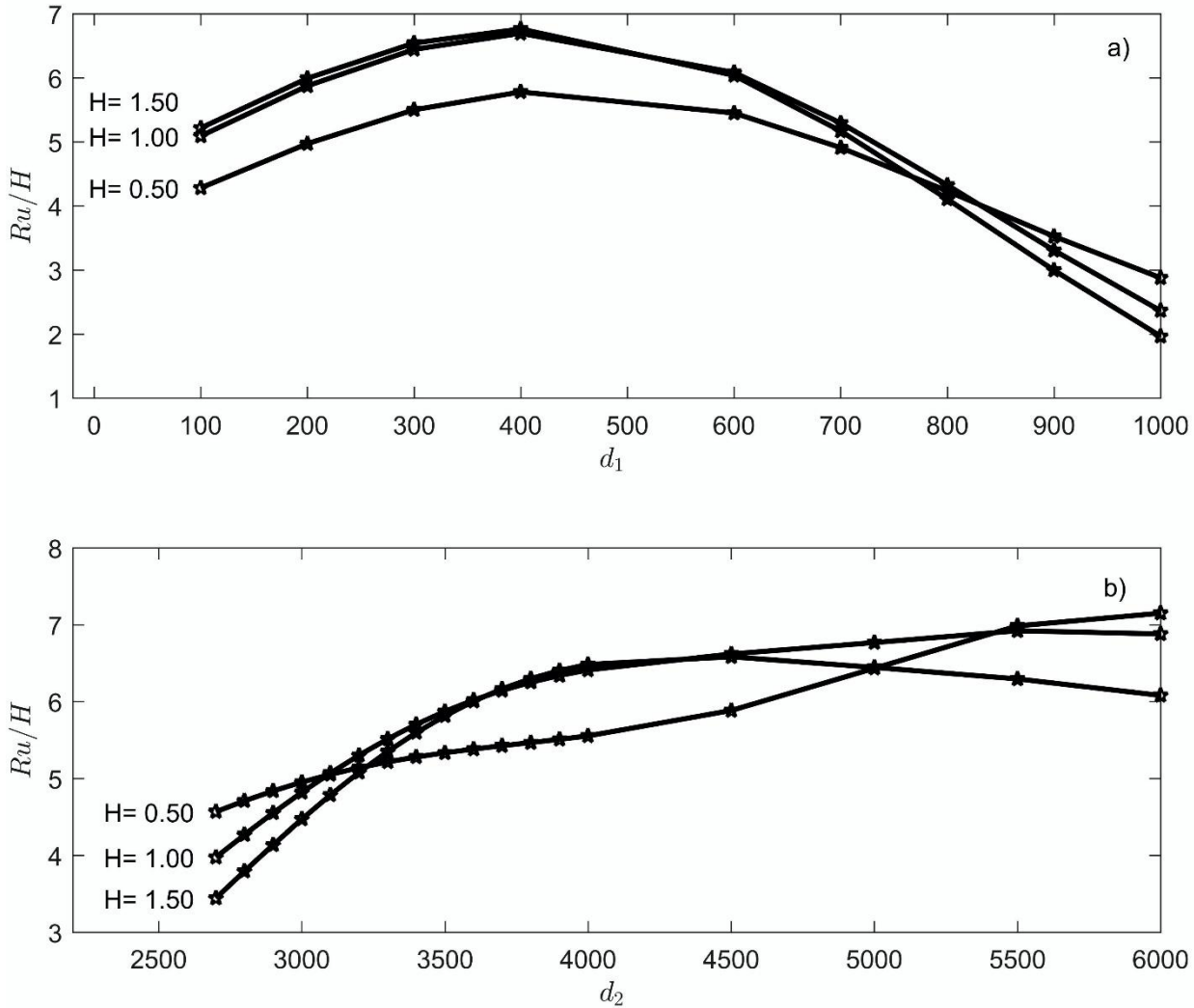
For each pair of values of  $H$  and  $T$ , 4 of the 5 profile parameters were kept constant, and the run-ups were calculated using IH-TRUST with the TRD by varying the 5<sup>th</sup> parameter.

The effect of the variation in  $Ru/H$  as a function of the parameters are shown in Fig. 14 and Fig. 15.



**Fig. 14. Maximum run-up (normalized by wave height) as a function of the profile slopes ( $\tan\beta_0$ ,  $\tan\beta_1$ , and  $\tan\beta_2$ ) of the parameterized profiles for different wave heights.**

The continental slope effect (Fig. 14), parameterized as  $\tan\beta_2$ , produces a maximum  $Ru$  when  $\tan\beta_2$  is close to  $\tan\beta_1$ , reproducing a single slope profile. For smaller  $\tan\beta_2$  values,  $Ru$  decreases rapidly due to wave shoaling. Low values of  $\tan\beta_2$  also indicate a large platform with a low slope, where the shoaling increases the wave height and the wave energy diminishes gradually due to bottom friction until wave breaking occurs. Thus, the energy flux that reaches the shore decreases with the run-up height. The profile typology characterized by a low value of  $\tan\beta_2$  is closer to Synolakis's canonical problem.



**Fig. 15. Maximum run-up heights as a function of the profile slopes ( $d_1$  and  $d_2$ ) of the parameterized profiles for different wave heights**

The higher the  $\tan\beta_2$  is, the shorter the platform, reducing the energy dissipation and allowing the slopes to have similar  $\tan\beta_0$  and  $\tan\beta_1$ ; this maximizes the run-up height.

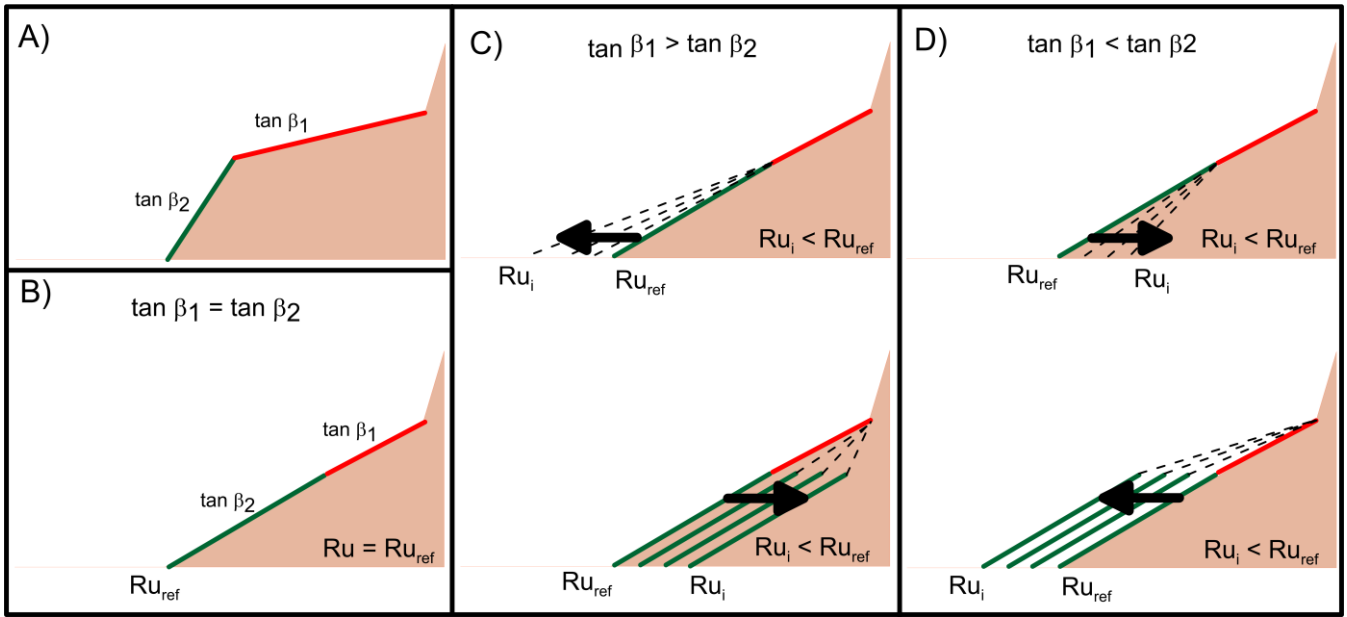
5 Regarding  $\tan\beta_1$ , when  $d_1$  is constant (Fig. 14), the higher  $\tan\beta_1$  is, the shorter the length of the shelf, reducing tsunami wave shoaling. In this case, the wave steepness increase drastically near the coast and breaks abruptly, triggering a considerable dissipation of energy within a short length; this effect reduces both the energy flux on the coast and the run-up.

10 K noğlu and Synolakis, (1998) shown that in the case of solitary waves the slope closest to the coast controls the run-up processes. In the case of tsunamis, in our geometry, although the influence of  $\tan\beta_1$  is indeed important, the influence of  $\tan\beta_2$  must not be neglected.

Finally, the influence of  $\tan\beta_0$  on the final value of the run-up is less important than those of  $\tan\beta_1$  and  $\tan\beta_2$ . The run-up decreases as  $\tan\beta_0$  increases. Due to the effect of gravity, the flow ascends less if greater slopes are present. This aspect is strengthened by the reflection of the energy.

15 The behavior described above can be summarized on three basic regimes, depicted in **¡Error! No se encuentra el origen de la referencia..** i) The larger run-up for any wave height is found if  $\tan\beta_1 = \tan\beta_2$  (see **¡Error! No se encuentra el origen de la referencia.b**). ii) When  $\tan\beta_1 > \tan\beta_2$  see **¡Error! No se encuentra el origen de la referencia.c**, the bigger the difference between slopes, the smaller the run-up run-up mainly because the increase of dissipation; and iii) when  $\tan\beta_1 < \tan\beta_2$ , see **¡Error! No se encuentra el origen de la referencia.d**, the run-up is also smaller, but in this case it is due to the effect of increasing reflection.

20



**Fig. 16. Influence of the variation of the  $\tan\beta_1$  and  $\tan\beta_2$  on the run-up.** A) shows a scheme of the profile where segment of  $\tan\beta_2$  is marked as green and  $\tan\beta_1$  as red. B) shows the situation where  $\tan\beta_1 = \tan\beta_2$ . C) depicts the situations where  $\tan\beta_1$  is larger than  $\tan\beta_2$ . And, finally, D) shows the situations where  $\tan\beta_1$  is smaller than  $\tan\beta_2$ .

5

The influence of depths  $d_1$  and  $d_2$  is shown in Fig. 15. For deeper continental shelf depths  $d_1$ , the shelf is wider and, consequently, the bottom friction affects the wave over a longer profile, creating a run-up smaller. For a constant  $\tan\beta_1$ , lower values of  $d_1$  represent a shorter continental shelf, and abrupt and dissipating wave breaking. Moderate values of  $d_1$  are characterized by a gradual tsunami wave shoaling, during which the bottom friction allows a maximum run-up. From that critical point, higher values of  $d_1$  mean a longer continental shelf, generating a larger frictional area, reducing the energy flux that reaches the shore and consequently diminishing the run-up.

10

In Fig. 15b, it can be observed how the run-up increases almost linearly with  $d_2$ . The effect of  $d_2$  in the run-up is similar to the effect of  $\tan\beta_2$ . The shallower  $d_2$  is, the greater the shoaling and the higher the wave. The wave energy diminishes gradually due to bottom friction until wave breaking, which depends on the tsunami wave height. In addition, it was found that although the variations in wave height produce different  $Ru/H$  values for the same profile, the influence of the variation in the wave period is negligible. Therefore, different wave heights but not different periods are shown in Fig. 14 and Fig. 15.

15

Finally, these results highlight the importance of using an accurate geometry to define the run-up. The influence of  $d_2$  and  $\tan\beta_2$  in the final run-up estimation is considerable, and the use of complete profiles, from the generation area to the coast, is necessary but not considered in traditional approaches and simplifications.

20

## 5. Validation of the methodology with numerical test results and observational data

The methodology presented here aims to calculate the tsunami run-up in coastal areas. This calculation can be applied to study the run-up of historic events but also to calculate the run-up of potential scenarios, which are the primary focus. These potential cases are used to evaluate tsunami hazard and the flooded area when a tsunami occurs. As mentioned in the introduction, run-up is commonly assessed by means of high resolution data.

Therefore, to validate this methodology/tool, the results of its application have been compared with both high-resolution numerical simulations of potential events and historical tsunami run-up scenarios. It is important to highlight that the presented tool allows to estimate tsunami run-up when these HR data are not available or accessible, what is a common situation.

The results of these comparisons are detailed in the following subsection, which is focused on describing the strengths and limitations of the methodology for each case.

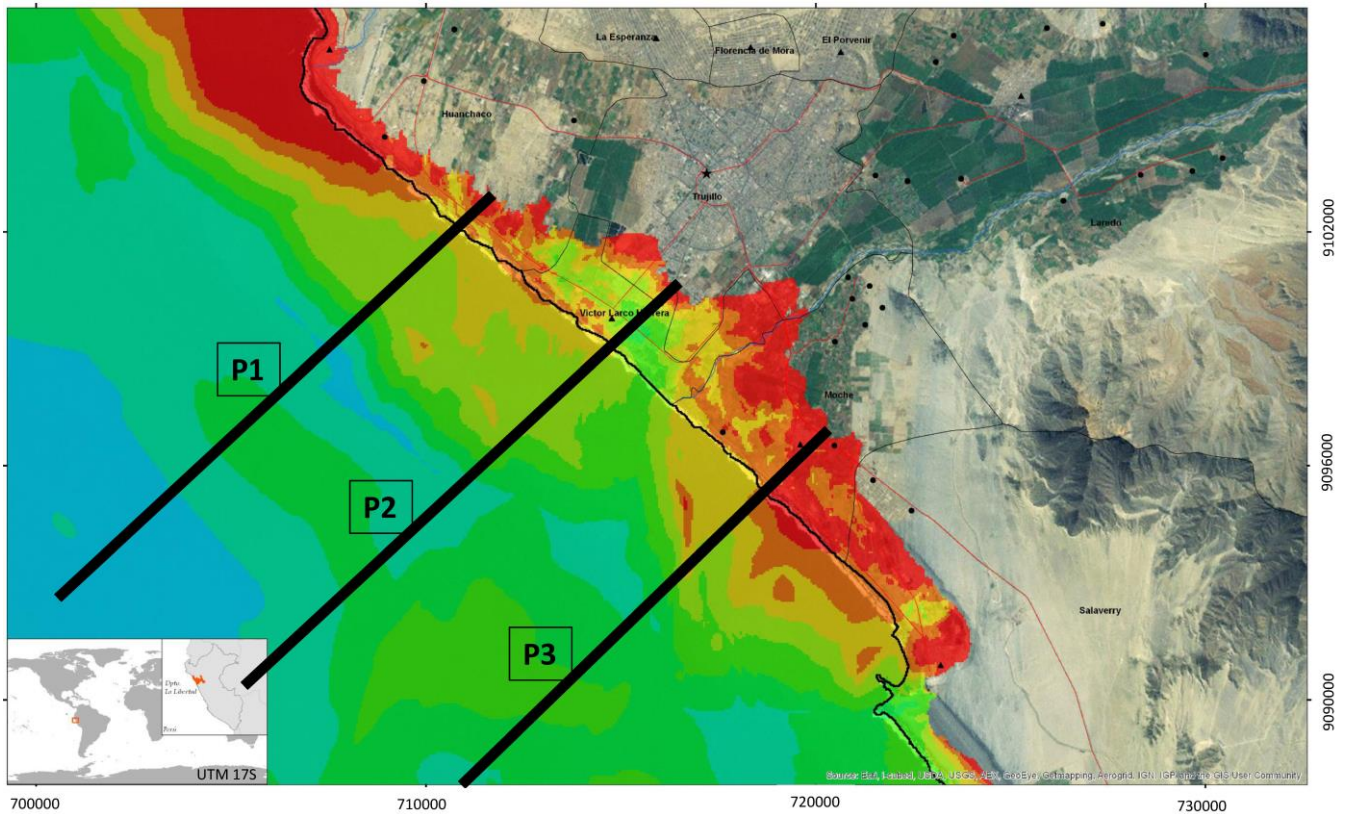
### 5.1 Validation with numerical model simulations

This validation was carried out as follows: first, a topobathymetric profile of the study area was obtained using the GEBCO database. On that profile, a point was selected offshore, and the time series of the tsunami was extracted at that point from the COMCOT numerical simulation of the event. Using the topobathymetric profile and the time series as input for the IH-TRUST tool, the run-up was interpolated by using the created database. The interpolated run-up was then compared to the run-up obtained by using the high-resolution numerical simulation of the potential scenario.

Three numerically simulated scenarios with high resolutions have been selected for the validation. All these scenarios are from real projects, studies and published papers that were focused on analyzing and assessing the tsunami hazard in coastal areas worldwide and characterizing the potential flooded areas due to tsunami events in the selected zones. These simulations used high-resolution topographic and bathymetric data to construct grids with 30 m cells.

#### 5.1.1 Tsunami scenario in Trujillo, Peru

The results of the application of the methodology were compared to the results of a high-resolution numerical simulation of a magnitude 8.5 event in the subduction zone located along the coast of Trujillo, a municipality in northern Peru. This synthetic scenario represents the event that occurred in this zone in 1619 and is part of the study *Probabilistic evaluation of the hazard and vulnerability under natural disasters in the metropolitan area of Trujillo*, funded by the Inter-American Development Bank (IHCantabria, 2013). The numerical simulation used a 30-m-resolution grid to accurately calculate the flooded area for a tsunami wave height and period of approximately 1.5 m and 400 s at a depth of 3000 m.

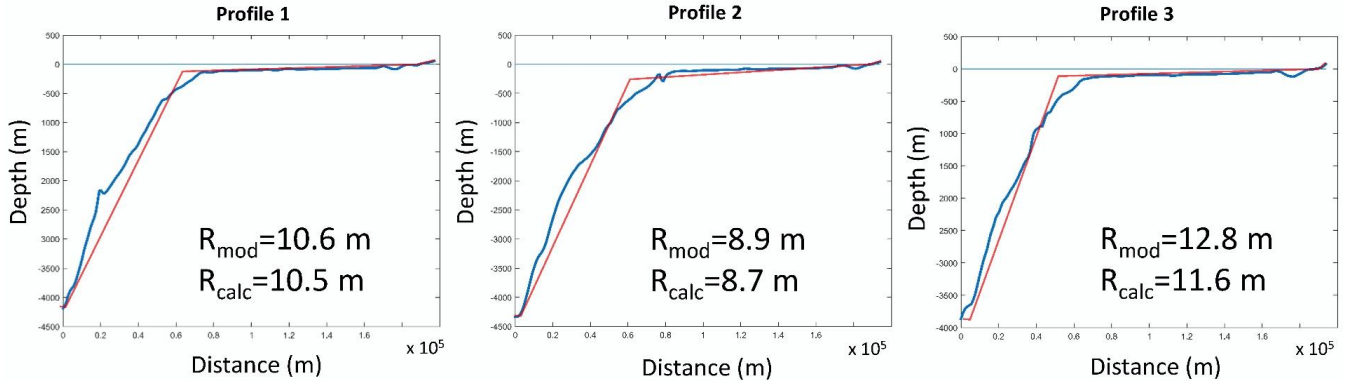


**Fig. 17. Flooded area in the municipality of Trujillo in Peru, due to a tsunami triggered by an 8.5 magnitude earthquake, including 3 selected profiles with the run-up obtained by using the numerical model. The coordinates of the exact locations where the run-up was estimated are provided in Table 4**

5

In Fig. 17, the flooded area map of Trujillo, as well as the selected profiles, are shown. In the study, the numerically calculated run-ups at those profiles (Fig. 18) were 8.9, 10.6 and 12.8 m. The corresponding values for the run-up obtained by interpolating the TRD with the IH-TRUST tool were 8.8, 10.5 and 11.6 m (see Table 4). Compared to the results of the numerical simulation, these 3 values from the 3 zones of the study area provide a good approximation of the tsunami flooding.



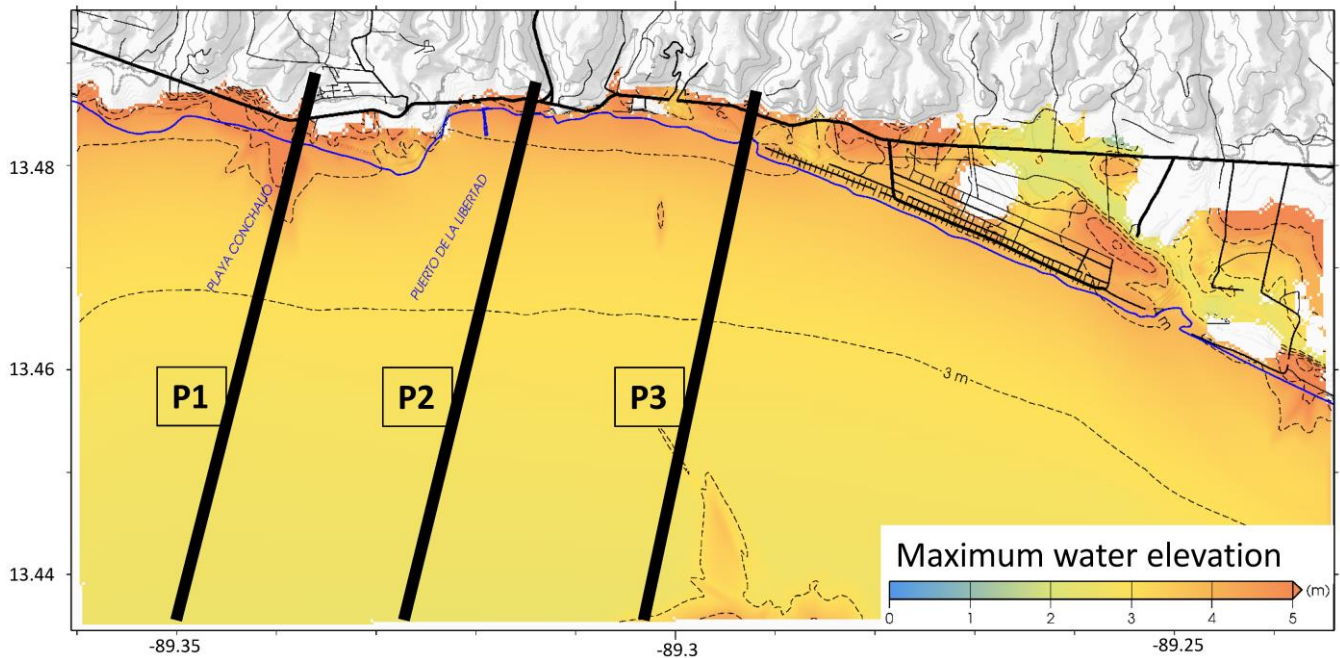


**Fig. 18.** Topobathymetric profiles selected in Trujillo, Peru to validate the methodology. The topobathymetric profile (blue) and the parameterized profile (red) are compared

### 5.1.2 Tsunami scenario in La Libertad, El Salvador

5 Following the same procedure, a validation case was addressed in El Salvador. The event is a potential scenario of an earthquake of magnitude 8.1 along the El Salvador thrust, which is in the subduction zone along the El Salvador coast. The study area is the flat area of La Libertad, on the western side of this Central American country. This high-resolution numerical simulation is part of the project *Tsunami Risk Assessment in El Salvador*, financed by AECID (Spanish Agency for International Cooperation and Development) during the period 2009–2012 (Álvarez-Gómez et al., 2013). The resolution of

10 the numerical simulation was 30 m, and the grid that was built for the propagation and inundation calculations used data from local bathymetric campaigns and high-resolution topographic studies. The tsunami wave height and period at a depth of 3000 m were approximately 0.9 m and 700 s.

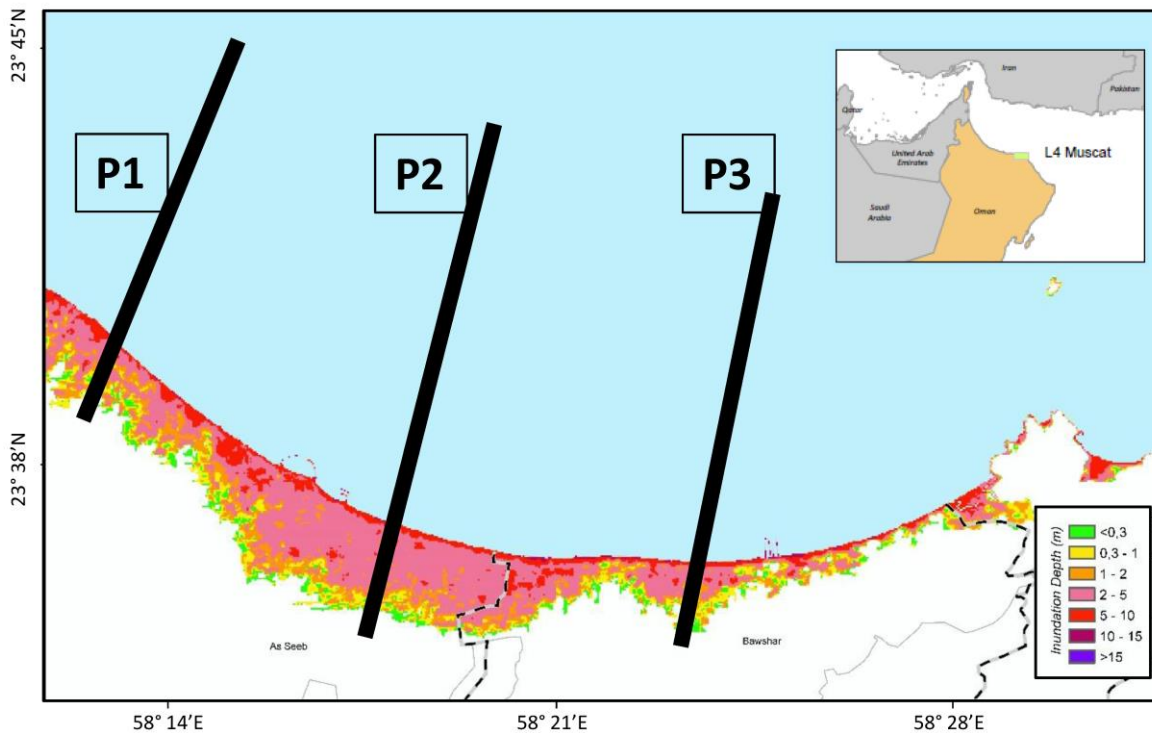


**Fig. 19. Flooded area in the municipality of La Libertad in El Salvador, due to an 8.1 magnitude event with epicenter along the coast of this Central American country. The exact locations where the run-up was estimated are provided in Table 4**

5 In Fig. 19, the flooding map that was part of this project is shown, and in the same figure, the selected profiles have been superimposed. In this simulation, the run-ups obtained at the three profiles in Fig. 19 were 5.2, 5.5, and 6.3 m. The corresponding run-ups obtained by interpolating the TRD with the IH-TRUST tool were 6.2, 6.1 and 7 m.

### 5.1.3 Tsunami scenario in Muscat, Oman

As part of the Multi Hazard Risk Assessment System of Oman (Aniel-Quiroga et al., 2015), more than 3000 potential tsunami events were numerically modeled. A selection of these events were selected to assess the tsunami hazard for some specific municipalities in Oman by means of high-resolution numerical simulations of the generation, propagation and inundation processes, with a 30 m grid. One of these cases was an extreme event of magnitude 9.0 with epicenter in the Makran Subduction Zone (MSZ). For the capital city area, Muscat, the resultant flooding map is shown in Fig. 20; the profiles that were selected for the validation are superimposed on this map. The tsunami wave height and period offshore were approximately 2 m and 2300 s. In these cases, the measured run-ups at each profile were 6.2, 8.7, and 7.7 m. The corresponding run-ups calculated with the new database were 6.3, 8.5, and 7.8 m.



**Fig. 20. Flooded area in the municipality of Muscat, capital city of Oman, due to a 9.0 magnitude event with epicenter in the Makran Subduction zone. The extracted locations where run-up was estimated and the run-up values, both modeled and estimated with the IH-TRUST tool, are provided in Table 4.**

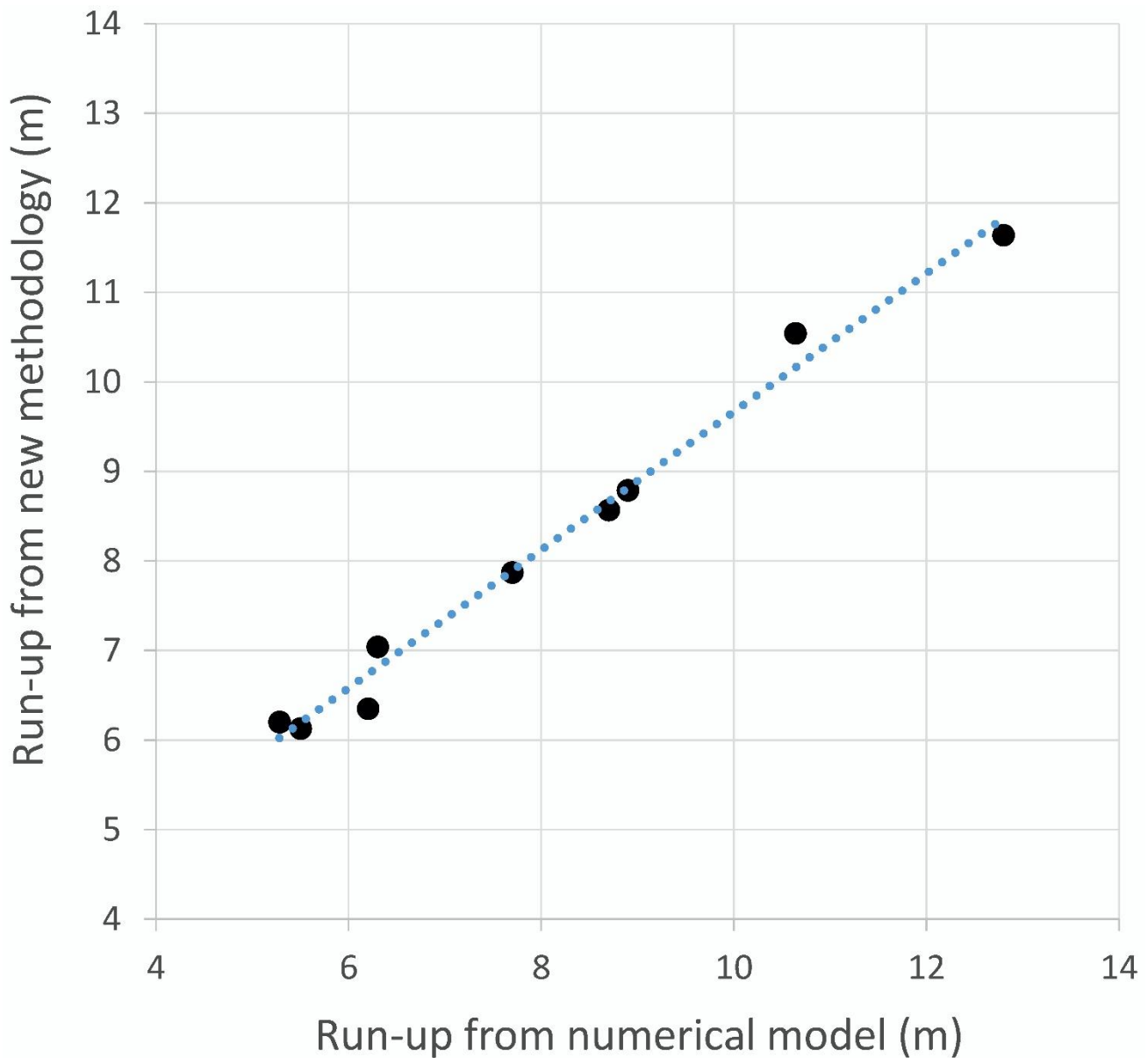
5 Table 4 summarizes the results obtained for the validation with the high-resolution simulations in the three scenarios. The run-up values, both those calculated with the numerical model and those estimated with the proposed database and detailed methodology described above, have a similar magnitude; in some cases, the result is accurate enough to rely on the results of the presented methodology. In addition, in Table 4, the estimated run-up is also compared to the result of applying 2 empirical run-up formulae. First, the Synolakis formula, that although it was created for the run-up of Solitary Waves it has been widely applied in the past for tsunamis, and second, Madsen and Schäffer expression for single waves (Madsen and Schäffer, 2010).

**Table 4. Tsunamis scenarios included in the validation process of the database and tool. The numerical model column includes the run-up obtained with the high-resolution numerical simulations and can be compared to the estimations from the application of the IH-TRUST and formulae of Synolakis (Synolakis, 1987) and Madsen and Schäffer (2010)**

Place	Coordinates of the run-up point			Run-up			
	lon	lat	z	From model	From Method	From Synolakis	From Madsen
Trujillo	-80.035	-9.07239	2000	10.64	10.54	8.2	18.0
(Perú)	-80.23769	-9.02625	2000	8.9	8.79	8.1	12.4

	-80.184535	-9.1122838	2000	12.8	11.64	9.9	18.5
Muscat (Oman)	58.252483	24.026722	1400	6.2	6.35	8.2	7.5
	58.3063	24.01996	1300	8.7	8.57	8.8	7.6
	58.446381	24.025935	1500	7.7	7.87	10.4	6.4
La Libertad (ElSalvador)	89.63894	12.78405	3000	5.28	6.2	6.9	10.2
	-89.595559	12.7768479	3000	5.5	6.13	7.0	10.3
	-89.575252	12.776232	3000	6.3	7.04	6.9	10.3

Fig. 21 shows this comparison in a plot, in which the fitting between the modeled and calculated run-up values is noted. In addition, the new methodology is better than the result of the Synolakis and Madsen and Schäffer formulae, which generally overestimates the run-up.



**Fig. 21.** Plot of the numerically modeled run-up against the calculated run-up values for the validation cases.

## 5.2 Validation with data recorded during field campaigns after real events

The low frequency of major tsunamis are invaluable to the field campaigns that are carried out immediately after a tsunami event occurs. These field campaigns allow the evaluation of the developed strategies of risk reduction and the creation of new, more accurate strategies. From a pragmatic point of view, the data collected during these campaigns allow scientists and engineers to validate or calibrate numerical models and methodologies. In this case, this type of validation has been addressed

5 using the available field data of the events in Japan (2011) and Chile (2010 and 1960). The bathymetric profiles used in the validation have been constructed using GEBCO. The tsunami wave time series have been obtained from the data available from DART buoys (Meinig et al., 2005) or numerical simulations of accurate sources; this process is explained in detail later in the paper. The results of the application of the methodology have been compared to observational data recordings and field survey papers.

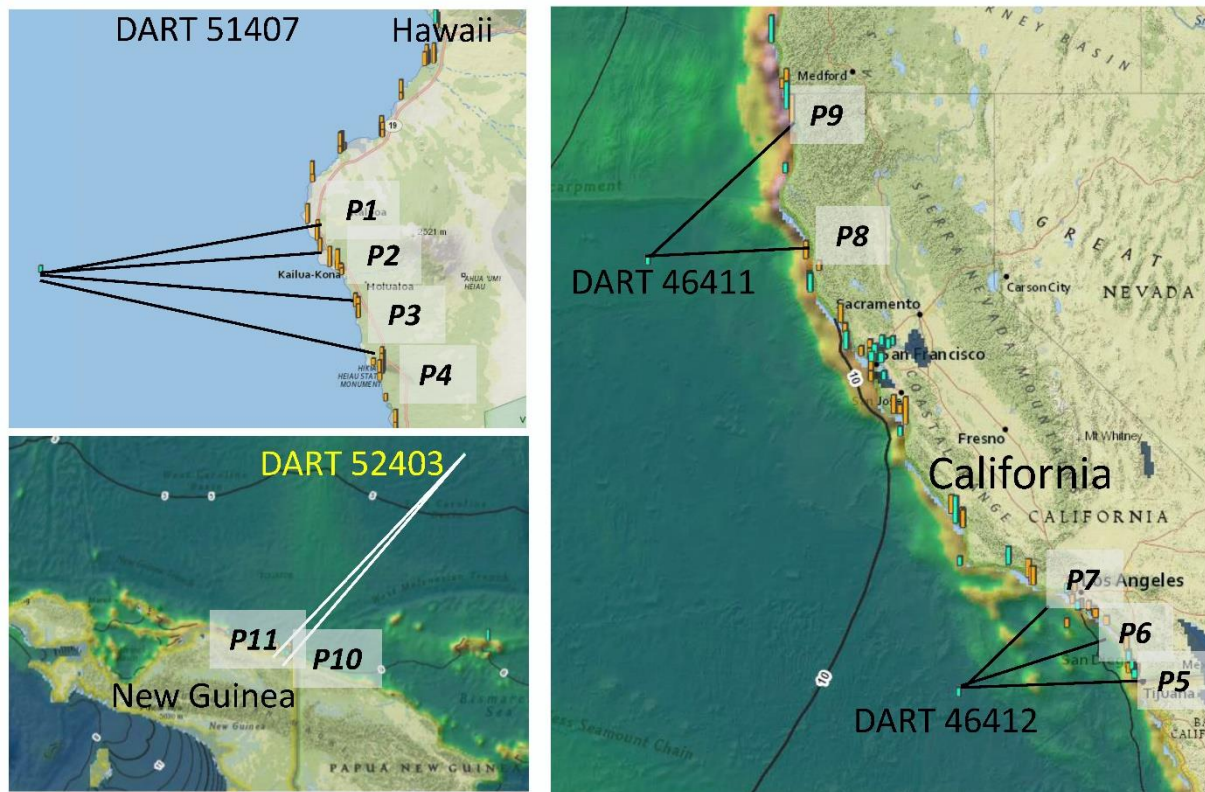
### 5.2.1 2011 tsunami on the coast of Japan affecting the Pacific basin

On the 11<sup>th</sup> of March, 2011, a 9.0 earthquake, which had an epicenter close to the coast of Japan, triggered a tsunami that reached the coast of Japan within one hour. This tsunami wave propagated across the Pacific Ocean, reaching the U.S. West Coast in 10 hours and the coast of Chile in 21 hours.

10 The tsunami wave time series used for this validation have been obtained from the data available from DART buoys (Meinig et al., 2005). The results were compared with the observed run-up (National Geophysical Data Center NOAA).

It is essential to highlight that the application of the new run-up estimation methodology is restricted to the profiles and wave shapes whose parameters fall inside the ranges covered by the database (see Table 1).

15 Therefore, the use of the methodology is limited to these cases. An example of non-applicability occurs when the tsunami height and period are obtained ( $d_2$ ) in a shallow area of the ocean or when the generation zone is too close to the study area and a complete time series of the tsunami wave cannot be properly recorded at an adequate depth.



**Fig. 22. Validation with DART buoy time series. 4 DART buoys were used, and their data were applied to several bathymetric profiles to validate the methodology. The locations of the points where the run-up was estimated are included in Table 5**

5 In the case of the Japan 2011 event, due to the proximity of the coast, it was not possible to obtain a complete time series between the epicenter and Japan, and the validation has been carried out in other areas of the Pacific Ocean, using four DART buoy records (near Hawaii, California, and Papua-New Guinea). The names and locations of the DART buoys used are given in Table 5. This table also includes the names and locations where the run-up was estimated with the data of each DART buoy, the run-up value recorded in the field surveys at those locations, and the estimated value of the run-up, both by using the new methodology and by applying the Synolakis formula (calculating the tsunami wave height at a depth of 10 m using Green's law) and Madsen and Schäffer formula. The buoy locations are also included in Fig. 22.

10

**Table 5. Validation with DART buoy time series of the Japan 2011 event. 4 DART buoy datasets were used on several bathymetric profiles to validate the methodology. Location names correspond to those given by the National Geophysical Data Center (NOAA). Synolakis run-up was estimated by applying the so-called Green's Law to the time series of the DART buoys to obtain the tsunami height near the coast.**

15



DART Buoy					Run-up (m)					
#	LON	LAT	DEPTH (m)	LOCATION	LON	LAT	SURVEY	Synolakis + Green	IH-TRUST	Madsen And Schäffer
51407	-156.5	19.620	4771	P1 Wawalolo	-156.05	19.71	2.4	3.8	2.0	1.67
				P2 OldAirport	-156.01	19.64	3.1	3.8	2.0	1.69
				P3 Kahaloo	-155.97	19.58	2.0	3.8	2.4	1.72
				P4 Keel	-155.93	19.46	3.0	3.8	2.8	1.74
46412	-120.7	32.250	3776	P6 Ocean Beach	-117.26	32.74	1.0	1.9	1.5	1.07
				P7 Marina del Rey	-118.45	33.98	1.0	1.9	1.9	1.13
				P5 Channel Islands	-119.22	34.15	1.2	1.9	1.4	1.21
46411	-127.0	39.340	4319	P9 Jenner River	-123.1	38.43	1.0	2.7	1.1	1.75
				P8 Dolphin isle	-123.8	39.43	1.0	2.7	1.3	1.98
52403	145.52	4.020	4474	P10 Holtekamp	140.779	-2.627	2.0	2.3	1.7	3.64
				P11 Pelabuhan	140.368	-2.461	1.3	2.3	1.4	3.67

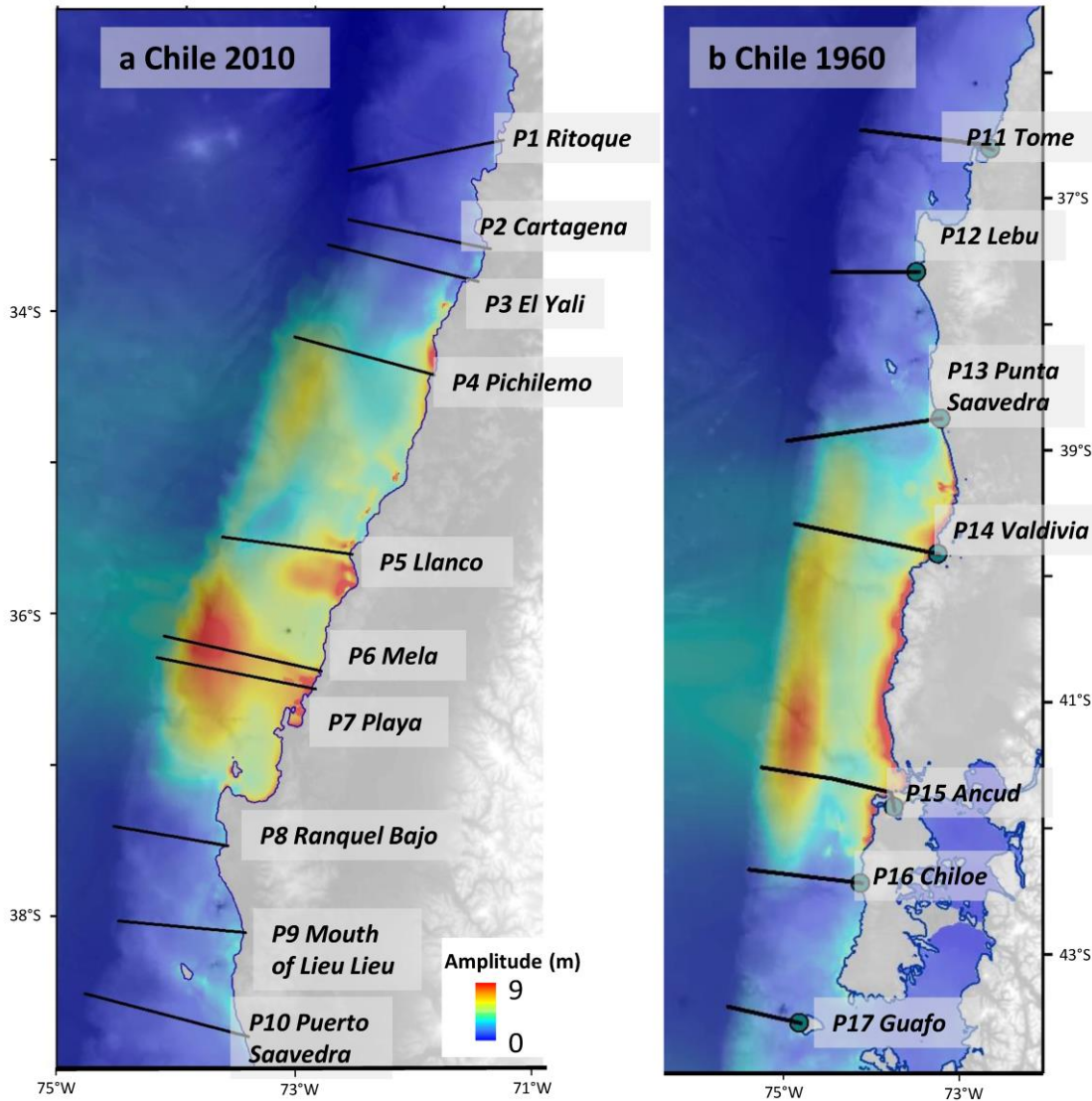
As it can be inferred from the application of the methodology, the run-up estimated values are on the same order of magnitude as the recorded inundation; generally, the results are accurate, and differences are normally lower than 20%. These results are

also closer to the observed run-ups than those obtained by applying Synolakis and Madsen and Schäffer formulae, which often overestimates the run-up.

### 5.2.2 Chilean coast tsunamis (2010 and 1960)

When no real record is available to determine the offshore wave shape (DART buoys), the main issue is the correct definition of the source to compute an accurate numerical simulation. Although there is no shortage of uncertainties in the determination of the source, the tsunami initial surface deformation models that have been developed are accurate (Barrientos and Ward, 1990). As an alternative validation approach, two of these models have been used to validate the new run-up estimation methodology with the events that occurred in Chile in 2010 and 1960.





**Fig. 23. Profiles and locations used in the validation of the new methodology by using the run-up recorded after the tsunami events of 2010 (a) and 1960 (b) in Chile. The locations of the points where the run-up was estimated are included in Table 6**

- 5 On the 27<sup>th</sup> of February, 2010, an 8.8 magnitude earthquake with epicenter on the coast of Chile triggered a tsunami that reached the Chilean coast in less than 30 minutes. In the Bio-Bio region, the run-up was recorded at several locations (Fritz et al., 2011). To apply this methodology, first, a rough numerical simulation of the generated tsunami was addressed. This simulation used the source definition by (Shao et al., 2010) and gridded the GEBCO data with 700 m cells (see Fig. 23). From this simulation, the profiles and wave amplitude time series in the generation area were obtained. The tsunami wave height

and period recorded at each location and the result of the interpolation from the further improved database for each corresponding profile are given in Table 6.

The optimal application of the run-up estimation methodology is achieved at the locations sufficiently far from the source, as explained in the previous subsection. The result at these points (1, 2, 3, 8, 9 and 10) have the same order of magnitude of the recorded run-up from Fritz’s survey. In the locations in front of the source, the initial deformation of the water surface did not allow a complete time series to be obtained to estimate the tsunami wave height and run-up.

Regarding the 1960 earthquake and tsunami in Chile (Lomnitz, 2004) (Fig. 23), this earthquake is considered the greatest earthquake ever recorded, and the numerical simulation computed for the validation used the source by Barrientos and Ward (1990). The run-up data for the validation was obtained from the NOAA global historic tsunami database. In this case, the data are mainly from eyewitness testimony.

In Table 6, the results of the application of the methodology at 7 locations and the recorded run-up are given. In this case, the tsunami wave height at 3 of the locations (P13, P14, and P15), was such that the profiles were not within the database application ranges. The other 4 locations provided results that are on the same order of magnitude as the observed run-up.

In the application of the new methodology to the Chile events, the tsunami wave height used for the interpolation came from a numerical simulation, and the results were compared to real run-up records. Although the validation inherits the uncertainties of the source, the results are sufficiently accurate, taking into account the limitations explained above.

**Table 6. Validation of the methodology with the results of numerical simulations of realistic sources of the 1960 and 2010 events on the coast of Chile. Fritz et al. (2011) survey results were used to validate the results from the new methodology for the Chile 2010 event. NOAA’s National Geophysical Data Center data were used to carry out the comparison with the 1960 event**

	PROFILE	LOCATION	LON	LAT	Run-up (m)	
					SURVEY	IH-TRUST
2010 event	P1	Ritoque	-71.528	-32.826	3.4	1.39
	P2	Cartagena	-71.602	-33.542	4	1.93
	P3	El Yali	-71.717	-33.751	2.1	3.47
	P4	Pichilemo	-72.005	-34.384	5	N/A
	P5	Llanco	-72.623	-35.584	11.4	N/A
	P6	Mela	-72.852	-36.36	3.1	N/A
	P7	Playa	-72.911	-36.478	6.6	N/A
	P8	Ranquil Bajo	-73.596	-37.526	5.7	2.2
	P9	Mouth of Lieu Lieu	-73.449	-38.097	2.3	2.2

	P10	<b>Puerto Saavedra</b>	-73.701	-38.783	2.5	2.3
	P11	<b>Tome</b>	-72.962	-36.619	2.5	3.4
	P12	<b>Lebu</b>	-73.674	-37.608	4	4.6
	P13	<b>Punta Saavedra</b>	-73.407	-37.608	11.5	N/A
<b>1960 event</b>	P14	<b>Valdivia</b>	-73.411	-39.844	10	N/A
	P15	<b>Ancud</b>	-73.828	-41.859	12	N/A
	P16	<b>Chiloe</b>	-74.176	-42.465	10	9
	P17	<b>Guafo</b>	-74.83	-43.578	10	10.2

## 6. Conclusions

The calculation of the flooding that a tsunami causes inland is addressed when a tsunami risk assessment is conducted. For a historical event, the assessment determines the limit of the affected area. In addition, the predictive evaluation of this flooded area, based on the potential tsunami scenarios that can affect it, allows prevention and mitigation measures to be established, helping to reduce the risk.

However, the calculation of this flooded area, particularly the assessment of the run-up, is not always direct. Occasionally, there are no high-resolution data that allow the application of numerical models. In addition, the accuracy of the existing empirical formulae can be improved, since they do not take into account natural topobathymetric profiles from the propagation to the inundation areas.

In this paper, an alternative methodology that complements the existing ones has been presented. This methodology consists of a numerical flume formed by the coupling of two numerical models (COMCOT and IH2VOF). The developed hybrid model is applied to each part of the generation-propagation-inundation process and this numerical model obtains a more accurate result; additionally, it is computationally affordable. The inputs for this hybrid model are the topobathymetric profile and the tsunami wave. The topobathymetric profiles were parameterized with 5 parameters (3 slopes and 2 depths), using a real profile sample to define the parameterization. In addition, the tsunami waves were parameterized with 2 parameters (tsunami height and period) using tsunami amplitude time series obtained by using numerical simulations of realistic tsunami events.

This methodology allows the accurate calculation of the run-up on along topobathymetric profile. Therefore, this methodology has been used to construct a tsunami run-up database. This database aimed to create an interpolation domain in which new run-up calculations could be carried out. The events of the database are combinations of a selection of bathymetric profiles

and tsunami waves that were simulated with the hybrid model to create the database of simulations from which an interpolation can be executed to calculate the run-up of new tsunami scenarios.

To easily address the interpolation, a tool called IH-TRUST was scripted. This tool uses real profile and wave data, parameterizes them to find their most similar parameters in the database, and interpolates the results to provide a run-up value.

Once the input parameters are given, the application of this interpolation provides results in just a few seconds, what shortens typical simulations of several nested grids, which commonly take several hours to provide results in all the computational domain.

To validate this new methodology and tool, the results of its application have been compared with both high-resolution numerical simulations and field survey data. The run-ups obtained with IH-TRUST are consistent and suggest that the tool can accurately calculate the run-up.

The assessment of the tsunami hazard begins by calculating the area affected by the tsunami. In those coastal areas where no other data are available, the detailed methodology and tool allow the run-up value of tsunami events to be determined without using high-resolution numerical simulations.

Therefore, to assess the hazard in a tsunami-prone area, this methodology can be applied to several profiles along the coastal area study. As a result, the methodology provides the run-up at each of the profiles, allowing an estimation of the flooded area from an area within the envelope of the run-up limits.

The application of the tool has some limitations; for example, the tool will indicate if the bathymetric profile or the tsunami wave parameters are not included within the range of values in the database, as explained for the case of Chilean Trench in 5.2.1.

New work in this field should take into account these difficulties to further develop the database with new parameter values that include these singularities.

The generation of the database and the values of run-up obtained from a combination of the bathymetric profiles and tsunami waves have provided a rich and populated space where the influence of each parameter on the final value of the run-up can be addressed. In this sense, which profiles are more prone to suffer higher run-ups in the case of a tsunami can be defined. For instance, profiles with high land slopes ( $\tan\beta_0$ ) are associated with higher run-up values than those with low land slopes. In addition, some combinations of offshore slopes and continental shelf slopes ( $\tan\beta_1$  and  $\tan\beta_2$ ) minimize the run-up value for

the same tsunami wave. In addition, the influence of  $\tan\beta_2$  is considerable and justifies inclusion of the deep-water area ( $d_2$ ) in the parameterized profile. On the other hand, when the profile is for a large continental shelf, the run-up increases; however, the run-up value decreases for gentle continental shelf slopes.

5    Traditionally, empirical methods, like the application of Synolakis’s formula, simplify the profile using one or two slopes (Park et al., 2015). However, this assumption is not accurate; in this study, the importance of using a complete profile, including the tsunami generation area, has been noted, as well as the influence of the profile parameters on the final run-up value.

**Acknowledgements**

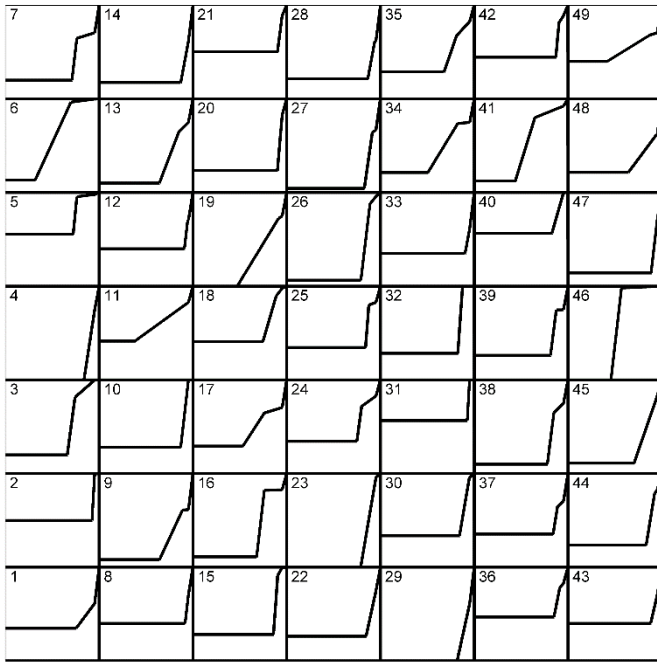
10   The research leading to these results has received funding from the European Union's Seventh Framework Programme (FP7/2007-2013) under grant agreement n° 603839 (Project ASTARTE - Assessment, Strategy and Risk Reduction for Tsunamis in Europe).

15

**Appendix A**

**Database profiles**

In this section, the 49 artificially generated profiles are shown in Figure A1. The five corresponding parameters are listed on  
20   Table A1.



**Fig. A1.** The 49 profiles used to generate the IH-TRUST database

**Table A1.** Synthetic profiles

#	$\tan\beta_0$	$\tan\beta_1$	$\tan\beta_2$	$d_1$	$d_2$
1	15.00%	2.50%	2.50%	0	2600
2	0.05%	20.00%	20.00%	0	3000
3	0.05%	1.00%	8.50%	1100	4800
4	8.00%	10.00%	10.00%	0	6000
5	0.50%	0.15%	12.00%	150	2600
6	0.05%	0.15%	2.50%	200	5200
7	10.00%	0.50%	14.00%	500	4300
8	10.00%	14.00%	14.00%	0	2600
9	15.00%	0.15%	4.00%	50	5200
10	0.05%	10.00%	10.00%	0	4300
11	5.00%	1.00%	1.00%	0	3000
12	10.00%	0.50%	14.00%	20	2600
13	8.00%	1.50%	4.00%	800	5200
14	15.00%	10.00%	10.00%	0	4300
15	0.05%	2.00%	16.00%	500	4300
16	5.00%	0.05%	12.00%	50	5200
17	10.00%	0.50%	2.50%	500	3500
18	0.50%	1.50%	4.00%	500	3500
19	8.00%	1.00%	2.50%	200	6000
20	5.00%	16.00%	16.00%	0	4300

21	3.00%	10.00%	10.00%	0	2600
22	8.00%	7.00%	7.00%	0	3900
23	1.00%	7.00%	7.00%	0	6000
24	5.00%	1.00%	10.00%	800	3500
25	5.00%	0.50%	18.00%	200	3500
26	1.00%	1.50%	10.00%	500	5600
27	12.00%	1.00%	12.00%	200	5600
28	15.00%	0.30%	10.00%	20	3900
29	15.00%	8.50%	8.50%	0	6000
30	1.00%	1.00%	7.00%	20	3900
31	0.05%	2.50%	20.00%	20	2600
32	0.05%	0.05%	16.00%	20	4300
33	15.00%	10.00%	10.00%	0	2600
34	8.00%	0.15%	2.50%	100	4300
35	5.00%	1.50%	4.00%	1100	3900
36	5.00%	1.50%	7.00%	300	2600
37	10.00%	1.50%	10.00%	500	3000
38	8.00%	1.50%	12.00%	800	5200
39	8.00%	0.05%	12.00%	20	3900
40	0.05%	4.00%	4.00%	0	2600
41	2.00%	0.50%	4.00%	800	5200
42	3.00%	2.00%	14.00%	500	3000
43	10.00%	7.00%	7.00%	0	2600
44	5.00%	2.00%	8.50%	300	4300
45	5.00%	4.00%	4.00%	0	5200
46	0.05%	0.05%	10.00%	100	6000
47	10.00%	14.00%	14.00%	0	4800
48	15.00%	0.30%	2.50%	20	3900
49	10.00%	0.50%	1.00%	200	2600

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