Dear Editor and Referees,

Thank you for your kindly providing all these helpful comments. Our manuscript was revised and asked a native speaker for editing. Besides, we added some paragraphs to explain our methods including how we retrieve data, analyze them and also how we define the thresholds. Our replies and the corresponding revisions were all listed below.

Sincerely yours, Lun-Wei Wei

Referee #1

Comments to author:

No.	Comment	Reply
	The main issue of the work is the lack of a	Thanks for the comment. We agree the real
	real validation, since authors consider only	validation is needed to evaluate whether this
	rainfall events that triggered landslides, but	EWS is effective or not. We defined red
	they should consider, if possible, even	(extreme danger level) and orange (high
	events that not triggered landslide, to	danger level) alerts as alarm zone, while
	validate the early warning system in terms	yellow (medium danger level) and green
	of false alarms, missed alarms and correct	(low danger level) alerts as no alarm zone in
	alarms. To identify these categories, they	page 7, line 27–30. After that, the numbers
1	should define a threshold to identify a "no	of True Positive (TP), True Negative (TN),
	alarm zone" and an "alarm zone" (e.g. green	False Positive (FP) and False Negative (FN)
	area of fig. 6, 8, 9 could be considered as no	were counted and the skill scores including
	alarm zone, while yellow to red areas as	the probability of detection (POD), the
	alarm zone). Without such a validation a	probability of false detection (POFD) and
	functional EWS cannot be considered as	the probability of false alarm (POFA) were
	effective or ineffective.	used to evaluate the effectiveness of this
		EWS in the last paragraph of section 5.3 in
		page 10.
	Another important point author should	Thanks for the comment. During field
	clarify is how they identified the exact time	investigations, we not only verified the
	of landslide, since it is necessary to	correctness of landslide inventories but also
2	calculate the 3-hours rainfall intensity. They	tried to acquire the exact time of landslides
Z	located landslide with several approaches as	from residents lived around. The accuracy of
	the use of SPOT5 satellite imagery, but in	exact time of landslide was hard to evaluate,
	this case is not possible to identify the exact	however, we focused on interviewing as
	occurrence time of the landslides.	many residents whose relatives were injured

	or houses	were damage	ed/destr	oyed by	the
	landslides	as possible, s	o that t	he qualit	y of
	landslide	occurrence	time	might	be
	improved.				

Referee #1

Comments in PDF file:

No.	Comment	Reply
1	[Page 2, line 8] Please add Rosi et al. 2012;	Thanks for the comment. We added this
1		reference in page 2, line 10.
	[Page 2, line 14] (1) Modified as "Segoni et	Thanks for the comment. We modified and
	al, 2014, 2015" (2) Add also "Rosi et al.	added these important references in page 2,
2	2016." Rainfall thresholds for	lines 15–16.
	rainfall-induced landslides in Slovenia.	
	https://doi.org/10.1007/s10346-016-0733-3	
3	[Page 2, line 27] "region" replace with	Thanks for the comment. We used "mosaic"
5	"mosaic."	instead of "region" in page 2, line 33.
	[Page 2, line 29] "Geological settings"	Thanks for the comment. We used
4	replace with "Lithological units."	"Lithological units" instead of "Geological
		settings" in page 2, line 35.
	[Page 3, line 1] I suggest splitting this	Thanks for the comment. We split "Data and
5	chapter into two chapters. 3: Available data.	methodology" into "Available data" and
5	4: Methodology. This will increase the	"Methodology" to increase the readability.
	readability of the document	
	[Page 3, line 3] Please change the number of	Thanks for the comment. We renewed the
6	the paragraphs according to the new chapter	numbers of each paragraphs.
	division	
	[Page 3, line 5] all the approaches you used	Thanks for the comment. We agree that it is
	to create a landslide DB are right, but they	impossible to get the exact time of landslide
	have a major issue: the date of the landslides	from landslide DB. Therefore, we tried to
	are approximated and this is will affect the	acquire the exact time of landslide from
	identification of the real rainfalls	residents lived around, especially whose
7	responsible of the initiation of the	relatives were injured or houses were
	landslides. If you use 3 hours rainfall you	damaged/destroyed by the landslides during
	need the exact time of landslide triggering.	our field investigations. We emphasized this
	Please clarify these points.	in section "3.2 Landslide occurrence time
		and field investigation", page 3, lines 19-
		23.

	[Page 3, line 21] Exact date is usually hard	Thanks for the comment. We believe that
	to identify and the exact hour is even more	the uncertainty of triggering time is hard to
	difficult. Do you consider the uncertainty of	evaluate due to the lack of video records.
	triggering time? How do you manage it?	However, we tried to interview residents,
		especially whose relatives were injured or
		houses were damaged/destroyed by the
0		landslides, to get the occurrence time of
8		landslides during field investigation. Given
		the deep impressions left by such memories,
		we believe that the quality of landslide
		occurrence times might be improved. We
		added these descriptions in section "3.2
		Landslide occurrence time and field
		investigation", page 3, lines 19-23.
	[Page 3, line 25] Please describe how you	Thanks for the comment. We developed a
	performed the reduction to 10 m resolution	Fortran program to obtain smoothed and
	and the smoothing. Did you use a simple	resolution-reduced 10×10m DEMs (10m
	GIS resample technique? Have you	DEMS) by calculating the average value of
	considered the effects of smoothing the	each 2 by 2 grid in the 5m DEMs. The
	DEM on the morphological analyses?	resolution-reduced 10m DEMs could
	Please clarify.	generate some differences in the
9		morphological analysis, but the expected
		scale of the landslide susceptibility in this
		study was set to 1:25,000, so differences
		smaller than 12.5m could be ignored
		according to the relationship between
		mapping scale and 5% acceptable error. We
		added these descriptions in section "3.3
		Slope units" (page 3, lines 30–34).
	[Page 3, line 26] the procedure you cited	Thanks for the comment. Slope units were
	(Xie et al, 2004) identify slope units from	delineated according to gullies and ridges.
	DEM, by the use of Arc Hydro tool. Each	First, gullies and watersheds were analyzed
	slope unit is characterized by several	by successively using spatial analysis tools
10	homogeneous parameters. I believe that a	in ArcGIS: fill, flow direction, flow
	more accurate description of the whole	accumulation, stream link (with 2,000 used
	procedure you used to identify slope units is	as the threshold) and watershed. Second,
	required, to better understand the paper.	reverse DEMs were generated by
		multiplying DEMs by -1. In the reverse

		DEMs, ridges became gullies and could be
		analyzed by the same methods used in the
		first step. Third, the watersheds of the
		DEMs and reverse DEMs were transformed
		from rasters to polygons for further editing
		by using the "Raster to Polygon" tool in
		ArcGIS and then cut by each other 5 to
		delineate the slope units. Finally, slope units
		were modified manually according to aspect
		and gradient. It is suggested that the aspect
		in a slope unit should be within three
		adjacent directions: e.g. northwest north
		and northeast. On the other hand, the
		difference in gradient should not be over 30
		degrees in a slope unit, and slope units
		situated on flat areas, including alluvial
		deposits and terraces, were deleted. In
		addition, the area of each slope unit was set
		to around 5 ha; therefore, slope units smaller
		than 5 ha were combined with adjacent
		slope 10 units and those larger than 5 ha
		were split into several smaller ones.
		Moreover, slope units delineated by parallel
		drainage on a dip slope were combined into
		one slope unit. We added these detailed
		procedure in section "3.3 Slope units" from
		page 3, line 35 to page 4, line 11.
	[Page 4, line 2] What do you mean with	Thanks for the comment. Whenever Taiwan
	total rainfall? How long is the period you	has a typhoon event, the Central Weather
	considered to calculate it? How did you	Bureau issues disaster prevention alerts. We
	decide to use 3 and 24 hours rainfall? Please	therefore counted the time that the first alert
	clarify.	was issued as the beginning of the rainfall
11		event and the time that the alert was
		cancelled as the end of the rainfall event to
		calculate rainfall amounts. We added this in
		page 5, lines 6–8.
		For the decision of 3-hour mean rainfall
		intensity and 24-hour accumulated rainfall,

		we calculated the triggering rainfall, including the rainfall intensity (I1, I2, I3, I4,
		I5, I6) and accumulated rainfall (R6, R12, R24,
		R48, R72) of different time windows of each
		landslide case according to the landslide
		occurrence time. The results revealed that
		218 landslides occurred within the 3 hours
		following the highest rainfall intensity, and
		242 occurred within the 3 hours following
		the 2 nd or 3 rd highest rainfall intensity (i.e.,
		induced by high rainfall intensity),
		accounting for nearly 49% of the landslide
		cases gathered in this study (shown in Table
		3). From these results, it became clear that
		in Taiwan, I3 is the most important index for
		landslides induced by rainfall of short
		duration but high intensity. On the other
		hand, 481 landslides occurred close to the
		end of the rainfall events (i.e., induced by
		high accumulated rainfall), accounting for
		about 51% of the total cases. Furthermore,
		analysis of the different accumulated rainfall
		indexes showed that 24-hour accumulated
		rainfall had the lowest coefficient of
		variation (shown in Table 4), indicating that
		this index was less dispersive than others
		and might be more suitable for serving as an
		accumulated rainfall index for establishing
		rainfall thresholds. We added these details in
		section 4.2 (page 6, lines 9–19).
	[Page 4, line 12] what do you mean "The	Thanks for the comment. As we know,
	ratio of steep slope was calculated by	shallow landslides are prone to occur on
12	dividing the area that greater than 30	steep slopes; therefore, we used the "ratio of
	degrees by total area of slope unit."?	steep slopes" to present how many steep
		slopes existed in a slope unit. It was found
		after trial and error that a threshold of
		gradient higher than 30 degrees had a higher
		relationship with landslide susceptibility.

		Thus, we calculated the area where the
		gradient was greater than 30 degrees ($A_{>30}$)
		as well as the total area (Atotal) of each slope
		unit Therefore the ratio of steen slope
		could be calculated by dividing $A_{>20}$ by
		A set We added these detailed descriptions
		Atotal. We added these detailed descriptions in section 2.4 (mass 4 lines $28, 22$)
		In section 3.4 (page 4, lines $26-32$).
	[Page 4, line 1/] Kriging interpolation	Thanks for the comment. We collected
	method is very effective, but it has to be	hourly rainfall data of 423 rain stations
	properly performed. You should describe	provided by Central Weather Bureau,
	how you applied it.	Taiwan and analyzed both the 3-hour mean
		rainfall intensity (I_3) and the 24-hour
13		accumulated rainfall (R ₂₄) of each station.
10		After that, we used the linear mode of
		ordinary kriging and applied the default
		setting in Surfer software to obtain the
		rainfall distribution of the whole study area.
		We added these descriptions in page 5, lines
		2–6).
14	[Page 4, line 23] & [Page 4, line 25]	Thanks for the comment. We corrected these
14	"required" \rightarrow "require"	sentences in page 5, lines 14–21.
	[Page 4, line 35] please clarify how you	Thanks for the comment. The index
	defined the coefficient w in LR function.	indicating landslide/non-landslide was set as
		the dependent variable, and all the landslide
		susceptibility factors were set as covariates
		in SPSS for training of the model. After
		iterative training, the regression coefficients
15		of each landslide susceptibility factor, as
		well as the success rate curve (SRC), the
		prediction rate curve (PRC), and the area
		under the curve (AUC), were reported in
		SPSS. We added a more detailed
		descriptions in page 5, lines 31–34
16	[Page 5, line 18] Why did you not use the	Thanks for the comment We agree that
	cumulative rainfall of 3 hours? It is the	using cumulative rainfall of 3 hours is
	same	similar to 3-hour mean rainfall intensity (Ia)
	Sumo.	We chose 3-hour mean rainfall intensity
		we chose 5-nour mean rannan mensity

		to focus on rainfall of short duration but
		high intensity. Similarly, we chose 24-hour
		accumulated rainfall to focus on rainfall of
		long duration but low intensity. We added
		these descriptions in page 6, lines 25–27.
	[Page 7, line 16] for a complete validation	Thanks for the comment and kindly
	you should use also rainfall events that not	providing relevant references. We defined
	triggered landslides, to calculate False	the no alarm zone from alarm zone and
	alarms, correct alarm and missed alarm.	calculate the numbers of false alarms,
17	See Segoni et al. 2014, Rosi et al, 2015, etc.	correct alarms and missed alarms to make a
		complete validation of our EWS. It was
		revised in section 4.3 (page 7, lines 16–34).
		On the other hand, the results were shown in
		page 10, lines 5–15.
	[Page 7, line 25] I believe this happened	Thanks for the comment. If rainfall stops,
	because you used rainfall intensity. If rain	not only 3-hour mean rainfall intensity (I ₃)
	stops, intensity decreases, but if you try to	but also 3-hours cumulative rainfall (R ₃)
	use 3-hours cumulative rainfall you should	decrease because only the rainfall in the
	avoid this problems.	nearest 3 hours (h, h-1, h-2) are taking into
		consideration. In this study, rainfall
		thresholds were set according to the I ₃ -R ₂₄
		diagram shown as Figure 4. If 3-hours
18		cumulative rainfall (R3) were used to
		replace 3-hour mean rainfall intensity (I ₃),
		the scale of y-axis and the value of new
		threshold will also be 3 times larger in the
		R ₃ -R ₂₄ diagram. It means that no matter in
		the I ₃ -R ₂₄ diagram or R ₃ -R ₂₄ diagram, for
		the same rainfall events, the snake line will
		all turned back to yellow when the rainfall
		fell.
	[Page 14, Figure 3] this Figure is missing of	Thanks for the comment. The other reviewer
10	some elements: scale bar, legend,	suggested deleting this figure because it was
19	orientation (North direction).	not useful for the discussion. We deleted it
		in this revised manuscript.

Referee #2

Comments to author:

No.	Comment	Reply
	The paper is very poorly written, with a bad	Thanks for the comment. We carefully
1	English. Several typos are present	checked again and asked for the editing by
	everywhere in the text. Moreover, the use of	an English native speaker throughout the
	past and present tenses is hardly	manuscript.
1	understandable. Several sentences are not	
	clear at all. I suggest a strong revision of the	
	paper in this view, possibly with an editing	
	by an English native speaker.	
	The introduction could be improved by	Thanks for the comment. We analyzed and
	reporting and analyzing some works that	added these important references in the
	have dealt with regional early warning	revised introduction.
2	models and early warning systems for	
	landslide occurrence, e.g. Segoni et al.	
	2014; Calvello et al. 2015, Devoli et al.	
	2015, Ficturio et al. 2017, Fumo et al. 2017.	Thanks for the commont. We split "Date and
	improved by adding more details on data	methodology" into "Available data" and
	gathering As an example, it is not clear how	"Methodology" so that more details can be
	Authors identified rock falls from the	described in each section
	landslide inventory. Moreover, Authors state	For the identification of rock fall from the
	that they gathered landslide occurrence time	landslide inventories, we deleted the
	by inquiry residents during field	polygons situated on slopes having
	investigations. This should be clarified, in	gradients greater than 55 degrees according
	particular because the occurrence time of	to the classification rules proposed by the
2	the landslides is very important for the	Central Geological Survey, Taiwan (Central
3	reconstruction of the 3-hour mean rainfall	Geological Survey, 2008). These
	intensity. In addition, more details on the	descriptions were added in page 3, lines 7-
	definition of landslide inventory would be	9.
	useful. Furthermore, it is not clear why the	For gathering landslide occurrence time by
	Authors calculated a precipitation map for	acquiring residents during field
	the whole study area. What is it for?	investigation, we focused on interviewing as
		many residents whose relatives were injured
		or houses were damaged/destroyed by the
		landslides as possible, so that the quality
		might be improved. These descriptions were

		added in page 3, lines 19–23.
		Detailed definitions including the
		classification and procedure for the
		generation of landslide inventory was added
		in section 3.1 (nage 3. lines $4-13$)
		The precipitation maps were produced and
		the triggering rainfalls of landslides were
		avtracted for the purpose of analyzing
		landalida augaantibility. This was revised in
		randshide susceptionity. This was revised in
		section 3.4 (page 3, lines $2-8$).
	Nothing is said about rainfall data. Did	Thanks for the comment. Yes, we used
	authors use rain gauge series? If yes, please	rainfall data from 423 rain gauges provided
4	explain how many rain gauges.	by Central Weather Bureau, Taiwan. Their
		distributions are shown in Figure 1. We
		added detailed descriptions in section 3.4
		(page 5, lines 2–8)
	The whole section regarding the landslide	Thanks for the comment, we revised this
	susceptibility analysis (section 3.2.1) should	section and asked for the editing by a native
5	be rewritten and increased by adding more	speaker again. We also added more detailed
	information. Several details on the adopted	procedures in this revised version in section
	method are missing.	3.4.
	In the section on rainfall thresholds, Authors	Thanks for the comment. We added the
6	refer to a coefficient of variation (also	equation of coefficient of variation in order
0	reported in Table 4); please explain how it	to explain how the calculation was made in
	was calculated.	page 6, lines 16–22.
	In the "3.2.3 landslide early warning model"	Thanks for the comment. The original
	section, it is very strange that 30%, 60% and	warning values of I3 and R24 of the 90%,
	90% thresholds correspond exactly to	60%, 30%, 15% thresholds were equal to
	integer values of I ₃ (30, 40, 60) and R ₂₄	the semi-minor axis and semi-major axis of
	(300, 400, 600). Is it just an example?	each threshold respectively. After that, I ₃
7	Please explain.	was rounded to the nearest 5 mm/h and R_{24}
	1	was rounded to the nearest 50 mm for
		operational purposes, such as the evacuation
		of residents. We added these explains in
		page 6, lines $32-35$ as well as page 8, lines
		26-27 and the caption of Figure 8.
	In the section related to the results of	Thanks for the comment. For a statistical
8	landslide susceptibility analysis the values	landslide suscentibility analysis it is
	initiasities subsceptionity unarysis, the values	initiasitice susceptionity analysis, it is

	of AUC are not so high to justify that "the	essential to use as many samples as
	results showed that LR model was stable	possible. However, we used slope units
	and nice in training as well as validation"	instead of grid units in this study for
	(Page 6, line 20). I suggest rephrasing this	application to disaster prevention. This led
	sentence, acknowledging that results could	to the reduction of samples, since one slope
	be better. Moreover, I suggest avoiding the	unit might equal hundreds of grids.
	word "nice", here and elsewhere in the text.	Therefore, our AUC might not be
		considered high in comparison to a
		grid-based landslide susceptibility model.
		We added these descriptions in page 8, lines
		12–15 and replaced the word "nice" with
		"acceptable" in this revised manuscript.
	At the end of section 4.2 (page 7, lines	Thanks for the comment. We agree these
	8-13), several actions to be performed in	suggested actions lead from a model to a
	case of different warning levels are reported.	EWS. Now we also develop a system
9	This step leads from an early warning model	connecting to the near real-time radar
	to an early warning system; therefore, it	rainfall data for disaster prevention. We
	should be remarked.	remarked these in section 5.2 (from page 8,
		line 32 to page 9, line 1).
	Regarding validation of the model (Section	Thanks for the comment. We agree that
	4.3), I would suggest using some indices or	quantitative evaluation of the performance
	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct	quantitative evaluation of the performance of early warning model is necessary. The
	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True
	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and
	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section
10	4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure.	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10.
10	 4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure. Conclusions section is very short! Authors 	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10. Thanks for the comment. We increased the
10	 4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure. Conclusions section is very short! Authors should add the main findings and the lesson 	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10. Thanks for the comment. We increased the contents of conclusion and all major
10	 4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure. Conclusions section is very short! Authors should add the main findings and the lesson learnt from their work. I suggest increasing 	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10. Thanks for the comment. We increased the contents of conclusion and all major findings were also included in this section.
10	 4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure. Conclusions section is very short! Authors should add the main findings and the lesson learnt from their work. I suggest increasing a lot this last section. 	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10. Thanks for the comment. We increased the contents of conclusion and all major findings were also included in this section.
10	 4.3), I would suggest using some indices or scores (e.g., count – and ratio – of correct and incorrect predictions, True Positive Rate, ROC analysis, etc.) to quantitatively evaluate the performance of the validation procedure. Conclusions section is very short! Authors should add the main findings and the lesson learnt from their work. I suggest increasing a lot this last section. Figure 1: add more descriptions in the 	quantitative evaluation of the performance of early warning model is necessary. The numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) were counted and the skill scores including the probability of detection (POD), the probability of false detection (POFD) and the probability of false alarm (POFA) were used to evaluate the effectiveness of this EWS. These results were shown in the last paragraph of section 5.3 in page 10. Thanks for the comment. We increased the contents of conclusion and all major findings were also included in this section.

12	Figure 3: not useful for the discussion. I	Thanks for the comment. We deleted this
15	suggest deleting it.	figure.
14	Figure 5: in the label of y-axis, pleas change	Thanks for the comment. We changed "hr"
14	"hr" into "h".	into "h" in texts, figures, and tables.
	Figure 6: it's a repetition of Figure 8b (for	Thanks for the comment. We deleted this
15	moderate susceptibility areas); I suggest	figure.
	deleting it.	
	Figure 7: I would suggest the following	Thanks for the comment. We changed the
	labels for x- and y-axes, respectively:	label in Figure 6 (Figure 7 in the original
16	"Portion of areas predicted as hazardous"	manuscript) according to the suggestion.
	for x-axis, and "portion of landslide	
	occurred" for y-axis.	
	Figure 8: I suggest enlarging it, and	Thanks for the comment. We enlarged
17	distribute the three panels vertically.	Figure 7 (Figure 8 in the original
1/	Moreover, please add a), b) an c) to the	manuscript) and distributed them vertically.
	three panels.	Besides, we added (a), (b), (c) in each panel.
	Tables 5 and 6: I'm not sure that colours can	Thanks for the comment. The colours are
10	be used in tables in NHESS journal. I	essential for understanding the alert.
18	suggest converting them into two figures, if	Therefore, we converted these tables into
	Authors want to maintain colours.	figures.
10	References: Please add DOI to each	Thanks for the comment. We added DOI for
19	reference in the list.	the references.
	As I already stated, the manuscript is full of	Thanks for the comment, we carefully
	technical and grammatical errors, typos, and	checked the manuscript again and asked for
20	incorrect use of words. Here I list just some	the editing by a native speaker.
20	suggestions of technical corrections, but	
	again I suggest a check and a language	
	revision of the whole text.	
	• Page 1, lines 29-31: please check this	Thanks for pointing out these unclear
	sentence and rewrite.	sentences and typos. We corrected all of
	• Page 3, line 9: correct "form".	them with caution in this revised
	• Page 3, lines 15, 22, 23, 30: please check	manuscript.
21	plurals (e.g., slope units, landslides,: : :).	Besides, we replaced the "rounded to" with
21	• Page 3, line 23: please check and correct	"rounded to the nearest 5 mm/h" and
	the sentence "This study used slope unit	"rounded to the nearest 50 mm" in page 6,
	that based on the features of: : :".	line 34; page 8, lines 26-27; the caption of
	• Page 4, lines 11-12: please reword.	figure 8.
	• Page 6, line 5: unclear, please rewrite.	On the other hand, we revised the words

• Page 6, lines 22-26: this sentence is	that expressing days in section 5.3, table 7
unclear, please reword.	and table 8.
• Page 7, line 3: replace "rounded to" with	
"rounded by".	
• Page 7, line 22: correct "form".	
• Page 7, line 25 and following: authors	
mention "14th", "15th", and others; if	
they are days, I suggest using the format	
dd-mm, which results more clear.	
• Page 8, line 4: "once landslide", what	
does it mean? Please correct.	

Adopting <u>the I</u>₃–R₂₄ rainfall index and landslide susceptibility <u>onfor</u> the establishment of <u>an</u> early warning model for rainfall-induced shallow landslides

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Abstract. Rainfall-induced landslide is one of landslides number among the most devastating natural hazards in the world, and the setup of early warning models is a pressing need for reducingare urgently needed to reduce losses and fatalities. Most part of landslide early warningswarning systems are based on rainfall thresholds defined aton the regional scale, regardless of the different landslide susceptibilitysusceptibilities of each slope various slopes.² Here we tried to divided value along units in

- 20 southern Taiwan into three categories (high, moderate_a, and low) according to their susceptibility. For each category, we established theirseparate rainfall thresholds separately so as to provide differentiated thresholds for different degrees of susceptibility. Logistic regression (LR) analysis was performed to evaluate the landslide susceptibility by using event-based landslide inventories and predisposing factors. Through the analysis Analysis of rainfall patterns of 941 more than 900 landslide cases gathered from field investigation, led to the recognition that 3-hour mean rainfall intensity (I₃) was recognized as a key
- 25 rainfall index for <u>rainfall of short duration but high intensity rainfall</u>; on the other hand, 24-hour accumulated rainfall (R₂₄) was recognized as a key rainfall index for <u>rainfall of long duration but low intensity rainfall</u>. Thus, the I₃-R₂₄ rainfall index was used for the establishment ofto establish rainfall thresholds in this study. Finally, an early warning model wasis proposed by setting warning signsalert levels including yellow (advisory), orange (watch) and red (warning) according to the concept of a hazard matrix. These differentiated thresholds and warning signsalert levels can provide essential information for local
- 30 government on evacuating decision of governments to use in deciding whether to evacuate residents.

Keywords: rainfall-induced landslide, landslide susceptibility analysis, rainfall threshold, early warning

1 Introduction

5

Rainfall-induced landslide is one of landslides number among the most perilous natural hazards, causing severe casualties and economic losses all over the worldworldwideworld (Ayalew, 1999; Evans et al., 2007; Tsou et al., 2011—; Petley, 2012; Wang

35 et al., 2015; Iverson et al., 2015; Sassa et al., 2015; Fan et al., 2017). Therefore, many efforts have been made to evaluate

the<u>landslide</u> susceptibility and thereby set criteria offor issuing early warning for the sake of savingalerts that can save lives and properties property.

Landslide couldLandslides can be triggered by either rainfall or <u>earthquakecarthquakes</u> in Taiwan (Dadson et al., 2003; Lee et al., 2004; Lin et al., 2008; Chen et al., 2011). In , especially the former. Taiwan, <u>monsoonswas invaded by</u>

- 5 several __and typhoons bring which brought great amounts of rainfall, up to 3,000 mm/yearevery year, eausing lots of and numerous landslides and cause casualties every year. Therefore, recognizing the area thatareas where rainfall-induced landslide landslide is an urgent issue, important and landslide susceptibility analysis was a general method. Here we(please check this sentence and rewrite). We adopted a statistical method for the analysis of landslide susceptibility __model in this study based on the assumption that the predisposing
- 10 factors <u>eausedthat cause which caused slope failurelandslides</u> in a region <u>wereare</u> <u>the samesimilar</u> and can be used for predicting the <u>locationlocations of as those which will generate</u> landslides in the future (Guzzetti et al., 1999). <u>There wereInIm</u> <u>previous researches</u>, several statistical models that have been proposed <u>as well asand</u> widely utilized in landslide susceptibility <u>analysis</u>, and <u>in recent years</u>, especially the logistic regression were one of the most used methods (Guzzetti et al., 1999; Lee et al., 2004, 2008a, 2008b, 2014). <u>Therefore</u>, we applied logistic regression (LR) in this study.
- 15 On the other hand, rainfall thresholds for landslides can be categorized as <u>either</u> statistical approaches <u>andor</u> deterministic approaches. In the former method, thresholds are decided by collecting historical landslide cases and analyzing their rainfall parameters <u>as well asand the</u> probability lines of rainfall conditions (Caine, 1980; Guzzetti et al., 2008). In the latter method, thresholds are decided by calculating <u>the</u> safety factors of each slope or grid with geomaterial and rainfall parameters (Terlien, 1998; Kim et al., 2010)._
- 20 Statistical rainfall thresholds for shallow landslidelandslides have been well discussed (Guzzetti et al., 2007). They can be classified mainly elassify into 5 categories including: intensity-duration (Brunetti et al., 2010; Zhou et al., 2014; Pradhan et al., 2017), accumulated rainfall-duration (Martelloni, 2011; Rosi et al., 2012; Vessia et al., 2014; Gariano et al., 2015; Rossi et al., 2017), accumulated rainfall (Corominas and Moya, 1999; Bell and Maud, 2000), intensity-accumulated rainfall (Hong et al., 2005)), and accumulated rainfall-accumulated rainfall (Osanai et al., 2010; Turkington et al., 2014).
- 25 Most of the studies mentioned above set up only one threshold for their study areaareas despite differences in spite of the difference in physical settings (geology, geomorphology, and meteorological conditionconditions) of that region the regions. Recently, some studies have subdivided their study areaareas into several homogeneous sub-zones in order to discuss the influence of physical settings on thresholds (Hong and Adler, 2008; Segoni et al., 2014, 2015; Lee et al., 2015; Segoni et al., 2015, 2016; Peruccacci et al., 2017). However, for a smaller area likesuch as slope units, the
- 35 different landslide susceptibility levels (high, moderate, and low). After that, we and establishingestablishedestablished their rainfall thresholds separately. BesidesFurthermore, we set warning signsalert levels by adopting the concept of a hazard matrix and examine if examined whether differentiated warning thresholds for different degrees of susceptibility existed. Moreover, it is essential to validate given the importance of validating the performance of a landslide early warning model, especially the false alarms and missed alarms, so as to make it feasible for further practical application (Calvello et al. 2015; Devoli et al.
- 40 2015; Piciullo et al. 2017; Segoni et al., 2018), therefore, we also adopted skill scores to verify our results.

2 Study area

Taiwan is located <u>atin</u> the western Pacific Ocean, on the convergent plate boundary zone of <u>the</u> Philippine Sea plate and <u>the</u> Eurasian plate. The orogenic uplift rate is 5 - 7 mm/year (Willett et al., 2003); however, the exhumation rate is also as high

5 as 3—_6 mm/year (Dadson et al., 2003) due to the fractured geological materials and the high mean annual precipitation up toof 2,500—_3,000 mm brought by typhoons and monsoons every year (Hsu, 2013). The frequent naturenatural disasters and high population density (23 million people overin 36,000 km²) makeof Taiwan make it one of the countries most exposed to multiple hazards (Dilley et al., 2005).

The study area-is, located in southern Taiwan (red box in Fig. 1), including theincludes a _-region-mosaic of 47 1:25,000
 scale maps (about 7,258.5 km²) and coveringcovers densely inhabited as well asand landslide-_threatening hillslopes. The elevation ranges from 3,243 meters in mountain areaareas to 0 meters in plain areaareas, while the gradient ranges from 87° 87° to 0°. Lithological0°. The lithological units units Geological settings are mainly sedimentary rocks composed of sandstone, shale, mudstone, and conglomerate in the Western Foothills, as well as metamorphic rocks composed of slate, argillite and metasandstone in the Central Range.

15 3 Available data Data and methodology

3.1 Data

3.23.1 Landslide inventory

Landslide <u>A-ILandslide</u> inventoriesy areis essential for the assessment of landslide susceptibility or spatio-temporal land changes (Van Westen et al., 2003; Guzzetti et al., 2012; Samia et al., 2017; Valenzuela et al., 2017). ThereIn this study, _____tThere
 were fourFour procedures for the construction generation of rainfall-induced landslide inventories were followed in this study. FirstlyFirst, barren lands the landslide inventories (Table 1) were interpreted manually from SPOT 5 images by drawing polygons in ESRI ArcGIS software. SecondlySecond, , the aerial photographs and the satellite images infrom Google Earth Ggoogle Eearth software were applied to identify if whether check identify if the barren lands landsareas were landslides or agricultural lands. Besidesland. In additionlands the locational correctness of inventories and confirm the type of landslides.

- 25 Besides, the polygons that situated in the slope whose gradient is higheron slopes having gradients greater than 55 degrees were marked as rockfallrockfalls according to the classification rules proposed by the Central Geological Survey, Taiwan (Central Geological Survey, 2008). Polygons which were marked as agricultural landsland or rockfallrockfalls were deleted from the inventories for the purpose of ensuring to ensure that This study used_only_only_shallow landslides werewould be analyzed in thisthe study. ThirdlyThird-study to analyze and the other types (e.g. rockfall) were filtered. _Thirdly, we randomly
- 30 selected landslides fromforom inventories randomly and verified to verified to verified the correctness of location the locations and boundary viaboundaries by fieldwork. Finally, event-based <u>triggered</u> landslide inventories inventories, including newnewly-generated landslides and expanded landslide and slide sdue to the event, were identified through the comparison of by comparing inventories before and after each rainfall event. In the end, there were 6 In this study, there were six 6 heavy rainfall events that triggered landslide landslide through the total of seven landslide.

3.33.2 Landslide occurrence time and field investigation

Rainfall conditions such as intensity, duration, <u>and accumulated rainfall that induced landslide landslides are</u> key data <u>while applying in the application of statistical method</u> to establish the rainfall thresholds for

- 5 landslideslandslides (Guzzetti et al., 2007, 2008; Brunetti et al., 2010; Peruccacci et al., 2017). In order to To analyze the rainfall conditions for each landslide case used in this study, a flowchart was proposed in this study (Fig. 2). -During field investigationinvestigations, we not only verified the correctness of the landslide inventories but also tried and interviewed local residents to try to inquireacquireidentify the landslide occurrence time from residents lived aroundtimes, since it is nearly impossible to get this information from is rarely included in landslide inventories. We gathered landslide occurrence time by
- 10 inquiring residents during field investigation or collecting reports in newspapers. Besides, detailed characteristics of landslides such as lithology, geological structure, joint, strength, area, depth and mechanism were also recorded during field work. Finally, 941 landslide cases including their occurrence time (date and hour) and characteristics of landslides were gathered for further analysis of the rainfall conditions. The accuracy of landslide occurrence times is hard to evaluate due to the lack of video records; however, we tried to interview focused on interviewing as many residents whose family was relatives were injured
- 15 or house washouses were damaged/-or-destroyed by the landslide as many-landslides as possible. Based on these impressiveGiven the deep impressions left by such memories, we believe that the quality of landslide occurrence timetimes might be improved. We gathered landslide occurrence time by inquiring residents during field investigation or collecting reports in newspapers. BesidesOn the other hand, detailed characteristics of the landslides, such as lithology, geological structure, joint, strength, area, depth and mechanism, were also recorded during the field work.
- 20 <u>Finally, there are 941 landslide cases, including their occurrence timetimes (date and hour) and the characteristics of the landslides, were gathered for further analysis of the rainfall conditions.</u>

3.4<u>3.3</u> **Slope units**

<u>Slope units were used for the analysis of landslide susceptibility in this study This study used slope unit that based on the features of geomorphology such as ridges and river valleys to analyze landslide susceptibility (Carrara, 1988; Carrara et al., 1991, 1995; Guzzetti et al., 1999; Schlögel et al., 2017; Yang, 2017). In order to To delineate the boundaryboundaries of slope unitsunitscorrectly, the 5×5m digital elevation models (5m 5m DEMs) were acquired from the Department-Ministry of the Interior, Taiwan. However, for the sake of reducing noises reduce noise, the DEM was smoothed and reduced to 10m resolution for the sake of reducing noises. we developed a Fortran program to obtain the smoothed and resolution-reduced 10×10m DEMS) by calculating the average value of each 2 by 2 grid in the 5×5m DEMs. The
</u>

30 resolution-reduced 10m DEM mightDEMs could generate some differences onin the morphological analysis, but the expected scale of the landslide susceptibility in this study iswas set to 1:25,000, so the differences that smaller than 12.5m could might be able to ignored according to the relationship between mapping scale and 5% acceptable error.-

This study followed the method proposed by Xie et al. (2004) to delineate delineating slope units according to the gullies and ridges. FirstlyFirst, gullies and watershedwatersheds were analyzed by successively using the spatial analystanalysis tools in ArcGIS-including: fill, flow direction, flow accumulation, stream link (we usedwith 2,000 used as the threshold) and watershed-successively. Secondly. Second, reverse DEMDEMs were generated by multiplying DEMs by -1-on DEM. Now. In the reverse DEMs, ridges became gullies in the reverse DEM and could be analyzed by the same methods used

in the first step. Thirdly, watershed Third, the watersheds of DEM the DEMs and reverse DEMDEMs were transformed from rasterrasterss to polygonpolygons for further editing by using the "Raster to Polygon" tool in ArcGIS and then cut by each other to delineate the slope units. Finally, slope units were modified manually according to Through the mapping concept proposed by Xie et al. (2004), the slope units were mapped automatically and modified manually (Fig. 3). Each slope unit was

- 5 given a unique code and separated into stable or unstable unit.Slope units were delineated according to the ridges and gullies as well as their aspect and gradient. aspect and gradient. It ishas been is suggested that the aspect in a slope unit should be within three adjacent directions; e.g., northwest, north-west, north, and north-eastnortheast. On the other hand, the difference in gradient should not be over 30 degrees in a slope unit, and the-slope units that-situated on flat areas, including alluviumalluvial deposits and terraceterraces, were deleted. BesidesIn addition, the area of each slope unit wasis set to around
- 10 <u>5 ha.,..; -t</u><u>Therefore, slope units that some smaller than 5 ha slope units were united to</u>combined with adjacent slope units and slope units that hose larger than 5 ha were split into several smaller ones. Moreover, <u>Besides</u>, slope units that delineated by parallel drainage on a dip slope should united aswere combined into one slope unit. After the editing, each slope unit <u>Moreover</u>, the area of each slope unit was given a unique code is set to around 5 ha. Therefore, some smaller slope units were united to adjacent slope units. Each slope unit was given a unique code and separated into stable or unstable unit for the sake of disaster

15 prevention.

Each slope units is characterized by several homogeneous parameters. We will add these parameters and a more detailed procedure.

3.53.4 Landslide Susceptibility Factors

- 20 Many factors could induce landslides, but each factor had different effect. This study initially selected some Several predisposing factors that might lead to landslides were selected initially in this study in order to suitable to construct a landslide susceptibility model for slope units.__These factors included included rock mass strength-size classification (RMSSC I-__VII), dip slope, average slope, variance of slope, ratio of steep slope, total slope highheight, average elevation, average curvature, variance of curvature, fault density, fold density, average wetness, rainfall intensity, total rainfall, 3-hour mean rainfall intensity (I₃), and 24-hour accumulated rainfall (R₂₄). The relationship between-relationships of these factors
- andto landslides Whenever a typhoon attacks Taiwan, Central Weather Bureau will issue alerts for typhoon. We therefore take the time of the first alert issued as the beginning of rainfall event and the time of canceling alert as the end of rainfall event to calculate the total rainfall. The reason we choose 3-hour mean rainfall intensity (I₂) and 24-hour accumulated rainfall (R₂₄) as factors will explain in section 4.2... They These factors could be analyzed through graphic discrimination, including that
- 30 included success rate curve, probability of failure curve, and difference between landslide and non-landslide groups (Lee, 2014). After that, Finally, we applied factor correlation analyses analysis analyses were applied to delete highlighly and deleted the high related ive factors for the sake ofto keep the factors used in the landslide susceptibility model-were as independent as possible (Table 2).

In terms of geological factors, the lithologyIn terms of geological factors, Lilithology iwithology of a aiss always a location is essential essential eritical <u>factorelement for the when analysis of of fingin landslide susceptibility.</u> Solve stability. However, in-our study area, there are hadwere more than 50 detailed types of lithology in this study area, which iswaswas unfavorable for the analysis. Therefore, we adopted the 1:25,000 rock mass strength-size classification (Franklin, 1975) maps from the Central Geological Survey, Taiwan, <u>were adopted</u> to replace the use of lithology (Franklin, 1975; Central Geological Survey, 2008). Besides, In addition, the dip Dip-slope inventory used in this study was-interpreted manually from

1:5,000 aerial photographs by the Central Geological Survey, Taiwan, waswa also adopted (Central Geological Survey, 2008). On the other hand, the The The fold density waswas also calculated with by dividing the total length of all the folds-divided by the total area in each slope unit.

For morphological factors, the The average slope and the variance of slope werewereas obtained by averaging and

- 5 calculating the standard deviationdeviations of all the grid cells in the slope unit separately. BesidesIn additionThe ratio of steep slope was calculated by dividing the area that greater than 30 degrees by total area of slope unit. Besides, shallow landslides are prone to occur on steep slopes; therefore, we also used "the "ratio of steep slopes" to present how many steep slopes are thereexisted in a slope unit. It was found after trial and error that a threshold of gradient higher than 30 degrees had a higher relationship with landslide susceptibility after trial and error. Thus, we calculated the area where the
- 10 gradient iswas greater than 30 degrees (A >30) as well as the total area (A total) of each slope unit-and. Therefore, the ratio of steep slope can thereforecould be calculated withby dividing A >30 divided by A total. On the other hand, the Tthe average curvature and variance of curvature could also be calculated by the same method as slope in the ArcGIS software. The average wetness iswas calculated by averaging the wetness index of grid cells in a slope unit. This factor represents the effect of morphology on soil wetness. When the drainage area is larger and the slope is gentler, the water content in the soil wouldwill also be higher and therefore make a slope more prone to failure. Wetness The wetness index can be calculated according to the method proposed by Wilson and Gallant (2000) as followedfollows:

$$\omega = ln(\frac{A_s}{\tan\theta})$$
(1)

where ω is wetness index, A_{s} is the drainage area of a specific grid cell, and θ is the slope of the grid.

For triggering factors, The fold density was the total length of all the folds divided by the total area in each slope unit.
 The average wetness was calculated by following the method proposed by Wilson and Gallant (2000).

The average wetness was calculated by following the method proposed by Wilson and Gallant (2000), we collected hourly the rainfall data offrom 423 rain stationsgauges provided by Central Weather Bureau, Taiwanin Taiwan (96 of them arewhich were located in our study area, shown in Fig. 1) and analyzed both the 3-hour mean rainfall intensity (1₃) as well asand the 24-hour accumulated rainfall (R₂₄) of each stationgauge infor each rainfalltwo typhoon events in Table 2. After that, we used the linear mode of ordinary kriging and applied the default setting in Surfer software to obtain the rainfall distribution of the whole study area. The hourly rainfall records during two typhoon events in this study were collected from the Central Weather Bureau, Taiwan. According to these rainfall records, the 3-hour mean rainfall intensity (1₂) and 24-hour accumulated rainfall (R₂₄) in each rainfall station were calculated. Besides, Kriging interpolation method was applied to generate the precipitation distribution map for the whole study area. Whenever Taiwan has a typhoon attacks Taiwan, event, the Central Weather Bureau will issue alerts for thetyphoonthe sake offissues disaster prevention alerts. We therefore takecounted the time of canceling that the alert was cancelled as the end of the rainfall event to calculate the total-rainfall amounts. The reason we chooseOur reasons for choosing 3-hour mean rainfall intensity (1₃) and 24-hour accumulated rainfall (R₂₄) as factors will explainbe explained in detail in section 4.2.

35 4 Methodology

4.1 Landslide susceptibility analysis

There are many predisposing factors that might lead to landslides, but the effectiveness of each factor is different. The main

purpose of landslide susceptibility analysis <u>is determining</u>to <u>determinei</u>was to <u>determining</u>e <u>the effectiveness of each</u> <u>predisposing factor and the</u> relative possibility of landslide occurrence <u>in a specific area</u>. There are severalSeveral methods for <u>the analysis ofcan be used to analyze landslide susceptibility</u>. The However, t<u>The</u> deterministic <u>methodmethods</u> uses a physical model and geotechnical material properties to determine the safety factor of safety of slopes;; however, precise parameters of

- 5 <u>materials required geotechnical material properties which wereare</u> difficult to obtain, <u>especially for on especially for for a</u>the regional <u>scallandslide susceptibilityscale</u> (Montgomery and Dietrich, 1994; Van Westen and Terlien, 1996). The qualitative <u>methodmethods and semi-quantitative methodmethods rely on depended on the experience and knowledge of the experts who carried out the analysis₁; however, these results might <u>be differentvary from one expert to another.</u> —The machine learning <u>methodmethods</u> uses <u>lots ofmultiple</u> samples to build <u>a model by trial and error</u>; however, it is <u>always time</u></u>
- 10 <u>consumingeonsuming</u> required more training time to build model by trial and error (Gorsevski and Jankowski, 2010; Yeon et al., 2010; Yilmaz, 2010; Marjanovic et al., 2011; Lee et al., 2012; Song et al., 2012). <u>The In order to avoid these difficulties</u>, this study adopted statistical method<u>methods</u>_-also requires lots ofnumerous samples for the training; however, it is more efficient, especially when dealing with regional--scale analyses, and can avoid the uncertainty of material parameters as well as the difference of differences in _expert experiences experience. Recently, nonlinear analysis, one of thea statistical method,
- 15 that suitablehas been used for the analysis of complex landslide phenomenon includingphenomena. Methods such as logistic regression (Yilmaz, 2010; Lee et al., 2012, ; Lee et al., 2014, ; Lee et al., 2015; Schlögel et al., 2017) and discriminant analysis (Lee et al., 2004, 2008a, 2008b) were often used to analyze landslide susceptibility. HereIn this study, weHere welogistic regression (Yilmaz, 2010; Lee et al., 2012; Lee et al., 2014; Lee et al., 2015; Schlögel et al., 2017) and discriminant analysis (Lee et al., 2004, 2008a, 2008b) were often used to analyze landslide susceptibility in the statistical methods Besides, due to analyze landslide susceptibility in the statistical methods.
- 20 landslide was a complex phenomenon, a nonlinear analysis was more suitable for this study. Recently, logistic regression (Yilmaz, 2010; Lee et al., 2012; Lee et al., 2014; Lee et al., 2015; Schlögel et al., 2017) and discriminant analysis (Lee et al., 2004, 2008a, 2008b) were often used to analyze landslide susceptibility in the statistical methods, therefore, LR was applied logistic regression (LR) to evaluate the susceptibility of each slope unit (Guzzetti et al., 1999; Ayalew and Yamagishi, 2005). The LR function was can be expressed as followedfollowsfolloweds:

25
$$P = \frac{1}{1 + e^{-z}} P = \frac{1}{1 + e^{-z}}$$

$$z = \sum_{i=1}^{m} L_i w_i + \sum_{j=1}^{n} F_j w_{m+j} + C z = \sum_{i=1}^{m} L_i w_i + \sum_{j=1}^{n} F_j w_{m+j} + C z$$

$$\frac{(3)}{P = \frac{1}{1 + 1}}$$

30 (1<u>2</u>)

35

$$z = \sum_{i=1}^{m} L_{i} w_{i} + \sum_{j=1}^{n} F_{j} w_{m+j} + C$$

where P isiwas the landslide susceptibility; L_i iswais RMSSC factor (L₀₁ to L₀₇ in Table 2); F_j was other factors (F_{01} to F_{10} in Table 2); w_i and w_{m+j} is are was regression coefficient coefficients, and C is a was constant. The sSix event-based triggered landslide inventories in this study were used to label if whether or not landslides occurred in the slope units occurred landslide or not.. After that, all the slope units were divided randomly into two parts randomly groups, T_{27} oOne was for training the model

-<u>(23)</u>

and the other was for the validation. T

he index that indicating landslide/non-landslide was set as the dependent variable, and all the landslide susceptibility factors were set as covariates underin SPSS software while for training of the model. After iteratediterative training, the regression coefficients of each landslide susceptibility factor, as well as the success rate curve (SRC), the prediction rate curve (PRC), and the area under the curve (AUC), were reported in SPSS software. The AUC can be used to examexamine if the model

- 5 predicts landslides well-or not, and the regression coefficients can be used for the prediction of landslide susceptibility. During the process of training, there are several details that should payrequired attention-to. Due to that. Because _____The six event-based triggered landslide inventories in this study were divided into two parts randomly. One was for training model and the other was for the validation. In addition, the non-landslide samplessamples data were manymanyuch more than outnumbered the landslide samples, samplesdata, so we randomly selected the same amounts equal numbers of non-landslide and landslide
- 10 samplessamples data randomly for the training so as to avoid the effect of difference in quantity. Besides In addition, quantity in SPSS software. Besides, different samples might leadcould have led to different results when selecting non-landslide samples randomly. In order to To reduce this effect, Besides, we prepared several sets of randomly selected samples (especially non landslide data) were also tested for the analysis of landslide susceptibility in order to test if ensure the model werewas stable enough; i.e., the AUC would not variate vary not variate geverely when validating models with different sets
- 15 of samples. with alternating samples. Finally, the individual landslide susceptibility susceptibilities distribution of each the the failure ratio and landslide susceptibility of in __each slope unit wasunits were calculated with this model and classified into plotted and then used in classifyingied into landslide susceptibility level (high, moderate and low susceptibility level levels...).

4.2 I₃-R₂₄ rainfall index and thresholds

- Rainfall-induced landslides are always triggered by either high intensity rainfall or high accumulated rainfall (Larsen and Simon, 1993; Corominas and Moya; 1999; Yu et al., 2006). In order to find outTo identify rainfall indexes responsible for landslides, the triggering rainfall, including the rainfall intensity (I₁, I₂, I₃, I₄, I₅, I₆) and accumulated rainfall (R₆, R₁₂, R₂₄, R₄₈, R₇₂) of different time windowwindows of each landslide case_x wereas analyzed according to the landslide occurrence time...It was found The results revealed that there were 218 landslide caseslandslides occurred within the 3 hours right afterfollowing
 the highest rainfall intensity, and 242 eases-occurred within the 3 hours right afterfollowing the 2nd or 3rd highest rainfall
- 2.5 the highest fundal meets(y, and 2-2 closes occurred within <u>me_</u>) hours fight after <u>the landslide</u> cases gathered in this study (<u>Table</u> 3). <u>This indicated</u> From these results, it became clear that in Taiwan, I₃ is the most <u>keyimportant index for landslides induced</u> by rainfall of short duration but high intensity rainfall in Taiwan. On the other hand, there were 481 landslide cases[andslides occurred at the time close to the end of <u>the</u> rainfall eventevents (i.e., induced by high accumulated rainfall), accounting for
- 30 about 51% of the total cases (Table 3). –Furthermore, analysis of the different accumulated rainfall indexes showed that 24-hour accumulated rainfall hashad the lowest coefficient of variation (Table 4), indicating that this index was less dispersive than others and might be more suitable for serving as an accumulated rainfall index infor establishing rainfall thresholds. Establishing rainfall threshold. The coefficient of variation can be calculated as followedfollows:

$$C_{v} = \frac{\sigma}{\mu} \tag{4}$$

- 35 where C_{ν} is the coefficient of variation; and σ and μ are the standard deviation and average of accumulated rainfall of all the cases used in this study respectively.
 - Based on these data and <u>literaturesprevious studies</u> (Cheung et al., 2006; Liao et al., 2010), 3-hour mean rainfall intensity (I₃) and 24-hour accumulated rainfall (R₂₄) were therefore respectively chosen as the short-term and long-term rainfall index

respectively indexes for the establishment of the rainfall threshold (Fig. <u>343</u>). We choose chose 3-hour mean rainfall intensity here instead of 3-hour accumulated rainfall for the purpose to focus on rainfall of emphasizing the short duration but high intensity rainfall. Similarly, we choose chose 24-hour accumulated rainfall for the sake to focus on rainfall of emphasizing the long duration but low intensity rainfall.

- 5 Finally, rainfall thresholds were decided by plotting the I₃ and R₂₄ rainfall index of historical landslides in the I₃–R₂₄ diagram (Fig. <u>454</u>). Here we used the ellipse as the threshold line, and the parameterparameters *a* (semi-major axis) as well as and *b* (semi-minor axis) of the ellipse were set according to the slope of best fit line gettingobtained from the least square method. Different thresholds Thresholds such as 90%, 60%, 30%, 15% were determined according to the percentage of historical cases that could be enveloped under the threshold line; e.g., the 90% threshold (T_{90%}) included 90% of the historical
- 10 cases<u>and aA</u> higher threshold <u>indicatesalso-indicates</u>d a more dangerous condition<u>for the occurrence of landslidelandslides</u>. <u>The original warning values of I₃ and R₂₄ of the 90%, 60%, 30%, 15% thresholds were equal to the semi-minor axis and semi-major axis of each threshold respectively. After that, I₃ was rounded to the nearestby 5 mm/h and R₂₄ was rounded byto the nearest 50 mm for operational purpose, e.g., purposes, such as the evacuation of residents.</u>

4.3 Landslide early warning model and validation

- 15 Landslide_The landslide_early warning model in this study considered both landslide susceptibility as well asand rainfall thresholds and was given warning signsalerts were determined by using the concept of a hazard matrix. As mentioned above, the LR method was applied to analyze the susceptibility of each slope unit. After that, all the slope units were categorized into high, moderate, and low susceptibility levellevels. We consequently established rainfall thresholds for each susceptibility level separately and then gave warning signs includingset alerts of red, orange, yellow and green according to the dangerous level
- 20 <u>of danger</u>.

<u>HighFor high High_susceptibility slopes (Table 5Fig. 5), they</u>) might be more susceptible to rainfall, <u>hence. Hence.</u> the <u>warning sign wasalerts were</u> set as red (extreme <u>dangerousdanger</u> level) <u>whenfor</u> rainfall <u>condition exceedsconditions</u> <u>exceeding</u> the 60% threshold line; orange (high <u>dangerousdanger</u> level) <u>when rainfall condition wasfor those</u> between the 60% and 30% threshold lines; yellow (medium <u>dangerousdanger</u> level) <u>when rainfall condition wasfor those</u> between the 30% and

- 25 15% thresholdsthreshold lines; and green (low dangerousdanger level) whenfor rainfall condition wasconditions lower than the 15% threshold line (Fig. 5). For moderate-susceptibility slopes, (Fig. 6), the warning sign wasalerts were set as red whenfor rainfall condition exceeds conditions exceeding the 90% threshold line; orange when rainfall condition wasfor those between the 90% and 60% threshold lines; yellow whenfor rainfall condition wasconditions between the 60% and 30% thresholds lines; and green whenfor rainfall condition wasconditions lower than the 30% threshold line. For low
- 30 Low-susceptibility slopes, they might should be less susceptible to rainfall, hence. Hence, there was no red sign and the warning sign wasalerts were set as orange whenfor rainfall condition exceeds conditions exceeding the 90% threshold line; yellow when rainfall condition for those wereas between the 90% and 60% thresholds threshold lines; and green when rainfall condition was for those lower than the <u>660%</u> threshold line.
- There are severSeveral methods can be used for the validation of a landslide early warning model (Segoni et al., 2014, 2018; Gariano, 2015; Rosi et al, 2015; Piciullo et al., 2017; Krøgli et al., 2018). According to the analysis of Segoni et al. (2018), compiling a contingency matrix and calculating skill scores areis the most commonly used method in recent years, therefore. Therefore, we applied this method and validatequantitatively validated our model with probability of detection (POD, also known as hit rate), probability of false alarm (POFA, also known as false alarm ratio) and probability of false detection (POFD, also known as false alarm ratio) and probability of false alarm ratio.

The contingency matrix is shownpresented as Table 5. There are four outcomes when comparingComparing the observed events and the forecasted event includingevents produces four outcomes: True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN). Due to that the Because in Taiwan, warnings of naturenatural hazards are always issued by taking a village as a unit-in Taiwan, here we validated our model with a villages on the village scale. TP indicates indicated

- 5 the number of villages thatfor which warnings of slope units arewere issued and landslides dodid occur, while TN indicates indicated the number of villages thatfor which no warning iswas issued and there is also no landslide occurred. On the other hand, FP indicates indicated the number of villages that do not occur landslides butfor which warnings arewere issued but no landslides occurred, also known as false alarms, while FN indicates indicated the number of villages that occur landslides butfor which no warnings are notwere issued, but landslides did occur; i.e., missed alarms. Besides In addition, in our study,
- 10 we define the warning sign of defined red (extreme dangerousdanger level) and orange (high dangerousdanger level) alerts as warnings issued for the evacuation; i.e., the alarm zone in our model. On the other hand, the warning sign of yellow (medium dangerousdanger level) and green (low dangerousdanger level) arealerts were considered asto indicate no need for the evacuation; i.e., the no alarm zone in our model. The POD, POFA, POFD, POFA can be calculated by the following equations:

$$POD = \frac{TP}{TP + FN}$$
(5)
15
$$POFD = \frac{FP}{TP + FP}$$
(67)

$$POFA = \frac{FP}{FP + TN}$$
(76)
20
$$POFD = \frac{FP}{TP + FP}$$
(7)

The ranges of POD, POFA, POFD, POFA are all between 0 to 1, and their optimal values are 1, 0 and 0, respectively.

5 Results and discussions

5.1 Landslide susceptibility analysis

- 25 After several times of calibration<u>calibrations</u>model calibration, the resultant model was obtained. The coefficients for each factor of LR wereare given in table<u>Table</u> 2, and the landslide susceptibility of each slope unit <u>waswasere</u> also <u>calculated</u>.defined<u>calculated</u>. In order to<u>To</u> evaluate the quality of a predicted model, the success rate curve (SRC) and prediction rate curve (PRC) (Chung and Fabbri, 1999) were mapped, and then the area under the curve (AUC) was used to describe the model's ability of distinguishingto distinguish landslide and non-landslide (Yesilnacar and Topal, 2005). A higher
- 30 AUC value indicated a better model for the prediction of landslides. If the AUC value was 0.5, it meant that the model didn'tdid not predict the occurrence of the landslide better than a random approach. If the AUC value was close to 1.0, the capability of the model that interpreting for predicting a landslide was perfect perf

<u>In our study, the Tthe AUC was AUCs were</u> 0.745 and 0.691 in training and <u>validating validation</u> model respectively, indicating that our LR model could identify 60% of the landslides in the top 25% and 30% of the highest susceptibility areas

during training and validation (Fig. <u>675</u>). These results showed that the LR model was <u>acceptableacceptablestable</u> and nice in <u>both the</u> training as well as and the validation. For a statistical landslide susceptibility analysis, it is essential to use as many samples as possible. However, we used slope units instead of grid units in this study for the purpose of the application onto disaster prevention. This leadedled to the reduction of samples, since one slope unit might equal to hundreds of grids. Therefore,

5 <u>our AUC might not be soconsidered high comparedin comparison to a grid-based landslide susceptibility model. It also</u> represented that LR was useful in landslide susceptibility analysis.

Having enough samples were the foundation of statistic method. Due to our using slope units, the amount of samples were less than a traditional grid method, so it was not easy to establish a well performed model compared with a grid based landslide susceptibility model. For the sake of increasing more samples for analysis, we integrated the data from several events, but it might also brought some noises for the training. Therefore, filtering unfavorable or unsuitable samples were required.

On the other hand, <u>in order to notavoid for avoiding over</u>-training, it was necessary to validate the capability of <u>the</u> model. One common method <u>was dividingis to divide the</u> study area into sub-regions such as left and right, one for training and the other for validation (Chung and Fabbri, 2008). But <u>itthis method</u> might <u>losecause the loss of alost a</u> training pattern in a small or particular geological region if the study area <u>wasis</u> extensive. To overcome this problem, we <u>suggested using used</u> multi-

15 event data infrom the same area for training and testing. The data used in this study were therefore <u>randomly</u> divided into two parts randomly portions, and several sets of data were tested. This <u>approach</u> would also solve the problems mentioned above.

5.2 I₃-R₂₄ rainfall threshold

10

We gathered totallya total of 941 landslide cases in this study and picked outselectusced 240 cases located in southern Taiwan, includingconsisting of 110 high-_susceptibility cases, 84 moderate-_susceptibility cases, and 46 low-_susceptibility cases, to

- 20 establish a susceptibility-based regional landslide early warning model. The ellipse-shaped I₃-R₂₄ rainfall thresholds for 3 different landslide susceptibility slopes were shownare presented in Table 6 and Fig. 876. For the purpose of practical use, the original threshold values of I₃ and R₂₄ (as shown in the parentheses in Fig. 8) were separately rounded to the nearest to by 5 mm/<u>hhr</u> and the nearest 50 mm separately, as shown in the parentheses in <u>_</u>Table 6<u>Fig. 7</u>. It could be was found that the threshold values gradually decreased as the susceptibility of the slope decreased for the same threshold (e.g., Took) and that the threshold the threshold values gradually decreased as the susceptibility of the slope decreased for the same threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold values are threshold (e.g., Took) and that the threshold (e.g., Took) are the took of the took of
- 25 values-also gradually indecreased as the susceptibility of slope units deincreased for the same warning-signalert level, indicating that greater rainfall amounts would be needed when issuing alters on less susceptible slope units. TheseTheseis results showed that establishing rainfall thresholds according to different landslide susceptibilitysusceptibilities and then setsetting warning signsalert levels by adopting the concept of a hazard matrix could not only provide provided differentiated thresholds but also avoid an overestimate avoided the over- or underestimate underestimation of the thresholds for slopes.-
- 30 After the establishment of the landslide early warning model, we leaded convertedled the model to an early warning system (EWS) which connected to the QPESUMS, a nearwhich provides nearly real-time radar rainfall data, for disaster prevention.
 <u>According toBased on the In addition, Table 7. showed the warning signsalertswwarning signs present in the system, and the corresponding dangerousdanger levels as well as and suggested actionactions for residents around the warning slope are shown in Table 6. During a yellow signalert, residents should pay attention to whether there are listen for further announcements or</u>
- 35 not and readyprepare for an evacuation if the sign turnsalert is raised to orange. WhileWhen an orange signalert is issued, residents should evacuate as quickly as possible because landslides are pronelikely to occur, according to the validations shown in the next section. LastlyFinally, when the warning sign goes to a red, forced alert is issued, evacuation mightmay need to preventbe enforced to protect residents from getting injuredinjury.

We validated our model with two kinds of data: (1) three disastrous shallow landslides <u>cases</u> in 2016 and the occurrence <u>timetimes</u> provided by witnesses; (2) <u>a</u> landslide inventory of two historical typhoon events and the occurrence <u>timetimes</u> reported by newspapers.

5 For the The first set of validation data-one, it showed that orange or redyellow alerts could have been issued in advance for all of the disastrous landslides cases could be warned with orange or red sign in advance before landslide the landslides occurred, according to the rainfall snake line in the I₃-R₂₄ diagram (Fig. 999-; Table $\frac{87}{7}$).

<u>The</u> Shihwen landslide occurred on a low-<u>-</u>susceptibility slope. <u>FromForem</u> the rainfall histogram and I_3 -R₂₄ diagram (Fig. 998(a)), we knew that the occurrence time was quite close to the end of <u>the</u> rainfall event, and <u>that</u> the I_3 was only 2.3

- 10 mm/<u>h</u>hr while the R₂₄ was 507.5 mm, indicating that accumulated rainfall might <u>behave been</u> the principal cause of this case. <u>Rainfall The rainfall</u> snake line showed that <u>the warning sign turned on September 14</u>, the alert was raised to yellow at 10:00, <u>14 Sep</u>, 14 th <u>(authors mention "14th", "15th", and others; if they are days, I suggest using the format dd mm, which results more clear.)</u> and <u>soon turnedthen</u> to orange at 11:00, <u>14 Sep</u> during the downpour; then it was a little bit let up. Then the precipitation-rate fell for several hours, and the <u>warning sign turned backalert was lowered</u> to yellow. However, when it rained
- again, the warning sign also turned alert was raised back to orange again at 18:00, 14-Sep as well as _____ and at 23:00, 14-Sep and finally the landslide occurred at 05:00, 00 September 15-Sep -15th.

<u>The</u> Zhongmin landslide occurred on a high-susceptibility slope. <u>FromForom</u> the rainfall histogram and I₃-R₂₄ diagram (Fig. 989(b)), it <u>could bewas</u> found that the occurrence time was also quite close to the end of <u>the</u> rainfall event and <u>that</u> the I₃ was 8.3 mm/<u>hhr</u> while the R₂₄ was 479 mm. The high rainfall intensity (74 mm/<u>hhr</u> at 04:00, on September 28th, 28-

- 20 SepSeptember) as well as <u>)</u> and accumulated rainfall might both result in <u>have contributed to</u> this <u>landslide</u>. Rainfall <u>_____cascevent</u>. The rainfall snake line showed that <u>on September 28</u>, the warning sign turned to yellow and then quickly turnedalert was raised to orange and red at 04:00, <u>28_Sep28th</u> during <u>the</u> high intensity rainfall mentioned above. After that, although it let up<u>the</u> rainfall soon_fell off, the landslide finally-occurred 6 hours later, at 10:00, <u>on September 28</u>, <u>Sep28th</u> during <u>thean</u> orange signalert.
- The Houcuo landslide also occurred on a low-_susceptibility slope. FromForom the rainfall histogram and I₃-R₂₄ diagram (Fig. 998(c)), we knewfound that the occurrence time was almost nearclose to the point of time that the highest rainfall intensity showed in the rainfall event, and the I₃ was 24.3 mm/hhr while the R₂₄ was 291.3 mm, indicating that high rainfall intensity might behave been responsible for this case. Due to this intensity, the rainfall snake line showed that the warning sign turnedalert was raised from green to yellow and then to orange inwithin just one hour-during, from 03:00 -to 04:00, on
 September 28-Sep, and 28th, the landslide also occurred at around 03:30, 28-Sep28th.

For the second <u>set of validation dataone</u>, we applied <u>Kriging the kriging</u> method to interpolate <u>spespatialeial</u> rainfall data and analyzed the <u>warning sign of alerts for</u> each slope unit hour-_by-_hour. <u>It The results</u> showed that the <u>hitting ratiohit ratesios</u> in <u>the</u> two historical typhoon events were all <u>higher enoughsufficiently high</u>, according to the accumulative warning numbers relative to the numbers of landslide slopes (Fig. 10.;;10;9.; Table 8.).).98.).

- During Typhoon Mindulle in 2004, there were landslides occurred in 10,911 slope units that occurred_once landslide, including 5,129 high-_susceptibility slopes, 2,750 moderate-_susceptibility slopes and 3,032 low-_susceptibility slopes. According to thenewspaper reports in newspapers, several landslides occurred at 10:00 and between 15:00 - and 16:00, on July2nd, July 2-Jul, 2004; however, most of the landslides occurred duringbetween 06:00 - and 13:00, the next day, 3rd Julyy 3-Jul, 2004 (blue dashed box in Fig. 10109(a)). From the warningalert history (Fig. 10109(a)), it could bewas found that the
- 40 peak number of orange and red signs fittedalerts matched the reported occurrence time quite well. Besides, there were times

<u>quite well. In addition, orange alerts, indicating the need for evacuation, had been issued for 8,283 slope units hadever been</u> warned as orange sign_during the whole event, which is the sign for evacuation, during the whole event, accounting for 75.9% of the slope units <u>that where landslides occurredonce landslide-in this event</u>.

- Typhoon Haitang in 2005 was another event of concern. There were-Landslides occurred in 10,804 slope units-once landslide, including 2,592 high-susceptibility slopes, 2,355 moderate-susceptibility slopes and 5,857 low-susceptibility slopes. According to thenewspaper reports-in newspapers, landslides occurred frombetween 05:00, on July 19-Jul19th-to and 06:00, on July 20-Jul20th, July, 2005 (blue dashed box in Fig. 10109(b)). From the warningalert history (Fig. 10109(b)), it could bewas found that landslidelandslides occurred rightimmediately after the number of orange and red signsalerts increased sharply, and the peak number of orange and red signsalerts also fittedmatched the reported occurrence timetimes quite well.
- 10 On the other hand, there wereorange alerts had been issued for 10,245 slope units hadever been warned as orange sign during the whole event, accounting for 94.8% of the slope units that where landslides occurred landslide in this event.once landslide. These results revealed that our model could provide valuable information for evacuation and disaster prevention.

BesidesIn addition, the second set of validation data werewas also used to validate the warnings issued for villages during two typhoon events by adopting the contingency matrix and skill scores. According to the event-based landslide inventories,

- 15 if thereany landslides were any landslide-located in a village, the village would be was classified as "Yes" offor observed events. On the other hand, if there were anyIf orange or red warning signalerts were issued for slope units in a village, the village would be was classified as "Yes" offor forecasted events. Based on these rules, the numbers of True Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) can be were counted and the skill scores can also be were calculated (Table 9). The probabilities of detection (PODPODs) of the two typhoon events were 0.961 and 0.874 respectively, indicating
- 20 that most of the villages that occurred where landslides occurred could behave been warned in advance. The probabilities of false detection (POFDPOFDs) of the two typhoon events were 0.280 and 0.667 respectively, suggesting that the model performed well infor Typhoon Mindulle but might not be so perfect inwellideal for Typhoon Haitang. Lastly, the probabilities of false alarm (POFAPOFAs) of the two typhoon events were 0.120 and 0.110 respectively, which meant that our model would not issue too manyan excessive number of false alarms and was feasible for disaster prevention.

25

6 Conclusions

In order to verify if the difference in susceptibility might lead to a difference in warning threshold, we divided slope units into three susceptibility levels (high, moderate, low) based on the results of Logistic Regression (LR) and established their rainfall thresholds separately in this study. I₂-R₂₄-rainfall index, a combination of short-term as well as long term rainfall index, were used for the establishment of rainfall thresholds. After that, three warning signs including yellow (advisory), orange (watch) and red (warning) were set by adopting the concept of hazard matrix. It was found that the warning thresholds were different for each susceptibility level and gradually decreased as the susceptibility of slope increased. Validations from three disastrous shallow landslides in 2016 showed that they can be warned in advance before landslide occurred and validations from two serious historical typhoon events also showed that the hitting ratio of our early waring model were 75.9% and 94.8% the able to provide differentiated warning thresholds for different susceptibility levels. This study triedattempted to establish regional rainfall thresholds for shallow landslides according to their landslide susceptibility levels and set warning signsalerts with the concept of a hazard matrix in order to provide -a more detailed results for disaster mitigation.

Logistic Regression (LR), one of thea statistical method, was applied in this study to analyze the landslide

susceptibilitysusceptibilities of slope units. The area areas under the curve (AUC) was were 0.745 and 0.691 in the training and validating validation respectively. Due to our using use of slope units instead of grid units in this study for the purpose of the application onto disaster prevention, the amounts number of our training samples were lesswas lessow, since one slope unit might equal to-hundreds of grids. Therefore, our AUC might not be soconsidered high as compared to a grid-based landslide susceptibility model, but it was still acceptable for practical use.

- This study also examined the relationship between rainfall indexes and the occurrence of landslides. From 941 landslide cases we gathered, it was found that 3-hour mean rainfall intensity (I_3) and 24-hour accumulated rainfall (R_{24}) were the most dominant short-term and long-term parameters that-responsible for rainfall-induced landslides in Taiwan. because there were $_3$ There were for 460 cases (about 49%) occurred within the 3 hours right afterfollowing the highest, 2^{nd}
- 10 and 3rd rainfall intensity intensities, while 24-hour accumulated rainfall had the lowest coefficient of variation than other of the long-term rainfall indexes. The I₃-R₂₄ rainfall index werewas therefore used for the establishment ofto establish rainfall thresholds.

We categorized the slope units into 3 landslide susceptibility levellevels (high, moderate, low) and then separately established a susceptibility-based regional rainfall threshold-separately. We also set three warning signs alert levels, including

- 15 red (extreme dangerousdanger level), orange (high dangerousdanger level), and yellow (medium dangerousdanger level)), by adopting the concept of a hazard matrix for the purpose of the application onto evacuation decisions. It was found that the threshold values gradually increased as the susceptibility of slope units decreased for the same warning signalert level, indicating that it needed moregreater rainfall amounts would be needed when issuing alters on a less susceptible slope units. Validations of using three disastrous shallow landslides in 2016 and two landslide inventories of historical typhoon events
- 20 showed that, for the landslide cases in 2016-could be warned with, orange or red sign in advance alerts could have been issued before landslidethe landslides occurred and the hitting ratiohit ratesios of the warningsalerts issued for slope units in the two historical typhoon events were 75.9% and 94.8% respectively, both of which were all higher enoughare sufficiently high for a landslide early warning model. BesidesIn addition, the skill scores that applied to the validation of warningsalerts issued for villages during two typhoon events showed that the probability probabilities of detection (PODPODs) were 0.961 and 0.874,
- 25 the probability probabilities of false detection (POFDPOFDs) were 0.280 and 0.667, while and the probability probabilities of false alarm (POFAPOFAs) were 0.120 and 0.110 respectively, indicating that our model might be able to put in use used for landslide early warning warnings.

It could becan concluded that classifying landslide susceptibility and establishing rainfall thresholds separately not only provide provides refined thresholds but also avoid an overestimate avoids over- or underestimate underestimate of the

- 30 thresholds for slopes, especially when considering the application onto disaster prevention. In order to verify if the difference in susceptibility might lead to a difference in warning threshold, we divided slope units into three susceptibility levels (high, moderate, low) based on the results of Logistic Regression (LR) and established their rainfall thresholds separately in this study. I₂ R₂₄ rainfall index, a combination of short term as well as long term rainfall index, were used for the establishment of rainfall thresholds. After that, three warning signs including yellow (advisory), orange (watch) and red (warning) were set
- 35 by adopting the concept of hazard matrix. It was found that the warning thresholds were different for each susceptibility level and gradually decreased as the susceptibility of slope increased. Validations from three disastrous shallow landslides in 2016 showed that they can be warned in advance before landslide occurred and validations from two serious historical typhoon events also showed that the hitting ratio of our early waring model were 75.9% and 94.8% respectively. It could be concluded that classifying landslide susceptibility and establishing rainfall thresholds separately might be able to provide differentiated
- 40 warning thresholds for different susceptibility levels.

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Figure 1: Geomorphological and geological settings of <u>the</u>study area. <u>The elevation ranges from 3,243 meters in the eastern</u> <u>mountain area to the sea level in the western plainplains area. Lithological units are mainly metamorphic rocks in <u>the</u> Central Range and sedimentary rocks in the Western Foothills. Rainfall data of 423 stations in Taiwan (96 of which are located in the study area)</u>

5 were collected for the interpolation and analysis of <u>the triggering rainfall of landslides.</u>



















Figure 2: Flowchart of landslide occurrence time gathering during field investigation (left), <u>locationlocations</u> of landslide cases with the occurrence <u>timetimes</u> used in this study (middle), and the pictures of <u>inquiringinterviewing</u> residents (right). <u>LandslideTo</u> <u>improve the quality of this key information, landslide occurrence time was inquired</u>times were obtained from local residents-lived <u>around</u> ______, especially those whose family wasrelatives were injured or <u>house washouses were damaged</u>/<u>-or</u>_destroyed by the landslide, in order to improve the quality of this key information landslides.



Figure 3:

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10 Figure 3: A display of Slope unit.



Figure 4<u>2</u>: 3-hour mean rainfall intensity (I₃) and 24-hour accumulated rainfall (R₂₄) were used as short-term and long-term rainfall indexindexes for the establishment of rainfall thresholds.





Figure <u>454</u>: Establishment of I₃–R₂₄ rainfall thresholds for shallow landslides<u>. The best fit line was derived</u> - by least square method, and the ratio of a and b werewas used as the ratio of the semi-major axis and semi-minor axis in the ellipse threshold line.

		Rainfall threshold (T)				
		T _{90%}	T _{60%}	T _{30%}	T _{15%}	
ity	High	Extreme	Extreme	High	Medium	
tibili	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	
nscep	Moderate	Extreme	High	Medium	Low	
andslide s	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	
	Low	High	Medium	Low	Low	
Τ	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	

		Rainfall threshold (T)				
		T _{90%}	T _{60%}	T _{30%}	T _{15%}	
ity	High	Extreme	Extreme	High	Medium	
tibil	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	
nscep	Moderate	Extreme	High	Medium	Low	
ide s	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	
andsl	Low	High	Medium	Low	Low	
Τ	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level	

		Rainfall threshold (T)				
		T _{90%}	T _{60%}	T _{30%}	T _{15%}	
tibility	High	Extreme	Extreme	High	Medium	
	susceptibility	danger level	danger level	danger level	danger level	
ide suscept	Moderate	Extreme	High	Medium	Low	
	susceptibility	danger level	danger level	danger level	danger level	
Landsli	Low	High	Medium	Low	Low	
	susceptibility	danger level	danger level	danger level	danger level	

Figure 5: Landslide early warning model and the warning signalert considering both landslide susceptibility and rainfall thresholds.







⁵ Figure 6: 57: Area under the curve (AUC) Result of training and validation of landslide susceptibility analysis.

(a) High susceptibility 120 Landslide cases 110 I₃ (3-hour mean rainfall intensity, mm/h) Extreme High 100 Medium Low 90 80 90% 70 60 60% 50 30% 40 15% 30 20 10 0 0 100 200 300 400 500 600 700 800 900 1000 1100 1200 R24 (24-hour accumlated rainfall, mm)

(b) Moderate susceptibility



(c) Low susceptibility





Figure 876: I3-R24 rainfall thresholds and alert of (a) high-susceptibility slope units (b) moderate susceptibility slope units and (c)

low-susceptibility slope units for southern Taiwan--

			Rainfall threshold (T)						
		T _{90%}		Τ _e	T _{60%}		30%	T _{15%}	
		I ₃	R ₂₄	I ₃	R ₂₄	I ₃	R ₂₄	I ₃	R ₂₄
tibility	High	70	750	55	600	40	450	30	300
	susceptibility	(68)	(745)	(56)	(610)	(40)	(438)	(27)	(291)
ide suscep	Moderate	60	650	45	500	35	350	20	250
	susceptibility	(61)	(657)	(46)	(498)	(34)	(368)	(22)	(236)
Landsli	Low	50	550	40	450	30	300	15	200
	susceptibility	(50)	(539)	(40)	(430)	(29)	(316)	(15)	(167)

5

Figure 87: Rainfall thresholds for southern Taiwan. The values were calculated as 90%, 60%, 30%, and 15% of the original threshold respectively. The original values were calculated from 30%, 60% and 90% thresholds respectively. After that, I₃ was rounded to the nearestby 5 mm/h and R₂₄ was rounded by to the nearest 50 mm for operational purpose (e.g., evacuation). The original values are shown in parentheses.



(b) Zhongmin landslide







(c) Houcuo landslide





(c) Houcuo landslide



Figure 998: Disastrous landslide cases in 2016, and their rainfall histogram as well as histograms, and their snake linelines in the I3-R24 diagram. it The results showed that these disastrous landslide cases could be warned with orange or red sign-alerts could have been issued in advance before landslide occurred for these disastrous landslides.



(b) Typhoon Haitang





(b) Typhoon Haitang



model issued warningalerts matched the landslide occurrence timetimes reported by newspapers.

Table 1: List The multi-year List of landslide inventories that generated in this studystudyy.

Year	Event
2004	Before Typhoon Mindulle
2004	After Typhoon Mindulle
2005	After Typhoon Haitang
2006	After 0609
	torrentialt Torrential rainfall
2007	After Typhoon Sepat
2008	After Typhoon Sinlaku
2009	After Typhoon Morakot

5 Table 2: <u>Predisposing Ffactors</u> items and their logistic function coefficient in logistic regression analysis.

Code	Factor item	coefficient
L01	RMSSC I	-
L02	RMSSC II	-
L03	RMSSC III	-0.874
L04	RMSSC IV	-0.099
L05	RMSSC V	0.314
L ₀₆	RMSSC VI	-0.384
L07	RMSSC VII	-
F01	dip slope	0.207
F02	average slope	0.265
F03	variance of slope	0.098
F04	ratio of steep slope	0.344
F05	average curvature	0.016
F06	variance of curvature	0.161
F07	fold density	0.013
F08	average wetness	0.061
F09	<u>3-hour mean rainfall intensity (I3)</u>	-0.817
F_{10}	<u>24-hour accumulated rainfall (R₂₄)</u>	0.665
С	Constant	0.057

Table 3: Type and the proportion of landslide occurrence timetimes.

type of landslide occurrence time	amount (percentage)
Type A: within the 3 hours right after following the highest rainfall	
intensity	218 (23%)
(landslide induced by high rainfall intensity)	
Type B: within the 3 hours right after following the 2 nd or 3 rd highest	
rainfall intensity	242 (26%)
(landslide induced by high rainfall intensity)	
Type C: near the end of the rainfall event	401 (710/)
(landslide induced by high accumulated rainfall)	481 (51%)

 Table 4: Coefficient of variation of different accumulated rainfall indexindexes.

Accumulated rainfall Indexes	Coefficient of Variation
6-hour accumulated rainfall (R6)	0.68
12-hour accumulated rainfall (R ₁₂)	0.47
24-hour accumulated rainfall (R ₂₄)	0.38
48-hour accumulated rainfall (R48)	0.41
72-hour accumulated rainfall (R72)	0.45

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 Table 5: Contingency matrix for the validation of landslide early warning model. FourShown are four outcomes-including: True

 Positive (TP), True Negative (TN), False Positive (FP) and False Negative (FN) are shown.).

		Observed events		
		Yes	No	
	Yes	<u>True Positive</u> (TP)	<u>False Positive</u> (FP)	
rorecasted events	<u>No</u>	<u>False Negative</u> (FN)	<u>True Negative</u> (TN)	

Table 6: Alerts and the corresponding dangerousdanger levels, as well as suggested actions.

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Table 5: Landslide early warning model and the warning sign considering both landslide susceptibility and rainfall thresholds

		Rainfall threshold (T)							
		T_{90%}	$T_{60\%}$ $T_{30\%}$ $T_{15\%}$						
urni. 18-	JLa <u>tLa</u> Jlide cept cept	Dangerous <u>Danger</u> Extreme	Buggested action Extreme						
₹ M	sig nds sus	dangerous-level	dangerous-level						
	Moderate	Extreme	High	Medium	Low				
	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level				
	Low	High	Medium	Low	Low				
	susceptibility	dangerous level	dangerous level	dangerous level	dangerous level				

Table 6: Rainfall thresholds for southern Taiwan. I3 was rounded to 5 mm/hr, R24 was rounded to 50 mm and parentheses referredto the original value.

Rainfall threshold (T)					
T_{90%}	Ŧ _{60%}	T_{30%}	$T_{15\%}$		

		I ₃	R ₂₄	I_3	R ₂₄	H3	R ₂₄	I 3	R ₂₄
	High	70	750	55	600	40	450	30	300
լ ≵։	susceptibility	(68)	(745)	(56)	(610)	(40)	(438)	(27)	(291)
slid. Hibili	Moderate	60	650	4 5	500	35	350	20	250
and	susceptibility	(61)	(657)	(46)	(498)	(34)	(368)	(22)	(236)
Ţ ¥	Low	50	550	40	4 50	30	300	15	200
	susceptibility	(50)	(539)	(40)	(430)	(29)	(316)	(15)	(167)

Table 76: Warning signs and the corresponding dangerous levels as well as suggested actions.

5

Warning sign	Dangerous level	Suggested action
Alert	Danger level	Suggested action
Green	Low	-
Yellow	Medium	Notice announcements
Orange	High	Evacuation
Red	Extreme	Forced evacuation

 Table 87: Disastrous landslide cases in 2016 and the results of validation-results. These-Warnings could have been issued for all landslide cases could all be warned-in advance or just onat the time of occurrence.

Landslide susceptibility Sh	Lithology ihwen landslide (S	Landslide area hihwen villia	Warning sign & warning <u>Alert</u> <u>& issuing</u> time age, Chunri Township	Occurrence of Landslide , Pingtung County)	Early (+) Late (-)		
Low	Weathered sandstone	61,500 m ²	Orange, firstly at-11:00, <u>14-</u> <u>Sep</u> 14 th September, 2016	05:00, <u>15-Sep</u> 15 th September , 2016	+18 hours		
Zhongmin landslide (Zhongmin Rd., Yanchao District, Kaohsiung City)							
High	Mudstone interbedded with thin sandstone	3,500 m ²	Orange and red, 04:00, 28th <u>28-SepSeptember</u> , 2016	10:00, <u>28-Sep</u> 28th September , 2016	+ 6 hours		
Houcuo landslide (Houcuo Ln., Qishan District, Kaohsiung City)							
Low	Conglomerate	4,000 m ²	Orange, 03:00	03:30, <u>28-Sep</u> 28 th September, 2016	0 hour		

 Table 98: Validation of warnings issued for slope units and the hitting ratio hit rate ios of during Typhoon Typhoons Mindulle and Haitang.

			Number of slope units	
Truch a on arrest	Landslide occurrence	Number of landslide slope units	ever had be warned	Hitting
i yphoon event	time reported by	(Number of high <u>;-</u> , moderate; <u>-</u> ,	as for which orange	Detiellit meteie
(year)	newspapers	lowsusceptibility slope units)	signalerts had	Kationi raicio
			beenwere issued	
Mindulle (2004)	Mainly 06:00	10.011		
	13:00,	10,911	8,283	75.9%
	<u>3</u> 3 rd , Jul <u>, 2004</u> y	(3,129; 2,730; 3,032)		
Haitang (2005)	From 05:00, <u>19-Jul</u> 19 th	10.904		
	to 06:00, <u>20-Jul</u> 20 th ,	10,804	10,245	94.8%
	July, 2005	(2,392; 2,333; 3,837)		

<u>Table 9: Validation of warningsalerts</u> issued for villages during Typhoons Mindulle and Haitang by using contingency matrix and skill scores.

	Observed events						
_	<u>Typho</u>	on Mindulle (ir	<u>n 2004)</u>	Typhoon Haitang (in 2005)			
_	<u>Yes</u> <u>No</u>				Yes	<u>No</u>	
Forecasted	Yes	220	<u>30</u>	Yes	<u>194</u>	<u>24</u>	
events	<u>No</u>	<u>9</u>	<u>77</u>	<u>No</u>	<u>28</u>	<u>12</u>	
POD		<u>0.961</u>			0.874		
POFD	0.280			0.667			
POFA	<u>0.120</u> <u>0.110</u>						