

Reply to JL Macías Comments

Please indicate in section 7.2 Lava Flow Scenarios which program was used to simulate lava flows it is not mentioned in the text. Done

Figures:

Figure 1. Legend Change the colour of the 2011-2012 eruption, it is not visible or it is too small?

The box of the Timanfaya eruption in the legend does not match that of the map
The triangle for the 2011-2012 was too small, we have enlarged the symbol.
We are not sure if you refer to the red box of the inset. We only wanted to show with this inset the location of Lanzarote (not Timanfaya) within the Canary Archipelago, as it is written in the Figure Caption.

Figure 2. Legend In the box labels you mixed eruptions with deposits so you need to define them in a homogeneous way for instance: Historical eruptions (1824) or Lava flows and pyroclastic (1824 eruption). We have corrected the legend
What is the meaning of subhistorical?? Those Holocene eruptions that took place before the last 600 years. We have modified it in the legend

Fig. 2b caption mention the diameter of the crater. Done

Fig. 2b and 2c please indicate the orientation of the photographs. Done

Figure 4. change obtained in a NE-SW area. for obtained along a NE-SW oriented area. Correction Done

Please also note the supplement to this comment:

<http://www.nat-hazards-earth-syst-sci-discuss.net/nhess-2017-2/nhess-2017-2-RC1-supplement.pdf>

Lines are referred to the new corrected manuscript

Line 24: correction done

Line 26: we do not understand why the word visitors, has been crossed out, we have left it.

Line 28: we have decided to leave “forget about” since the meaning is the same one

Line 29: we have left the sentence as it is since it does not change the meaning

Lines 43, 44: correction done

Line 84: correction done

Line 111: correction done

Lines 143, 149, 153: correction done

Line 164: correction done

Lines 239, 240: correction done

Lines 253, 254, 255: corrections done

Lines 265, 266: corrections done

Line 271: we have included the tool used for simulation lava flows (VORIS 2.0.1)

Line 276: correction done

Line 288: corrections done

Reply to JL Sophie Mossoux

Specific Comments:

How do you explain the difference that can be noticed between this manuscript and the susceptibility map published in 2013 by Bartolini et al.? Why did you had to produce a new version?

The previous susceptibility map created by Bartolini et al 2013 is only an example of the capabilities of QVAST. They used the geological information available at that time in the literature. In this work we have done a comprehensive analysis of the volcano-structural and additional geological information, adding new structural data and also the stress field model for the island, therefore obtaining a most complete susceptibility map.

Explain why it has been decided to only focus, for some hazards, to specific scenarios instead of hazard assessment on the whole island. The ash fallout is in the paper limited on one specific scenario meanwhile the susceptibility map could be used to extent the analysis to the whole island. Lava flows have been simulated for the whole island.

The particularity of each process (hazard), as for example ashfall, forces to simulate individual scenarios. Ashfall process does not depend on the topography (DEM). Therefore, it is not possible to use the volcanic susceptibility map as base map for simulating ashfall. It would be necessary to do almost 150000 simulations that correspond to the number of pixels of the susceptibility model. All these simulations together would give us a superposition of many plumes that would cover the entire island...so not having much sense. For this reason we decided to simulate ashfall only in the highest probability vent assuming a Strombolian eruption similar to 1824 eruption. This approach has been used also in other long-term volcanic hazard assessment for ashfall hazard (see Cioni et al., 2003; Orsi et al., 2004; Rolandi 2010)".

How would you combine all the hazards together?

In this case we have preferred to show individual scenarios that can be useful during volcanic crises for Civil Protection, since each hazard has to be managed in a different way during an emergency. Nevertheless, if we have to combine all of them, we will follow the methodology that we used for El Hierro Island (Becerril et al., 2014). We combined the most probable scenarios to create a qualitative hazard map of El Hierro constructed from the combination of all them. We did map algebra and distinguished four levels of hazard, from very low to high hazard, depending on the number of individual hazards that overlapped on each point (pixel) of the map (Becerril et al. (2014); <http://www.nat-hazards-earth-syst-sci.net/14/1853/2014/nhess-14-1853-2014.pdf>).

TECHNICAL CORRECTIONS/TYPING ERRORS

TITLE

- wouldn't it more suitable to write "hazards". Three different hazards are studied in the paper.

Correction done

L13-14: Lanzarote is an active volcanic island that has hosted the largest (>1.5 km³ DRE) and longest (6 years) eruption, the Timanfaya eruption, on the Canary Islands in historical times (last 600 years):

- the largest and the longest compared to?

It is referred to the rest of the Archipelago. It is written in the text: Lanzarote is an active volcanic island that has hosted the largest (>1.5 km³ DRE) and longest (6 years) eruption, the Timanfaya eruption, **on the Canary Islands in historical times** (last 600 years).

L14: the Timanfaya eruption

- give the exact date of the eruption

Correction done

L21: rational land planning

- would you mean rational or national?

Rational (Based on or in accordance with reason or logic)

L24: ... the main aspects...

- give more information about what you mean with “main aspect”

We have added to the text the following sentence: such as the extension, the magnitude or the impact of hazards on an area ...

L27: ... in those places...:

- is Lanzarote also included in the THOSE places? If not I would suggest to skip the “those”

We have deleted “those”

L30: ... due to the increase of exposition of most places ...

- due to urban sprawl . We have changed exposition by urban sprawl

L33: ... despite having hosted 15 eruptions in historical times

- you might refer to your table 1 and to add, in this table, information about the years of these eruptions. The 15 historical eruptions took place in the whole Archipelago, so none of the eruptions of table 1 correspond to any historical eruption in Lanzarote. Those of the table are pre-historical eruptions.

L34: ... one of the most important touristic destinations

- What is your reference? Wouldn't it better to be more general and say that “Lanzarote is an important touristic destination”? We are talking about the Canary Islands in this sentence. The reference has been taken from the fact that Teide National Park is the most visited in Europe (<http://www.gobiernodecanarias.org/istac/jaxi-istac/tabla.do>; <http://dx.doi.org/10.1016/j.jnc.2016.03.001> see table 5)

L35: ... has abandoned some traditional livelihoods

- Was the traditional livelihood more adapted to the volcano to mention this in the paper? Not really, we have deleted this part of the sentence in order to not create confusion.

L35-38: Tourism has had a considerable economic impact on the region that has abandoned some traditional livelihoods and has suffered a tremendous demographic expansion. The latter, not always well planned and without considering potential natural hazards, may now interfere with the effective management of future volcanic crisis.

- It would be great to support your statement of growth with some numbers or with a map where you would see the urban sprawl. So we can have an idea of the importance of the change (big/small change)

We have added some numbers that indicate the total inhabitants of the archipelago in 1970 and in 2016.

L38-42: The last eruption, that occurred in El Hierro (Fig. 1 Inset) in 2011-2012, is a good example of the implications of not having conducted a previous hazard assessment...

- The link with the demographic expansion is not clear enough. Could you please give us additional information: did the urban expansion mainly occur in the last eruption period? We have written additional information which is possible to compare the growth of the population in The Canary Islands during the last 50 years.

Did the eruption cause socio-economic disaster because no plan was existing?

There was an emergency plan at that moment, but the management of the eruption was based on potential eruptive scenarios that may occur in similar volcanoes than those from el Hierro but not on a specific knowledge of the past volcanic activity on the island.

What was the human component/implications/problems of this eruption on the Island? One year before the eruption, the island was already severely impacted by a drastic drought that caused famine, and half of the population emigrated to other areas. When the eruption occurred, many of the Lanzarote inhabitants, who decided to remain there, finally emigrated too.

L55: Timanfaya type eruption

- Define the Timanfaya type eruption

We have added to the text the following sentence: (see the geological setting description to obtain more information about this eruption) to not describe again the Timanfaya type

L57: ...the first susceptibility map

- ...the first volcanic eruption susceptibility map

Correction done

L59: ... a spatial probability map:

- Of what?

We have written volcanic spatial probability map

L61: ...these previous studies tackles a proper volcanic hazard

- What do you mean with “proper”

We have changed proper by thorough

L65: based on a review of these previous studies...

- Did you only used the above mentioned studies or did you used also additional information. If so, I would suggest deleting the word “previous”.

We have added “new generate information” in order to clarify that we used previous and new information.

L66: Matri et al (2016a)

- et al.

Correction done

L68 : due to the scarce available information

- Gives a “bad” impression. Wouldn’t it more suitable to present the think like if there are out of the scoop of this paper for example

We can change the adjective, but it is true that not much geochronological information, essential to obtain a right recurrence period and to evaluate volcanic hazard on the island, is available. We prefer to state the sentence that it is, due to if the paper is read in the future, perhaps someone decides to invest money for completing the dating catalogue.

L70 : recognised in the Holocene

- recorded ? in the Holocene

We have written “documented” instead of “recognised or recorded”

L84-L89:

- Does this paragraph give an added value to this work?

This paragraph contextualises the geological setting of the island. Since the volcanic hazard evaluation has as basis the geology, we consider that it is important to make a general framework of the island, and afterwards to go into more in detail.

L95: a really high eruptive rate L105: eruptive rates of 0.013-0.027 km³/ka

- Pay attention to the fact that the second eruptive rate is higher than the first one that you mention. Maybe it could be a good idea to avoid the use of “really high” for the first eruptive rate or to also give this impression for the second one.

We have deleted “really”

L108: Tao, Nuevo Fuego and Tinguaton eruption

- To keep as singular or to put as a plural?

This is the name of the eruption because 3 cones were formed along a fissure during 1824. Nevertheless, we have deleted “the” to not create confusion

L115: Methodology

- Isn’t this paragraph more related to the data that you are using?

In this first paragraph we are explaining the data and how we have used these data for conducting our analyses. For this reason we consider it is part of the methodology.

L116: the first step in any long-term...

- Could you mention more precisely why you are collecting your data “to produce a volcanic susceptibility map”?

We explain which kind of data and why we use them to produce the volcanic susceptibility map further in the text (lines 200-210 in the new manuscript).

L117: Holocene period

- **Justify the choice of this period in the frame of your work.**

We have added to the sentence that products are better preserved in the Holocene period. But also, and this is inherent to the meaning of Holocene, it is common to focus the volcanic hazard assessment during this epoch, since the volcanos from the Holocene are considered as active.

L120: “we”

- **It is a personal choice but may I suggest to use an impersonal form in the whole manuscript. To avoid the use of “we” and to adapt the sentence like this “Previous geological ... have been taken into account”.**

Correction done. In the rest of the paper we have preferred to use the active “we” form, due to NHES has not the rule for writing in impersonal or passive style.

L127: ... in the computation of volcanic susceptibility

- **State clearly why the previous volcanic susceptibility maps that have been realized in Lanzarote have to be updated for this work.**

We have preferred to include a sentence in the susceptibility analyses section (lines 221-222) to state clearer why this new susceptibility map has been updated on the previous ones.

L129: previous information

- **Please, provide additional information about which previous information you are talking about**

We refer to all above mentioned information (maps, ortophotos, structural analyses, etc.). We have changed previous information by “all above mentioned information”

L132: for the spatial analysis (volcanic susceptibility).

- **Are you producing one susceptibility map or are you making a distinction between lava flows and explosive eruptions? I guess the probability to have these kind of eruption differ within the island (presence or not of water).**

The susceptibility map shows the probability of hosting new eruptions regardless of the type of the process such as: lava flows, ashfall, PDCs..... or other hazardous processes. Therefore, we have developed a susceptibility map that is the basis of the lava flow scenarios, since it is the most expected process on the island. For the other two processes, ashfall and PDCs, we have preferred to not use the susceptibility map, because our intention was to show how these processes would affect if there were an eruption from the highest probability area of the island.

L133:... to simulate lava flows, fallout and pyroclastic...

- **Use the same sequence as the one presented in the manuscript.**

Correction done

L115-L134:

- **When stating the data that you are using, I think you forgotten to mention the wind information you’ve collected.**

You are right!. We have added data wind collection in this section.

L140: most sub-historical

- **What do you mean with this?** It is the period of time before the Spanish conquer, that is, before 1405.

L141: Guatiza map)

- **No parenthesis needed**

Done

L146: Guatiza map)

- **No parenthesis needed**

Done

L140-L146 and L157-L166:

- **Clearly state the difference between both paragraphs.**

In the first paragraph we are talking about the sub-historical eruptions (before the last 600 years), meanwhile in the second one, we are talking about the historical ones (last 600 years). We have included before what means historical and sub-historical eruptions.

L163: 226km² of the Lanzarote's surface

- **Replace with 226km² of Lanzarote surface**

Correction done

L165: Some of the stages...

- **Simply refer to these studies**

Correction done

L167: the consequences of 6 years

- **All number under 10 have to be completely spelled. Replace with "six years"**

Correction done

L173: 14km in length

- **Where is it on the map?**

We have added to figure 1, dashed lines to indicate the total length of the historical fissures.

L174: ...the SW coast

- **Isn't it the NW coast?** We have changed SW by NW

L190: ...we obtained Miocene-Pliocene, Pleistocene

- **Give some arguments why are you here extending your time frame.** Done

- **Be consistent with your figures and give the same terminology in your text and on your figures (table 2= mio-pliocene)** Correction done

L199: ...eruption will start...

- **Replace "will" with "may".** Susceptibility maps are still probabilities and some nuance has to be given to this sentence.

Correction done

L201: This volcano-structural information...

- **Based on the premise that new vents will not form far from the previous ones, this volcano-structural information is used...**

Correction done

L214: ...the distribution of volcanism...

- Replace with: the volcanism distribution

Done

L216: ... taking into consideration Geyer et al. (2016)

- the contribution of Geyer et al. (2016) is not clear enough. If I'm right, they produced the regional stress field. Refer to them after mentioning the regional stress field or state clearly that they produced it: "the regional stress field produced by..."

Correction done

L218: to generate quantitative ... in the island

- to generate a quantitative... on the island

Correction done

L219: ...method that uses the calculation of a ...

- method that calculates a kernel function...

Correction done

L220: The method is based on the distance...

- rephrase your sentence. It has the exact same structure as the previous one.

We have changed the previous sentence, therefore we have left this sentence as it was.

L223: onshore and offshore

- it is quite confusing, you are giving us the impression that offshore eruptive fissures may be observed and used but later on you state that they can't be used for the analysis (L315). The information present on line 315 has to be given before to clearly state that even though offshore eruptions are highly probable, they can't for the moment be included.

We are talking about two different things. Volcano-structural information offshore (vents and eruptive fissures) have been considered for the spatial analysis but due to the fact that they have not ages (geochronological dates), it is not possible to conduct a temporal hazard assessment using such information.

L224: ...and reliability values

- refer to table 3

Done

L225: GVB-CSIC

- define the acronym

Done

L227: LSCV:

- define the acronym

Done

L225 and L228:

- avoid repetition of the same information or clearly state the difference between both panels. Elicitation of expert judgment procedure ... meanwhile a group of expert...

We have deleted the sentence to avoid duplicity and rephrased the sentence

L217-239:

- the sequence of the manipulation is not clear. It gives a repetition feeling. Re-structure the paragraph.

Done

L231: the bandwidth parameter

- could I advise to be consistent in the terminology that you choose. Pick up one word: smoothing parameter or bandwidth parameter.

We have stated that there is the possibility to call this parameter in four different ways: smoothing parameter, smoothing factor, parameter h or bandwidth to make readers clear that commonly in the literature it could appear with different names. In the rest of the text we have called it as Bandwidth, maintaining this term throughout the text, even in Table 3.

L235: considering the regional stress field model

- clearly state that the stress field is not use as input in QVAST.

We have included a sentence to clarify this issue.

L249: ...size were inferred from data published from historical eruptions

- replace with ... inferred from the historical eruptions published data.

Correction done

L250: ... and references therein

- please provide all main references you are using.

We have provided the first three references as main ones, but inside the map's memories there are more that can be consulted.

L251: ...parameters of 1824

- replace with: "parameters as the 1824... "

Correction done

L252: ...since this scenario can be...

- since these parameters can be?

We have slightly changed the sentence

L260-264:

- wouldn't this be more suitable for a caption?

We are describing the figure 5, but anyway we have changed a bit the sentence.

L258: ... the entire wind rose directions and for the NE direction...

- refer to the figures (Fig. 5b) and (Fig. 5a)

Correction done

L260: figure 5A

- Please provide, such as for figure 5B, the parameters that have been used.

They are the same parameters. We have changed the sentence into the text to make it clearer.

L269: ... as single vent scenarios reproducing lava flows of 1730

- I don't see the added value of showing these results. Where you calibrating the model using these lava flows? The overall map is more interesting.

We wanted to show the extension of the lava flows from both historical eruptions. The 1730-1736 eruption was longer in time, and therefore more volume was emitted, invading a greater area than the 1824's eruption. If we only show the total hazard map, we are not giving the opportunity to the reader to see the extension differences between these two historical eruptions.

L273: ...35km, since 1730-36 eruption poured out lavas... 25km.

- Not clear which final length you are using for the simulation. Two lengths are mentioned. Which has been used to model the lava flows? The parameters used for the simulation could be added in the caption of the figure.

We have clarify the sentence adding more information

L286: ...in areas close to the previous

- Are you still using the susceptibility map?

In this case we are simulating only in areas close to previous eruptions that have generated PDCs, without considering the susceptibility map. We consider more interesting to show the reach of PDCs with different characteristics.

L295: in the range of around 5-29°

- Be more precise

We have added some more information to the text

L296: ...areas with different Heim

- Heim? We have change heim coefficients by collapse equivalent angles

- Give the exact values that you used. We have added a sentence clarifying that each of the simulation is associated with previous occurred PDCs on the island. Numbers in Figure 7 are related with those from Table 1.

L300: ... has hosted important eruptive

- Has hosted an important...

Correction done

L321: ... Timanfaya eruption

- Reference to fig 4

We have reference fig 4

L322: ...in this zone (Figs 1,4)

- Reference only to figure 3

Correction done

L326-328:

- The argument that is given is weak. Give more arguments. State clearly how does that method of Cappello et al. have been proved.

We have preferred to use a method such as the one of Capello et al. (2013) as it has been successfully tested in volcanic fields similar to Lanzarote, rather than to develop a new method, whose the lack of testing could imply a higher uncertainty in the results obtained.

L329-333:

- **Keep in mind that the conclusion that is made is only valid for one case scenario.**

But this scenario has been done with the prevalent winds of the Canary Islands. In the model it is not possible to take into account all possible winds for one scenario.

L338: ...National Park and Natural Park

- **Show these areas on a map**

We have included in Figure 2 two dashed lines to show the extension of these areas

L341: ...would be practically unaffected by lava flows

- **Would have a lower chance to be inundated by lava flows.**

Correction done

L343: ... to areas close to the coast

- **All areas in Lanzarote are close from the coast.**

Those closer areas to the coast are more suitable to the occurrence of hydromagmatic events, since the water from the sea can play an important role in eruptions located there. Nevertheless we have clarified the sentence adding “more”.

TABLES:

Table 1: add the starting and end year of each eruptions

They are not historical eruptions; they have been recorded in geological times. They are not dating for any of them.

Table 2: Are the faults onshore or offshore

We have modified the table to clarify they are onshore faults

Table 2-3: use similar names

Correction done

Table 4:

- **The mean is usually associated with one value. Clearly state which is your mean length: 5 or 7.**

- **Put the table in the same sequence as the manuscript.** We are not sure if you mean format. We have copied the format from previous tables.

- **Column height: which is the unit of it?** km- added

- **Size particles: unit?** Phi scale (ϕ)

FIGURES:

GENERAL REMARQUES

- Put the names of the islands, cities, volcanoes, parks that you mention on a map.

Done

- Use **colorbrewer2.org** to pick up colors that have a good contrast. Sometimes, some colors have not been selected properly provoking some confusion in the figure.
- Writing “**legend**” as title of the legend is not really needed. People know that symbols are the legend. Ok

Figure 1

- The 2011-12 eruption is quite hard to see on your map. Make the symbol or the frame **bigger**. The triangle for the 2011-2012 was too small, we have enlarged the symbol.
- **colorbrewer2.org** We have decided to leave figure as it is but we really appreciate your suggestion that will serve us for future figures.
- “**historical eruptions 1824**” to be replaced with **eruption**. Correction done
- Use in the sea the same blue as the one used in figure 2. The blue that is use for the moment make the message more difficult to see. The geological information is the message of this figure and not the sea. We have applied the same transparency to both figures. They show now the same blue sea colour.
- **Source: <http....> : wouldn’t it adapted to propose your own version of the map?**
We have clarified the source of the figure.

Figure 2:

- **Historical eruptions 1824: to replace with eruption**. Correction done
- If a name is associated to some eruptions, I would suggest adding the names in the legend. We have placed the corresponding names
- **colorbrewer2.org**
- Some elements of the map are not present in the legend (the symbol of how the pictures have been taken for example). We have indicated in the figure the orientation in the figure. Therefore, we have not inserted the symbol into the legend.
- Show only the elements you mention in the text. Some cones and eruptions that are present in the figure are never mentioned in the text: remove them from the figure. All names of the figure caption have been referred in the text of tables at least once.

Figure 3:

- **colorbrewer2.org**
- Use similar colors for the vents and fissures of one same period. We had already done it.

Figure 4:

- **Susceptibility? Give more information: susceptibility of... .** We have written “Volcanic Susceptibility”
- “**Value**” can be removed. Done
- The colors of the legend are not the same as the one in the figure. We have changed them
- **Timanfaya park: where is the park?** We have changed in the caption Timanfaya Park by south of the Island.

Figure 5:

- **Scale: put everything to 10km and use the same extent for 5a and 5b.**

Done

Figure 6:

- Remove the “legend” title and the “lava_35km”, “lava_7km”, “value”, “lava flows.tif” Done

- For a and b: place the eruption location you used for the simulation

- Inform the reader the values are probabilities to be inundated by lava flows.
Done

- For c: Color choices: Look at Thompson et al. (2015). She is giving useful tips for the selecting the right color range for maps. Use the same color range as for a and b or adapt all of them based on Thompson et al. Green is usually considered as safe however, you have still some probabilities.

You are right; therefore we have used the same colour for a, b and c figures.

Figure 7:

- VEI: no additional information is given about that in your text. You are right. We have deleted it.

- Give some extra information about all the symbols of your map. What are the lines for. Done (We have used different patterns to show the limits of all PDC scenarios).

Assessing qualitative long-term volcanic hazards at Lanzarote Island (Canary Islands)

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Abstract. Conducting long-term hazard assessment in active volcanic areas is of primordial importance for land-use planning and to define emergency plans able to be applied in case of a crisis. Definition of scenario hazard maps helps to mitigate the consequences of future eruptions by anticipating to the events that may occur. Lanzarote is an active volcanic island that has hosted the largest ($>1.5 \text{ km}^3$ DRE) and longest (6 years) eruption, the Timanfaya eruption (1730-36), on the Canary Islands in historical times (last 600 years). This eruption brought severe economic losses and forced local people to migrate. In spite of all these facts, no comprehensive hazard assessment neither hazard maps have been developed for the island. In this work, we present an integrated long-term volcanic hazard evaluation using a systematic methodology that includes spatial analysis and simulations of the most probable expected eruptive scenarios.

1 Introduction

Active volcanic areas require conducting long-term hazard assessment in order to ensure a rational land planning and to elaborate precise emergency plans that can be applied in case of a crisis. ~~Moreover, the~~ Long-term hazard assessment is important to identify the main aspects related to volcanic hazards, ~~such as the extension, -the magnitude or the~~ ~~that may~~ potential hazards impact ~~zones hazards~~ -on an area, ~~-and that which~~ should be known by local population and potential visitors, especially when these may potentially affect touristic destinations. Unfortunately, this is not the case of many active volcanic areas around the World, particularly in ~~those~~ places with a lower eruption frequency, thus making the historical memory of local societies to rapidly forget about past events. Also, ~~even~~ when ~~the impact~~ past eruptions ~~impacts~~ ~~have~~ ~~s~~ not been very significant, without causing a serious damage on human life and properties, they might become nowadays a socio-economic disaster due to ~~the increase of exposition~~ ~~urban sprawl~~ of most places and vulnerability of exposed elements.

This is, for example, the case of the Canary Islands where, despite having hosted 15 eruptions in historical times, volcanic hazard assessment is still a pending task for most of the islands. This volcanic archipelago, which includes four National Parks, is one of the most important touristic destinations in Europe. Tourism has had a considerable economic impact on the region that ~~has abandoned some tradi-~~ ~~tional livelihoods and~~ has suffered a tremendous demographic expansion ~~in the last 50 years (ca 1 million inhabitants in 1970 and more than 2 million people in 2016; <http://www.gobiernodecanarias.org/istac/>).~~

The latter, not always well planned and without considering potential natural hazards, may now interfere with the effective management of future volcanic crisis. The last eruption, that occurred in El Hierro (Fig. 1 Inset) in 2011-2012, is a good example of the implications of not having conducted a previous hazard assessment. Despite having an emergency plan that was correctly applied during the crisis, the occurrence of a submarine eruption was not considered as a probable scenario, ~~having been afterwards having that was later on~~ proved ~~this that it was as one of~~ the most probable scenarios ~~by further studies~~ (Becerril et al., 2013, 2014, 2015).

Here, we concentrate our attention on Lanzarote (Fig. 1), the easternmost island of the Canary archipelago. It has hosted the largest historical eruption of the Canaries (Timanfaya, 1730-1736) and one of the largest occurred on European territory. Lanzarote, declared Biosphere Reserve by UNESCO (1993, <http://www.lanzarotebiosfera.org/>) and Global Geopark (2015, <http://www.geoparquelanzarote.org/>), is an important touristic destination with 12 natural protected areas (<http://www.gobiernodecanarias.org/cmayer/espaciosnaturales/>) and a National Park (1974, <http://www.gobiernodecanarias.org/parquesnacionalesdecanarias/es/Timanfaya/>) that receives near 1.5 million visitors per year. As in the rest of the Canary Islands, local economy is tourism based and volcanism is regarded as an attraction and not as a potential problem for both local population and visitors.

During the last two decades, several attempts have been ~~carried out~~made to analyse volcanic hazard in Lanzarote. The first published works correspond to Felpeto (2002) and Felpeto et al. (2001, 2007) who presented a new methodology for the evaluation of the lava flow hazard on Lanzarote. However, these studies only focused on simulating lava flows related to a Timanfaya-type eruption (~~see the geological setting description to obtain more information about this eruption~~) without performing a general susceptibility analysis or a lava flow map for the whole island. Bartolini et al. (2013) presented the first susceptibility map of Lanzarote as an example of application of the QVAST tool, using the volcano-structural information available at that time. More recently, Galindo et al. (2016) published a spatial probability map of Lanzarote and Chinijo Islands and their submarine flanks. Their analyses were based on kernel density estimation via a linear diffusion process, using chronostratigraphic, volcano-structural and geomorphological data. However, none of these previous studies tackles a ~~proper thorough~~ volcanic hazard assessment for Lanzarote, although the information they provide should contribute to accomplish such task.

In this study, we applied a systematic methodology to conduct long-term volcanic hazard assessment at Lanzarote, based on a review of these previous studies, ~~new generate information,~~ -and the application of the methodology and e-tools described by Martí et al. (2016a) (see ~~also~~ www.vetools.eu/), ~~which-It~~ includes the sequential application of spatial analysis, temporal analysis, simulation of most probable scenarios, and vulnerability analysis. In the case of Lanzarote and due to the scarce available information (e.g.: lack of geochronological data), we only conducted the spatial analysis and the simulation of eruptive scenarios. The latter included the main volcanic hazards (fallout, lava flows, and pyroclastic density currents) ~~recognised~~documented in the Holocene volcanism in Lanzarote. Results obtained are volcanic hazard scenario maps, which should be considered for land-use planning, elaboration of emergency plans, and for managing a volcanic crisis, in order to protect people, their properties and the geological heritage of the island.

2 Geographical and Geological Setting

The island of Lanzarote (Canary Archipelago, Spain) is the north-easternmost island of the Canaries, located 125 km far from the western African coast and just 7 km towards the north of Fuerteventura (Fig. 1). It has an irregular morphology elongated NE-SW, with a maximum altitude of 671 m (Macizo de Famara) and covers an area of 846 km², which includes some islets located to the North. It rises approximately 2500 m from the sea bottom, being most part of the volcanic edifice submerged. Actually, ~~it is connected in its submerged part is connected part~~ with the island of Fuerteventura, both constituting the same volcanic edifice (Banda et al., 1981).

The basement of the island was constructed during the Oligocene above oceanic sediments of 65-55 Ma old, formed by submarine volcanic materials, plutonic rocks and sediments. It is located on an atypical oceanic crust, at least 11 km thick (Banda et al., 1981), or up to 15 km (Ortiz et al., 1986; Camacho et al., 2001). The subaerial volcanic history of Lanzarote started about 15.5 Ma ago (Coello et al., 1992) (Fig. 1). In addition to the volcanic materials, there are sedimentary formations, represented by aeolian sands, alluvial and colluvial deposits, mainly Pliocene and Quaternary (Fig. 1) (IGME, 2005).

Two major volcanic cycles have been established during its growth. The first cycle corresponds to the old buildings construction (between 11 and 3 Ma) and was characterised by the emission of important volumes of basaltic materials that formed a complex tabular sequence of lavas and pyroclasts gently dipping to the SE and ESE, with isolated outcrops of differentiated trachybasalts and trachytes (Fig. 1) (IGME, 2005). This first stage represents the maximum subaerial growing period (Ancochea et al., 2004), characterised by a ~~really~~ high eruptive rate, approximately 0.01-0.02 km³/ka (Coello et al., 1992). Los Ajaches, Famara and Tías Massifs are part of this cycle (Fig. 1) (Carracedo and Badiola 1993). The second stage (3 Ma - present) was characterised by a period of Pleistocene-Holocene eruptions and historical eruptions (last 600 years) (IGME, 2005). This second subaerial cycle includes the recent activity of Lanzarote and the growth of the small islands located to the North, the Chinijo Archipelago (Fig. 1) (Ancochea et al., 2004). It was characterised by the formation of widespread lava fields covering the materials of the first stage, and by the alignment of most vents trending NE-SW. On the other hand, the Chinijo Archipelago was also constructed by hydromagmatic eruptions (De la Nuez et al. 1997). It is marked by the emission of alkaline rocks that evolved to basaltic magmas, with a decrease of the alkalinity, and finally the emission of tholeiitic olivine basalts (Armienti et al. 1991; Carracedo and Badiola 1993). This second cycle of growth is characterised by continuous volcanic activity with eruptive rates of 0.013-0.027 km³/ka (Coello et al. 1992).

Two historical eruptions took place on the island: ~~the~~ Timanfaya (1730-1736) and ~~the~~ Tao, Nuevo Fuego and Tinguatón eruption (1824). Both were multiple-fissure type eruptions but quite different in size and duration. The Timanfaya eruption lasted ~~for~~ 6 years and formed hundreds of vents aligned along a 13-15 km long fissure, from where lava flows covered almost one-third of the island, erupting a total of > 1.5 km³ of magma (Romero, 1991; Carracedo et al., 1992) (Fig. 2). During the 1824 eruption, three eruptive fissures were formed emitting few pyroclasts and some lava flows, with lengths in the order of hundred meters (Romero, 1991; Carracedo et al., 1992) (Fig. 2).

3 Methodology

The first step in any long-term volcanic hazard assessment is the reconstruction of the past eruptive history of the volcano or volcanic area. In this sense, we based our analysis on the Holocene period from where we identified the different eruptive episodes and their products, since they are better preserved and established a relative volcano-stratigraphy for all of them. To accomplish this task, ~~we took into account~~ previous geological and volcanological studies of Lanzarote ~~have-were~~was taken into account (Romero, 1991; Carracedo et al., 1992; Ancochea et al. 2004; IGME maps (2004), and references therein) and ~~completed~~ing them with new field work when necessary. We ~~have~~ also conducted a structural analysis of the island based on previous geological maps at 1:25000 scale (MAGNA, GEODE) and structural studies (Marinoni and Pasquare, 1994; Galindo et al., 2016), and on remote sensing and morpho-tectonic analysis of orthophotos (GRAFCAN (<http://www.grafcan.es/>), topography (LIDAR Digital Elevation Model (1:5000), GRAFCAN ©) and bathymetry (1:100.000, IEO). In addition to these volcano-structural features, we also took into account in the computation of volcanic susceptibility the recently modelled regional stress field for the Canary Islands (Geyer et al., 2016).

~~All above mentioned~~ ~~The previous~~ information was used to define the input parameters necessary to run the different tools we ~~have~~ applied to conduct the systematic hazard assessment. These form part of the methodology described by Martí et al. (2016a), (<http://www.vetools.eu/>), i.e. QVAST (Bartolini et al., 2013) for the spatial analysis (volcanic susceptibility), and VORIS (Felpeto et al., 2007), a GIS-based tool that allows users to simulate ~~lava flows~~, fallout, lava flows and pyroclastic density current scenarios. For ashfall simulations, wind data was compiled from the University of Wyoming Department of Atmospheric Science sounding database (<http://weather.uwyo.edu/upperair/sounding.html>).

4 Holocene volcanism

Holocene eruptions in Lanzarote are restricted to a few sub-historical ~~fissure~~events (before the last 600 years) at the northeast (Guatiza area), and the historical eruptions located towards the western-central part of the island (Timanfaya area) (Fig. 2d).

Most sub-historical eruptions are fissure type, basic in composition (olivine basalts), with clear Strombolian character, (IGME, 2004); Guatiza map). Their main products are proximal fallout pyroclastic deposits and lava flows, mainly of 'aa' type, which reached the sea generating a platform, so having at least 5 km in length. Lava flows from Mt. de Guenia, Las Calderas de Guatiza, Las Calderas and Las Calderetas (Fig. 2d) come from fissures with trending N30°E - N37°E, being from 1-1.5 m to several meters wide. They have associated several scoria cones showing a great range of particle sizes (IGME, 2004); Guatiza map).

Hydrovolcanic events also occurred on Lanzarote during the Holocene and previous times. They include both Surtseyan eruptions, caused by the interaction of magma with water in coastal or shallow offshore settings, and inland phreatomagmatic eruptions generated by interaction of erupting magmas with groundwater (Pedrazzi et al., 2013). Several well preserved hydrovolcanic edifices are identified on

the island and islets (Fig. 2b). El Golfo (Martí and Colombo, 1990; Pedrazzi et al. (2013), La Caldera del Cuchillo, Mt. Cavera and Mt. Chica are some examples of hydromagmatic coastal edifices (Fig. 2b, Table 1) (Aparicio et al., 1994). The main characteristics of these eruptions and their subsequent deposits have been gathered from geological maps (IGME (2004)) and some previous studies (Martí and Colombo (1990); Carracedo and Badiola (1991); Aparicio et al. (1994); Pedrazzi et al. (2013); IGME (2004)-Geological Maps). They are summarised in Table 1.

Historical eruptions (both 1730-36 and 1824) were also of basaltic character. Timanfaya eruption differs from the rest of the Canary Islands historical eruptions, mainly because of its long duration, magnitude, type and evolution of magmas (Carracedo et al., 1992). It is the second largest historical effusive eruption in Europe (last 600 years) after Laki (1783-85) in Iceland (Thordarson and Self, 1993). A complex fissural volcanic system of approximately 13-15 km length, with more than 30 cones, was formed during this eruption (Fig. 2c), that produced lava flows and pyroclastic fallout that covered approximately 226 km² of the Lanzarote's surface (Hernández Pacheco, 1960; Carracedo et al., 1992). The total volume expelled was between 3 and 5 km³ (>1.5 km³ DRE). Lava flows reached the coast, and maximum onshore paths reached up to 21 km (Figure 1). Some of the stages of this eruption have been studied in detail by Romero et al. (1991), Carracedo et al. (1992) and Solana et al. (2004).

The consequences of 6-six years of activity were that more than one-third of farmland and numerous villages of the island were buried by ash and the accompanying degassing resulted in acidic rain fall, which triggered the evacuation and economic collapse of the island (Carracedo et al., 2012; Solana et al., 2004).

The 1824 eruption was characterized by basanitic products. Three cinder cones were formed during three months of activity (Tinguatón, Tao and Nuevo del Fuego; Fig. 2), generating an intermittent fissure almost 14 km in length (Figure 1). They produced a small lava flow, with a total on land length of 7-8 km that reached the SW coast of the island.

5 Volcano-tectonics

To identify the different structural elements that we will-considered in the susceptibility analysis, we defined vents and eruptive fissures following the same criteria established by Becerril et al. (2013, 2014, and 2015) on El Hierro. Thus, we recognised: (i) craters of isolated cinder cones, (ii) craters of coalescent cinder cones belonging to the same eruptive fissure, and (iii) craters without an associated cinder cone, both, submarine and subaerial. We discarded hornitos and rootless vents as volcanic vents to avoid over-value susceptibility analysis, since they are not lava emissions centres. Submarine eruptive vents morphologically recognisable were considered as volcanic cinder cones, including those located at the north of Fuerteventura, due to the proximity to Lanzarote and also because they belong to the same volcanic edifice.

From the volcano-structural study, we have-obtained different datasets that correspond to vents and eruptive fissures, both onshore and offshore the island, and onshore faults (Table 2). To identify on-shore structures we have-considered the complete emerged history of the island (from Miocene to Holocene). Volcano-structural datasets were divided according to the age of the structures and their location

(onshore or offshore) (Table 2). Thus, we obtained Miocene-Pliocene, Pleistocene and Holocene onshore vents, and eruptive fissures respectively, besides offshore vents and eruptive fissures (Fig. 3, Table 2). Only 6 faults ~~have~~were ~~been~~ identified on the island. The majority of the linear structures (eruptive fissures and faults) follow the NE-SW direction and they are from less than 1 km to 15 km length (Table 2).

6 Susceptibility analyses

The spatial probability of a future vent opening, given the past eruptive activity of a volcanic system, is a crucial step for simulating possible future eruptive scenarios, as it will provide indication from where the eruption ~~will~~may start, and how the corresponding hazards will distribute (Martí and Felpeto, 2010). The information required to perform this susceptibility analysis is the distribution of the past volcano-structural elements, their age, and the regional stress field. The first assumption is that the regional stress field has not changed since the last eruption. Based on this premise, new vents will not form far from the previous ones, and consequently, ~~This~~ this volcano-structural information ~~is~~can be used to pinpoint areas where next eruptions may most likely occur since they represent the sites where previous eruptions have taken place, ~~based on the premise that new vents will not form far from the previous ones~~ (Connor, 1990; Connor et al., 1992, 2000; Ho, 1992, 1995; Martin et al., 1994; Ho and Smith, 1998; Connor and Conway, 2000; Gaffney et al. 2017; Martí and Felpeto, 2010; Bebbington and Cronin, 2011, Capello et al., 2012; Selva et al., 2012; le Corvec et al., 2013a; Bartolini et al., 2013; Bevilacqua et al., 2015; Martí et al., 2016b). ~~This reasoning is based on the assumption that the regional stress field has not changed since the last eruption. Therefore,~~ Other kind of data such as geophysical information or the stress field configuration of a volcanic area, if available, should be also used to forecast more precisely the most probable areas to host future vents (Martí and Felpeto, 2010; Martí et al., 2016b). In particular, the stress field is a key parameter controlling magma generation, magma migration and magma accumulation inside the volcanic system, as well as the location, geometry and the distribution of the resulting volcanism at surface (Martí et al., 2016b). Therefore, knowing the stress configuration in the lithosphere at any scale (i.e. local, regional and plate-scale) is important to understand ~~volcanism the distribution of volcanism~~ and, subsequently, to predict the location of future eruptions (Martí et al., 2016b). For that reason, in this work we also considered the regional stress field configuration ~~under in~~ Lanzarote, (Geyer et al. 2016), taking into consideration Geyer et al. (2016) that which updates the previous susceptibility maps developed by Bartolini et al. (2013) and Galindo et al. (2016).

We used the QVAST tool (QGIS for VolcAnic SuscepTibility; Bartolini et al., 2013), to generate a quantitative assessment of volcanic susceptibility in the island. This tool is backed on a probabilistic method that ~~uses the calculation of~~ a kernel function at each data location, based on the distance from nearby volcanic structures, to estimate probability density functions (PDFs). ~~The method is based on the distance from nearby volcanic structures and a).~~ One of the most important factors to determine this density distribution is the smoothing parameter, also known as smoothing factor, ~~parameter h~~ or bandwidth, which represents the degree of randomness in the distribution of past events. ~~Our volcano-structural data sets (vents, and eruptive fissures onshore and offshore, and faults) plus the stress field (Fig. 3), which was the isan~~ were combined by assigning to each of them the corresponding relevance and

reliability values (Martí and Felpeto, 2010) through an elicitation of expert judgment procedure (Aspinall, 2006) among the members of the GVB-CSIC and external collaborators (14 experts in total).

We applied the LSCV method (Cappello et al. 2012) to evaluate the bandwidth of each dataset; meanwhile a group of experts evaluated the relevance of the datasets following the methodology proposed by Aspinall (2006).

The bandwidth parameter (h) obtained for each of the defined datasets were (Table 3): i) 2,527 m for vents and fissures of the Miocene-Pliocene; ii) 2,808 m for vents and fissures of the Pleistocene; iii) 560 m for the vents and fissures of the Holocene; iv) 6,508 m for vents and fissures offshore; and v) 20,808 m for faults (Table 3).

Considering the regional stress field model of Geyer et al. (2016) and the different ages of the volcano-structural elements, the expert judgement elicitation assigned the following weights to each data set: i) 0.107 for vents and fissures of the Miocene-Pliocene; ii) 0.207 for the vents and fissures of the Pleistocene; iii) 0.357 for the vents and fissures of the Holocene; iv) 0.193 for offshore vents and fissures; and v) 0.136 for faults (Table 3).

The total susceptibility map was obtained by assigning different weights to each of the PDFs, which are then combined via a weighted sum and modelled in a non-homogeneous Poisson process (Fig. 4).

In this study, we applied the Least Square Cross Validation (LSCV) method to evaluate the bandwidth of each dataset (Cappello et al. 2012, 2013; Del Negro et al., 2013), as it better represents the geometry of the vents distribution, NE-SW elongated. The dataset used is our volcano-structural information: vents, eruptive fissures onshore and offshore, and faults (Fig. 3). The bandwidth parameter (h) obtained for each of the defined datasets were (Table 3): i) 2,527 m for vents and fissures of the Miocene-Pliocene; ii) 2,808 m for vents and fissures of the Pleistocene; iii) 560 m for the vents and fissures of the Holocene; iv) 6,508 m for vents and fissures offshore; and v) 20,808 m for faults (Table 3).

Considering the regional stress field model by Geyer et al. (2016) and the different ages of the volcano-structural elements, the expert judgement elicitation assigned the following weights to each data set: i) 0.107 for vents and fissures of the Miocene-Pliocene; ii) 0.207 for vents and fissures of the Pleistocene; iii) 0.357 for vents and fissures of the Holocene; iv) 0.193 for offshore vents and fissures; and v) 0.136 for faults (Table 3). In detail, the relevance and reliability values (Table 3) (Martí and Felpeto, 2010) have been assigned as follow: relevance was given through an elicitation of expert judgment procedure (Aspinall, 2006) among the members of the Group of Volcanology of Barcelona (GVB-CSIC) and external collaborators (14 experts in total); reliability was considered as maximum in all the datasets (value of 1), since all of them come from previously published volcano-structural studies and direct field observations.

The total susceptibility map was thus obtained via a weighted sum and modelled in a non-homogeneous Poisson process (Fig. 4).

7 Eruptive scenarios

7.1 Fallout Scenarios

Fallout scenarios were obtained using VORIS 2.0.1 tool (Felpeto et al., 2007). The input data regarding the eruptive column and ash particle size were inferred from ~~data published from the~~ historical eruptions ~~published data~~ (Romero 1991; Carracedo et al. 1992; Ancochea et al., 2004; IGME maps (2004), and references therein). We simulated one scenario with the same eruptive parameters ~~of as the~~ 1824 eruption considering a maximum column height of 3 km and a total emitted volume of 0.02 km³ (Table 4) ~~since because assuming~~ this scenario ~~can may be as~~ the most ~~expected probable~~ in the near future in the island.

All simulations were conducted from one of the pixels located in the highest spatial probability area, and data inputs of wind velocities were compiled from the University of Wyoming Department of Atmospheric Science sounding database (<http://weather.uwyo.edu/upperair/sounding.html>) at different vertical heights (500, 1500, 2500 ~~and~~, 3500 ~~and 5000~~-m). We focused the attention of our study on the fallout scenarios ~~for the NE direction (Fig. 5a) which represents the typical north-east trade wind that characterises the Canary Islands latitude, and for the entire wind rose directions (Fig. 5b) and for the NE direction, which represents the typical north-east trade wind that characterises the Canary Islands latitude.~~ Results are shown in Figure 5. ~~Particle sizes (-6 to 2 ϕ) were considered in all simulations, thereby covering the entire range of particle sizes observed in the field.~~

~~Figure 5a shows the ash fall distribution, considering a strombolian eruption similar than the 1824 one, has been simulated from the highest susceptibility pixel, taking the average winds of the Canary Islands in any given day (Fig. 5a). The parameters from the 1824 eruption are: column height of 3 km and a total emitted volume of 0.02 km³. Figure 5b shows the distribution of the fallout from the same pixel considering the entire wind rose directions the parameters of the 1824 eruption: column height of 3 km and a total emitted volume of 0.02 km³ is shown in Figure 5b. Particle sizes in all simulations were considered in a range from -6 to 2 ϕ , thereby covering the entire range of particle sizes observed in the field.~~

~~In the case of the fallout scenarios we have only reproduced two scenarios (NE wind direction and entire wind rose directions) from a single vent located in the the area with highest susceptibility values, instead of making the calculation from all pixels of the map with a susceptibility value higher than 0. The reason is that ashfall process does not depend on the topography (DEM), but only on the position of the vent and wind direction, in addition to all eruptive parameters. Therefore, the use of the volcanic susceptibility map as base map for simulating ashfall would have required almost 150000 simulations that correspond to the number of pixels of the susceptibility model. All these simulations together would have given a superposition of many plumes that would cover the entire island, not having much sense for the purposes of this study.~~

7.2 Lava Flow Scenarios

The most expected processes associated with an effusive eruption in Lanzarote are lava flows. Lava flow scenarios were performed for the whole island ~~using VORIS 2.0.1 tool (Felpeto et al., 2007)~~, and as sin-

gle vent scenarios reproducing [the](#) lava flows of [the](#) 1730-36 and 1824 eruptions ([Fig. 4a, b](#)). For the first case, we used the whole susceptibility map (Fig. 4), only taking into account the on-land pixels. For single vent scenarios, we used only those pixels with the highest spatial probability values. Lava flow input parameters were constrained by maximum flow lengths and thicknesses taken from historical eruptions and field measurements. We assumed flow lengths up to 35 km, ~~since because of the~~ 1730-36 eruption poured out lavas that reached the sea after paths of 21 km onshore. Maximum lava flow length considered for the 1824 eruption was 7 km, while for the whole lava flow map, ~~a~~ maximum length was 25 km, [taking into account lava lengths considering a scenario more similar than this from the 1730-1736 eruption](#). The thickness used as input for all the models was 10 m. The results provide two single vent scenario maps and a total map that gives the probability that any particular cell is invaded by a lava flow (Fig. 6). The total lava flow map was performed with a cell size of 75 m, thus optimizing the result and computed time.

7.3 Pyroclastic Density Current Scenarios

Hydromagmatic eruptions have also occurred on Lanzarote in recent times and have generated a wide variety of pyroclastic density currents (PDCs) deposits. It is possible to recognise pure hydromagmatic edifices and also Strombolian edifices with phreatomagmatic phases (García-Cacho and Romero, 2000). For that, we have mainly simulated hydromagmatic ~~edifices-eruptions~~ in areas close to the previous ~~ones~~ [vents](#) but also some phreatomagmatic phases that could occur together with Strombolian activity. PDCs were simulated with an energy cone model (Sheridan and Malin, 1983) using as input parameters topography, the collapse equivalent height (H) and the collapse equivalent angle (θ), which is obtained through the arctangent of the ratio between Hc and L, where L represents the run-out length (Felpeto et al., 2007; Toyos et al., 2007).

L values were considered to be equivalent to the most distal exposure of PDC deposits found on the island (Tables 1 and 4), which correspond to lengths from 0.5 to 3 km. H was assumed to be 250 m for all simulations, considering similar kind of eruptive styles for these hydromagmatic eruptions (Toyos et al., 2007). We simulated PDCs with θ in the range of around 5–29° [\(low values for base surge type explosions and high values for PDCs derived from column collapse phases\)](#) (Sheridan and Malin, 1983) [\(Tables 1 and 4\)](#). Figure 7 shows coverage areas with different [column collapse equivalent angles](#) ~~Helm coefficients and VEI values~~, reaching the deposits up to almost 15 km. ~~Each of the simulation is associated with previous occurred PDCs occurred on the island, that is, similar parameters and close areas of previous PDCs deposits have been considered. Numbers in Figure 7 are related with~~ [those from Table 1](#).

8 Discussion and conclusions

Lanzarote is one of the four islands of the Canary Archipelago that has hosted [an](#) important eruptive activity during the last 600 years (historical period), being the Timanfaya eruption in 1730-1736 the second largest historical eruption occurred on a European territory. This, together with the fact that it is the third

preferred touristic destination of the Canary Islands, classifies Lanzarote as an active volcanic island for which a precise hazard assessment is urgently required.

Past on-land volcanism has been mainly characterised by multiple-fissure type eruptions of basaltic magmas, generating lava flows of variable length and small to medium sized cinder cones, so we should expect future eruptions being of the same type. A few hydromagmatic eruptions have also been recognised along the coast line or close to it, which generated Surtseyan activity when eruptive magma interacted directly with sea water (e.g.: El Golfo, Pedrazzi et al., 2013) or phreatomagmatic pulses when magma interacted with a saltwater intrusion near the coast (e.g.: El Cuchillo, Aparicio et al., 1994), respectively. In this case, different types of dilute PDC deposits were produced, together with ballistics and fallout, reaching distances up to 15 km from the vent. Moreover, the large number of well-preserved cones observed on the submerged slopes of the island suggests that the number of submarine eruptions in recent times may be similar or significantly higher than those from on-land. This suggests that a submarine eruption scenario should be considered as highly probable. Unfortunately, the lack of geochronological data precludes establishing the eruption recurrence in Lanzarote, so not allowing to conduct a temporal hazard assessment and to quantitatively identify the most probable eruptive scenarios. Therefore, our hazard assessment is restricted to the on-land volcanism, without this implying that a subaerial eruption is the one with the highest probability of occurrence on Lanzarote in the near future.

The spatial analysis revealed that the area with the highest probability of hosting a new subaerial eruption is mainly located in the same area than the previous 1824 and Timanfaya eruptions (Fig. 4). This is mainly due to the fact that the best preserved vents are concentrated in this zone (Fig. 3, 4), but also that the current stress field is compatible with orientation of fractures that governed these most recent eruptions (Fig. 3). Our results slightly contrast ~~a little bit~~ with those recently presented by Galindo et al (2016). The differences observed for the on-land areas may be due to the different method used in both studies, ~~but we have preferred to stay with~~. Our study follows the method of Cappello et al. (2013) ~~method, as since~~ it is a well tested method successfully applied to ~~has been proved successfully in many~~ volcanic fields such as Etna, El Hierro, Deception Island or Pico (Cappello et al., 2012; Becerril et al., 2013; Bartolini et al., 2014; Cappello et al., 2015), which show similar behavior ~~behaviour~~ than Lanzarote, and we considered it was more appropriate to model volcanic susceptibility in this particular case, rather than to develop a new model as it was done by Galindo et al (2016). ~~-, rather than to try a new one as done by Galindo et al (2016).~~

Simulation of the different volcanic hazards that may be produced in subaerial eruptions on Lanzarote revealed ~~that the~~ opening of new eruptive fissures in the highest probability areas, Assuming ~~assuming~~ a new typical Strombolian eruption and the typical winds of the Canary Islands (NE-SE winds), would imply the dispersion of the volcanic ash mainly towards the southern part of the island. As mentioned before, this area hosts a high number of tourist resorts, ~~so therefore, in case of an eruption,~~ probably a large number of people should be evacuated ~~in case of an eruption~~ (Fig. 5).

Lava flows are ~~more rather~~ constrained to the area around their vents. This implies that, according to the hazard map, if we expect a typical Strombolian eruption with lava flow emission, those areas that could be affected by this process, are mainly located surrounding the Timanfaya National Park. This

area includes two protected figures (a National Park and a Natural Park), but it does not host too many towns or infrastructures. If, on the contrary, we expect larger eruptions, in terms of emitted volume, the runout distances of the lava flows would be ~~bigger~~longer, affecting numerous towns and villages around the Timanfaya area, and others located to the north (Guatiza, Mala in Fig. 6). The rest of the island would ~~have a lower chance to be inundated~~be practically unaffected by lava flows.

Finally, the occurrence of PDC is more restricted to areas close to the coast, where the majority of the identified past hydromagmatic events are concentrated, being in age older than the most recent eruptions. However, such scenarios must be also considered as they may imply larger impacts than normal Strombolian eruptions.

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References

- Ancochea, E., Barrera, J.L., Bellido, F., Benito, R., Brändle, J.L., Cebriá, J.M., Coello, J., Cubas, C.R., De La Nuez, J., Doblas, M., Gómez, J.A., Gómez, J.A., Hernán, F., Herrera, R., Huertas, M.J., López-Ruiz, J., Martí, J., Muñoz, M. and Sagrado, J.: Canarias y el vulcanismo neógeno peninsular. En: *Geología de España*, (J.A. Vera, ed.), SGE-IGME, Madrid, 635-682, 2004.
- Aparicio, A., Araña, V., and Díez-Gil, J.L.: Una erupción hidromágmaica en la isla de Lanzarote: La Caldera de El Cuchillo. *Elementos de Volcanología* nº 3. Serie Casa de Los Volcanes. Excmo. Cabildo Insular de Lanzarote, 109–120, 1994.
- Armienti, P., Innocenti, F., Pareschi, M. T., Pompilio, M., and Rocchi, S. Crystal Population Density in not Stationary Volcanic Systems: Estimate of Olivine Growth Rate in Basalts of Lanzarote (Canary Islands). *Mineral. and Petrol.*, 44, 181-196, 1991.
- Aspinall, W.P.: Structured elicitation of expert judgment for probabilistic hazard and risk assessment in volcanic eruptions. In: Mader, H.M., Coles, S.G., Connor, C.B., Connor, L.J. (Eds.), *Statistics in Volcanology: Special Publication of IAVCEI*, 1. Geol.Soc. London, pp. 15–30, 2006.
- Banda, E., Dañobeitia, J. J. Suriñach, E. and Ansorge, J.: Features of crustal structure under the Canary Islands, *Earth Planet. Sci. Lett.*, 55, 11–24, 1981.
- Bartolini, S., Cappello, A., Martí, J. and Del Negro, C.: QVAST: A new Quantum GIS plugin for estimating volcanic susceptibility. *Nat. Hazards Earth Syst. Sci.*, 13(11), 3031–3042, 2013.
- Bartolini, S., Geyer, A., Martí, J., Pedrazzi, D. and Aguirre-Díaz, G. Volcanic hazard on deception Island (South Shetland Islands, Antarctica), *J Volcanol Geotherm Res.*, 285:150–168, 2014. doi:10.1016/j.jvolgeores.2014.08.009
- Bebbington, M. S. and Cronin, S.: Spatio-temporal hazard estimation in the Auckland Volcanic Field, New Zealand, with a new event-order model. *Bull. Volcanol.*, 73, 55–72, 2011, DOI 10.1007/s00445-010-0403-6
- Becerril, L., Cappello, A., Galindo, I., Neri, M., and Del Negro, C.: Spatial probability distribution of future volcanic eruptions at El Hierro Island (Canary Islands, Spain). *J. Volcanol. Geoth. Res.*, 257, 21–30, 2013.
- Becerril, L., Bartolini, S., Sobradelo, R., Martí, J., Morales, J.M. and Galindo, I.: Long-term volcanic hazard assessment on El Hierro (Canary Islands). *Nat. Hazards Earth Syst. Sci.*, 2, 1799–1835, 2014.
- Becerril, L., Galindo, I., Martí, J. and Gudmundsson, A.: Three-Armed Rifts or Masked Radial Pattern of Eruptive Fissures? The Intriguing Case of El Hierro Volcano (Canary Islands). *Tectonophysics*, 33–47, 647–648, 2015.
- Bevilacqua, A., Isaia, R., Neri, A., Vitale, S., Aspinall, W.P., Bisson, M., Flandoli, F., Baxter, P.J., Bertagnini, A., Esposti Ongaro, T., Iannuzzi, E., Pistolesi, M. and Rosi, M.: Quantifying volcanic hazard

at Campi Flegrei caldera (Italy) with uncertainty assessment: 1. Vent opening maps, *J. Geophys. Res. Solid Earth*, 2015, doi:10.1002/2014JB011775

Camacho, A. G., Montesinos, F. G., Vieira, R., and Arnosó J.: Modelling of crustal anomalies of Lanzarote (Canary Islands) in light of gravity data, *Geophys. J. Int.*, 147, 403–414, 2001.

Cappello, A., Neri, M., Acocella, V., Gallo, G., Vicari, A., and Del Negro, C.: Spatial vent opening probability map of Mt Etna volcano (Sicily, Italy), *Bull. Volcanol.*, 74, 2083–2094, 2012.

Cappello, A., Bilotta, G., Neri, M., and Del Negro, C.: Probabilistic modeling of future volcanic eruptions at Mount Etna, *J. Geophys. Res.*, 118, 1925–1935, doi:10.1002/jgrb.50190, 2013.

Cappello, A., Zanon, V., Del Negro, C., Ferreira, T. J. L. and Queiroz, M. G. P. S.: Exploring lava-flow hazards at Pico Island, Azores Archipelago (Portugal). *Terra Nova*, 27, 156–161, 2015.

Carracedo, J. C., Badiola, E. R. and Soler, V.: The 1730–1736 eruption of Lanzarote, Canary Islands: a long, high-magnitude basaltic fissure eruption. *J. Volcanol. Geotherm. Res.*, 53, 239–250, 1992.

Carracedo, J. C. and Rodríguez Badiola, E.: Evolución geológica y magmática de la isla de Lanzarote, Islas Canarias. *Rev. Acad. Canaria Ciencias*, 4, 25–58, 1993.

Coello J. Cantagrel, J. M., Hernan, F., Fuster, J.M., Ibarrola E., Ancochea, E., Casquet, C., Jamond, C., Diaz, de Teran, J.R and Cendrero, A.: Evolution of the eastern volcanic ridge of the Canary Islands based on new K-Ar data. *J. Volcanol. Geotherm. Res.* 53, 251–274, 1992.

Connor, C.B.: Cinder Cone Clustering in the TransMexican Volcanic Belt: Implications for Structural and Petrologic Models: *J. Volcanol. Geotherm. Res.* 95, 395–319–405, 1990.

Connor, C.B., Condit, C.D., Crumpler, L.S., and Aubele, J.C. Evidence of Regional Structural Controls on Vent Distribution: Springerville Volcanic Field, Arizona: *Journal of Geophysical Research*, v. 97, no. 12, p. 12349–12359, 1992.

Connor, C.B. and Conway, F.M.: Basaltic volcanic fields. In: Sigurdsson, H. (Ed.), *Encyclopedia of Volcanoes*. Academic Press, New York, pp. 331–343, 2000.

Connor, C., Stamatakis, J.A., Ferrill, D.A., Hill, B.E., Ofoegbu, G.I., Conway, F.M., Sagar, B. and Trapp, J.: Geologic factors controlling patterns of small-volume basaltic volcanism: application to a volcanic hazards assessment at Yucca Mountain, Nevada. *J Geophys Res.*, 105, 417–432, 2000.

De la Nuez, J., Quesada, M.L. and Alonso, J.J.: Los volcanes de los islotes al norte de Lanzarote, Fundación César Manrique, Teguiise, Lanzarote, 233 pp., 1997.

Felpeto, A.: Modelización física y simulación numérica de procesos eruptivos para la generación de mapas de peligrosidad volcánica, Ph.D. thesis, University of Madrid

Felpeto, A., Araña, V., Ortiz, R., Astiz, M., and García, A.: Assessment and modelling of lava flow hazard on Lanzarote (Canary Islands), *Nat. Hazards*, 23, 247–257, 2001.

Felpeto, A., Martí, J., and Ortiz, R.: Automatic GIS-based system for volcanic hazard assessment, *J. Volcanol. Geoth. Res.*, 166, 106–116, doi:10.1016/j.jvolgeores.2007.07.008, 2007.

Gaffney, E. S., Damjanac, B. and Valantine, G. A. Localization of volcanic activity: 2. Effects of pre-existing structure. *Earth Planet. Sci. Lett.*, 263, 323–338, 2007.

Galindo, I., Romero, M.C, Sánchez, N., Morales, J.M.: Quantitative volcanic susceptibility analysis of Lanzarote and Chinijo Islands based on kernel density estimation via a linear diffusion process. *Sci. Rep.* 6, 27381; doi: 10.1038/srep27381, 2016.

Geyer, A., Martí, J. and Vilaseñor, A.: First-order estimate of the Canary Islands plate-scale stress field: Implications for volcanic hazard assessment. *Tectonophysics*, 679, 125–139, 2016.

Hernández-Pacheco, E.: En relación a las grandes erupciones volcánicas del Siglo XVIII y 1824, en Lanzarote. *El Museo Canario*, 73, 239-254, 1960.

Ho, C.H.: Risk assessment for the Yucca Mountain high-level nuclear waste repository site: estimation of volcanic disruption. *Math Geol*, 24, 347–364, 1992.

Ho, C.H.: Sensitivity in volcanic hazard assessment for the Yucca Mountain high-level nuclear waste repository site: the model and the data. *Math Geol* 27,239–258, 1995.

Ho, C.H. and Smith, E.I. A spatial-temporal/3-D model for volcanic hazard assessment: application to the Yucca Mountain region, Nevada. *Math Geol*, 30,497–510, 1998.

IGME: Memorias mapas MAGNA Lanzarote 1:25000. Hojas: Guatiza, Arrecife, Yaiza, Haría, Teguiise, Soo, Tinajo, Femés, Graciosa, Alegranza, Caleta de Sebo. Instituto Geológico y Minero de España, Madrid, 2004.

IGME: Mapa geológico de España escala 1:100.000, 88, Lanzarote. Instituto Geológico y Minero de España, Madrid, 2005.

Le Corvec, N., Spörli, K. B., Rowland, J., and Lindsay, J. Spatial distribution and alignments of volcanic centers: Clues to the formation of monogenetic volcanic fields, *Earth Sci. Rev.*, 124, 96–114, 2013.

Marinoni, L.B. and Pasquarè, G.: Tectonic evolution of the emergent part of a volcanic ocean island: Lanzarote, Canary Islands. *Tectonophysics*, 239 (1-4), 111-137, 1994.

Martí, J. and Colombo, F.: Estratigrafía, sedimentología y mecanismos eruptivos del edificio hidromagmático de El Golfo (Lanzarote). *Bol. Geol. Min.*, 101, 560-579, 1990.

- Martí, J. and Felpeto, A.: Methodology for the computation of volcanic susceptibility. An example for mafic and felsic eruptions on Tenerife (Canary Islands), *J. Volcanol. Geoth. Res.* 195, 69–77, doi:10.1016/j.jvolgeores.2010.06.008, 2010.
- Martí, J., Bartolini, S. and Becerril, L.: The challenge of conducting volcanic hazard assessment and risk management. What the VeTOOLS project can offer us?, *EOS*, 2016a.
- Martí, J., López, C., Bartolini, S., Becerril, L. and Geyer, A.: Stress controls of monogenetic volcanism: a review, *Front. Earth Sci.* 4:106. doi:10.3389/feart.2016.00106, 2016b.
- Martin, A.J., Umeda, K., Connor, C.B., Weller, J.N., Zhao, D., Takahashi, M.: Modeling long-term volcanic hazards through Bayesian inference: an example from the Tohoku volcanic arc, Japan. *J. Geophys. Res.* 109, B10208, 2004.
- Ortiz, R., Araña, V. and Valverde, C.: Aproximación al conocimiento del mecanismo de la erupción de 1730–1736 en Lanzarote: *Anales de Física Serie B.* 82 Especial Issue. “Física de los Fenómenos Volcánicos”, 127–142 pp., 1986.
- Pedrazzi, D., Martí, J., and Geyer, A.: Stratigraphy, sedimentology and eruptive mechanisms in the tuff cone of El Golfo (Lanzarote, Canary Islands). *Bull Volcanol* 75:740, 2013.
- Romero, C.: La erupción de Timanfaya (Lanzarote, 1730–1736). Análisis documental y estudio geomorfológico (Universidad de La Laguna, secretariado de publicaciones, La Laguna, 1991.ç
- Selva, J., Orsi, G., Di Vito, M., Marzocchi, W. and Sandri, L.: Probability hazard map for future vent opening at the Campi Flegrei caldera, Italy, *Bull. Volcanol.*, 74, 497-510, 2012.
- Sheridan, M. F. and Malin, M. C.: Application of computer-assisted mapping to volcanic hazard evaluation of surge eruption: Vulcano, Lipari, Vesuvius. *Explosive Volcanism*, *J. Volcanol. Geoth. Res.*, 17, 187–202, 1983.
- Solana, C., Kilburn, C.R.J. Rodriguez Badiola, E. and Aparicio, A.: Fast emplacement of extensive pahoehoe flow-fields: the case of the 1736 flows from Montaña de las Nueces, Lanzarote, *J. Volcanol. Geoth. Res.*, 132, 189–207, 2004.
- Thordarson, Th., and Self, S.: The Laki (Skaftar Fires) and Grimsvotn eruptions in 1783-1785, *Bull. Volcanol.*, 55, 233–263, 1993.
- Toyos, G. P., Cole, P. D., Felpeto, A., and Martí, J.: A GIS-based methodology for hazard mapping of small pyroclastic density currents, *Nat. Hazards*, 41, 99–112, 2007.

Tables

Island	Map Nº	Local Name	X	Y	Start Simulation point Height (m.a.s.l.)	Collapse equivalent height (Hc) (m)	Run-out (L) (m)	Collapse Equivalent angle (Θ) (°)	Basal Diameter (km)	Type/Characteristics	Trend
LANZAROTE	1	El Golfo	614214	3205971	0	250	2500	5.71	1	Tuff Cone	N50°E
	2	Caldera Blanca	623734	3213091	142	250	3000	7.44	1.8	Maar	N85°E
	3	El Cuchillo	631054	3218877	42	250	3000	5.56	1.4	Tuff Ring	N65°E
	4	Mt. Cavera	637305	3222578	40	250	1500	10.94	0.185	Coastal Eruption;	N33°E
	5	Mt. Chica	636346	3222139	65	250	1500	11.86	0.175-0.25	Wet-surges. Last phases: Strombolian +	N50°E
	6	Mt. Mosta	632977	3219146	87	250	1000	18.62	>0.065	Coastal Eruption	N96°E
	7	Mt. Roja	611455	3193167	13	250	500	27.74	1.4	Tuff Cone	N75°E
	8	Mt. Mojón	623996	3202946	318	250	1000	29.60	0.8*0.625	Tuff-ring	N60°E
	9	Mt. Guatisea/Mt. Blanca	633449	3208190	378	250	1500	22.72	-	Strombolian cone with hydro-magmatic intercalations	N006°E

	10	Mt. Corona	646191	3211411	115	250	1500	13.68	1.2	Strombolian cone with hydro-magmatic intercalations	
	13	Mt. Ubigue	639999	3211732	231.5	250	1500	17.80	1*1	Strombolian cone with hydro-magmatic intercalations	N50°E
	14	Mt. Tinaché	629288	3214639	291	250	1500	19.83	1.25	Strombolian cone with hydro-magmatic intercalations	
	15	Mt. de Halcones	615178	3209072	63	250	1000	17.38	0.65	Strombolian cone with hydro-magmatic intercalations	N50°E
	16	Caldera Riscada	621975	3201907	322	250	5000	6.53	1*0.9	Strombolian and hydromagmatic phases	N60°E
	17	Caldera Gritana	621228	3201274	343	250	5000	6.76	0.65*0.6	Part of a hydromagmatic edifice	N60°E
LA GRACIOSA	18	Mt. Amarilla	642207	3233381	30	250	1000	15.64	0.9*0.65	Strombolian and phreatomagmatic phases	N45°E
	19	Mt. Aguja Grande	645018	3236401	82	250	600	28.96		Strombolian and phreatomagmatic phases	N45°E
MT. CLARA	20	Mt. Clara	642579	3242537	34	250	500	29.60	-	Wet surges	
ALEGRANZA	21	La Rapadura	646207	3252803	2	250	500	26.75	0.42*0.41	First phases Hydromagmatic	
	22	Mt. Lobos	645019	3251867	24	250	1250	12.36	1.2*0.87	First phases hydromagmatic	

	23	La Caldera	643151	3252587	16	250	3000	5.07	2.6 *1.75	Tuff Cone	N65°E
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Table 1. Main characteristics of hydromagmatic eruptions of Lanzarote. Run out distances correspond to minimum L due to these distances have been taken from the maximum exposure deposits on the geological maps. Different parameters have been chosen to simulate PDCs on the island (See section 7.3 for more information).

Volcano-structures	Onshore			Offshore
	Miocene-Pliocene	Pleistocene	Holocene	
Vents	23	419	171	102
Eruptive Fissures	1	69	25	9
Faults	6 (no associated age)			-

Table 2. Number of identified volcanic structures on Lanzarote Island, according to their ages and location

Nº	Structural Datasets	Age	Bandwidth	Weight
1	Miocene-Pliocene Vents and Eruptive Fissures	15 Ma- 2.5Ma	2527	0.107
2	Pleistocene Vents and Eruptive Fissures	2.5 Ma- 11.7 ka	2808	0.207
3	Holocene Vents and Eruptive Fissures	last 11.7 ka	560	0.357
4	Offshore vents and eruptive Fissures	Unknown ages	6508	0.193
5	Faults	Unknown ages	20808	0.136

Table 3. Parameters used for performing susceptibility analysis.

GEOLOGICAL PROCESS- HAZARD	These parameters are mainly derived from 1730-36 and 1824 eruptions				
	Max. Length (km)	Mean Length (km)	Min. Length (km)	Mean Thickness (m)	Total emitted volume (km ³)

Lava Flow	35/25	5-7	1.5	10	0.02-4
	Run out (km); from hydromagmatic eruptions or phreatic phases		Collapse Equivalent angle (°)		
Pyroclastic Density Current	0.5-3		5-29		
	Column height (km)		Size particles (Φ)		
Fallout	3-5		From -6 to 2		

Table 4. Main characteristics of the historical and Holocene eruptions and parameters used for scenario simulations.

Figures

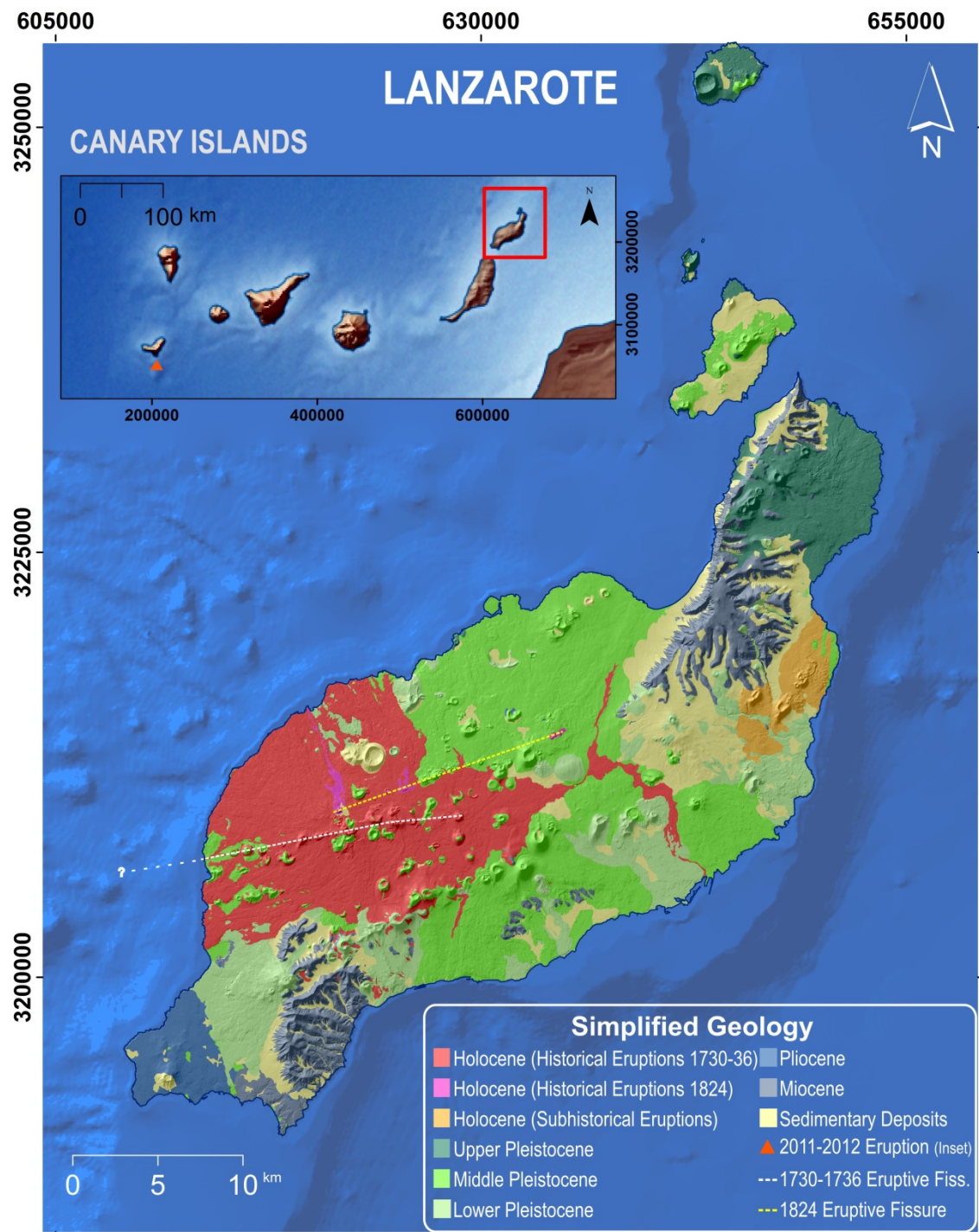


Figure 1. Simplified geological map of Lanzarote Island. The top left inset displays the location of Lanzarote within the Canary Archipelago. (Original geological map can be found in: <http://info.igme.es/cartografiadigital/geologica/Geode.aspx> Sources: <http://info.igme.es/cartografiadigital/geologica/Geode.aspx>).

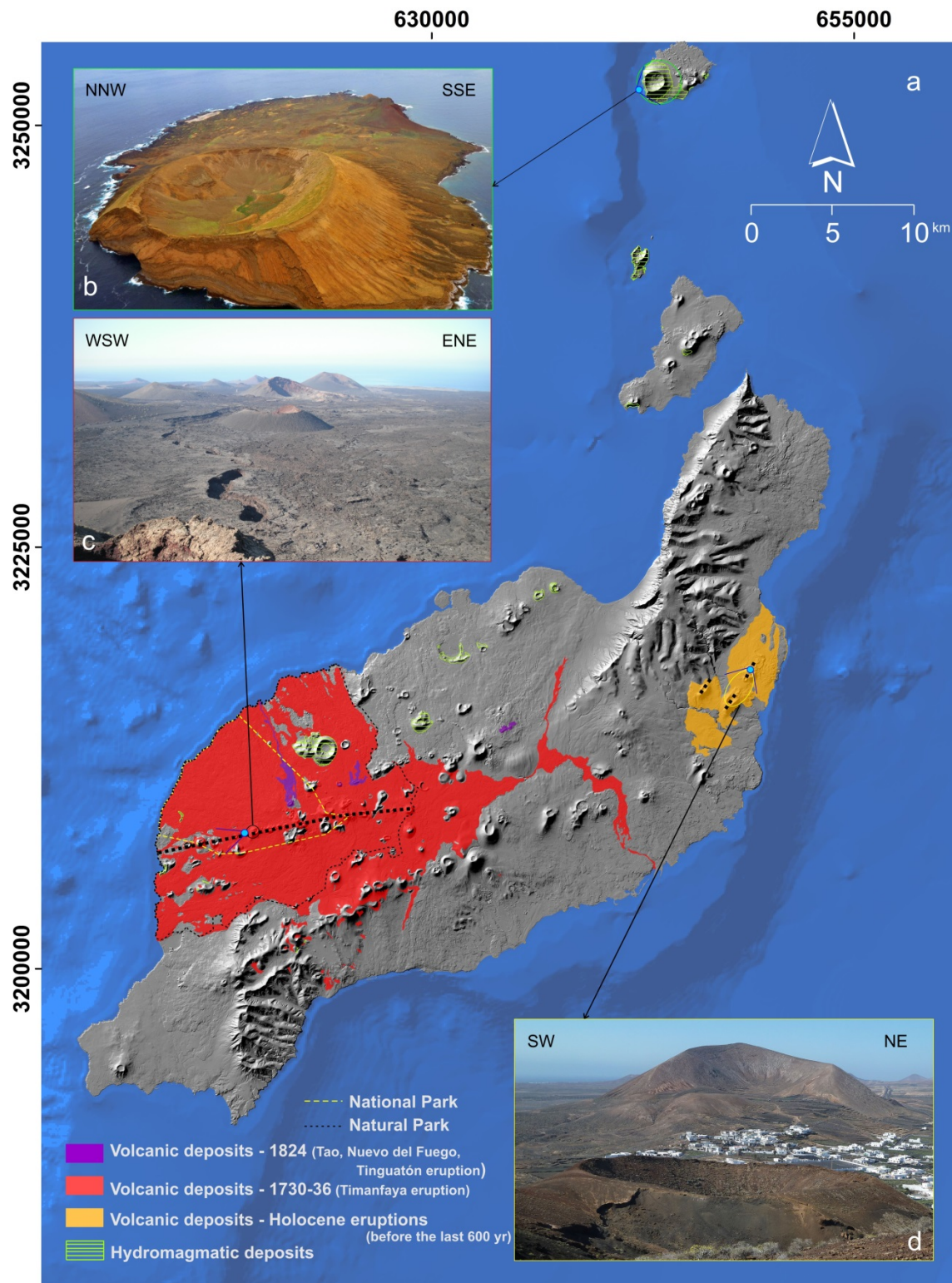


Figure 2. a) Historical eruptions (red, pink and yellow), and hydromagmatic edifices (green) on Lanzarote; b) Alegranza hydromagmatic cone with a diameter of 1.2 km; c) Timanfaya cones; d) Mt. Guenia and La Caldereta cones. Yellow and black dashed lines define the limits of the Timanfaya National Park and the Natural park, respectively.

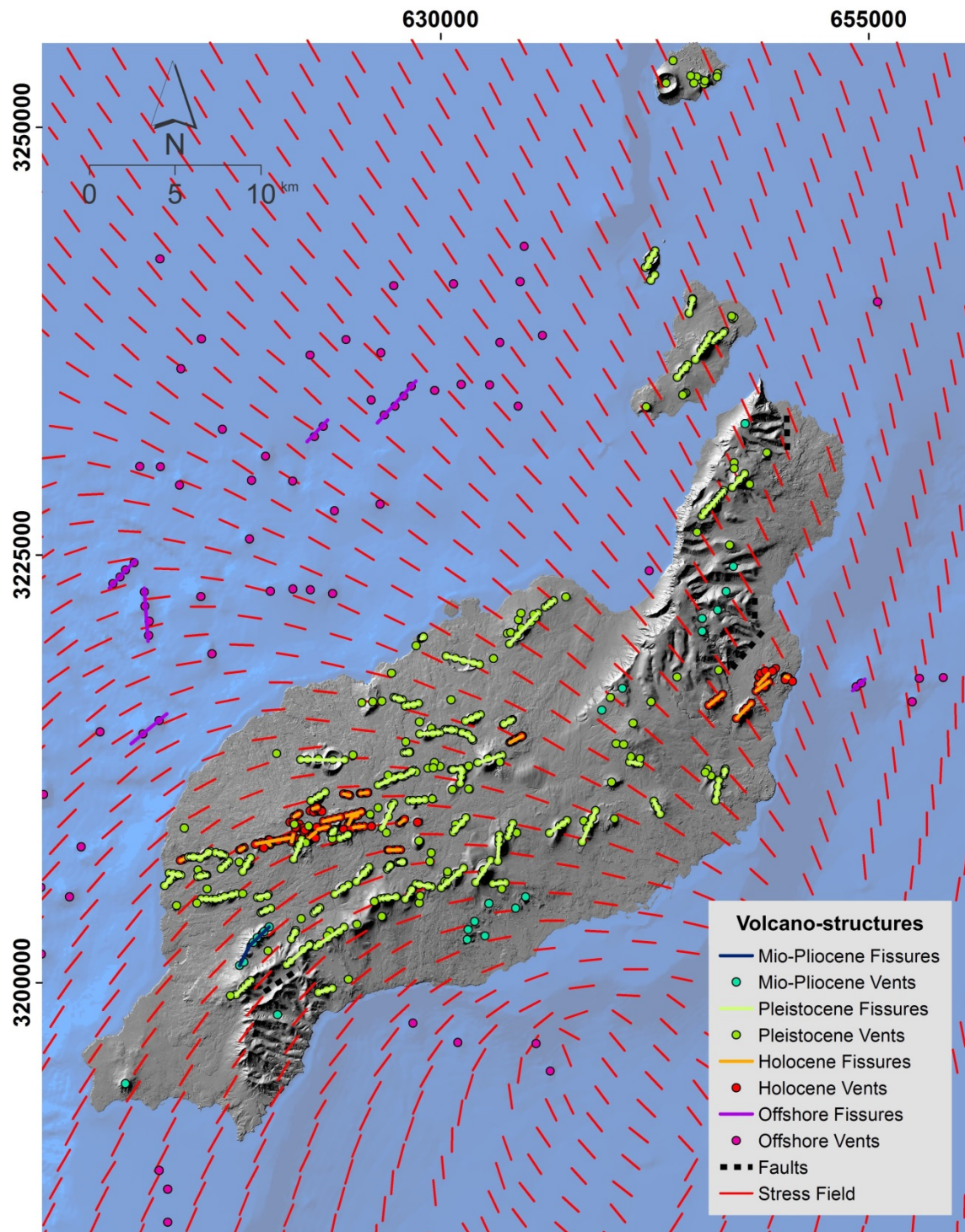


Figure 3. Volcano-structural datasets defined for Lanzarote and used for evaluating spatial probability. Maximum compressive horizontal stress trajectories are also indicated (red lines).

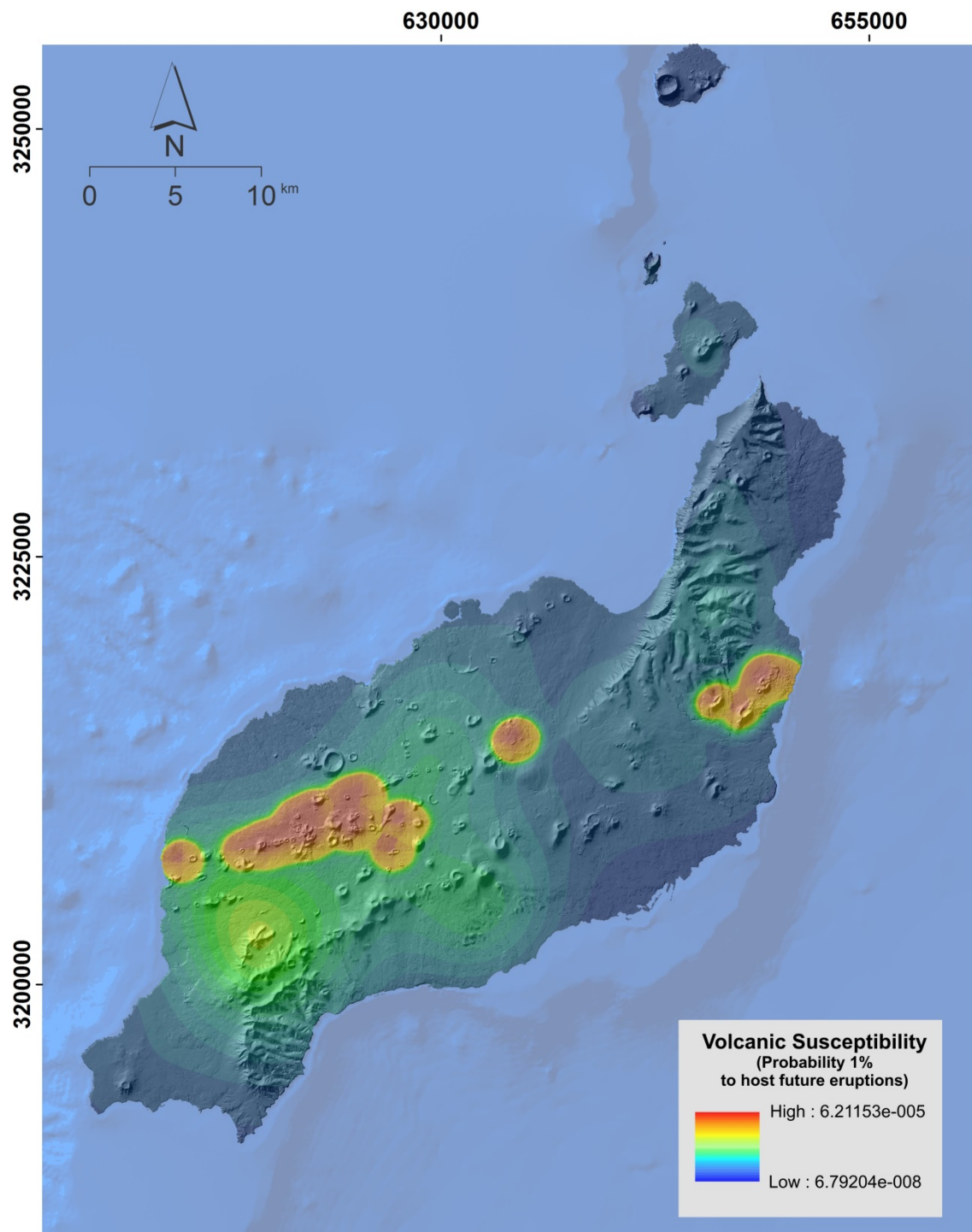
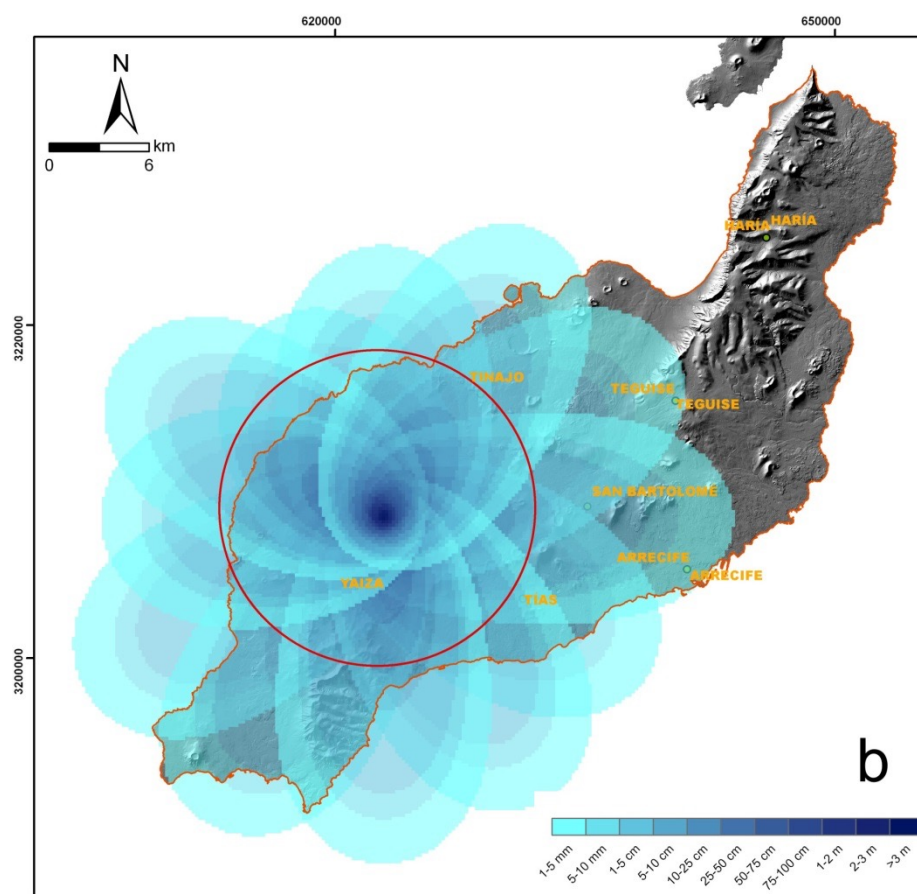
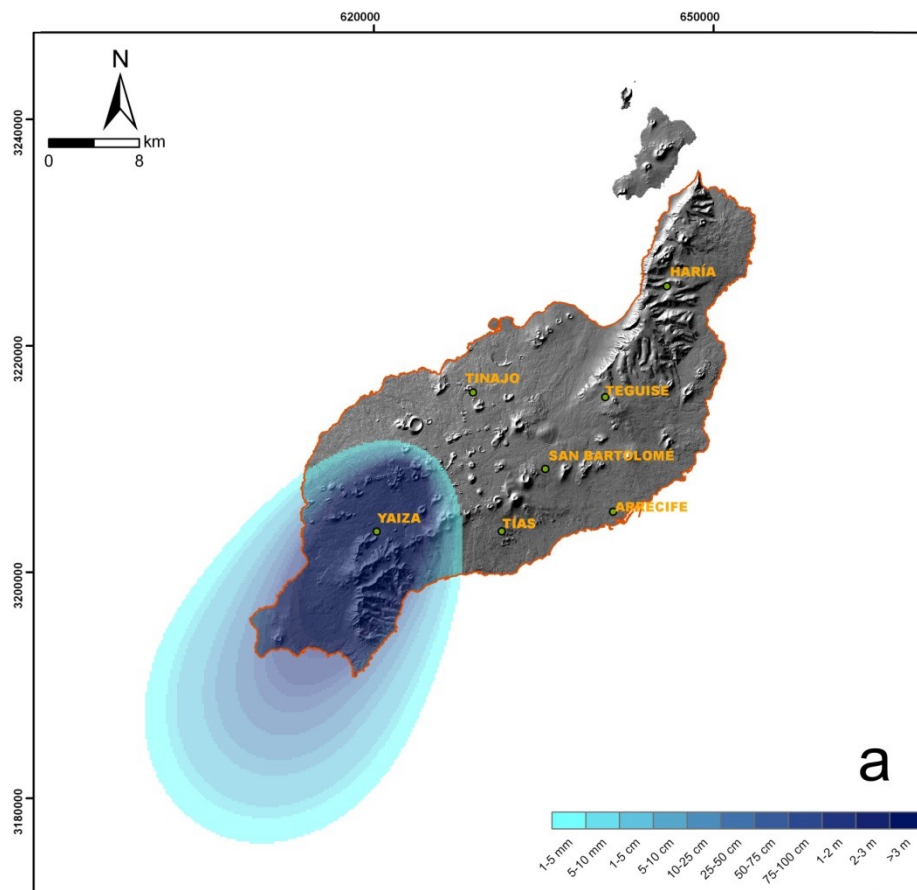


Figure 4. Volcanic Susceptibility map of Lanzarote Island. The highest probability (0.00006) of new vent opening is obtained in a along a NE-SW area. High probabilities are also observed in the South of Timanfaya Park the island.



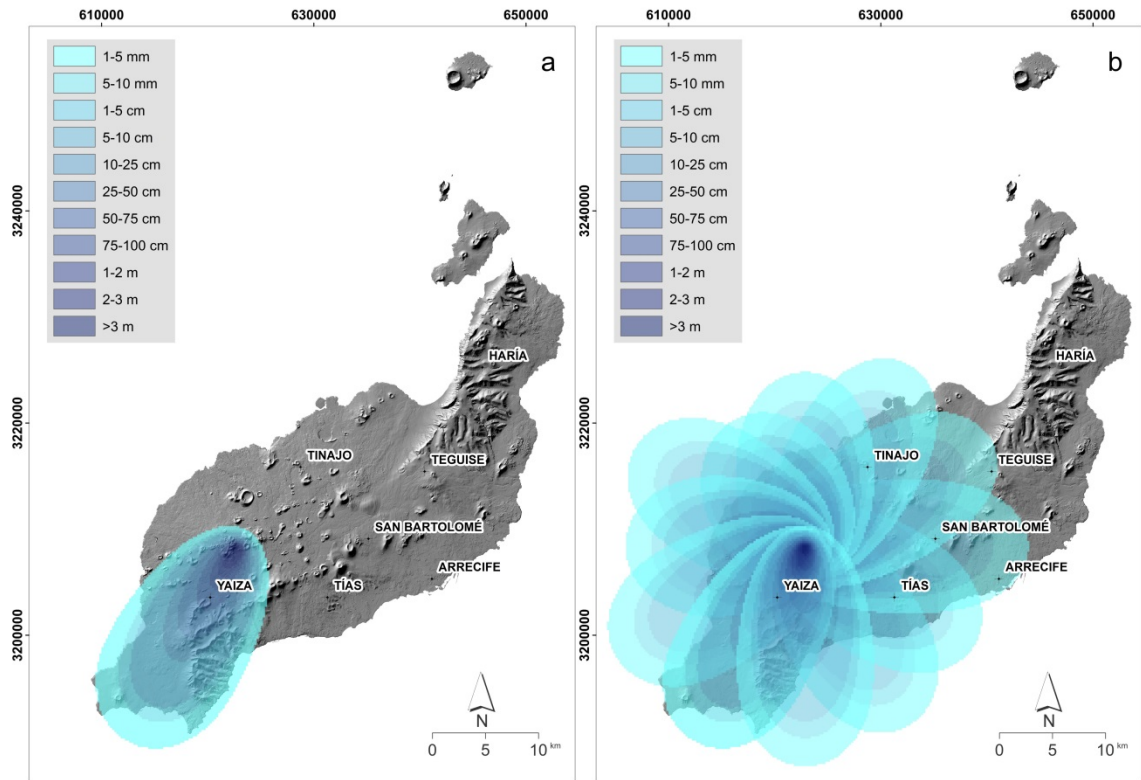
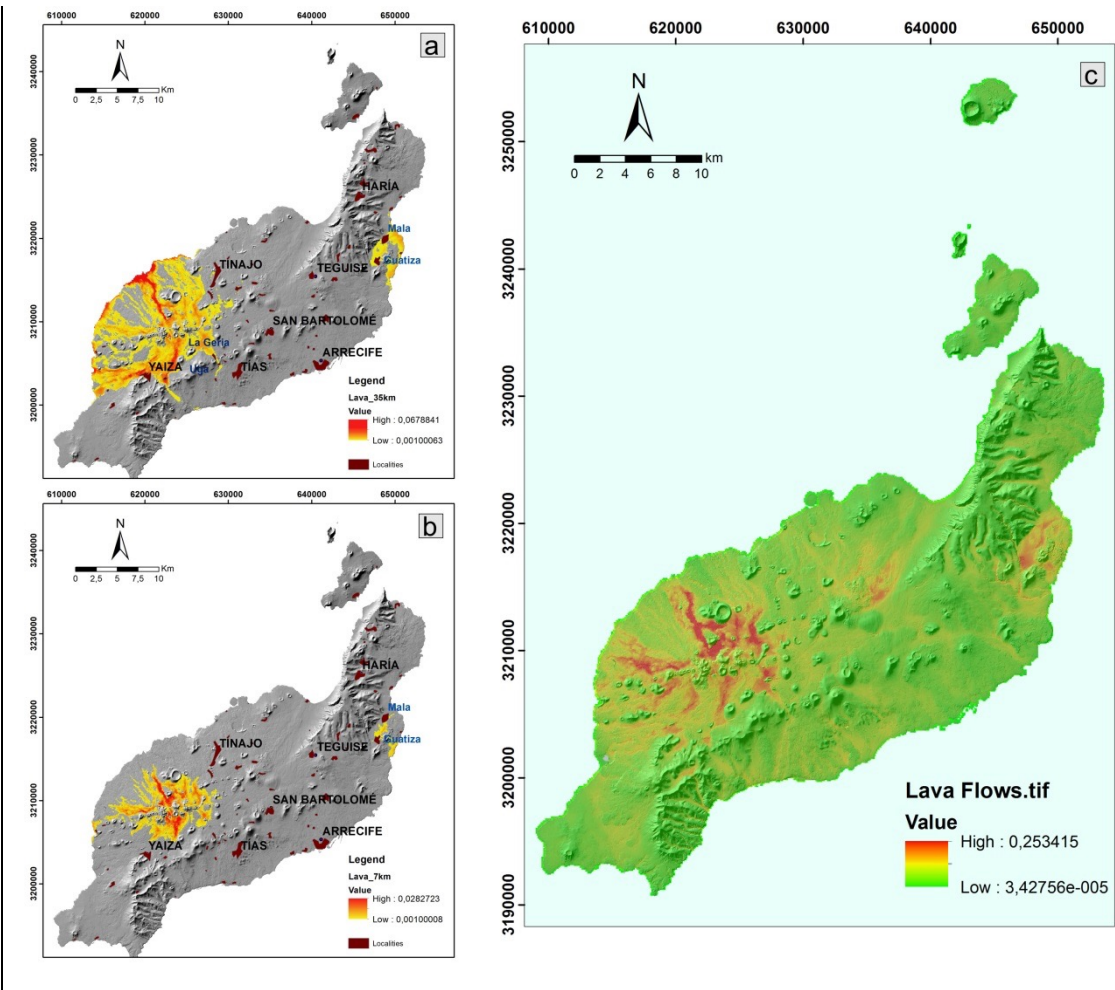


Figure 5. Fallout scenarios at the highest probability vent for the NE wind direction and for the entire wind rose directions performed with VORIS 2.0.1. a) NE wind simulation assuming a Strombolian eruption; b) 1824 eruption. Main localities have been placed in order to show which ones would be affected by the ashfall dispersion.



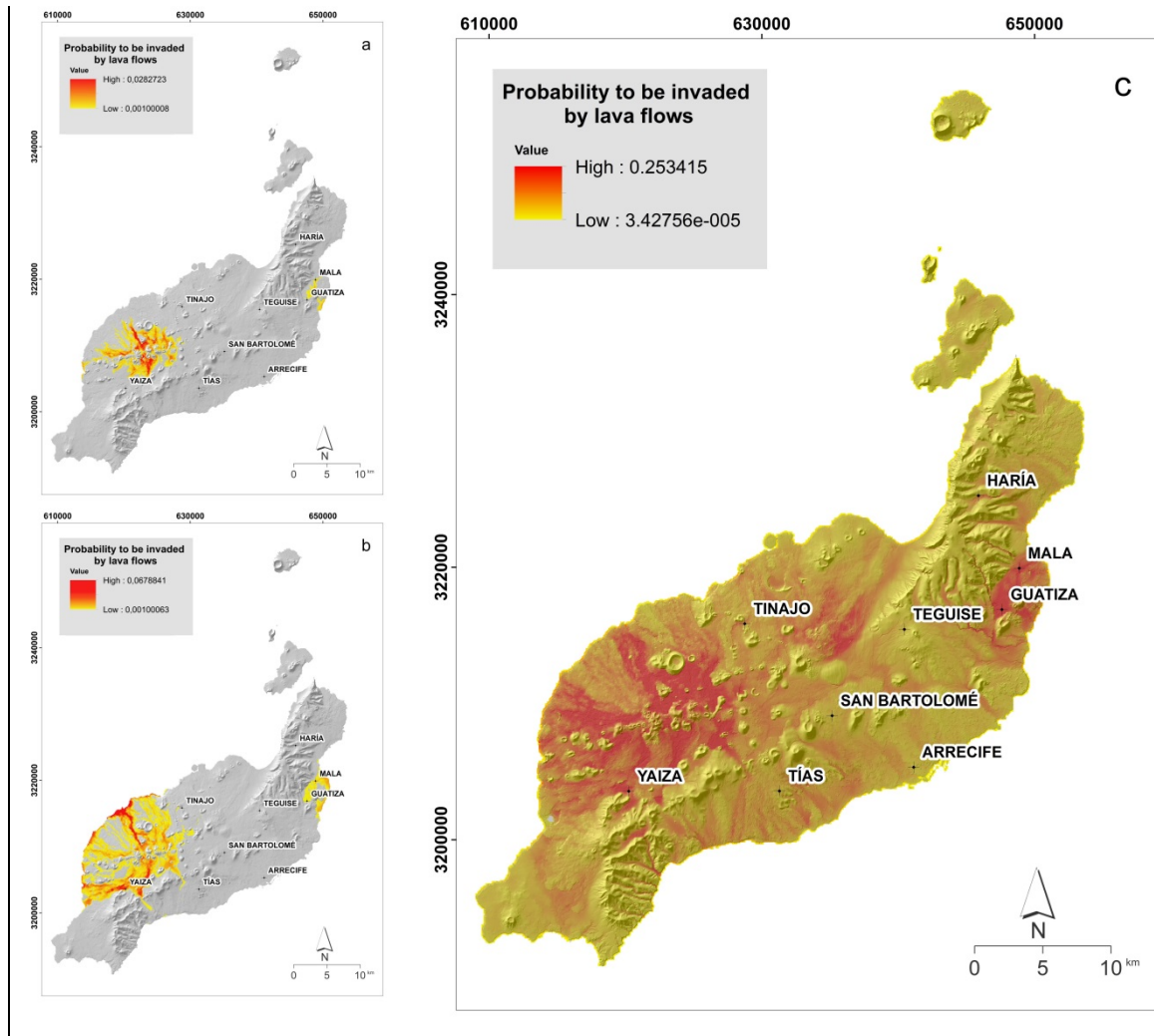
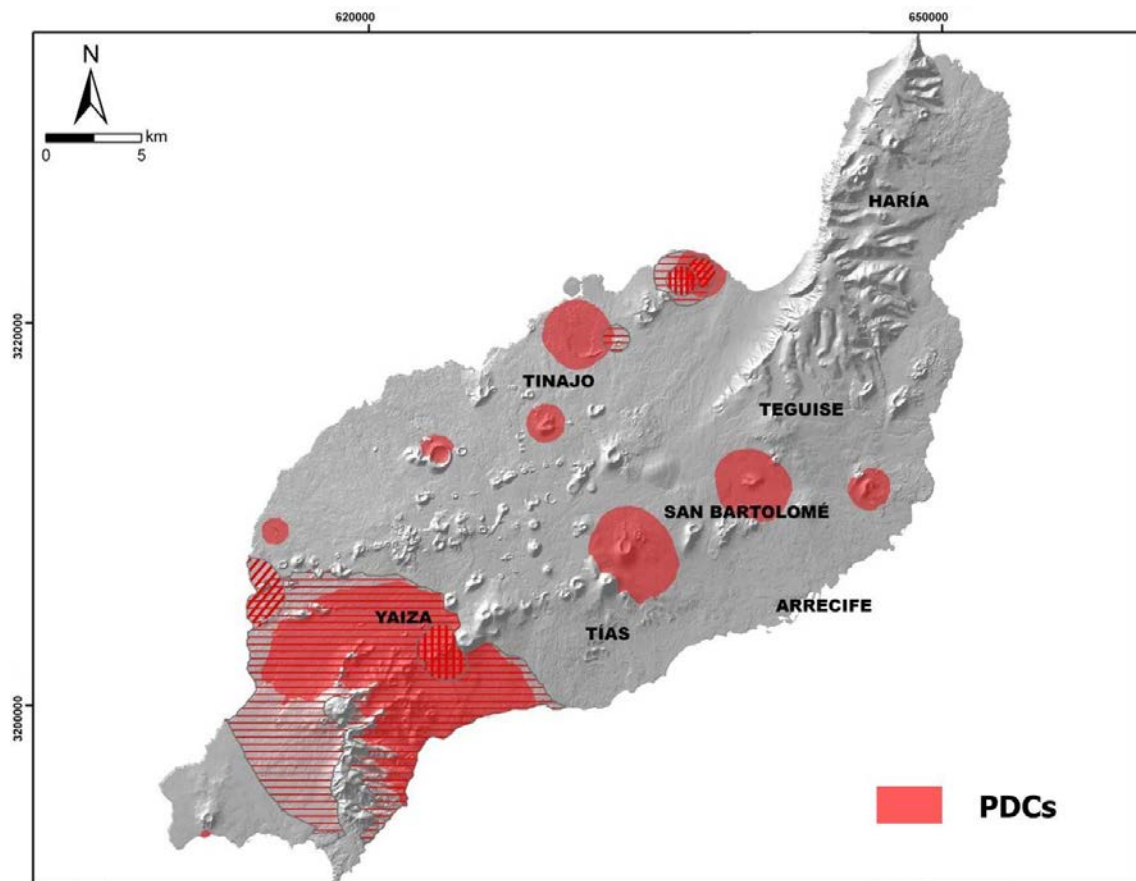


Figure 6. Lava flow scenarios for Lanzarote performed with VORIS 2.0.1. a) Timanfaya scenario; b) 1824 eruption scenario; c) Total lava flow map. Red colours are those areas with the highest probability to be invaded by lava flows.



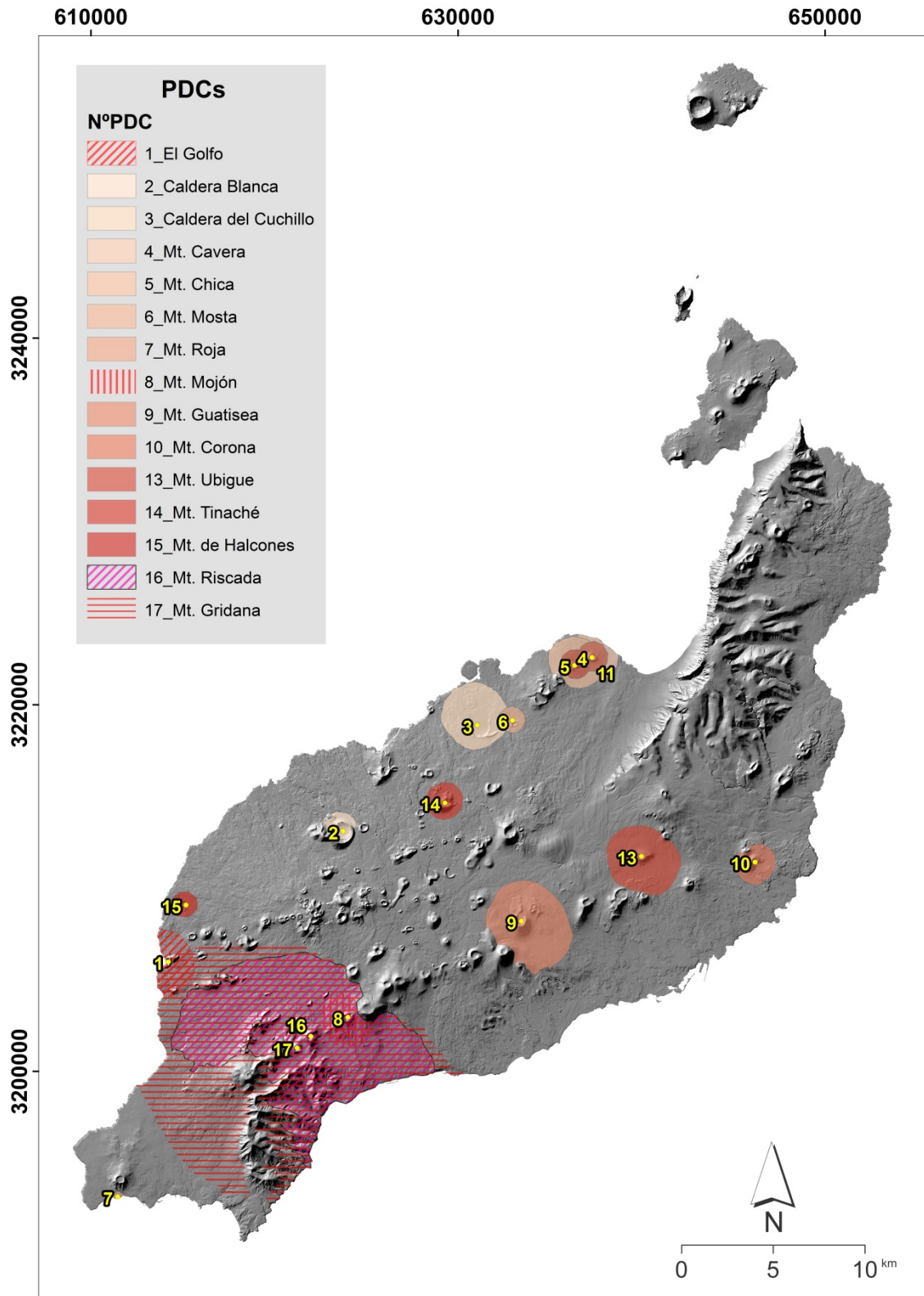


Figure 7. PDC scenarios performed with VORIS 2.0.1. Covered areas with different collapse equivalent heights (H_c), and collapse equivalent angles (θ) and VEI values (see the text for more detail). Different symbols (dashed, filled and coloured) have been used to show the limits of each PDC.