



1 The analysis of H/V curve from different ellipticity retrieval 2 technique for a single 3c-station recording.

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7 **Abstract.** In the last two decades or so the H/V (Horizontal-to-Vertical Spectral Ratio) technique remained very popular,
8 and are extensively used for the site fundamental frequency estimation. H/V curve are also used with dispersion curve to
9 jointly invert and retrieved the shear wave velocity of relatively deep soil deposit. Although a full theoretical explanation of
10 H/V technique is not been presented yet, there are two main assumptions used generally that H/V curves can be explained
11 by considering Rayleigh wave of noise wave field only while the other newly presented approach utilize the whole noise
12 wave field known as diffuse field approach (DFA). However in case of Rayleigh wave approach for H/V, it is almost
13 impossible to remove the fraction of Love wave horizontal component of H/V. Here in this study we aim to test
14 different approaches adopted for the removal of Love wave fraction from horizontal component for a borehole test site at
15 University of Sao Paulo. The result from different approaches are compared with borehole ellipticity curve. The result
16 shows that around the fundamental frequency of Love obtained in either way (DFA or ellipticity approach) is dominated by
17 Rayleigh waves.

19 1. Introduction.

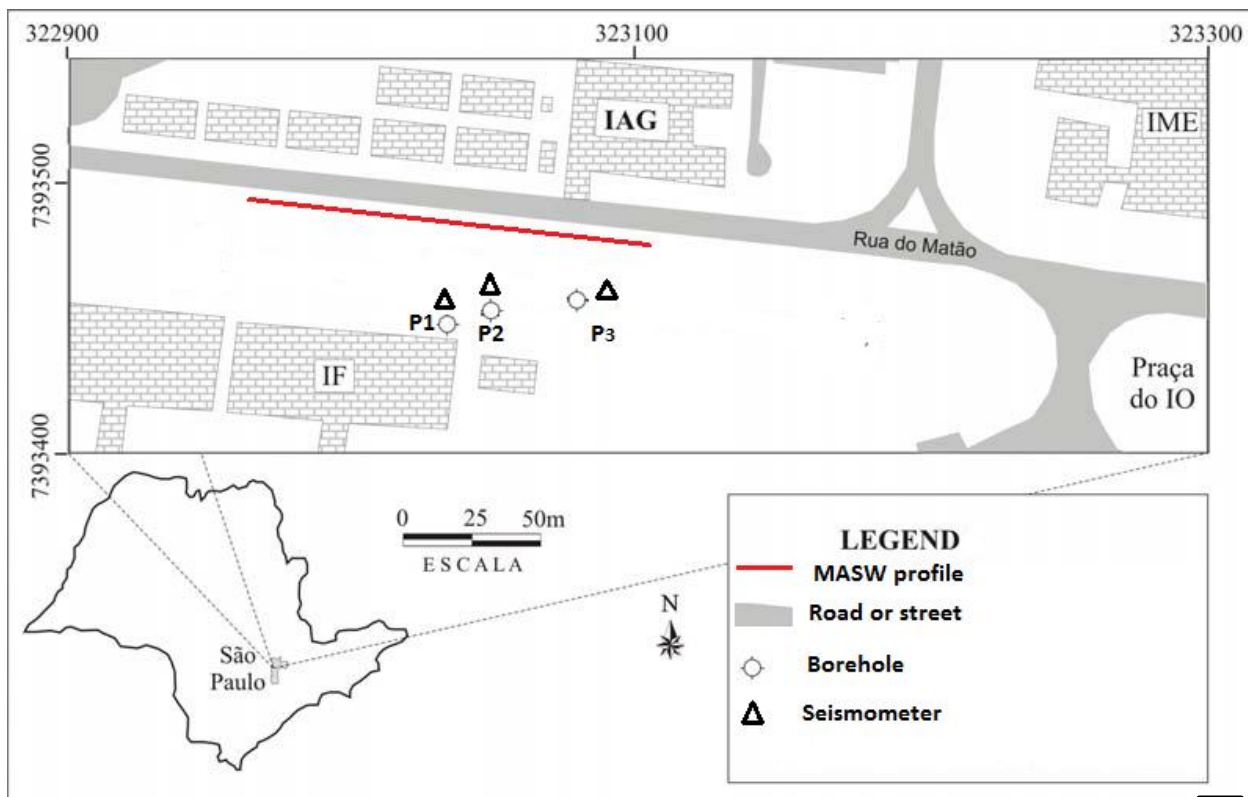
20
21 H/V (Horizontal-to-Vertical spectral ratio) is fast and quick way to get properties of a site for
22 engineering interest, by the measurement of ambient noise wave field with a single 3-component sensor
23 on the earth surface. The method is used for rapid estimation of fundamental resonance frequency (f_0)
24 of a site where the maximum displacement amplification is expected in case of an earthquake. This
25 technique is also used to retrieve the shear wave velocity of the geological structure in a joint inversion
26 with dispersion curve (Scherbaum et al, 2003 Picozi et al, 2005 Hobiger et al, 2013). However, some
27 controversies about the nature of ambient noise wave field and sources exist, which most of the time
28 make the results of H/V curve questionable and hence debatable. Apart from the controversies exist in
29 nature of ambient noise wave field M. Mucciarelli et al (2001) described some problem regarding the



30 acquisition and processing of H/V spectral ratio. Due to extensive utilization of H/V technique, a
31 commission was established to test the acquisition, processing and interpretation of this technique (site
32 effects assessment using ambient excitations) – the SESAME project (2001-2004), The guidelines
33 reports are published for acquisition, processing and interpretation of ambient noise wave field and have
34 addressed all the points raised and discussed by M. Mucciarelli et al (2001) in extensive details. The
35 H/V curve modeled with different approaches and each modeling approach might have the effect on
36 the retrieved soil profile (Sánchez-Sesma et al, 2011, Lunedei & Malischewsky 2015). The acquisition
37 of the data is made with a three component sensor placed on the surface of ground which record the
38 seismic noise wave field. Fourier spectra of the recorded seismic noise for all the three component (east-
39 west, north-south and vertical) are measured. The two horizontal component Fourier spectra are properly
40 averaged and then divided by the vertical Fourier spectra. This division of averaged horizontal and
41 vertical component result a curve (H/V) as function of frequency. H/V curve usually result a peak
42 depending on the subsoil stratigraphical profile, this peak correspond to fundamental resonance
43 frequency (f_0) of the site (Tokimatsu, 1997; Bard, 1999; Bonnefoy-Claudet et al., 2006). The averaged
44 horizontal component of H/V curve contains the contribution from both Rayleigh and Love wave and
45 some fraction of body waves as well. In a joint inversion of H/V curve with dispersion curve this other
46 elastic wave effect presence in the H/V curve might bias the retrieved s-wave velocity profile.
47 Therefore, the presence of Love and other elastic waves existence in the H/V curve must be assumed or
48 estimated before the inversion process (Bonnefoy-Claudet et al., 2006). Here in this communication we
49 will try to list the different approaches used for the refining H/V curve by removing unwanted
50 fraction (Love wave effect presence) prior to the joint inversion with dispersion curve. At present there
51 are two main research lines describing the H/V curve by taking in account the whole ambient-vibration
52 wave field, and another just studies the surface wave and Rayleigh wave dominance in noise wave field
53 (Lunedei & Malischewsky 2015). Sánchez-Sesma et al (2011) proposed that seismic noise field can be
54 consider as diffusion-like situation which contain all type of elastic wave (surface and body waves). He
55 suggested that the average autocorrelation of the motions at a given receiver, in the frequency domain,
56 measures average energy density (DEL) and is proportional to the imaginary part of the Green function
57 (GF) when both source and receiver are the same. The surface wave dominance opinion of noise wave



58 field is in favor of Rayleigh wave dominance (Yamamoto, 2000 Boremann 2002, Cornou, 2002 ,
59 Okada ,2003). We will try to check both these assumptions here with the borehole data.
60 The site for which this analysis are made is a borehole site at university of Sao Paulo shown in Fig.1.
61 The noise measurement were made with broadband 3 component seismometer nanometrics Trillium
62 Compact 120-s. Ambient noise wave field measurements were made for 24 hours on weekend night to
63 minimized cultural noise influence. The data acquisition of ambient noise has been done following the
64 guidelines developed under the SESAME (2004) recommendations. To obtain the fundamental
65 frequency of the site a window of one-hour recording were processed and the reliability condition
66 proposed by SESAME (2004) for the H/V curve and peak were followed.



67
68 Fig. 1 shows the location map studied area ,legends explain the different symbols. (modified from Porsani 2004)
69 We will try to discuss briefly here and analyzed our seismic noise data with this diffuse-field
70 assumption (DFA). Later we will analyzed our data with assumption that H/V curve basically reflect the
71 Rayleigh wave ellipticity.

72
73 **2. Diffuse field assumption technique.**



74
75 **Sánchez-Sesma et al (2011)** proposed to considered ambient noise wave field as diffuse wave field
76 which contain all different types of waves (surface and body). The ambient noise wave field is
77 generated by multiple random uncorrelated forces/sources near to or at the earth surface. The wave
78 field may contain the scattering effect of various elastic mode. The field intensities could be in a better
79 way described by diffuse like situation. To assume that the noise wave field is diffuse, the H/V curve
80 can be estimated for a receiver at earth surface in term of green tensor imaginary part at the source
81 (source and receiver are assumed to be at same location). The work of Sánchez-Sesma provides an idea
82 of linkage between energy density and imaginary part of GF in 3D (energy densities of the noise wave
83 field is proportional to the imaginary part of green tensor). The H/V curve obtained from the square
84 root ratio of imaginary parts of GF (horizontal and vertical components) Eq.2 serve as intrinsic
85 property of medium therefore its inversion can be used to retrieved subsurface soil profile. The detailed
86 analysis of the method is beyond the scope of this ~~communication~~ interested readers are referred to
87 Sánchez-Sesma et al (2011) for ~~detailed procedure~~. The summary of this procedure is that
88 autocorrelation of motion at a receiver sensor in a given direction is proportional to directional energy
89 density (DED),and this DED is proportional to the imaginary part of Green tensor at that sensor
90 (Sánchez-Sesma et al 2011). Patroal,(2009) showed that in case of 3D homogeneous elastic half
91 space, the horizontal displacement (radial and transverse) have fix energy proportion (e.g $E_1(x, x, \omega) =$
92 $E_2(x, x, \omega)$ and also $ImG_{11}(x, x, \omega) = ImG_{22}(x, x, \omega)$). For a diffused wave field the H/V can be
93 represented in term of directional energy densities assuming source and receiver lies at same location
94 (x) on the surface of earth as

95

$$96 \frac{H}{V}(\omega) = \sqrt{\left(\frac{E_1(x,x,\omega)+E_2(x,x,\omega)}{E_3(x,x,\omega)}\right)} \quad (1)$$

97

$$98 \frac{H}{V}(\omega) = \sqrt{\left(\frac{ImG_{11}(x,x,\omega)+ImG_{22}(x,x,\omega)}{ImG_{33}(x,x,\omega)}\right)} \quad (2)$$


99 where in (1)


$$E_m(x, \omega) = \rho\omega^2 \langle u_m(x, \omega) u_m^*(x, \omega) \rangle \quad \text{where } m = 1,2,3$$

100

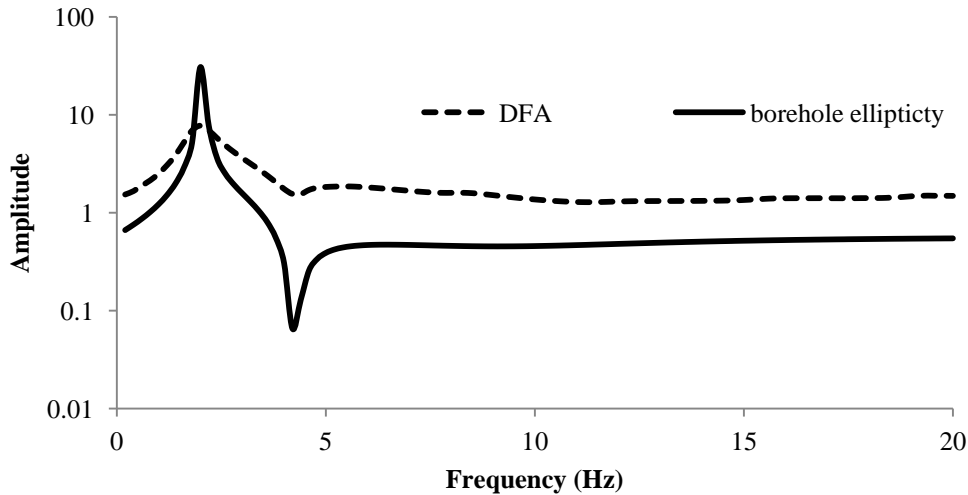
$$= -2\pi\mu E_s k_s^{-1} Im[G_{mm}(x, x, \omega)]$$



101
102 where energy density is find out at point x in direction m . ω , ρ and u_m are angular frequency , layer
103 density and displacement at point x respectively . $E_s = \rho\omega^2 s^2$ is the strength of diffuse illumination in
104 term of shear wave average energy density μ is shear wave modulus $\langle \dots \rangle$ bracket shows the azimuthal
105 average, $k_s = \frac{\omega}{V_s}$ shear wave number V_s shows medium S-wave velocity. The symbol $(.*)$ show complex
106 conjugate process, the medium response in a direction m (of impulse load and acting in same direction)
107 is indicated by G_{mm} . The H/V curve obtained in this manner at  linked to the intrinsic property of
108 medium , The resulted H/V curve from the diffuse-field approach might allow its inversion without
109 considering any supplemented information (dispersion curve).

110
111 For our analysis, the data of seismic noise recorded at borehole test site were analyzed with DFA
112 (diffuse field assumption) frame work of Eq.2 , $\frac{H}{V}(\omega)$ result  are obtained from the data with an
113 integration step of 1000 and window length 40s. The curve obtained by this directional energy density
114 approach is compared with borehole ellipticity Fig.2. The peak and trough of H/V curve obtained
115 (Fig.2) correspond to peak and trough of ellipticity. However the shape of H/V curve is generally higher
116 except the peak , ~~it is~~ because of other elastic wave phases contribution. This peak and trough
117 correlation of the H/V curve with borehole ellipticity shows that at these singularities (peak and trough)
118 Rayleigh wave contribution dominate the wave-field. It is important to note here the inversion of
119 ellipticity curve with dispersion curve are recommended around the peak region of H/V curve (Picozi
120 et al,2005 Hobiger et al, 2013). In next section we will focus our attention on the other line of research
121 which is in the opinion of, that noise wave field is dominated by surface wave especially Rayleigh
122 wave (see section 3) , and H/V curve can be explained by its correlation with Rayleigh wave ellipticity
123 curve (Bard 1999).

124



125

126

127 Fig.2 H/V curve obtained through DFA technique at the test site (IAG-USP), the Borehole ellipticity curve is plotted in solid
128 for comparison at the site.

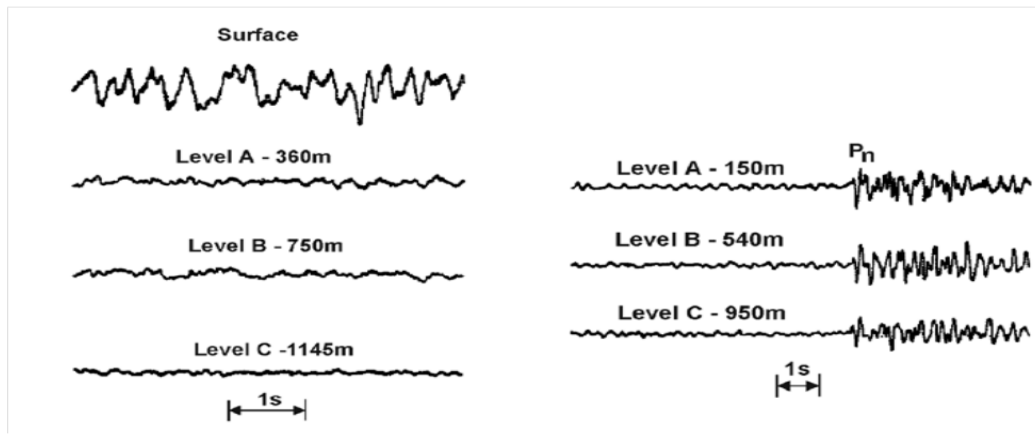
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130 3. Surface wave dominance of seismic noise wave field.

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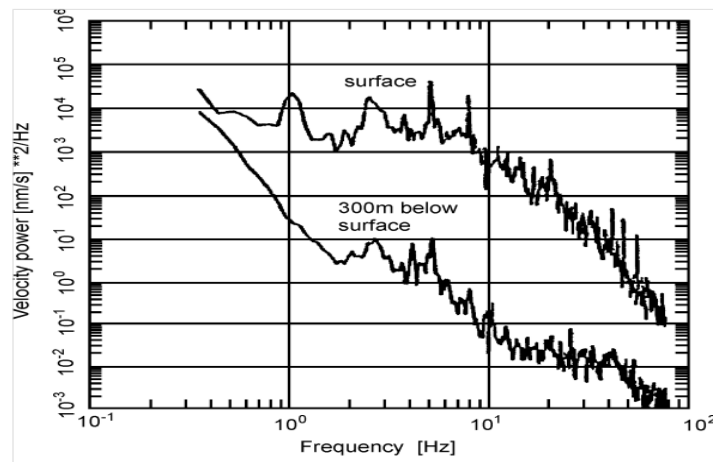
132 To find out that whether body or surface waves dominate the noise wave field is analyzed by Bormann
133 (2002), he used sensors for earthquake and seismic noise recording both at the surface and in the
134 boreholes at different depth levels and concluded the surface-wave nature of seismic noise Fig. 3.
135 Bormann (2002) showed that, the penetration depth of surface waves increases with wavelength, high
136 frequency noise attenuates more rapidly with depth. In case of Fig.3 the noise power at 300 m depth in a
137 borehole was reduced, as compared to the surface, by about 10 dB, at $f = 0.5$ Hz, 20 dB at 1 Hz and 35
138 dB at 10 Hz. This continuous amplitude decline with frequency is in accord with the surface waves
139 nature of seismic noise.

140



141

142 Fig.3 Recording of seismic noise (left) and earthquake signals (right) at the surface and at different depth levels of a
143 borehole. (Bormann 2002).



144

145 Fig. 4 Velocity power density spectra as obtained for noise records at the surface (top) and at 300 m depth in a borehole
146 (below) near Gorleben, Germany (Bormann , 2002).

147 If the noise wave field is mainly dominated by surface wave then another question arises that what is
148 the fraction contribution of Rayleigh and Love waves to the noise wave field. Most of the researchers
149 focused their attention to find the fraction Rayleigh to Love ratio from the analysis of noise wave field
150 recorded on vertical component (Li et al., 1984; Horike, 1985; Yamanaka et al., 1994). The results of
151 these studies showed an agreement in one aspect that microseism (<1Hz) are mainly dominated by
152 Rayleigh waves however at high frequency (> 1Hz) a combination of P and Rayleigh wave exists

153

154 Table 1 summarize the result of previous studies on this issue.



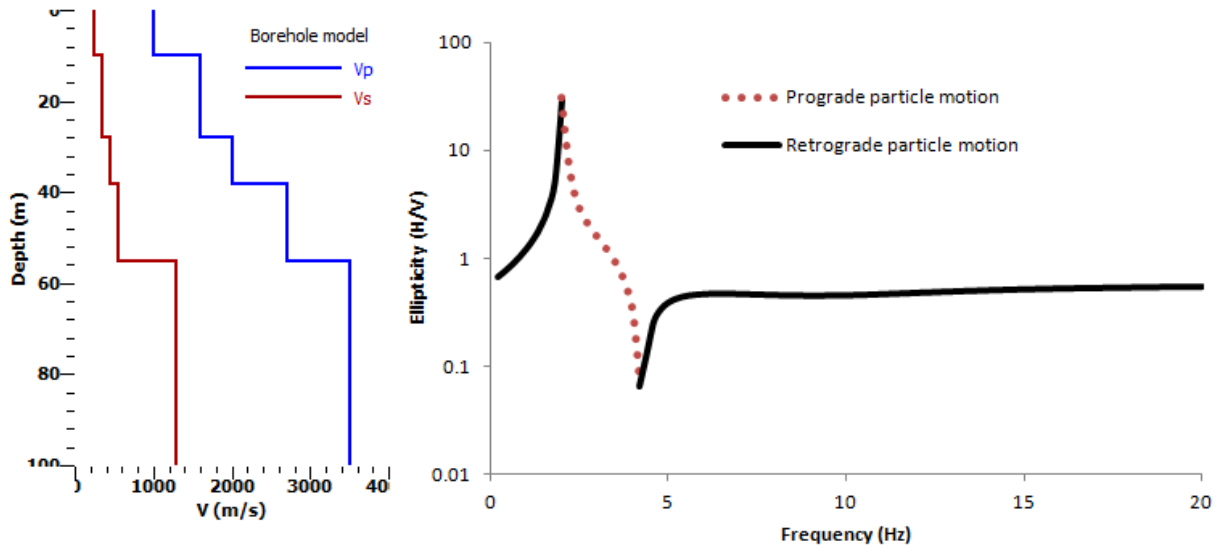
	Rayleigh waves(%)	Love waves(%)	frequency range(%)
Chouet et al.,1998	30%	70%	>2Hz
Yamamoto, 2000	<50%	>50%	3-10 Hz
Arai et al., 1998	30%	70%	1-12 Hz
Cornou, 2002	60%	40%	< 1 Hz
Okada (2003)	<50%	>=50%	0.4-1 HZ
Köhler(2006)	10–35%	65–90%	0.5–1.3 Hz

155

156 Table.1 Summary conclusions about the proportion of Rayleigh and Love waves in noise, after different authors (from
 157 Chouet et al., 1998; Yamamoto, 2000; Arai et al. 1998; Cornou 2002, Okada 2003, Köhler2006)

158 **4. Removal of Love wave from horizontal component.**

159 Rayleigh waves are formed by the linear pairing of P (primary waves) and Sv (vertically polarized
 160 shear waves) waves (Aki, 2002). This pairing of vertical and horizontal components have a phase shift
 161 of $\pm \frac{\pi}{2}$, the particle motion induced by Rayleigh waves will depict an ellipse, elliptical motion will
 162 either be retrograde or prograde depending on the sign of phase shift. Similarly Love waves composed
 163 of horizontally polarized shear waves (SH). The horizontal over vertical axes of ellipse described by
 164 particle motion under the Rayleigh wave influence is termed ellipticity. At situation of homogeneous
 165 half-space the particle motion is retrograde at all frequencies and ellipticity is constant . However in
 166 case of layered structure ellipticity exhibit a peak and trough and the particle motion switch from
 167 retrograde to prograde and then to retrograde with the frequency, depending on the velocity contrast
 168 between the soil and bedrock (Konno and Ohmachi, 1998).



169

170

171 Fig.5 Borehole model at IAG-USP . The ellipticity curve of fundamental Rayleigh wave of the borehole model , marking the
172 frequency ranges where the retrograde and prograde motion might occur.

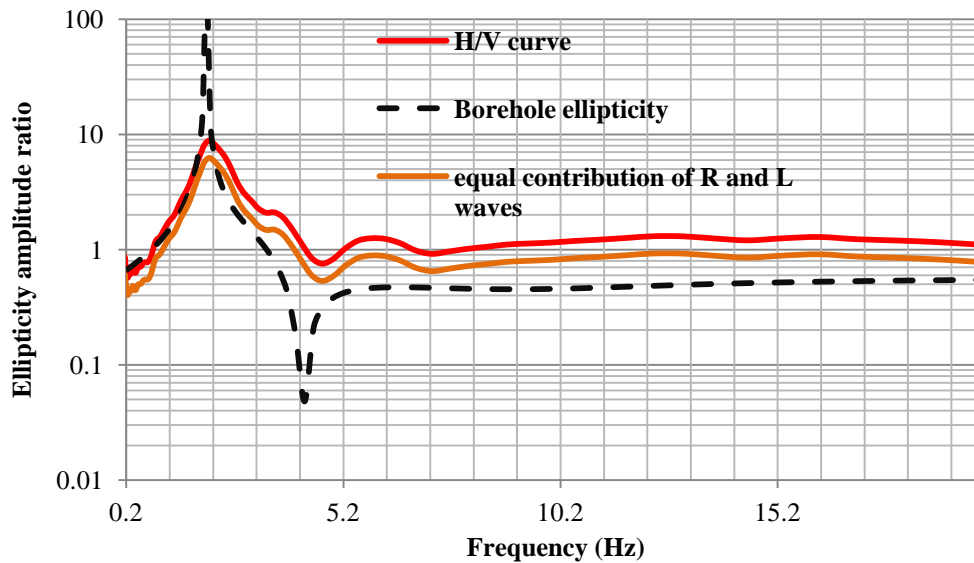
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174 The conclusion from the preceding section can be drawn that, the contribution of Love wave to the
175 horizontal is not predictable and fluctuate with frequency and from site to site. The **H/V** linked to
176 the ellipticity of Rayleigh wave , in situation where the high shear wave contrast exist between soil and
177 bedrock (Bard,1999).The H/V curve corresponds nicely to the peak of ellipticity curve (Fig.6). The
178 deviation between curves can be easily linked to the presence of Love wave contribution to noise wave
179 field at the horizontal component. We will try to review all the available technique for the this task and
180 compared its result with borehole ellipticity.

181 H/V curves are obtained and compared with the elliptiity curve of borehole at the same site. The
182 deviation between the curves are due to Love wave contribution. H/V curve is generally higher than the
183 ellipticity curve except at peak frequency of Rayleigh wave ellipticity Fig.6. Three different polarization
184 techniques are used to minimized the effect of Love wave. The first technique is the simple H/V of
185 seismic noise (Fig.6) , In this technique the polarization mean the division of Fourier spectral amplitude
186 of averaged horizontal component over vertical component with the assumption that around the
187 fundamental resonance frequency the vertical component is dominated by Rayleigh only. Generally it is



188 believed that Rayleigh (P-S) and Love wave (SL) contribute equally to the horizontal component. Fäh
189 et al (2001) proposed the division of H/V spectral amplitude by $\sqrt{2}$ for Love wave effect minimization
190 from horizontal component Fig.6. However this is not a wise approach because the energy partition of
191 horizontal component between Rayleigh and Love is not constant and varies with frequency and site
192 (Köhler *et al.* 2006; Endrun 2011) Table There are two other polarization approaches employed for
193 this job recently are: time-frequency analysis (Fah et,al 2009) and RayDec (Hobiger et, al, 2009)
194 followed by a concise introduction, the interested readers are referred for detail of these approaches
195 (Fah et,al 2009) and (Hobiger 2009).



196
197 Fig.6 shows the comparison of H/V curve of experimental data recorded at test borehole site at university of Sao Paulo.

198
199 **5. Time-frequency analysis.**

200
201 In time-frequency analysis the vertical component of noise wave field is considered as a trigger and its
202 energy level are estimated at given time and frequency and are correlated with horizontal components
203 (east-west, and north-south). A brief description of the technique is that continuous wavelet
204 transformation (CWT) are performed on the eastern $e(t)$, northern $n(t)$ and vertical $v(t)$ components of
205 noise wave field. CWT (continuous wavelet transform) transform a signal to time-scale plane. The
206 scale is a single parameter which controls both the duration and bandwidth. CWT is defined as.

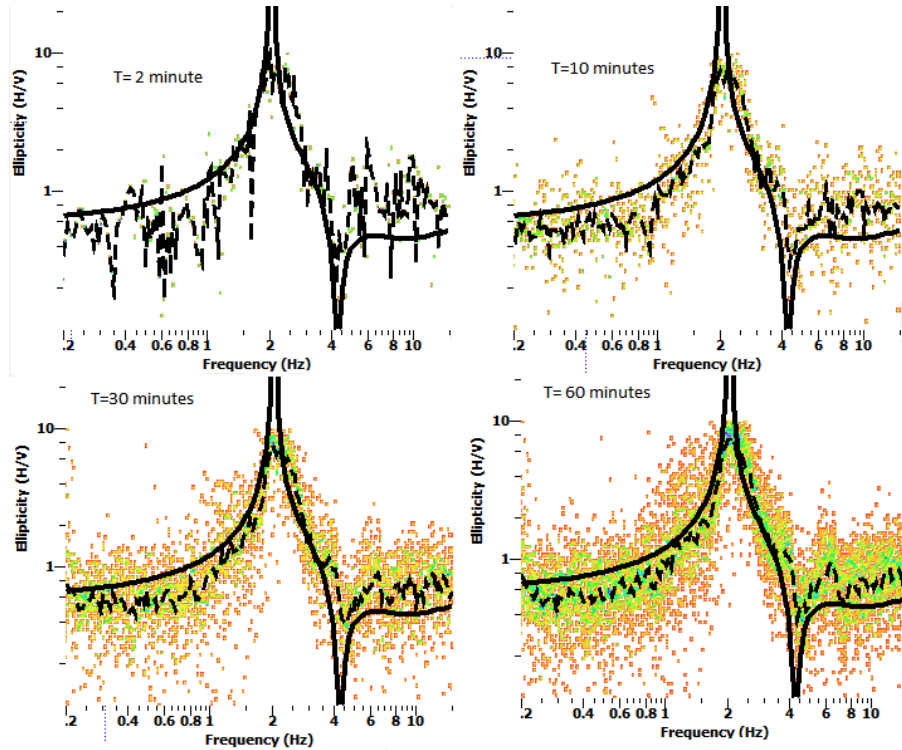


207
$$\text{CWT}_{x(t)}(a, b) = \frac{1}{\sqrt{a}} \int_{-\infty}^{\infty} x(t) \psi^* \left(\frac{a-t}{a} \right) dt \quad (3)$$

208 where $x(t)$ is real-valued signal component {where $x(t) = e(t)$ or $n(t)$ or $v(t)$ }, $\psi(t)$ shows the mother
209 wavelet, * shows a complex conjugation process, a scale parameter (scale control both the duration and
210 bandwidth) b is translation in time. Fah et,al (2009) used a modified Morelet wavelet in a code written
211 for this job due the reason that traditional Morelet wavelet does not act well for H/V analysis. The
212 modified Morelet transform is defined as

213
$$\psi(t) = \exp(i2\pi ft) e^{-\frac{t^2}{m}} \quad (4)$$

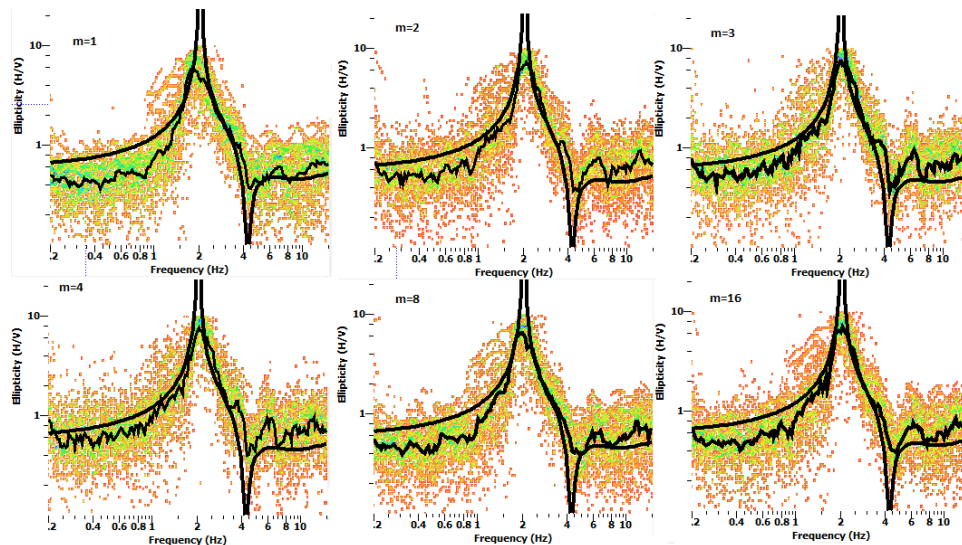
214 here f is the central frequency of the wavelet, m control the resolution of both frequency and time, low
215 m mean more time localization on the expense of frequency resolution and higher m result on the
216 contrary i.e increase frequency resolution at the expense of time resolution. The classical Morelet
217 wavelet is achieved when $m=1/2$. After the application of CWT to each component of a single 3-c
218 sensor recording { $e(t), n(t)$ or $v(t)$ } give rise to three signals amplitude which is both the function of
219 time and frequency. Two horizontal ($e(t), n(t)$) component are merged together to a single component.
220 As Rayleigh wave are the result of the coupling of horizontally and vertically polarized S-waves, the vertical
221 component is sorted out for amplitude maxima at each frequency and time translation . Similarly, the
222 horizontal component is analyzed for amplitude maximum at given frequency and time, the horizontal
223 components are phase shifted \pm at a quarter of the period. This shift of period is done because of
224 theoretical phase shift between $\pm \frac{\pi}{2}$ vertical and horizontal component particle motion. For each
225 maximum on the vertical component, the corresponding maximum on the horizontal component are
226 stored and the ratio of H/V are estimated. There is only one tuning parameter m ,the effect of it
227 different values of m are shown in Fig.8 , Also the effect on the length of the recorded signal is shown
228 in Fig.7. This whole process is statistically analyzed by histogram for all the frequency and translation
229 of time and H/V is obtained by the maximum of each histogram. For details theoretical outline of this
230 analysis please read (Fah et,al, 2009).



231

232 Fig.7 Ellipticity (H/V) obtained from continuous wavelet transformation CWT for different length of the recorded signal; a
233 histogram is drawn for each frequency, the color within histogram indicates the energy level. (dashed line shows ellipticity
234 obtained from CWT while solid line shows the ellipticity curve obtained from borehole data at same location - IAG-front)

235



236



237

238 Fig.8 Ellipticity (H/V) obtained from continuous wavelet transformation CWT for different value of m; (curvy line shows
239 ellipticity obtained from CWT while solid line shows the ellipticity curve obtained from borehole data at same location -
240 IAG-front)

241

242 Fig.7 and 8 show a better result when compared with the borehole ellipticity curve ~~of borehole~~. The
243 result of wavelet transform Fig.7&8 shows that the ellipticity of borehole is retrieved in better way(
244 especially at right limb of the ellipticity) when the recorded length of the signal is greater than 30
245 minute ,and when the value of $m \geq 8$.

246

247 **6 Random decrement technique (RAYDEC).**

248

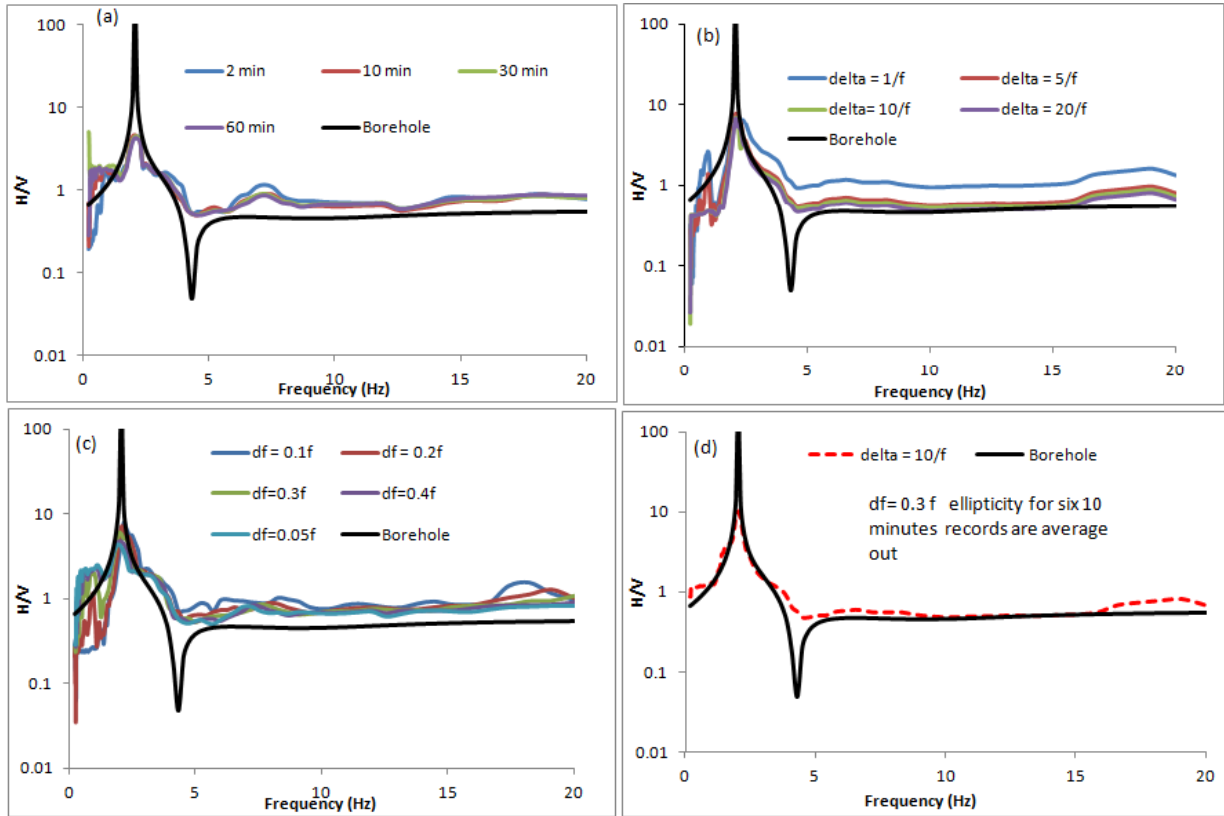
249 Another polarization technique use for the effect of Love ~~(S_H)~~ wave effect minimization for a single 3
250 component sensor recording is term as RayDec technique (Hobiger et.al, 2009) The vertical component
251 is taken as master trigger ~~as~~ because the ~~S_H~~ arrival occur only on the vertical component, while stacking
252 a large number of horizontal component to obtained ellipticity of Rayleigh curve (Hobiger et al., 2009)
253 showed that the obtained ellipticity will be closer to the true ellipticity curve rather than the H/V curve.
254 To elaborate RayDec techniques consider a signal {where $x(t) = e(t)$ or $n(t)$ or $v(t)$ }. These three
255 time series has N number of data points ~~and~~ having length T. To obtain a Rayleigh wave ellipticity
256 curve, the main idea of this method is to obtain only Rayleigh waves in comparison ~~with~~ other waves type
257 by the addition of a large number of filtered signal windows Δ , estimates the energy level of horizontal
258 averaged and vertical signal in a location where the vertical component change its ~~sign~~ from -ve to +ve.
259 Due to the phase shift of $\frac{\pi}{2}$ between Rayleigh wave vertical and horizontal component, both the
260 horizontal component of EW and NS are projected by factor ϕ with north direction in such a way to
261 maximized the correlation between the summed horizontal and vertical component. The Rayleigh wave
262 ellipticity is obtained latterly from Eq.5 . The ellipticity E is calculated as the square root of the ratio of
263 the energies in the signal window. Δ .

264
$$E = \sqrt{\frac{\int_0^{\Delta} hf^2 \cdot s(t) \cdot dt}{\int_0^{\Delta} vf^2 \cdot s(t) \cdot dt}} \quad (5)$$



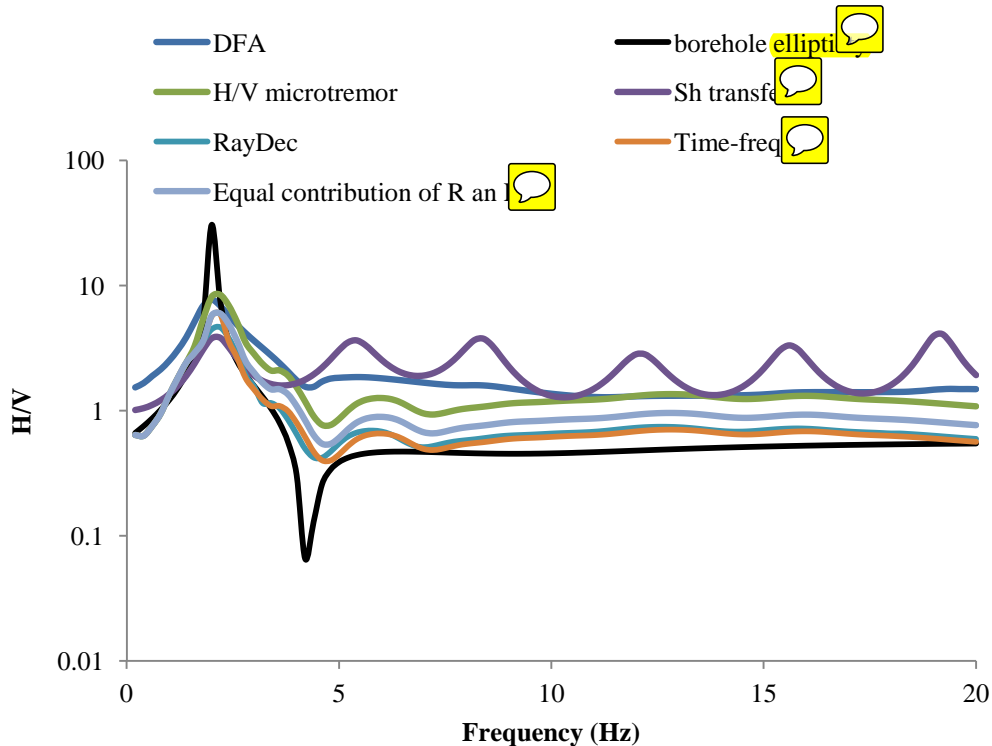
265 where $h_{f.s}(t)$ is the horizontal average (NS north-south , EW east-west) signal and $V_{f.s}(t)$ is the
266 vertical component. This process is repeated for a whole record for an increment of window length Δ
267 .The window length is taken as function of frequency such that to ensure 10 significant cycle at chosen
268 frequency ($\Delta=10/f$ where f is the analyzing frequency from 0.2 to 20 Hz are used). There are two tuning
269 parameter for this technique , Δ (window length) and df (width of frequency filter). The effect of these
270 two tuning param^{er} on the result is shown Fig.9. For our analysis we used one-hour records of noise,
271 which were divided into 6 segments of 10 minutes each , afterward, the result is average out for all the
272 six segments with the aim ^{of} minimization of misfit (see for detail procedure Hobiger et al., 2009). The
273 result of ellipticity obtained via RayDec show a better match with borehole curve especially at right and
274 left limb.
275 The analysis of both time-frequency and RayDec for H/V give satisfactory result for the ellipticity
276 retrieval of Rayleigh waves. Fig 10 shows the comparison of all the technique retrieved curves with that
277 of borehole ellipticity.

278



279
 280

281 Fig.9 Ellipticity obtained via **RADH** technique , (a) shows the effect signal duration on the result.(b) shows the effect of
 282 using different width of window (c) shows the effect of different filter width on the result (d) shows six 10-minute windows
 283 (ellipticity for each 10 minute is obtained at the end the result is average and compared with borehole data black line).





284 Fig.10 Show the comparison of borehole ellipticity curve combine with that of all ellipticity retrieval technique.
285


286 7. Discussion and conclusion.

287 Deep soil shear wave velocity information can be retrieved from the joint inversion of H/V curve with
288 dispersion curve. The H/V curve at assumed to be linked with Rayleigh wave ellipticity for a site.
289 However in real situation the H/V curve technique can not completely replicate the Rayleigh wave
290 ellipticity. It is because of the presence of different elastic wave influence presence in the H/V curve
291 retrieval. Diffuse field approach is a better way to consider the effect of all elastic seismic phases (body
292 and surface wave). However the analysis of our data shows that at singularities (peak and trough)
293 especially at peak of the H/V curve obtained via DFA is very well correlated with Rayleigh ellipticity
294 curve in term of these singularities frequency. This matching can be assumed that around the peak
295 frequency even in case of diffuse like field situation the seismic noise field is dominated by surface wave
296 especially Rayleigh waves.

297 The other approach used for the H/V curve technique is linked to the surface wave dominance
298 especially Rayleigh wave around the singularities. Three different polarization technique for retrieval of



299 ellipticity curve by minimizing the  contribution to the horizontal component  are simple H/V with
300 equal contribution of Rayleigh and Love wave, time-frequency analysis and RayDec technique. The
301 result of equal assumption of Rayleigh/Love wave fraction is unable to match with parts of borehole
302 ellipticity except at peak. However ellipticity curve retrieved from noise analysis through time-
303 frequency and RayDec shows a good replication of borehole ellipticity around the peak, right and left
304 limb.

305 For the deep soil deposit the joint inversion of ellipticity and dispersion curve of Rayleigh wave are
306 recommended. However due to the presence of effect of other seismic elastic phases especially Love
307 wave may produced some bias result. Therefore it is extremely necessary to retrieved a H/V curve
308 where the effect of Love wave contribution are minimized prior to the inversion with dispersion curve.
309 For our test site we found that time-frequency and RayDec show better result ,by replication left and
310 right limb of ellipticity curve of the borehole at the  e.

311 **Acknowledgement.**

312 We are thankful to the Sanchez-Sesma (Instituto de Ingeniería Universidad Nacional Autónoma de
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319

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