

1 **Title**

2

3 Adapting the EDuMaP method to test the performance of the Norwegian early
4 warning system for weather-induced landslides

5

6 **Authors**

7

8 Piciullo Luca⁽¹⁾, Dahl Mads-Peter⁽²⁾, Devoli Graziella^(2,3), Colleuille Hervé⁽²⁾, Calvello Michele⁽¹⁾

9 ⁽¹⁾ Department of Civil Engineering, University of Salerno, Italy

10 ⁽²⁾ Norwegian Water Resources and Energy Directorate, Oslo, Norway

11 ⁽³⁾ Department of Geosciences, University of Oslo, Oslo, Norway

12

13

14 **Abstract**

15 The Norwegian national landslide early warning system (LEWS), operational since 2013, is
16 managed by the Norwegian Water Resources and Energy Directorate and has been designed for
17 monitoring and forecasting the hydro-meteorological conditions potentially triggering slope
18 failures. Decision-making in the EWS is based upon rainfall thresholds, hydro-meteorological and
19 real-time landslide observations as well as on landslide inventory and susceptibility maps. Daily
20 alerts are issued throughout the country considering variable size warning zones. Warnings are
21 issued once per day for the following 3 days and can be updated according to weather forecasts and
22 information gathered by the monitoring network. The performance of the LEWS operational in
23 Norway has been evaluated applying the EDuMaP method, which is based on the computation of a
24 duration matrix relating number of landslides and warning levels issued in a warning zone. In the
25 past, this method has been exclusively employed to analyse the performance of regional early
26 warning model considering fixed warning zones. Herein, an original approach is proposed for the
27 computation of the elements of the duration matrix in the case of early warning models issuing
28 alerts on variable size areas. The approach has been used to evaluate the warnings issued in Western
29 Norway, in the period 2013-2014, considering two datasets of landslides. The results indicate that
30 the landslide datasets do not significantly influence the performance evaluation, although a slightly
31 better performance is registered for the smallest dataset. Different performance results are observed
32 as a function of the values adopted for one of the most important input parameters of EDuMaP, the
33 landslide density criterion (i.e. setting the thresholds to differentiate among classes of landslide
34 events). To investigate this issue, a parametric analysis has been conducted; the results of the
35 analysis show significant differences among computed performances when absolute or relative
36 landslide density criteria are considered.

37 **Keywords:** EDuMaP method, rainfall-induced landslides, warning zones, alert, landslide density.

38 **1. Introduction**

39 In the last decades, natural hazards caused an increased number of consequences in terms of
40 economic losses (Barredo, 2009) and fatalities throughout Europe (European Environment Agency,
41 2010; CRED, 2011). Most natural disasters are related to extreme rainfall events, which are
42 expected to increase with climate change (Easterling et al., 2000; Morss et al., 2011). The European
43 Commission, following an increase in human and economic losses due to natural hazards,
44 developed legal frameworks such as the Water Framework Directive 2000/60/EC (2000) and the
45 Floods Directive 2007/60/EC (2007), to increase prevention, preparedness, protection and response
46 to such events and to promote research and acceptance of risk prevention measures within the
47 society (Alfieri et al., 2012). Among the many mitigation measures available for reducing the risk to
48 life related to natural hazards, early warning systems (EWSs) constitute a significant option
49 available to authorities in charge of risk management and governance.

50 Within the landslide risk management framework proposed by Fell et al. (2005), landslide EWSs
51 (LEWSs) may be considered a non-structural passive mitigation option to be employed in areas
52 where risk, occasionally, rises above previously defined acceptability levels. According to Glade
53 and Nadim (2014), the installation of an EWS is often a cost-effective risk mitigation measure and
54 in some instances the only suitable option for sustainable management of disaster risks. Rainfall-
55 induced warning systems for landslides are, by far, the most diffuse class of landslide EWS
56 operating around the world. LEWSs can be employed at two distinct scales of analysis: “local” and
57 “regional” (ICG 2012; Thiebes et al. 2012; Calvello et al. 2015, Stähli et al., 2015). EWSs at a
58 regional scale for rainfall-induced landslides have become a sustainable risk management approach
59 worldwide to assess the probability of occurrence of landslides over appropriately-defined wide
60 warning zones. In fact during the last decades, several systems have been designed and improved,
61 not only in developing countries (UNISDR 2006; Chen et al., 2007; Huggel et al., 2010; among
62 others) but also in developed countries (NOAA-USGS, 2005; Badoux et al., 2009; Baum and Godt,
63 2010; Osanai et al., 2010; Lagomarsino et al., 2013; Tiranti and Rabuffetti, 2010; Rossi et al., 2012;
64 Staley et al., 2013; Calvello et al., 2015; Segoni et al., 2015). As a recent example, the Norwegian
65 landslide EWS was launched in autumn 2013 by the Norwegian Water Resources and Energy
66 Directorate (NVE). The regional system has been developed for monitoring and forecasting the
67 hydro-meteorological conditions triggering landslides and to inform local emergency authorities in
68 advance about the occurrence of possible events (Devoli et al., 2014). Daily alerts are issued
69 throughout the country in variable size warning zones. The evaluation of the alerts issued, i.e., the

70 performance of the early warning model is not a trivial issue, and regular system testing and
71 performance assessments (Hyogo Framework for Action, 2005) are fundamental steps.
72 The performance analysis of LEWSs can be an awkward process, particularly for systems employed
73 at regional scale, because many aspects are important for the analyst to consider. Most typically,
74 the performance evaluation is based on 2 by 2 confusion matrices computed for the joint frequency
75 distribution of landslides and alerts, both considered as dichotomous variables, and the evaluation
76 of statistical indicators (e.g., Cheung et al., 2006; Godt et al., 2006; Martelloni et al., 2012; Staley et
77 al., 2013; Segoni et al., 2014; Lagomarsino et al., 2015; Gariano et al., 2015; Stähli et al., 2015).
78 The method employed herein, which is called EDuMaP (Calvello and Piciullo, 2016), allows to
79 consider aspects peculiar to territorial LEWSs that are not considered by the joint frequency
80 distribution approach. In particular, the EDuMaP method takes into account: the occurrence of
81 concurrent multiple landslides in the warning zone; the duration of the warnings in relation to the
82 landslides; the issued warning level in relation to the landslide spatial density in the warning zone;
83 the relative importance attributed, by system managers, to different types of errors. Up to now, this
84 method has been applied exclusively to evaluate the performance of regional warning models
85 designed for issuing alerts in fixed warning zones (Calvello and Piciullo, 2016; Piciullo et al.,
86 2016a,b; Calvello et al., 2016). In the present study the EDuMaP method has been adapted to
87 evaluate the performance of the alerts issued for variable size warning zones. To this purpose, the
88 procedure has been tested on the Norwegian landslide EWS in the period 2013-2014. The Western
89 Norway is the area most prone to landslides in Norway and it has been chosen as test area because
90 the landslide database was more reliable and complete than for the rest of Norway.
91

92 **2. The national landslide early warning system for rainfall-and snowmelt-** 93 **induced landslides in Norway**

94 **2.1 Physical setting**

95 Norway covers an area of $\sim 324,000 \text{ km}^2$. With its elongated shape of 1800 km, the country reaches
96 from latitude 58°N to 71°N. Approximately 30% of the land area are mountainous, with the highest
97 peaks reaching up to 2500 m. a.s.l and slope angles over 30 degrees covering 6,7% of the country
98 (Jaedicke *et al.*, 2009). In geological terms, Norway is located along the western margin of the
99 Baltic shield with a cover of Caledonian nappes in the western parts of the country (Etzelmüller *et*
100 *al.*, 2007; Ramberg *et al.*, 2008). The Caledonian nappes are dominated by Precambrian rocks and
101 metamorphic Cambro-Silurian sediments, while the bedrock in the Baltic shield is dominated by

102 Precambrian basement rocks. Cambro-Silurian sediments and Permian volcanic rocks are found in
103 the Oslo Graben (Ramberg *et al.*, 2008).

104 Recurrent glaciations, variations in sea level and land subsidence/uplift, as well as weathering,
105 transport and deposition processes have created the modern Norwegian landscape (Ramberg *et al.*,
106 2008). Thus, dominating quaternary deposits include various shallow (in places colluvial) soils, as
107 well as moraine and marine deposits.

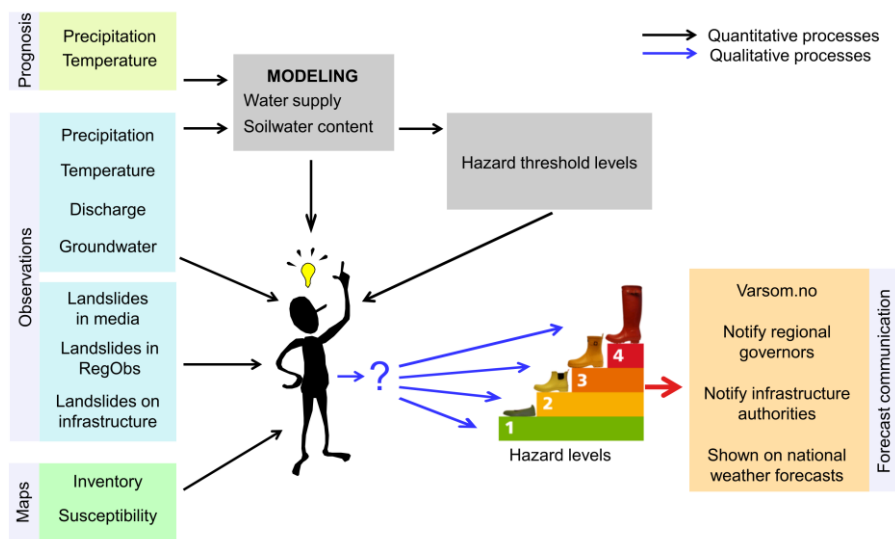
108 Because of the latitudinal elongation and the varied topography, the Norwegian climate displays
109 large variations. Along the Atlantic coast, the North Atlantic Current influences the climate whereas
110 the inland areas experiences a more continental climate. Based on the Köppen classification
111 scheme, the Norwegian climate can be classified in three main types: warm temperate humid
112 climate, cold temperate humid climate and polar climate. Precipitation types can be divided into
113 three categories: frontal, orographic and showery. The largest annual precipitation values are found
114 near the coast of Western Norway (herein also called Vestlandet) with up to 3575 mm/year. On the
115 other hand, the driest areas receiving <500 mm/year are found in parts of South-Eastern Norway
116 (Østlandet) and Finnmark county.

117 Steep landforms in combination with various soil and climatic properties provide a basis for several
118 types of shallow landslides in non-rock materials. These slope failures include slides in various
119 materials, debris avalanches, debris flows and slush flows. Landslides are mostly triggered by
120 rainfall, often in combination with snowmelt. Some events are also triggered from/initiated as
121 rockfall or slush flows, developing into, for example, debris flows as they propagate downslope.
122 Shallow landslides constitute a substantial threat to the Norwegian society. According to Furseth
123 (2006), at least 230 people have been killed by such slope failures during the latest approximately
124 500 years. In the period 2000-2009, road authorities registered more than 1800 shallow landslides
125 along Norwegian roads.

126 2.2 The national landslide early warning system (LEWS)

127 In order to mitigate the risk from shallow landslides, a national EWS has been developed at the
128 Norwegian Water Resources and Energy Directorate (NVE) as part of the national responsibility on
129 landslide risk management. The system is established to warn about the hazard of debris flows,
130 debris slides, debris avalanches and slush flows at regional scale. The EWS, operative since 2013,
131 has been developed in cooperation with the Norwegian Meteorological Institute (MET), Norwegian
132 Public Road Administration (SVV) and the Norwegian National Rail Administration (JBV).

133



134
135 **Figure 1.** Organization of the landslide early warning system in Norway.
136

137 Decision-making in the EWS is based upon hazard threshold levels, hydro-meteorological and real-
138 time landslide observations as well as landslide inventory and susceptibility maps (**Fig. 1**). In the
139 development phase of the EWS, hazard threshold levels have been investigated through statistical
140 analyses of historical landslides and modelled hydro-meteorological parameters. Daily hydro-
141 meteorological conditions such as rainfall, snowmelt, runoff, soil saturation, groundwater level and
142 frost depth have been obtained from a distributed version of the hydrological HBV-model (Beldring
143 *et al.*, 2003).

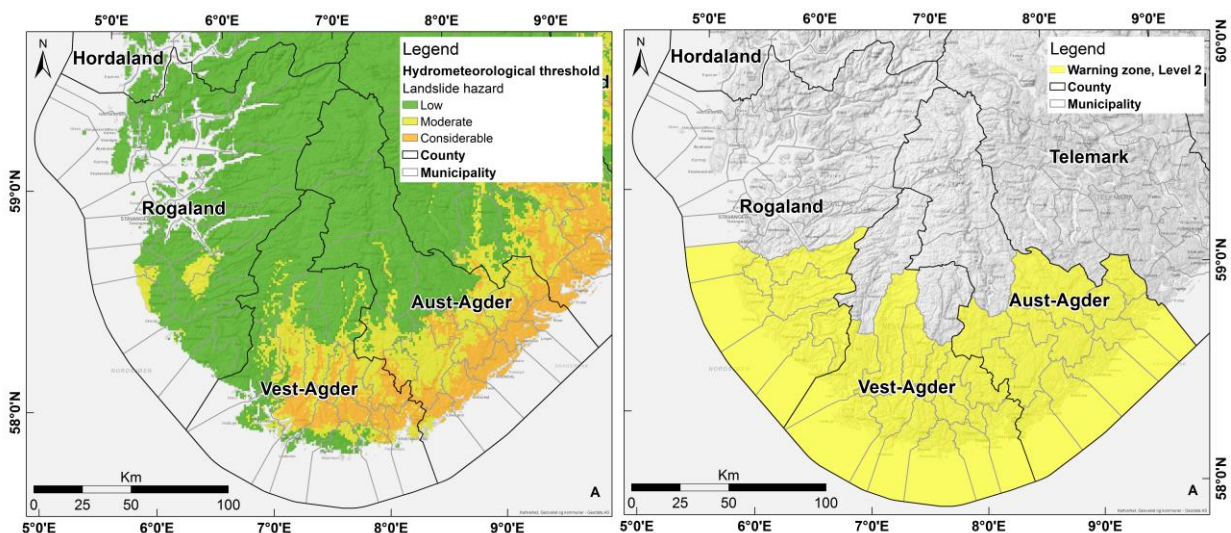
144 Hazard threshold levels presently used in the EWS were proposed by Colleuille *et al.* (2010). The
145 thresholds, combining simulations of relative water supply of rain or snowmelt and relative soil
146 saturation/groundwater conditions, were derived from empirical tree-classification using 206
147 landslide events from different parts of the country. Later analyses, summarized by Boje *et al.*
148 (2014), confirm the good performance of combining soil water saturation degree and normalised
149 rainfall and snowmelt.

150 Two different landslide susceptibility maps are used as supportive data in the process of setting
151 daily warning levels. One map indicates initiation and runout areas for debris flows at slope scale
152 (Fischer *et al.*, 2012), while another indicates susceptibility at catchment level, based upon
153 Generalized Additive Models (GAM) statistics (Bell *et al.*, 2014).

154 Susceptibility maps, hazard threshold levels and other relevant data are displayed in real-time in a
155 webpage, www.xgeo.no, which is used as decision expert tool to forecast various natural hazards
156 (floods, snow avalanches, landslides). Landslide hazard threshold levels and hydrometeorological
157 forecasts are displayed as raster data with 1 km² resolution, whereas susceptibility maps, landslide

158 information (historical and real-time) and hydrometeorological observations are shown as either
 159 raster, polygon or point data.

160 A landslide expert on duty (as member of a rotation team) uses the information from forecasts,
 161 observations, maps and uncertainty in weather forecasts to qualitatively perform a nationwide
 162 assessment of landslide warning levels (**Fig. 1**). Four warning levels are defined: green (1), yellow
 163 (2), orange (3), and red (4) showing the level of hazards, or more exactly the recommended
 164 awareness level (**Tab. 1**). The warning period follows the time steps of quantitative precipitation
 165 and temperature forecasts used to simulate other hydro-meteorological parameters, and thus lasts
 166 from 06:00 UTC to 06:00 UTC each day. Warning levels are updated minimum twice during the 24
 167 hour warning period (morning and afternoon) as a function of the weather forecast. Weather
 168 forecast updates are received 4 times per day and warning messages are sent as soon as possible,
 169 from 66 hours to few hours ahead. Warning messages are published in a publicly accessible
 170 webpage (www.varsom.no). Yellow, orange and red levels of warning are also sent to emergency
 171 authorities (regional administrative offices, roads and railways authorities) and media. Warning
 172 zones are not static geographical warning areas. Instead they vary from a small group of
 173 municipalities to several administrative regions, depending on current hydro-meteorological
 174 conditions (**Fig. 2**). Thus, extent and position of warning zones are dynamic and change from day to
 175 day.



176
 177 **Figure 2.** a) Hydrometeorological thresholds indicating potential landslide hazard in the counties of
 178 Rogaland, Vest-Agder, Aust-Agder and Telemark in South-Eastern Norway on 15.02.2014. b) The
 179 resultant early warning zone, on warning level 2 (“yellow level”) issued on 15.02.2014 for the same
 180 area and including about 32 municipalities.

181 **2.3 Current performance evaluation of the EWS**

182 To evaluate the performance of a regional landslide early warning model, a comparison of warning
183 levels issued and landslides occurred is carried out on a weekly basis. Event information is reported
184 by Roads/Railways Authorities or municipalities, as well as obtained from media and from a real-
185 time database to register observations. The latter has been designed as a public tool supporting
186 crowd sourcing (Ekker et al. 2013), and is currently available to the public as a telephone
187 application and a website (www.regobs.no). Categorization of issued warning levels into false
188 alarms, missed events, correct and wrong levels is based on semi-quantitative classification criteria
189 for each warning level. The principle behind the criteria is that rare hydro-meteorological conditions
190 are expected to cause more landslides and possibly higher damages (**Tab. 1**). As an example, the
191 warning level Red corresponds to an extreme situation that occurs very rarely. It requires immediate
192 action and may cause severe damages within a large extent of the warning area. The criteria contain
193 information on the expected number of landslides per area, as well as hazard signs indicating
194 landslide activity. As seen in **Table 1** the ranges chose for the number of expected landslides and
195 the size of the hazardous areas at each warning level are quite wide. This choice is due to the fact
196 that the EWS is relatively new and still in a phase of continuous development.

197

198 **Table 1.** Criteria for evaluating daily warning levels in the Norwegian EWS.

Warning level	Classification criteria
4 (Red)	> 14 landslide (per 10-15.000 km ²) Hazard signs: Several road blockings due to landslides or flooding Extreme situation that occurs very rarely, requires immediate action and may cause severe damages within a large extent of the warning area. This level corresponds to a >50 years return period flood warning.
3 (Orange)	6-10 landslides (per 10-15.000 km ²) Hazard signs: Several road blockings due to landslides or flooding Severe situation that occurs rarely, require contingency preparedness and may cause severe damages within some extent of the warning area. This level corresponds to 5-50 years return period flood warning.
2 (Yellow)	1-4 landslides (per 10-15.000 km ²) Hazard signs: flooding/erosion in streams Situation that requires monitoring and may cause local damages within the warning area. Expected some landslide events, certain large events may occur.
1 (Green)	No landslides 1-2 landslide caused by local rain showers 1 small debris slide if in area with no signs of elevated warning level Man-made events (from e.g. leakage, deposition, construction work or explosion)

199 **2.4 The EDuMaP method**

200 The paper proposes the evaluation of the performance of the landslide early warning system
201 operational in Norway by means of the “Event, Duration Matrix, Performance (EDuMaP) method”

202 (Calvello & Piciullo, 2016). This method has been principally employed to analyse the performance
203 of regional early warning model considering fixed warning zones for issuing alerts. The method
204 comprises three successive steps: identification and analysis of landslide and warning Events (E),
205 from available databases; definition and computation of a Duration Matrix (DuMa), and evaluation
206 of the early warning model Performance (P) by means of performance criteria and indicators.

207 The first step requires the availability of landslides and warnings databases for the preliminary
208 identification of “landslide events” (LEs) and “warning events” (WEs). A landslide event is defined
209 as one or more landslides grouped on the basis of their spatial and temporal characteristics. A
210 warning event is defined as a set of warning levels issued within a given warning zone, grouped
211 considering their temporal characteristics. The parameters which need to be defined to carry on the
212 events analysis are ten: 1) warning levels, W_{lev} ; 2) landslide density criterion, $L_{den(k)}$; 3) lead time,
213 t_{LEAD} ; 4) landslide typology, L_{typ} ; 5) minimum interval between landslide events, Δt_{LE} ; 6) over time,
214 t_{OVER} ; 7) area of analysis, A; 8) spatial discretization adopted for warnings, $\Delta A_{(k)}$; 9) time frame of
215 analysis, ΔT ; 10) temporal discretization of analysis, Δt . For more details see Calvello and Piciullo,
216 2016. The second step of the method is the definition and computation of a “duration matrix”,
217 whose elements d_{ij} , report the time associated with the occurrence of landslide events in relation to
218 the occurrence of warning events, in their respective classes. The element d_{11} of the matrix
219 expresses the number of hours when no warnings are issued and no landslides occur (Fig. 4). The
220 number of rows and columns of the matrix is equal to the number of classes defined for the warning
221 and landslide events, respectively (**Fig. 3**). The final step of the method is the evaluation of the
222 duration matrix based on a set of performance criteria assigning a performance meaning to the
223 element of the matrix. Two criteria are used for the following analyses (**Fig. 3**), respectively
224 indicated as criterion 1 and criterion 2. The first criterion employs an alert classification scheme
225 derived from a 2x2 contingency table, thus identifying: correct predictions, CPs; false alerts, FAs;
226 missed alerts, MAs; true negatives, TNs. The second criterion assigns a color code to the elements
227 of the matrix in relation to their grade of correctness, classified in four classes as follows: green, G,
228 for the elements which are assumed to be representative of the best model response; yellow, Y, for
229 elements representative of minor model errors; red, R, for elements representative of a significant
230 model errors; purple, P, for elements representative of the worst model errors. Both criteria
231 purposefully neglect element d_{11} , whose value is typically orders of magnitude higher than the
232 values of the other elements of the matrix because it also includes all hours without rainfall, for
233 which a LEWS is not designed to deal with, specifically. Thus, d_{11} element is neglected in order to
234 avoid an overestimation of the performance and to allow a more useful relative assessment of the

235 information located in the remaining part of the duration matrix. A number of performance
 236 indicators may be derived from the two performance criteria described. **Table 2** reports the name,
 237 symbol, formula and value of the performance indicators considered herein.

1) Alert classification criterion		Landslide events			
		no	S	I	L
Warning events	no	TN	CP	MA	MA
	M	CP	CP	MA	MA
	H	FA	CP	CP	CP
	VH	FA	FA	CP	CP

2) Grade of correctness criterion		Landslide events			
		no	S	I	L
Warning events	no	G	Y	R	P
	M	Y	G	R	P
	H	R	R	G	Y
	VH	P	P	Y	G

238
 239 **Figure 3.** Performance criteria used for the analyses performed herein (modified from Calvello &
 240 Piciullo, 2016). Four classes of warning events (key: no, no warning; M, moderate warning; H, high
 241 warning; VH, very high warning) and four classes of landslide events (key: no, no landslides; S,
 242 small event, few landslides; I, intermediate event, several landslides; L, large events, many
 243 landslides).

244
 245

Table 2. Performance indicators used for the analysis.

Performance indicator	Symbol	Formula
Efficiency index	I_{eff}	$CP/\sum_{ij}d_{ij}$ (excluding d_{11})
Hit rate	HR_L	$CP/(CP+MA)$
Predictive power	PPW	$CP/(CP+FA)$
Threat score	TS	$CP/(CP+MA+FA)$
Odds ratio	OR	$CP/(MA+FA)$
Miss classification rate	MR	$1 - I_{eff}$
Missed alert rate	R_{MA}	$MA/(CP+MA)$
False alert rate	R_{FA}	$FA/(CP+FA)$
Error Rate	ER	$(Red\&Pur)/\sum_{ij}d_{ij}$ (excluding d_{11})
Missed and false alerts balance	MFB	$MA/(MA+FA)$
Probability of serious mistakes	P_{SM}	$Pur/\sum_{ij}d_{ij}$ (excluding d_{11})

246 2.5 Adaptation of the EDuMaP method

247 LEWSs may adopt a fixed or a variable spatial discretization for warnings ($\Delta A(k)$). In the first case
 248 the warning zones are univocally defined with fixed extents. For each warning zone, the warnings
 249 are issued over the whole zone according to site specific rainfall thresholds and decisional

250 algorithms. Thus, only one level of warning can be issued in each warning zone in the minimum
 251 temporal discretization adopted for warnings (Δt). The performance analysis with the EDuMaP
 252 method is carried out separately for each warning zone. Therefore, in this case, the d_{ij} components
 253 of the duration matrix represent the time evaluation of the combination of warning levels issued and
 254 landslide events occurred in a specific warning zone in a period of analysis.

255 In the case of a variable spatial discretization for warnings the number and extent of the warning
 256 zones vary in time in the period of analysis (ΔT). The number of warning zones is defined by the
 257 number of warning levels issued in the minimum temporal discretization (Δt). For instance, if only
 258 two levels (e.g. green and orange) are issued in a given Δt , the area of analysis (A) would be
 259 divided into two warning zones. The extent of the warning zones is obtained grouping together all
 260 the territorial units (TUs) alerted with the same level of warning (see **Fig. 4**). In a given Δt , the
 261 Event analysis phase is carried out for all the warning zones simultaneously. The time evaluation of
 262 the elements of the duration matrix in a given Δt ($time_{ij}$) for the area of analysis (A) is carried out
 263 by weighting the spatial contribution of each warning zone in relation to the total area. In particular,
 264 the values of $time_{ij}$, for variable size warning zones, are computed as follows:

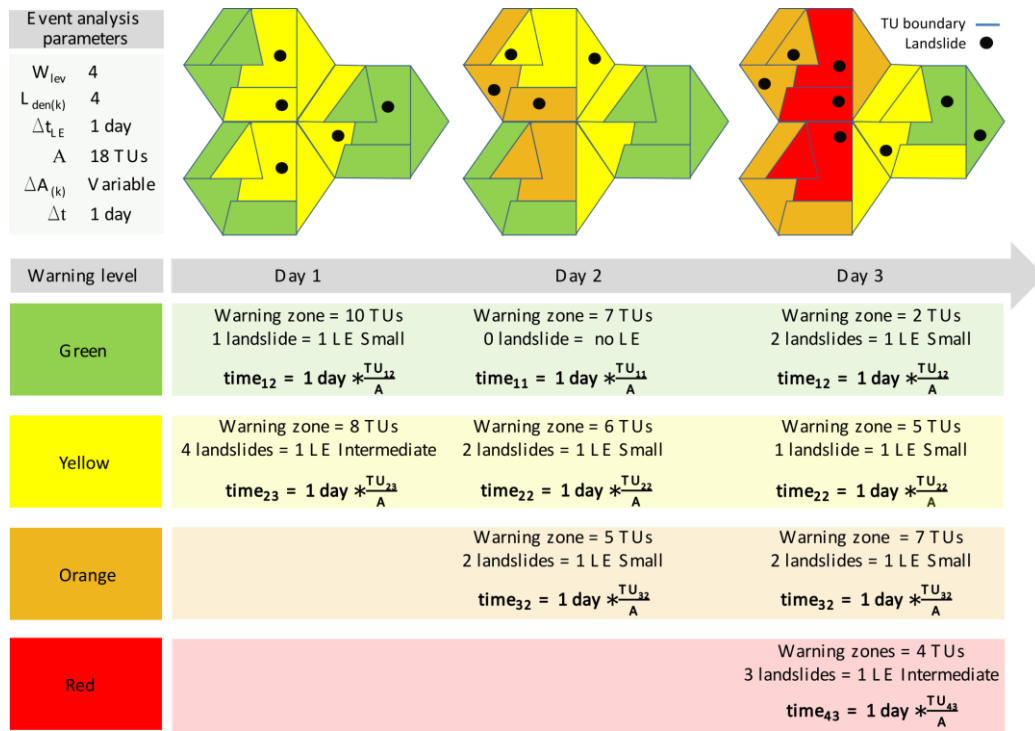
$$265 \quad time_{ij} = \Delta t \frac{(TU_{ij})}{A} \quad (\text{Eq. 1})$$

266 where: Δt is the minimum temporal discretization adopted for warnings (for the Norwegian EWS,
 267 equal to 1 day); A is the area of analysis; TU_{ij} is the extent of the territorial units alerted with a
 268 warning level i , and class of the landslide event, j , in a given Δt . Each element of the duration
 269 matrix, d_{ij} , is evaluated for the whole area of analysis, A , in a period of analysis, ΔT , summing the
 270 $time_{ij}$ computed within the different warning zones, for each temporal discretization Δt , as follows:

$$271 \quad d_{ij} = \sum_{\Delta T} (time_{ij}) \quad (\text{Eq. 2})$$

272 The evaluation of landslide and warning events and the definition and computation of a the duration
 273 matrix is herein exemplified for three hypothetical days (**Fig. 4**). For instance, on Day 1 two distinct
 274 LEs appear, containing 4 and 1 landslides, respectively. The first event belongs to the warning zone
 275 alerted with level 2 and the latter to the warning zone alerted with level 1. In Day 3 there are 4
 276 warning zones, each one alerted with a different warning level and 4 distinct LEs can be identified,
 277 one per warning zone. A landslide density criterion, $L_{den(k)}$ in four classes has been considered for
 278 the example of **Figure 4**: 0 (no landslides), small (1-2 landslides), Intermediate (3-4 landslides) and
 279 Large (≥ 5 landslides); together with four warning levels, W_{lev} : green, yellow, orange and red. At
 280 “day 1” two different warning zones can be defined grouping together the TUs (blue boundary in
 281 **Fig. 4**) with the same warning level. The warning zones are composed by 10 and 8 TUs, and they

282 are alerted with two different warning levels: green and yellow. In the two warning zones, a “small”
 283 LE and an “Intermediate” LE, respectively, are occurred. Once the warning levels and the LEs
 284 within each warning zone have been defined, $time_{12}$ and $time_{23}$ are evaluated for each TU using
 285 **Equation 1**. At “day 2” three warning zones and two “Small” LEs have been identified. At “day 3”
 286 LEs occurred in each of the four warning zones identified. Finally, the evaluation of elements d_{ij}
 287 of the duration matrix, is carried out following **Equation 2**, over the time frame of the analysis, ΔT .
 288



289 **Figure 4:** Computation of $time_{ij}$ elements as a function of warning levels and LEs occurred for each
 290 warning zone for three hypothetical days of warning.
 291

292

293 3. Performance evaluation of the LEWS in Western Norway for the period 294 2013-2014

295 3.1 Study area and landslide data

296 The study area includes the four administrative regions of Møre og Romsdal, Sogn og Fjordane,
 297 Hordaland and Rogaland located on the Norwegian west-coast. A common name for the entire area
 298 is Vestlandet (i.e. Western Norway). The area is dominated by narrow fjords and steep
 299 mountainsides reaching from sea level to 1000 m a.s.l. or more, and high annual precipitation of up
 300 to ~3500 mm. Shallow quaternary deposits cover locally weathered and altered bedrock of mainly

301 precambrian and Caledonian metamorphic and magmatic origin. As a result, Vestlandet is highly
302 prone to landslides, in particular, debris avalanches, debris flows and slush flows.

303 Vestlandet is the rainiest area of Norway with many annual precipitation events bringing high
304 amounts of rain and/or snow. Precipitation patterns and spatial distribution display large variations
305 within the study area. The precipitation patterns are described based on the main spatial
306 distribution:

- 307 a) NNW precipitation only in the region of Møre og Romsdal;
- 308 b) NW precipitation mainly in the regions of Møre og Romsdal and Sogn og Fjordane, or
309 sometimes in the northern part of Hordaland;
- 310 c) WNW precipitation in the entire study area;
- 311 d) W precipitation distributed mainly in Sogn og Fjordane, Hordaland and Rogaland;
- 312 e) SW precipitation distributed mainly in Rogaland and Hordaland, or sometimes also in Sogn
313 of Fjordane;
- 314 f) SSW precipitation only in Rogaland, or sometimes in Hordaland and rarely in the southern
315 part of Sogn og Fjordane;
- 316 g) S and SE with precipitation mainly in South-Eastern Norway (in summer) and not in the
317 study area, however because of size of the systems, precipitation can spread to Møre og
318 Romsdal or to eastern Sogn og Fjordane or Hordaland, depending on trajectory;
- 319 h) Local showers (mostly in summer), with clusters of maximum precipitation distributed
320 randomly within the study area;
- 321 i) Southern Norway, with precipitation distributed in the entire southern part of the country
322 and consequently in the entire study area.

323 During the years 2013 and 2014 more than 70 precipitation events, i.e. rain and/or snow records
324 with more than 30 mm/24h, were registered, with some episodes bringing more than 75-150
325 mm/24h of rain/snow to the entire study area or part of it, following the patterns indicated above.
326 Duration of precipitation events ranged from 1 day to 14-18 consecutive days, particularly during
327 autumn.

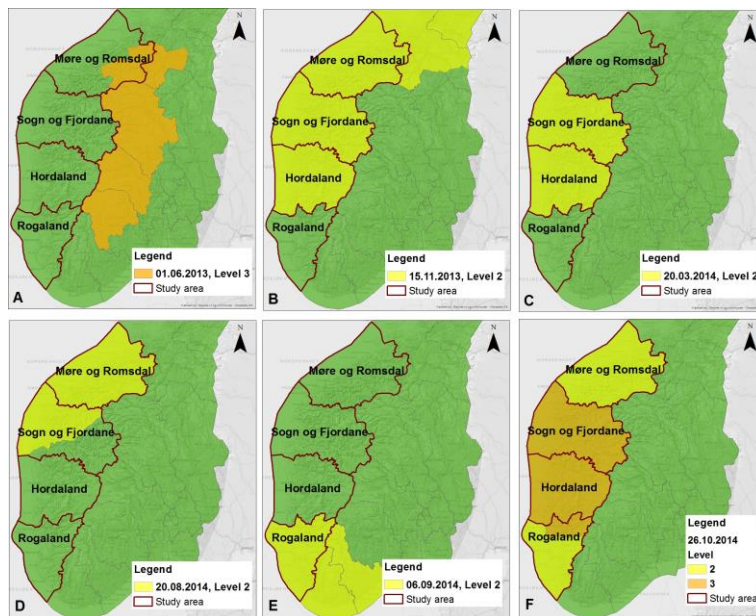
328 Landslide early warnings higher than green level were issued for 49 days during the two-year
329 period (**Tab. 3**). Most of these were at yellow level, however five warnings at orange level were
330 issued in 2014 in 3 consecutive days. In 12 cases, the yellow warnings issued during the morning
331 evaluation was downgraded to green later the same day. The most significant precipitation events
332 recorded in 2013-2014 are 11 and occurred in the following days: 14-15/04/13, 12-13/08/13,
333 7/10/13, 22/10/13, 15/11/13, 28/12/13, 23/02/14, 20/03/14, 14/07/14, 18-19/08/14, 27-28/10/14.

334

335 **Table 3.** Significant rainfall, number of days with at least one warning, number of warnings and
 336 landslides in the period 2013-2014.

	2013	2014	tot
Precipitation events, i.e. rainfall and/or snow > 30 mm/24h	41	32	73
Number of days with at least one warning	20	29	49
Number of warnings	21	39	60
red warnings	0	0	
orange warnings	0	5	
yellow warnings	21	34	
Number of landslides	204	181	385

337
 338
 339 Examples of warnings issued during 2013 and 2014 are shown in **Figure 5**. Most of the alerted
 340 warning zones were completely included in the study area (**Fig. 5 c, d, f**). However, some warnings
 341 were mainly issued for neighboring areas to the 4 regions chosen as case study (**Fig. 5 a, b, e**). The
 342 examples of **Figure 5** also illustrates the diversity in having variable instead of fixed size warning
 343 zones.
 344



345
 346 **Figure 5.** Examples of early warning areas and levels during 2013-2014.
 347

348 Within the study area, for the period 2013-2014, the Norwegian national landslide database
 349 (www.skrednett.no) lists 476 landslides in soils and/or slush flows. Due to errors and double

350 registration, 385 of these slope failures were considered valid for the current analyses: 249 (65%)
351 are categorized as landslide in soil, not otherwise specified due to lack of further documentation; 65
352 (17%) are categorized as debris avalanches, following Hungr et al. (2014), in many cases initiated
353 as small debris slides; 27 (7%) are classified as debris flows, following Hungr et al. (2014); 20 (5%)
354 are soil slides in artificial slopes (cuts and fillings along roads or railway lines); 19 (4%) are slush
355 flows and the remaining 5 (1%) are rock falls developing into debris avalanches.

356 The EDuMaP method was applied to two different datasets of phenomena: Set A and Set B. The
357 first set includes all 385 slope failures, while the second included only 131 phenomena, as
358 “landslide in soil not specified” and “rock fall/debris avalanches” were removed from this dataset.
359 The removal of non-specified landslides was due to the questionable quality of these registrations in
360 the national landslide database, while the exclusion of rock falls inducing debris avalanches was
361 due to uncertainty on whether precipitation can indeed be considered their triggering cause.

362 3.2 Events analysis

363 In earlier studies, the EDuMaP method has been applied to analyse the performance of regional
364 landslide EWSs adopting a fixed spatial discretization for warnings. In contrast, the Norwegian
365 landslide EWS employs variable size warning zones. This characteristic influences the first two
366 phases of the EDuMaP method: identification and analysis of landslide and warning events from
367 available databases; definition and computation of a duration matrix.

368 The values of the ten input parameters, cf. section 3, for the two analyses carried out, i.e. case A and
369 case B, are representative of the structure and operational procedures of the warning model
370 employed in the Norwegian EWS. It adopts four warning levels: green (no warning), yellow (WL_1),
371 orange (WL_2), red (WL_3). Daily warnings are issued throughout the country (i.e., Δt , is set to 1 day)
372 considering municipalities as the minimum warning territorial unit (TU). Hence, municipalities
373 alerted with the same warning level define a warning zone of level i . Therefore, on a day of alert, up
374 to four warning zones alerted with different warning levels can be issued (e.g., day 3 in **Fig. 4**).
375 Parameters t_{LEAD} and t_{OVER} are both set to zero. LEs are defined by grouping together landslides
376 occurred within each warning zone considering a Δt_{LE} of 1 day. The four classes of LEs are defined
377 employing a relative landslide density criterion, $L_{den(k)}$, as a function of both number of landslides
378 and territorial extensions. The values have been derived by the criteria for the daily warning levels
379 evaluation in the Norwegian EWS (see **Tab. 1**).

380 The only difference between case A and case B has to do with the type of landslides used for the
381 analyses, which respectively refer to the datasets A and B. Dataset A is composed by 385 rainfall-

382 and snowmelt-induced landslides occurring within the study area. These slope failures have been
 383 grouped into 137 LEs. The majority of LEs belong to class “Small” (133 events), while the rest of
 384 them (4 events) belong to class “Intermediate”; no “Large” LEs have been recorded in the period of
 385 analyses (**Tab. 4**). For case B, the 131 considered phenomena have been grouped into 57 LEs, 54
 386 “Small” and 3 “Intermediate” events (**Tab.4**). A total of 60 warnings were issued in the period of
 387 analysis; none of these were “Red”. Five warning zones received the level “Orange” and 55 zones
 388 received the warning level “Yellow”. In the period of analysis 37 different warning zones have been
 389 alerted (**Tab. 4**).

390

391 **Table 4:** Number of landslides, landslides, warning events issued and warning zones alerted in
 392 2013-2014 in the area of analysis.

	Case A	Case B
Landslide	385	131
Landslide events, LE	137	57
Small	132	54
Intermediate	5	3
Large	0	0
Warning events, WE	60	60
Warning zones alerted	37	37

393

394 3.3 Duration matrices and performance indicators

395 Two different sets of landslides have been considered in the performance of the Norwegian EWS
 396 for the Vestlandet area: Set A and Set B. The duration matrices obtained are shown in **Table 5**.
 397 Both cases refer to the years 2013-2014, thus, the sum of matrix elements is always equal to 730
 398 days.

399

400 **Table 5:** Duration matrices for cases A and B, units of time expressed in days.

CASE A		LE class			
		1	2	3	4
WE level	1	600,48	107,62	0,00	0,00
	2	9,88	8,47	1,80	0,00
	3	0,00	1,16	0,58	0,00
	4	0,00	0,00	0,00	0,00

CASE B		LE class			
--------	--	----------	--	--	--

401

		1	2	3	4
WE level	1	671,55	36,56	0,00	0,00
	2	11,32	7,90	0,93	0,00
	3	1,16	0,00	0,58	0,00
	4	0,00	0,00	0,00	0,00

402

403

404

405

406

407

408

409

410

411

412

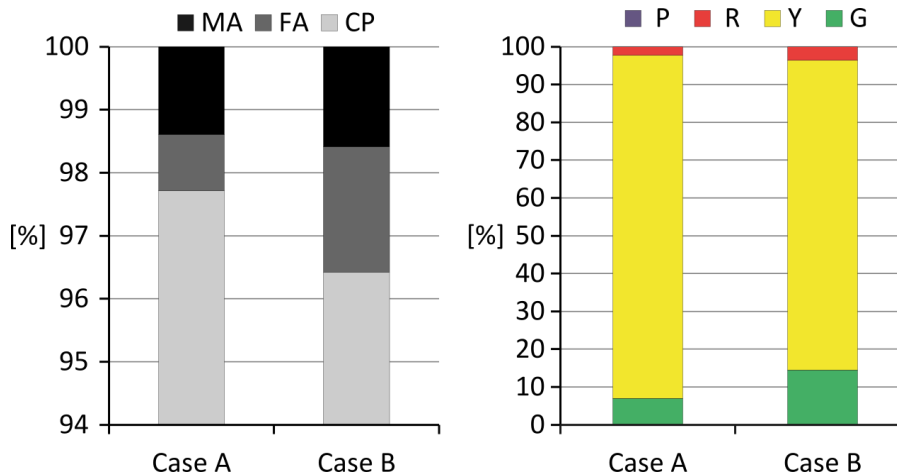
413

414

415

416

The duration matrices have been analysed considering two different performance criteria (see **Fig. 4**). The first one is derived by a contingency table scheme (criterion 1), the other one is based on a colour code assigning a grade of correctness to each matrix cell (criterion 2). The results obtained considering criterion 1 for both Case A and B (**Fig. 6 a**) show a very high percentage of correct predictions (CPs), over 96%, and around 1,5% of missed alerts (MAs). The amount of false alerts (FAs) are 1% and 2% respectively for Case A and B. Following criterion 2 (**Fig. 6 b**) differences, among Case A and B, can be observed in terms of greens (G), that are respectively equal to 7% and 14,5%, and yellows (Y) that are respectively equal to 91% and 82%. No P and just few R, equal to 2,3% and 3,6%, are observed in Case A and Case B, respectively. Following criterion 1, the differences among the two cases analysed are not significant. In terms of criterion 2, Case B shows slightly higher values of G (14%) than Case A (7%). This means that considering the reduced set of landslides (Set b), there is a slightly better correspondence between the LE classes and the corresponding warning levels issued.



417

418

Figure 6: Duration matrix results in terms of: criterion 1 (a) and criterion 2 (b).

419

420

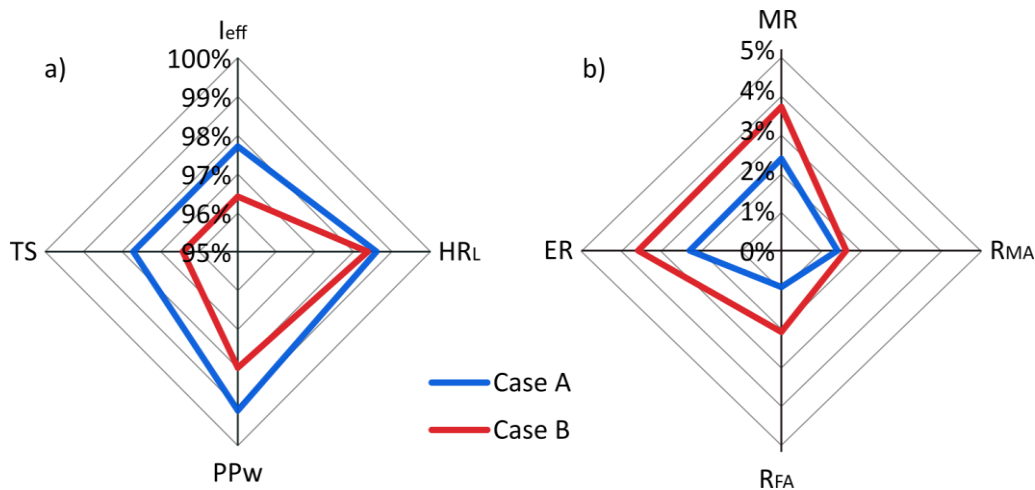
421

422

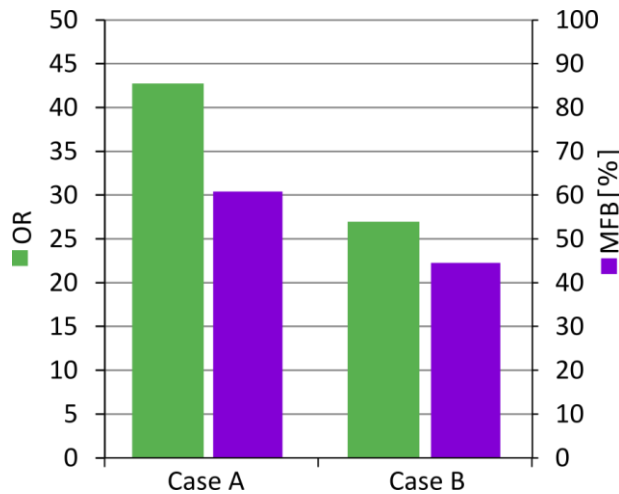
423

The performance indicators used to analyse the duration matrices (**Tab. 3**) are grouped into two subsets of indicators, respectively evaluating success and error (**Fig. 7**). Excluding the odds rate (OR), the remaining success indicators have a percentage higher than 95% for both cases, due to the high value of CPs that is orders of magnitude higher than MAs and FAs. Therefore the OR, that

424 indicates the correct predictions relative to the incorrect ones, assumes a very high value for both
 425 cases, although slightly higher for Case A (Fig. 8). The error indicators MR, ER, RMA and RFA
 426 assume very low values and the differences between the two cases are around 1% (Fig. 7 b). The
 427 MFB, which represents the ratio of MAs over the sum of MAs and FAs, is around 60% and 45%
 428 respectively for Cases A and B (Fig. 8).
 429



430 **Figure 7:** Performance indicators quantifying the landslide early warning performance of Case A
 431 (in blu) and Case B (in red) in terms of: success (a) and error (b).
 432



434 **Figure 8:** Odds Ratio (OR) and Missed and False alerts Balance (MFB) performance indicators,
 435 quantifying the landslide early warning performance of Case A and Case B.
 436

437 In this performance analysis the high value of I_{eff} , (>95%) and ORs, could be interpreted as an
 438 excellent result but, in contrast, the high value of MFB highlights some issues related to the
 439 duration of MAs in relation to the total duration of wrong predictions. In general, this could be a
 440 serious problem because MAs mean that no warnings or low level warnings have been issued
 441 during the occurrence of one or more LEs of the highest two classes (“Intermediate” and “Large”).
 442

443 In particular for Case A, 4 out of 5 LE of class “Intermediate” have to be considered MAs because
444 they occurred when the warning was set to level 2. Following the previous considerations, Case B
445 shows the best performance in terms of both success and error indicators, with a lower value of
446 MFB and a high value of OR. Case B uses a landslide dataset composed of rainfall-induced
447 landslides with a higher accuracy of information than Case A. As stated in Piciullo et al., (2016),
448 the result of a performance evaluation is strictly connected to the availability of a landslide
449 catalogue and to the accuracy of the information included in it.

450 Finally, it is important to stress the use of both success and error indicators to carry out a complete
451 performance analysis. As in this case, dealing with some indicators neglecting others could cause a
452 wrong evaluation of the early warning model performance. For instance, in the period of analysis,
453 no LEs of class 4 and only few LEs of class 3, occurred. However, the majority of durations of
454 these LEs have been missed. This means that the landslide early warning model was mostly able to
455 predict LEs of class “Small”. A possible solution to obtain a better model performance, reducing
456 MAs and simultaneously increasing CPs and G, could be to decrease the thresholds employed to
457 issue the warning level “High”.

458 3.4 Parametric analysis: the landslide density criterion

459 A parametric analysis on the landslide density criterion, $L_{den(k)}$, has been herein conducted with a
460 twofold purpose: to compare the performance of different early warning models, and to evaluate the
461 effect of the choices that the analyst makes when defining landslide event (LE) classes on the
462 performance indicators computed according to the EDuMaP method. The landslide density, $L_{den(k)}$,
463 represents the criterion used to differentiate among n classes of landslide events. The classes may be
464 established using an absolute (A) or a relative (R) criterion, i.e., simply setting a minimum and
465 maximum number of landslides for each class or defining these numbers as landslide spatial
466 density, i.e. in terms of number of landslides per unit area. Six landslide density criteria have been
467 considered in the performed parametric analysis (**Tab. 6**) referring to the criteria used in the
468 Norwegian EWS (**Tab. 1**). Two of them employ an absolute criterion using different numbers of
469 landslides per LE class the other four simulations, obtained considering the relative criterion, vary
470 as a function of both number of landslides and territorial extensions (10.000 km² and 15.000 km²).
471 Changing the definition of LE classes, the duration matrix and the performance indicators vary
472 because of relocation of the d_{ij} elements. In particular the $time_{ij}$ element, which is the amount of
473 time for which a level i -th warning event is concomitant with a class j -th landslide event, may vary
474 the j -th index causing a movement of the element along the i -th row. The parametric analysis has

475 been performed using the landslide dataset A, which includes 385 landslides. **Table 7** reports the
 476 classification of the LEs in the 6 combination of landslide density criteria.

477

478 **Table 6.** Parametric analysis: landslide density criteria considered to classify the LEs.

LE class	Absolute criterion [No. of landslides] and number of LEs		Relative criterion [No. of landslides / Area] and number of LEs			
	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
0	0	1	0	0	0	0
SMALL	1 to 4	2 to 4	(1 to 4)/15'000 km ²	(1 to 4)/15'000 km ²	(1 to 4)/10'000 km ²	(1 to 4)/10'000 km ²
INTERMEDIATE	5 to 14	5 to 18	(5 to 14)/15'000 km ²	(5 to 10)/15'000 km ²	(5 to 14)/10'000 km ²	(5 to 10)/10'000 km ²
LARGE	> 14	> 18	> 14/15'000 km ²	> 10/15'000 km ²	> 14/10'000 km ²	> 10/10'000 km ²

479

480 **Table 7.** Classification of LEs for the 6 simulations reported in table 8.

LE class	Absolute criterion [No. of landslides] and number of LEs		Relative criterion [No. of landslides / Area] and number of LEs			
	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
SMALL	124	32	132	132	133	133
INTERMEDIATE	9	9	5	3	4	4
LARGE	4	4	0	2	0	0

481

482 As an example, the simulations R-15K_{0,10} and R-15K_{0,14} differ for the definition of both LE classes
 483 Large and Intermediate. By comparing the two respective duration matrices (**Tab. 8 a, b**) a
 484 movement of the durations from d₂₄ and d₃₄ to respectively d₂₃ and d₃₃ is evident. This behaviour is
 485 due to the increase of spatial density for LE class Large, in particular from 0,67 landslides per 1000
 486 km² to 0,93 landslides per 1000 km² (**Tab. 6**), which causes a relocation of time_{i4} along the rows.

487 **Table 8.** Duration matrix results for simulations R-15_{0,10}, R-15_{0,14}.

R-15K _{0,10}		LE duration (h)			
		1	2	3	4
WE duration (h)	1	600,48	107,62	0,00	0,00
	2	9,88	8,47	0,98	0,82
	3	0,00	1,16	0,00	0,58
	4	0,00	0,00	0,00	0,00

488

R-15K _{0,14}		LE duration (h)			
		1	2	3	4
WE	1	600,48	107,62	0,00	0,00

duration (h)	2	9,88	8,47	1,80	0,00
	3	0,00	1,16	0,58	0,00
	4	0,00	0,00	0,00	0,00

489

490 Changes within the duration matrix mean that the value of the performance indicators may change.
 491 **Table 9** presents a summary of performance indicators for all six simulations of the landslide
 492 density criteria used in the parametric analysis.

493

494 **Table 9.** Performance indicators for the six simulations of landslide density criteria considered in
 495 the parametric analysis.

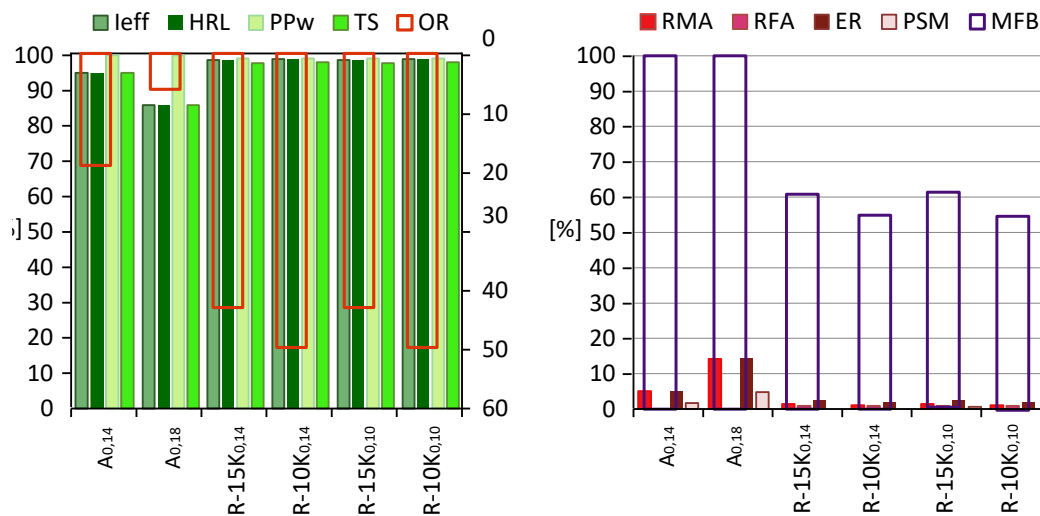
Performance indicator	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
I _{eff}	0,95	0,86	0,98	0,98	0,98	0,98
HR _L	0,95	0,86	0,99	0,99	0,99	0,99
PP _w	1,00	1,00	0,99	0,99	0,99	0,99
TS	0,95	0,86	0,98	0,98	0,98	0,98
OR	18,98	6,07	42,75	42,75	49,43	49,43
MR	0,05	0,14	0,02	0,02	0,02	0,02
R _{MA}	0,05	0,14	0,01	0,01	0,01	0,01
R _{FA}	0,00	0,00	0,01	0,01	0,01	0,01
ER	0,05	0,14	0,02	0,02	0,02	0,02
MFB	1,00	1,00	0,61	0,61	0,55	0,55

496

497 The results show similar performance for the four simulations derived using a relative criterion
 498 (R15-C_{0,14} R15-C_{0,10} R10-C_{0,14} R10-C_{0,10}). The values of the success indicators are always high:
 499 well above 95%, for I_{eff}, HR, TS, PP_w, while OR ranges between 42 and 49 (**Fig. 9 a**). This is due to
 500 the high value of CPs compared to those of MAs and FAs, underlining a good performance of the
 501 early warning model for these four simulations. In fact, also the error indicators are very low in
 502 terms of percentage, around 1-2% (**Fig. 9 b**). Lower values are observed for the combination
 503 obtained considering the absolute criterion, and in particular for A_{1,18}, with MR, R_{MA} and ER
 504 around 14%. The MFB is generally high for all simulations denoting a bad capability of the model
 505 to predict LEs of classes 3 and 4. Anyway, it must be emphasized that, considering these landslide
 506 density criteria, only the simulations R-15K_{0,10}, A_{0,14} and A_{1,18} have LEs of class 4 in the period of
 507 the analysis (**Tab. 7**).

508 In conclusion, the parametric analysis shows significant differences between the absolute and
 509 relative criterion simulations. For this case study, absolute criterion simulations have lower success
 510 performance indicators, in particular for the values of odds ratio (OR) and, very high values of
 511 missed and false alert balance (MFB) compared to the performance indicators obtained for relative
 512 criterion simulations. Moreover, the absolute criterion simulations produce a number of purple
 513 errors that increase the PSM (Fig. 9 b).

514



515

516

517

518

Figure 9. Performance indicators related to the success (a) and to the errors (b) of the warning model, evaluated for the six simulations of landslide density criteria considered in the parametric analysis.

519

520 4. Conclusions

521 The main aim of regional landslide early warning systems is to produce alert advices within a
 522 specific warning zone and to inform local authorities and the public of landslide hazard at a given
 523 level. To evaluate the performance of the alerts issued by such systems several aspects need to be
 524 considered, such as: the possible occurrence of multiple landslides in the warning zone, the duration
 525 of warnings in relation to the time of occurrence of landslides, the level of the issued warning in
 526 relation to spatial density of landslides in the warning zone and the relative importance system
 527 managers attribute to different types of errors. To solve these issues, the EDuMaP method can be
 528 seen as a useful tool for testing the performance of regional landslide warning models. Up to now,
 529 the method has been applied exclusively to systems that issue alerts on fixed warning zones. By
 530 using data from the Norwegian landslide EWS this study has extended the applicability of the
 531 EDuMaP method to warning systems that uses variable size warning zones. In this study, the

532 EDuMaP method has been used to evaluate the performance of the Norwegian landslide early
533 warning system for Vestlandet (Western Norway) for the period 2013-2014. The results show an
534 overall good performance of the system for the area analyzed. Two datasets of landslide
535 occurrences have been used in this study: the first one including all the slope failures registered and
536 gathered in the NVE database within the test area; the second one excluding the phenomena whose
537 typology was either not determined or is not typically associated to rainfall. The results are not too
538 sensitive to the dataset of landslides, although slightly better results are registered with the smallest
539 (i.e. more accurate) dataset. In both cases, the high value of the MFB highlights a high number of
540 MAs compared to the FAs. A recommendation could be to have a MFB lower than 25%, which
541 means that only 1 wrong alert out of 4 is a MA. Following this reasoning, a reduction of the
542 warning level “High” is recommended in order to reduce the MAs and to increase the performance
543 of the Norwegian EWS.

544 A parametric analysis was also conducted for evaluating the performance sensitivity, to the
545 landslide density criterion, $L_{den}(k)$, used as an input parameter with EDuMaP. This parameter
546 represents the way landslide events are differentiated in classes. In the analysis the classes were
547 established considering both absolute (2 simulations) and relative (4 simulations) criteria. The
548 parametric analysis shows how the variation of the intervals of the LE classes affects the model
549 performance. The best performance of the alerts issued in Western Norway was obtained applying a
550 relative density criterion for the definition of the LE classes. The parametric analysis shows only
551 minor differences in the performance analysis among the four cases considered with the relative
552 density criteria. In conclusion, this study highlights how the definition of the density criterion to be
553 used in defining the LE classes is a fundamental issue that system managers need to be take into
554 account in order to give an idea on the number of landslides expected for each warning level over a
555 given warning zone.

556 **Acknowledgement**

557 This work was carried out during a research period of LP as visiting PhD student at NVE, Oslo. The
558 authors are grateful to two NVE employees: Søren Boje for his criticism and comments and, Julio
559 Pereira for GIS data sharing.

560 **References**

- 561 Alfieri L, Salamon P, Pappenberger F, Wetterhall F, Thielen J: Operational early warning
562 systems for water-related hazards in Europe. *Environ Sci Pol* 15(1):35–49.
563 doi:10.1016/j.envsci.2012.01.008, 2012.
- 564 Barredo, J I: Normalised flood losses in Europe: 1970–2006. *Natural Hazards and Earth*
565 *System Sciences* 9, 97–104, doi:10.5194/nhess-9-97-2009, 2009.
- 566 Badoux A, Graf C, Rhyner J, Kuntner R, McArdell B W: A debris-flow alarm system for the
567 Alpine Illgraben catchment: design and performance. *Nat Hazards* (2009) 49:517–539, doi
568 10.1007/s11069-008-9303-x, 2009.
- 569 Baum R L and Godt J W: Early warning of rainfall-induced shallow landslides and debris
570 flows in the USA. *Landslides*, 7: 259–27, doi: 10.1007/s10346-009-0177-0, 2010.
- 571 Beldring S, Engeland K, Roald LA, Sælthun NR, Voksø, A: Estimation of parameters in a
572 distributed precipitation-runoff model for Norway. *Hydrology and earth system sciences*, 7: 304-
573 316, 2003.
- 574 Bell R, Cepeda J, Devoli G: Landslide susceptibility modeling at catchment level for
575 improvement of the landslide early warning system in Norway. *Conference Proceedings of the*
576 *World Landslide Forum*. Vol. 3. 2014. 3, 2-6th June 2014, Beijing, 2014.
- 577 Boje S, Colleuille H, Cepeda J, Devoli G : Landslide thresholds at regional scale for the
578 early warning system in Norway. *Conference Proceedings of the World Landslide Forum*. 3, 2-6th
579 June 2014, Beijing, 2014.
- 580 Calvello M and Piciullo L: Assessing the performance of regional landslide early warning
581 models: the EDuMaP method. *Natural Hazards Earth System Sciences*, 16, 103–122, 2016.
582 www.nat-hazards-earth-syst-sci.net/16/103/2016/doi:10.5194/nhess-16-103-2016, 2016.
- 583 Calvello M, d’Orsi RN, Piciullo L, Paes N, Magalhaes MA, Lacerda WA: The Rio de
584 Janeiro early warning system for rainfall-induced landslides: analysis of performance for the years
585 2010–2013. *Int J Disast Risk Reduc* 12:3–15. doi:10.1016/j.ijdrr.2014.10.005, 2015.
- 586 Chen C Y, Lin L Y, Yu F C, Lee C S, Tseng CC, Wang AH, Cheung KW: Improving debris
587 flow monitoring in Taiwan by using high-resolution rainfall products from QPESUMS. *Nat*
588 *Hazards*, 40: 447. doi:10.1007/s11069-006-9004-2, 2007.
- 589 Cheung, P. Y., Wong, M. C., and Yeung, H. Y.: Application of rainstorm nowcast to real-
590 time warning of landslide hazards in Hong Kong, in: WMO PWS, Workshop on Warnings of Real-
591 Time Hazards by Using Nowcasting Technology, 9–13 October 2006, Sydney, Australia, 2006.
- 592 Colleuille H, Haugen LE, Beldring S: A forecast analysis tool for extreme hydrological
593 conditions in Norway. Poster presentation, 6th World FRIEND Conference, Fez, Morocco, 2010.
- 594 CRED: EM-DAT. In: The OFDA/CRED International Disaster Database, Universite´
595 Catholique de Louvain, Brussels, Belgium. www.emdat.be, 2011.

596 Devoli G, Kleivane I, Sund M, Orthe N-K, Ekker R, Johnsen E, Colleuille H: Landslide
597 early warning system and web tools for real-time scenarios and for distribution of warning
598 messages in Norway, in: Engineering Geology for Society and Territory “Landslide Processes”,
599 Proc. XII International IAEG Congress, Torino, Italy, 625–629, doi:10.1007/978-3-319-09057-
600 3_104, 2014.

601 Easterling D R, Meehl G A, Parmesan C, Changnon S A, Karl T R, Mearns L O: Climate
602 extremes: observations, modeling, and impacts. *Science*, Vol. 289, Issue 5487, pp. 2068-2074, doi:
603 10.1126/science.289.5487.2068, 2000.

604 Ekker R, Kværne K, Os A, Humstad T, Warttinen A, Eide V, Hansen RK: regObs – public
605 database for submitting and sharing observations. Proceedings of International Snow Science
606 Workshop, 7-11th October 2013 Grenoble - Chamonix Mont-Blanc, 2013.

607 Etzelmüller B, Romstad B, Fjellanger J: Automatic regional classification of topography in
608 Norway. *Norwegian Journal of Geology*, 87: 167-180, 2007.

609 European Commission: Directive 2000/60/EC of the European Parliament and of the
610 Council of 23 October 2000 establishing a framework for Community action in the field of water
611 policy, 2000.

612 European Commission: Directive 2007/60/EC of the European Parliament and of the
613 Council of 23 October 2007 on the assessment and management of flood risks, 2007.

614 European Environment Agency: Mapping the impacts of natural hazards and technological
615 accidents in Europe – an overview of the last decade. EEA Technical Report No 13/2010, isbn: 978-
616 92-9213-168-5, 2010.

617 Fell R, Ho K K S, Lacasse S, and Leroi E: A framework for landslide risk assessment and
618 management, in: *Landslide Risk Management*, edited by: Hungr, O, Fell, R, Couture, R, and
619 Eberhardt, E, Taylor and Francis, London, 3–26, 2005.

620 Fischer L, Rubensdotter L, Stalsberg K, Melchiorre C, Horton P, Jaboyedoff M: Debris flow
621 modeling for susceptibility at regional to national scale in Norway. In Eberhardt et al. *Landslides
622 and Engineered Slopes: Protecting Society through Improved Understanding*, p. 723-729, 2012.

623 Furseth A: Slide accidents in Norway. Oslo, Tun Forlag, 207 p. (in Norwegian), 2006.

624 Gariano, S. L., Brunetti, M. T., Iovine, G., Melillo, M., Peruccacci, S., Terranova, O.,
625 Vennari, C., and Guzzetti, F.: Calibration and validation of rainfall thresholds for shallow landslide
626 forecasting in Sicily, southern Italy, *Geomorphology*, 228, 653–665,
627 <http://dx.doi.org/10.1016/j.geomorph.2014.10.019>, 2015.

628 Glade T and Nadim F: Early warning systems for natural hazards and risks, *Nat. Hazards*,
629 70, 1669–1671, doi:10.1007/s11069-013-1000-8, 2014.

630 Godt, J. W., Baum, R. L., and Chleborad, A. F.: Rainfall characteristics for shallow
631 landsliding in Seattle, Washington, USA, *Earth Surf. Proc. Land.*, 31, 97–110,
632 doi:10.1002/esp.1237, 2006.

- 633 Hyogo Framework for Action: Building the Resilience of Nations and Communities to
634 Disasters, World Conference on Disaster Reduction 18–22 January 2005, Kobe, Hyogo, Japan, 22
635 pp., 2005.
- 636 Huggel C, Khabarov N, Obersteiner M, Ramirez J M: Implementation and integrated
637 numerical modelling of a landslide early warning system: a pilot study in Colombia. *Nat Hazards*
638 (2010) 52:501–518. doi:10.1007/s11069-009-9393-0, 2010.
- 639 Hungr O, Leroueil S, Picarelli L: The Varnes classification of landslide types, an update.
640 *Landslides* 11(2):167–194. doi:10.1007/s10346-013-0436-y, 2014.
- 641 ICG: Guidelines for landslide monitoring and early warning systems in Europe - Design and
642 required technology. Project Safe Land "Living with landslide risk in Europe: Assessment, effects
643 of global change, and risk management strategies". D4.8, 153p. Available from:
644 <http://www.safeland-fp7.eu/results/Documents/D4.8.pdf>, 2012.
- 645 Jaedicke C, Lied K, Kronholm K: Integrated database for rapid mass movements in Norway.
646 *Natural Hazards and Earth System Sciences*, 9: 469-479, doi:10.5194/nhess-9-469-2009, 2009.
- 647 Lagomarsino, D, Segoni, S, Fanti, R, and Catani, F: Updating and tuning a regional-scale
648 landslide early warning system. *Landslides*, 10, 91–97. doi:10.1007/s10346-012-0376-y, 2013.
- 649 Lagomarsino, D., Segoni, S., Rosi, A., Rossi, G., Battistini, A., Catani, F., and Casagli, N.:
650 Quantitative comparison between two different methodologies to define rainfall thresholds for
651 landslide forecasting, *Nat. Hazards Earth Syst. Sci.*, 15, 2413–2423, doi:10.5194/nhess-15-2413-
652 2015, 2015.
- 653 Martelloni, G., Segoni, S., Fanti, R., and Catani, F.: Rainfall thresholds for the forecasting of
654 landslide occurrence at regional scale, *Landslides*, 9, 485–495, doi:10.1007/s10346-011-0308-2,
655 2012.
- 656 Morss, R, Wilhelmi, O, Meehl, G, Dilling, L: Improving societal outcomes of extreme
657 weather in a changing climate: an integrated perspective. *Annual Review of Environment and*
658 *Resources*, 36, 1–25, doi: 10.1146/annurev-environ-060809-100145, 2011.
- 659 NOAA-USGS Debris Flow Task Force: NOAA-USGS debris-flow warning system. Final
660 report, US Geological Survey Circular 1283, US Geological Survey, Reston, Virginia, USA: 47.
661 Available at: <http://pubs.usgs.gov/circ/2005/1283/pdf/Circular1283.pdf> (last access: November
662 2014), 2005.
- 663 Osanai N, Shimizu T, Kuramoto K, Kojima S, Noro T: Japanese early-warning for debris
664 flows and slope failures using rainfall indices with Radial Basis Function Network, *Landslides*, 7,
665 325–338, doi:10.1007/s10346-010-0229-5, 2010.
- 666 Piciullo L, Siano I, Calvello M: Calibration of rainfall thresholds for landslide early warning
667 purposes: applying the EDuMaP method to the system deployed in Campania region (Italy). In:
668 *Proceedings of the International Symposium on Landslides 2016-Landslides and Engineered*
669 *Slopes. Experience, Theory and Practice*. Napoli, Italy, 3:1621–1629. ISBN 978–1–138-02988-0,
670 2016a.

671 Piciullo L, Gariano S L, Melillo M, Brunetti M T, Peruccacci S, Guzzetti F., Calvello M:
672 Definition and performance of a threshold-based regional early warning model for rainfall-induced
673 landslides. *Landslides*. doi:10.1007/s10346-016-0750-2, 2016b.

674 Ramberg IB, Bryhni I, Nøttvedt A, Rangnes K: The making of a land – geology of Norway.
675 Trondheim. Norsk Geologisk Forening, 624 p., 2008.

676 Rossi, M, Peruccacci, S, Brunetti, M T, Marchesini, I, Luciani, S, Ardizzone, Balducci, S V,
677 Bianchi, C, Cardinali, M, Fiorucci, F, Mondini, A C, Reichenbach, P, Salvati, P, Santangelo, M,
678 Bartolini, D, Gariano, S L, Palladino, M, Vessia, G, Viero, A, Antronico, L, Borselli, L, Deganutti,
679 A M, Iovine, G, Luino, F, Parise, M, Polemio, M, Guzzetti, F: SANF: National warning system for
680 rainfall-induced landslides in Italy, in: *Landslides and Engineered Slopes: Protecting Society
681 through Improved Understanding*, edited by: Eberhardt, E., Froese, C., Turner, K., and Leroueil, S.,
682 Taylor& Francis, London, 1895–1899, isbn 978-0-415-62123-6, 2012.

683 Segoni, S., Rossi, G., Rosi, A., and Catani, F.: Landslides triggered by rainfall: a
684 semiautomated procedure to define consistent intensity-duration thresholds, *Comput. Geosci.*, 3063,
685 123–131, <http://dx.doi.org/10.1016/j.cageo.2013.10.009>, 2014.

686 Segoni S, Battistini A, Rossi G, Rosi A, Lagomarsino D, Catani F, Moretti S, Casagli N:
687 Technical note: an operational landslide early warning system at regional scale based on space–
688 time-variable rainfall thresholds. *Nat Hazards Earth Syst Sci* 15:853–861. doi:10.5194/nhess-15-
689 853-2015, 2015.

690 Stähli M, Sättele M, Huggel C, McArdell BW, Lehmann P, Van Herwijnen A, Berne A,
691 Schleiss M, Ferrari A, Kos A, Or D, Springman SM: Monitoring and prediction in early warning
692 systems for rapid mass movements. *Nat Hazards Earth Syst Sci* 15:905–917. doi:10.5194/nhess-15-
693 905-2015, 2015.

694 Staley D M, Kean J W, Cannon S H, Laber J L, Schmidt K M: Objective definition of
695 rainfall intensity-duration thresholds for the initiation of post-fire debris flows in southern
696 California. *Landslides* 10: 547-562. doi: 10.1007/s10346-012-0341-9, 2013.

697 Thiebes B, Glade T, Bell R: Landslide analysis and integrative early warning-local and
698 regional case studies, in: *Landslides and Engineered Slopes: Protecting Society through Improved
699 Understanding*, edited by: Eberhardt, E., Taylor& Francis Group, London, 1915–1921, isbn 978-0-
700 415-62123-6, 2012.

701 Tiranti D, and Rabuffetti D: Estimation of rainfall thresholds triggering shallow landslides
702 for an operational warning system implementation. *Landslides*, 7: 471–481, doi:10.1007/s10346-
703 010-0198-8, 2010.

704 United Nations Inter-Agency Secretariat of the International Strategy for Disaster Reduction
705 (UN ISDR): *Global Survey of Early Warning Systems: An assessment of capacities, gaps and
706 opportunities towards building a comprehensive global early warning system for all natural hazards*,
707 available at:[http://www.unisdr.org/2006/ppew/info-resources/ewc3/Global-Survey-of-Early-
708 Warning-Systems.pdf](http://www.unisdr.org/2006/ppew/info-resources/ewc3/Global-Survey-of-Early-Warning-Systems.pdf) (last access:November 2014), 2006.