

1 **A Procedure to Select Earthquake Time Histories for Deterministic Seismic Hazard Analysis**
2 **from the Next Generation Attenuation (NGA) database**

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12
13 **Abstract:** In performance-based seismic design, ground-motion time histories are needed for analyzing
14 dynamic responses of nonlinear structural systems. However, the number of strong-motion data at
15 design level is often limited. In order to analyze seismic performance of structures, ground-motion time
16 histories need to be either selected from recorded strong-motion database, or numerically simulated
17 using stochastic approaches. In this paper, a detailed procedure to select proper acceleration time
18 histories from the Next Generation Attenuation (NGA) database for several cities in Taiwan is presented.
19 Target response spectra are initially determined based on a local ground motion prediction equation
20 under representative deterministic seismic hazard analyses. Then several suites of ground motions are
21 selected for these cities using the Design Ground Motion Library (DGML), a recently proposed
22 interactive ground-motion selection tool. The selected time histories are representatives of the regional
23 seismic hazard, and should be beneficial to earthquake studies when comprehensive seismic hazard
24 assessments and site investigations are yet available. Note that this method is also applicable to site-
25 specific motion selections with the target spectra near the ground surface considering the site effect.

26
27 **Keywords:** Ground motion selection, Seismic hazard analysis, NGA database, DGML tool

28 **1 Introduction**

29 In performance-based earthquake engineering, ground-motion time histories are usually needed for
30 analyzing the distribution of dynamic responses of nonlinear systems, such as site response or structural
31 analysis. In such an analysis, it is one of the key aspects to use appropriate acceleration time histories,
32 which should realistically reflect regional seismology and site conditions.

33 Understandably, the selected time histories should reasonably respond to seismic hazards at a given
34 site. For example, a recent technical guideline implemented by the U.S. Nuclear Regulatory
35 Commission (USNRC, 2007) prescribed the probabilistic seismic hazard analysis (PSHA) as the
36 underlying approach to generate time histories for future earthquake-resistant designs. Many studies
37 have highlighted the importance of matching a target response spectrum in the ground-motion selection
38 and modification process (e.g., Bommer and Acevedo, 2004). The target spectrum can be obtained by
39 deterministic seismic hazard analysis (DSHA), probabilistic seismic hazard analysis (PSHA) or seismic
40 design codes. A classic example is SIMQKE, which generates synthetic time histories to match a target
41 response spectrum with an iterative process using Gaussian random process and a time-varying
42 modulating function (Gasparini and Vanmarcke, 1976).

43 Recently, some scholars studied that a well-selected ground motion suite should match not only the
44 target mean, but also the variation of the target spectrum (Jayaram et al., 2011; Wang, 2011). In other
45 words, a suite of ground motions should be selected in performance-based earthquake engineering; the
46 resulting ground motion suite should properly capture the statistical distribution of ground motions
47 under the given earthquake scenario, which is commonly specified by means, standard deviations, and
48 inherent correlations (e.g., Baker and Jayaram, 2008; Wang and Du, 2012) of a target spectrum. There
49 are several ground motion selection algorithms available in the literature (Baker, 2010; Jayaram et al.,
50 2011; Wang, 2011). One of the recently proposed interactive tools is the Design Ground Motion Library
51 (DGML), which allows for selecting a suite of modified ground motions (multiple by scale factors) on
52 the basis of response spectral shape, as well as the characteristics of the recordings such as magnitude,
53 distances, faulting types and site conditions (Wang et al., 2015).

54 This paper aims at presenting a detailed procedure in selecting ground-motion time histories for
55 major cities of Taiwan using the DGML interactive tool. With deterministic seismic hazard analysis for

56 these cities, several suites of time histories are selected from the Pacific Earthquake Engineering
57 Research Center's Next Generation Attenuation (NGA) strong-motion database (Chiou et al., 2008).
58 Those selected motion suites are appropriate for general seismic designs, e.g., dynamic analysis of
59 structures in these cities.

60 **2 Deterministic Seismic Hazard Analyses (DSHA) for Major Cities in Taiwan**

61 **2.1 Overview of DSHA**

62 Seismic hazard analysis is an approach to describe the potential shaking intensity for future earthquakes,
63 which can be estimated by deterministic or probabilistic approaches. The deterministic approach
64 estimates the intensity measure amplitude (e.g., peak ground acceleration PGA as 0.2 g) under an
65 assigned earthquake scenario, while the probabilistic approach estimates the annual rate of exceeding
66 specific level of earthquake shaking at a site (e.g., PGA=0.2 g corresponding to 10% probability of
67 exceedance in 50 years).

68 **Compared to the complicated probabilistic approach, DSHA is an analysis accounting for a worst-**
69 **case scenario in terms of earthquake size and location. Specifically, DSHA utilizes the maximum**
70 **magnitude and shortest source-to-site distance to evaluate the ground motion intensities under such a**
71 **worse-case scenario. The basic steps are listed as follows: (1) Identify all possible fault sources of**
72 **earthquakes around a given site; (2) Define the maximum magnitude and closest distance for each fault;**
73 **(3) Compute the ground motion intensities based on attenuation relationships; (4) Take the maximum**
74 **intensity amplitudes as the final DSHA estimate.** Figure 1 shows a schematic diagram illustrating the
75 framework and the algorithm for DSHA. Seismic source models, the maximum earthquake of each
76 source, and ground motion prediction equations (GMPEs) are key inputs for DSHA. The detailed source
77 models and GMPEs used in this study would be introduced in this following subsection.

78 **2.2 Seismic source model and ground-motion model**

79 Figures 2 and 3 show the up-to-date seismic source models for Taiwan (Cheng et al., 2007), which have
80 also been used in a few seismic hazard studies by several authors (Cheng et al., 2007). It includes 20
81 area sources, in addition to 49 line sources associated with each active fault on this island. Table 1

82 summarizes the best-estimated maximum magnitude for each source from the literature (Cheng et al.,
83 2007). The seismic zonation used in this study (from Zone A to Zone T) is categorized as shallow
84 crustal regional source following previous researchers' work (e.g., Tsai 1986; Cheng 2002; Cheng et al.
85 2007). The maximum earthquake magnitude reflects a combined effect of regional seismology
86 regarding historical earthquakes, focal mechanism, and source zonation, etc. Thus, the maximum
87 magnitude of these seismogenic zones (e.g. M7.1 for source C) is adopted as the worst-case scenario
88 during DSHA calculations. The worst-case scenario was used for identifying the earthquake scenario
89 considered in DSHA analysis; for each area source considered, the closest source-to-site distance is
90 assigned accordingly, as listed in Table 3. The response spectra for major cities in Taiwan are also
91 presented in this section with DSHA calculations.

92 Ground motion prediction equations (GMPEs) are commonly used to predict ground motion
93 intensities (e.g., PGA) as a function of earthquake magnitude, source-to-site distance, site parameters,
94 etc. A few regional GMPEs models have been developed based on local strong-motion data in Taiwan
95 (Cheng et al., 2007; Lin et al., 2011). For example, Lin and Lee (2008) developed local GMPEs for
96 subduction earthquakes. Considering this subduction attenuation relation is incompatible with the NGA
97 database which primarily contains records from shallow crustal earthquakes, the model is not adopted in
98 the current study. On the other hand, the recent GMPE developed by Lin et al. (2011) is capable of
99 predicting PGA and response spectra for periods ranging from 0.01 s to 5 s, and therefore it is adopted
100 in this study, to develop the target response spectra for selecting earthquake time histories.

101 The function form of the adopted model (Lin et al. 2011) is expressed as follows:

$$102 \quad \ln Y = c_1 + c_2 M_w + c_3 \ln(R + c_4 e^{c_5 M_w}) \quad \sigma_{\ln Y} = \sigma^* \quad (1)$$

103 where Y denotes PGA or spectral accelerations in unit of g ; M_w refers to moment magnitude; R is the
104 rupture distance (closest distance from the rupture surface to site) in km ; c_1 to c_5 are regressed
105 coefficients. The model's coefficients are summarized in Table 2, and $\sigma_{\ln Y}$ denotes the model's standard
106 deviation. It is noted that this model was developed using around 5,000 earthquake records, 98% of
107 which are taken from Taiwan. It should be also noted that the attenuation adopted in this study is for the
108 hanging-wall and rock sites. It is agreeable that the worst-case scenarios considered in this manuscript

109 may not be the hanging wall case. However, since the Lin et al. (2011) model is the only available
110 regional-specific response spectral attenuation model for shallow crustal earthquakes to the authors'
111 best knowledge, this hanging-wall attenuation model is adopted in the current study to construct the
112 target response spectra with reasonably conservative results provided.

113 It is also worth noting that we only employ the local ground motion model in this study. It is
114 understood that logic-tree analyses can be used to quantify the so-called epistemic uncertainty in PSHA.
115 But as studied by some scholars (e.g., Krinitzsky, 2003), the weights in logic-tree analyses cannot be
116 scientifically verified. Therefore, this study used one local model available as the best estimate. When
117 new local models are developed, the update of seismic hazards or sensitivity analyses will be worth
118 conducting in future.

119

120 **2.3 DSHA-based response spectra for major cities in Taiwan**

121 The aforementioned DSHA procedures can be performed for major cities in Taiwan, with the adopted
122 seismic source models (Figures 2 and 3) and attenuation relationship introduced in previous subsections.
123 Six major cities are chosen for such calculations, including site (a) Taipei, site (b) Kaohsiung, site (c)
124 Taichung, site (d) Chiayi, site (e) Pingtung and site (f) Hualien. Figure 2 shows location of the six study
125 sites, with their coordinates (i.e., the city's geographical centers) summarized in Table 3. As is well
126 know that Taiwan is located at the boundary between the Philippine Sea Plate to the East and the
127 Eurasian Plate to the West, the six study sites are intentionally selected to represent a variety of
128 geological components of the island: site (a) Taipei, site (c) Taichung and site (d) Chiayi are located at
129 the Western Foothills mainly composed of Oligocene to Pleistocene clastic sediments, where syn-
130 orogenic sediments of the foreland basin have been accreted and deformed.; site (b) Kaohsiung is
131 located at the Coastal Plain as a part of the foreland basin of Taiwan; site (e) Pingtung is located within
132 the West Central Range with mostly Miocene to Eocene slates, corresponding to the area of highest
133 altitudes in Taiwan; site (f) Hualien is located at the Longitudinal Valley which is believe as the suture
134 zone between the Luzon arc and the Chinese continental margin (CES, 2017). For each study site, the
135 worse-case scenario was firstly identified, and then the corresponding response spectrum was
136 determined by using the adopted local GMPE.

137 Figure 4 shows the resulting response spectra from DSHA calculations for the six considered cities
138 in Taiwan. Table 3 also summarizes the controlling seismic source for each site. For example, the
139 DSHA seismic hazard at the center of **site (a) Taipei** is governed by Area Source C. In other words, the
140 Area Source C, rather than the other line sources or active faults, contributes to the deterministic seismic
141 hazard for the center of Taipei. The same situation is occurring to other cities with an area source being
142 the controlling source. This is expected, since the DSHA seismic hazard from an area source could be
143 commonly higher than a line source due to the relatively closer source-to-site distance.

144 It should be noted that the adopted local GMPE has been thoroughly compared with the globally
145 NGA GMPEs (Abrahamson and Silva, 2008; Boore and Atkinson, 2008; Campbell and Bozorgnia,
146 2008; Chiou and Youngs, 2008). In general, the PGA amplitudes predicted by the adopted model is
147 generally comparable to those of the NGA models, except that for scenarios with distances greater than
148 20 km the estimated PGAs of the local model attenuate faster. The steeper slope of the local attenuation
149 curves could be due to the fact that the local crust is relatively weak, given that Taiwan is a very young
150 orogeny (Lin et al., 2011). This implies that a design or target spectrum derived from local GMPEs is
151 particularly necessary for selecting suitable ground-motion time histories for local engineering practice.

152 **3 Selection of Ground-Motion Time Histories**

153 **3.1 The NGA database and Design Ground Motion Library (DGML)**

154 The source for ground-motion selection in this study is the PEER-NGA strong motion database, which
155 contains 3,551 three-component recordings from 173 earthquakes (Chiou et al., 2008). Various subsets
156 of the database have been used to develop GMPE models for various ground motion intensities in
157 earthquake engineering (e.g., Du and Wang, 2013; Foulser-Piggott and Stafford, 2012). Figure 5 shows
158 the moment magnitude-rupture distance distribution of the ground motions in the NGA database. The
159 aforementioned interactive tool, DGML, is used to search ground-motion time histories in the NGA
160 database on the basis of similarity of a record's response spectrum to the target response spectrum over
161 a use-defined range of period (Wang et al., 2015). The DGML has the broad capability of searching for
162 ground-motion time histories in the library database on the basis of response spectral shape,
163 characteristics of the recordings in terms of earthquake magnitude and type of faulting distance, site

164 characteristics, duration, and presence of velocity pulses in near-fault time histories. These ground
165 motions intensity measures have been found important in liquefaction and seismic assessment of a
166 variety of geotechnical systems (Wang and Wei, 2016; Ye and Wang, 2016). The DGML Search Engine
167 window is shown in Figure 6. It contains the following major parts: (1) Inputs for searching criteria; (2)
168 Prescribed range of scale factor; (3) Prescribed weight factor for spectral period; (4) Spectrum plot of
169 selected motions; (5) MSE (mean squared error) of each individual selected ground-motion record; (6)
170 Scale factor of each record; (7) Event name and (8) station name of each record.

171 To be more specific, it is requested to specify the seismological parameter bounds (e.g., range of
172 considered M_w and distance R) as inputs, which can implicitly constrain the ground motion
173 characteristics in addition to the explicit target spectrum. Given the fact that the target spectra from
174 DSHA are a result of the maximum earthquake and the closest source-to-site distance, a relatively large
175 magnitude bound ($5.5 < M_w < 8$) and a narrow distance range ($0 \text{ km} < R_{rup} < 30 \text{ km}$) have been employed as
176 the searching criteria. Since all the six cities are located at soil sites, a V_{s30} (time-averaged shear-wave
177 velocity down to 30 m) bound in the range of 0-450 m/s is also applied. Other causal parameters, such
178 as the category of fault types or the range of duration parameters, are not particularly specified.

179 Scaling factor is another key input for selecting ground motions, but has been subjected to intense
180 debate over the past decades. Previous researchers pointed out that improper scaling of a record can lead
181 to bias estimates of structural responses (Luco and Bazzurro 2007). For example, if an excessive range
182 of scale factors is applied, the selected ground motion suite might result in drastically biased
183 distribution of the other ground-motion characteristics, such as duration, Arias intensity, that cannot be
184 represented by the target response spectrum. Therefore, we follow the general practice of the Design
185 Ground Motion Library (DGML) and assign a relative narrow range of scale factors (0.4-2.5)
186 throughout the selection procedure in this study (Wang et al. 2015). After searching for properly
187 matched time histories with target spectrum and magnitude and distance thresholds, the ranking of
188 earthquake motions is tabulated after spectral matching process. The motions of interest can be
189 downloaded from the list, as well as their descriptions such as fault types, earthquake magnitudes,
190 rupture distances, durations, scaling factors, and V_{s30} values (V_{s30} is commonly employed site condition

191 indicator). Note that DGML is also capable of performing weight-matching when a specific range of the
192 motion's frequencies is of more interest in follow-up applications.

193 **3.2 Time history recommendations for major cities of Taiwan**

194 With the target spectra from DSHA calculations, the selection procedures in DGML are performed to
195 select a suite of time histories from the NGA database for each city. **The DGML search engine adopted**
196 **in this study searches the NGA database for ground-motion waveforms that satisfy the general criteria**
197 **(i.e. $5.5 < M_w < 8$, $0 < R_{rup} < 30$ km) and then ranks these records in an order of an increasing MSE (mean**
198 **squared error). It means that the ground-motion waveform that matches the target RS best has the**
199 **lowest MSE and will be ranked No. 1. To be more specific, the MSE is defined using the following**
200 **equation (Wang et al. 2015):**

$$201 \quad \text{MSE} = \frac{\sum_i w(T_i) \left\{ \ln \left(Sa^{target}(T_i) \right) - \ln \left(f \times Sa^{record}(T_i) \right) \right\}^2}{\sum_i w(T_i)} \quad (2)$$

202 where T_i denotes considered spectral periods, $w(T_i)$ denotes a weight function that allows for assigning
203 weights to different period ranges so that the periods of more interest can be emphasized in the ground-
204 motion selection process, f represents a scale factor to linearly scale the whole ground-motion time
205 history. It should be also noted that the MSE does not vary too much in some cases. For example, as
206 shown in Figure 6, the MSE ranges from 0.023-0.035, indicating that the selected scaled ground
207 motions are almost equally good and compatible with the target response spectrum. Therefore, in this
208 study, we intentionally select some other ground-motion waveforms if some of them have been
209 recommended in the other study cities. We expect, by doing so, more flexibility and options can be
210 provided for time-history analyses in engineering practice. It should be also noted that although
211 different ground motions are selected for various sites, they are statically consistent and compatible with
212 the corresponding DSHA spectrum.

213 Figure 7 shows the selected response spectra for the six study cities. The median and median \pm one
214 standard deviation of the selected SA ordinates are also compared to the target spectrum in each plot. It
215 can be seen that the selected ground motion suites can properly match the target spectra over a wide

216 period range. Table 4 summarizes the time histories selected from the database. Figures 8-14 show the
217 selected time histories for the six cities in Taiwan with seismic hazards calculated with DSHA
218 calculations. Note that two sets of selections were given for **site (a) Taipei**, with and without the
219 consideration of basin effect. It should also be noted that for each site the best-matching motions were
220 selected regardless of local earthquakes or not, in addition to one or two best-matching local motion (i.e.,
221 the Chi-Chi earthquake). The multiple time histories in each suite are considered as a measure to
222 account for the variability or natural randomness of ground motion characteristics under a considered
223 scenario, which, for example, is considered as mandatory for probabilistic site response analyses
224 prescribed in a technical reference (USNRC, 2007).

225 **4 Discussions**

226 **4.1 DSHA versus PSHA**

227 PSHA and DSHA are the two representative approaches in assessing earthquake hazards. Over the past
228 decades, numerous seismic hazard studies have been conducted with the two methods (e.g., Joshi et al.,
229 2007; Kolathayar and Sitharam, 2012; Moratto et al., 2007; Sitharam and Vipin, 2011; Stirling et al.,
230 2011). The two methods have also been prescribed in various technical references. As mentioned
231 previously, a technical reference (USNRC, 2007) prescribes PSHA as the underlying approach, in
232 contrast to another guideline implemented by Department of California Transportation prescribing
233 DSHA for bridge designs under earthquake loadings (Mualchin, 2011).

234 It is worth noting that extensive discussions over the pros and cons of the two methods have been
235 reported in the literature (e.g., Bommer, 2003; Castanos and Lomnitz, 2002; Krinitzsky, 2003; Klugel,
236 2008). In general, DSHA is a simple approach that earthquake scenarios are considered logically
237 understandably, but the uncertainties in DSHA may not be well quantified. On the other hand, PSHA is
238 capable of quantifying the uncertainties associated with earthquake scenarios via a probabilistic
239 approach; however, some scholars (e.g., Krinitzsky, 2003) pointed out the shortcomings in PSHA, such
240 as the uniform assumption in the occurrences of earthquakes. It is not this paper's purpose to argue
241 which seismic hazard method is superior. But with all that in mind, it should come to a logical
242 understanding that both the deterministic and probabilistic analyses are needed and useful in

243 engineering applications. The use of the DSHA approach in this study is primarily due to its analytical
244 simplicity and transparency. Since it has been reported that DSHA rather than PSHA is more
245 appropriate for design of critical structures (Bommer et al., 2000), the selected ground motion suites,
246 with a representative seismic hazard analysis and a reputable earthquake database, are then
247 recommended for such applications.

248

249 **4.2 Site-specific time histories**

250 This paper presents an option to select earthquake time histories from the reputable NGA database. But
251 strictly speaking, those time history recommendations are not site-specific, because the site condition is
252 not carefully taken into account with a comprehensive site investigations and site response analyses. In
253 other words, the site-specific motions are those from seismic hazard analyses, to site response studies
254 (e.g., Du and Pan, 2016).

255 As a result, this study refers to those time-history recommendations as “tentative site-specific,”
256 because the site effect is not comprehensively characterized with a more detailed site response analysis,
257 but with a soil-site ground motion prediction model. Therefore, the selected ground motion time-
258 histories could be recommended for general earthquake analytical cases, where specific site
259 investigations are not performed. Since the recommended time-histories can reasonably reflect the local
260 seismic hazards at these cities, they should be used as basic results and then be serviceable for common
261 engineering practice.

262

263 **4.3 Basin effect**

264 Basin effect is another important issue to estimate the seismic hazards for **site (a) Taipei**. From
265 analyzing the recorded time histories around Taipei (Sokolov et al., 2009; 2010), some suggestions were
266 made to up-scale low-frequency spectral accelerations to incorporate the basin effect in Taipei.
267 Following this suggestion, Figure 15 shows the response spectra with/without considering basin effects
268 for Taipei by DSHA calculations. Likewise, the time histories matching the up-scaled spectra (with
269 basin effects) as the target are selected from the database, as summarized in Table 4.

4.4 Why local earthquake's motions are not selected for all cases?

It somewhat comes to as a surprise that the motions of the local earthquake were “out-performed” by non-local motions in matching the response spectra with local ground motion models. This might be due to the following reasons. First, apart from the Chi-Chi earthquake, most events used for developing the local GMPE are not included in the NGA database. Besides, the employed searching process does not specify more weights or preferences to local earthquakes. As discussed previously, the search criterion are only associated with the spectral shape, as well as seismological parameters such as magnitude, distance, site condition, *etc.* The search engine searches the database and ranks the records based on a quantitative measure: the MSE. Table 4 summarizes MSE for each single selected record. It can be observed that the MSEs range from 0.023-0.046 for different study sites. These ground-motion waveforms have been recommended in this study based on their compatibility with the target response spectrum, and such compatibility is parametrized as the MSE. It is yet agreeable that local ground motions may contain intrinsic correct path effects at the site of interest. However, the principle of current ground-motion selection practice is searching for time history record sets in the database on the basis of the similarity of a record's response spectral shape to a design response spectrum over a user-defined period range. In such a case, local earthquake records are not always selected and recommended for the study sites because they might not conform to the target spectra. With this in mind, as long as the size of the database is sufficient, it is not surprising that a non-local ground motion can be found better matching the target spectra.

It is also worth mentioning that the local GMPE adopted in this study is a “generalized” one, and is primarily consistent with global GMPEs in terms of spectral shape and response acceleration amplitudes. Figure 16 compares spectral accelerations predicted using the local model with those computed using other four widely used NGA global GMPEs, namely, AS08, BA08, CB08 and CY08, for several earthquake scenarios (Abrahamson and Silva 2008; Boore and Atkinson 2008; Campbell and Bozorgnia 2008; Chiou and Youngs 2008). It can be seen that under earthquake scenario $M=7$, $R_{rup}=30$ km, $V_{S30} = 760$ m/s, the spectral accelerations predicted by local attenuation agree well with the BA08 and CY08 models across a wide range of periods (i.e. from 0.01 s to 5 s). As for the scenarios $M=7$, $R_{rup}=10$ km, $V_{S30} = 760$ m/s, the spectral accelerations predicted by local GMPE again corresponds well with those

298 computed using the CY08 model, as demonstrated in Figure 16 (a). Apart from the consistency with
299 global GMPEs, it is also worth mentioning that the functional form of the local model is based on
300 Campbell (1981), which is a quite generic and widely adopted one. Therefore, in this study, ground
301 motions that are selected from the comprehensive NGA database based on compatibility with the target
302 response spectra can be either local records or global ones, but not necessarily local motions, given the
303 “generic target GMPE”.

304

305 **5 Conclusions**

306 The paper presented the procedures to select earthquake time histories with target response spectra from
307 deterministic seismic hazard analysis (DSHA), using the recently proposed DGML selection tool. The
308 worst-case earthquake scenarios were first defined for six major cities in Taiwan, and the response
309 target spectra were computed by employing a regional attenuation model under these defined scenarios.
310 Finally, a suite of time histories are selected for each city by matching the calculated target spectra. The
311 selected suites of time histories can properly represent the regional seismic hazards, which are then
312 recommended and used for seismic analyses in these cities. The similar ground motion selection
313 approaches can also be applicable to selecting appropriate time histories at bedrock layers, as input
314 motions for a more comprehensive site investigations and site response analysis.

315 Given the limited understandings of the earthquake process and the randomness in nature, some
316 scholars have pointed out the importance of analytical simplicity to earthquake studies. Among several
317 approaches to define the target spectra, the ones from DSHA calculations are logically transparent and
318 simple, and therefore they are adopted in this study for selecting hazard-consistent time histories.

319 **Acknowledgments**

320 The first author acknowledges financial support provided by Jockey Club Institute for Advanced Study
321 at the Hong Kong University of Science and Technology for this work.

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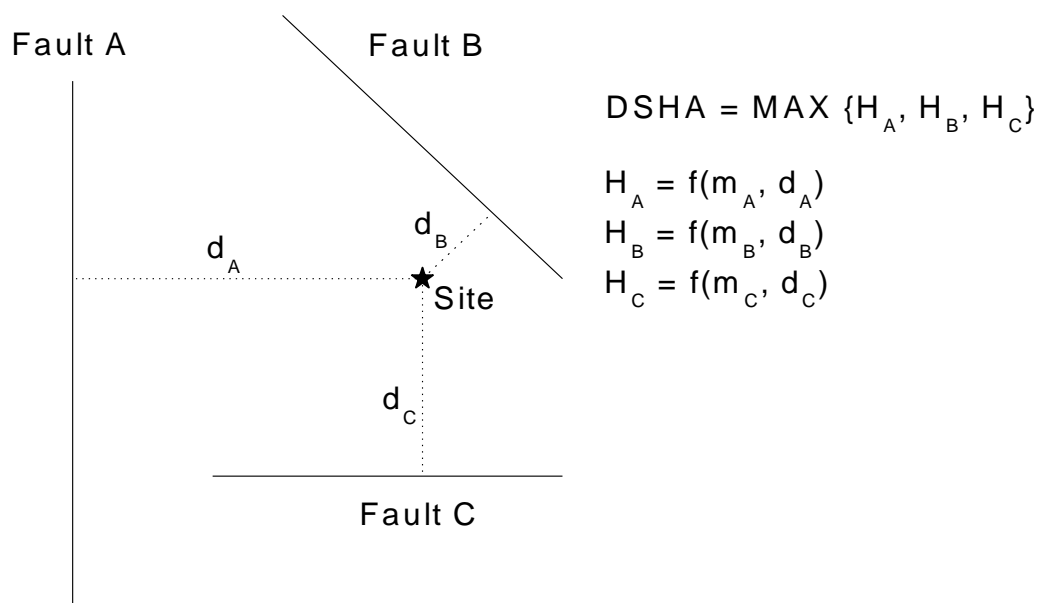


Figure 1. Schematic diagram illustrating the analytical framework of DSHA, where the solid star represents a study site, H denotes the seismic hazard induced by each source, m and d are the maximum earthquake magnitude and shortest source-to-site distance, and f is the function of a ground motion model

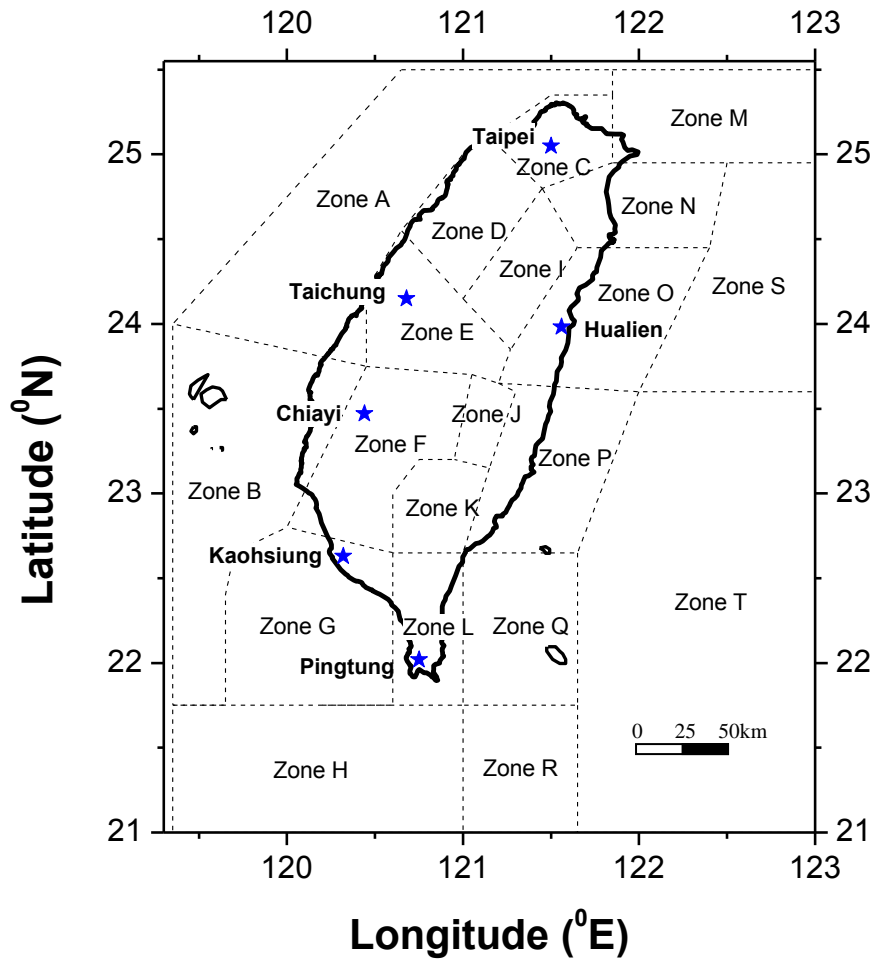


Figure 2. The area seismic source model for Taiwan (after Cheng et al., 2007); six study cities including (a) Taipei, (b) Kaohsiung, (c) Taichung, (d) Chiayi, (e) Pingtung and (f) Hualien, are plotted as blue stars.

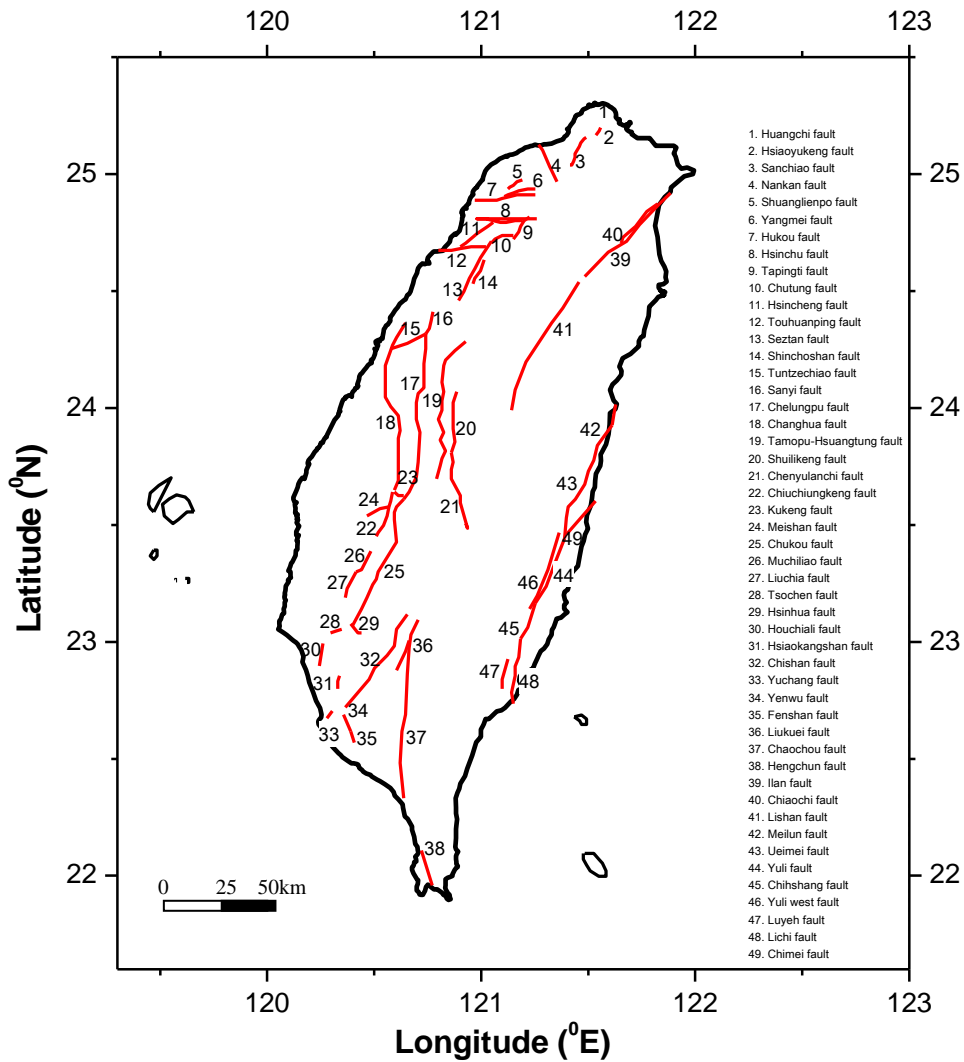


Figure 3. The line source model of the active faults in Taiwan (after Cheng et al., 2007)

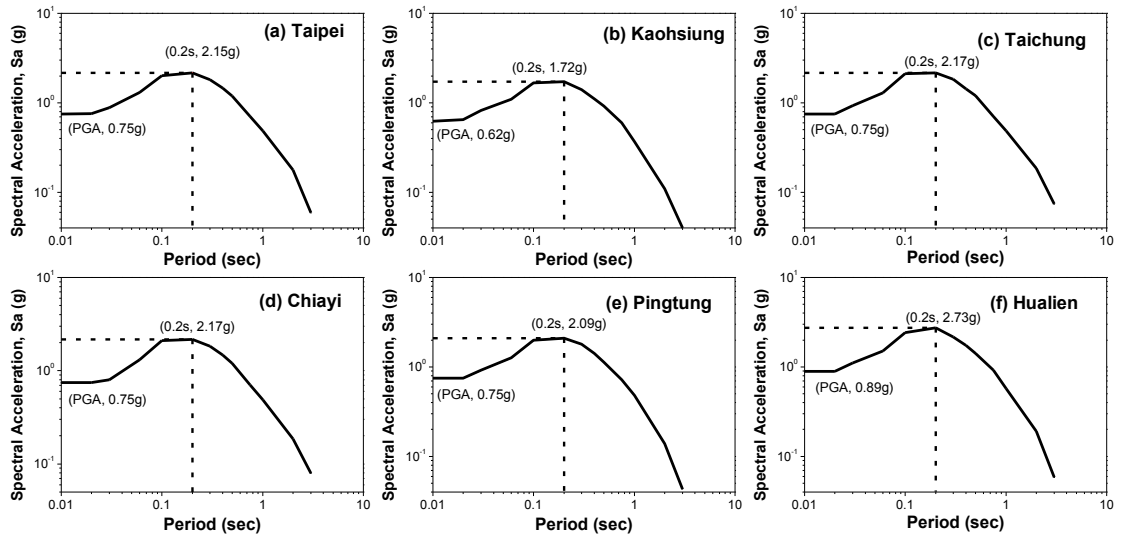


Figure 4. The response spectra for major cities in Taiwan with DSHA calculations

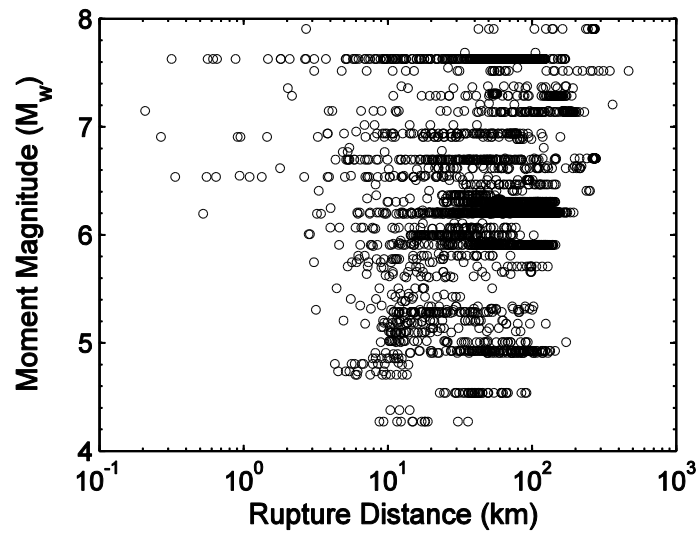


Figure 5. Moment magnitude and rupture distance distribution for PEER NGA records used in this study

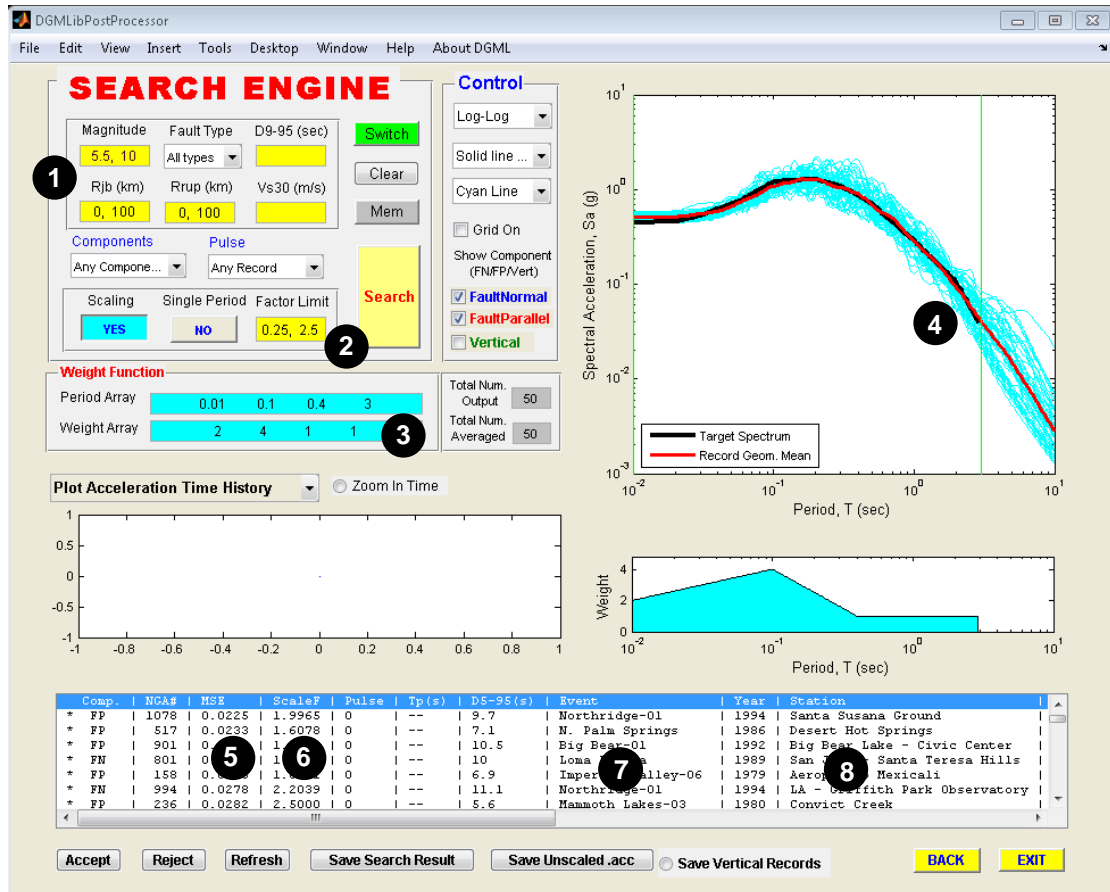


Figure 6. The screenshot of the DGML searching interface: (1) Searching criteria; (2) Prescribed range of scale factor; (3) Prescribed weight factor for spectral period; (4) Spectrum plot of selected motions; (5) MSE of each individual selected ground-motion record; (6) Scale factor of each record; (7) Event name of each record and (8) station name of each record.

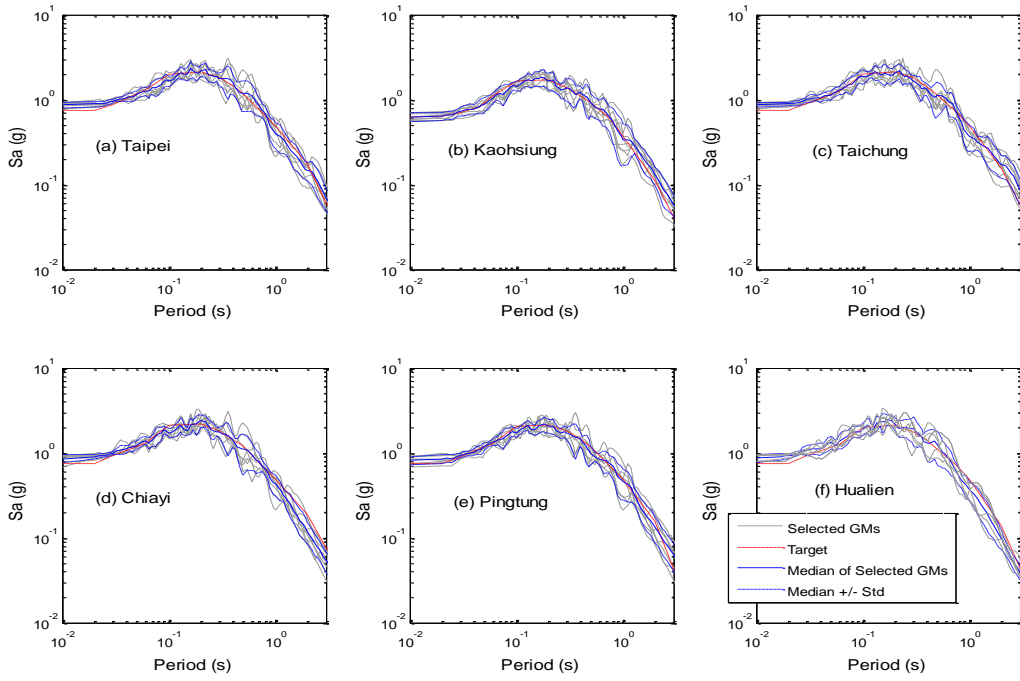


Figure 7. The target spectrum, individual and average response spectrum of selected records for six major cities in Taiwan

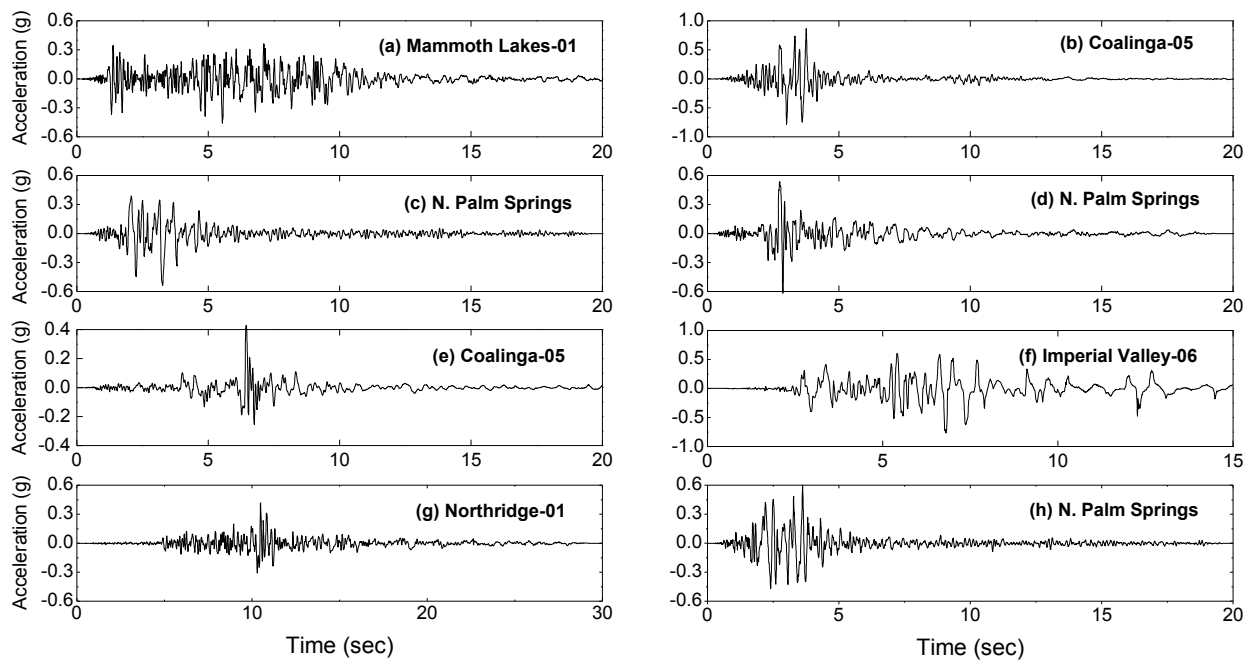


Figure 8. Eight time history recommendations for site (a) Taipei with DSHA calculations and the NGA strong-motion database

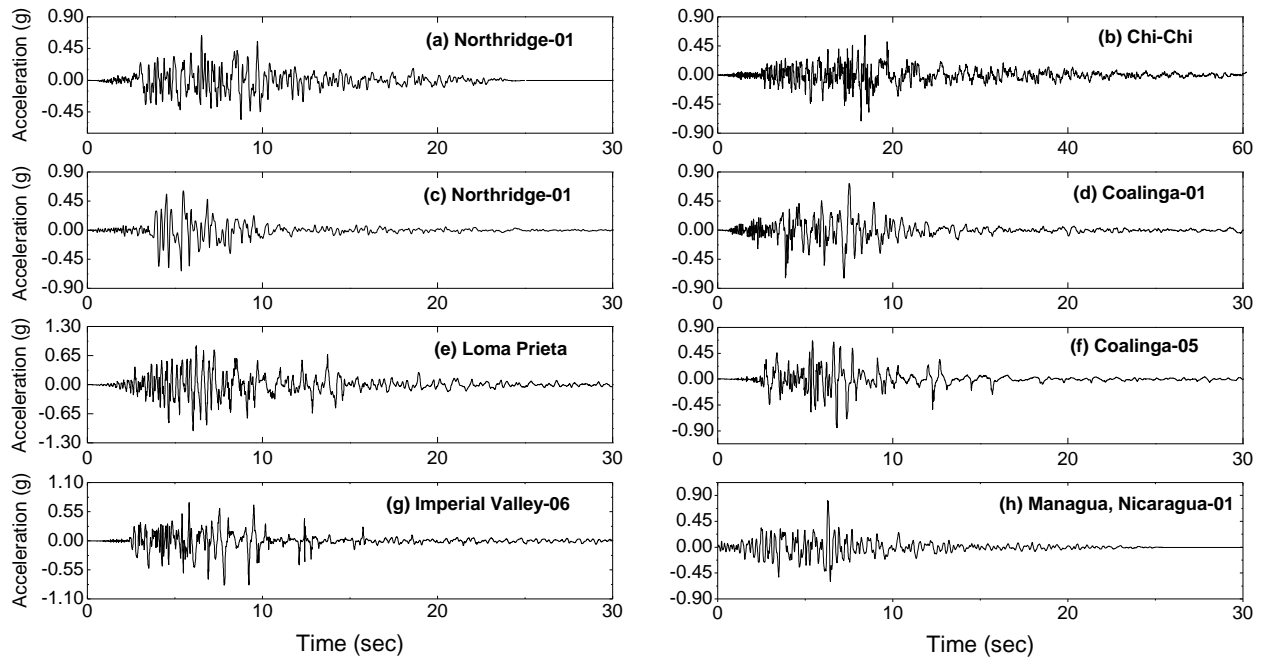


Figure 9. Another set of time history recommendations for site (a) Taipei with the basin effect taken into account

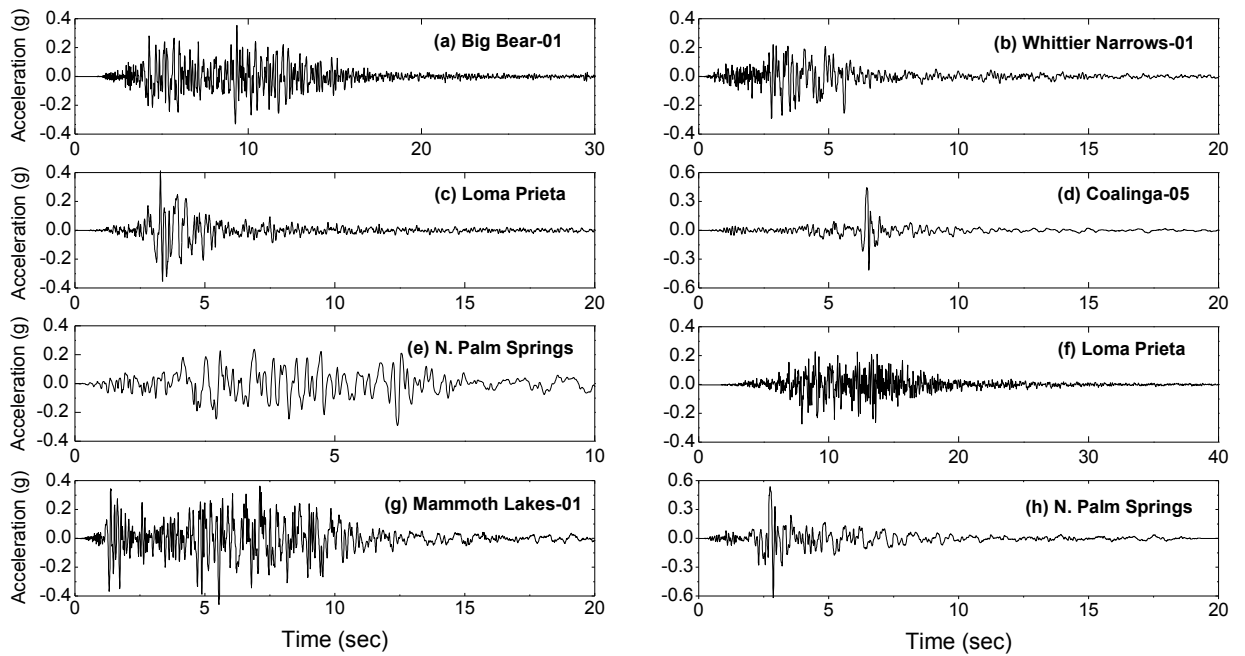


Figure 10. Eight time history recommendations for site (b) Kaohsiung with DSHA calculations and the NGA strong-motion database

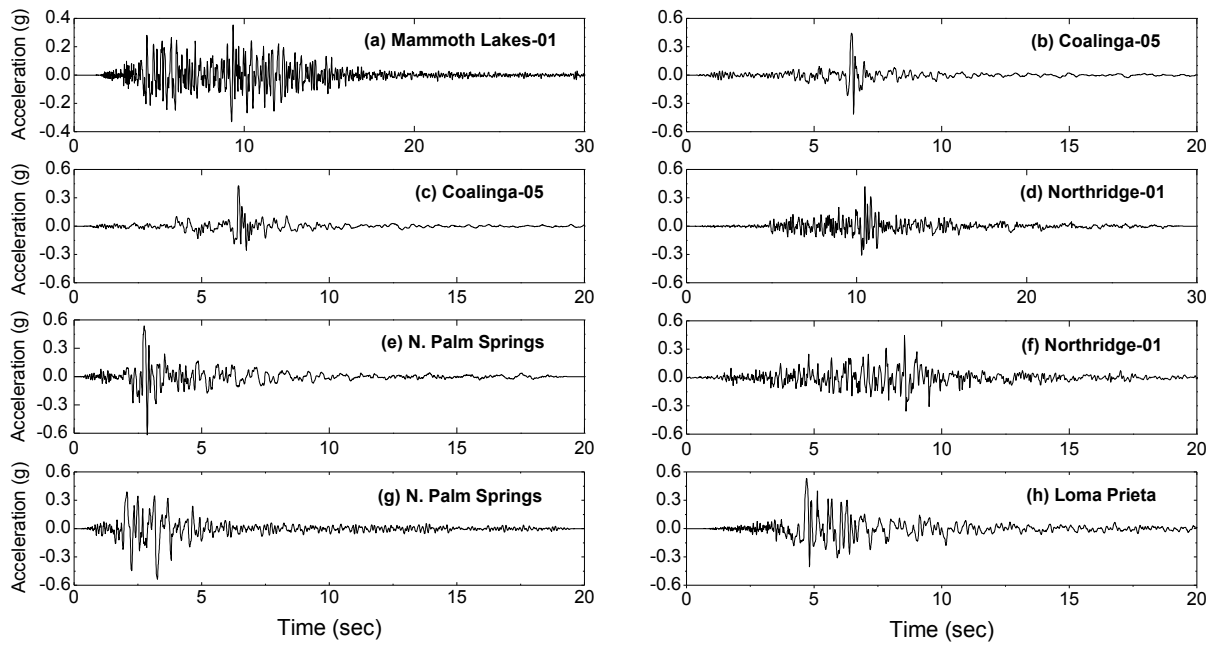


Figure 11. Eight time history recommendations for site (c) Taichung with DSHA calculations and the NGA strong-motion database

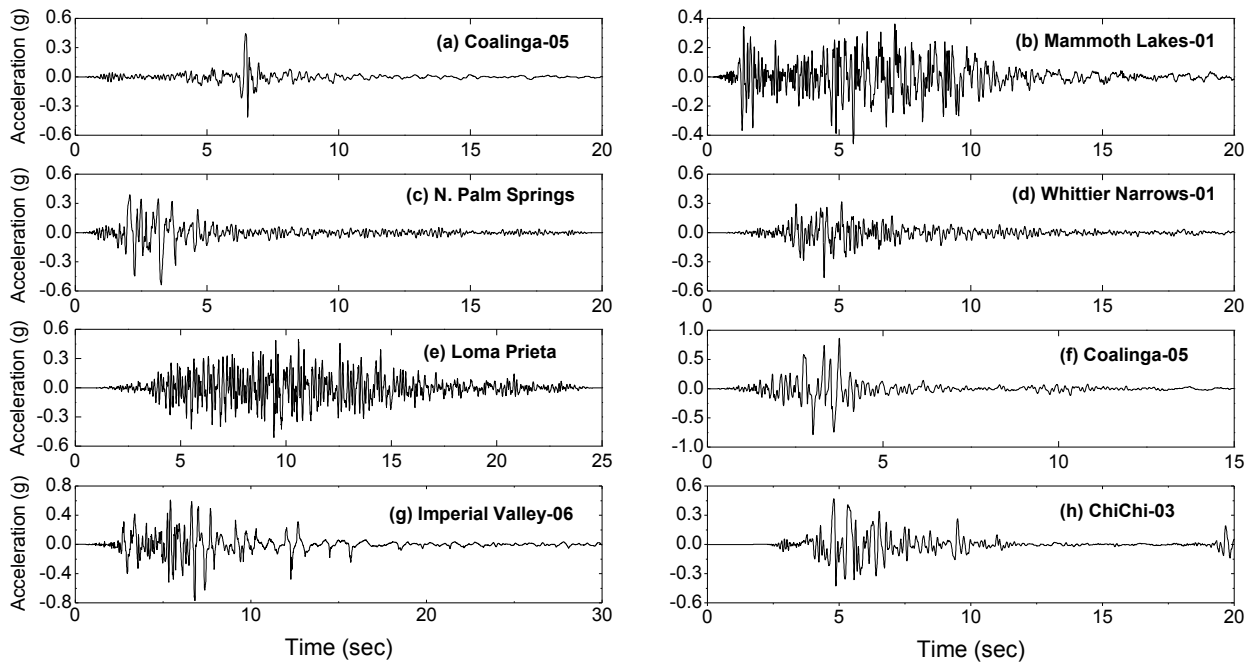


Figure 12. Eight time history recommendations for site (d) Chaiyi with DSHA calculations and the NGA strong-motion database

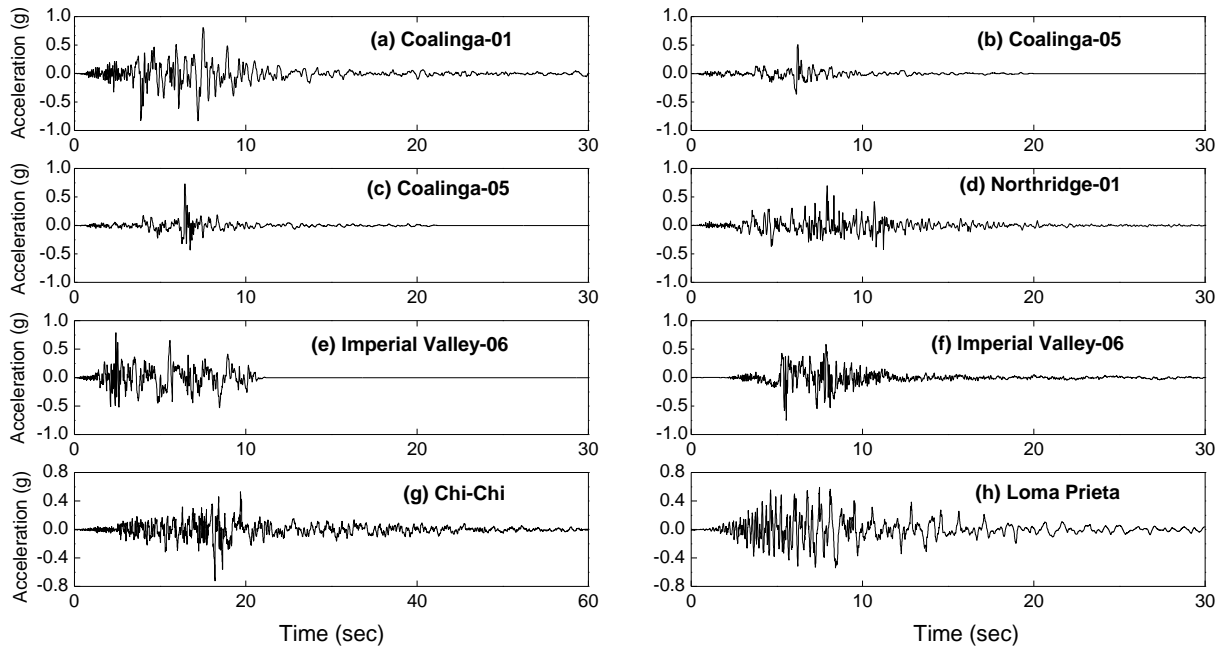


Figure 13. Eight time history recommendations for site (e) Pingtung with DSHA calculations and the NGA strong-motion database

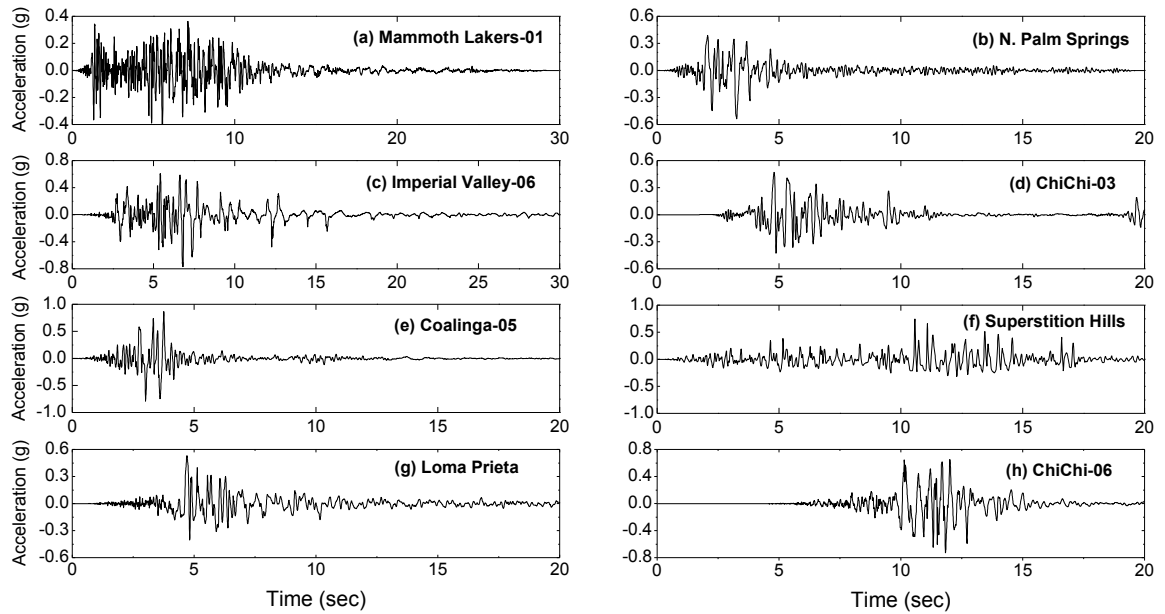


Figure 14. Eight time history recommendations for site (f) Hualien with DSHA calculations and the NGA strong-motion database

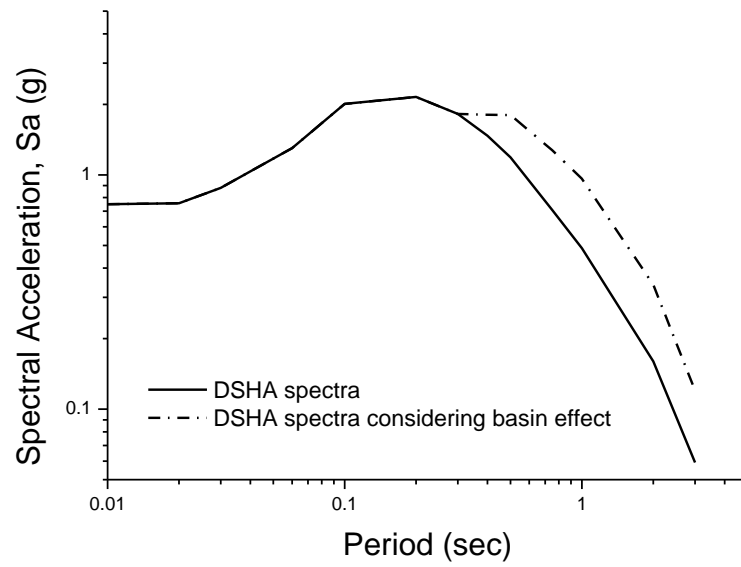


Figure 15. The basin effect in site (a) Taipei on response spectra; the spectra scaling follows the suggestions of Solokov et al. (2009, 2010)

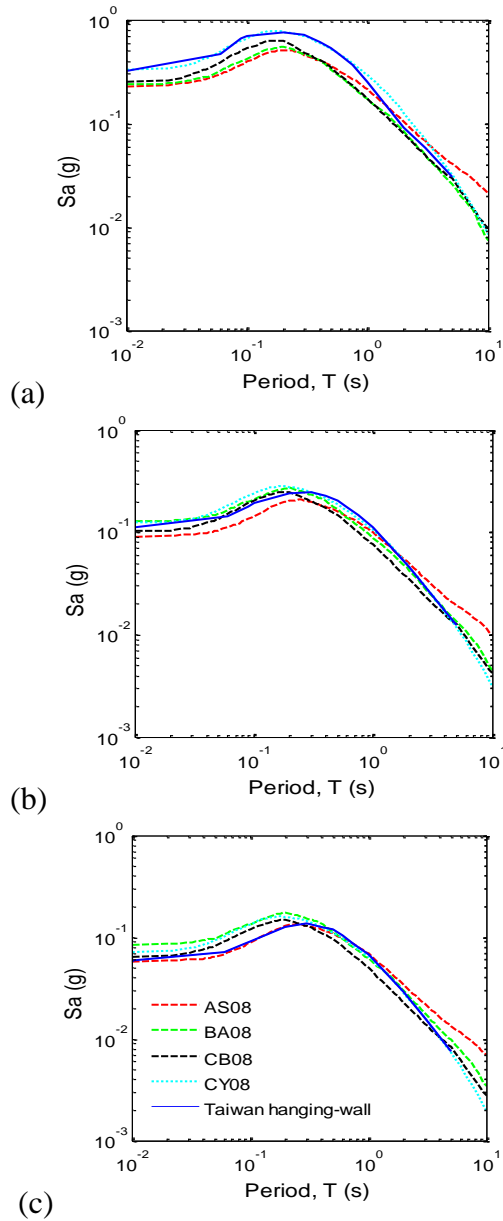


Figure 16. Comparison of spectral acceleration predicted using the local GMPE (Lin et al 2011) and NGA global ground motion prediction equations (GMPE) under three earthquake scenarios (a) $M=7$, $R_{rup}=10$ km, $V_{s30} = 760$ m/s, (b) $M=7$, $R_{rup}=30$ km, $V_{s30} = 760$ m/s; (c) $M=7$, $R_{rup}=50$ km, $V_{s30} = 760$ m/s

Table 1. Summary of Maximum Earthquake Magnitudes (in M_w) of Each Seismic Source around Taiwan

Area source	Max. magnitude	Line source (active fault)	Max. magnitude	Fault Mechanism
Zone A	6.5	Huangchi	7.0	Normal & Sinistral
Zone B	6.5	Hsiaoyukeng	7.0	Normal & Sinistral
Zone C	7.1	Sanchiao	7.0	Normal & Sinistral
Zone D	7.3	Nankan	6.5	Normal & Dextral
Zone E	7.3	Shuanglienpo	6.2	Reverse
Zone F	7.3	Yangmei	6.6	Reverse
Zone G	6.5	Hukou	6.9	Thrust
Zone H	7.3	Hsinchu	6.8	Thrust
Zone I	6.5	Tapingti	6.5	Thrust
Zone J	6.5	Chutung	6.5	Reverse
Zone K	6.5	Hsincheng	6.7	Thrust
Zone L	7.3	Touhuanping	6.7	Dextral
Zone M	6.5	Seztan	6.8	Reverse
Zone N	8.0	Shinchoshan	6.5	Reverse
Zone O	8.3	Tuntzechiao	6.5	Dextral
Zone P	7.8	Sanyi	6.9	Thrust
Zone Q	7.8	Chelungpu	7.7	Thrust
Zone R	7.8	Changhua	7.6	Thrust
Zone S	8.0	Tamopu-Hsuangtung	7.4	Thrust
Zone T	7.8	Shuilikeng	7.0	Thrust

Table 1. Summary of Maximum Earthquake Magnitudes (in M_w) of Each Seismic Source around Taiwan (Continued)

Line source (active fault)	Max. magnitude	Fault Mechanism	Line source (active fault)	Max. magnitude	Fault Mechanism
Chenyulanchi	7.0	Thrust	Lishan	6.9	Normal
Chiuchiungkeng	7.0	Thrust	Meilun	7.3	Norma & Sinistral
Kukeng	6.3	Sinistral	Ueimei	7.5	Norma & Sinistral
Meishan	6.5	Dextral	Yuli	7.5	Norma & Sinistral
Chukou	7.5	Thrust	Chihshang	7.3	Norma & Sinistral
Muchiliao	7.1	Thrust	Yuli west	7.3	Norma & Sinistral
Liuchia	7.1	Thrust	Luyeh	6.9	Reverse
Tsochen	6.4	Sinistral	Lichi	7.1	Norma & Sinistral
Hsinhua	6.4	Dextral	Chimei	7.2	Norma & Sinistral
Houchiali	6.4	Thrust			
Hsiaokangshan	6.5	Reverse			
Chishan	7.3	Thrust			
Yuchang	6.4	Reverse			
Yenwu	6.7	Reverse			
Fenshan	6.7	Reverse			
Liukuei	6.7	Reverse			
Chaochou	7.3	Reverse			
Hengchun	7.2	Reverse			
Ilan	6.9	Normal			
Chiaochi	6.8	Normal			

Table 2. Summary of the Coefficients of the Local Ground Motion Models used in This Study (Lin et al. 2011)

Periods (s)	c_1	c_2	c_3	c_4	c_5	$\sigma_{\ln Y}$
PGA	-3.279	1.035	-1.651	0.152	0.623	0.651
0.01	-3.253	1.018	-1.629	0.159	0.612	0.647
0.06	-1.738	0.908	-1.769	0.327	0.502	0.702
0.09	-1.237	0.841	-1.750	0.478	0.402	0.748
0.1	-1.103	0.841	-1.765	0.455	0.417	0.750
0.2	-2.767	0.980	-1.522	0.097	0.627	0.697
0.3	-4.440	1.186	-1.438	0.027	0.823	0.685
0.4	-5.630	1.335	-1.414	0.014	0.932	0.683
0.5	-6.746	1.456	-1.365	0.006	1.057	0.678
0.6	-7.637	1.557	-1.348	0.0033	1.147	0.666
0.75	-8.641	1.653	-1.313	0.0015	1.257	0.652
1	-9.978	1.800	-1.286	0.0008	1.377	0.671
2	-12.611	2.058	-1.261	0.0005	1.497	0.706
3	-13.303	2.036	-1.234	0.0013	1.302	0.702

Table 3. Summary of the Site's Coordinates, along with Respective Controlling Seismic Sources for Each Site in DSHA Computations

City	Latitude (° N)	Longitude (° E)	Controlling source	Maximum magnitude	Closest source-to-site distance (km)
(a) Taipei	25.05	121.50	Zone C	7.1	2
(b) Kaohsiung	22.63	120.32	Zone G	6.5	2
(c) Taichung	24.15	120.68	Zone E	7.3	2
(d) Chiayi	23.47	120.44	Zone F	7.3	2
(e) Pingtung	22.02	120.75	Zone L	7.3	2
(f) Hualien	23.98	121.56	Zone O	8.3	2

Table 4. Summary of the Earthquake Time History Recommendations from the NGA Database with DSHA Calculations

City	Earthquake motion	Year	Magnitude	Rupture Distance (km)	Station	Fault Mechanism	D ₅₋₉₅ (s)	V _{s30} (m/s)	Scale Factor	MSE
Taipei	Mammoth Lakes-01	1980	6.06	4.0	Convict Creek	N-O***	9.1	338	1.67	0.023
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.-FP	Reverse	5.0	257	1.93	0.023
	N. Palm Springs	1986	6.06	6.0	Whitewater Trout Farm	R-O**	5.1	345	1.67	0.023
	N. Palm Springs	1986	6.06	11.2	North Palm Springs	R-O**	5.6	345	1.48	0.030
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.-FN	Reverse	5.0	257	2.00	0.031
	Imperial Valley-06	1979	6.53	2.7	Bonds Corner	Reverse	9.7	223	1.05	0.031
	Northridge-01	1994	6.69	28.3	LA – Centinela St.	Reverse	13.0	235	0.98	0.031
	N. Palm Springs	1986	6.06	6.0	Whitewater Trout Farm	R-O**	5.1	345	1.52	0.032
Taipei (with basin effect)	Northridge-01	1994	6.69	14.7	Canoga Park	Reverse	11.1	268	0.50	0.030
	Chi-Chi	1999	7.62	10.0	CHY101	R-O**	29.0	259	1.16	0.033
	Northridge-01	1994	6.69	28.3	LA – Centinela St.	Reverse	13.0	235	0.98	0.039
	Coalinga-01	1983	6.36	8.4	Pleasant Valley P.P.	Reverse	8.0	257	1.32	0.039
	Loma Prieta	1989	6.93	15.2	Capitola	R-O**	14.7	289	1.50	0.039
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.	Reverse	5.0	257	1.22	0.041
	Imperial Valley-06	1979	6.53	2.7	Bonds Corner	Strike-Slip	9.7	223	1.35	0.041
	M. - N.* -01	1972	6.24	4.1	Managua- ESSO	Strike-Slip	9.0	289	2.00	0.043
Kaohsiung	Big Bear-01	1992	6.46	9.4	Big Bear Lake	Strike-Slip	10.5	338	1.88	0.025
	Whittier Narrows-01	1987	5.99	14.5	Garvey Res	R-O**	5.9	468	2.02	0.027
	Loma Prieta	1989	6.93	10	Gilroy-Gavilan Coll	R-O**	4.7	729	1.66	0.027
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.	Reverse	4.9	257	1.54	0.029
	N. Palm Springs	1986	6.06	6.8	Desert Hot Springs	R-O**	7.1	345	2.03	0.029
	Loma Prieta	1989	6.93	14.7	Santa Teresa Hills	R-O**	10	271	2.03	0.030
	Mammoth Lakes-01	1980	6.06	4	Convict Creek	N-O***	9.1	338	1.30	0.030
	N. Palm Springs	1986	6.06	11.2	North Palm Springs	R-O**	5.6	345	1.19	0.031

Table 4. Summary of the Earthquake Time History Recommendations from the NGA Database with DSHA Calculations (Continued-I)

City	Earthquake motion	Year	Magnitude	Rupture Distance (km)	Station	Fault Mechanism	D ₅₋₉₅ (s)	V _{s30} (m/s)	Scale Factor	MSE
Taichung	Mammoth Lakes-01	1980	6.06	6.6	Convict Creek	N-O**	9.1	338	1.69	0.030
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.-FP	Reverse	4.9	257	1.96	0.034
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.-FN	Reverse	5.0	257	1.99	0.040
	Northridge-01	1994	6.69	28.3	LA – Centinela St.	Reverse	11.9	235	1.99	0.041
	N. Palm Springs	1986	6.06	16.1	North Palm Springs	R-O**	5.6	345	1.51	0.041
	Northridge-01	1994	6.69	22.5	LA-UCLA	Reverse	9.4	398	2.00	0.041
	N. Palm Springs	1986	6.06	6.0	Whitewater Trout Farm	R-O**	25.8	345	1.70	0.043
	Loma Prieta***	1989	6.93	12.8	Gilroy Array #3	R-O**	7.7	349	1.63	0.045
Chiayi	Coalinga-05	1983	5.77	2.7	Pleasant Valley P.P	Reverse	4.9	257	1.92	0.032
	Mammoth Lakes-01	1980	6.06	6.6	Convict Creek	N-O**	9.1	338	1.66	0.034
	N. Palm Springs	1986	6.06	6.0	Whitewater Trout Farm	R-O**	5.1	345	1.67	0.039
	Whittier Narrows-01	1994	6.69	28.3	LA – Obregon Park	R-O**	7.8	349	2.00	0.040
	Loma Prieta	1989	6.93	17.5	WAHO	R-O**	11.1	376	1.30	0.045
	Coalinga-05***	1983	5.77	8.5	Oil City	Reverse	2.8	376	1.03	0.046
	Imperial Valley-06	1979	6.53	2.7	Bonds Corner	Reverse	9.7	223	1.04	0.046
	Chi-Chi-03	1989	6.2	7.6	TCU078	Reverse	6.7	443	1.66	0.046
Hualien	Mammoth Lakes-01	1980	6.06	6.6	Convict Creek	N-O**	9.1	338	2.01	0.035
	N. Palm Springs	1986	6.06	6.0	Whitewater Trout Farm	R-O**	5.1	345	2.00	0.035
	Imperial Valley-06	1979	6.53	2.7	Bonds Corner	Reverse	9.7	223	1.26	0.036
	Chi-Chi-03	1989	6.2	7.6	TCU078	Reverse	6.7	443	2.00	0.036
	Coalinga-05***	1983	5.77	8.5	Oil City	Reverse	2.8	376	1.24	0.038
	Superstition Hills-02	1987	6.54	5.6	Superstition Camera	Strike-Slip	12.1	362	1.53	0.038
	Loma Prieta***	1989	6.93	12.8	Gilroy Array #3	R-O**	7.7	349	1.92	0.040
	Chi-Chi-06	1989	6.3	10.1	TCU079	Reverse	4.0	443	1.28	0.040

Table 4. Summary of the Earthquake Time History Recommendations from the NGA Database with DSHA Calculations (Continued-II)

City	Earthquake motion	Year	Magnitude	Rupture Distance (km)	Station	Fault Mechanism	D ₅₋₉₅ (s)	V _{s30} (m/s)	Scale Factor	MSE
Pingtung	Imperial Valley-06	1979	6.53	0.3	Aeropuerto Mexicali	Strike-Slip	7.1	274	2.02	0.023
	Imperial Valley-06	1979	6.53	3.9	EL Centro Array #8	Strike-Slip	5.8	206	1.38	0.028
	Coalinga-01	1983	6.36	8.4	Pleasant Valley P.P.	Reverse	8.0	257	1.30	0.029
	Coalinga-05	1983	5.77	16.1	Pleasant Valley P.P.	Reverse	5.0	257	1.50	0.032
	Coalinga-05	1983	5.77	23.5	Bonds Corner	Reverse	5.0	257	1.26	0.033
	Northridge-01	1994	6.69	15.6	Tarzana-Cedar Hill A	Reverse	10.3	257	2.00	0.034
	Chi-Chi	1999	7.62	10.0	CHY101	R-O**	29.0	258	1.59	0.035
	Loma Prieta	1989	6.93	15.2	Capitola	R-O**	14.7	288	1.43	0.035

1. * M. - N. = Managua – Nicaragua
2. R-O** = Reverse – oblique
3. N-O** = Normal – oblique
4. *** refers to pulse-like record