

1 **Title**

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4 **Adapting the EDuMaP method to test the performance of the Norwegian early**
5 **warning system for weather-induced landslides**

6

7 **Authors**

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14

15 **Abstract**

16 The Norwegian national landslide early warning system (LEWS), operational since 2013, is
17 managed by the Norwegian Water Resources and Energy Directorate and has been designed for
18 monitoring and forecasting the hydro-meteorological conditions potentially triggering slope
19 failures. Decision-making in the EWS is based upon rainfall thresholds, hydro-meteorological and
20 real-time landslide observations as well as on landslide inventory and susceptibility maps. Daily
21 alerts are issued throughout the country considering variable size warning zones. Warnings are
22 issued once per day for the following 3 days and can be updated according to weather forecasts and
23 information gathered by the monitoring network. The performance of the LEWS operational in
24 Norway has been evaluated applying the EDuMaP method, which is based on the computation of a
25 duration matrix relating number of landslides and warning levels issued in a warning zone. In the
26 past, this method has been exclusively employed to analyse the performance of regional early
27 warning model considering fixed warning zones. Herein, an original approach is proposed for the
28 computation of the elements of the duration matrix in the case of early warning models issuing
29 alerts on variable size areas. The approach has been used to evaluate the warnings issued in Western
30 Norway, in the period 2013-2014, considering two datasets of landslides. The results indicate that
31 the landslide datasets do not significantly influence the performance evaluation, although a slightly
32 better performance is registered for the smallest dataset. Different performance results are observed
33 as a function of the values adopted for one of the most important input parameters of EDuMaP, the
34 landslide density criterion (i.e. setting the thresholds to differentiate among classes of landslide
35 events). To investigate this issue, a parametric analysis has been conducted; the results of the
36 analysis show significant differences among computed performances when absolute or relative
37 landslide density criteria are considered.

38 **Keywords:** EDuMaP method, rainfall-induced landslides, warning zones, alert, landslide density.

39 **1. Introduction**

40 In the last decades, natural hazards caused an increased number of consequences in terms of
41 economic losses (Barredo, 2009) and fatalities throughout Europe (European Environment Agency,
42 2010; CRED, 2011). Most natural disasters are related to extreme rainfall events, which are
43 expected to increase with climate change (Easterling et al., 2000; Morss et al., 2011). The European
44 Commission, following an increase in human and economic losses due to natural hazards,
45 developed legal frameworks such as the Water Framework Directive 2000/60/EC (2000) and the
46 Floods Directive 2007/60/EC (2007), to increase prevention, preparedness, protection and response
47 to such events and to promote research and acceptance of risk prevention measures within the
48 society (Alfieri et al., 2012). Among the many mitigation measures available for reducing the risk to
49 life related to natural hazards, early warning systems (EWSs) constitute a significant option
50 available to authorities in charge of risk management and governance.

51 Within the landslide risk management framework proposed by Fell et al. (2005), landslide EWSs
52 (LEWSs) may be considered a non-structural passive mitigation option to be employed in areas
53 where risk, occasionally, rises above previously defined acceptability levels. According to Glade
54 and Nadim (2014), the installation of an EWS is often a cost-effective risk mitigation measure and
55 in some instances the only suitable option for sustainable management of disaster risks. Rainfall-
56 induced warning systems for landslides are, by far, the most diffuse class of landslide EWS
57 operating around the world. LEWSs can be employed at two distinct scales of analysis: “local” and
58 “regional” (ICG 2012; Thiebes et al. 2012; Calvello et al. 2015, Stähli et al., 2015). EWSs at a
59 regional scale for rainfall-induced landslides have become a sustainable risk management approach
60 worldwide to assess the probability of occurrence of landslides over appropriately-defined wide
61 warning zones. In fact during the last decades, several systems have been designed and improved,
62 not only in developing countries (UNISDR 2006; Chen et al., 2007; Huggel et al., 2010; among
63 others) but also in developed countries (NOAA-USGS, 2005; Badoux et al., 2009; Baum and Godt,
64 2010; Osanai et al., 2010; Lagomarsino et al., 2013; Tiranti and Rabuffetti, 2010; Rossi et al., 2012;
65 Staley et al., 2013; Calvello et al., 2015; Segoni et al., 2015). As a recent example, the Norwegian
66 landslide EWS was launched in autumn 2013 by the Norwegian Water Resources and Energy
67 Directorate (NVE). The regional system has been developed for monitoring and forecasting the
68 hydro-meteorological conditions triggering landslides and to inform local emergency authorities in
69 advance about the occurrence of possible events (Devoli et al., 2014). Daily alerts are issued

70 throughout the country in variable size warning zones. The evaluation of the alerts issued, i.e., the
71 performance of the early warning model is not a trivial issue, and regular system testing and
72 performance assessments (Hyogo Framework for Action, 2005) are fundamental steps.
73 The performance analysis of LEWSs can be an awkward process, particularly for systems employed
74 at regional scale, because many aspects are important for the analyst to consider. Most typically,
75 the performance evaluation is based on 2 by 2 confusion matrices computed for the joint frequency
76 distribution of landslides and alerts, both considered as dichotomous variables, and the evaluation
77 of statistical indicators (e.g., Cheung et al., 2006; Godt et al., 2006; Martelloni et al., 2012; Staley et
78 al., 2013; Segoni et al., 2014; Lagomarsino et al., 2015; Gariano et al., 2015; Stähli et al., 2015).
79 The method employed herein, which is called EDuMaP (Calvello and Piciullo, 2016), allows to
80 consider aspects peculiar to territorial LEWSs that are not considered by the joint frequency
81 distribution approach. In particular, the EDuMaP method takes into account: the occurrence of
82 concurrent multiple landslides in the warning zone; the duration of the warnings in relation to the
83 landslides; the issued warning level in relation to the landslide spatial density in the warning zone;
84 the relative importance attributed, by system managers, to different types of errors. Up to now, this
85 method has been applied exclusively to evaluate the performance of regional warning models
86 designed for issuing alerts in fixed warning zones (Calvello and Piciullo, 2016; Piciullo et al.,
87 2016a,b; Calvello et al., 2016). In the present study the EDuMaP method has been adapted to
88 evaluate the performance of the alerts issued for variable size warning zones. To this purpose, the
89 procedure has been tested on the Norwegian landslide EWS in the period 2013-2014. The Western
90 Norway is the area most prone to landslides in Norway and it has been chosen as test area because
91 the landslide database was more reliable and complete than for the rest of Norway.
92

93 **2. The national landslide early warning system for rainfall-and snowmelt-** 94 **induced landslides in Norway**

95 **2.1 Physical setting**

96 Norway covers an area of ~ 324,000 km². With its elongated shape of 1800 km, the country reaches
97 from latitude 58°N to 71°N. Approximately 30% of the land area are mountainous, with the highest
98 peaks reaching up to 2500 m. a.s.l and slope angles over 30 degrees covering 6,7% of the country
99 (Jaedicke *et al.*, 2009). In geological terms, Norway is located along the western margin of the
100 Baltic shield with a cover of Caledonian nappes in the western parts of the country (Etzelmüller *et*
101 *al.*, 2007; Ramberg *et al.*, 2008). The Caledonian nappes are dominated by Precambrian rocks and

102 metamorphic Cambro-Silurian sediments, while the bedrock in the Baltic shield is dominated by
103 Precambrian basement rocks. Cambro-Silurian sediments and Permian volcanic rocks are found in
104 the Oslo Graben (Ramberg *et al.*, 2008).

105 Recurrent glaciations, variations in sea level and land subsidence/uplift, as well as weathering,
106 transport and deposition processes have created the modern Norwegian landscape (Ramberg *et al.*,
107 2008). Thus, dominating quaternary deposits include various shallow (in places colluvial) soils, as
108 well as moraine and marine deposits.

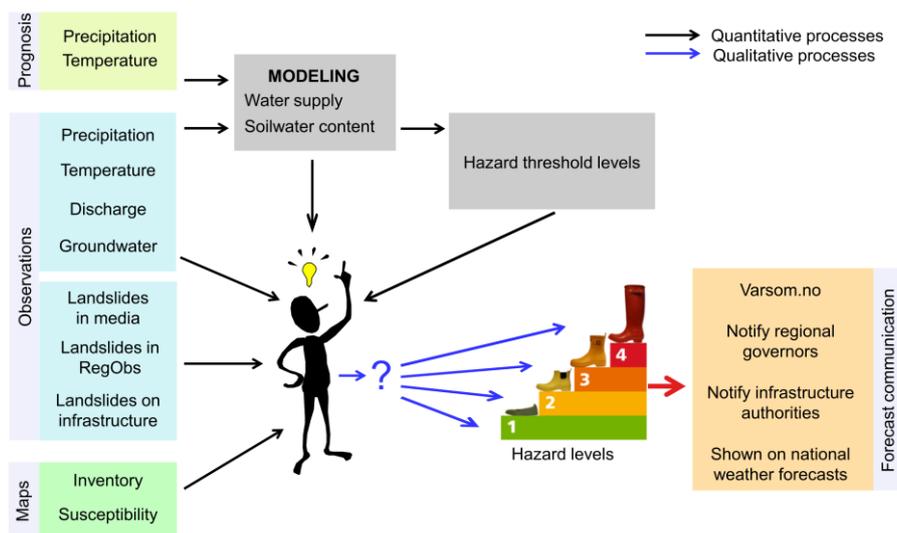
109 Because of the latitudinal elongation and the varied topography, the Norwegian climate displays
110 large variations. Along the Atlantic coast, the North Atlantic Current influences the climate whereas
111 the inland areas experiences a more continental climate. Based on the Köppen classification
112 scheme, the Norwegian climate can be classified in three main types: warm temperate humid
113 climate, cold temperate humid climate and polar climate. Precipitation types can be divided into
114 three categories: frontal, orographic and showery. The largest annual precipitation values are found
115 near the coast of Western Norway (herein also called Vestlandet) with up to 3575 mm/year. On the
116 other hand, the driest areas receiving <500 mm/year are found in parts of South-Eastern Norway
117 (Østlandet) and Finnmark county.

118 Steep landforms in combination with various soil and climatic properties provide a basis for several
119 types of shallow landslides in non-rock materials. These slope failures include slides in various
120 materials, debris avalanches, debris flows and slush flows. Landslides are mostly triggered by
121 rainfall, often in combination with snowmelt. Some events are also triggered from/initiated as
122 rockfall or slush flows, developing into, for example, debris flows as they propagate downslope.
123 Shallow landslides constitute a substantial threat to the Norwegian society. According to Furseth
124 (2006), at least 230 people have been killed by such slope failures during the latest approximately
125 500 years. In the period 2000-2009, road authorities registered more than 1800 shallow landslides
126 along Norwegian roads.

127

128 2.2 The national landslide early warning system

129 In order to mitigate the risk from shallow landslides, a national EWS has been developed at the
130 Norwegian Water Resources and Energy Directorate (NVE) as part of the national responsibility on
131 landslide risk management. The system is established to warn about the hazard of debris flows,
132 debris slides, debris avalanches and slush flows at regional scale. The EWS, operative since 2013,
133 has been developed in cooperation with the Norwegian Meteorological Institute (MET), Norwegian
134 Public Road Administration (SVV) and the Norwegian National Rail Administration (JBV).



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Figure 1. Organization of the landslide early warning system in Norway.

139 Decision-making in the EWS is based upon hazard threshold levels, hydro-meteorological and real-
 140 time landslide observations as well as landslide inventory and susceptibility maps (**Fig. 1**). In the
 141 development phase of the EWS, hazard threshold levels have been investigated through statistical
 142 analyses of historical landslides and modelled hydro-meteorological parameters. Daily hydro-
 143 meteorological conditions such as rainfall, snowmelt, runoff, soil saturation, groundwater level and
 144 frost depth have been obtained from a distributed version of the hydrological HBV-model (Beldring
 145 *et al.*, 2003).

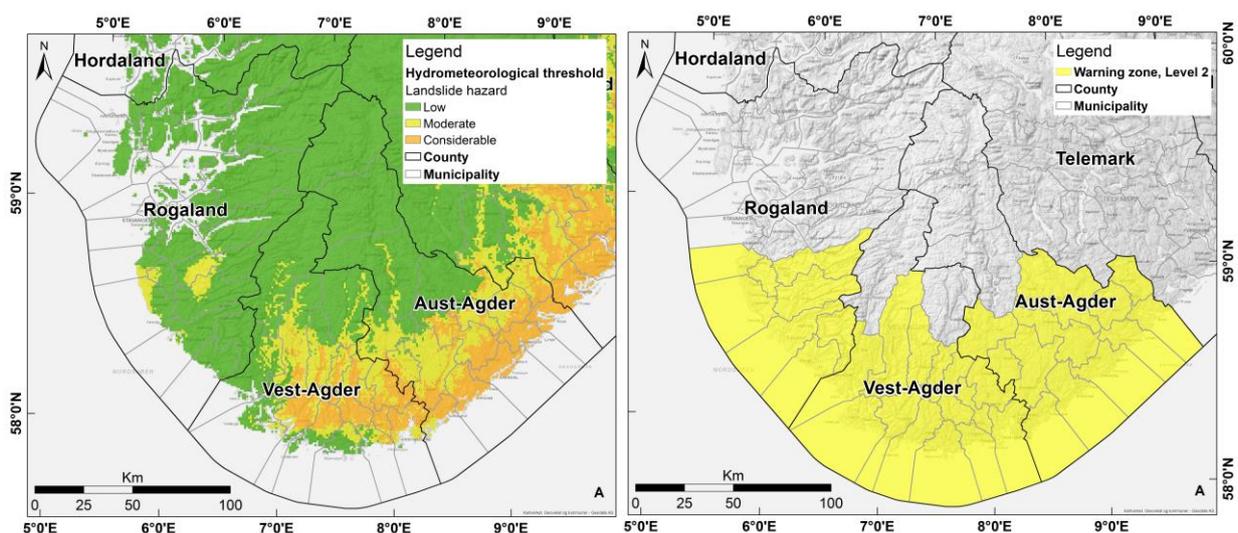
146 Hazard threshold levels presently used in the EWS were proposed by Colleuille *et al.* (2010). The
 147 thresholds, combining simulations of relative water supply of rain or snowmelt and relative soil
 148 saturation/groundwater conditions, were derived from empirical tree-classification using 206
 149 landslide events from different parts of the country. Later analyses, summarized by Boje *et al.*
 150 (2014), confirm the good performance of combining soil water saturation degree and normalised
 151 rainfall and snowmelt.

152 Two different landslide susceptibility maps are used as supportive data in the process of setting
 153 daily warning levels. One map indicates initiation and runout areas for debris flows at slope scale
 154 (Fischer *et al.*, 2012), while another indicates susceptibility at catchment level, based upon
 155 Generalized Additive Models (GAM) statistics (Bell *et al.*, 2014).

156 Susceptibility maps, hazard threshold levels and other relevant data are displayed in real-time in a
 157 webpage, www.xgeo.no, which is used as decision expert tool to forecast various natural hazards
 158 (floods, snow avalanches, landslides). Landslide hazard threshold levels and hydrometeorological

159 forecasts are displayed as raster data with 1 km² resolution, whereas susceptibility maps, landslide
160 information (historical and real-time) and hydrometeorological observations are shown as either
161 raster, polygon or point data.

162 A landslide expert on duty (as member of a rotation team) uses the information from forecasts,
163 observations, maps and uncertainty in weather forecasts to qualitatively perform a nationwide
164 assessment of landslide warning levels (**Fig. 1**). Four warning levels are defined: green (1), yellow
165 (2), orange (3), and red (4) showing the level of hazards, or more exactly the recommended
166 awareness level (**Tab. 1**). The warning period follows the time steps of quantitative precipitation
167 and temperature forecasts used to simulate other hydro-meteorological parameters, and thus lasts
168 from 06:00 UTC to 06:00 UTC each day. Warning levels are updated minimum twice during the 24
169 hour warning period (morning and afternoon) as a function of the weather forecast. Weather
170 forecast updates are received 4 times per day and warning messages are sent as soon as possible,
171 from 66 hours to few hours ahead. Warning messages are published in a publicly accessible
172 webpage (www.varsom.no). Yellow, orange and red levels of warning are also sent to emergency
173 authorities (regional administrative offices, roads and railways authorities) and media. Warning
174 zones are not static geographical warning areas. Instead they vary from a small group of
175 municipalities to several administrative regions, depending on current hydro-meteorological
176 conditions (**Fig. 2**). Thus, extent and position of warning zones are dynamic and change from day to
177 day.



178
179 **Figure 2.** a) Hydrometeorological thresholds indicating potential landslide hazard in the counties of
180 Rogaland, Vest-Agder, Aust-Agder and Telemark in South-Eastern Norway on 15.02.2014. b) The

181 resultant early warning zone, on warning level 2 (“yellow level”) issued on 15.02.2014 for the same
 182 area and including about 32 municipalities.

183

184 2.3 Current performance evaluation of the EWS

185 To evaluate the performance of a regional landslide early warning model, a comparison of warning
 186 levels issued and landslides occurred is carried out on a weekly basis. Event information is reported
 187 by Roads/Railways Authorities or municipalities, as well as obtained from media and from a real-
 188 time database to register observations. The latter has been designed as a public tool supporting
 189 crowd sourcing (Ekker et al. 2013), and is currently available to the public as a telephone
 190 application and a website (www.regobs.no). Categorization of issued warning levels into false
 191 alarms, missed events, correct and wrong levels is based on semi-quantitative classification criteria
 192 for each warning level. The principle behind the criteria is that rare hydro-meteorological conditions
 193 are expected to cause more landslides and possibly higher damages (**Tab. 1**). As an example, the
 194 warning level Red corresponds to an extreme situation that occurs very rarely. It requires immediate
 195 action and may cause severe damages within a large extent of the warning area. The criteria contain
 196 information on the expected number of landslides per area, as well as hazard signs indicating
 197 landslide activity. As seen in **Table 1** the ranges chose for the number of expected landslides and
 198 the size of the hazardous areas at each warning level are quite wide. This choice is due to the fact
 199 that the EWS is relatively new and still in a phase of continuous development.

200

201 **Table 1.** Criteria for evaluating daily warning levels in the Norwegian EWS.

Warning level	Classification criteria
4 (Red)	> 14 landslide (per 10-15.000 km ²) Hazard signs: Several road blockings due to landslides or flooding Extreme situation that occurs very rarely, requires immediate action and may cause severe damages within a large extent of the warning area. This level corresponds to a >50 years return period flood warning.
3 (Orange)	6-10 landslides (per 10-15.000 km ²) Hazard signs: Several road blockings due to landslides or flooding Severe situation that occurs rarely, require contingency preparedness and may cause severe damages within some extent of the warning area. This level corresponds to 5-50 years return period flood warning.
2 (Yellow)	1-4 landslides (per 10-15.000 km ²) Hazard signs: flooding/erosion in streams Situation that requires monitoring and may cause local damages within the warning area. Expected some landslide events, certain large events may occur.
1 (Green)	No landslides 1-2 landslide caused by local rain showers 1 small debris slide if in area with no signs of elevated warning level

202

203 **3. Performance evaluation of the LEWS in Western Norway for the period**
204 **2013-2014**

205 **3.1 Study area and landslide data**

206 The study area includes the four administrative regions of Møre og Romsdal, Sogn og Fjordane,
207 Hordaland and Rogaland located on the Norwegian west-coast. A common name for the entire area
208 is Vestlandet (i.e. Western Norway). The area is dominated by narrow fjords and steep
209 mountainsides reaching from sea level to 1000 m a.s.l. or more, and high annual precipitation of up
210 to ~3500 mm. Shallow quaternary deposits cover locally weathered and altered bedrock of mainly
211 precambrian and Caledonian metamorphic and magmatic origin. As a result, Vestlandet is highly
212 prone to landslides, in particular, debris avalanches, debris flows and slush flows.

213 Vestlandet is the rainiest area of Norway with many annual precipitation events bringing high
214 amounts of rain and/or snow. Precipitation patterns and spatial distribution display large variations
215 within the study area. The precipitation patterns are described based on the main spatial
216 distribution:

- 217 a) NNW precipitation only in the region of Møre og Romsdal;
- 218 b) NW precipitation mainly in the regions of Møre og Romsdal and Sogn og Fjordane, or
219 sometimes in the northern part of Hordaland;
- 220 c) WNW precipitation in the entire study area;
- 221 d) W precipitation distributed mainly in Sogn og Fjordane, Hordaland and Rogaland;
- 222 e) SW precipitation distributed mainly in Rogaland and Hordaland, or sometimes also in Sogn
223 of Fjordane;
- 224 f) SSW precipitation only in Rogaland, or sometimes in Hordaland and rarely in the southern
225 part of Sogn og Fjordane;
- 226 g) S and SE with precipitation mainly in South-Eastern Norway (in summer) and not in the
227 study area, however because of size of the systems, precipitation can spread to Møre og
228 Romsdal or to eastern Sogn og Fjordane or Hordaland, depending on trajectory;
- 229 h) Local showers (mostly in summer), with clusters of maximum precipitation distributed
230 randomly within the study area;
- 231 i) Southern Norway, with precipitation distributed in the entire southern part of the country
232 and consequently in the entire study area.

233 During the years 2013 and 2014 more than 70 precipitation events, i.e. rain and/or snow records
234 with more than 30 mm/24h, were registered, with some episodes bringing more than 75-150

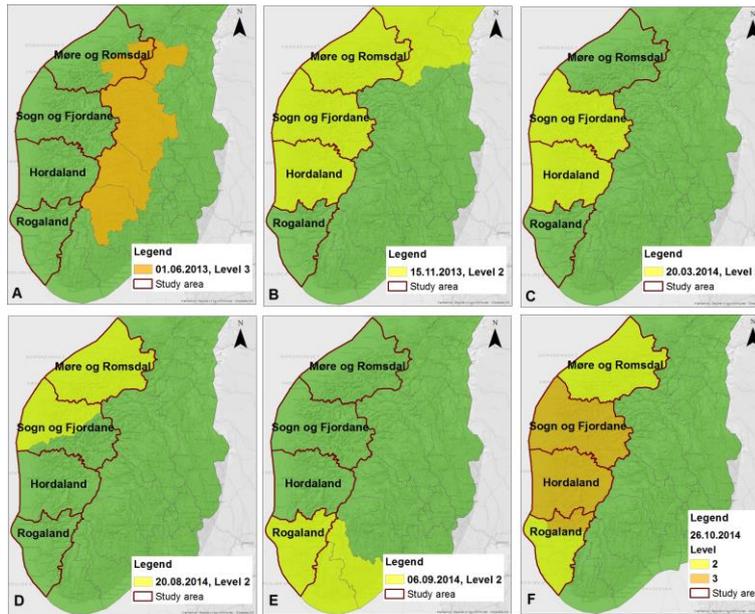
235 mm/24h of rain/snow to the entire study area or part of it, following the patterns indicated above.
 236 Duration of precipitation events ranged from 1 day to 14-18 consecutive days, particularly during
 237 autumn.

238 Landslide early warnings higher than green level were issued for 49 days during the two-year
 239 period (**Tab. 2**). Most of these were at yellow level, however five warnings at orange level were
 240 issued in 2014 in 3 consecutive days. In 12 cases, the yellow warnings issued during the morning
 241 evaluation was downgraded to green later the same day. The most significant precipitation events
 242 recorded in 2013-2014 are 11 and occurred in the following days: 14-15/04/13, 12-13/08/13,
 243 7/10/13, 22/10/13, 15/11/ 13, 28/12/ 13, 23/02/ 14, 20/03/14, 14/07/14, 18-19/08/14, 27-28/10/14.
 244

245 **Table 2.** Significant rainfall, number of days with at least one warning, number of warnings and
 246 landslides in the period 2013-2014.

	2013	2014	tot
Precipitation events, i.e. rainfall and/or snow > 30 mm/24h	41	32	73
Number of days with at least one warning	20	29	49
Number of warnings	21	39	60
red warnings	0	0	
orange warnings	0	5	
yellow warnings	21	34	
Number of landslides	204	181	385

247
 248
 249 Examples of warnings issued during 2013 and 2014 are shown in **Figure 3**. Most of the alerted
 250 warning zones were completely included in the study area (**Fig. 3 c, d, f**). However, some warnings
 251 were mainly issued for neighboring areas to the 4 regions chosen as case study (**Fig.3 a, b, e**). The
 252 examples of **Figure 3** also illustrates the diversity in having variable instead of fixed size warning
 253 zones.
 254



255
256 **Figure 3.** Examples of early warning areas and levels during 2013-2014.
257

258 Within the study area, for the period 2013-2014, the Norwegian national landslide database
259 (www.skrednett.no) lists 476 landslides in soils and/or slush flows. Due to errors and double
260 registration, 385 of these slope failures were considered valid for the current analyses: 249 (65%)
261 are categorized as landslide in soil, not otherwise specified due to lack of further documentation; 65
262 (17%) are categorized as debris avalanches, following Hungr et al. (2014), in many cases initiated
263 as small debris slides; 27 (7%) are classified as debris flows, following Hungr et al. (2014); 20 (5%)
264 are soil slides in artificial slopes (cuts and fillings along roads or railway lines); 19 (4%) are slush
265 flows and the remaining 5 (1%) are rock falls developing into debris avalanches.

266 The EDuMaP method was applied to two different sets of phenomena: Set A and Set B. The first set
267 includes all 385 slope failures, while the second included only 131 phenomena, as “landslide in soil
268 not specified” and “rock fall/debris avalanches” were removed from this dataset. The removal of
269 non-specified landslides was due to the questionable quality of these registrations in the national
270 landslide database, while the exclusion of rock falls inducing debris avalanches was due to
271 uncertainty on whether precipitation can indeed be considered their triggering cause.

272 3.2 The EDuMaP method

273 The paper proposes the evaluation of the performance of the landslide early warning system
274 operational in Norway by means of the “Event, Duration Matrix, Performance (EDuMaP) method”
275 (Calvello & Piciullo, 2016). This method has been principally employed to analyse the performance
276 of regional early warning model considering fixed warning zones for issuing alerts. The method

277 comprises three successive steps: identification and analysis of landslide and warning Events (E),
278 from available databases; definition and computation of a Duration Matrix (DuMa), and evaluation
279 of the early warning model Performance (P) by means of performance criteria and indicators.
280 The first step requires the availability of landslides and warnings databases for the preliminary
281 identification of “landslide events” (LEs) and “warning events” (WEs). A landslide event is defined
282 as one or more landslides grouped on the basis of their spatial and temporal characteristics. A
283 warning event is defined as a set of warning levels issued within a given warning zone, grouped
284 considering their temporal characteristics. The parameters which need to be defined to carry on the
285 events analysis are ten: 1) warning levels, W_{lev} ; 2) landslide density criterion, $L_{den(k)}$; 3) lead time,
286 t_{LEAD} ; 4) landslide typology, L_{typ} ; 5) minimum interval between landslide events, Δt_{LE} ; 6) over time,
287 t_{OVER} ; 7) area of analysis, A; 8) spatial discretization adopted for warnings, $\Delta A_{(k)}$; 9) time frame of
288 analysis, ΔT ; 10) temporal discretization of analysis, Δt . For more details see Calvello and Piciullo,
289 2016. The second step of the method is the definition and computation of a “duration matrix”,
290 whose elements d_{ij} , report the time associated with the occurrence of landslide events in relation to
291 the occurrence of warning events, in their respective classes. The element d_{11} of the matrix
292 expresses the number of hours when no warnings are issued and no landslides occur (Fig. 4). The
293 number of rows and columns of the matrix is equal to the number of classes defined for the warning
294 and landslide events, respectively (**Fig. 4**). The final step of the method is the evaluation of the
295 duration matrix based on a set of performance criteria assigning a performance meaning to the
296 element of the matrix. Two criteria are used for the following analyses (**Fig. 4**), respectively
297 indicated as criterion 1 and criterion 2. The first criterion employs an alert classification scheme
298 derived from a 2x2 contingency table, thus identifying: correct predictions, CPs; false alerts, FAs;
299 missed alerts, MAs; true negatives, TNs. The second criterion assigns a color code to the elements
300 of the matrix in relation to their grade of correctness, classified in four classes as follows: green, G,
301 for the elements which are assumed to be representative of the best model response; yellow, Y, for
302 elements representative of minor model errors; red, R, for elements representative of a significant
303 model errors; purple, P, for elements representative of the worst model errors. Both criteria
304 purposefully neglect element d_{11} , whose value is typically orders of magnitude higher than the
305 values of the other elements of the matrix because it also includes all hours without rainfall, for
306 which a LEWS is not designed to deal with, specifically. Thus, d_{11} element is neglected in order to
307 avoid an overestimation of the performance and to allow a more useful relative assessment of the
308 information located in the remaining part of the duration matrix. A number of performance

309 indicators may be derived from the two performance criteria described. **Table 3** reports the name,
 310 symbol, formula and value of the performance indicators considered herein.

1) Alert classification criterion		Landslide events			
		no	S	I	L
Warning events	no	TN	CP	MA	MA
	M	CP	CP	MA	MA
	H	FA	CP	CP	CP
	VH	FA	FA	CP	CP

2) Grade of correctness criterion		Landslide events			
		no	S	I	L
Warning events	no	G	Y	R	P
	M	Y	G	R	P
	H	R	R	G	Y
	VH	P	P	Y	G

311
 312 **Figure 4.** Performance criteria used for the analyses performed herein (modified from Calvello &
 313 Piciullo, 2016). Four classes of warning events (key: no, no warning; M, moderate warning; H, high
 314 warning; VH, very high warning) and four classes of landslide events (key: no, no landslides; S,
 315 small event, few landslides; I, intermediate event, several landslides; L, large events, many
 316 landslides).

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Table 3. Performance indicators used for the analysis.

Performance indicator	Symbol	Formula
Efficiency index	I_{eff}	$CP/\sum_{ij}d_{ij}$ (excluding d_{11})
Hit rate	HR_L	$CP/(CP+MA)$
Predictive power	PPW	$CP/(CP+FA)$
Threat score	TS	$CP/(CP+MA+FA)$
Odds ratio	OR	$CP/(MA+FA)$
Miss classification rate	MR	$1 - I_{eff}$
Missed alert rate	R_{MA}	$MA/(CP+MA)$
False alert rate	R_{FA}	$FA/(CP+FA)$
Error Rate	ER	$(Red\&Pur)/\sum_{ij}d_{ij}$ (excluding d_{11})
Missed and false alerts balance	MFB	$MA/(MA+FA)$
Probability of serious mistakes	P_{SM}	$Pur/\sum_{ij}d_{ij}$ (excluding d_{11})

319 3.3 Adaptation of the EDuMaP method to variable size warning zones

320 In earlier studies, the EDuMaP method has been applied to analyse the performance of regional
 321 landslide EWSs adopting a fixed spatial discretization for warnings. In contrast, the Norwegian
 322 landslide EWS employs variable size warning zones. This characteristic influences the first two
 323 phases of the EDuMaP method: identification and analysis of landslide and warning events from

324 available databases; definition and computation of a duration matrix. This section explains how to
325 define LEs and WEs and how to compute the duration matrix in case of variable size warning
326 zones.

327 The Norwegian EWS adopts four warning levels. Daily warnings are issued throughout the country
328 considering municipalities as the minimum warning territorial unit (TU). Hence, municipalities
329 alerted with the same warning level define a warning zone of level i (e.g., green, yellow, orange, red
330 in **Fig. 5**). Therefore, on a day of alert, up to four warning zones alerted with different warning
331 levels can be issued (e.g., day 3 in **Fig. 5**). LEs are defined by grouping together landslides occurred
332 within each warning zone. The class each LE belong to, as defined in **section 3.2**, depends on the
333 landslide density criterion, $L_{den(k)}$, chosen for the analyses.

334 The duration matrix is evaluated for the whole area of analysis, A , in a period of analysis, ΔT ,
335 summing the $time_{ij}$ computed within the different warning zones, for each temporal discretization
336 Δt . In particular, the values of $time_{ij}$, for variable size warning zones, are computed as follows:

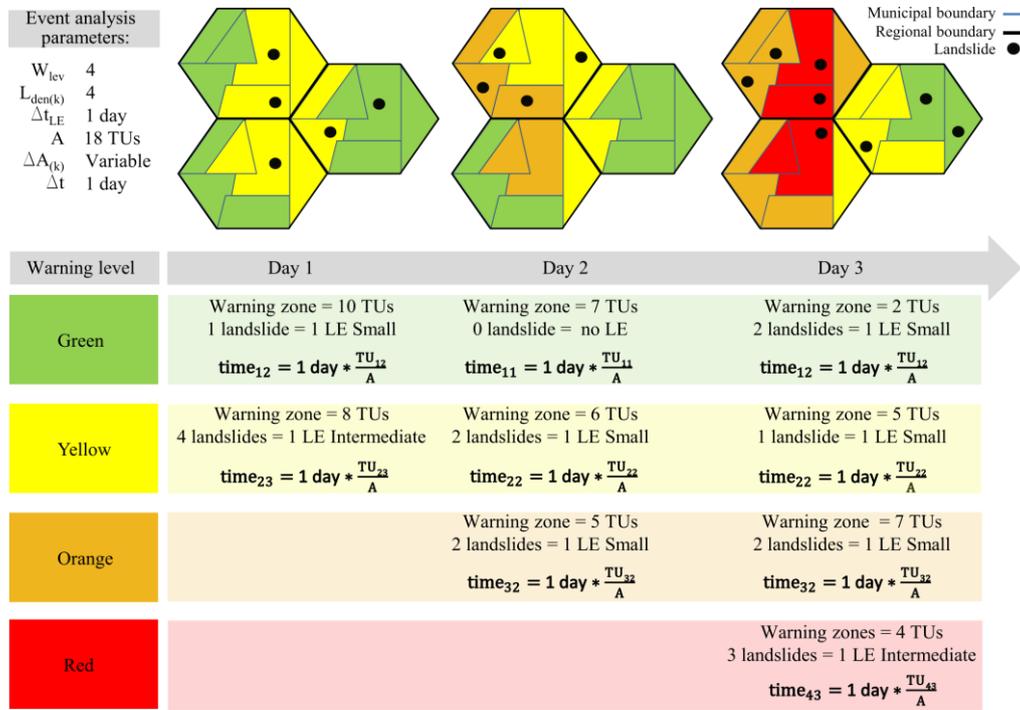
$$337 \quad time_{ij} = \sum_{\Delta t} \frac{(TUA_{ij})}{A} \quad (\text{Eq. 1})$$

338 where: Δt is the minimum temporal discretization adopted for warnings (for the Norwegian EWS,
339 equal to 1 day); A is the area of analysis; TUA_{ij} is the extent of the territorial unit alerted with a
340 warning level i , and class of the landslide event, j , per day of alert. Each element of the duration
341 matrix, d_{ij} , is then computed, within the time frame of the analysis, ΔT , as follows:

$$342 \quad d_{ij} = \sum_{\Delta T} (time_{ij}) \quad (\text{Eq. 2})$$

343 The evaluation of landslide and warning events and the definition and computation of a the duration
344 matrix is herein exemplified for three hypothetical days (Fig. 8). For instance, on Day 1 two distinct
345 LEs appear, containing 4 and 1 landslides, respectively. The first event belongs to the warning zone
346 alerted with level 2 and the latter to the warning zone alerted with level 1. In Day 3 there are 4
347 warning zones, each one alerted with a different warning level and 4 distinct LEs can be identified,
348 one per warning zone. A landslide density criterion, $L_{den(k)}$ in four classes has been considered for
349 the example of **Figure 5**: 0 (no landslides), small (1-2 landslides), Intermediate (3-4 landslides) and
350 Large (≥ 5 landslides); together with four warning levels, W_{lev} : green, yellow, orange and red. At
351 “day 1” two different warning zones can be defined grouping together the TUs (blue boundary in
352 **Fig. 5**) with the same warning level. The warning zones are composed by 10 and 8 TUs, and they
353 are alerted with two different warning levels: green and yellow. In the two warning zones, a “small”
354 LE and an “Intermediate” LE, respectively, are occurred. Once the warning levels and the LEs
355 within each warning zone have been defined, $time_{12}$ and $time_{23}$ are evaluated for each TU using

356 **Equation 1.** At “day 2” three warning zones and two “Small” LEs have been identified. At “day 3”
 357 LEs occurred in each of the four warning zones identified. Finally, the evaluation of elements d_{ij} of
 358 the duration matrix, is carried out following **Equation 2**, over the time frame of the analysis, ΔT .
 359



360
 361 **Figure 5:** Computation of $time_{ij}$ elements as a function of warning levels and LEs occurred for each
 362 warning zone for three hypothetical days of warning.

363

364 4. Results and discussion

365 4.1 Events analysis

366 As previously mentioned, the events analysis phase of the EDuMaP method depends on the values
 367 assumed by a series of well-identified parameters, which are defined to allow the analyst to make
 368 choices on how to select and group landslides and warnings.

369 The values of the ten input parameters, cf. section 3, for the two analyses carried out, i.e. case A and
 370 case B, are representative of the structure and operational procedures of the warning model
 371 employed in the Norwegian EWS. The period of analysis, ΔT , is 2013-2014, while Δt , is set to 1
 372 day. Parameters t_{LEAD} and t_{OVER} are both set to zero. The four warning levels, W_{lev} , are: green (no
 373 warning), yellow (WL_1), orange (WL_2), red (WL_3). The landslides used for the analyses are grouped
 374 into landslide events considering a Δt_{LE} of 1 day. The four classes of LEs are defined employing a
 375 relative landslide density criterion, $L_{den(k)}$, as a function of both number of landslides and territorial

376 extensions. The values have been derived by the criteria for the daily warning levels evaluation in
 377 the Norwegian EWS (see **Tab. 1**). The only difference between case A and case B has to do with
 378 the type of landslides used for the analyses, which respectively refer to the datasets A and B.
 379 Dataset A is composed by 385 rainfall- and snowmelt-induced landslides occurring within the study
 380 area. These slope failures have been grouped into 137 LEs. The majority of LEs belong to class
 381 “Small” (133 events), while the rest of them (4 events) belong to class “Intermediate”; no “Large”
 382 LEs have been recorded in the period of analyses (**Tab. 4**). For case B, the 131 considered
 383 phenomena have been grouped into 57 LEs, 54 “Small” and 3 “Intermediate” events (**Tab.4**). A
 384 total of 60 warnings were issued in the period of analysis; none of these were “Red”. Five warning
 385 zones received the level “Orange” and 55 zones received the warning level “Yellow”. In the period
 386 of analysis 37 different warning zones have been alerted (**Tab. 4**).

387

388 **Table 4:** Number of landslides, landslides, warning events issued and warning zones alerted in
 389 2013-2014 in the area of analysis.

	Case A	Case B
Landslide	385	131
Landslide events, LE	137	57
Small	132	54
Intermediate	5	3
Large	0	0
Warning events, WE	60	60
Warning zones alerted	37	37

390

391 4.2 Performance evaluation for the years 2013-2014

392 Two different sets of landslides have been considered in the performance of the Norwegian EWS
 393 for the Vestlandet area: Set A and Set B. The duration matrices obtained are shown in **Table 5**.
 394 Both cases refer to the years 2013-2014, thus, the sum of matrix elements is always equal to 730
 395 days.

396

397 **Table 5:** Duration matrices for cases A and B, units of time expressed in days.

CASE A		LE class			
		1	2	3	4
WE level	1	600,48	107,62	0,00	0,00
	2	9,88	8,47	1,80	0,00
	3	0,00	1,16	0,58	0,00

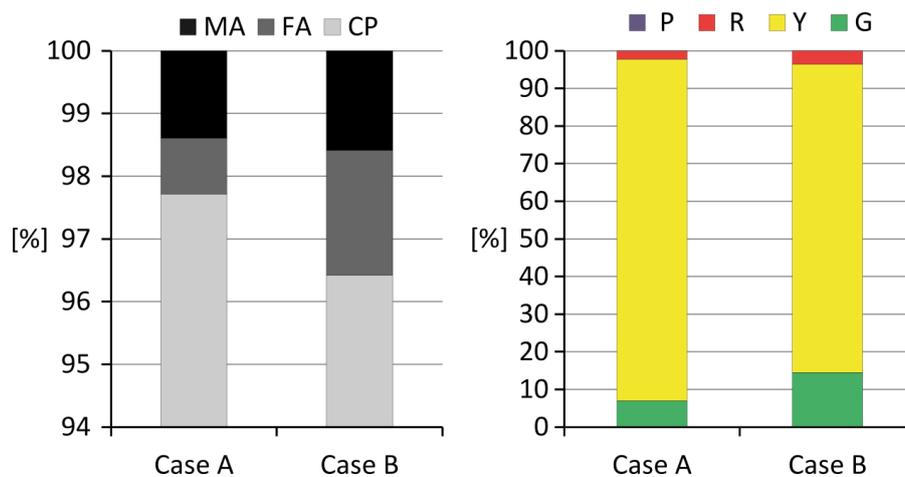
398

		4	0,00	0,00	0,00	0,00
CASE B		LE class				
		1	2	3	4	
WE level	1	671,55	36,56	0,00	0,00	
	2	11,32	7,90	0,93	0,00	
	3	1,16	0,00	0,58	0,00	
	4	0,00	0,00	0,00	0,00	

399

400 The duration matrices have been analysed considering two different performance criteria (see **Fig.**
 401 **4**). The first one is derived by a contingency table scheme (criterion 1), the other one is based on a
 402 colour code assigning a grade of correctness to each matrix cell (criterion 2). The results obtained
 403 considering criterion 1 for both Case A and B (**Fig. 6 a**) show a very high percentage of correct
 404 predictions (CPs), over 96%, and around 1,5% of missed alerts (MAs). The amount of false alerts
 405 (FAs) are 1% and 2% respectively for Case A and B. Following criterion 2 (**Fig. 6 b**) differences,
 406 among Case A and B, can be observed in terms of greens (G), that are respectively equal to 7% and
 407 14,5%, and yellows (Y) that are respectively equal to 91% and 82%. No P and just few R, equal to
 408 2,3% and 3,6%, are observed in Case A and Case B, respectively. Following criterion 1, the
 409 differences among the two cases analysed are not significant. In terms of criterion 2, Case B shows
 410 slightly higher values of G (14%) than Case A (7%). This means that considering the reduced set of
 411 landslides (Set b), there is a slightly better correspondence between the LE classes and the
 412 corresponding warning levels issued.

413



414

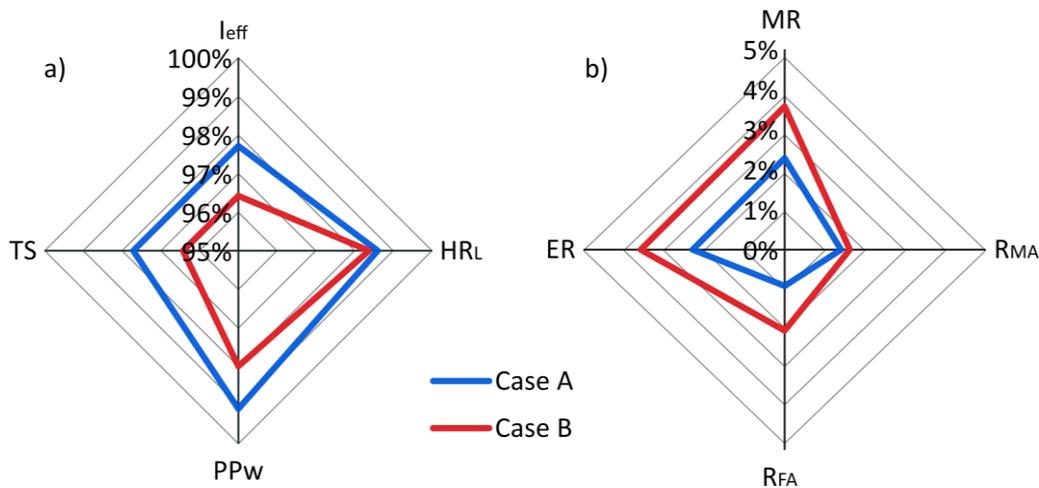
415

Figure 6: Duration matrix results in terms of: criterion 1 (a) and criterion 2 (b).

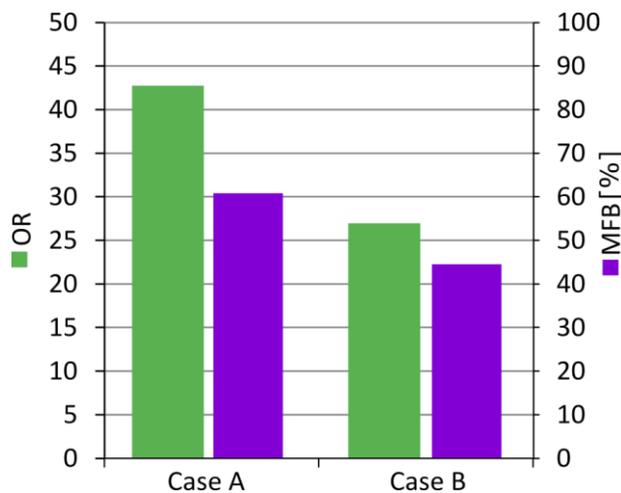
416

417 The performance indicators used to analyse the duration matrices (**Tab. 3**) are grouped into two
 418 subsets of indicators, respectively evaluating success and error (**Fig. 7**). Excluding the odds rate

419 (OR), the remaining success indicators have a percentage higher than 95% for both cases, due to the
 420 high value of CPs that is orders of magnitude higher than MAs and FAs. Therefore the OR, that
 421 indicates the correct predictions relative to the incorrect ones, assumes a very high value for both
 422 cases, although slightly higher for Case A (**Fig. 8**). The error indicators MR, ER, RMA and RFA
 423 assume very low values and the differences between the two cases are around 1% (**Fig. 7 b**). The
 424 MFB, which represents the ratio of MAs over the sum of MAs and FAs, is around 60% and 45%
 425 respectively for Cases A and B (**Fig. 8**).
 426



427
 428 **Figure 7:** Performance indicators quantifying the landslide early warning performance of Case A
 429 (in blue) and Case B (in red) in terms of: success (a) and error (b).



431
 432 **Figure 8:** Odds Ratio (OR) and Missed and False alerts Balance (MFB) performance indicators,
 433 quantifying the landslide early warning performance of Case A and Case B.

434
 435 In this performance analysis the high value of I_{eff} , (>95%) and ORs, could be interpreted as an
 436 excellent result but, in contrast, the high value of MFB highlights some issues related to the
 437 duration of MAs in relation to the total duration of wrong predictions. In general, this could be a

438 serious problem because MAs mean that no warnings or low level warnings have been issued
439 during the occurrence of one or more LEs of the highest two classes (“Intermediate” and “Large”).
440 In particular for Case A, 4 out of 5 LE of class “Intermediate” have to be considered MAs because
441 they occurred when the warning was set to level 2. Following the previous considerations, Case B
442 shows the best performance in terms of both success and error indicators, with a lower value of
443 MFB and a high value of OR. Case B uses a landslide dataset composed of rainfall-induced
444 landslides with a higher accuracy of information than Case A. As stated in Piciullo et al., (2016),
445 the result of a performance evaluation is strictly connected to the availability of a landslide
446 catalogue and to the accuracy of the information included in it.
447 Finally, it is important to stress the use of both success and error indicators to carry out a complete
448 performance analysis. As in this case, dealing with some indicators neglecting others could cause a
449 wrong evaluation of the early warning model performance. For instance, in the period of analysis,
450 no LEs of class 4 and only few LEs of class 3, occurred. However, the majority of durations of
451 these LEs have been missed. This means that the landslide early warning model was mostly able to
452 predict LEs of class “Small”. A possible solution to obtain a better model performance, reducing
453 MAs and simultaneously increasing CPs and G, could be to decrease the thresholds employed to
454 issue the warning level “High”.

455 4.3 Parametric analysis: the landslide density criterion

456 A parametric analysis on the landslide density criterion, $L_{den(k)}$, has been herein conducted with a
457 twofold purpose: to compare the performance of different early warning models, and to evaluate the
458 effect of the choices that the analyst makes when defining landslide event (LE) classes on the
459 performance indicators computed according to the EDuMaP method. The landslide density, $L_{den(k)}$,
460 represents the criterion used to differentiate among n classes of landslide events. The classes may be
461 established using an absolute (A) or a relative (R) criterion, i.e., simply setting a minimum and
462 maximum number of landslides for each class or defining these numbers as landslide spatial
463 density, i.e. in terms of number of landslides per unit area. Six landslide density criteria have been
464 considered in the performed parametric analysis (**Tab. 6**) referring to the criteria used in the
465 Norwegian EWS (**Tab. 1**). Two of them employ an absolute criterion using different numbers of
466 landslides per LE class the other four simulations, obtained considering the relative criterion, vary
467 as a function of both number of landslides and territorial extensions (10.000 km² and 15.000 km²).
468 Changing the definition of LE classes, the duration matrix and the performance indicators vary
469 because of relocation of the d_{ij} elements. In particular the $time_{ij}$ element, which is the amount of

470 time for which a level i -th warning event is concomitant with a class j -th landslide event, may vary
 471 the j -th index causing a movement of the element along the i -th row. The parametric analysis has
 472 been performed using the landslide dataset A, which includes 385 landslides. **Table 7** reports the
 473 classification of the LEs in the 6 combination of landslide density criteria.

474

475 **Table 6.** Parametric analysis: landslide density criteria considered to classify the LEs.

LE class	Absolute criterion [No. of landslides] and number of LEs		Relative criterion [No. of landslides / Area] and number of LEs			
	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
0	0	1	0	0	0	0
SMALL	1 to 4	2 to 4	(1 to 4)/15'000 km ²	(1 to 4)/15'000 km ²	(1 to 4)/10'000 km ²	(1 to 4)/10'000 km ²
INTERMEDIATE	5 to 14	5 to 18	(5 to 14)/15'000 km ²	(5 to 10)/15'000 km ²	(5 to 14)/10'000 km ²	(5 to 10)/10'000 km ²
LARGE	> 14	> 18	> 14/15'000 km ²	> 10/15'000 km ²	> 14/10'000 km ²	> 10/10'000 km ²

476

477 **Table 7.** Classification of LEs for the 6 simulations reported in table 8.

LE class	Absolute criterion [No. of landslides] and number of LEs		Relative criterion [No. of landslides / Area] and number of LEs			
	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
SMALL	124	32	132	132	133	133
INTERMEDIATE	9	9	5	3	4	4
LARGE	4	4	0	2	0	0

478

479 As an example, the simulations R-15K_{0,10} and R-15K_{0,14} differ for the definition of both LE classes
 480 Large and Intermediate. By comparing the two respective duration matrices (**Tab. 8 a, b**) a
 481 movement of the durations from d_{24} and d_{34} to respectively d_{23} and d_{33} is evident. This behaviour is
 482 due to the increase of spatial density for LE class Large, in particular from 0,67 landslides per 1000
 483 km² to 0,93 landslides per 1000 km² (**Tab. 6**), which causes a relocation of $time_{i4}$ along the rows.

484 **Table 8.** Duration matrix results for simulations R-15_{0,10}, R-15_{0,14}.

R-15K _{0,10}		LE duration (h)			
		1	2	3	4
WE duration (h)	1	600,48	107,62	0,00	0,00
	2	9,88	8,47	0,98	0,82
	3	0,00	1,16	0,00	0,58
	4	0,00	0,00	0,00	0,00

R-15K _{0,14}		LE duration (h)			
-----------------------	--	-----------------	--	--	--

485

		1	2	3	4
WE duration (h)	1	600,48	107,62	0,00	0,00
	2	9,88	8,47	1,80	0,00
	3	0,00	1,16	0,58	0,00
	4	0,00	0,00	0,00	0,00

486

487 Changes within the duration matrix mean that the value of the performance indicators may change.

488 **Table 9** presents a summary of performance indicators for all six simulations of the landslide
489 density criteria used in the parametric analysis.

490

491 **Table 9.** Performance indicators for the six simulations of landslide density criteria considered in
492 the parametric analysis.

Performance indicator	A _{0,14}	A _{1,18}	R-15K _{0,14}	R-15K _{0,10}	R-10K _{0,14}	R-10K _{0,10}
I _{eff}	0,95	0,86	0,98	0,98	0,98	0,98
HR _L	0,95	0,86	0,99	0,99	0,99	0,99
PP _w	1,00	1,00	0,99	0,99	0,99	0,99
TS	0,95	0,86	0,98	0,98	0,98	0,98
OR	18,98	6,07	42,75	42,75	49,43	49,43
MR	0,05	0,14	0,02	0,02	0,02	0,02
R _{MA}	0,05	0,14	0,01	0,01	0,01	0,01
R _{FA}	0,00	0,00	0,01	0,01	0,01	0,01
ER	0,05	0,14	0,02	0,02	0,02	0,02
MFB	1,00	1,00	0,61	0,61	0,55	0,55

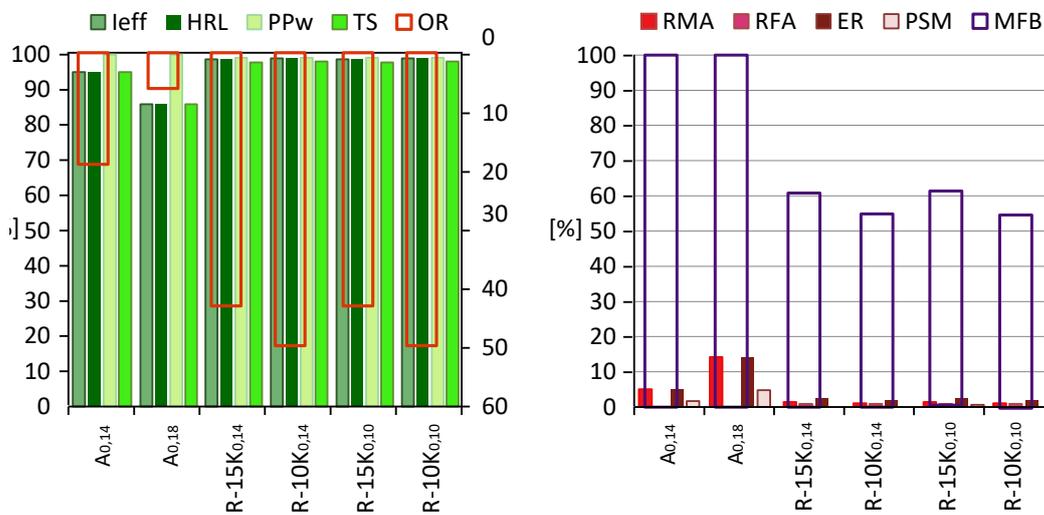
493

494 The results show similar performance for the four simulations derived using a relative criterion
495 (R15-C_{0,14} R15-C_{0,10} R10-C_{0,14} R10-C_{0,10}). The values of the success indicators are always high:
496 well above 95%, for I_{eff}, HR, TS, PP_w, while OR ranges between 42 and 49 (**Fig. 9 a**). This is due to
497 the high value of CPs compared to those of MAs and FAs, underlining a good performance of the
498 early warning model for these four simulations. In fact, also the error indicators are very low in
499 terms of percentage, around 1-2% (**Fig. 9 b**). Lower values are observed for the combination
500 obtained considering the absolute criterion, and in particular for A_{1,18}, with MR, R_{MA} and ER
501 around 14%. The MFB is generally high for all simulations denoting a bad capability of the model
502 to predict LEs of classes 3 and 4. Anyway, it must be emphasized that, considering these landslide

503 density criteria, only the simulations R-15K_{0,10}, A_{0,14} and A_{1,18} have LEs of class 4 in the period of
 504 the analysis (**Tab. 7**).

505 In conclusion, the parametric analysis shows significant differences between the absolute and
 506 relative criterion simulations. For this case study, absolute criterion simulations have lower success
 507 performance indicators, in particular for the values of odds ratio (OR) and, very high values of
 508 missed and false alert balance (MFB) compared to the performance indicators obtained for relative
 509 criterion simulations. Moreover, the absolute criterion simulations produce a number of purple
 510 errors that increase the PSM (**Fig. 9 b**).

511



512

513

514

515

Figure 9. Performance indicators related to the success (a) and to the errors (b) of the warning model, evaluated for the six simulations of landslide density criteria considered in the parametric analysis.

516

517 5. Conclusions

518 The main aim of regional landslide early warning systems is to produce alert advices within a
 519 specific warning zone and to inform local authorities and the public of landslide hazard at a given
 520 level. To evaluate the performance of the alerts issued by such systems several aspects need to be
 521 considered, such as: the possible occurrence of multiple landslides in the warning zone, the duration
 522 of warnings in relation to the time of occurrence of landslides, the level of the issued warning in
 523 relation to spatial density of landslides in the warning zone and the relative importance system
 524 managers attribute to different types of errors. To solve these issues, the EDuMaP method can be
 525 seen as a useful tool for testing the performance of regional landslide warning models. Up to now,
 526 the method has been applied exclusively to systems that issue alerts on fixed warning zones. By

527 using data from the Norwegian landslide EWS this study has extended the applicability of the
528 EDuMaP method to warning systems that uses variable size warning zones. In this study, the
529 EDuMaP method has been used to evaluate the performance of the Norwegian landslide early
530 warning system for Vestlandet (Western Norway) for the period 2013-2014. The results show an
531 overall good performance of the system for the area analyzed. Two datasets of landslide
532 occurrences have been used in this study: the first one including all the slope failures registered and
533 gathered in the NVE database within the test area; the second one excluding the phenomena whose
534 typology was either not determined or is not typically associated to rainfall. The results are not too
535 sensitive to the dataset of landslides, although slightly better results are registered with the smallest
536 (i.e. more accurate) dataset. In both cases, the high value of the MFB highlights a high number of
537 MAs compared to the FAs. A recommendation could be to have a MFB lower than 25%, which
538 means that only 1 wrong alert out of 4 is a MA. Following this reasoning, a reduction of the
539 warning level “High” is recommended in order to reduce the MAs and to increase the performance
540 of the Norwegian EWS.

541 A parametric analysis was also conducted for evaluating the performance sensitivity, to the
542 landslide density criterion, $L_{den}(k)$, used as an input parameter with EDuMaP. This parameter
543 represents the way landslide events are differentiated in classes. In the analysis the classes were
544 established considering both absolute (2 simulations) and relative (4 simulations) criteria. The
545 parametric analysis shows how the variation of the intervals of the LE classes affects the model
546 performance. The best performance of the alerts issued in Western Norway was obtained applying a
547 relative density criterion for the definition of the LE classes. The parametric analysis shows only
548 minor differences in the performance analysis among the four cases considered with the relative
549 density criteria. In conclusion, this study highlights how the definition of the density criterion to be
550 used in defining the LE classes is a fundamental issue that system managers need to be take into
551 account in order to give an idea on the number of landslides expected for each warning level over a
552 given warning zone.

553 **Acknowledgement**

554 This work was carried out during a research period of LP as visiting PhD student at NVE, Oslo. The
555 authors are grateful to two NVE employees: Søren Boje for his criticism and comments and, Julio
556 Pereira for GIS data sharing.

557 **References**

- 558 Alfieri L, Salamon P, Pappenberger F, Wetterhall F, Thielen J: Operational early warning
559 systems for water-related hazards in Europe. *Environ Sci Pol* 15(1):35–49.
560 doi:10.1016/j.envsci.2012.01.008, 2012.
- 561 Barredo, J I: Normalised flood losses in Europe: 1970–2006. *Natural Hazards and Earth*
562 *System Sciences* 9, 97–104, doi:10.5194/nhess-9-97-2009, 2009.
- 563 Badoux A, Graf C, Rhyner J, Kuntner R, McArdell B W: A debris-flow alarm system for the
564 Alpine Illgraben catchment: design and performance. *Nat Hazards* (2009) 49:517–539, doi
565 10.1007/s11069-008-9303-x, 2009.
- 566 Baum R L and Godt J W: Early warning of rainfall-induced shallow landslides and debris
567 flows in the USA. *Landslides*, 7: 259–27, doi: 10.1007/s10346-009-0177-0, 2010.
- 568 Beldring S, Engeland K, Roald LA, Sælthun NR, Voksø, A: Estimation of parameters in a
569 distributed precipitation-runoff model for Norway. *Hydrology and earth system sciences*, 7: 304-
570 316, 2003.
- 571 Bell R, Cepeda J, Devoli G: Landslide susceptibility modeling at catchment level for
572 improvement of the landslide early warning system in Norway. *Conference Proceedings of the*
573 *World Landslide Forum*. Vol. 3. 2014. 3, 2-6th June 2014, Beijing, 2014.
- 574 Boje S, Colleuille H, Cepeda J, Devoli G : Landslide thresholds at regional scale for the
575 early warning system in Norway. *Conference Proceedings of the World Landslide Forum*. 3, 2-6th
576 June 2014, Beijing, 2014.
- 577 Calvello M and Piciullo L: Assessing the performance of regional landslide early warning
578 models: the EDuMaP method. *Natural Hazards Earth System Sciences*, 16, 103–122, 2016.
579 www.nat-hazards-earth-syst-sci.net/16/103/2016/doi:10.5194/nhess-16-103-2016, 2016.
- 580 Calvello M, d’Orsi RN, Piciullo L, Paes N, Magalhaes MA, Lacerda WA: The Rio de
581 Janeiro early warning system for rainfall-induced landslides: analysis of performance for the years
582 2010–2013. *Int J Disast Risk Reduc* 12:3–15. doi:10.1016/j.ijdr.2014.10.005, 2015.
- 583 Chen C Y, Lin L Y, Yu F C, Lee C S, Tseng CC, Wang AH, Cheung KW: Improving debris
584 flow monitoring in Taiwan by using high-resolution rainfall products from QPESUMS. *Nat*
585 *Hazards*, 40: 447. doi:10.1007/s11069-006-9004-2, 2007.
- 586 Cheung, P. Y., Wong, M. C., and Yeung, H. Y.: Application of rainstorm nowcast to real-
587 time warning of landslide hazards in Hong Kong, in: WMO PWS, Workshop on Warnings of Real-
588 Time Hazards by Using Nowcasting Technology, 9–13 October 2006, Sydney, Australia, 2006.
- 589 Colleuille H, Haugen LE, Beldring S: A forecast analysis tool for extreme hydrological
590 conditions in Norway. Poster presentation, 6th World FRIEND Conference, Fez, Morocco, 2010.
- 591 CRED: EM-DAT. In: The OFDA/CRED International Disaster Database, Universite´
592 Catholique de Louvain, Brussels, Belgium. www.emdat.be, 2011.

593 Devoli G, Kleivane I, Sund M, Orthe N-K, Ekker R, Johnsen E, Colleuille H: Landslide
594 early warning system and web tools for real-time scenarios and for distribution of warning
595 messages in Norway, in: Engineering Geology for Society and Territory “Landslide Processes”,
596 Proc. XII International IAEG Congress, Torino, Italy, 625–629, doi:10.1007/978-3-319-09057-
597 3_104, 2014.

598 Easterling D R, Meehl G A, Parmesan C, Changnon S A, Karl T R, Mearns L O: Climate
599 extremes: observations, modeling, and impacts. *Science*, Vol. 289, Issue 5487, pp. 2068-2074, doi:
600 10.1126/science.289.5487.2068, 2000.

601 Ekker R, Kværne K, Os A, Humstad T, Warttinen A, Eide V, Hansen RK: regObs – public
602 database for submitting and sharing observations. Proceedings of International Snow Science
603 Workshop, 7-11th October 2013 Grenoble - Chamonix Mont-Blanc, 2013.

604 Etzelmüller B, Romstad B, Fjellanger J: Automatic regional classification of topography in
605 Norway. *Norwegian Journal of Geology*, 87: 167-180, 2007.

606 European Commission: Directive 2000/60/EC of the European Parliament and of the
607 Council of 23 October 2000 establishing a framework for Community action in the field of water
608 policy, 2000.

609 European Commission: Directive 2007/60/EC of the European Parliament and of the
610 Council of 23 October 2007 on the assessment and management of flood risks, 2007.

611 European Environment Agency: Mapping the impacts of natural hazards and technological
612 accidents in Europe – an overview of the last decade. EEA Technical Report No 13/2010, isbn: 978-
613 92-9213-168-5, 2010.

614 Fell R, Ho K K S, Lacasse S, and Leroi E: A framework for landslide risk assessment and
615 management, in: *Landslide Risk Management*, edited by: Hungr, O, Fell, R, Couture, R, and
616 Eberhardt, E, Taylor and Francis, London, 3–26, 2005.

617 Fischer L, Rubensdotter L, Stalsberg K, Melchiorre C, Horton P, Jaboyedoff M: Debris flow
618 modeling for susceptibility at regional to national scale in Norway. In Eberhardt et al. *Landslides
619 and Engineered Slopes: Protecting Society through Improved Understanding*, p. 723-729, 2012.

620 Furseth A: Slide accidents in Norway. Oslo, Tun Forlag, 207 p. (in Norwegian), 2006.

621 Gariano, S. L., Brunetti, M. T., Iovine, G., Melillo, M., Peruccacci, S., Terranova, O.,
622 Vennari, C., and Guzzetti, F.: Calibration and validation of rainfall thresholds for shallow landslide
623 forecasting in Sicily, southern Italy, *Geomorphology*, 228, 653–665,
624 <http://dx.doi.org/10.1016/j.geomorph.2014.10.019>, 2015.

625 Glade T and Nadim F: Early warning systems for natural hazards and risks, *Nat. Hazards*,
626 70, 1669–1671, doi:10.1007/s11069-013-1000-8, 2014.

627 Godt, J. W., Baum, R. L., and Chleborad, A. F.: Rainfall characteristics for shallow
628 landsliding in Seattle, Washington, USA, *Earth Surf. Proc. Land.*, 31, 97–110,
629 doi:10.1002/esp.1237, 2006.

630 Hyogo Framework for Action: Building the Resilience of Nations and Communities to
631 Disasters, World Conference on Disaster Reduction 18–22 January 2005, Kobe, Hyogo, Japan, 22
632 pp., 2005.

633 Huggel C, Khabarov N, Obersteiner M, Ramirez J M: Implementation and integrated
634 numerical modelling of a landslide early warning system: a pilot study in Colombia. *Nat Hazards*
635 (2010) 52:501–518. doi:10.1007/s11069-009-9393-0, 2010.

636 Hungr O, Leroueil S, Picarelli L: The Varnes classification of landslide types, an update.
637 *Landslides* 11(2):167–194. doi:10.1007/s10346-013-0436-y, 2014.

638 ICG: Guidelines for landslide monitoring and early warning systems in Europe - Design and
639 required technology. Project Safe Land "Living with landslide risk in Europe: Assessment, effects
640 of global change, and risk management strategies". D4.8, 153p. Available from:
641 <http://www.safeland-fp7.eu/results/Documents/D4.8.pdf>, 2012.

642 Jaedicke C, Lied K, Kronholm K: Integrated database for rapid mass movements in Norway.
643 *Natural Hazards and Earth System Sciences*, 9: 469-479, doi:10.5194/nhess-9-469-2009, 2009.

644 Lagomarsino, D, Segoni, S, Fanti, R, and Catani, F: Updating and tuning a regional-scale
645 landslide early warning system. *Landslides*, 10, 91–97. doi:10.1007/s10346-012-0376-y, 2013.

646 Lagomarsino, D., Segoni, S., Rosi, A., Rossi, G., Battistini, A., Catani, F., and Casagli, N.:
647 Quantitative comparison between two different methodologies to define rainfall thresholds for
648 landslide forecasting, *Nat. Hazards Earth Syst. Sci.*, 15, 2413–2423, doi:10.5194/nhess-15-2413-
649 2015, 2015.

650 Martelloni, G., Segoni, S., Fanti, R., and Catani, F.: Rainfall thresholds for the forecasting of
651 landslide occurrence at regional scale, *Landslides*, 9, 485–495, doi:10.1007/s10346-011-0308-2,
652 2012.

653 Morss, R, Wilhelmi, O, Meehl, G, Dilling, L: Improving societal outcomes of extreme
654 weather in a changing climate: an integrated perspective. *Annual Review of Environment and*
655 *Resources*, 36, 1–25, doi: 10.1146/annurev-environ-060809-100145, 2011.

656 NOAA-USGS Debris Flow Task Force: NOAA-USGS debris-flow warning system. Final
657 report, US Geological Survey Circular 1283, US Geological Survey, Reston, Virginia, USA: 47.
658 Available at: <http://pubs.usgs.gov/circ/2005/1283/pdf/Circular1283.pdf> (last access: November
659 2014), 2005.

660 Osanai N, Shimizu T, Kuramoto K, Kojima S, Noro T: Japanese early-warning for debris
661 flows and slope failures using rainfall indices with Radial Basis Function Network, *Landslides*, 7,
662 325–338, doi:10.1007/s10346-010-0229-5, 2010.

663 Piciullo L, Siano I, Calvello M: Calibration of rainfall thresholds for landslide early warning
664 purposes: applying the EDuMaP method to the system deployed in Campania region (Italy). In:
665 *Proceedings of the International Symposium on Landslides 2016-Landslides and Engineered*
666 *Slopes. Experience, Theory and Practice*. Napoli, Italy, 3:1621–1629. ISBN 978–1–138-02988-0,
667 2016a.

668 Piciullo L, Gariano S L, Melillo M, Brunetti M T, Peruccacci S, Guzzetti F., Calvello M:
669 Definition and performance of a threshold-based regional early warning model for rainfall-induced
670 landslides. *Landslides*. doi:10.1007/s10346-016-0750-2, 2016b.

671 Ramberg IB, Bryhni I, Nøttvedt A, Rangnes K: The making of a land – geology of Norway.
672 Trondheim. Norsk Geologisk Forening, 624 p., 2008.

673 Rossi, M, Peruccacci, S, Brunetti, M T, Marchesini, I, Luciani, S, Ardizzone, Balducci, S V,
674 Bianchi, C, Cardinali, M, Fiorucci, F, Mondini, A C, Reichenbach, P, Salvati, P, Santangelo, M,
675 Bartolini, D, Gariano, S L, Palladino, M, Vessia, G, Viero, A, Antronico, L, Borselli, L, Deganutti,
676 A M, Iovine, G, Luino, F, Parise, M, Polemio, M, Guzzetti, F: SANF: National warning system for
677 rainfall-induced landslides in Italy, in: *Landslides and Engineered Slopes: Protecting Society
678 through Improved Understanding*, edited by: Eberhardt, E., Froese, C., Turner, K., and Leroueil, S.,
679 Taylor& Francis, London, 1895–1899, isbn 978-0-415-62123-6, 2012.

680 Segoni, S., Rossi, G., Rosi, A., and Catani, F.: Landslides triggered by rainfall: a
681 semiautomated procedure to define consistent intensity-duration thresholds, *Comput. Geosci.*, 3063,
682 123–131, <http://dx.doi.org/10.1016/j.cageo.2013.10.009>, 2014.

683 Segoni S, Battistini A, Rossi G, Rosi A, Lagomarsino D, Catani F, Moretti S, Casagli N:
684 Technical note: an operational landslide early warning system at regional scale based on space–
685 time-variable rainfall thresholds. *Nat Hazards Earth Syst Sci* 15:853–861. doi:10.5194/nhess-15-
686 853-2015, 2015.

687 Stähli M, Sättele M, Huggel C, McARDell BW, Lehmann P, Van Herwijnen A, Berne A,
688 Schleiss M, Ferrari A, Kos A, Or D, Springman SM: Monitoring and prediction in early warning
689 systems for rapid mass movements. *Nat Hazards Earth Syst Sci* 15:905–917. doi:10.5194/nhess-15-
690 905-2015, 2015.

691 Staley D M, Kean J W, Cannon S H, Laber J L, Schmidt K M: Objective definition of
692 rainfall intensity-duration thresholds for the initiation of post-fire debris flows in southern
693 California. *Landslides* 10: 547-562. doi: 10.1007/s10346-012-0341-9, 2013.

694 Thiebes B, Glade T, Bell R: Landslide analysis and integrative early warning-local and
695 regional case studies, in: *Landslides and Engineered Slopes: Protecting Society through Improved
696 Understanding*, edited by: Eberhardt, E., Taylor& Francis Group, London, 1915–1921, isbn 978-0-
697 415-62123-6, 2012.

698 Tiranti D, and Rabuffetti D: Estimation of rainfall thresholds triggering shallow landslides
699 for an operational warning system implementation. *Landslides*, 7: 471–481, doi:10.1007/s10346-
700 010-0198-8, 2010.

701 United Nations Inter-Agency Secretariat of the International Strategy for Disaster Reduction
702 (UN ISDR): *Global Survey of Early Warning Systems: An assessment of capacities, gaps and
703 opportunities towards building a comprehensive global early warning system for all natural hazards*,
704 available at:[http://www.unisdr.org/2006/ppew/info-resources/ewc3/Global-Survey-of-Early-
705 Warning-Systems.pdf](http://www.unisdr.org/2006/ppew/info-resources/ewc3/Global-Survey-of-Early-Warning-Systems.pdf) (last access:November 2014), 2006.

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