

1 **On the drought in the Balearic Islands during the hydrological year 2015-2016**

2 Climent Ramis, Romualdo Romero, Víctor Homar, Sergio Alonso, Agustí Jansà and Arnau
3 Amengual

4 Meteorology Group. Department of Physics. University of the Balearic Islands- 07122 Palma.
5 Spain.

6

7 **Abstract**

8 During the hydrological year 2015-16 (September to August) a severe drought affected the
9 Balearic Islands, with substantial consequences (alleviated partially by desalination plants) on
10 water availability for consumption from reservoirs and aquifers and also on the vegetation
11 cover. In particular, a plague of 'Xylella fastidiosa' reached a relatively alarming level for the case
12 of the almond and olive trees. The expansion of this infestation could be attributed to, or at least
13 favored by, the extreme drought. In this paper we analyze this anomalous episode in terms of
14 the corresponding water balance in comparison with the balance obtained from long-term
15 climatological data. It is shown that the drought was the result of a lack of winter precipitation,
16 the lowest in 43 years, which led to a shortage of water storage in the soil. In several
17 meteorological stations analyzed, evaporation was greater than precipitation during all the
18 months of the year. In terms of attribution, it is found that during the 2015-16 winter the
19 atmospheric circulation over the North Atlantic was largely westerly and intense, with high
20 values of the NAO index that were reflected in high pressures over the Iberian Peninsula and the
21 western Mediterranean.

22

23 Keywords: drought, Balearic Islands, water balance, Mediterranean Sea.

24

25 **1. Introduction.**

26

27 The Balearic Islands are located in the central part of the Western Mediterranean basin (Fig. 1).
28 The archipelago presents a well-marked interannual variability in the annual precipitation as it
29 was shown by Homar et al. (2010). Within this interannual variability, a particularly severe
30 drought episode occurred during the hydrological year (September to August) 2015-2016.
31 Actually, the drought affected the eastern part of the Iberian Peninsula as it was reported by the
32 Spanish Meteorological Agency (AEMET,
33 http://www.aemet.es/es/serviciosclimaticos/vigilancia_clima/). However, we restrict this study
34 to the Balearic Islands, where the population of perennials suffered a remarkable mortality,
35 especially among almond, olive and other fruit trees. Mostly in the southern part of the
36 archipelago shrubs and other plants such as bushes and steppes also perished, especially young
37 individuals which have very shallow roots. In addition, a plague of 'Xylella fastidiosa' reinforced
38 after the summer of 2016, and this could be attributed to, or at least favored by, the drought
39 and further hydrological stress suffered by the almond and olive trees. Although it is difficult to
40 assess quantitatively the total losses resulting from the drought (these could have reached more
41 than 10 million euros in the livestock breeding owing to loss of up to 90% of the production of
42 forage, according to the media), different lines of funding were issued by the regional

43 government. Besides the impacts on the natural and agricultural systems, the demand of water
44 for personal and leisure consumption reached its historical maximum during the summer of
45 2016 ($36.5 \times 10^6 \text{ m}^3$ during August 2016 in Mallorca, according to media sources), when the
46 islands registered a record number of tourists to date (more than 10.9 million in Mallorca; media
47 sources). All together left the reservoirs and aquifers of the islands at levels of great concern,
48 putting at serious risk the supply for the following months in case of drought persistence without
49 the help of the desalination plants. This severe drought can be framed in the context of the
50 observed increase in the frequency of droughts in the Mediterranean area (Hoerling et al. 2012)
51 and in particular in the Spanish eastern lands (Vicente-Serrano et al. 1994).

52 The lands surrounding the north, east, and west of the Mediterranean Sea have a climate that
53 is characterized by a mild and rainy winter and a warm and dry summer. According to the
54 classification of Köppen these are thus considered to have a Csa type climate (Peel et al. 2007).
55 This type takes the generic name of Mediterranean climate. The Köppen classification global
56 map is determined from gross climatic features; when analyzing the data at higher resolution,
57 noticeable differences are found, even between contiguous areas of reduced extent. The
58 Balearic Islands (Fig. 1), with a typical Mediterranean climate, is a specific example of a context
59 exhibiting notable climatic differences within a relatively small region. Given the size of the
60 islands (Mallorca, the largest, extends over 3640 km^2), among all the influencing factors we must
61 attribute to the orography the greatest part of observed climatic differences over the territory.
62 These contrasts are indeed quite accentuated in the archipelago. The four major islands of the
63 Balearics have similar patterns of mean monthly rainfall but the spatial distribution of annual
64 totals is quite heterogeneous. Menorca and Ibiza-Formentera show a remarkable spatial
65 uniformity, with mean annual values higher in Menorca than in Ibiza-Formentera (Guijarro 1986,
66 Jansà 2014, López et al. 2017). These wetter conditions are attributed to the higher latitude of
67 Menorca, being the island more frequently affected by the fronts linked to the low-pressure
68 disturbances that evolve through Central Europe and by the lows developed over the Genoa
69 Gulf. In Mallorca there is high spatial contrast in the mean annual distribution of precipitation.
70 Along the southern coasts, where the orography is practically absent, annual precipitation
71 values are of the order of 350 mm on average, while in the zones with the highest mountains
72 (Tramuntana range, heights up to 1500 m, see Fig. 1) in the northwest of the island, the average
73 annual rainfall reaches 1400 mm (Guijarro 1986). These large contrasts occur within a distance
74 of about 50 km. In fact, attending to the climatic characteristics of the south of Mallorca, it rather
75 conforms to BSk type from the classification of Köpen, that is, winters not excessively dry,
76 temperate, but with very dry and torrid summers. The northern and northeastern zones of
77 Mallorca receive precipitations of the same order as those of Menorca, once again clearly above
78 the accumulations of the southern region.

79 Another characteristic of the rainfall over the Balearic Islands is its marked seasonality. The
80 ombrothermic diagram for the Mallorca airport (Fig. 2; Jansà et al., 2016) shows the most
81 outstanding feature of the Mediterranean climate: the above mentioned scarcity of
82 precipitation during the summer as well as relatively high temperatures during this period of the
83 year; also that autumn and winter are mild and relatively wet. The fact of reaching the end of a
84 hot summer after two months with almost no precipitation, somehow characterizes the type of
85 vegetation present in the lowlands (pines, shrubs, bushes and steppes but also almond trees).
86 At the same time, the islands have an economy fundamentally dependent on tourism (in 2016,
87 Balearic airports received 36.8 million passengers, according to the official web pages of the
88 three airports) that is mainly concentrated in the summer months. The supply of drinking water
89 during this period depends critically on underground aquifers (and on the supplementary action

90 of desalination plants) since existing reservoirs in the rainiest mountainous area of Mallorca are
91 too small. After the long and extreme summer, the recovery of the aquifers is strongly
92 determined by the amount of rainfall received during the autumn and the following winter. The
93 flora will be subjected to greater or lesser hydric stress depending mainly on the behavior of
94 autumn rainfall. The occurrence of large water stress situations is not uncommon, given the high
95 interannual variability that characterizes annual precipitations in the Balearic Islands (Homar et
96 al., 2010). Extreme manifestations of such variability are not new; there are written references
97 about important droughts affecting the archipelago during the Middle Age (Barceló, 1991), as
98 well as many oral references to the hazardous drought occurred during 1912-1913 in Mallorca,
99 a time when the local economy was almost exclusively dependent on agriculture.

100 Given the strong water deficit imposed to the vegetation by the end of the summer and also the
101 natural cycle of the underground aquifers, it may be more suitable to analyze precipitation in
102 terms of the hydrological year (September to August). Additionally, in order to account for the
103 vegetation stress in more detail, it becomes more informative to calculate the annual water
104 balance in which precipitation and evaporation are presented together (considering for the
105 latter the potential evapotranspiration) and to compare it with the climatic water balance for
106 which the local vegetation has adapted.

107 This paper presents in section 2 the interannual variability of the precipitation regime in the
108 Balearic Islands, both from the standard and hydrological year perspectives, as well as the
109 climatic water balance of the region. Section 3 discusses the water balance for the hydrological
110 year 2015-16 in detail. In section 4 the circulation pattern of the exceptional context that led to
111 the severe drought of that year is analyzed and compared with the pattern of an illustrative wet
112 year. Finally, section 5 presents the main findings and conclusions of the study.

113

114 **2. Precipitation variability and climatic water balance.**

115

116 Monthly precipitation values at Mallorca, Menorca and Ibiza airports from 1973 to 2016 (44
117 years) have been analyzed. These are the longest climatic series without gaps in the Balearic
118 Islands. From the monthly values, annual accumulations as well as those corresponding to the
119 43 hydrological years from 1973-74 to 2015-16 have been calculated.

120 The anomalies of the annual rainfall with respect to the average of the reference period 1981-
121 2010 for the airports of Mallorca, Menorca and Ibiza have been considered (not shown). The
122 yearly mean for the reference period at Mallorca is 411.3 mm and the interannual variability of
123 the series is large enough as to yield a standard deviation of 100.9 mm (coefficient of variation
124 $CV=24.5\%$). The average for Menorca airport is 548.6 mm and the standard deviation is 132.8
125 mm ($CV=24.2\%$). These values for Ibiza are 411.1 mm and 117.3 mm ($CV=28.5\%$). These
126 relatively large values of the coefficients of variation reveal the high interannual variability of
127 precipitation in the islands, which is itself related to the variability of the atmospheric patterns,
128 as shown in Section 4.

129 As revealed by the CV values, the variability is greater in Ibiza than in Menorca, although there
130 are anomalies in both stations that occasionally exceed 200 mm. It is noteworthy the relatively
131 low correlation (0.54) that exists between the time series of Mallorca and Menorca, but
132 especially low is the correlation between the time series of Menorca and Ibiza (0.30). For
133 Mallorca and Menorca there are few cases in which a positive anomaly in one station does not

134 correspond to the same sign in the other. One of such cases is 2016, when the intense rainfall
135 recorded in Mallorca during the months of October and December (107.6 and 150.4 mm,
136 respectively) explains the positive anomaly of its airport; however, this event did not affect
137 Menorca (13.2 mm and 79.8 mm, respectively).

138 It should be noted that a wet/dry year at the airports tends to be accompanied by greater/lower
139 than normal annual precipitation in the rest of each respective island. Figure 3 shows this kind
140 of distribution for the years 1996 and 1999, considered as wet and dry years, respectively.
141 However some kind of objective index should be applied to analyze the representativeness of
142 the interannual variability of the rainfall captured by the airports, especially in Mallorca where
143 the spatial variability of the annual rainfall is very high as previously indicated. An analysis of the
144 spatial representativeness of the interannual variability captured by the Mallorca airport has
145 been performed using two methodologies. First, the time series of the relative annual anomalies
146 (anomaly divided by the corresponding annual average) have been calculated for five
147 meteorological stations located in Mallorca, and the resulting mean time series (of the five
148 individual series) has been determined. The five stations are representative of different
149 pluviometric regimes of the island: mountainous area, north, center, east and south. For this
150 analysis, the period 1981-2010 has been considered. The time series of annual relative
151 anomalies at Mallorca airport has been compared against the above mean time series. The time
152 series exhibit a correlation coefficient as high as 0.9. The second method is analogous to the
153 previous one but uses the precipitation analyses across the island of Mallorca that were derived
154 in the PREGRIDBAL project (López et al. 2017). These analyses have a resolution of 100 m and
155 use all available observed data for each product requested. Annual precipitation grid data has
156 been considered for each of the years 1980-2009, together with the grid analysis of mean
157 precipitation corresponding to these 30 years. For each grid point and for each year the relative
158 annual anomalies have been determined and a time series expressing the spatial average of
159 annual anomalies has been calculated. Finally, this time series has been compared against the
160 relative anomalies at Mallorca airport, yielding in this case a correlation coefficient of 0.86 (Fig.
161 4). Thus, it seems well justified the assumption that the spatial-temporal variability in the island
162 of Mallorca is correctly captured by the series of precipitations at the Mallorca airport.

163 Due to their relatively small size and moderate orography, the spatial variabilities of the annual
164 mean precipitation in Menorca and Ibiza are much lower than in Mallorca, therefore it seems
165 clear that the corresponding time series at the airports are even more representative of the
166 corresponding interannual variability of the whole islands.

167 Figure 5 shows the precipitation anomalies at the airports of Mallorca, Menorca and Ibiza for
168 the hydrological years 1973-74 to 2015-16 (43 years) with respect to the reference period 1980-
169 81 to 2009-10. Recall the hydrological year comprises from September to August. The mean
170 precipitation for the reference period in Mallorca is 409.5 mm, with a standard deviation of
171 119.2 mm (CV=29.1%). For Menorca these values are 544.3 mm and 120.5 mm (CV=22.1%). For
172 Ibiza 413.0 mm and 116.6 mm (CV=28.2%) respectively. Mean values are very similar with
173 respect to the observations derived from the "standard" or natural years but the interannual
174 variability is higher now in Mallorca and lower in Menorca. In the present case there is a greater
175 correlation (0.68) between the anomalies of these two rainfall stations. In Ibiza the values are
176 very similar to those obtained for the natural year. The correlation between the time series of
177 Menorca and Ibiza is identically low (0.33 vs 0.30 for the natural years). These low correlations
178 values are a clear manifestation that the rain bearing meteorological systems for the north and
179 south of the archipelago do not respond to the same circulation patterns, as previously reported

180 by Guijarro (2002, 2003). A detailed study on the surface circulation related with daily rainfall
181 patterns in Mallorca can be found in Sumner et al (1995).

182 Looking at Fig. 5, it can be observed that dry hydrological years leading to water stress on the
183 flora, and probably on the aquifers, become clearly distinguishable. The periods 1981 to 1984,
184 1991 to 1994 and to 1998 to 2001 are noteworthy. It can be observed that in 2015-16 there are
185 also negative anomalies, much more important in Menorca.

186 Although there are several indices to characterize a drought (e.g. the Palmer Drought Severity
187 Index (PDSI), Palmer (1965); the Standardized Precipitation Index (SPI), McKee et al. (1993);
188 Supply Demand Index (SDDI), Rind et al. (1990)), from an ecological point of view and in order
189 to account for the possible water stress on the flora, it is interesting to analyze the water balance
190 directly, in which the precipitation is compared against the evaporation, month by month, and
191 from this balance to evaluate the periods of the year in which there is an excess or lack of water
192 in the soil. In this sense there are studies on the effects of droughts on the Mediterranean flora
193 in Spain (e.g. Peñuelas et al. 2001). The determination of the potential evapotranspiration (PET)
194 is an important step when estimating soil water deficit or excess. However, empirical formulas
195 for estimating PET have their limitations, the results cannot be considered at the same level of
196 exactitude as precipitation measurements. In consequence, the comparison between
197 precipitation and PET has to be regarded as an approximation to the reality. The existence of
198 several analytical expressions to calculate PET using different variables, also demonstrate the
199 difficulty to determine this magnitude accurately.

200 Estimation of the climatic water balances at the three airports was carried out using the
201 Thornthwaite method (1948) for the determination of monthly potential evapotranspiration,
202 using monthly mean temperature and precipitation values referred to the reference period
203 1981-2010. In our analysis, actual evaporation is considered to coincide with calculated potential
204 evapotranspiration if monthly precipitation is greater than potential evapotranspiration, and in
205 these circumstances the remaining precipitation is converted to water stored in the soil. These
206 amounts can be cumulative through the year and if the total storage reaches a value which is
207 considered to be the maximum capacity of the soil, the excess becomes surface runoff and
208 infiltration. The maximum storage of the soil depends on several factors, e.g. the texture, land
209 use and slope of the terrain. Botey and Moreno (2015) have produced a map of the soil
210 maximum storage for the Iberian Peninsula and the Balearic Islands. From the information
211 displayed in their map, for the low lands of the Balearic Islands, where the used meteorological
212 stations are located, 100 mm can be considered a reasonable value. If the monthly precipitation
213 is less than the potential evapotranspiration, the actual evaporation is equal to the precipitation
214 plus the reserve portion of the soil moisture that is needed, until it is exhausted. The remaining
215 difference between potential evapotranspiration and actual evaporation is indicative of the
216 water deficit that has to be overcome by vegetation. Balance calculations begin in the month of
217 September, considering that the soil does not contain any water after the dry summer.

218 Figure 6 shows the climatic water balance (1981-2010) during the hydrological year, according
219 to the indicated method, for the airports of Mallorca, Menorca and Ibiza. Climatologically, there
220 is deficit in Mallorca for the first month of September. There is storage of water in the soil from
221 October to February, which is totally exhausted by the end of June. During the summer (June-
222 August) the deficit is very large, reaching 150 mm. At the Menorca airport there is also a deficit
223 in September, the accumulation of water in the soil begins in October, and there is runoff or/and
224 infiltration during January, February and March. The water stored in the soil of Menorca allows
225 for evaporation to be larger than precipitation even in June, with a total lack of soil water

226 observed only in July and August. The maximum deficit also reaches 150 mm. At Ibiza the water
227 balance is very similar to Mallorca but the storage of water in the soil during the winter is lower
228 and therefore it is consumed more quickly, inducing a large deficit during all the summer.

229 The climatic water balance at Menorca and Ibiza airports can be considered representative of
230 the whole islands. In contrast, for the larger and more complex island of Mallorca it is evident,
231 bearing in mind Fig. 3 and the results of Guijarro (1986) and Jansà (2014), that the water balance
232 of the airport cannot, in any way, be extended to the whole island. The water balance shown is
233 representative of the south of Mallorca. It is also indicative of the situation in the western and
234 eastern coastal zones and in the center of the island, although the latter zone tends to store a
235 little more of water in the soil during the winter as consequence of the higher precipitation
236 (recall Fig. 3). For the northern and northeastern zones of Mallorca the water balance is
237 expected to be much more similar to that at the Menorca airport, as the rainfall regimes are
238 quite similar in monthly distributions and amounts. In the mountainous area of Mallorca the
239 water balance is certainly very different to that at the airport, as the climatological annual
240 precipitation is almost four times greater. In this zone there are two reservoirs dedicated to the
241 supply of water to the population, which of course rely on the regular runoff of the fall and
242 winter. In any case, some drought also exists on the mountains during the summer, since
243 precipitation in this season is basically absent as in low lands.

244 In order to validate the previous water balance in terms of precipitation and potential
245 evapotranspiration, the results of a more sophisticated method have been examined. The data
246 provided at the web site <https://wci.earth2observe.eu/portal/> (which collects data from the
247 European Earth2Observe project) at the three grid points (resolution 0.25°) closest to the
248 airports of Mallorca, Menorca and Ibiza have been obtained. Monthly total precipitation values
249 (PCP) for each grid point have been extracted from 1981 to 2010 and the mean monthly values
250 have been computed (these monthly data originally come from the ECMWF data set). Monthly
251 total values of evapotranspiration (EVT; i.e. surface evaporation, interception and transpiration)
252 and monthly total runoff (R; i.e. surface runoff, sub-surface flow and deep percolation) provided
253 by the 8 available models have also been obtained from the Earth2Observe web site. For each
254 model and variable, the mean monthly values with reference to 1981-2010 have been
255 calculated. Finally, the 8 models-ensemble mean and inter-model standard deviation of the
256 previous monthly values were obtained. With these values the water balance was estimated at
257 each of the three mesh points considered. This balance is built as the precipitation minus the
258 actual evapotranspiration minus the losses ($WB = PCP - EVT - R$).

259 These results reveal:

260

261 a) For the mesh point near the Mallorca airport (components of the water balance and the
262 balance itself are displayed in Fig. 7) the precipitation values used by the models are much higher
263 than those observed at the airport. As an example, the observed mean annual value (1981-2010)
264 is 411.3 mm, while the same rainfall product used by the models is 597.4 mm. Regarding EVT,
265 the model ensemble mean values are higher than those obtained at the Mallorca airport for PET
266 using the Thornthwaite formula, especially in summer. The monthly standard deviations are very
267 high (that is, large differences among the different models). The PET for the Mallorca airport lies
268 within the ensemble spread region. Regarding the water balance, and accepting 100 mm as
269 saturation threshold for the soil, saturation in the Earth2Observe data is obtained during
270 December, January and February that may be due to the high precipitation values ingested in
271 the models. Dryness is obtained in July and August and very low water reserve values in June

272 and September. In the former results a remarkable water deficit is obtained in September (Fig
273 6), because the temperatures are still high.

274 b) For the grid point near Menorca airport (not shown), the monthly values of precipitation used
275 by the models are much more similar to those observed at the airport (registered annual average
276 of 548.6 mm versus 601.2 mm in the models). The EVT shows a behavior similar to that at the
277 grid point near the airport of Mallorca: values are greater than those of PET obtained from the
278 Menorca airport data using the Thornthwaite expression, and there is a large spread among the
279 8 models. The calculated PET values are also well encompassed by the ensemble dispersion
280 band. Regarding the water balance, saturation of the soil is obtained in January and February
281 and a value close to saturation in December. Dryness is also obtained in July and August. These
282 results are very similar to those obtained directly in our study for the Menorca airport.

283 c) For the grid point close to the Ibiza airport (not shown), the average monthly
284 precipitation values used by the models are also significantly higher than those registered at the
285 Ibiza airport (observed annual mean of 411.1 mm versus 497.2 mm in the models). Again the
286 average EVT values of the ensemble are larger than PET values given by the Thornthwaite's
287 expression at the Ibiza airport. The inter-model spread is very high. Regarding the water balance,
288 saturation of the soil is not reached in any month; in contrast dryness is present during May,
289 June, July and August. These results are in agreement with the results obtained directly with the
290 airport data.

291

292 In conclusion, it seems that the simple method used in the paper is sufficient to obtain a clear
293 representation of the drought object of the study.

294

295

296 **3. Hydrologic Year 2015-16.**

297

298 As already mentioned, the hydrological year 2015-16 was characterized by a negative anomaly
299 with respect to the reference period (Fig. 5). Other hydrological years exhibit greater negative
300 anomalies, but it was the widespread deficit of precipitation during 2015-16 what characterizes
301 the hazardous effects of this drought event. Figure 8 presents the hydrological balance for
302 Mallorca, Menorca and Ibiza airports corresponding to that hydrological year. For these water
303 balances, daily potential evapotranspiration has been calculated using the Hargreaves method
304 (Hargreaves and Samani, 1985). The monthly values have been obtained from the daily values.
305 The distribution of rainfall shows significant accumulations in September, due to the convective
306 rains that affected the islands (176.4 mm in Mallorca, 181.1 in Menorca and 139.6 in Ibiza). It is
307 also evident the quite low rainfall recorded during the rest of the hydrologic year, particularly
308 during the rest of the autumn and the whole winter. At Mallorca airport the precipitation during
309 November 2015 to January 2016 was 25.6 mm, which represents the lowest value among the
310 43 considered hydrologic years (Fig. 9). Similarly, the total precipitation recorded during
311 December 2015 was 0.2 mm, the lowest of this month for the whole period 1973-2016. At
312 Menorca airport the accumulated precipitation from November 2015 to January 2016 was 45
313 mm, also the lowest quantity recorded in a hydrologic year. The precipitation for December
314 2015 was 2.1 mm, again the minimum record for this month during the period 1973-2016. In
315 Ibiza the situation was similarly extreme, since 35.2 mm was the precipitation recorded for

316 November-January, the lowest for the 43 analyzed hydrologic years, and only 0.7 mm were
317 registered in December 2015 (only surpassed by the 0.2 mm recorded in 1974).

318 For the Mallorca airport (Fig. 8) it is observed that already during the month of October, the
319 water that was stored in the soil as consequence of the heavy precipitation events of September,
320 was already consumed; during the rest of the year there is deficit. The lack of precipitation
321 during the winter months implies a very dry soil when the sunny days and rise of temperatures
322 establish in spring.

323 Something similar happens in Ibiza, where the water deficit starts a bit later than in Mallorca as
324 a consequence of the rainy early autumn (September and October) but where the abnormal lack
325 of winter rains is also quite remarkable. In Menorca the situation is to some extent similar: the
326 deficit begins in March, although the winter precipitation was also very scarce.

327 The Thornthwaite approach applied for obtaining PET monthly climatic values uses directly the
328 monthly mean temperatures provided by AEMET for the period 1981-2010. The aim is to build
329 a reference water balance for a comparison with the particular water balance of the hydrological
330 year 2015-16. For this hydrological year, the PET monthly values have been calculated from the
331 daily values obtained by the Hargreaves method. Some comparison between the two methods
332 for this year is necessary to fully justify the reference to the climatic water balance. A
333 comparison between both methods was made. Specifically, monthly PET values using
334 Thornthwaite method were calculated for the hydrological years 2014-15 and 2015-16 at the E1
335 site in Mallorca (see Fig. 1). Analogous monthly values were obtained from the daily PET values
336 given by the Hargreaves formula. The E1 station is located in the most arid region of the island.
337 The two time series show a correlation coefficient of 0.9 (see Fig. 10). For the warmer/colder
338 months the Thornthwaite method reveals larger/lower monthly PET values than the other
339 approach. In any case, given the high value of the correlation coefficient, the obtained reference
340 or climatic water balance can be effectively compared with the one calculated for the 2015-16
341 hydrological year.

342 Comparing the water balances of 2015-16 (Figure 8) with the climatic water balances (Figure 6)
343 at the three airports, notable differences during the fall and winter are found. In the climatic
344 balance the beginning of autumn shows a water deficit that is rapidly reversed during the rest
345 of autumn and winter. Winter rains develop the reserves for the ground, since the summer is
346 really dry. Only at the Menorca airport this storage exceeds the 100 mm threshold and therefore
347 surface runoff and infiltration are produced. The lack of rainfall in the Balearic Islands, especially
348 during the extreme winter of 2015-16, gives an idea, when analyzed in terms of the water
349 balance, of the hydrological stress to which the local vegetation was subjected. This deficit of
350 precipitation during the winter in the Mediterranean area has been related to some more
351 general droughts observed in Europe (Vautard et al. 2007).

352 It is interesting to display some other areas of Mallorca that were affected by a still more intense
353 drought, again in terms of their water balances. Figure 11 shows the water balance for 2015-16
354 obtained from the data at three automatic meteorological stations located in the south, central
355 and northern parts of Mallorca (see Fig. 1). It can be observed that at the southernmost station
356 (E1) the precipitation throughout the year was lower than the potential evapotranspiration,
357 motivating that the water deficit was accumulating during the whole hydrological year. The
358 intense rains that affected the airport location in September did not occur in this area. The lack
359 of precipitation in winter is remarkable. The accumulated drought that reached the always dry
360 summer was very severe and had dramatic consequences on the vegetation types possessing

361 shallow roots. But also on some trees, especially almond trees, whose fruit maturation had to
362 develop under absolutely unfavorable conditions.

363 The precipitation regime in the north of Mallorca (E3) was very similar to that of the south
364 region. Rainfall was also lower than potential evapotranspiration during all the months of the
365 hydrological year. In the center of the island (E2) the situation was not very different, although
366 during the month of September the precipitation was enough to surpass the potential
367 evapotranspiration. The rainfall and the evaporation regimes resemble those at the airport. The
368 convective rains of early fall also reached the center of the island, but the profound lack of
369 rainfall in winter was a constant that is repeated at all locations, supposing that evaporation
370 rates permanently exceed precipitation, a feature clearly divergent from what is climatologically
371 expected.

372 The hydrological year 2015-16 was characterized by very intense rainfall events in September
373 followed by a continuous deficit for the rest of the rainy period. This begs the question of the
374 role of runoff, especially when the season starts with heavy precipitation, on soil dried out by
375 the summer; in these conditions much less water will infiltrate and thus recharge soil moisture.
376 There is an added problem for an accurate computation of the water balance when
377 measurements on the runoff are not available, as it is the case of our study. There are very few
378 measurements of runoff in streams of the Balearics, all of them belonging only to special
379 campaigns and always before 2014. Note that in the Balearic Islands there are not permanent
380 rivers. In addition, no information about the episode can be obtained from the Earth2Observe
381 web page, since model data extends only till 2012.

382 However, the runoff coefficient for a nearby stream basin to the Mallorca airport (few
383 kilometers away) was estimated by García et al. (2017), based on observed stream flows for the
384 1977-2009 period. The estimated runoff coefficient was as low as 0.03. This result ensures that
385 the conversion of precipitation into surface runoff is quite low for this nearby basin.
386 Furthermore, no substantial changes are found in the spatial distribution of the physiography
387 and hydrology of the stream basin where the meteorological station is located. We can safely
388 assume that almost all precipitation is infiltrated and that the P-E balance is quite realistic when
389 assessing the climatic water balance for the 1981-2010 period. Even for the heavy precipitation
390 event at the end of the 2015 warm season, most of precipitation would have infiltrated owing
391 to the high infiltration capacity of the soil and its low water content.

392

393 **4.- Circulation patterns.**

394

395 During the winter of 2015-16 the North Atlantic was especially active cyclonically speaking.
396 Many deep depressions developed above 45^o-50^o of latitude and affected Europe. Particularly
397 deep was the impact on Ireland and England, especially in December, where very intense rains
398 (up to 200% of the climatic value referred to 1981-2010 for that winter (McCarthy et al., 2016))
399 resulted in floods. The substantial westerly flow also advected warm air along that latitude belt
400 and the mean winter climatic temperature values (period 1981-2010) were largely exceeded in
401 Ireland and England, up to two degrees in the south of England. In December this warm anomaly
402 in the south of England reached 5 degrees (McCarthy et al., 2016). This situation was caused by
403 a strong zonal circulation of the jet stream over the North Atlantic; the jet basically pointed
404 directly to Ireland from the coasts of America during that winter (Burt and Kendon 2016). For

405 latitudes below 50^o, the westerly flow was also maintained during that winter. Figure 12 shows
406 the average geopotential structure at 500 hPa for Europe and the Mediterranean for November
407 2015 to January 2016. High geopotential values over the Iberian Peninsula and the western
408 Mediterranean that extend towards Central Europe are evident. For these months the NAO
409 index was 3.56 for November, 4.22 for December and 1.16 for January
410 (<https://crudata.uea.ac.uk/~timo/datapages/naoi.htm>), thus reflecting a strong westerly
411 circulation.

412 The above meteorological situation is essentially unfavorable for any significant occurrence of
413 rainfall in the western Mediterranean and particularly in the Balearic Islands. The most favorable
414 rainfall conditions in the islands are linked with the evolution of cyclonic disturbances at mid-
415 upper tropospheric levels which give rise to secondary depressions at surface over the
416 Mediterranean and easterly moist flows impinging over the Balearic Islands (Romero et al.
417 1999). Atlantic disturbances crossing central Europe, even involving active fronts, generally
418 produce little precipitation along the Spanish Mediterranean coast and in the Balearic Islands,
419 in any case just affecting the northern half of the islands. Figure 13 shows that during the months
420 of November 2015 to January 2016, when the precipitation in the Balearics was practically null,
421 there was a strong positive anomaly of geopotential at 500 hPa over the western
422 Mediterranean, a circulation pattern entirely inhibiting the generation of any type of
423 precipitation system.

424 It was previously reported that during September 2015 intense precipitation happened on all
425 three islands. Figure 10 shows that the atmospheric circulation during this month was
426 characterized by the presence of lows at 500 hPa, indicated by the nucleus of negative anomaly
427 affecting Western Europe and the Western Mediterranean. This pattern is dynamically favorable
428 for the generation of heavy rainfall situations slightly downstream, over the Spanish
429 Mediterranean coast and the Balearic Islands (Romero et al., 1999).
430 The average conditions displayed in Figure 13 show the radical change of the circulation that
431 occurred between September and November 2015. The pattern of September would
432 correspond, at low levels, with the persistence of meridional flows over the north Atlantic and
433 low NAO values (-1.65 for September and -1.13 In October), the opposite pattern found during
434 the period from November 2015 to January 2016. The occurrence of rainfall in the Balearic
435 Islands could be better correlated with high values of the Scandinavian Index (September 1.09,
436 October 0.62, November -1.4, December 0.08, January -0.68, normalized to the period 1981-
437 2010; <http://www.cpc.ncep.noaa.gov/data/teledoc/scand.shtml>).

438 As a contrasting situation, the hydrological year 2008-09 can be considered a wet case (see
439 Figure 5). During the months of November to January, 214 mm at the Mallorca Airport, 303 mm
440 at the Menorca Airport and 187 mm at the Ibiza Airport were recorded. Figure 11 shows the
441 geopotential anomaly at 500 hPa from November 2008 to January 2009. A notable negative
442 anomaly centered over the western Mediterranean can be observed, resulting in a completely
443 opposite pattern to that of 2015-16 (Figure 14). The values of the NAO index for these months
444 were negative or low (November -1.30, December -0.58, January 0.6).

445

446 **5. Conclusions.**

447 The characteristics of the recent drought that occurred in the Balearic Islands during the 2015-
448 16 hydrological year (September to August) have been presented. The analysis was carried out
449 in terms of the particular hydrologic balance for this year using data from six meteorological

450 stations to determine the potential evapotranspiration and to estimate the actual evaporation.
451 These water balances have been compared against those corresponding to the long-term
452 climatic conditions for the reference period 1981-2010. Comparison of the climatic water
453 balance calculated with the empirical expressions against the balance deduced from 8 models
454 used by the European Earth2Observe project show some differences. Most of these differences
455 have to be attributed to the greater values of precipitation ingested in the models and the high
456 variability of the simulated evaporation and runoff. However, the calculated values of PET lie
457 within the spread interval of the models.

458

459 The analyzed hydrologic year reveals a profound precipitation deficit during the winter, such
460 that the potential evapotranspiration surpassed the precipitation practically the whole year,
461 except in September when at some stations the precipitation exceeded the evaporation. The
462 recorded precipitation from November 2015 to January 2016 was the lowest for this period at
463 the three airports of the Balearic Islands for the 43 considered hydrologic years. The
464 precipitation of December was also unappreciable in all the islands. Accordingly, the soil could
465 not store any water to face the spring, when insolation hours and temperatures increased. This
466 resulted in the lack of any water reserves during 2015-16, an aspect totally anomalous compared
467 with an average winter, for which certain levels of moisture can be maintained in the soil until
468 June in Mallorca and Ibiza, and until July in Menorca.

469

470 We verified that the meteorological situation during the anomalous 2015-16 winter was
471 dominated by a very marked westerly flow over the North Atlantic, with high values of the NAO
472 index. This situation caused intense precipitations and anomalously warm temperatures in
473 Ireland and England. On the contrary, precipitations at lower latitudes, and particularly in the
474 western Mediterranean, were very scarce.

475

476 The identification of anomalous circulation patterns in seasonal or climate prediction models
477 can be a mechanism for anticipating drought situations and stimulate planning and mitigation
478 measures in a region like the Mediterranean, where water demand is high, especially at the
479 time of the year when precipitation is scarce. It is also a promising line of research for purposes
480 of agricultural planning and conservation of the current vegetation.

481

482 **Acknowledgments**

483 Temperature and precipitation data were recorded and provided by the Spanish Meteorological
484 Agency (AEMET). The weather analyses correspond to the NCEP/NOAA reanalysis database
485 (<https://www.esrl.noaa.gov/psd/cgi-bin/data/composites/printpage.pl>). Figure 3 comes from
486 the PREGRIDBAL project (<http://pregridbal-v1.uib.es/>). References to media correspond to
487 Diario de Mallorca. This research was sponsored by CGL2014-52199-R (EXTREMO) project, which
488 is partially supported with FEDER funds, an action funded by the Spanish Ministerio de Economía
489 y Competitividad.

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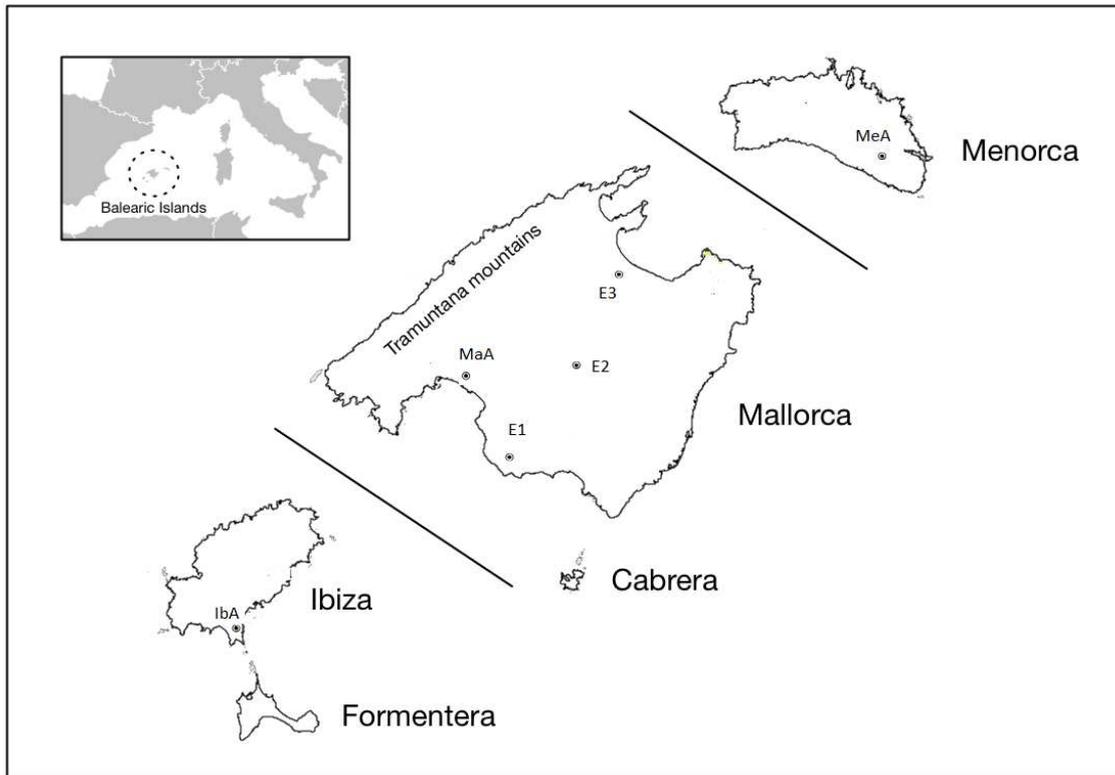
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558 Figure 1.- The Balearic Islands. MaA, Mallorca airport; MeA, Menorca airport; IbA, Ibiza airport.
559 Locations of the other climatological stations analyzed in the text are also indicated.

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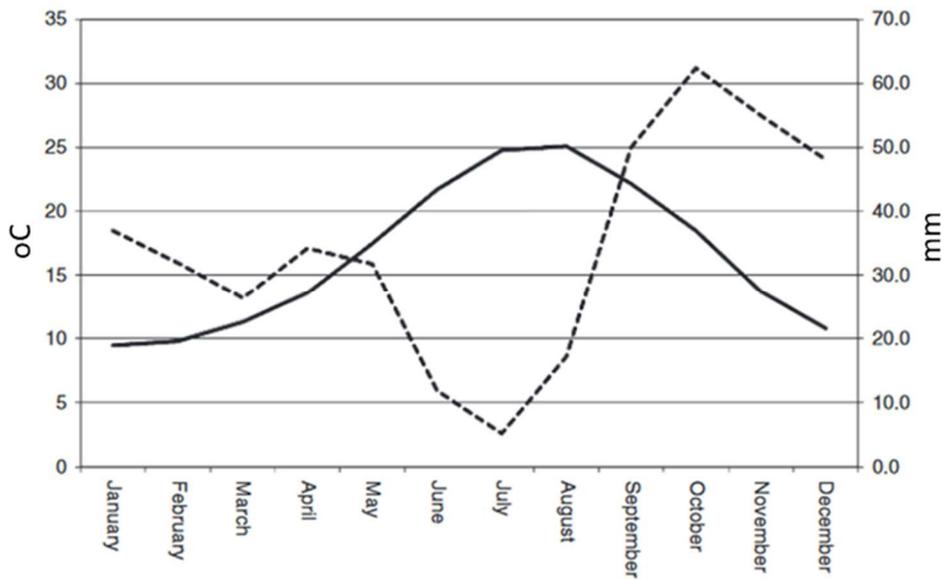
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576 Figure 2. Ombrothermic diagram (Gausson 1955) for Mallorca airport (1981-2010) (after Jansà
577 et al., 2016). Continuous line: mean temperature. Dashed line: mean precipitation.

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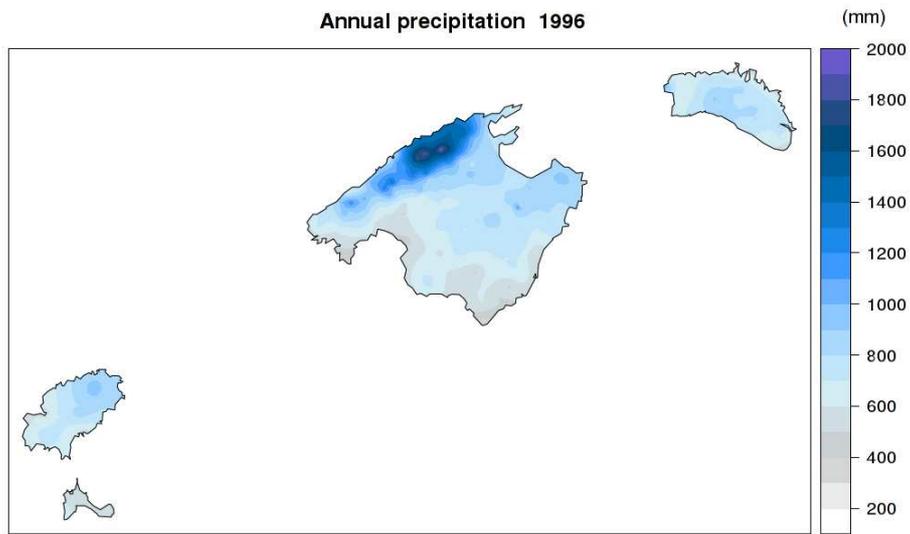
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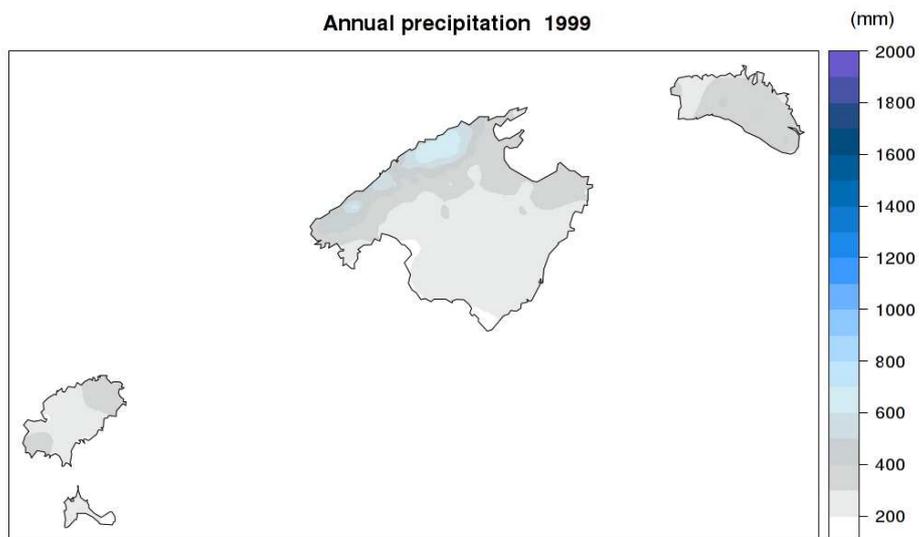
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587 Figure 3.- Spatial distribution of accumulated precipitation for 1996 (wet year) and 1999 (dry
 588 year). The same scale is used. (from <http://pregridbal-v1.uib.es/>).

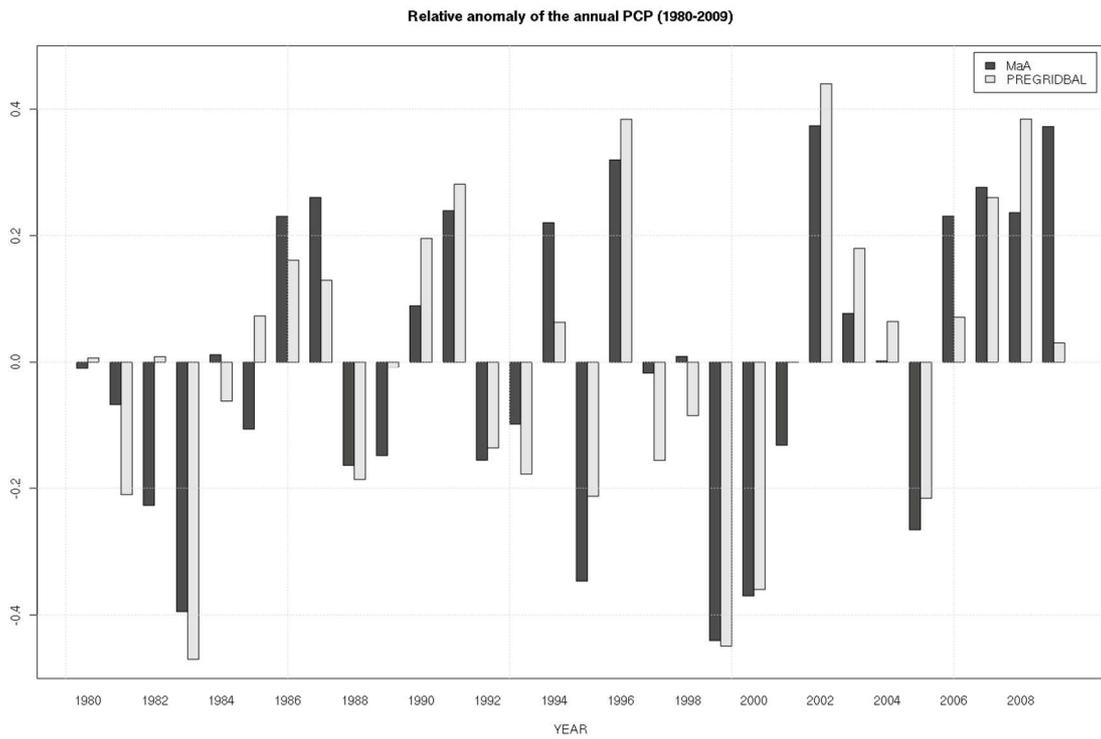
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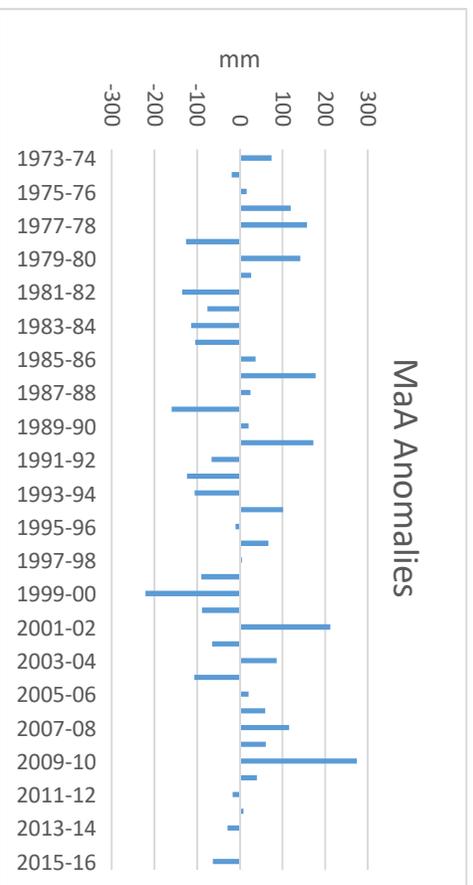
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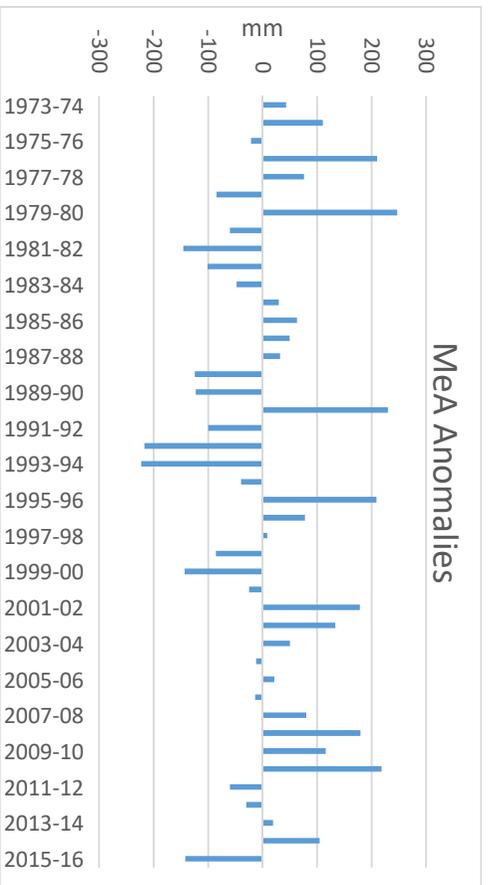
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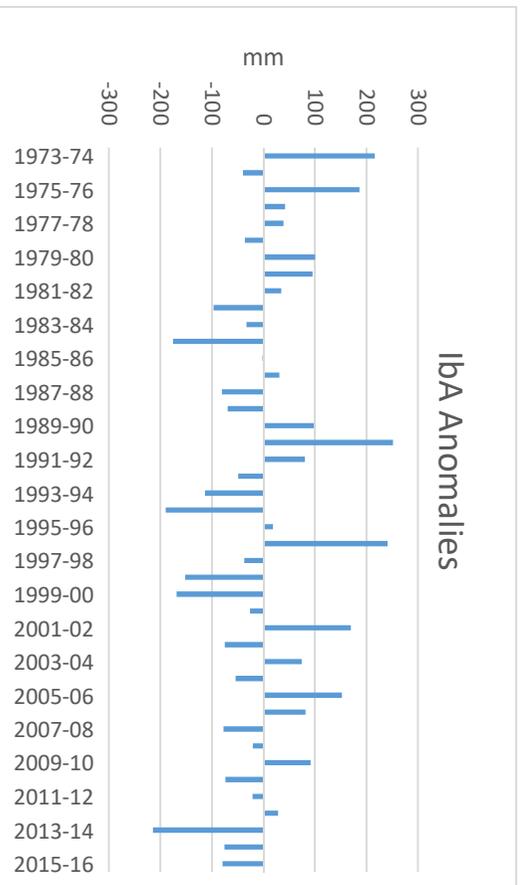
597 Figure 4.- Time series of the relative annual precipitation anomalies at the Mallorca airport and
 598 for Mallorca as a whole derived from the PREGRIDBAL project.



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Figure 5.- Anomalies of the precipitation for the hydrological year at the airports of Mallorca Menorca and Ibiza with respect to the respective averages calculated for the reference period 1980-81 to 2009-10.

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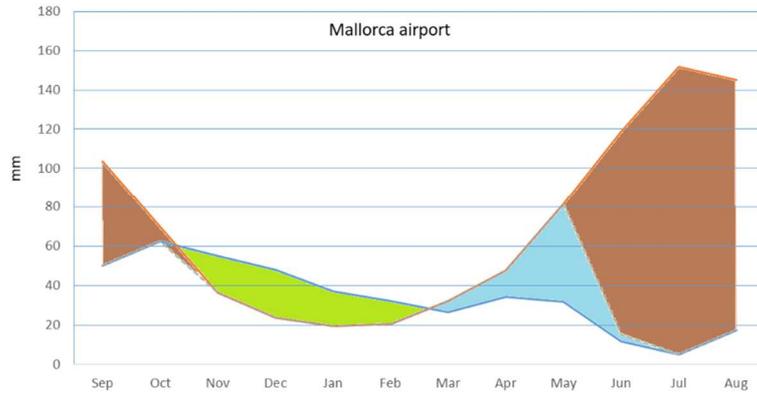
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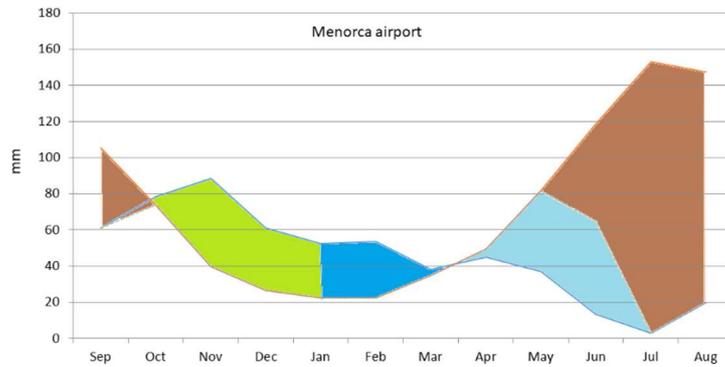
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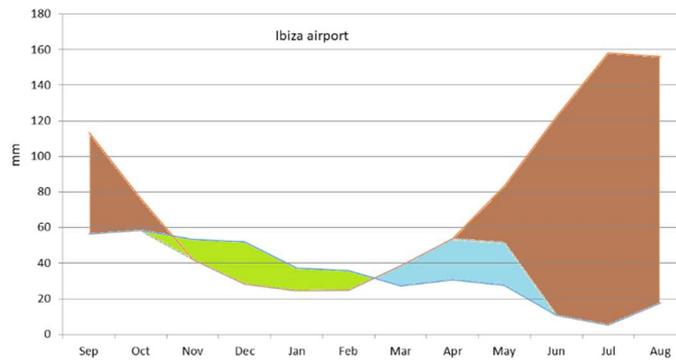
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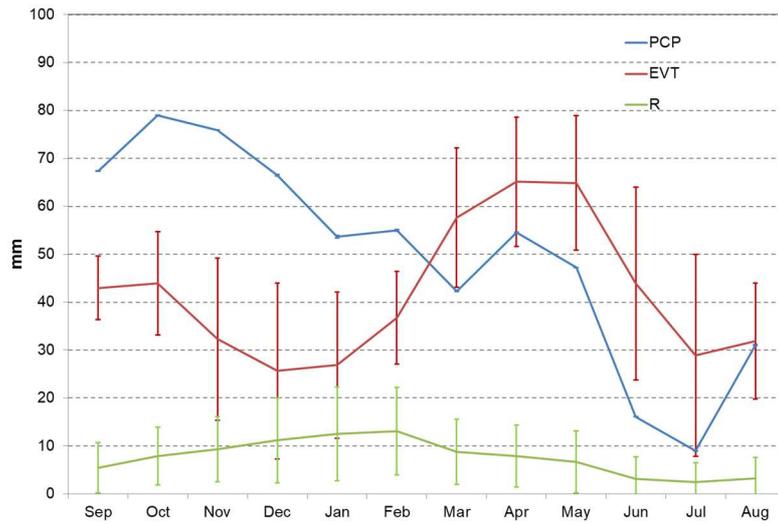
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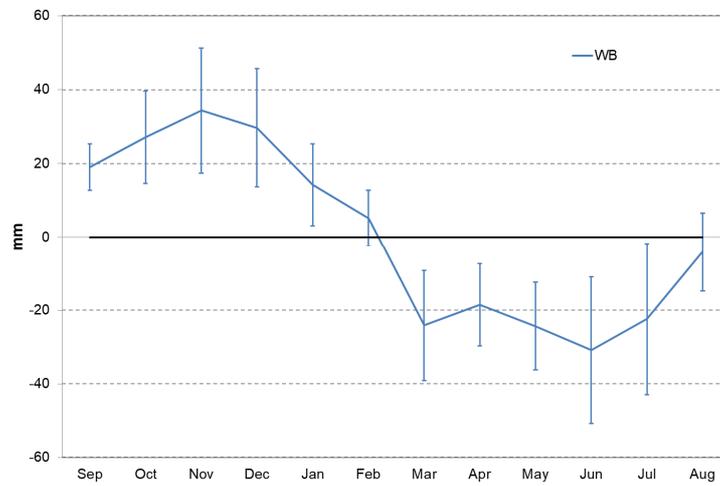
612 Figure 6.- Climatic water balance (1981-2010) at the airports of Mallorca, Menorca and Ibiza
613 (MaA, MeA and IbA in Fig. 1). Lines indicate: blue, precipitation (mm); brown, potential
614 evapotranspiration (mm); dashed green, evaporation (mm). Colored areas indicate: green color,
615 accumulation of water in the soil; cyan, evaporation of water stored in the soil; blue, runoff;
616 brown, water deficit in the soil.

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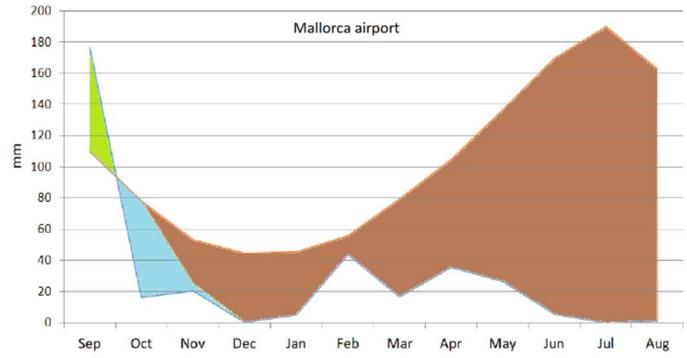
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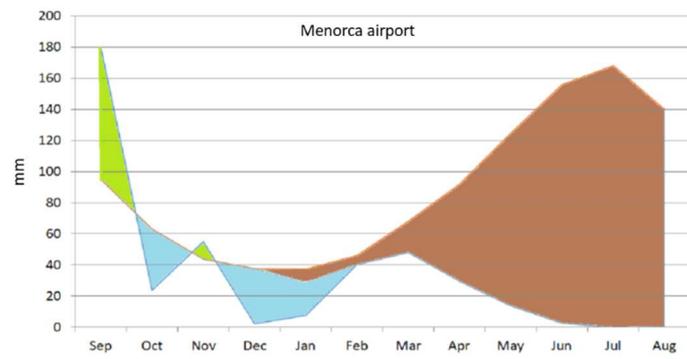
622 Figure 7.- Top: Components of the climatic water balance (1981-2010) at the grid point nearest
 623 to the Mallorca airport, deduced using data from the European Earth2Observe project. Vertical
 624 bars represent standard deviation among the 8 models. Bottom: Water balance at the same grid
 625 point (see text for details).

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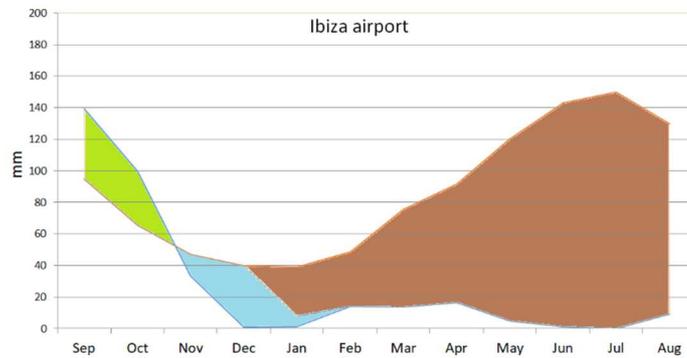
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Figure 8.- As in Figure 6 but for the hydrologic year 2015-16.

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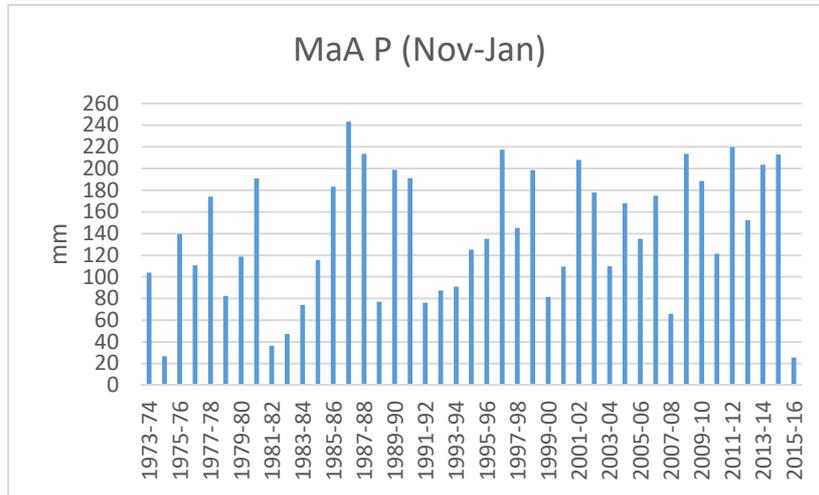
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643 Figure 9.- Accumulated precipitation from November to January at the Mallorca airport.

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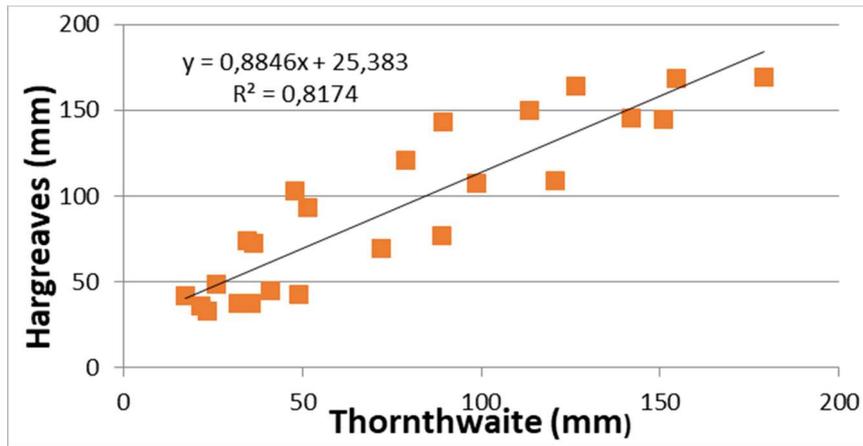
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662 Figure 10.- Comparison of the monthly PET obtained by the Thornthwaite and Hargreaves
 663 methods at E1 station (see Fig. 1) for the hydrological years 2014-15 and 2015-16.

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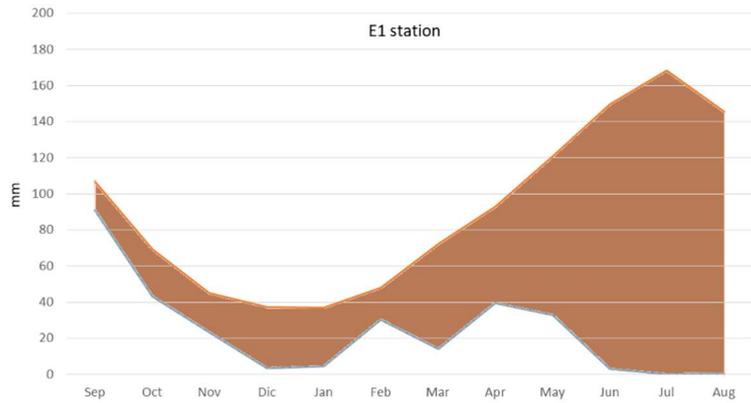
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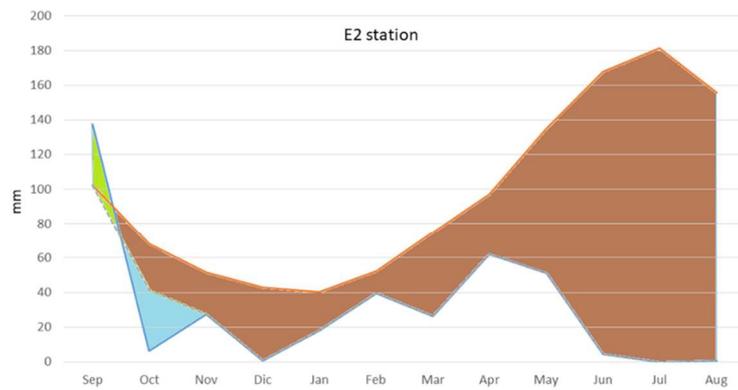
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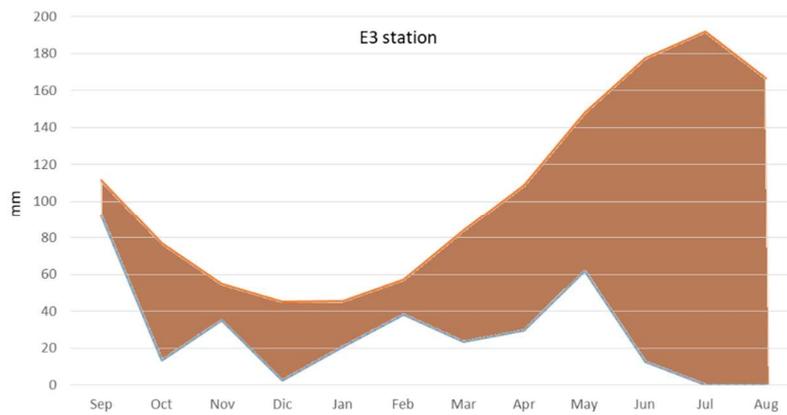
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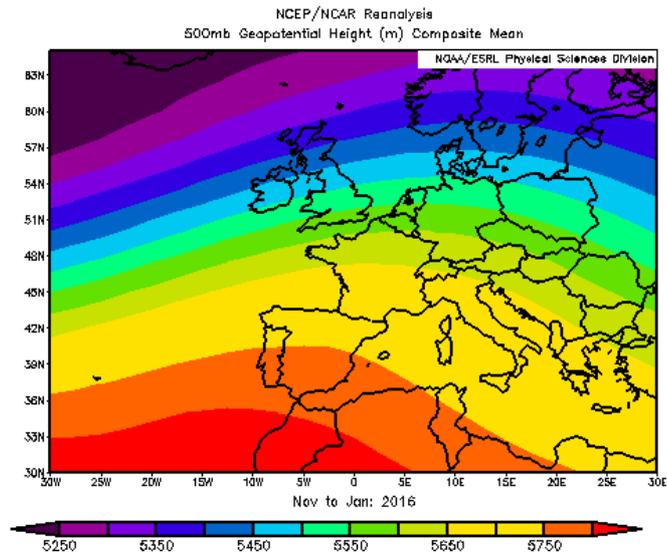


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683 Figure 11.- As in Figure 8 but for the three additional locations in Mallorca (E1, E2 and E3 in
 684 Figure 1, respectively).

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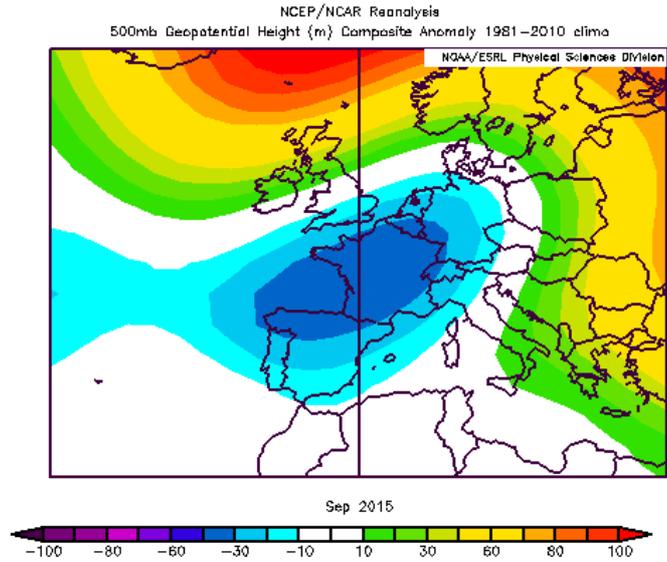


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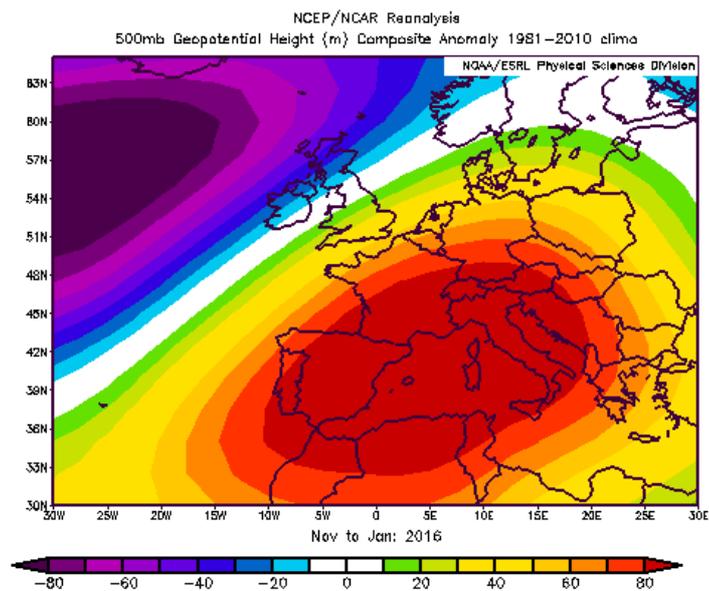
688 Figure 12.- Mean geopotential height at 500 hPa for November 2015 - January 2016 (source
689 NCEP/NOAA reanalysis)

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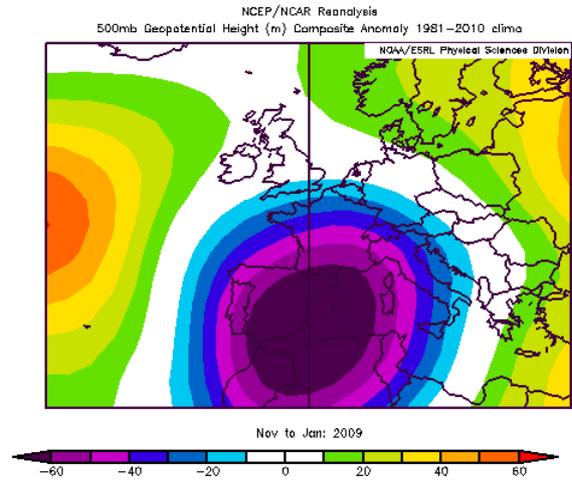


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695 Figure 13.- Geopotential height anomalies at 500 hPa for September 2015 and for November
696 2015 - January 2016, referring to the reference period 1981-2010. (source NCEP / NOAA
697 reanalysis)

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702 Figure 14.- Geopotential height anomalies at 500 hPa for November 2008 - January 2009 with
703 respect to the reference period 1981-2010. (source NCEP/NOAA reanalysis).

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