

1 **Brief Communication**

2 **Co-seismic displacement on October 26 and 30, 2016 ( $M_w$  5.9 and 6.5) -**  
3 **earthquakes in central Italy from the analysis of a local GNSS network**

4  
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15  
16 **1 - Abstract**

17 On August 24<sup>th</sup> 2016 a strong earthquake ( $M_w = 6.0$ ) affected Central Italy starting and an intense seismic  
18 sequence. Field observations, DInSAR analyses, preliminary focal mechanisms as well as the distribution of  
19 aftershocks suggested the reactivation of the northern sector of the Laga Fault, whose southern sector was  
20 already reactivated rebooted during the 2009 L'Aquila sequence, and of the southern segment of the Monte  
21 Vettore Fault System (MVFS). Based on these preliminary information and following the stress-triggering  
22 concept (Stein et al., 1999; Steacy et al., 2005), we tentatively identified a potential fault zone most  
23 vulnerable to future seismic events just north of the first earlier epicentral area. Accordingly, we planned a  
24 local geodetic network consisting of five new GNSS (Global Navigation Satellite System) stations located at  
25 few km of distance on both sides of the MVFS. This was devoted to picture out, at least partially but in some  
26 detail, the possible northward propagation of the crustal network ruptures. The building of the stations and  
27 a first set of measurements were carried out during a first campaign (September 30<sup>th</sup>-October 2<sup>nd</sup>, 2016). On  
28 October 26<sup>th</sup> 2016, immediately north of the epicentral area of the August 24<sup>th</sup> event, a moderate another  
29 earthquake ( $M_w = 5.9$ ) indeed occurred, followed four days later (October 30<sup>th</sup>) by the mainshock ( $M_w = 6.5$ )  
30 of the whole 2016 Summer-Autumn seismic sequence. Our local geodetic network was fully affected by the  
31 new events and therefore we performed a second campaign soon after (November 11<sup>th</sup>-13<sup>th</sup>, 2016). In this  
32 brief note, we provide the results of our geodetic measurements that registered the co-seismic and  
33 immediately post-seismic deformation of the two major October shocks documenting in some detail the  
34 surface deformation close to the fault trace. We also compare our results with the available surface  
35 deformation field of the broader area, obtained on the basis of the DInSAR technique, and showing an overall  
36 good fit.

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38 **2 - Geological framework**

39 The central Apennines are characterized by northeast-verging thrust-propagation folds, involving Mesozoic-  
40 Tertiary sedimentary successions. During the 2016 sequence, coseismic deformation was recorded at the  
41 rear of the Sibillini Thrust that separates the homonymous mountain chain from the Marche-Abruzzi foothills  
42 (Fig. 1). According to many studies in the area affected by the recent earthquakes, the main thrust-related

43 anticlines and associated reverse faults have been dissected and/or inverted by NNW-SSE trending  
44 Quaternary normal and oblique-slip faults (Figs. 1 and 2), in particular by the Norcia Fault System (NFS)  
45 (Calamita and Pizzi, 1992; Calamita et al., 1982; 1995; 1999; 2000; Blumetti et al., 1990; Blumetti, 1995;  
46 Brozzetti and Lavecchia, 1994; Cello et al., 1998; Galadini and Galli, 2000; Pizzi and Scisciani, 2000; Pizzi et al.,  
47 2002; Boncio et al., 2004 Galadini, 2006; Gori et al., 2007) and the Mt. Vettore Fault System (MVFS) (Calamita  
48 and Pizzi, 1991; Coltorti and Farabollini, 1995; Cello et al., 1997; Pizzi et al., 2002; Galadini and Galli, 2003;  
49 Pizzi and Galadini, 2009) (Figs. 1 and 2). Conversely, Pierantoni et al. (2013) suggest that the major Mt. Sibillini  
50 Thrust has **not** been **not** yet dissected by **q**uaternary normal faulting, though some fresh morphological  
51 scarps with free faces in the carbonate bedrock and/or affecting recent slope deposits have been observed  
52 and attributed to the local seismic activity.

53 Within a distance of few tens of kilometers, large evidence **of this Quaternary seismotectonic behaviour**  
54 **ground deformation** has been provided by several recent earthquakes, like the 1979 Norcia event ( $M_w$  5.9,  
55 reactivating the Norcia Fault; e.g. Deschamps et al., 2000), the 1984 Gubbio ( $M_w$  5.6, Gubbio Fault; e.g. Boncio  
56 et al., 2004), the 1997 Colfiorito ones ( $M_w$  5.7, 6.0 and 5.6, Calfiorito-Cesi-Costa fault system; e.g. Cello et al.,  
57 1997), the 2009 L'Aquila mainshock and the Campotosto aftershock ( $M_w$  6.3 and 5.4, Upper Aterno Valley-  
58 Paganica fault system and Gorzano Fault; Blumetti et al. 2013) and basically the same occurred with the  
59 2016 seismic sequence.

60 Surface evidence of the August 24<sup>th</sup> (e.g., EMERGEO WG, 2016; Livio et al., 2016; Aringoli et al., 2016) **was**  
61 mainly **occurred** **observed** in the area of the Laga basin (Gorzano Fault), which corresponds to the footwall  
62 block of Sibillini Thrust, while debated ground ruptures (e.g. Valensise et al., 2016) also occurred in the  
63 southern sector of the MVFS, which belongs to the hanging-wall block of the orogenic structure. In contrast,  
64 as a consequence of the mainshock of October 30<sup>th</sup>, the entire western flank of the Monte Vettore was  
65 affected by impressive geological effects and clear coseismic ruptures mapped for a minimum length of 15  
66 km, between the Castelluccio di Norcia and Ussita (EMERGEO WG, 2016) (Fig. 2). The surface ruptures  
67 occurred along distinct fault splays of the fracture system. For example, along the western slope of Monte  
68 Vettore three main west dipping splays were activated together with two antithetic branches (Figs. 1 and 2).  
69 The observed vertical offset reached 2 m along the main west dipping fault segment, where the slickensides  
70 show a prevalent dip-slip component of motion. Vertical displacements of a few centimetres were also  
71 recorded along an antithetic surface rupture bordering to the west the Castelluccio plain, about 6-7 km far  
72 from the main ground rupture, possibly connected to a secondary fault (Figs. 1B and 2).

73 It is worth to note that **the** August-October earthquakes occurred in a sector of the Central Apennines  
74 characterized by high geodetic strain-rates (e.g., Devoti et al., 2011; D'Agostino 2014), where several  
75 continuous GNSS stations are operating.

76

## 77 **Implementation and Analysis of UNICT discrete GPS stations**

78 Following the August, 24<sup>th</sup>,  $M_w$  6.0 earthquake, the GEOMATIC Working Group of the Catania University  
79 (UNICT) in collaboration with the SpinOff EcoStat s.r.l. and researches of the Ferrara University, started a  
80 detailed monitoring of ground deformation in the epicentral area using the Global Navigation Satellite System  
81 (GNSS) technique. The GNSS measurement has been made in static mode, setting the time at 6 hours and  
82 post-processing position, in order to reduce tropospheric error and using IGS precise products for orbits. The  
83 IGS station coordinates were kept fixed in order to align the **final** **final** velocity field with the WGS84 reference  
84 frame. The measurement mode, adopted for receiver-satellite range determination, is performed with a  
85 double frequency receiver, allowing phase and code measurements on the signal carrier (L1, L2, C1, P1, P2,  
86 S1, S2). The coordinates estimation is based on the principle of minimum squares.

87 For this aim, five GNSS stations have been installed on new benchmarks purposely built by the working group  
88 and here referred to as UNICT network (Fig. 3). These new stations have been realized taking into account  
89 the following criteria:

- 90 I. the distribution of the existing permanent and discrete measurement benchmarks belonging to  
91 different networks that were active before the event of 24 August (IGM; RING; CAGEONET; DPC;  
92 ISPRA) (Fig. 2B).
- 93 II. the seismotectonic setting of the area in relation to the macroseismic data and to the reactivated  
94 structures (Figs. 1 and 2);
- 95 III. surface and deep geometry of the major faults related to tectonic setting (Fig. 1B).
- 96 IV. the lack of possible gravitational instabilities in both static and dynamic conditions in sites where the  
97 new benchmarks are built.

98 Based on the above criteria, the working group installed the benchmarks at the bottom of both western and  
99 eastern slopes of Mt. Vettore, within an area about 8 km-long and 5 km-wide in the N-S and E-W directions,  
100 respectively. The distribution of the benchmarks was planned for ~~depicting~~ ~~reconstructing~~ the principal  
101 deformation zone developed as a consequence of the August 24<sup>th</sup> event (Fig. 2 ~~and 3~~) and particularly with  
102 points:

- 103 I. much closer to the epicentral area than the already existing ones belonging to other networks (Fig.  
104 2B);
- 105 II. characterized by equivalent distances from the reactivated Mt. Vettore Fault segments (Fig. 2);
- 106 III. within a distance of 30 km from the closest permanent network points that have been not affected  
107 by deformation, therefore allowing a rigorous elaboration during the post processing phases ~~(Fig. 3)~~.

108 The ~~realization~~ ~~building~~ of GNSS monument on the UNICT benchmarks consists of the following steps ~~(Fig.~~  
109 ~~3)~~:

- 110 I. selection of a suitable ~~point~~ ~~site~~, corresponding to a massive rocky outcrop or a man-made  
111 monument with foundation; these sites must be also free of structures or other natural elements in  
112 the surroundings that may constitute a perturbation during recording;
- 113 II. testing of GPS signal reception by short-term exams, and control of parameters set ~~by~~ ~~through~~ the  
114 quality check carried out by software TEQC ([http://www.unavco.org/software/data-  
115 processing/teqc/teqc.html](http://www.unavco.org/software/data-processing/teqc/teqc.html));
- 116 III. implementation of the hole for housing the bushing and check of its verticality; the hole has a  
117 diameter of 35 mm and a depth of 100 mm, it is realized through small-sized battery-powered  
118 equipment (Makita DHR243 hammer drill) ~~(Fig. 3)~~;
- 119 IV. fixing and anchoring of the knurled steel bushing (length 67 mm and diameter 20 mm), with bi-  
120 component resins or quick-setting cements (Fig. 3);
- 121 V. following the cementation to the artefact or to rocky outcrop, a male-male threaded bar can be  
122 screwed in until end of stroke; the height could be variable and this fact ~~will be~~ ~~is~~ considered in the  
123 data processing. We have used a threaded bar 670 mm-high. ~~(Fig. 3)~~.

124 The GPS monument is thus completed with a GNSS receiver TOPCON, mounting a HiPer V antenna,  
125 characterized by 226 channels and position accuracy with band L1+L2 in Static mode of 3 mm + 0.1 ppm  
126 (horizontal) and 3.5 mm + 0.4 ppm (vertical). All registrations last six hours in static mode.

127 Following the August 24<sup>th</sup> event, at the end of September 2016 the working group ~~have started~~ ~~curried out~~  
128 the first survey campaign with the installation of five UNICT benchmarks: two stations were located east of

129 the Mt. Vettore fault (VTE1,VTE2), the other three (VTW3,VTW4, VTW5) west of the fault (Figs. 4 and 5).  
 130 During November 2016 (*i.e.* after the October 30<sup>th</sup> event), a second field campaign was carried out following  
 131 the same procedure and using the same instrumentation. The second set of measurements allowed  
 132 recording us to record the co-seismic displacement caused by both the  $M_w$ 5.9 and  $M_w$ 6.5 events of October  
 133 26<sup>th</sup> and 30<sup>th</sup>, respectively (doy (day of year) 2016/274 and doy 2016/318).  
 134 The data from survey-mode GNSS stations have been downloaded and processed using TOPCON Magnet  
 135 analysis software evaluating co-seismic solutions and comparing with AUSPOS web-based online services for  
 136 GPS data processing (Ocalan et al., 2013), whose engine is based on Bernese 5.2 software. In the software  
 137 TOPCON, the baseline is automatically created for any pair of static occupations, where we set up six hours  
 138 for Minimum Duration and the baselines max length of 50 km, cut-off angle of 15°, troposphere model Goad-  
 139 Goodman and, finally, meteo model NRLMSISE (neutral temperature and densities in Earth's atmosphere).  
 140 For the analyses we referred to the measurement of a stable reference frame of five GNSS stations belonging  
 141 to the RING (Rete Integrata Nazionale GPS) network, with a maximum baseline length of 50 km, using stations  
 142 CESI, GNAL, GUMA, MTER and MTT0 (Figs. 4 and 5). Data processing has been carried out with adjustment  
 143 by Least Squares and a TAU Criterion.  
 144

ID	Station	Longitudine	Latitudine	disp <sub>N-S</sub>	disp <sub>E-W</sub>	disp <sub>UP</sub>	unc <sub>N-S</sub>	unc <sub>E-W</sub>	unc <sub>UP</sub>
VTE1	FOCE_SENTIERO	13° 15' 57,45166"	42° 51' 57,04340"	141	312	29	15.5	16.5	44.0
VTE2	PRETARE	13° 16' 33,20959"	42° 47' 56,56780"	60	282	67	19.0	16.5	46.0
VTW3	QUARTUCCIOLO	13° 14' 46,41153"	42° 47' 56,57032"	198	26	-349	15.5	14.5	36.0
VTW4	COLLE_CURINA	13° 13' 55,01245"	42° 48' 59,62491"	102	288	-769	15.5	15.0	36.0
VTW5	CASTELLUCCIO_VALLE	13° 12' 56,20423"	42° 49' 54,89014"	353	418	-707	15.0	13.5	37.5

145  
 146 *Tab 1 - Three components co-seismic displacements and relative uncertainties estimated for the GNSS stations of the*  
 147 *UNICT network. Coordinates are WGS84 east and north, respectively. All displacement and uncertainty values are in*  
 148 *millimeters. For all stations, the cut-off angle is 15°, the troposphere model is the Goad-Goodmar and the meteo model*  
 149 *used is NRLMSISE. The table can be download as ASCII file on the INGVRING web page (<http://ring.gm.ingv.it>).*  
 150

## 151 Concluding remarks

152 Using the GNSS technique, we investigated the ground deformation that occurred in the surroundings of the  
 153 Mt. Vettore Fault System during the 2016 central Italy seismic sequence. This foresight action allowed us to  
 154 record the co-seismic and part of the post-seismic deformation of the second and third (strongest) events  
 155 ( $M_w$  5.9 and  $M_w$  6.5) on October 26<sup>th</sup> and October 30<sup>th</sup>, 2016, respectively. Taking into account the geometry  
 156 of the fault system in the broader epicentral area and following the stress-triggering concept (Stein et al.,  
 157 1999; Steacy et al., 2005), we have identified a potential fault zone most vulnerable to future seismic events  
 158 just north of the fault segment reactivated during the August 24<sup>th</sup> earthquake (Figs. 2B and 5). With this in  
 159 mind, in order to measure the post seismic deformation and to possibly record the potential migration of the  
 160 co-seismic process, we selected some sites and built five new GNSS benchmarks, distributed east and west  
 161 of the northern-central segment of the Mt. Vettore Fault System. For site selection we also considered the  
 162 presence and distribution of other benchmarks located before the second seismic event by other research  
 163 groups (IGM; RING; CAGEONET; DPC; ISPRA). The epicentral location of the October events eventually  
 164 confirmed our guess and we then we performed soon after a second campaign of measurements for  
 165 quantifying the relative motion of the stations.  
 166 The measured deformation (with 95% confidence errors) is characterised by both horizontal and vertical  
 167 movements. In particular, the east benchmark VTE1 documents recorded 312 mm of eastward horizontal  
 168 displacement and 29 mm of upward motion, while the VTE2 recorded 282 mm of eastward horizontal

169 displacement and 67 mm of upward component of motion. On the contrary, all three western benchmarks  
170 recorded westward horizontal displacements (419, 288 and 26 mm) and subsidence (707, 288 and 769 mm)  
171 for stations VTW5, VTW4 and VTW3, respectively. In conclusion, we documented ca. 730 mm of ENE-WSW  
172 lengthening on a distance of 7 km in correspondence of the northern sector of the Mt. Vettore Fault Segment,  
173 while the off-fault vertical displacement between footwall and hanging-wall blocks was 736 mm.

174 We also compared our results with the displacement distribution obtained by other research group with  
175 DInSAR techniques, recorded between October 26th 2016 (pre-event images) and November 1st 2016 (post-  
176 event images), and other GNSS stations, active before the second seismic event. In Fig. 5 we may observe the  
177 overall consistency of the different approaches and datasets.

178  
179

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 279 *Ann. Geophysics*, 59, fast track 5, 2016; doi: 10.4401/ag-7215.

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281 **Website links**

282 <http://iside.rm.ingv.it>

283 <http://ring.gm.ingv.it>

284 <http://www.igmi.org/geodetica/>

285 [http://www.irea.cnr.it/index.php?option=com\\_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1](http://www.irea.cnr.it/index.php?option=com_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1)

287 [http://www.irea.cnr.it/index.php?option=com\\_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1](http://www.irea.cnr.it/index.php?option=com_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1)

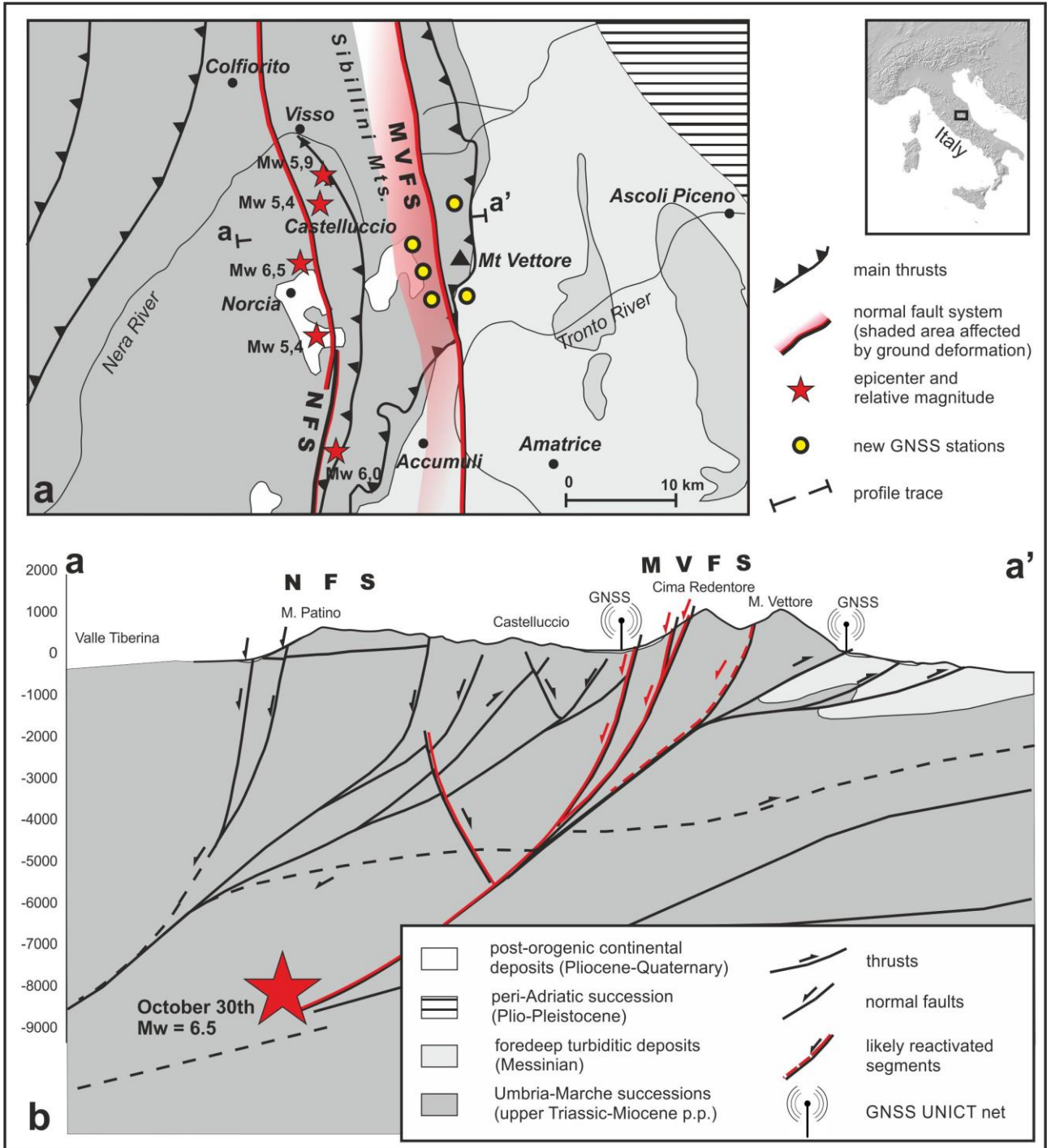
289 <http://www.isprambiente.gov.it>

290 <http://www.unavco.org/software/data-processing/teqc/teqc.html>

291 Rete Integrata Nazionale GPS - <http://ring.gm.ingv.it/>

292 <http://terremoti.ingv.it/it/ultimi-eventi/1001-evento-sismico-tra-le-province-di-rieti-e-ascoli-p-m-6-0-24-agosto.html>; Sequenza sismica di Amatrice, Norcia, Visso: approfondimenti e report scientifici

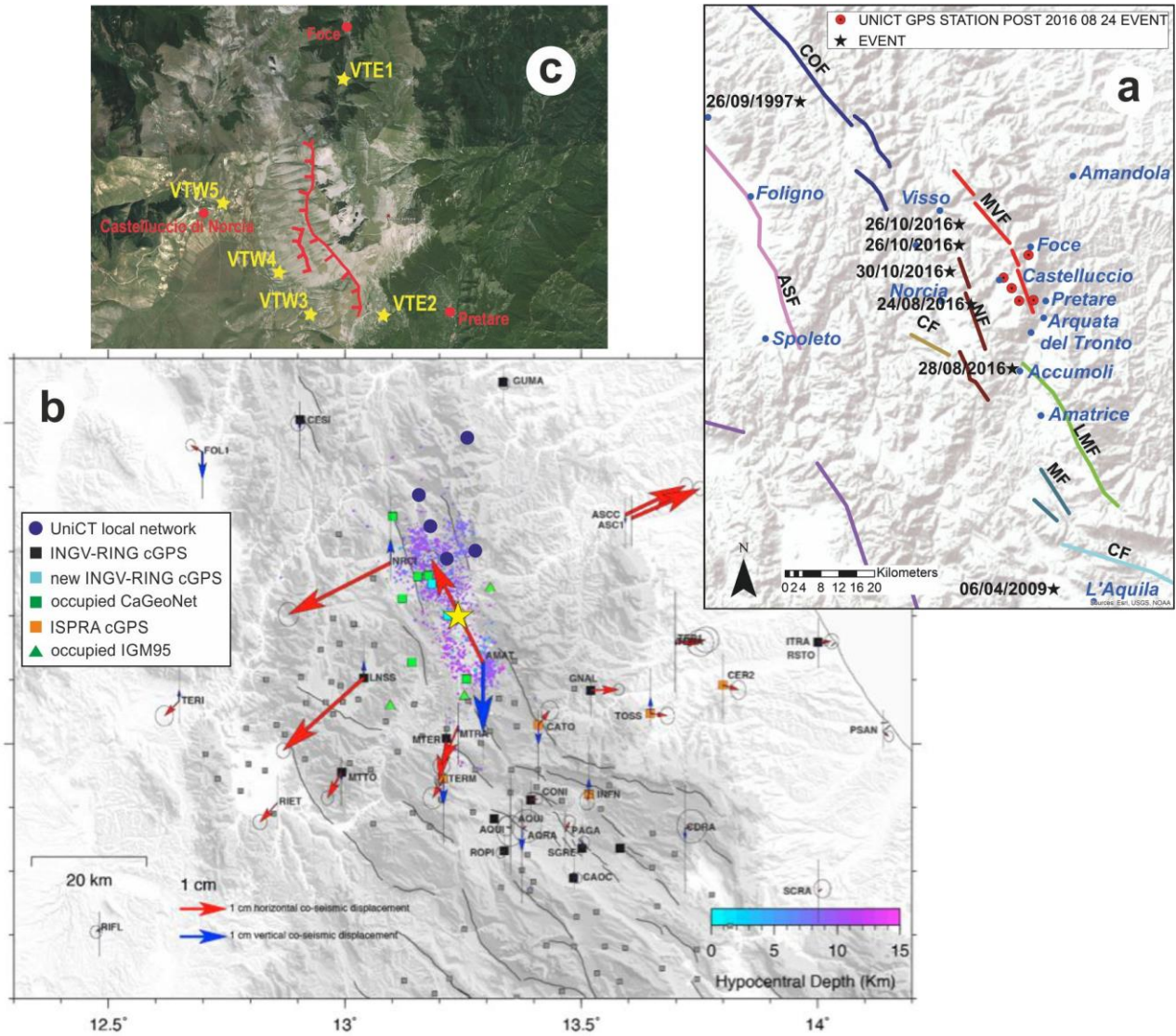
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Fig. 1 Simplified seismotectonic map of central Apennines (A) and geological profile across the epicentral area (B). The location of the major event (October 30th) is from GdL INGV (2016), while the main geostructural features from Pierantoni et al. (2013) and Mantovani et al. (2011) modified).





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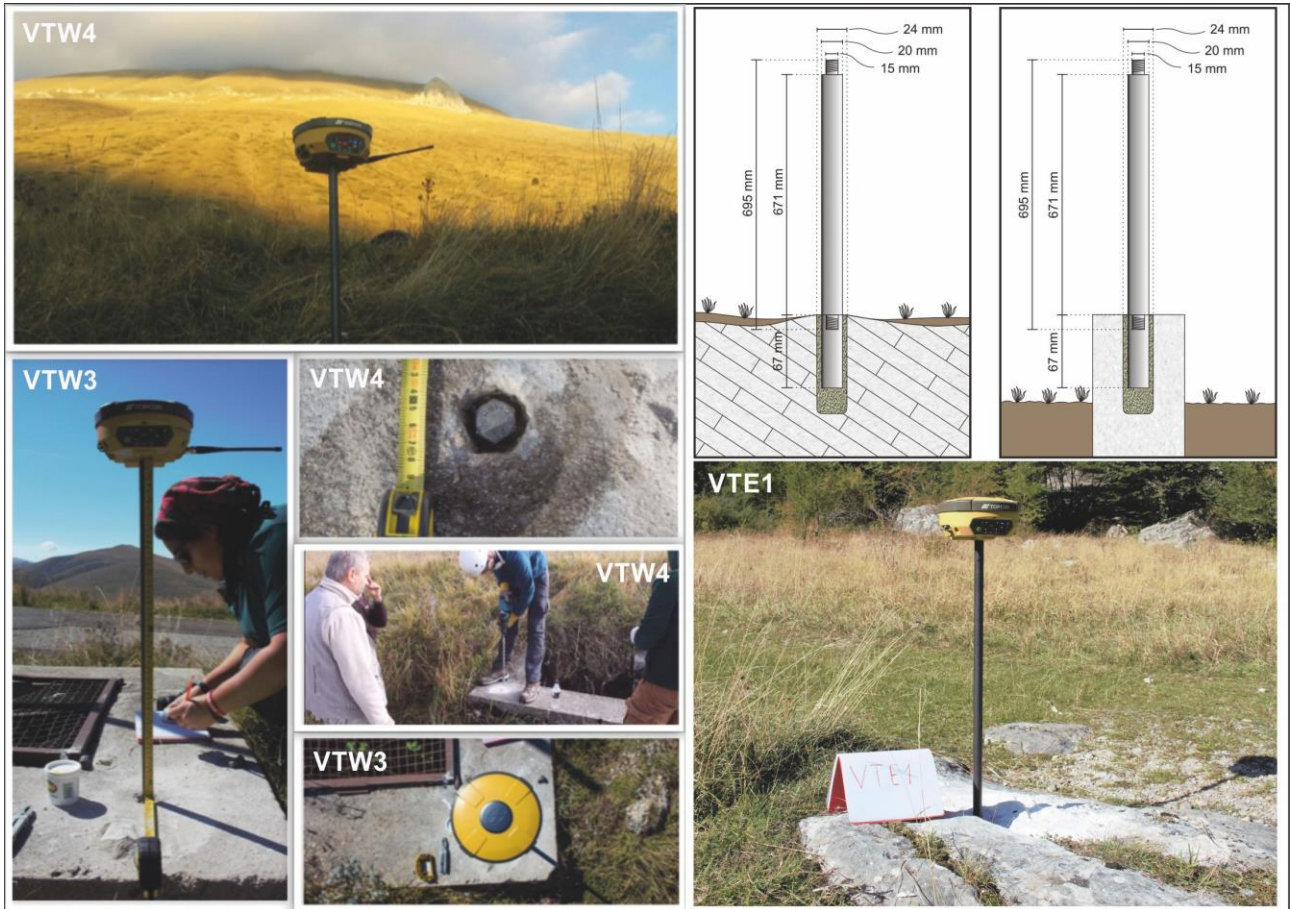
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Fig. 2 – a) Digital Elevation Model with shaded relief of central Apennines showing the active fault system and the major events since 1997 (ASF: Assisi Fault; COF: Colfiorito Fault; CF: Cascia Fault; MVF: Mt. Vettore Fault; NF: Norcia Fault; LMF: Laga Mts. Fault; MF: faults of the Montereale basin). b) Horizontal (red arrows) and vertical (blue arrows) consensus co-seismic displacements (with 68% confidence errors), and the local UniCT GPS network. The aftershocks of the August 24<sup>th</sup>, Mw 6.0 main event (yellow star) are colored as a function of depth (from <http://iside.rm.ingv.it>); c). GoogleEarth map showing the new five GNSS stations (yellow stars) located in the near field of (and surrounding) the October 30<sup>th</sup> coseismic ground ruptures (red lines).

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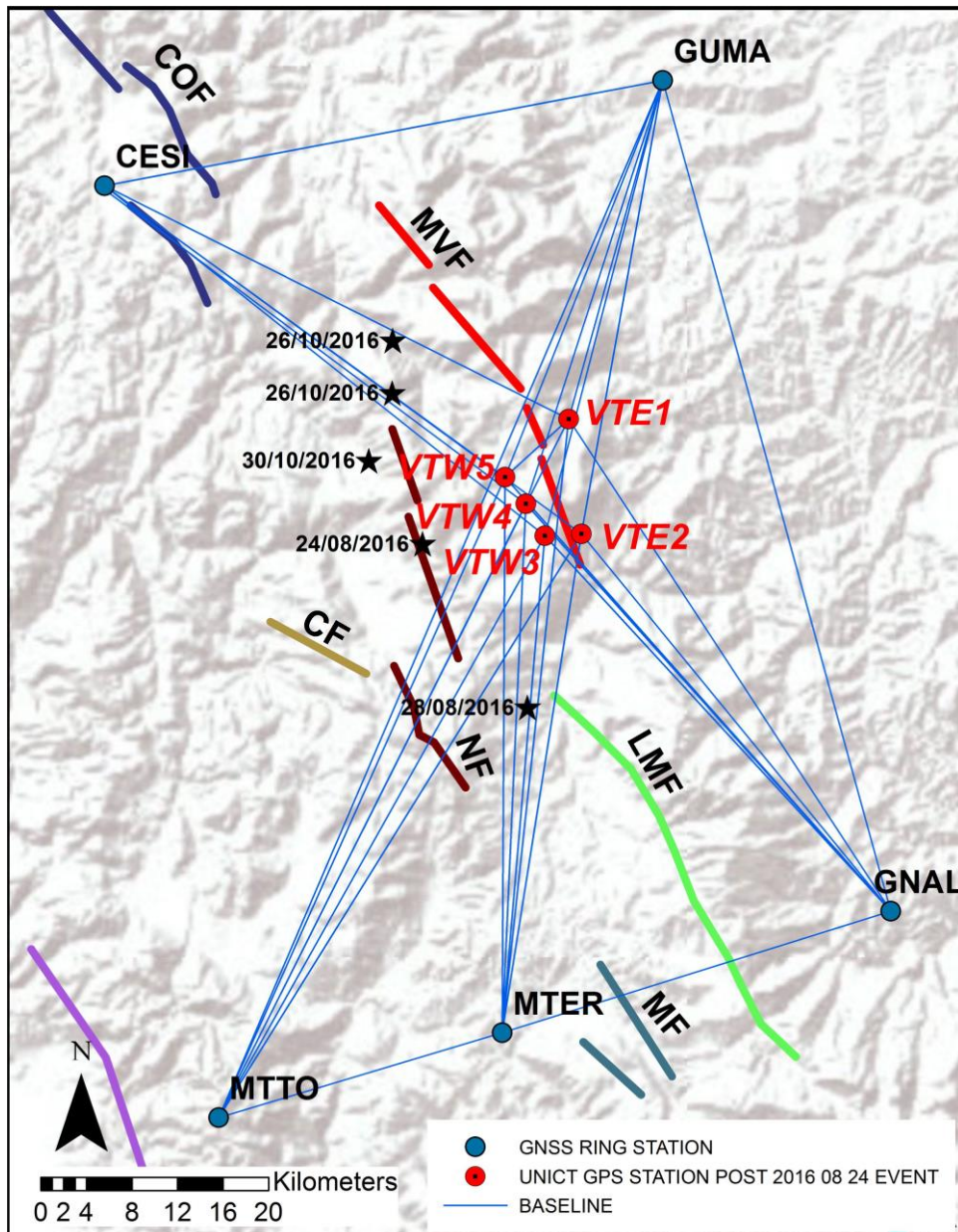


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Fig. 3 Synoptic picture showing installation of the new GNSS stations, measurement and processing phases.



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318 *Fig. 4 Baselines obtained by combining the new GPS UNICT stations with selected GNSS ones from the RING Network.*

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324 *Fig. 5 Color-coded maps showing the E-W (a) and vertical (b) displacement distribution obtained by the DInSAR*  
 325 *technique ([http://www.irea.cnr.it/index.php?option=com\\_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-](http://www.irea.cnr.it/index.php?option=com_k2&view=item&id=761:nuovi-risultati-sul-terremoto-del-30-ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1)*  
 326 *ottobre-2016-ottenuti-dai-radar-dei-satelliti-sentinel-1) recorded On October 26th 2016 (pre-event images) and on*  
 327 *November 1st 2016 (post-event images). The red and blu arrows represent the consensus pre-, co-, and post-seismic*  
 328 *displacements (with 95% confidence errors) on the basis of the GNSS UNICT network. Epicenters of major shocks are*  
 329 *from <http://ring.gm.ingv.it>.*

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