Dear Samuele Segoni,

Thank you very much for everything you have done for us about our manuscript entitled “**Rainfall threshold calculation for debris flow early warning in areas with scarcity of data**” (No.nhess-2017-333). We truly appreciate all of the thoughtful comments from you and the Referee, and we have now revised our manuscript accordingly with a list of changes detailed below.

Response to Referee 1#:

In general, all the arisen questions have been fulfilled. On my view, the choice of proxies could be not fully agreeable. About item 18), I confirm that, on my view, rainfall intensity should be reported on x-axis in terms of cumulative values over time unit (also if for 1 hour time reference, the values are the same). Moreover the readability of x-axis in Graphs forming Figure 13 could be improved; for example, reporting only hours and putting date as x-label.

Thanks a lot for your suggestions and advices, and we fully considered your advices and redrew the figures in Figure 13.

Response to Editor:

1) *I agree with the reviewer that the methodology has some major weaknesses. The choice of calculating API using 20 days is subjective and arbitrary. Why 20 and not 21, 30, 15, 10 or 5? Moreover, in several parts of the manuscript you state that in the literature other parameters are more frequently and conveniently used for debris flow modelling. Indeed, an overwhelming literature demonstrated that in similar cases of study I-D thresholds are more appropriate than thresholds accounting for antecedent rainfall conditions.*

Thank you very much for your kindly advices on this problem of this manuscript. As you suggested, firstly, we made 4 versions of the counting days when calculate API, 3days, 10days, 20days and 30days. The comparison among all the versions is shown in Table 6 (Line 469). It indicates that the value of the effective antecedent precipitations (P_{a0}) were increasing from 3 days to 20 days, while with the time last to 30 days, the value of P_{a0} was barely changed anymore. Therefore, it can be considered that the effect of a rainfall event usually diminished in 20 days. Hence, the numbers of previous indirect rainfall days (n) is identified as 20. At the same time, we analyzed the trend of the trigger rainfall of debris flow events in Guojuanyan gully, and made a debris flow triggering thresholds for the gully (Line 485, Figure 14). The comparison result shows that the laws of the two threshold curve are the same and validates that the calculated method of rainfall threshold proposed in this work is reasonable. We also discussed this in the Discussions part of the manuscript(Line 488-497).

2) *Fig. 7 and lines 234-260. This part is troublesome. You actually say that some kind of data are better than the ones you are going to use, then you state that you will use the worst data. You need to change this part because it introduces a major weakness in your research. Moreover, I do not agree with your conclusions. Fig.7 clearly shows that DF initiation cannot be characterized us-*

ing only rainfall intensity, even if different time spans are taken into account (10', 1h and 1d intensity). You can get to this conclusion by observing that many DF are below the annual average values. This encourages you to use antecedent rainfall… HOWEVER, THESE PLOTS WOULD BE MORE INFORMATIVE IF YOU PLOT DURATION ON THE X AXYS, to get I-D thresholds as mentioned in the former comment (1b).

We rewrote the whole paragraph about figure 7 to try to make it concise and clearly. As you suggested, the main conclusion of Fig.7 is that the trigger rainfall of the debris flow events had decreased obviously after the earthquake.

3) The section 4.2.1 should be moved into the materials and methods section. API should be explained the first time you mention it.

This part has been moved to the materials and methods part as section 3.2 (Line 301-339).

4) In the discussion section, you should clearly state that the threshold cannot be applied elsewhere: the proposed approach is based on a procedure that can be exported elsewhere only if a site-specific calibration is needed to develop specific thresholds for other test sites.

This has been highlighted (Line 494-497).

5) Others:

We have checked the whole manuscript thoroughly again. Some additional and important references have also been added.

We wish that with the above revisions made, our manuscript can now be accepted for publication on *Nature Hazards and Earth System Sciences* soon. Please do not hesitate to contact me if you have any additional questions or comments.

Looking forward to hearing from you.

Regards

JIANG Yuanjun

Rainfall threshold calculation for debris flow early warning in areas with scarcity of data

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Abstract: Debris flows are one of the natural disasters that frequently occur in mountain areas, usually accompanied by serious loss of lives and properties. One of the most used approaches to mitigate the risk associated to debris flows is the implementation of early warning systems based on well calibrated rainfall thresholds. However, many mountainous areas have little data regarding rainfall and hazards, especially in debris flow forming regions. Therefore, the traditional statistical analysis method that determines the empirical relationship between rainstorm and debris flow events cannot be effectively used to calculate reliable rainfall thresholds in these areas. After the severe Wenchuan earthquake, there were plenty of diposits deposited in the gullies which resulted in lots of debris flow events subsequently. The triggering rainfall threshold has decreased obviously. To get a reliable and accurate rainfall threshold and improve the accuracy of debris flow early warning, this paper developed a quantitative method, which is suit for debris flow triggering mechanism in meizoseismal areas, to identify rainfall threshold for debris flow early warning in areas with scarcity of data based on the initiation mechanism of hydraulic-driven debris flow. First, we studied the characteristics of the study area, including meteorology, hydrology, topography and physical characteristics of the loose solid materials. Then, the rainfall threshold was calculated by the initiation mechanism of the hydraulic debris flow. The comparison with other models and with alternate configurations demonstrates that the proposed rainfall threshold curve is a function of

the antecedent precipitation index (*API*) and 1-h rainfall. To test the proposed method, we selected the Guojuanyan gully, a typical debris flow valley that during the 2008-2013 period experienced several debris flow events and that is located in the meizoseismal areas of Wenchuan earthquake, as a case study. The comparison with other threshold models and with configurations shows that the selected approach is the most promising to be used as a starting point for further studies on debris flow early warning systems in areas with scarcity of data.

Keywords: Debris flow; rainfall threshold curve; rainfall threshold; areas with scarcity of data

1 Introduction

Debris flow is rapid, gravity-induced mass movement consisting of a mixture of water, sediment, wood and anthropogenic debris that propagate along channels incised on mountain slopes and onto debris fans (Gregoretti et al., 2016). It has been reported in over 70 countries in the world and often causes severe economic losses and human casualties, seriously retarding social and economic development (Imaizumi et al., 2006;Tecca and Genevois, 2009; Dahal et al., 2009; Liu et al., 2010; Cui et al., 2011; McCoy et al., 2012; Degetto et al., 2015; Tiranti and Deangeli, 2015; Hu et al., 2016). Rainfall is one of the main triggering factors of debris flows and is the most active factor when debris flows occur, which also determines the temporal and spatial distribution characteristics of the hazards. As one of the important and effective means of non-engineering disaster mitigation, much attention has been paid to debris flow early warning by researchers (Pan et al., 2013; Guo et al., 2013; Zhou et al., 2014; Wei et al., 2017). For rainstorm triggered debris flows, the precipitation and intensity of rainfall are the decisive factors of debris flow initiation, and a reasonable rainfall threshold target is essential to ensure the accuracy of debris flow early warning. However, if there are some extreme events occurred, such as an earthquake, the rainfall threshold of debris flow may change a lot. Tang et al. (2012) analyzed the critical rainfall of Beichuan city and found that the cumulative rainfall triggering debris flow decreased by 14.8%-22.1% when compared with the pre-earthquake period, and the critical hour rainfall decreased by 25.4%-31.6%. Chen et al. (2013)analyzed the pre- and post-earthquake critical rainfall for debris flow of Xiaogangjian gully and found that the critical rainfall for debris flow in 2011 was approximately 23% lower than the value during the pre-earthquake period. Other researches, such as Chen et al. (2008) and Shied et al. (2009) has reached similar conclusions that the post-earthquake critical rainfall for debris flow is markedly lower than that of the pre-earthquake period. The Guojuanyang gully, a small gully located in the meizoseismal areas of the big earthquake, has no debris flows under the annual average rainfall before 2008, but it became a debris flow gully after the earthquake under the same conditions, even the rainfall was smaller than the annual average rainfall. These indicated that earthquakes have a big influence on debris flow occurrence. The earthquake triggered many unstable slopes, collapses, and landslides, which have served as the source material for debris flow and shallow landslide in the years after the earthquake (Tang et al. 2009, 2012; Xu et al. 2012; Hu et al. 2014). Therefore, the rainfall threshold of debris flow post-earthquake is an important and urgent issue to study for debris flow early warning and mitigation.

As an important and effective means of disaster mitigation, debris flow early warning have received much attention from researchers. The rainfall threshold is the core of the debris flow early warning , on which have a great deal of researches yet (Cannon et al., 2008; Chen and Huang 2010; Baum and Godt, 2010;Staley et al., 2013; Winter et al., 2013; Zhou and Tang, 2014; Segoni et al., 2015; Rosi et al 2015). Although the formation mechanism of debris flow has been extensively studied, it is difficult to perform distributed physically based modeling over large areas, mainly because the spatial variability of geotechnical parameters is very difficult to assess (Tofani et al., 2017). Therefore, many researchers (Wilson and Joyko, 1997; Campbell, 1975; Cheng et al., 1998) have had to determine the empirical relationship between rainfall and debris flow events and to determine the rainfall threshold depending on the combinations of rainfall parameters, such as antecedent rainfall, rainfall intensity, cumulative rainfall, et al.. Takahashi (1978), Iverson (1989)and Cui (1991) predicted the formation of debris flow based on studies ofslope stability, hydrodynamic action and the influence of pore water pressure on the formation process of debris flow. Caine (1980) first statistically analyzed the empirical relationship between rainfall intensity and the duration of debris flows and shallow landslides and proposed an exponential expression($I = 14.82D^{-0.39}$). Afterwards, other researchers, such as Wieczorek (1987), Jison (1989), Hong et al. (2005), Dahal and Hasegawa (2008), Guzzetti et al. (2008) and Saito et al. (2010), carried out further research on the empirical relationship between rainfall intensity and the duration of debris flows, established the empirical expression of rainfall intensity - duration ($I = D$) and proposed debris flow prediction models. Although I-D is the most used approach, other rainfall parameters have been considered as well for debris flow thresholds. Shied and Chen (1995) established the critical condition of debris flow based on the relationship between cumulative rainfall and rainfall intensity. Zhang (2014) developed a model for debris flow forecasting based on the water-soil coupling mechanism at the watershed scale. In addition, some researchers have highlighted the importance to find more robust hydrological bases to empirical rainfall thresholds for landslide initiation (Bogaard et al., 2018; Canli et al., in review; Segoni et al., 2018). When data are scarce, a robust validation of a threshold model can be based on a quantitative comparison with alternate versions of the threshold (Althuwaynee et al.,2015) or with thresholds calculated with completely different approaches (Frattini et al., 2009; Lagomarsino et al., 2015). Zhenlei Wei et al. (2017) investigated a rainfall threshold method for predicting the initiation of channelized debris flows in a small catchment, using field measurements of rainfall and runoff data.

Overall, the studies on the rainfall threshold of debris flow can be summarized as two methods: the demonstration method and the frequency calculated method. The demonstration method employs statistical analysis of rainfall and debris flow data to study the relationship between rainfall and debris flow events and to obtain the rainfall threshold curve (Bai et al., 2008; Tian et al., 2008; Zhuang, et al., 2009). The I-D approaches would be this kind of method. This method is relatively accurate, but it needs very rich, long-term rainfall database and disaster information; therefore, it can be applied only to areas with a history of long-term observations. The frequency calculated method, assuming that debris flow and torrential rain have the same frequency, and thus, debris flow rainfall threshold can be calculated based on the rainstorm frequency in the mountain towns where have abundant rainfall data but lack of disaster data (Yao, 1988; Liang and Yao, 2008). Researchers have also analyzed the relationship between debris flow occurrences and precipitation and soil moisture content based on initial debris flow conditions (Hu and Wang, 2003). However, this approach is rarely applied to the determination of debris flow rainfall thresholds because it needs series of rainfall data. Pan et al. (2013) calculated the threshold rainfall for debris flow pre-warning by calculating the critical depth of debrisflow initiation combined with the amount and regulating factors of runoff generation.

Most mountainous areas have little data regarding rainfall and hazards, especially in Western China. Neither the traditional demonstration method nor frequency calculated method can satisfy the debris flow early warning requirements in these areas. Therefore, how to calculate the rainfall threshold in these data-poor areas has become one of the most important challenges for the debris flow early warning systems. To solve this problem, this paper developed a quantitative method of calculating rainfall threshold for debris flow early warning in areas with scarcity of data based on the initiation mechanism of hydraulic-driven debris flows.

2 Study site

2.1 Location and gully characteristics of the study area

The Guojuanyan gully in Du Jiangyan city, located in the meizoseismal areas of the Wenchuan earthquake, China, was selected as the study area (Fig. 1). It is located at the Baisha River, which is the first tributary of the Minjiang River. The seismic intensity of the study area was XI, which was the maximum seismic intensity of the Wenchuan earthquake. The Shenxi Gully Earthquake Site Park is at the right side of this gully. The area extends from 31°05′27″ N to 31°05′46″ N latitude and 103°36′58″ E to 103°37′09″ E longitude, covering an area of 0.15 km^2 with a population of 20 inhabitants. The elevation range is from 943 m to 1222 m, the average gradient of the main channel is 270‰ (the average slope angle is 15.1°), and the length of the main channel is approximately 580m.

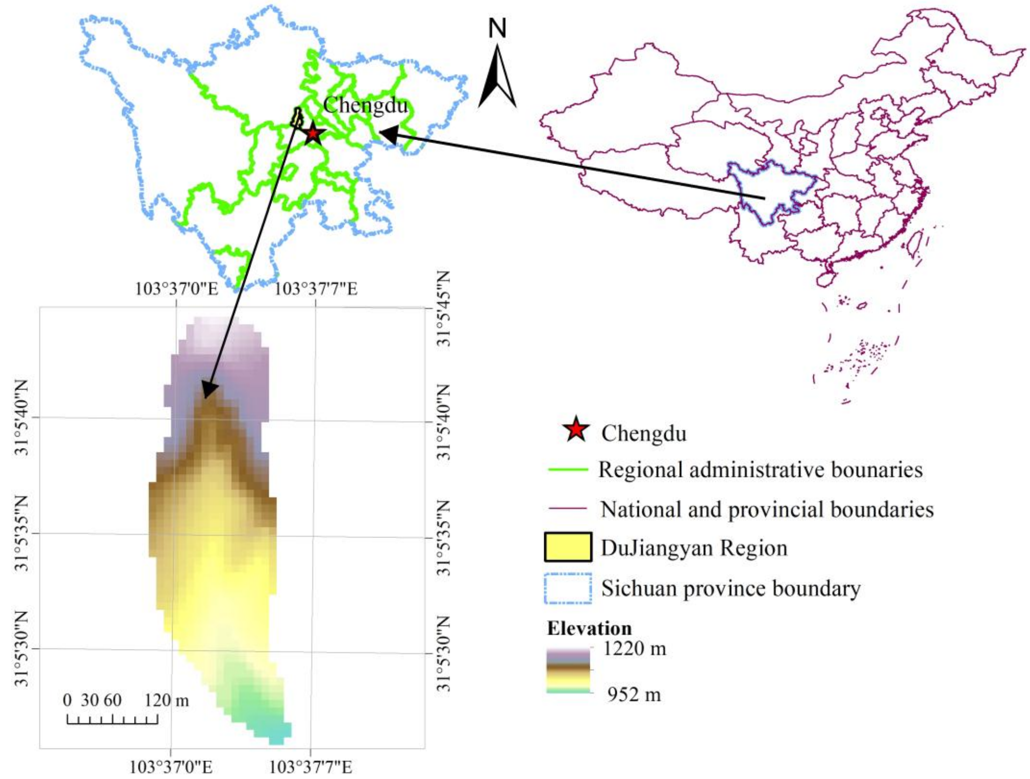


Figure 1. The location of the Guojuanyan gully

Geologically, the Guojuanyan gully is composed of bedrock and Quaternary strata. The bedrock is upper Triassic Xujiahe petrofabric (T_3x) whose lithology is mainly sandstone; mudstone; carbonaceous shale belonging to layered, massive structures; and semi solid-solid petrofabric. The Quaternary strata are alluvium (Q_4^{el+pl}), alluvial materials (Q_4^{pl+dl}), landslide accumulations and debris flow deposits ($Q_4^{sef+del}$). The thickness of the Quaternary strata ranges from 1 m to 20 m and varies greatly. The strata profile of the Guojuanyan gully is shown in Fig. 2.

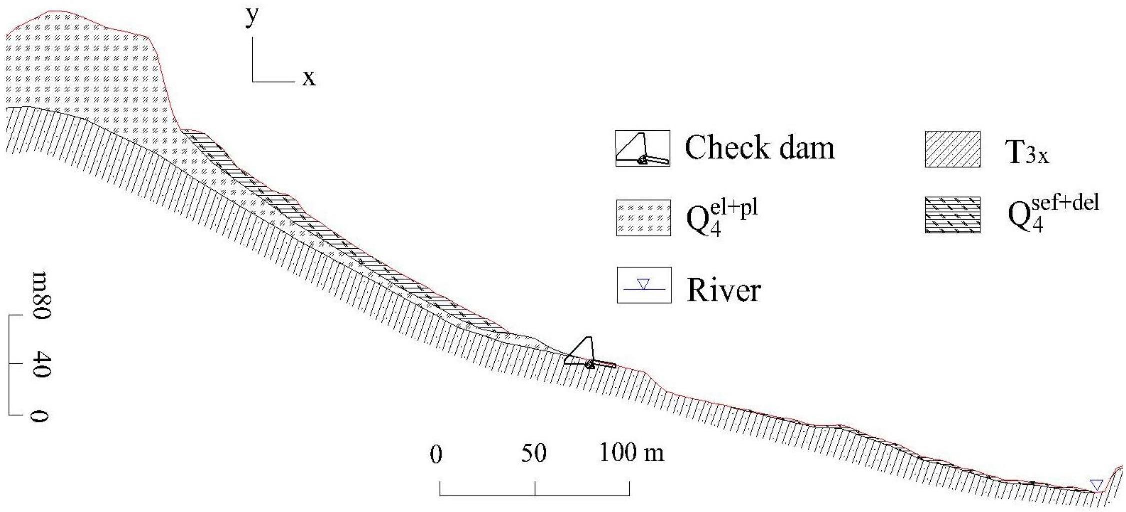


Figure 2. The strata profile of the Guojuanyan gully (Jun Wang et al, 2017)

Geographically, the study area belongs to the Longmenshan Mountains. The famous Longmenshan tectonic belt has a significant effect on this region, especially the Hongkou-Yinxiu fault. The study area has strong tectonic movement and strong erosion, and the main channel is “V”-shaped. The area is characterized by a rugged topography, and the main slope gradient interval of the gully is 20° to 40°, accounting for 52.38% of the entire study area.

Climatically, this area has a subtropical and humid climate, with an average annual temperature of 15.2°C and an average annual rainfall of 1200 mm (Wang et al., 2014).

2.2 Materials and debris flow characteristics of the study area

The Wenchuan earthquake generated a landslide in the Guojuanyan gully, leading to an abundance of loose deposits that have served as the source materials for debris flows. A comparison of the Guojuanyan gully before and after the Wenchuan earthquake is shown in Fig. 3. According to the field investigation and field tests, the landslide 3D characteristics induced by the earthquake and the infiltration characteristics of the loose materials are shown in Table 1 and Table 2 (Wang et al., 2016). They indicate that the volume of materials is more than 20×10^4 m^3,and the infiltration capable of the earth surface have much increased. Therefore, the trigger rainfall for debris flow has decreased greatly. The Guojuanyan gully had no debris flows before the earthquake because of the lack of loose solid materials before the earthquake; however, it became a debris flow gully after the earthquake, and debris flows occurred in the following years (Table 3). The specific conditions of these debris flow events were collected through field investigations and interviews. The field investigations and experiments determined that the density of the debris flow was between 1.8 and 2.1 g/cm^3. Unfortunately, there were no rainfall data before 2011, when we started field surveys in the Guojuanyan gully.

(a) 14 September, 2006 (b) 28 June, 2008

Figure 3. The Guojuanyan gully before (a) and after the Wenchuan earthquake (b) (from Google Earth)

Table 1. The landslide 3D characteristics induced by the earthquake in the study area

Average length /m	Average width /m	Average Height /m	Average depth /m	Slope /°	Volume /×10^{4}m^3
160	80	180	15	≧30	20

Table 2. The infiltration characteristics of solid materials in the study area

Infiltration curve	Infiltration rate	
	Initial infiltration /cm/min	Stable infiltration /cm/min
f= 0.6529*exp(-0.057*t)	3.52	0.34

Table 3. The specific conditions of debris flow events in the Guojuanyan gully after the earthquake

Time	Volume (10^4 m^3)	Surges	Rainfall data record
24 September, 2008	0.6	1	No
17 July, 2009	0.8	1	No
13 August, 2010	4.0	3	No
17 August, 2010	0.4	1	No
1 July, 2011	0.8	1	Yes
17 August, 2012	0.7	1	Yes
9 July, 2013	0.4	1	Yes
26 July, 2013	2.0	2	Yes
18 July, 2014	1.5	1	Yes

2.3 Debris flow monitoring and streambed survey of the study area

After the Wenchuan earthquake, continuous field surveillance was undertaken in the study area. A debris flow monitoring system was also established in the study area. To identify the debris flow events, this monitoring system recorded stream water depth, precipitation and real-time video of the gully (Fig. 4). The water depth was measured using an ultrasonic level meter, and precipitation was recorded by a self-registering rain gauge. The real-time video was recorded onto a data logger and transmitted to the monitoring center, located in the Institute of Mountain Hazards and Environment, Chinese Academy of Sciences. When a rainstorm or a debris flow event occurs, the realtime data, including rainfall data, video record, and water depth data, can be observed and queried directly in the remote client computer in the monitoring center. Fig. 5 shows images taken from the recorded video. These data can be used to analyze the rainfall or other characteristics, such as the 10-min, 1- and 24-h critical rainfall. The recorded video is usually used to analyse the whole inundated process of debris flow events and to identify debris flow events as well as the data from rainfall, flow depth, and field investigation.

(a)Real-time camera and rain gauge (b) Ultrasonic level meters

Figure 4. Debris flow monitoring system in the study area

Figure 5. Real-time images from video taken during the debris flow movement

2.4 Data collection and the characteristics of rainfall

The Wenchuan earthquake occurred in the Longmenshan tectonic belt, located on the eastern edge of the Tibetan plateau, China, which is one of three rainstorm areas of Sichuan Province (Longmen mountain rainstorm area, Qingyi river rainstorm area and Daba mountain rainstorm area). Heavy rainstorms and extreme rainfall events occur frequently. Because there were few data in the mountain areas, we collected the rainfall data from 1971- 2000 and 2011-2012 (from our own on-site monitoring); the characteristics of the rainfalls are as following:

(1) Abundant precipitation: The average annual precipitation was 1177.3 mm from 1971 to 2000, and the average monthly precipitation is shown in Fig. 6. From 1971 to 2000, the minimum annual precipitation of 713.5 mm occurred in 1974, and the maximum annual precipitation of 1605.4 mm occurred in 1978. The total precipitation in 2012 is 1148mm, in the trend range of the historical data.

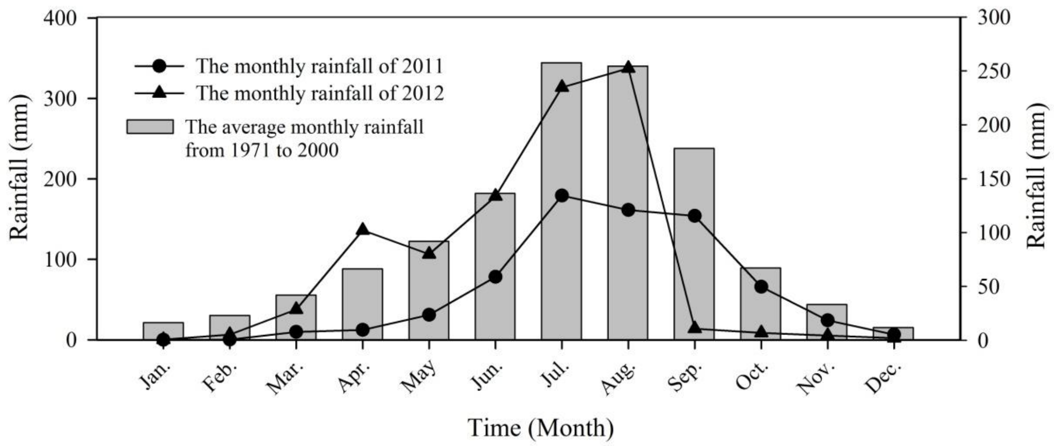


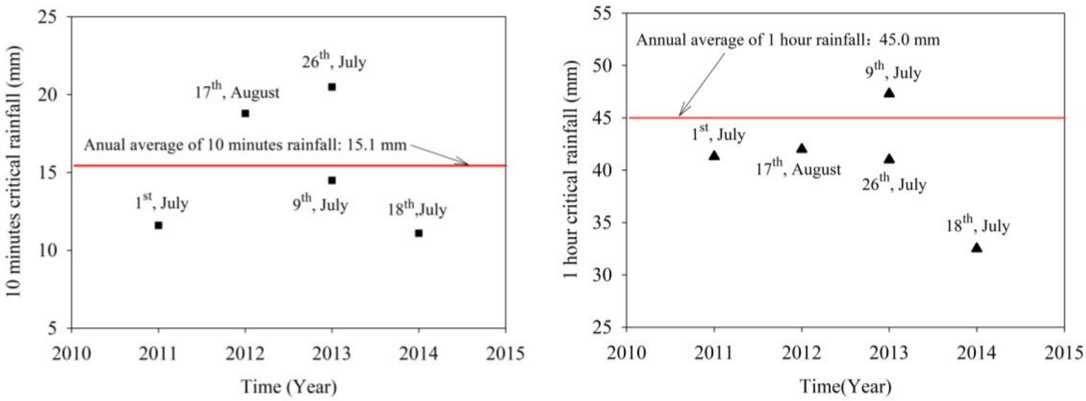
Figure 6. The average monthly precipitation of the Guojuanyan gully from 1971 to 2000 and the monthly rainfall of 2011 and 2012

(2) Seasonality of the distribution of precipitation: from Fig. 6 we can observe that rainfall is seasonal, with approximately 80% of the total rainfall occurring during the monsoon season (from June to September) and the other 20% in other seasons. And the laws of monthly rainfall in 2011 and 2012 coincide to the historical data. For instance, in 2012, the

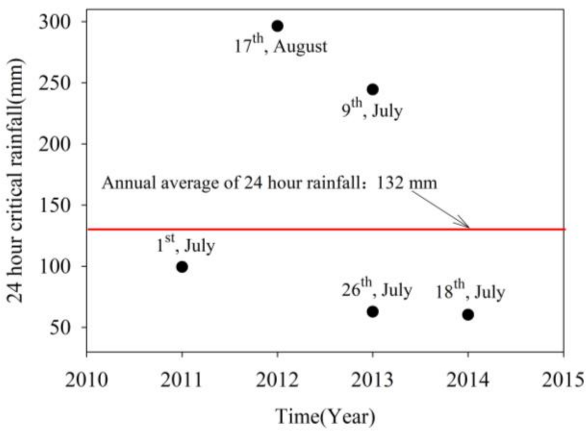
total annual rainfall in this area was approximately 1148 mm, and rainfall in the monsoon season from June to September was 961 mm, accounting for 83.7% of the annual total.

(3) The rainfall intensity has great differences. From 1971 to 2000, the maximum monthly rainfall was 592.9 mm, the daily maximum rainfall was 233.8 mm, the hourly maximum rainfall was 83.9 mm, the 10 minute maximum rainfall was 28.3 mm, and the longest continuous rainfall time was 28 days.

Debris flow field monitoring data and on-site investigation data were used to identify the debris flow events and to analyze the characteristics of the rainfall pattern and the critical rainfall characteristics. Analyzing the typical rainfall process curves (Fig. 13), we can find that the hourly rainfall pattern of the Guojuanyang gully is the peak pattern, displaying the single peak and multi-peak, a characteristic of short-duration rainstorms. Through the statistical analysis of the 10-min, 1-, and 24-h critical rainfall of debris flow events after the earthquake, their characteristics can be obtained, as shown in Fig. 7.



(a) The 10-min critical rainfall (b) The 1-h critical rainfall



(c) The 24-h critical rainfall

Figure 7. The critical rainfall of debris flows in the Guojuanyan gully

According to the Sichuan Hydrology Record Handbook (Sichuan Water and Power Department 1984), during 1940-1975, the annual average of maximum 10-min rainfall of the study area is approximately 15.1 mm, the maximum 1-h rainfall is 45.0 mm and the annual average of maximum 24-h rainfall is 132 mm. Fig. 7 shows that the majority of the debris flow events in 2011-2014 occurred in a rainfall below the annual average values. It convinced that the rainfall threshold of debris flow was decreased obviously after the earthquake.

3 Materials and methods

This study makes an attempt to analyze the trigger rainfall threshold for debris flow by using the initiation mechanism of debris flow. Firstly, to analyze the rainfall characteristics of the watershed by using the field monitoring data; then to calculate the runoff yield and concentration progress based on field observation. Additionally, the critical runoff depth to initiate debris flow was calculated by the initiation mechanism with the underlying surface condition (materials, longitudinal slope, etc.) of the gully. Then, the corresponding rainfall for the initiation of debris was back-calculated based on the stored- full runoff generation. At last, these factors were combined to build the rainfall threshold model. This method can be applied to the early warning system in the areas with scarcity of rainfall data.

The flow chart of the research is shown in Fig. 8.

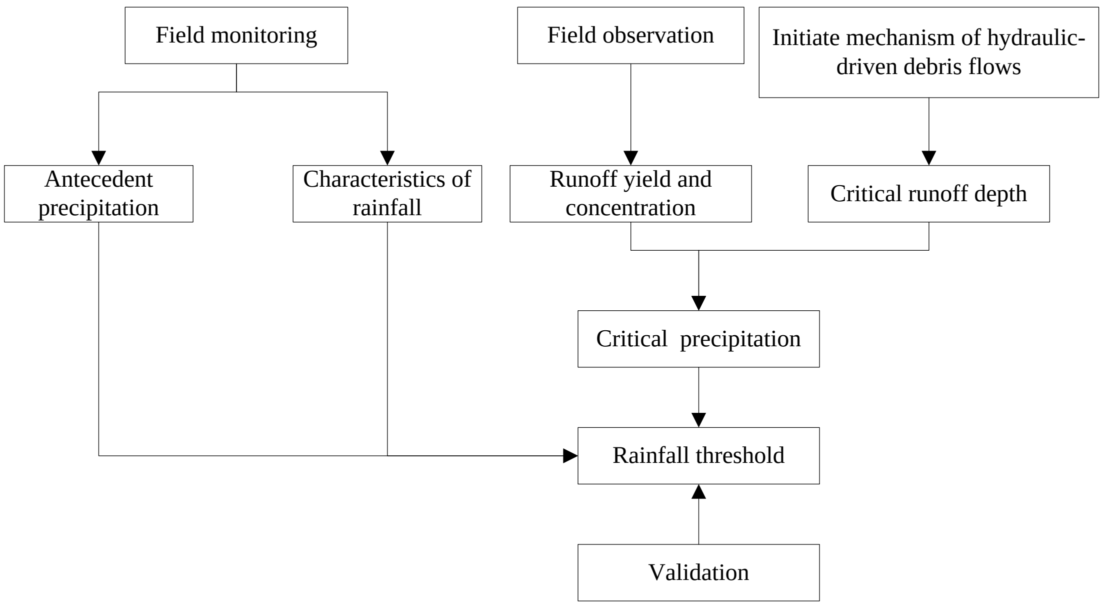


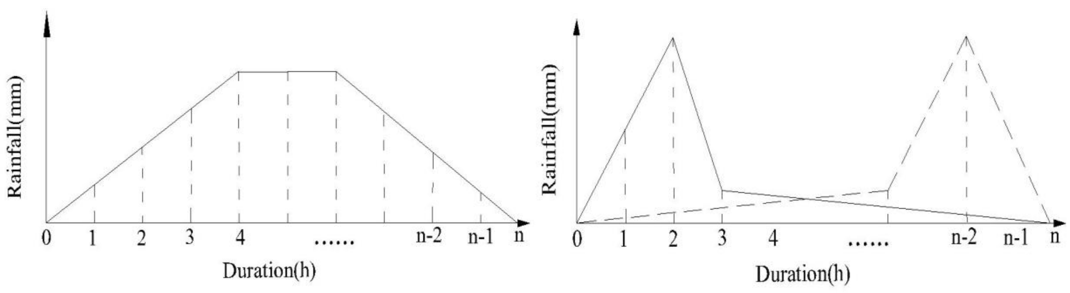
Figure 8. The flow chart of the research

The main influence factors for the formation of debris flow event include three parts: a steep slope of the gully (served as potential energy condition), abundant solid materials (source condition) and water source condition (usually is rainfall condition for rainstorm debris flow). For rainstorm debris flow events, the precipitation and intensity of rainfall are the decisive factors of debris flow initiation. If there is no earthquakes or other extreme events, the topography of the gully can be considered relatively stable. In contrast, rainfall conditions and the distribution of solid materials that determine the occurrence of debris flows can display temporal and spatial variation within the same watershed. Therefore, it is common to provide warning of debris flows based rainfall data after assessing the supply and distribution of loose solid materials. In Takahashi's model, the characteristics of soil, such as the porosity and the hydraulic conductivity of soils, are not considered, and considered the characteristic particle size and the volume concentration of sediment; while the characteristics of topography is mainly represented by the longitudinal slope of the gully. Furthermore, in the stored-full runoff model, the maximum storage capacity of watershed, which mainly decided by the porosity and permeability of the soil, may represent the characteristic of the hydraulic conductivity of solid material to a certain extent. Therefore, this study wouldn't consider the hydraulic conductivity any more.

3.1 Rainfall pattern and the spatial-temporal distribution characteristics

Mountain hazards such as debris flows are closely related to rainfall duration, rainfall amount and rainfall pattern (Liu et al., 2009). Rainfall pattern not only affects the formation of surface runoff but also affects the formation and development of debris flows. Different rainfall patterns result in different soil water contents; thus, the internal structure of the soil, stress conditions, shear resistance, slip resistance and removable thickness can vary. The initiation of a debris flow is the result of both short-duration heavy rains and the antecedent rainfall (Cui et al., 2007; Guo et al., 2013). Many previous observational data have shown that the initiation of a debris flow often appears at a certain time that has a high correlation with the rainfall pattern (Rianna et al., 2014; Mohamad Ayob Mohamadi, 2015).

The precipitation characteristics not only affect the formation of runoff, also affect the formation and development of the debris flow. Different rainfalls result in different soil water contents, and thus the internal structure of the soil, stress conditions, corrosion resistance and slip resistance can vary (Pan et al., 2013). Based on the rainfall characteristics, rainfall patterns can be roughly divided into two kinds, the flat pattern and the peak pattern, as shown in Fig. 9. If the rainfall intensity has little variation, there is no obvious peak in the whole rainfall process; such rainfall can be described as flat pattern rainfall. If the soils characterized by low hydraulic conductivity, this kind of rainfall can hardly trigger a debris flow separately, and the debris flows will mainly be triggered by the great amount of effective antecedent precipitation. While if the rainfall intensity increases suddenly during a certain period of time, the rainfall process will have an obvious peak and is termed peak pattern rainfall. If the hydraulic conductivity is high enough, the rainfall can totally entering the soil and mass can move easily. These debris flows are mainly controlled by the short-duration heavy rains. Peak pattern rainfall may have one peak or multi-peak (Pan, et al., 2013).



(a) Flat pattern rainfall (b) Peak pattern rainfall

Figure 9. The diagram of rainfall patterns

Through analyzing the rainfall data of the Guojuanyan gully, the rainfall pattern and the spatial-temporal distribution characteristics can be obtained.

3.2 The calculation of the antecedent precipitation index (*API*)

The rainfall factor influencing debris flows consists of three parts: indirect antecedent precipitation (IAP) (it is P_{a0} in this paper), direct antecedent precipitation (DAP) (it is R_t in this paper), and triggering precipitation (TP) (it is I_{60} in this paper). The relationships among them are shown in Figure 10. Obviously, IAP increases soil moisture and decreases the soil stability, and DAP saturates soils and thus decrease the critical condition of debris flow occurrence. Although TP is believed to initiate debris flows directly, its contribution amounts to only 37% of total water (Cui et al. 2007). Guo et al (2013) analyzed the rainstorms and debris flow events during June and September in 2006 and 2008, there were 208 days with antecedent rainfall more than 10mm, approximately 57% days of the rain season. Among them, there were 66 days with antecedent rainfall between 10-15mm, and 1 debris flow event happened; 53 days between 15-20 mm and 4 debris flow events happened; 28 days between 20-25 mm and 4 debris flow events happened; 30 days between 25-33 mm and 5 debris flow happened; and 35 days more than 33mm and 9 debris flow events happened. So this group of data can specifically illustrate the importance of the antecedent rainfall to the debris flow events.

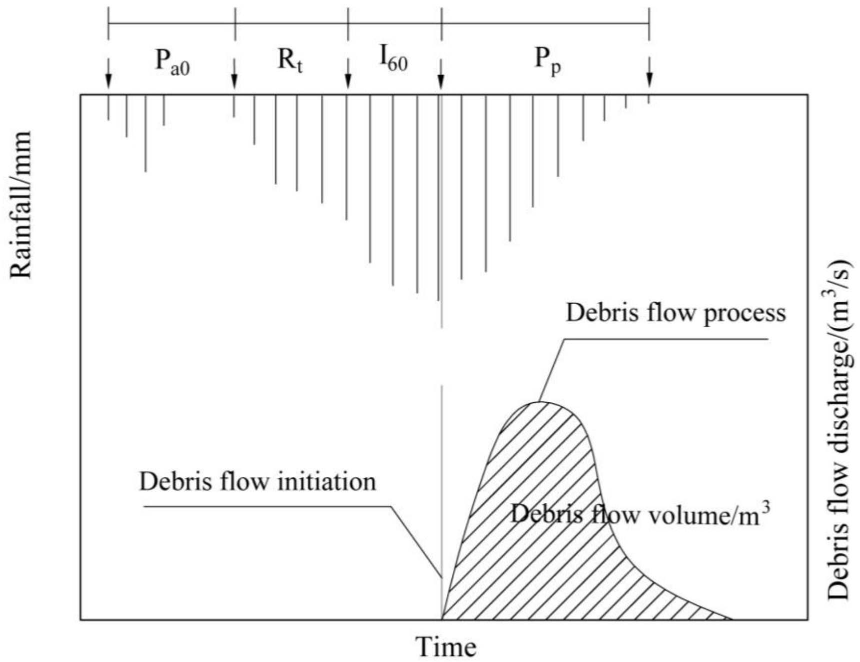


Figure 10. Rainfall index classifications

As Fig. 10 shows, take 1-h rainfall (I_{60}) that obtained from the observed data of the Guojuanyan gully for the TP. The antecedent precipitation index (API) includes IAP and DAP, calculated as the following expression (Zhao, 2011; Guo, 2013; Zhuang, 2015):

$$API = P_{a0} + R_t \qquad (1)$$

where P_{a0} is the effective antecedent precipitation (mm) and R_t is the direct antecedent precipitation (mm), which is the precipitation from the beginning of the rainfall that trigger debris flow to the 1 hour before the debris flow.

It's difficult to study the influence of antecedent rainfall to debris flow as it mainly relies on the heterogeneity of soils (strength and permeability properties), which makes it hard to measure the moisture. Usually, the frequently used method for calculating antecedent daily rainfall is the weighted sum equation as below (Crozier and Eyles 1980; Glade et al. 2000):

$$P_{a0} = \sum_{1}^{n} P_i \cdot K_i \qquad (2)$$

Where P_i is the daily precipitation in the i-th day proceeding to the debris flow event ($1 \le i \le n$) and K_i is a decay coefficient due to evaporation and geomorphological conditions of the soil. The value of the K, is typically 0.8-0.9, can be determined by the test of soil moisture content based on Eq.2 in the watershed. The effect of a rainfall event usually diminishes with the time going forward. Different patterns of storm debris flow gullies require different numbers of previous indirect rainfall days (n), which can be determined by the relationship between the triggering rainfall and the antecedent rainfall of a debris flow (Pan, et al., 2013). If the rainfall is sharp and heavy, the initiation of debris flow would mainly be determined by DAP and TP, while the influence of the antecedent precipitation would be decreased, and vice versa.

3.3 The rainfall threshold curve of debris flows

3.3.1 The initiation mechanism of hydraulic-driven debris flows

When the watershed hydrodynamics, which include the runoff, soil moisture content and the discharge, reach to a certain level, the loose deposits in the channel bed will initiate movement and the sediment concentration of the flow will increase, leading the sediment laden flow to transform into a debris flow. The formation of this kind of debris flow is a completely hydrodynamic process. Therefore, it can be regarded as the initiation problem of debris flow under hydrodynamic force. The forming process of hydraulic-driven debris flows is shown in Fig. 10.

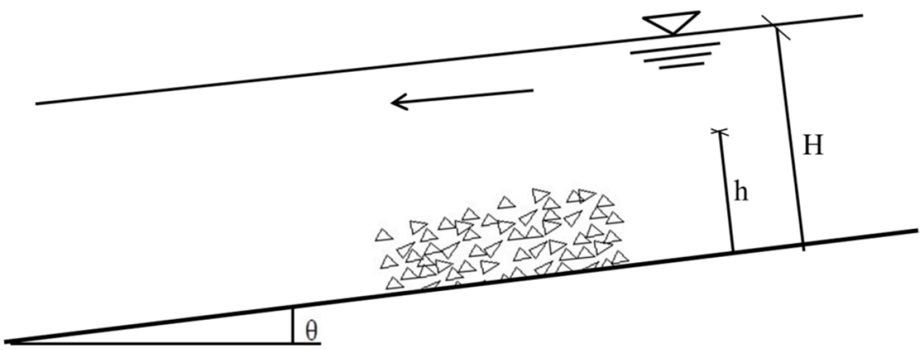


Figure 11. The typical debris flow initiate model

According to Takahashi's model, the critical depth for hydraulic-driven debris flows is:

$$h_0 = \left[\frac{C_*(\sigma - \rho)\tan\phi}{\rho \tan\theta} - \frac{C_*(\sigma-\rho)}{\rho} - 1 \right] d_m \tag{3}$$

where C_* is the volume concentration obtained by experiments(0.812); σ is the unit weight of loose deposits (usually is 2.65 g/cm3); ρ is the unit weight of water,1.0 g/cm3; θ is the channel bed slope (°); ϕ is the internal friction angle (°) and can be measured by shear tests ; And d_m is the average grain diameter (mm), which can be expressed as:

$$d_m = \frac{d_{16} + d_{50} + d_{84}}{3} \tag{4}$$

where d_{16} , d_{50} and d_{84} are characteristic particle sizes of the loose deposits (mm), whose weight percentage are 16%, 50% and 84% separately.

Takahashi's model became one of the most common for the initiation of debris flow after it was presented. A great deal of related studies was published based on Takahashi's model later. Some discussed the laws of debris flow according to the geomorphology and the water content (Sassa et al., 2010; Wang, 2016), while others examined the critical conditions of debris flow with mechanical stability analysis (Cao et al., 2004; Jiang et al., 2017). However, Takahashi's relation was determined for debris flow propagating over a rigid bed, hence, with a minor effect of quasi-static actions near the bed. Lanzoni et al. (2017) slightly modified the Takahashi's formulation of the bulk concentration, which considered the long lasting grain interactions at the boundary between the upper, grain inertial layer and the underlying static sediment bed, and validated the proposed formulation with a wide set of experimental data (Takahashi, 1978, Tsubaki et al., 1983, Lanzoni, 1993, Armanini et al., 2005). The effects of flow rheology on the basis of velocity profiles are analyzed with attention to the role of different stress-generating mechanisms.

This study aims to the initiation of loose solid materials in the gully under surface runoff; the interactions on the boundary are not involved. Therefore, Takahashi's model can be used in this study.

3.3.2 Calculation of watershed runoff yield and concentration

The stored-full runoff, one of the modes of runoff production, is also called as the super storage runoff. The reason of the runoff yeild is that the aeration zone and the saturation zone of the soil are both saturated. In the humid and semi humid areas where rainfall is plentiful, because of the high groundwater level and soil moisture content, when the losses of precipitation meet the plant interception and infiltration, it would not increase anymore with the rains continuous. The Guojuanyan gully is located in Du Jiangyan city, which is in a humid area. Therefore, stored-full runoff can be used to calculate the watershed runoff. That is, it can be supposed that the water storage can reach the maximum storage capacity of the watershed in each heavy rain event. Therefore, the rainfall loss in each time I is the difference between the maximum water storage capacity I_m and the soil moisture content before the rain P_a. The water balance equation of stored-full runoff is expressed as follows (Ye, et al., 1992):

$$R = P - I = P - (I_m - P_a) \tag{5}$$

where R is the runoff depth (mm); P is the precipitation of one rainfall (mm); I is the rainfall loss (mm); I_m is the watershed maximum storage capacity (mm) for a certain watershed, it is a constant for a certain watershed that can be calculated by the infiltration curve or infiltration experiment data. In this study, I_m has been picked up from Handbook of rainstorm and flood in Sichuan (Sichuan Water and Power Department 1984); and P_a is the antecedent precipitation index, referring to the total rainfall prior to the 1 hour peak rainfall leading to debris flow initiation.

Eq. 5 can be expressed as follows:

$$P + P_a = R + I_m \tag{6}$$

The precipitation intensity is a measure of the peak precipitation. At the same time, the duration of the peak precipitation is generally brief, lasting only up to tens of minutes. Therefore, 10-minute precipitation intensity (maximum precipitation over a 10-minute period during the rainfall event) is selected as the triggering rainfall for debris flow, which is appropriate and most representative. However, it is difficult to obtain such short-duration rainfall data in areas with scarcity of data. Therefore, in this study, P and P_a are replaced by I_{60} (1 hour rainfall) and API (the antecedent precipitation index), respectively; thus, Eq. 6 is expressed as:

$$I_{60} + API = R + I_m \tag{7}$$

In the hydrological study, the runoff depth R is:

$$R=\frac{W}{1000F}=\frac{3.6\sum Q\cdot\Delta t}{F}=\frac{3.6Q}{F} \tag{8}$$

where R is the runoff depth (m); W is the total volume of runoff (m^3); F is the watershed area (km^2); Δt is the duration time, in this study it is 1 hour; and Q is the average flow of the watershed (m^3/s), which can be expressed as follows:

$$Q = BVh_0 \tag{9}$$

where B is the width of the channel (m), V is the average velocity (m/s) and h_0 is the critical depth (m).

Eq. 7 is the expression of the rainfall threshold curve for a watershed, which can be used for debris flow early warning. This proposed rainfall threshold curve is a function of the antecedent precipitation index (API) and 1 hour rainfall (I_{60}), which is a line and a negative slope.

4 Results

4.1 The rainfall threshold curve of debris flow

4.1.1 The critical depth of the Guojuanyan gully

The grain grading graph (Fig. 11) is obtained by laboratory grain size analysis experiments for the loose deposits of the Guojuanyan gully. Figure 11 shows that the characteristic particle sizes d_{16}, d_{50}, d_{84} and d_m are 0.18 mm, 1.9 mm, and 10.2 mm, 4.1 mm, respectively. According to Eq. (1), the critical depth (h_0) of the Guojuanyan gully is 7.04 mm.

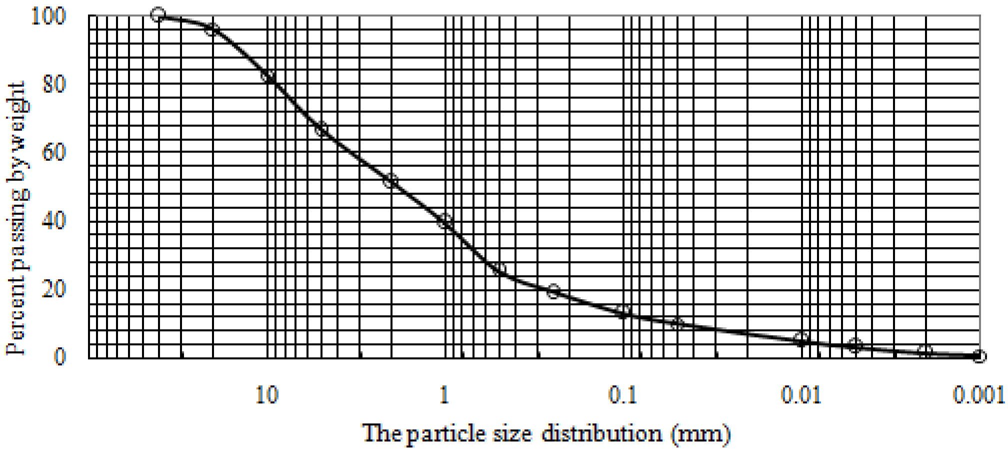


Figure 11. The grain grading graph of the Guojuanyan gully

Table 4. Critical water depth of debris flow triggering in Guojuanyan gully

C_*	σ (g/cm³)	ρ (g/cm³)	$\tan\theta$	d_{16} (mm)	d_{50} (mm)	d_{84} (mm)	d_m (mm)	ϕ (°)	$\tan\phi$	h_0 (mm)
0.812	2.67	1.0	0.333	0.18	1.9	10.2	4.1	21.21	0.388	7.04

4.1.2 The rainfall threshold curve of debris flow

Taking the cross-section at the outlet of the debris flow formation region as the computation object, based on the field investigations and measurements, the width of the cross-section is 20 m, and the average velocity of debris flows which is calculated by the several debris flow events, is 1.5m/s. Based on the Handbook of rainstorm and flood in Sichuan (Sichuan Water and Power Department 1984), the watershed maximum storage capacity (I_m) of the Guojuanyan gully is 100mm. According to Eq. (5) - Eq. (7), the calculated rainfall threshold curve of debris flow in the Guojuanyan gully is shown in Table 5.

Table 5. The calculated process of the rainfall threshold

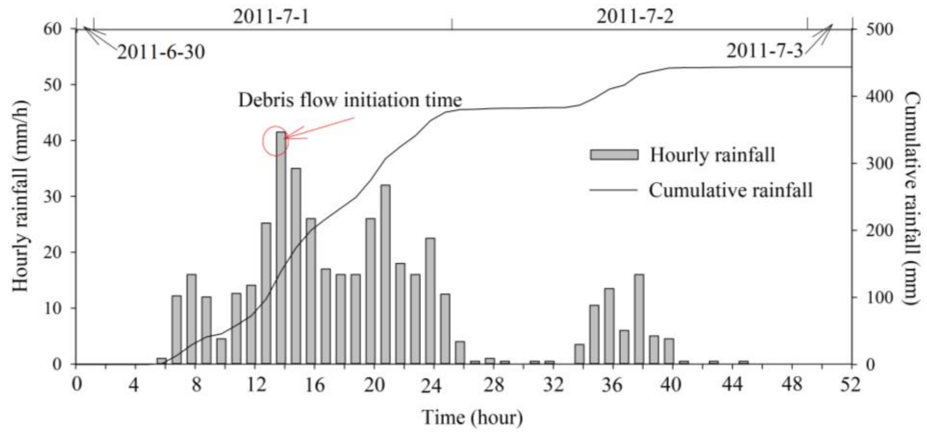
Watershed	h_0 (mm)	B (m)	V (m/s)	Q (m³/s)	Δt (h)	F (km²)	R (mm)	I_m (mm)	$R+I_m$ (mm)
Guojuanyan	7.04	20.0	1.5	0.197	1	0.11	6.9	100	106.9

From the calculated results, we can conclude the rainfall threshold of the debris flow is $I_{60} + API = R + I_m = 106.9 \approx 107$ mm; that is, when the sum of the antecedent precipitation index (API) and the 1 hour rainfall (I_{60}) reaches 107 mm (early warning area), the gully may trigger debris flow.

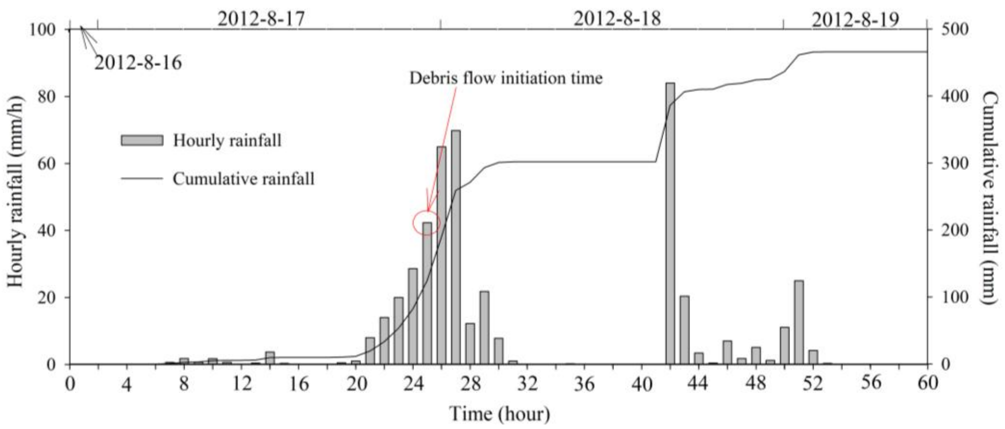
4.2 Validation of the results

4.2.1 The typical debris flow events in the Guojuanyan gully after earthquake

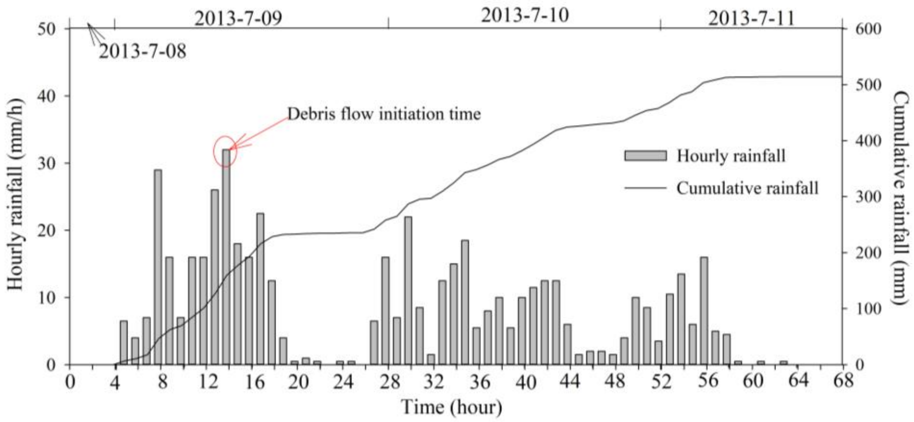
Five typical debris flow events and the corresponding rainfall processes are showed in Figure 13. The debris flow initiation time and the rainfall, both hourly rainfall and cumulative rainfall, have been recorded. From Fig.13, the five debris flows were triggered by torrential rains.



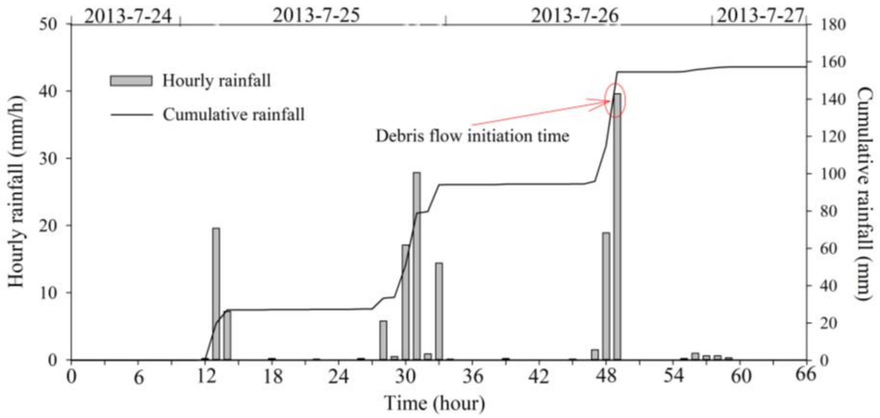
(a)



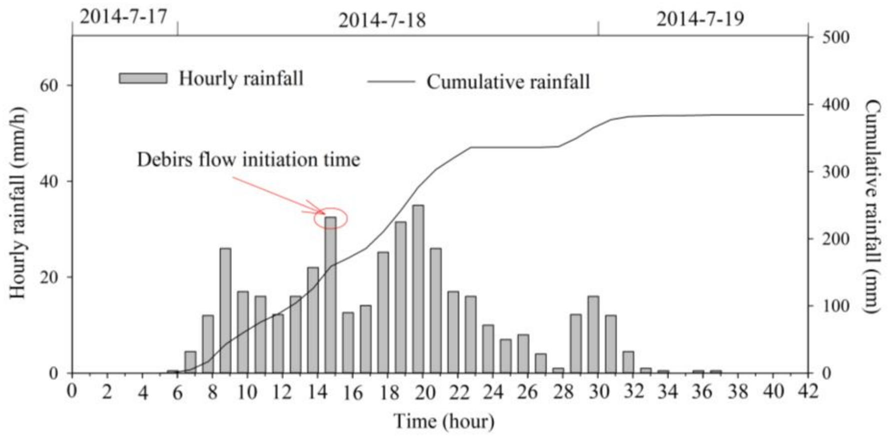
(b)



(c)



(d)



(e)

Figure 13. The rainfall process of debris flow events in the Guojuanyan gully from 2011 to 2014 (a, July 1, 2011; b, August 17, 2012; c, July 9, 2013; d, July 26, 2013; e, July 18, 2014)

4.2.2 The calculation of *API* and 1-h triggering rainfall of the typical rainstorms during 2010-2014

Based on the field tests and experiences, the value of K in Eq.2 is identified as 0.8 (Cui et al. 2007). To determine the numbers of previous indirect rainfall days (n), a comparison among 3 days, 10days, 20days and 30 days were showed in Table 6. It indicates that the value of the effective antecedent precipitations (P_{a0}) were increasing from 3 days to 20 days, while with the time last to 30 days, the value of P_{a0} was barely changed. Therefore, it can be considered that the effect of a rainfall event usually diminished in 20 days. Hence, the numbers of previous indirect rainfall days (n) is identified as 20.

Table 6. The comparisons of P_{a0} when n have different values

Time	Pa_0(mm)			
	n=3	n=10	n=20	n=30
July 1, 2011	3.4	5.2	9.7	9.7
August 17, 2012	2.3	4.7	12.1	12.1
July 9, 2013	0.8	2.5	5.7	5.7
July 26, 2013	6.2	10.8	22.4	22.6
July 18, 2014	0	6.2	10.7	10.7
August 20, 2011	8.3	0	8.5	8.6
September 5, 2011	21.3	45.9	48.7	48.8
June 16, 2012	0	2.7	5.6	5.6
August 3, 2012	5.6	6.1	7.5	7.5
August 18, 2012	10.2	18.4	54.3	54.3
June 18, 2013	0	2.8	6.2	6.2
July 28, 2013	0.2	1.7	13.4	13.5
August 6, 2013	0.2	6.6	12.4	12.4

Thus, the intensity of the 1-h triggering rainfall I_{60} and cumulative rainfall for the typical rainstorms are shown in Table 7. In addition to the rainfall process of the 5 debris flow events (Fig. 13), some typical rainfalls whose daily rainfall were greater than 50 mm but did not trigger a debris flow were also calculated as a contrast; the greatest 1-h rainfall is considered as I_{60} .

Table 7. The data of typical rainfall in the Guojuanyan gully after the earthquake

Time	Daily rainfall (mm)	Pa_0 (mm)	R_t (mm)	*API* (mm)	I_{60} (mm)	*API*+I_{60} (mm)	Location to the threshold line	Triggered debris flow
1 July, 2011		9.7	97.6	107.3	41.5	148.8	Above	Yes
17 August , 2012		12.1	81.9	94.0	42.3	136.3	Above	Yes
9 July , 2013		5.7	127.5	133.2	32	165.2	Above	Yes
26 July , 2013		22.4	96.0	118.4	18.9	137.3	Above	Yes
18 July, 2014		10.7	116.2	126.9	32.5	159.4	Above	Yes
20 August , 2011	82.8	8.5	19.0	27.5	26.8	54.3	Below	No
5 September , 2011	52.1	48.7	1.2	49.9	16.2	66.1	Below	No

16 June , 2012	55.8	5.6	6.6	12.2	27.0	39.2	Below	No
3 August , 2012	148.3	7.5	84.3	91.8	26.7	118.5	Above	No
18 August , 2012	125.7	54.3	0	54.3	65.0	119.3	Above	No
18 June , 2013	50.6	6.2	3.8	10.0	40.0	50.0	Below	No
28 July , 2013	59.4	13.4	30.0	43.4	29.4	72.8	Below	No
6 August , 2013	56.1	12.4	34.0	46.4	17.1	63.5	Below	No

The proposed rainfall threshold curve is shown in Figure 14, in which the red real line defines the threshold relationship. It shows that the calculated values $I_{60} + API$ of debris flow events in the Guojuanyan gully are all above the rainfall threshold curve, while most of the rainstorms that did not trigger debris flow are lay below the curve. Therefore, it indicates that the rainfall threshold curve calculated by this work is reasonable through the validation by rainfall and hazards data of the Guojuanyan gully.

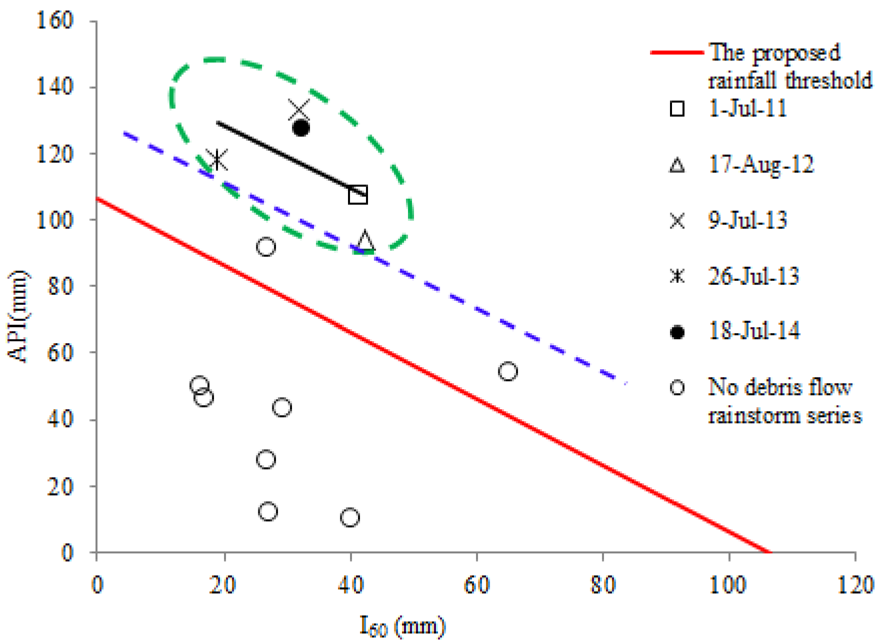


Figure 14. The calculated rainfall threshold curve (red real line), the trend line (black real line) of the debris flow events and the debris flows triggering thresholds (dashed line) in Guojuanyan gully

5 Discussions

The trend of the debris flow events as well as the debris flow thresholds were analyzed in Fig. 14 by using the monitoring rainfall data. A comparison between the thresholds and the calculated threshold curve indicates that they have the same laws. Therefore, the threshold calculated method proposed in this work is reasonable and can be used in the areas with scarcity of data. The proposed rainfall threshold curve is a function of the antecedent precipitation index (*API*) and the 1-h rainfall (I_{60}), which has been validated by rainfall and hazards data. It should be noted that the proposed approach is based on a procedure that can be exported elsewhere only if a site-specific calibration is needed to develop specific thresholds for other test sites. Therefore, the specific value of the threshold should be calculated by the initiation conditions of the debris flow in specific gully.

However, this work still has two limitations. In Figure 14, there are two points above the curve that did not trigger debris flow at all. Although we have highlighted the significance and interconnect of antecedent rainfall, critical rainfall, 1-h triggering rainfall, as well as their accurate determination before the hour of debris flow triggering, it should be noticed that the rainfall is only the triggering factor of debris flows. A comprehensive warning system must contain more environmental factors, such as the geologic and geomorphologic factors, the distribution of material source. In addition, the special and complex formative environment of debris flow after earthquake caused the rainfall threshold is much more complex and uncertain. The rainfall threshold of debris flow is influenced by the antecedent precipitation index (*API*), rainfall characteristics, amount of loose deposits, channel and slope characteristics, and so on. Therefore, we should further study the characteristics of the movable solid materials, the shape of gully, and so on to modify the rainfall threshold curve. But, on the other hand, if given the two rainstorms under the threshold, all the debris flow events points will still locate above the threshold and there will have no missed alarms. Therefore, the threshold established in this work is a conservative one and respect safety.

On the other hand, restricted by the limited rainfall data, this study was validated by only 5 debris flow events. Another limitation of this work is that the approach proposed in this study hasn't been validated by other gullies except the Guojuanyan gully so far. Figure 13 and Figure 14 indicated that the only 5 debris flow events all triggered by the rainfalls with high-intensity and short-duration. In the future, the value of the curve should be further validated and continuously corrected with more rainfall and disaster data in later years.

6 Conclusions

(1) In the Wenchuan earthquake affected areas, loose deposits are widely distributed, causing dramatic changes on the environmental development for the occurrence of debris flow; thus, the debris flow occurrence increased dramatically in the subsequent years. The

characteristics of the 10-min, 1-h and 24-h critical rainfalls were represented based on a comprehensive analysis of limited rainfall and hazards data. The statistical results show that the 10-min and 1-h critical rainfalls of different debris flow events have minor differences; however, the 24 hour critical rainfalls vary widely. The 10-min and 1-h critical rainfalls have a notably higher correlation with debris flow occurrences than to the 24-h critical rainfalls.

(2) The rainfall pattern of the Guojuanyan gully is the peak pattern, both single peak and multi-peak. The antecedent precipitation index (API) was fully explored by the antecedent effective rainfall and triggering rainfall.

(3) As an important and effective means of debris flow early warning and mitigation, the rainfall threshold of debris flow was determined in this paper, and a new method to calculate the rainfall threshold is put forward. Firstly, the rainfall characteristics, hydrological characteristics, and some other topography conditions were analyzed. Then, the critical water depth for the initiation of debris flows is calculated according to the topography conditions and physical characteristics of the loose solid materials. Finally, according to the initiation mechanism of hydraulic-driven debris flow, combined with the runoff yield and concentration laws of the watershed, this study promoted a new method to calculate the debris flow rainfall threshold. At last, the hydrological condition for the initiation of a debris flow is the result of both short-duration heavy rains (I_{60}) and the antecedent precipitation index (API). The proposed approach resolves the problem of debris flow early warning in areas with scarcity data, can be used to establish warning systems of debris flows for similar catchments in areas with scaricty data although it still need further modification. This study provides a new thinking for the debris flow early warning in the mountain areas.

Acknowledgments

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References

Althuwaynee OF, Pradhan B, Ahmad N (2015) Estimation of rainfall threshold and its use in landslide hazard mapping of Kuala Lumpur metropolitan and surrounding areas. Landslides 12:861-875. doi:10.1007/s10346-014-0512-y

Bai LP, Sun JL, Nan Y (2008) Analysis of thecritical rainfall thresholds for mudflow in Beijing, China. Geological Bulletin of China 27(5): 674-680.(in Chinese)

Baum RL, Godt JW (2010). Early warning of rainfall-induced shallow landslides and debris flows in the USA. Landslides, 7(3):259–272.

Bogaard, T. and Greco, R.(2018) Invited perspectives: Hydrological perspectives on precipitation intensity-duration thresholds for landslide initiation: proposing hydro-meteorological thresholds, Nat. Hazards Earth Syst. Sci., 18, 31–39, https://doi.org/10.5194/nhess-18-31-2018

Caine, N (1980) The rainfall intensity-duration control of shallow landslides and debris flows. Physical Geography 62A (1-2):23-27

Campbell RH (1975) Debris Flow Originating from Soil Slipduring Rainstorm in Southern California. Q. Engineering Geologist7: 339–349. DOI:10.1144/GSL.QJEG.1974.007.04.04

Canli, E., Mergili, M., and Glade, T.(2017) Probabilistic landslide ensemble prediction systems: Lessons to be learned from hydrology, Nat. Hazards Earth Syst. Sci. Discuss., https://doi.org/10.5194/nhess-2017-427, in review

Cannon, Susan H., et al.(2008) Storm rainfall conditions for floods and debris flows from recently burned areas in southwestern Colorado and southern California. Geomorphology 96(3): 250-269.

Cao Z., Pender G., Wallis S., Carling P.(2004) Computational dam-break hydraulics over erodible sediment bed. Journal of hydraulic engineering 130.7 (2004): 689-703.

Chen, Su-Chin, and Bo-Tsung Huang (2010) Non-structural mitigation programs for sediment-related disasters after the Chichi Earthquake in Taiwan. Journal of Mountain Science 7(3): 291-300.

Chen YS (2008) An influence of earthquake on the occurrence of landslide and debris flow. Taipei: National Cheng Kung University.

Chen YJ, Yu B, Zhu Y, et al. (2013) Characteristics of critical rainfall of debris flow after earthquake - a case study of the Xiaogangjian gully. Journal of Mountain Science 31(3): 356-361. (in Chinese)

Cheng ZL, Zhu PY, Liu LJ (1998) The Relationshipbetween Debris Flow Activity and Rainfall Intensity.Journal of Natural Disasters7 (1): 118–120. (in Chinese)

Chen NS, Yang CL, Zhou W, et al. (2009)The Critical Rainfall Characteristics for Torrents andDebris Flows in the Wenchuan Earthquake Stricken Area. Journal of Mountain Science 6: 362-372. DOI: 10.1007/s11629-009-1064-9

Cui P (1991) Experiment Research of the Initial Condition andMechanism of Debris Flow. Chinese Science Bulletin 21:1650–1652. (in Chinese)

Cui P, Hu KH, Zhuang JQ Yang Y, Zhang J (2011) Prediction of debris-flow danger area by combining hydro-logical and inundation simulation methods. Journal of Mountain Science 8(1): 1-9. doi: 10.1007/s11629-011-2040-8

Cui P, Zhu YY, Chen J, et al. (2007) Relationships between antecedant rainfall and debris flows in Jiangjia Ravine, China. In: Chen C L and Majir JJ (eds.), Debris flow hazard mitigation mechanics, Prediction , and Assessment. Millpress, Rotterdam: 1-10.

Dahal RK, Hasegawa S, Nonomura A, et al. (2009) Failure characteristics of rainfall-induced shallow landslides in granitic terrains of Shikoku Island of Japan. Environmental geology 56(7): 1295-1310. DOI: 10.1007/s00254-008-1228-x

Degetto M, Gregoretti C, Bernard M (2015) Comparative analysis of the differences between using LiDAR cotour-based DEMs for hydrological modeling of runoff generating debris flows in the Dolomites. Front. Earth Sci. 3, 21. doi: 10.3389/feart.2015.00021

Frattini P, Crosta G, Sosio R (2009) Approaches for defining thresholds and return periods for rainfall - triggered shallow landslides. Hydrol Proc 23(10):1444-1460. doi:10.1002/hyp.7269

Gregoretti C, Degetto M, Boreggio M (2016) GIS-based cell model for simulating debris flow runout on a fan. Journal of Hydrology 534: 326-340. doi: 10.1016/j.jhydrol.2015.12.054

Guido Rianna, Luca Pagano, Gianfranco Urciuoli (2014) Rainfall patterns triggering shallow flowslides in pyroclastic soils. Engineering Geology, 174: 22-35 doi: 10.1016/j.enggeo.2014.03.004

Guo, X.J., Cui, P., Li, Y., 2013. Debris flow warning threshold based on antecedent rainfall: a case study in Jiangjia Ravine, Yunnan, China. J. Mt. Sci. 10 (2), 305–314.

Guzzetti, F., Peruccacci, S., Rossi, M., & Stark, C. P. (2008). The rainfall intensity–duration control of shallow landslides and debris flows: an update. Landslides, 5(1), 3-17.

Hu M J, Wang R (2003) Testing Study of the Correlation among Landslide, Debris Flow and Rainfall in Jiangjia Gully. Chinese Journal of Rock Mechanics and Engineering, 22(5): 824–828 (in Chinese)

Hong Y, Hiura H, Shino K, et al. (2005) The influence of intense rainfall on the activity of large-scale crystalline schist landslides in Shikoku Island, Japan. Landslides 2(2): 97-105. DOI: 10.1007/s10346-004-0043-z

Hu W, Dong XJ, Wang GH, van Asch TWJ, Hicher PY (2016) Initiation processes for run-off generated debris flows in the Wenchuan earthquake area of China. Geomorphology 253: 468–477. doi: 10.1016/j.geomorph.2015.10.024

Iverson RM, Lahusen RG (1989) Dynamic Pore-PressureFluctuations in Rapidly Shearing Granular Materials. Science 246 (4931): 796–799. DOI: 10.1126/science.246.4931.796

Jianqi Zhuang, Peng Cui, Gonghui Wang, et al. (2015) Rainfall thresholds for the occurrence of debris flows in the Jiangjia Gully,Yunnan Province, China.Engineering Geology, 195: 335-346.

Jiang X., Cui P., Chen H., Guo Y. (2016) Formation conditions of outburst debris flow triggered by overtopped natural dam failure. Landslides 14(3):1-11.

Jibson RW (1989) Debris flows in southern Puerto Rico. Geological Society of America Special Papers 236: 29-56. DOI: 10.1130/SPE236-p29

Jun Wang, Shun Yang, Guoqiang Ou, et al. (2017) Debris flow hazards assessment by combing numberical simulation and land utilization. Bulletin of Engineering Geology and the Environment, 1-15. Doi: 10.1007/s10064-017-1006-7

Lagomarsino D, Segoni S, Rosi A, Rossi G, Battistini A, Catani F, Casagli N (2015) Quantitative comparison between two different methodologies to define rainfall thresholds for landslide forecasting. Nat Hazards Earth Syst Sci 15:2413–2423. doi:10.5194/nhess-15-2413-2015

Liang GM, Yao LK (2008) Study on the critical rainfall for debris flows. Lu Ji Gongcheng 6: 3-5. (in Chinese)

Liu YH, Tang C, Li TF, et al. (2009) Statistical relations between geo-hazards and rain type. Journal of Engineering Geology 17(5): 656-661. (in Chinese)

Liu JF, You Y, Chen XZ, Fan JR (2010) Identification of potential sites of debris flows in the upper Min River drainage, following environmental changes caused by the Wenchuan earthquake. Journal of Mountain Science 3: 255-263. doi: 10.1007/s11629-010-2017-z

Lanzoni, S., C. Gregoretti, and L. M. Stancanelli (2017) Coarse-grained debris flow dynamics on erodible beds, J. Geophys. Res. Earth Surf., 122, doi:10.1002/2016JF004046.

McCoy SW, Kean JW, Coe JA, Tucker GE, Staley DM, Wasklewicz WA (2012) Sediment entrainment by debris flows: In situ measurements from the head waters of a steep catchment. J. Geophys. Res.117, F03016. doi: 10.1029/2011JF002278

Mohamad Ayob Mohamadi, Ataollah Kavian (2015) Effects of rainfall patterns on runoff and soil erosion in field plots. International Soil and Water Conservation Research 3: 273-281. http://dx.doi.org/10.1016/j.iswcr.2015.10.001Imaizumi F, Sidle RC, Tsuchiya S, Ohsaka O (2006) Hydrogeomorphic processes in a steep debris flow initiation zone. Geophys. Res. Lett. 33, L10404. doi: 10.1029/2006GL026250

Pan HL, Ou GQ, Hang JC, et al.(2012)Study of rainfall threshold of debris flow forewarning in data lack areas. Rock and Soil Mechanics 33(7): 2122-2126. (in Chinese)

Pan HL, Huang JC, Wang R, et al. (2013) Rainfall Threshold Calculation Method for Debris FlowPre-Warning in Data-Poor Areas. Journal of Earth Science 24(5): 854–862. DOI:10.1007/s12583-013-0377-3

Rosi A, Lagomarsino D, Rossi G, Segoni S, Battistini A, Casagli N (2015) Updating EWS rainfall thresholds for the triggering of landslides. Nature Hazard 78:297–308

Saito H, Nakayama D, Matsuyama H (2010) Relationship between the initiation of a shallow landslide and rainfall intensity—duration thresholds in Japan. Geomorphology 118(1): 167-175. DOI: 10.1016/j.geomorph.2009.12.016

Sassa, K., Nagai, O., Solidum, R., Yamazaki, Y., & Ohta, H. (2010) An integrated model simulating the initiation and motion of earthquake and rain induced rapid landslides and its application to the 2006 Leyte landslide. Landslides, 7(3), 219-236.

Segoni S, Battistini A, Rossi G, Rosi A, Lagomarsino D, Catani F, Moretti S, Casagli N (2015) Technical note: an operational landslide early warning system at regional scale based on space–time variable rainfall thresholds. Nat Hazards Earth Syst Sci 15: 853–861

Segoni S, Rosi A, Lagomarsino D, Fanti R, and Casagli N (2018) Brief Communication: Using averaged soil moisture estimates to improve the performances of a regional-scale landslide early warning system. Nat. Hazards Earth Syst. Sci.

Shied CL, Chen LZ (1995) Developing the critical line of debris –flow occurrence. Journal of Chinese Soil and Water Conservation 26(3):167-172. (in Chinese)

Shieh CL, Chen YS, Tsai YJ, et al (2009) Variability in rainfall threshold for debris flow after the Chi-Chi earthquake in central Taiwan, China. International Journal of Sediment Research 24(2): 177-188.

Staley, D.M., Kean, J.W., Cannon, S.C., Schmidt, K.M., Laber, J.L. (2013) Objective definition of rainfall intensity–duration thresholds for the initiation of post-fire debris flows in southern California, Landslides 10, 547–562

Takahashi T (1978) Mechanical Characteristics of Debris Flow. Journal of the Hydraulics Division 104:1153–1169

Tang C, Zhu J, Li WL (2009) Rainfall-triggered debris flows following the Wenchuan earthquake. Bull Eng Geol Environ 68(2):187–194. DOI: 10.1007/s10064-009-0201-6

Tang C, Van Asch TWJ, Chang M, et al.(2012)Catastrophic debris flows on 13 August 2010 in the Qingping area, southwestern China: the combined effects of as trong earthquake and subsequent rainstorms. Geomorphology139–140:559–576. DOI: 10.1016/j.geomorph.2011.12.021

Tang C, Zhu J, Chang M, et al. (2012) An empirical–statistical model for predicting debris-flow runout zones in the Wenchuan earthquake area. Quaternary International 250:63–73. DOI:10.1016/j.quaint.2010.11.020.

Tecca PR, Genevois R (2009) Field observations of the June 30, 2001 debris flow at Acquabona (Dolomites, Italy). Landslides 6(1): 39-45. doi: 10.1007/s10346-009-0145-8

Tian B, Wang YY, Hong Y (2008) Weighted relation between antecedent rainfall and processprecipitation in debris flow prediction—A case study ofJiangjia gully in Yunnan province. Bulletin of Soil and Water Conservation28(2): 71-75.(in Chinese)

Tiranti D, Deangeli C (2015) Modeling of debris flow depositional patterns according to the catchment and sediment source area characteristics. Front. Earth Sci. 3, 8. doi: 10.3389/feart.2015.00008

Tofani et al., Soil characterization for shallow landslides modeling: a case study in the Northern Apennines (Central Italy). 2017. Landslides 14:755–770, DOI 10.1007/s10346-017-0809-8

Y.Zhao, F. Wei, H.Yang, et al. (2011) Discussion on Using Antecedent Precipitation Index to Supplement Relative Soil Moisture Data Series. Procedia Environment Sciences 10: 1489-1495.

Wang EC, Meng QR (2009) Mesozoic and cenozoic tectonic evolution of the Longmenshan fault belt. Science in China Series D: Earth Sciences 52(5): 579-592. DOI:10.1007/s11430-009-0053-8

Wang G., Furuya G., Zhang F., Doi I., Watanabe N. (2016) Layered internal structure and breaching risk assessment of the Higashi-Takezawa landslide dam in Niigata, Japan. Geomorphology, 267:48-58.

Wang J, Ou GQ, Yang S, Lu GH, et al. (2013) Applicability of geomorphic information entropy in the post-earthuake debris flow risk assessment. Jounal of Mountain Science 31(1): 83-91. (in Chinese)

Wang J, Yu Y, Yang S, et al.(2014) A Modified Certainty Coefficient Method (M-CF) for Debris Flow Susceptibility Assessment: A Case Study for the Wenchuan Earthquake Meizoseismal Areas. Journal of Mountain Science11(5): 1286-1297. DOI: 10.1007/s11629-013-2781-7.

Wang J, Yu Y, Ou GQ, et al.(2016) Study on the Geotechnical Mechanical Characteristics of Loose Materials in the Wenchuan Earthquake-hit Areas. Science Technology and Engineering 16(5): 11-18. (in Chinese)

Wieczorek GF (1987) Effect of rainfall intensity and during in debris flows in central Santa Cruz Mountain. California. Engineering Geology 7: 93-104. DOI: 10.1130/REG7-p93

Wilson, RC, Jayko AS (1997) Preliminary Maps ShowingRainfall Thresholds for Debris-Flow Activity, San Franciscoby Region, California. U.S. Geological SurveyOpen-File Report 97-745 F

Winter, M. G., et al. (2010) Debris flow, rainfall and climate change in Scotland. Quarterly Journal of Engineering Geology and Hydrogeology 43(4): 429-446.

Xu ZQ, Ji SC, Li HB, et al. (2008)Uplift of the Longmen Shan range and the Wenchuan earthquake. Episodes 31(3): 291-301

Xu Q, Zhang S, Li WL, et al.(2012) The 13 August 2010catastrophic debris flows after the 2008 Wenchuan earthquake, China. Natural hazards and earth system sciences 12(1):201–216. DOI: 10.5194/nhess-12-201-2012

Yao LK (1988) A research on the calculation of criticalrainfall with frequency of debris flow and torrentialrain. Journal of Soil and Water Conservation 2(4): 72-78 (in Chinese)

Ye SZ (1992) Hydrological calculation. Water conservancy and Hydropower Press, 111.

Zhou, W., & Tang, C. (2014). Rainfall thresholds for debris flow initiation in the Wenchuan earthquake-stricken area, southwestern China. Landslides, 11(5), 877-887.

Zhuang JQ, Cui P, Ge YG, et al. (2009) Relationship between rainfall characteristics and totalamount of debris flow. Journal of Beijing Forestry University 31(4): 77-83 (in Chinese)

Zhang SJ, Yang HJ, Wei FQ, et al. (2014)A Model of Debris Flow Forecast Based on the Water-SoilCoupling Mechanism. Journal of Earth Science, 25(4): 757-763. DOI:10.1007/s12583-014-0463-1

Zhenlei Wei, Yuequan Shang, Yu Zhao, et al. (2017) Rainfall threshold for initiation of channelized debris flows in a small catchment based on in-site measurement. Engineering Geology, 217, 23-34.