

1 Meteorological factors driving glacial till variation and the associated
2 ~~periglacial-periglacial~~ debris flows in Tianmo Valley,
3 southeast Tibetan Plateau

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8 Abstract: Meteorological studies have indicated that high ~~a~~Alpine ~~environments~~ are strongly
9 affected by climate warming, ~~and~~ ~~p~~Periglacial debris flows are frequent in deglaciated regions.
10 The combination of rainfall and air temperature controls the initiation of periglacial debris flows,
11 and the addition of melt-water due to higher air temperatures enhances the complexity of the
12 triggering mechanism compared to ~~that of~~ storm-induced debris flows. ~~In~~ ~~On the~~ south-eastern
13 Tibetan Plateau, where temperate glaciers are widely distributed, numerous periglacial debris
14 flows have occurred ~~in over~~ the past 100 years, but none ~~had happened~~ ~~occurred~~ in the Tianmo
15 watershed until 2007. In 2007 and 2010, three large-scale debris flows occurred in the
16 Tianmo ~~watershed~~ ~~Valley~~. In this ~~research~~ ~~study~~, these three debris flow events were chosen to
17 ~~analyze~~ ~~analyse~~ the impacts of the annual meteorological conditions, ~~including~~ the antecedent air
18 temperature and meteorological triggers. TM images and field measurements of the nearby glacier
19 suggested that ~~a~~ sharp glacier retreats ~~had existed~~ ~~occurred~~ in the ~~previous one~~ ~~or to~~ two years
20 preceding the events, which coincided with ~~the~~ ~~spikes~~ ~~in~~ ~~the~~ mean annual air temperature.
21 ~~Besides~~, ~~Moreover~~, ~~changing of~~ glacial till ~~changes~~ driven by ~~a~~ prolonged increase in the air
22 temperature ~~is are the a~~ prerequisite of periglacial debris flows. ~~Different factors can trigger The~~
23 ~~Triggers of~~ periglacial debris flows ~~are multiplied often coupled~~, and they ~~could be may include~~
24 high intensity rainfall, ~~as in the first debris flows and the third debris flows~~, or continuous,
25 long-term ~~rising increases in~~ air temperature, ~~s~~ as in the second debris flow ~~events~~.

26 **Key words:** glacial till variation; meteorological factors; periglacial debris flows;
27 southeast Tibetan Plateau

28 1. Introduction

29 ~~The Alpine~~alpine environments are ~~strongly~~ vulnerable to climate changes, ~~of~~
30 ~~which the~~and alpine glaciers and permafrost are the most sensitive ~~in the form of~~
31 ~~glacier and permafrost to~~ degradation (Harris et al., 2009; IPCC, 2013). Glacier and

批注 [A1]: Answer to reviewer #2: the three debris flows is of large magnitude while debris flows with small magnitude occurred followed each of the three debris flows; while debris flow with small magnitude is hard to be determined compared with the larger ones.

批注 [A2]: Answer to reviewer #2: Line 18 spikedm?? What is it?

批注 [A3]: Answer to reviewer #2: Triggers of periglacial debris flows are multiplied....What does it mean?

批注 [A4]: Answer to reviewer #2: as in the first and third debris flow is better than as in the first debris flows and third debris flows

带格式的: 缩进: 首行缩进: 5 字符

32 permafrost retreat can induce mass movements, such as landslides, shallow slides,
33 debris slides, moraine collapses, etc. (Cruden and Hu, 1993; ~~Korup and Clague~~ Korup,
34 2009; McColl, 2012; Stoffel and Huggel, 2012; Fischer et al., 2012). These
35 movements, that will be expelled out of the watershed would bring the material out of the
36 expelled material from watersheds in the form of debris flows or sediment fluxes. The Debris
37 flows in alpine ~~areas~~ regions can often bury residential areas, cut off main roads, and
38 block rivers (Shang et al., 2003; Cheng et al., 2005; Deng et al., 2013) and destroy
39 basic facilities ~~located in~~ downstream; ~~thus, they posing a great~~ pose a considerable
40 threat to the local economy and social development. In undeveloped alpine areas
41 where the transportation system is particularly poor or limited, such as in the
42 south-eastern Tibet, where the transportation system is particularly poor or limited, the
43 negative effects produced by debris flows, such as cutting off main roads, such as
44 cutting off main roads are can be serious (Cheng et al., 2005).

批注 [A5]: Answer to reviewer #2:
that will be expelled out... what does it mean?

45 Periglacial debris flows occurs ~~in the~~ high alpine areas where there is with large
46 areas of glaciers, such as on the Tibetan Plateau in China (Shang et al., 2003; Ge et al.,
47 2014), in the Alps in Europe (Sattler et al., 2011; Stoffel and Huggel, 2012), in the
48 Caucasus Mountains in Russia (Evans et al., 2009) and in northern Canada
49 (Lewkowicz and Harris, 2005). Periglacial debris flows ~~were reported to be~~ can be
50 initiated by rainfall (Stoffel et al., 2011; Schneuwly-Bollsweiler and Stoffel, 2012),
51 glacial melt-water flow ~~of glacier~~ or ice particle ablation (Arenson and Springman,
52 2005; Decaulne et al., 2005), or outburst floods from glacier lakes (Chiarle et al., 2007)
53 in different parts of the world; however, while the multiple triggers of a single event
54 have rarely been studied for the case is rarely to be read. Because debris flows are
55 commonly triggered by rainfall (Sassa and Wang, 2005; Decaulne et al., 2007; Kean
56 et al., 2013; Takahashi, 2014), the rainfall threshold, intensity and duration ~~has~~ have
57 been widely used for debris flow monitoring and giving to provide event warnings in
58 non-glacier areas (Guzzetti et al., 2008).

批注 [A6]: Answer to reviewer #2:
occurs without s

批注 [A7]: Answer to reviewer #2:
for the case is rarely to be read. What does it mean?

59 In deglaciated ~~edion~~ areas, the debris flow threshold can be more difficult to
60 determine. Periglacial debris flows tend to occur in the summer when the thawing of
61 glaciers and glacial tills predominates and melt-water penetrates ~~into~~ the glacial tills at

62 a constant and successive flow rate. The effect of melt-water appears is similar to that
63 of antecedent rainfall (Rahardjo et al., 2008) and is variable in different periods,
64 considering snow and glacier shrinkage and air temperature fluctuations. In the Swiss
65 Alps, the melt-water volume is high in early summer,r and as-debris flows can be
66 initiated by low total-intensity rainstormrainfall. However, whereas higher totallarger
67 rainstorms are required to produce debris flows in late summer or-and early autumn
68 when the melt-water volume is low (Stoffel et al., 2011; Schneuwly-Bollscheweiler and
69 Stoffel, 2012). ~~In~~On the south-eastern Tibetan Plateau, the rainfall threshold given by
70 Chen et al., (2011) is quite-relatively wide (0.2~2.0 mm/10min, 0.6~6.3 mm/h or
71 3.0~19.4 mm/24h), ~~and~~ the small rainfall threshold of which in-particular is likely to
72 contain the effect of affected by the air temperature. Moreover, periglacial debris flows
73 induced by a-sudden releases of water from glacier lakes have a close relationship
74 withare closely relatedthe rising to increasing air temperature (Liu et al., 2014).

批注 [A8]: Answer to reviewer #2: the small rainfall threshold which is the reference of this?

75 Fluctuation of Air temperature is-fluctuations are likely to be quite important in
76 triggeringtriggers of periglacial debris flows. Compared with-to the storm-induced
77 debris flows, the addition-ofincreased air temperature can greatly enhance the
78 complexity of the initiation of periglacial debris flows. It is of high-difficulty to
79 simulate the triggering process by-via experiments or mathematical simulation;
80 thus, -and instead,case studies of natural debris flows eases in the natural environment
81 could be appliedmust be explored. In this researchstudy, three debris flow events,after
82 a debris flow free period of nearly 100 year, in the Tianmo watershed of-on the
83 southeastern of the-Tibetan Plateau as-deglaciation continued are used as examples
84 after a debris flow-free period of nearly 100 years as deglaciation continues. -and
85 theThe annual meteorological conditions, antecedent air temperature and triggering
86 conditions prior to debris flows are analyzed-analysed to further understand the
87 meteorological triggers and their roles in glacier retreat, glacial till variation and
88 debris flow initiation.

批注 [A9]: Answer to reviewer #2: Lines 69-71 Confused text

89 2. Background

90 (1) Study area

91 ~~The T~~temperate glaciers ~~in-on~~ the Tibetan Plateau ~~is-are~~ primarily distributed in
92 the Parlung_Zangbo Basin, and ~~they~~ covered a total landmass of 2381.47 km² in 2010
93 based on TM images ~~(taken by the No. 4 or 5 thematic mapperson the Landsatsatellite~~
94 ~~with a spatial resolution of 30 m)~~ (Liu, 2013). Historically, the movement of
95 temperate glaciers has produced ~~a large amount of numerous~~ moraines, the depth of
96 which can reach ~~up to~~ 500 m locally (Yuan et al., 2012). In recent decades, ~~there has~~
97 ~~been~~ a dynamic significant ~~increase in~~ temperature ~~increase has occurred~~, and
98 ~~according to statistics~~ the temperature at the Bomi meteorological station (~~midstream~~
99 ~~central in the~~ Parlung_Zangbo Basin) ~~has rose~~ ~~increased~~ by 0.23°C/10a from 1969 to
100 2007, resulting in remarkable ~~glacial~~ shrinkage ~~of the glacier~~ (Yang et al., 2010).

101 Tianmo Valley, ~~which is~~ located in Bomi County and to the south of the
102 ParlungZangbo River, covers an area of 17.76 km² (29°59'N/95°19'E; Figure 1). This
103 valley has a northeast-southeast orientation and is surrounded by high mountains
104 reaching 5590 m a.s.l. at the southernmost ~~site-location~~ and 2460 m a.s.l. at the
105 ~~junction-offork in~~ the ParlungZangbo River. The TM image ~~in-from~~ 2013 ~~showed~~
106 ~~illustrated~~ the presence of a hanging glacier with an area of 1.42 km² in the upper
107 concave area at an ~~altitude-elevation~~ of 4246 m to 4934 m. Bared rock, dipping at an
108 angle of ~~around 6~~ ~~approximately~~ 60°, ~~emerged~~ below and above the hanging glacier
109 and ~~is~~ often covered by ~~everlasting~~ snow. Below 3800m ~~ma~~.s.l., vegetation,
110 including forest and shrubs, occupies most of the area (Table 1).

111 The river channel in the watershed is sheltered by shade and not directly affected
112 by sunlight, resulting in less solar radiation and a location at which a small trough
113 glacier can form. In the main channel, the trough glacier extended to 2966 m a.s.l. in
114 2006. The lower part of the trough glacier has been eroded by glacier melt-water flow,
115 and an arch glacier that is vulnerable to high pressure was formed (Figure 2). The
116 remnants of the landslide deposits ~~are~~ approximately 10-~~meter~~ ~~metres~~ high ~~can be~~

117 observed on both sides of the channel. These deposits, ~~which~~ consist of low-stability
118 sediment and can be easily entrained by debris flows, ~~can be observed in both sides of~~
119 ~~the channel.~~

120 Tianmo Valley is located on the north side of the bend in the YarlungZangbo
121 River and is strongly affected by ~~the~~ new tectonic movement. An inferred normal fault
122 vertical to the channel cuts through the valley and is only 30 km ~~away~~ from the
123 Yarlung_Zangbo fault. In 1950, a rather significant earthquake (Ms. 8.6) hit Zayu,
124 which is only 200 km away, and local records reported that a large amount of rock
125 collapsed and landslides were produced at that time. The whole valley is located in a
126 strong ductile deformation zone and is dominated by gneissic lithology belonging to a
127 Presinian System.

128 (2) Disaster history

129 According to our field interviews with local residents, there were no debris flows
130 in the approximately 100 years prior to 2007 in Tianmo Valley. The channel was ~~quite~~
131 relatively narrow before 2007, and the local people could walk across via a wooden
132 bridge to live and farm on the terrace on the west side. On the morning of September-
133 4th, 2007, after ~~the a~~ rainfall which even that did not hit the downstream area ceased,
134 the local forest guard heard a loud noise coming from the upstream area at
135 approximately 18:00, with Rrainfall which later began in the upstream area at
136 approximately 19:00. Then, following this rainfall was debris flows which rushed out
137 of began to occur in the Tianmo Channel, and they subsequently blocked the Parlung
138 Zangbo River. This debris flow event is listed as DF1 in this paper. It is told by the
139 local citizen that several debris flows occurred during that entire night and
140 lasted, lasting the entire night while we cannot separate them. According to the field
141 measurements, approximately 1,340,000 m³ of sediment was transported during this
142 event, resulting in 8 missing persons and deaths. ~~Concurrently within this same~~
143 ~~time~~, debris flows occurred in the four nearby 4 valleys (Table 2). According to the
144 size classification proposed by Jakob (2005), which is based on the total volume, peak
145 discharge and inundated area, the sSize classes of the debris flows in the five valleys

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批注 [A10]: Answer to reviewer #2:
Line 125-128 Sentence a bit confused.

批注 [A11]: Answer to reviewer #2:
DF2 stand for the debris flows occurred
on 19:00, September 4, 2007. The other
debris flows took place in the nearby 4
valleys as in Table 2.

146 ~~is-are~~ given in Table 2. ~~The Tianmo Valley is debris flow events is listed as DF1 in~~
147 ~~this paper.~~

148 At 11:30 on July ~~25th~~, 2010, debris flows were again triggered in Tianmo
149 Valley that traced the path of the preceding debris flow deposits and reached the other
150 side of the ParlungZangbo River. According to Ge et al., (2014), ~~a solid mass~~
151 sediment ~~mass~~ of approximately 500,000 m³ was carried ~~to out~~ (Table 1) and
152 deposited ~~on the cone in the channel and blocked to block~~ the main river. A barrier lake
153 was formed, and the rising water destroyed the roadbed of G318. ~~The following week~~
154 ~~also experienced~~ ~~D~~dozens of ~~debris flows in~~ small-magnitude ~~debris flows occurred in~~
155 ~~the following week~~. This debris flow ~~events~~ is listed as DF2 in this paper.

156 Debris flows occurred again two months later on Sep. 6th (The Ministry of Land
157 and Resources P. R. C., 2010), although we could not determine the exact times
158 sequence of ~~the events~~, ~~but a~~ according to speculation, ~~these the debris flows~~ could
159 have occurred in the early morning before dawn ~~and~~ when the rainfall intensity ~~has~~
160 reached its maximum (Figure 9). ~~This theory, which~~ agrees with the findings of Chen
161 (1991), ~~who found~~ that periglacial debris flows ~~have~~ historically occurred between
162 18:00~24:00 in this area. The debris barrier in the main river was consequently
163 increased by an additional 450,000 m³, and the barrier lake was enlarged to ~~hold~~
164 ~~maintain~~ 9,000,000 m³ of water. This debris flow ~~events~~ is listed as DF3 in this paper.

165 A field investigation revealed that a high percentage of boulders in the
166 downstream area and glacial tills above the trough glacier ~~were are quite~~ loose and ~~of~~
167 high porosity ~~rocks~~ (Figure 2); hence, they have low density and can be easily
168 entrained. Our particle size tests ~~of~~ the glacial tills and debris flow deposits indicate
169 a ~~lower~~ clay (d<0.005 mm) content, whereas the debris flow deposits contain more
170 fine particles that are smaller than 10 mm (Figure 4), suggesting that ~~the~~ entrainment
171 ~~supplied-accounted for~~ a considerable amount of fine particles.

172 (3) Meteorological data

173 The study area is located in a high alpine area where the economy is ~~quite~~
174 ~~relatively~~ undeveloped ~~with and only~~ few meteorological stations ~~exist~~. Before 2011,

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批注 [A12]: Answer to reviewer #2:
These debris flows or this debris flow?

175 the Bomi meteorological station (~~since established in~~1955) was the only station in the
176 area. ~~It is,~~ located 54 km ~~away~~ from Tianmo Valley at an ~~altitude elevation~~ of 2730 m,
177 and other stations were located more than 200 km away.

178 The Tibetan Plateau is a massive terrace that obstructs the Indian monsoon,
179 causing it to travel through the Yarlung Zangbo Canyon and its tributaries. As the
180 Indian monsoon is transported to higher ~~altitudes altitudes~~, a rainfall gradient emerges
181 in the Parlung Zangbo Basin. However, according to ~~our statistics on the~~ rainfall data
182 ~~in from~~ the area, ~~the~~ rainfall often ~~enjoys exhibits~~ ~~the a~~ similar intensity ~~for as that of~~
183 the long-term rainfall process from Guxiang to Songzong, which ~~means suggests that~~
184 ~~the there is no a~~ large rainfall gradient ~~does not exist~~ between Tianmo Valley and
185 Bomi meteorological station; therefore, the rainfall data from ~~the~~ Bomi
186 meteorological station ~~can could~~ be used ~~for in~~ our study. ~~In order to To~~ conduct
187 ~~further additional studies study~~, another meteorological station was built in 2011 near
188 Tianmo Valley.

189 It has been established that the air temperature decreases with altitude; ~~therefore;~~
190 ~~therefore,~~ the air temperature in the source area of Tianmo Valley is lower than that in
191 Bomi County. According to the research by Li and Xie (2006), the air temperature
192 decreases at a rate of 0.46~0.69°C/100m over the ~~whole entire~~ Tibetan Plateau,
193 and the rate in the study area is 0.54°C/100 m. Because the glacier and permafrost in
194 the source area have ~~a~~ planar distributions, the air temperature at the geometric centre
195 of the glacier and permafrost can be used to ~~analyze analyse~~ the temperature process.

196 3. Analysis and results

197 (1) ~~Changing of~~ Air temperature and rainfall ~~changes~~

198 The mean annual air temperature is ~~usually generally~~ used to reflect the ~~tendency~~
199 ~~trends~~ of glacier change (Yang et al., 2015). We collected ~~the~~ mean annual air
200 temperature and annual rainfall data from 1970 to 2014 from the Bomi meteorological
201 station (Figure 5). The records showed that the mean air temperature has increased by
202 approximately 1.5 °C ~~in over~~ the ~~last past~~ 45 years ~~at a rate of ,accounting for~~

批注 [A13]: Answer to reviewer #1:
not really the right verb !

批注 [A14]: Answer to reviewer #2:
please substitute can with could

203 0.033 °C/a. This air temperature increase was particularly ~~more~~ rapid between 2005
204 ~~and ~2007, an at~~ approximately 0.7 °C/3a, which is 7 times the average value ~~of over~~
205 the ~~last past~~ 45 years. ~~On the other hand~~ However, the annual rainfall from 2000 to
206 2010 was low and ~~it was~~ estimated ~~at as~~ 828.2 mm per year. From 2000 to 2004, the
207 rainfall during summer (July to September) accounted for approximately 50% of the
208 total annual rainfall; however, only 32% of the rainfall occurred in the summer of
209 2005~2006, even though the annual rainfall exhibited ~~the a same similar~~ trend. In
210 2007, rainfall in the summer and the entire year returned to the mean rainfall state.

211 ~~According to~~ Figure 5 ~~shows a~~ similar ~~trend in the~~ air temperature and rainfall
212 ~~was observed trends~~ before DF2 and DF3. The air temperature ~~elevated increased~~ in
213 2009 ~~to reach and reached~~ the maximum ~~temperature over of~~ the ~~last past~~ 45 years
214 ~~period of, accounting for~~ 10.2 °C; however, the annual rainfall, ~~was~~ only 65% of the
215 average amount; and the summer rainfall, ~~which was,~~ lower than that in 2005 and
216 2006, reached ~~their a~~ minimum values. In 2010, ~~the~~ rainfall was abundant, and the
217 annual rainfall increased to 1080.6 mm, which is approximately 30% more than the
218 average value and close to the maximum.

219 The following common traits can be identified ~~from by~~ comparing the annual
220 meteorological conditions of DF1, DF2 and DF3: 1) ~~o~~ One or two years before the
221 debris flows, the mean annual temperature ~~elevated increased~~ and the annual rainfall
222 and summer rainfall ~~declined decreased~~. ~~The Additionally, the~~ climate was in a
223 "hot-dry" state; 2) As the temperature gradually decreased, the annual rainfall
224 returned to normal or increased, and ~~the a~~ "hot-wet" climate ~~state~~ contributed to debris
225 flow initiation (Lu and Li, 1989).

226 (2) Changing of glacier in Tianmo Valley

227 In our ~~research study,~~ remote images ~~is were~~ collected to ~~analyze analyse~~ the
228 ~~changing of~~ glacier ~~changes~~ in the source area ~~during the past in recent~~ years. ~~In order~~
229 ~~to~~ To eliminate the effect of snow cover, images were taken in the thawing seasons
230 when ~~the~~ snow cover is limited, ~~enabling to enable an~~ easy detection of ~~the~~ glaciers
231 ~~from and~~ snow. ~~Besides, Moreover,~~ an image taken on a bright, cloudless day is still

232 needed to show the watershed clearly; ~~however;~~ however, a difficult case ~~ensues is~~
233 encountered when the rainy season ~~comes in between~~ begins during the thawing
234 season, ~~as when~~ the atmosphere is often covered by thick clouds. Further more, ~~in order~~
235 ~~to show illustrate~~ glacier retreat and its impact on debris flows properly, the images
236 should be within similar time intervals, ~~like~~ such as 3 years, before and after debris
237 flow events. ~~As the H~~ high-resolution images are rare, ~~to obtain~~ and we could only
238 collect one SPOT image (~~ta~~ Taken ~~by~~ the ~~satellite~~ of ~~Systeme~~
239 ~~Probatoired~~ Observation de la Terre ~~satellite~~ with a ~~space spatial~~ resolution of 55m m)
240 in 2008. To achieve ~~consistency of the~~ image consistency, we collected 5 TM images
241 ~~image~~ (~~t~~ Taken ~~by~~ the No. 4 or 5 thematic mapper ~~carried~~ on the ~~satellite~~ Landsat
242 satellite with a ~~space spatial~~ resolution of 300m m); taken on ~~September~~ 17th, 2000,
243 ~~July~~ 24th, 2003, ~~September~~ 21st, 2006, ~~September~~ 24th, 2009 and ~~August~~ 4th,
244 2013, ~~respectively~~.

245 Based on the 5 TM images, we classified the area as glacier, snow, ~~bared~~ barred land,
246 gully ~~deposition~~ and vegetation in time series (Figure 6), and the area of each is given
247 in Table 1. Figure 6 show ~~sed~~ that deglaciation ~~was taking place~~ occurred in Tianmo
248 V ~~valley; and in particular~~ notably, the eastern branch ~~had has~~ experienced ~~the~~
249 ~~sharpest considerable~~ deglaciation. ~~In order to~~ To show clearly show the rapid rate of
250 glacial ~~er~~ retreat in the entire basin and eastern branch, a graph of retreat was plotted,
251 ~~to show the changing of glacier and the eastern branch~~ as shown in Figure 7.

252 Figure 7 shows that the glacier in Tianmo V ~~valley~~ had shrank from ~~been in~~
253 ~~shrinkage since~~ 2000 to 2013, with variable rates of ~~tion in~~ glacier retreat ~~rate~~. In
254 2000~2003, 2003~2006, 2006~2009 and 2009~2013, the glacier retreat rates in
255 Tianmo V ~~valley~~ corresponds to ~~were~~ 0.02, 0.06, 0.027 and 0.007 55km km²/a,
256 respectively, and those of the eastern branch were 0.0033, 0.01, 0.008 and 0.002
257 km²/a, respectively for the eastern branch. According to ~~these the~~ figures, the largest
258 glacier retreat rate was observed in 2003~2006, followed by that in 2006~2009. ~~It is~~
259 ~~important that~~ The glacier area at the beginning should be ~~taken into~~
260 ~~consideration~~ noted to ~~judge assess~~ the ~~changing~~ rate of change of the glacier. The
261 glacier retreat rate is can be normalized, and the relative glacier retreat rate can be

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批注 [A15]: Answer to reviewer #2:
Confused sentence

262 calculated based on ~~the~~this area-change in areaing.

263 The relative glacier retreat rates~~are—were~~ 11.30, 35.09, 17.43 and 5.17
264 $10^{-3}\text{km}^2/\text{a}/\text{km}^2$ ~~during—in~~ 2000~2003, 2003~2006, 2006~2009 and 2009~2013,
265 respectively; ~~whereas, it is~~and the corresponding values were 20.83, 66.67, 66.67 and
266 20.83 $10^{-3}\text{km}^2/\text{a}/\text{km}^2$ for the eastern branch. ~~These figures show that the~~The relative
267 glacier retreat rate ~~for—of~~ the eastern branch ~~had shrunk much more—decreased~~ sharply
268 between 2000 ~2013.

269 In this ~~research~~study, TM images ~~with—over~~ 3-year intervals were applied ~~can~~
270 ~~only get to obtain~~ the mean glacier retreat rate. As glacier retreat rate in the 3 three
271 years could be either high or low, field measurements of the nearby glacier ~~is—were~~
272 used to show the glacier retreat condition before debris flows occurred. Yang et al.
273 (2015) ~~had~~ conducted field measurements of the No.94 Glacier in the Parlung_Zangbo
274 Basin since 2006, and the field measurements suggests it ~~was—had in—a~~ negative
275 balance ~~in—from~~ 2006~2010 (Figure 7). The negative balance reached a maximum
276 level ~~the—maximal~~ in 2009, followed by 2008 and 2006, indicating ~~sharp—rapid~~
277 deglaciation in these three years.

278 When we combined the results of the TM image analysis and field—field
279 measurements of the No. 94 Glacier, we observed ~~that it is right before debris flows~~
280 that the glacier in Tianmo Valley experienced the ~~sharpest—most rapid~~ deglaciation
281 prior to debris flows in 2006, 2008 and 2009, which ~~was also coincidental~~coincided
282 with ~~the elevated~~an increase in the mean annual air temperature (Figure 5). ~~Besides,~~
283 Moreover, the maximum glacier retreat in 2009 ~~could be also—was potentially~~ related
284 to the decline ~~of—in~~ snowfall in the preceding winter and early spring.:

285 (3) Antecedent air temperature and rainfall ~~process~~

286 The air temperature in the source area can be obtained ~~using—based on the—a~~
287 vertical rate of vertical decline rate ($0.54\text{ }^{\circ}\text{C}/100\text{ m}$). ~~According—Based on to~~ this
288 method, the air temperature in the source area was $9.8\text{ }^{\circ}\text{C}$ lower than that at the Bomi
289 meteorological station. We collected ~~the—daily temperature; that is~~ the lowest
290 temperature, ~~the the~~ mean temperature and daily rainfall from June to September ~~in in~~

批注 [A16]: Answer to reviewer #2:
field instead of file

批注 [A17]: Answer to reviewer #2:
confused sentence

291 2007 and 2010 (Figure 8).

292 ~~According to~~ Figure 8 ~~shows that~~, the lowest air temperature was below 0 at the
293 end of June, 2007. At the beginning of July, the air temperature started to rise quickly,
294 which continued until early September when DF1 occurred. ~~this~~ ~~This demonstrates~~
295 ~~trend suggests~~ that the high air temperatures in July and August contributed to DF1.

296 ~~According to Figure 8~~ ~~Additionally~~, the air temperature was high from early July
297 to late August, and another high air temperature period ~~emerged~~ ~~was observed~~ in early
298 September. When DF2 occurred in late July, the air temperature had reached the
299 maximum for that year, which suggests that the air temperature in early and
300 ~~middle~~ July was responsible for DF2. After DF2 occurred, the air temperature in
301 August ~~began to prepare for~~ ~~varied towards the conditions that caused~~ DF3.

302 Antecedent air temperature fluctuations includes the air temperature and ~~its~~ ~~the~~
303 duration ~~of variations~~. The air temperatures and durations before debris flows are
304 variable ~~and~~ ~~making them~~ difficult to evaluate. The accumulation of positive air
305 temperature is ~~usually often applied~~ ~~used~~ to ~~analyze~~ ~~analyse~~ the ~~impact effect~~ of air
306 temperature on glacier melting (Rango and Martinec, 1995) ~~and~~ ~~which~~ can be
307 expressed as ~~follows~~:

308
$$T_{PT} = \sum_{i=-n}^0 T_i (T_i > 0) \quad (1)$$

309 ~~Where~~ ~~the~~ T_{PT} is the positive air temperature accumulation, ($^{\circ}\text{C}$) and T_i is the
310 average daily air temperature; (only $T_i > 0$ is included).

311 Because air temperature is successive, it is difficult to determine the beginning of
312 positive air temperature accumulation. Glacial tills can ~~lessen~~ ~~decrease~~ the heat that
313 penetrates into them, and the low air temperature ~~is can~~ only ~~contribute to~~ ~~observed in~~
314 the upper thin layer. ~~Moreover~~, freeze-thaw cycles exist when the lowest air
315 temperature is less than 0°C . From this ~~point of view~~ ~~perspective~~, the beginning of
316 positive air temperature accumulation is defined as the time at which the lowest air
317 temperature exceeds 0°C for two or three successive days or ~~since~~ the last debris flow.

318 Based on the above method, we can deduce that ~~the~~ positive air temperature

319 accumulation began when the lowest air temperature exceeded 0°C for several
320 successive days ~~beginning, starting~~ on June 28th, 2007, ~~and~~ June 9th, 2010, ~~and July~~
321 ~~26th, 2010, which correspond~~~~corresponding~~ to DF1, ~~and~~ DF2, ~~and~~ DF3, respectively,
322 ~~and on July 26th, 2010 for DF3, following DF2.~~ The duration and T_{PT} were calculated
323 for each debris flow event, ~~the~~ ~~The~~ results ~~was~~ ~~were~~ 69 days and 517.9 °C, 47 days
324 and 332.1 °C ~~and~~, 42 days and 320.4 °C (Figure 8) for DF1, DF2, and DF3,
325 respectively. The results showed that T_{PT} for DF1 ~~is~~ ~~was~~ much larger than the other
326 two T_{PT} values, ~~and~~ which ~~is~~ ~~coincidence~~ ~~coincides~~ with ~~to~~ the fact that there was no
327 debris flows in the past dozens of years, and extraordinary external forces such as
328 large T_{PT} ~~is~~ ~~are~~ required to ~~destroyed~~ ~~disrupt~~ the long-term balance.

329 (4) Triggering conditions

330 The continuous nature of ~~the~~ air temperature limits the possibility for debris
331 flows triggered by a sole abrupt increase in air temperature. ~~and since~~ ~~Since~~ ~~the~~
332 previous air temperature trends cannot be neglected, it is of no sense to study air
333 temperature triggers.

334 Antecedent rainfall is a factor that favours debris flows. In our analysis, the
335 rainfall over the three days preceding a debris flow event is given in Figure 9.

336 Before DF1, the air temperature was high, ~~and~~ ~~which~~ continued through July
337 and August. ~~The~~ ~~Notably,~~ ~~the~~ T_{PT} reached 517.9°C. According to the local forest
338 guard, an isolated convective storm occurred prior to DF1, ~~al~~ though no rainfall was
339 recorded at the Bomi meteorological station or in the downstream area at that time. In
340 Figure 9, as the rainfall right before DF1 occurred was not recorded by ~~the~~ Bomi
341 metrological station, we added ~~to the rainfall intensity (about 5~~ ~~approximately 5~~ mm/h
342 ~~of rainfall intensity~~ (according to the description ~~of~~ ~~provided by~~ the forest guard)
343 before DF1 to account for the storm, which might ~~nt~~ ~~not~~ reflect the real rainfall
344 process. We can therefore conclude that this isolated convective storm initiated DF1,
345 while the long-term high air temperature trend ~~had~~ paved the ~~road~~ ~~way~~ for DF1.

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域代码已更改

批注 [A18]: Answer to reviewer #1:
"to" is not needed, right ?

346 Considering a large deglaciated area, several other periglacial debris flows
347 | simultaneously ~~also~~ occurred near Tianmo Valley (Deng et al., 2013), which suggests
348 the advantageous meteorological conditions for debris flow initiation.

349 | DF2 ~~took place~~occurred when the air temperature reached ~~the a~~ peak in 2010.
350 The thaw season began in the middle of June, and ~~the~~ T_{PT} reached 332.1 °C. On July
351 24th, one day before DF2, the air temperature reached ~~the a~~ maximum value for that
352 year. The rainfall record at the Bomi meteorological station shows that there had been
353 no rainfall several days preceding DF2, and the local citizens also ~~did not observe~~
354 any observed no rain ~~either~~. The trigger of DF2 was likely the continuous percolation
355 of melt-water due to the long-term rising increase in air temperature.

356 | According to field interviews, several debris flows of small magnitude ~~had also~~
357 occurred before DF3. The air temperature decreased in late August but increased to
358 another high peak value before DF3, and the T_{PT} reached 320.4 °C. Rainfall began 2
359 days prior to DF3 and was steady the entire day before DF3. According to the rainfall
360 trend at the Bomi meteorological station, the rapid increase in rainfall intensity started
361 4 hours before DF3 and reached 3.8 mm/h, which was responsible for the initiation of
362 DF3.

363 4. Discussion

364 | Debris flows initiation is the process when of a water source provokes provoking
365 the movement of soil sediments. In this research study, we found that the triggering
366 factors of the three debris flows were high air temperature and rainfall for DF1, high
367 air temperature for DF2 and storm for DF3, respectively. ~~three debris flows were~~
368 ~~triggered by high air temperature and rainfall in during DF1, high air temperature in~~
369 ~~during DF2, and rainfall in during DF3 respectively.~~ When we ~~analyzed~~ analysed the
370 ~~dates~~ and ~~the~~ triggers for of these events, various questions ~~came to mind~~ should be
371 settled first that gave reasons to doubts: 1) wWhy did debris flows ~~did~~ not occur in
372 2006 or 2009 when deglaciation reached its peak and more ice meltwater was present;
373 2) wWhy did DF1 and DF3 occurred in September when the air temperature and

批注 [A19]: Answers to reviewer #1:
this definition is a bit rough. It could be
skipped.

批注 [A20]: Answers to reviewer #2: it
should be the triggering factors of the
three debris flows, were.....

批注 [A21]: Answers to reviewer #1:
too gergal

374 ~~volume of the~~ ice meltwater ~~was were~~ decreasing; and 3) ~~w~~Why ~~was were~~ there ~~is~~ no
375 large-scale debris flows triggered by ~~the~~ previous ~~heavier heavy~~ storms. ~~It makes~~
376 ~~us~~Based on our results, we believe that the impact of ~~the~~ water source on the
377 magnitude and frequency of debris flows is ~~quite relatively small~~low, or ~~there could~~
378 ~~be much~~more debris flows ~~would form during the early larger storm;~~and instead, ~~the~~
379 soil source, including ~~its the associated~~ magnitude and activity, ~~should may~~ be the
380 predominant~~te~~ controller, ~~just~~as ~~reported by~~ Jakob et al., (2005), ~~who pointed~~
381 ~~out~~noted that ~~the recharge of channel recharge is a~~should be the prerequisite for debris
382 flows. However, in most situations, we cannot reach the source area to detect the soil
383 source, and ~~the~~high-tech remote sensing can ~~just only~~ distinguish the boundary of ~~the~~
384 soil source. In the periglacial area where ~~the~~glacial till is often covered by glacier or
385 everlasting snow, ~~changing of a change in the~~ soil source ~~seems to be of~~would be
386 highly difficult~~ty~~ to detect. In this ~~research~~study, we ~~try to~~combine the
387 meteorological conditions and ~~the~~literature ~~reports~~ to discuss the ~~probable likely~~
388 variations~~of in~~ glacial tills before debris flows.

389 (1) ~~Annual v~~Variations~~of in~~ glacial till ~~in annual years~~

390 Climate warming is a global trend (IPCC, 2013), and the Tibetan Plateau, as the
391 third pole, is no exception ~~to climate change~~. According to our statistics, the air
392 temperature in Bomi County has increased by 1.5 °C ~~in~~over the ~~last past~~ 45 years
393 (1970~2014). Glacier retreat induced by climate warming has been widely accepted,
394 and recent research suggests ~~that~~ the weaker Indian monsoon could be another reason
395 ~~for such retreat~~ (Yao et al., ~~—~~2012). Glaciers are always located in concave ground
396 ~~areas~~ and cover a large ~~amount~~ volumes of glacial tills. Glacial pressure can generate
397 normal stress vertical to the slope, which can strengthen the slope stability. The effect
398 of glaciers on slope stability is called glacial debuitressing (Cossart et al., 2008). As
399 deglaciation continues, the result could lead to ~~the~~ exposure of the frozen glacial tills
400 (Figure 10, ~~A~~ A to B) and smaller glacial debuitressing.

401 The retreats of glaciers and glacial tills ~~with due to~~ climate warming ~~is are~~ quite
402 different. Deglaciation is accompanied by ~~the~~ melting of internal ice particles,

批注 [A22]: Answers to reviewer #2:
Confused sentence

403 ~~which. The melt of internal ice particles~~ can produce an active surface layer ~~which that~~
 404 can obstruct heat fluxes from penetrating into the deep layer ~~and~~, result ~~into in~~ the
 405 melting of internal ice particles ~~lagging at a rate slower than that of behind~~ glacial
 406 retreat (Takeuchi et al., 2000). ~~As~~ Because a strong heat gradient ~~is~~
 407 ~~existed exists occurs~~ at the surface ~~while but is quite~~ limited in deep layers, glacial tills
 408 with thicker coverage always ~~has have~~ relatively thinner thawed ~~ing~~ layers, and the
 409 ablation rates of glaciers and internal ice particles ~~can enjoy the same are similar value~~
 410 ~~to that~~ at the glacier surface ~~and~~ close to the moraine slope. ~~The newly formed bared~~
 411 glacial till is frozen with a high ice content. ~~The~~ cohesion of the ice particles ~~renders~~
 412 ~~creates the a~~ bared glacial till with high shearing strength and stability, and only the
 413 surface layer is ~~of high highly active activity~~. ~~Therefore Thus, we often see many bare~~
 414 ~~moraine slopes near glaciers. Therefore, for this reason there were~~ no debris flows of
 415 large magnitude ~~were observed~~ in 2006 and 2009 when glacier retreat reached ~~the a~~
 416 ~~maximal maximum~~ and the active glacial till is quite limited.

417 **(2) Variation of in glacial till ~~in on~~ antecedent days**

418 After the long ~~term~~ cold winter, ~~the whole~~ glacial tills ~~would~~ become frozen. If
 419 ~~the a~~ regressive glacier ~~was does~~ not recovered in the winter, ~~the~~ glacial tills ~~would~~
 420 ~~often be are~~ covered by snow. As ~~the~~ air temperature increases again, the surface snow
 421 ~~would~~ melts first, followed by the internal ice particles. The thawing of internal ice
 422 particles ~~would~~ induces a series of changes in the glacial till, which include the
 423 following: 1) the thawing will break the bonds of ice particles and increase the
 424 instability between ice cracks (Ryzhkin and Petrenko, 1997; Davies et al., 2001); 2)
 425 the sharp air temperature fluctuations in high alpine, mountainous areas induces a
 426 repeated cycle of expansion and contraction in the glacial till that can destroy the
 427 mass structure to some extent; 3) the seepage of ice melt-water can ~~deliver transport~~
 428 fine-grained sediments that were formerly frozen in the ice matrix (Rist, 2007); and 4)
 429 the ice melt-water can result in a higher water content and pore water pressure
 430 (Christian et al., 2012). These changes in glacial till can sharply ~~decline decrease~~ the
 431 soil strength, shifting to an active mass from ~~the an~~ uncovered and frozen moraine

批注 [A23]: Answer to reviewer #1:
occurs ?

批注 [A24]: Answer to reviewer #1:
similarly to above, not really the correct
verb

批注 [A25]: Answers to reviewer #2:
Confused sentence

批注 [A26]: Answers to reviewer
#2: ...there were no debris flows of large
magnitude... explain why—the newly
formed glacial till is of high shearing
strength and low activity, leading to the
low possibility of shallow landslide
occurring.

432 (Figure 10-B to C). Because ~~the~~ heat conduction in glacial till is ~~quite~~ relatively slow,
433 this process may last for a very long time and ~~also~~ requires a high antecedent air
434 temperature.

435 Heat conduction via the percolation of rainfall and ice melt-water can amplify
436 the depth of ~~active~~ active-of glacial till (Gruber and Haeberli, 2007), whereas ~~the~~
437 ~~shelter-of~~ covering the surface glacial till can hinder ~~the-a~~ heat flux from penetrating
438 into the deep layers (Noetzli et al., 2007). At a low air temperature, the heat flux
439 should be constrained to the surface layer, and a large heat gradient due to a high air
440 temperature would contribute much more to the heat flux and ice melt in the deep
441 mass. ~~Thus, meaning that~~ the long-term effect of a high air temperature can amplify
442 the active glacial till (Noetzli et al., 2007; Åkerman and Johansson, 2008; Åkerman et
443 al., 2008), under which lies frozen glacial till with a high ice content. The activity of
444 glacial till ~~variations-varies~~ with depth ~~from~~ high ~~in-at~~ the surface ~~and-to~~ low in the
445 deep layers, and landslide failure can take place on glacial till slopes in a retrogressive
446 manner, coinciding with long-term air temperature fluctuations, ~~although-as~~ active the
447 glacial till is ~~significantly~~ largely ~~unlimited~~ relatively limited in deglaciation
448 deglaciated areas.

449 (3) Failure of glacial tills

450 ~~Failure of Gglacial could be diversity failure can vary~~ Different factors can lead
451 to glacial till failure. Active glacial till slopes with low strength are usually vulnerable,
452 and their failure can occur when the air temperature is above 0 -0°C (Arenson and
453 Springman, 2005). ~~Either R-rainfall or ice melt water induced by air temperature can~~ ;
454 ~~the seepage flow of a glacier or ice particle melt-water could percolate the through tills~~
455 ~~and~~ trigger the failure (Figure 10-C to D). This ~~kind-type of event~~ is called ~~the-a~~
456 shallow landslide ~~type~~, and the failure mechanism lies in the ablation of internal ice
457 particles and the percolation of melt-water, which can initially decrease the soil
458 strength ~~at first~~ (Arenson and Springman, 2005; Decaulne et al., 2005). ~~Later~~, the
459 subsequent rapid percolation of ice melt-water or heavy rainfall can saturate the till
460 debris, decrease soil suction and shearing strength, and form seepage flows that

批注 [A27]: Answers to reviewer #2:
retrogressive manner means failure will
take place at the surface active layer,
followed by the deeper layer with
increase of its activity

批注 [A28]: Answers to reviewer #1:
what do you mean ? "different factors
can lead to glacial till failure" ?

批注 [A29]: Answers to reviewer #2:
Authors should distinguish triggering
factors from the triggering mechanisms.

461 ~~can matrix. The glacial till decrease soil suction and shearing strength of the glacial till~~
462 ~~decrease and then~~ trigger the shallow landslide failure (Springman et al., 2003;
463 Decaulne and Sæmundsson, 2007; Chiarle et al., 2007). Whether the failure can
464 induce debris flows is ~~still still~~ dependent on ~~the its~~ ability ~~that it can to~~ entrain the
465 debris layer, ~~in which case the otherwise, it can~~ debris is deposited ~~and~~ ~~as the flow~~
466 ~~moves through~~ ~~large~~ the channel.

467 ~~Another kind type~~ of failure ~~can~~ ~~might~~ take place ~~when a peaked runoff flows~~
468 ~~over and entrains debris deposits in the charged channel and reach a critical discharge~~
469 ~~(Berti and Simoni, 2005; Gregoretti and Dalla Fontana, 2008; Kean et al., 2013;~~
470 ~~Takahashi, 2014), which is more determined by channel bed slope and grain size of~~
471 ~~debris (Tognacca et al., 2000; Gregoretti, 2000; Armanini and Gregoretti, 2005; Kean~~
472 ~~et al., 2013).~~ ~~by~~ This ~~kind type~~ of ~~water stream~~ ~~channelized runoff~~ could be ~~the a~~
473 combination of ~~the~~ three factors: ~~including~~ rainfall, melting ice or the overflow ~~that~~
474 ~~forms~~ when ~~the a~~ glacier collapses ~~falling down into the downwards~~ ~~downward~~ into a
475 water pool. ~~Mechanism of this process lies in the hydrodynamic forces, created by the~~
476 ~~channelized runoff, (Kean et al., 2013; Gregoretti et al., 2016) and pose~~
477 ~~create hydrodynamic forces acting that acting~~ on the surface elements of the debris
478 layer ~~and surpassing~~ resistance of the sediment (Tognacca et al., 2000; Gregoretti,
479 2000; Armanini and Gregoretti, 2005; Prancevic et al., 2014). The concentration of
480 runoff in the channel bottom causes ~~the~~ erosion of the debris surface layer ~~forming a~~
481 ~~solid-liquid current at first, and then extends to the layers below with whole or partial~~
482 ~~mobilization and debris flows was generated (Gregoretti and Dalla Fontana, 2008).~~ ~~;~~
483 ~~resulting in with whole or partial mobilization of the bed material. The inclusion of~~
484 ~~bed material in the flowing water~~ ~~water stream~~ generates a debris flow ~~(Gregoretti and~~
485 ~~Fontana, 2008).~~

486 ~~Therefore, debris flows initiated from seepage flow that leads to a landslide~~
487 ~~failure or channelized runoff that entrain sediments in the periglacial area is similar~~
488 ~~with the mechanism of debris flows initiation in non-glacier areas. The fluctuation of~~
489 ~~in air temperature within a specific low range can result into in limited seepage flow.~~
490 ~~As Because the glacier in is limited to one valley is limited, it is unlikely for that~~

批注 [A30]: Answer to reviewer 2#: reference of Gregoretti (2008) is missing and that at line 426 it should be Gregoretti and Dalla Fontana (2008).

批注 [A31]: Answer to reviewer 2#: I suggest the authors a better description of triggering mechanism.....

批注 [A32]: Answer to reviewer 2#: reference of Gregoretti (2008) is missing and that at line 426 it should be Gregoretti and Dalla Fontana (2008).

批注 [A33]: Answer to reviewer 2#: reference of Gregoretti (2008) is missing and that at line 426 it should be Gregoretti and Dalla Fontana (2008).

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491 failure to ~~can be triggered by a the limited amount of ice meltwater in due to~~
492 short term increases of ~~in air temperature; instead, prolonged air temperature~~
493 increases it ~~is are~~ needed to generate more water flow. Rainfall can initiate debris
494 flows from active glacial tills with ~~via a mechanism similar to that of associated with~~
495 storm induced debris flows in non glacier areas (Iverson et al., 1997; Springman et al.,
496 2003; Sassa and Wang, 2005; ~~Gregoretto and Dalla Fontana, Gregoretto, 2008; Kean et~~
497 al., ~~2012~~2013), while the difference lies in the activity of debris and the source of
498 water. In the European Alps, periglacial debris flows are mainly provoked by rainfall,
499 which is also related ~~with to~~ air temperature fluxes (Stoffel et al., ~~2011~~).
500 ~~The~~ Additionally, the ~~portion of values of~~ rainfall and air temperature required ~~for to~~
501 ~~trigger~~ debris flows ~~triggering~~ could be ~~negative~~ inversely correlated. Air temperature
502 increases ~~causes can cause~~ melting and water runoff; ~~thus, and~~ the rainfall ~~needed~~
503 ~~required to create for providing the~~ percolating flows ~~percolation flows~~ or ~~exact~~ critical
504 discharge ~~for to trigger a~~ debris flow ~~triggering~~ would be much less. ~~Beside~~ In addition,
505 the ~~required rainfall, like the~~ intensity and duration ~~of the required rainfall,~~ may also
506 require other preconditions, such as ~~the those associated with the~~ distributions of
507 glaciers and frozen glacial tills and the terrain of the source area, to enhance the debris
508 flow (Lewkowicz and Harris, 2005).

509 The three debris flow events ~~possess were associated with~~ similar annual
510 meteorological conditions, except that the positive air temperature accumulation prior
511 to DF1 was ~~significantly large~~ largest. DF1 occurred at the end of a prolonged period
512 of high air temperature, prior to this, there were instances of failure but no large-scale
513 debris flows. On July 25th, 2010, when the daily rainfall ~~particularly~~ reached 20.7
514 mm, no debris flows were generated because thick active glacial till was still lacking
515 after small failure events. In 2010, the largest daily rainfall occurred on June 7th,
516 accounting for 37.5 mm, at the beginning of an air temperature increase when the
517 glacial till was frozen and had low activity. The lack of glacial till activity was the
518 likely cause of the absence of debris flows. On August 23rd, the daily rainfall was 20.3
519 mm, the antecedent air temperature accumulation ~~dated had remained stable since~~
520 ~~from~~ DF2; and ~~the~~ active glacial till was still ~~under development~~ developing. On

批注 [A34]: Answer to reviewer 2#:
Delete the sentence at line 420-423 and
write Therefore, runoff provided by
rainfall, seepage flow and melting ice or
glacier collapse can initiate debris flow
with the same mechanism of the runoff
generated debris flows in non-glacier
areas (Iverson et al., 1997, Kean et al.,
2012).

批注 [A35]: Answer to reviewer 1#:
negative ? what does this mean ?

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521 September 6th, the antecedent positive air temperature accumulation was ~~smaller~~small,
522 and a low air temperature ~~had emerged~~was observed previously; however, the high
523 rainfall intensity supplemented this lack of prolonged high air temperature.

524 5. Conclusion

525 Climate changes have serious effects ~~on~~in high mountainous areas, and the mass
526 movement of sediments such as periglacial debris flows ~~is~~has become increasingly
527 frequent. Prolonged increases in the mean annual air temperature are regarded as very
528 favourable for periglacial debris flows. In particular, the annual “hot-dry” weather
529 conditions one or two years~~earlier~~prior ~~was~~were responsible for ~~the~~ three debris flow
530 events in Tianmo Valley. Debris flows ~~is~~usually are generally not initiated in the year
531 when the mean annual air temperature spikes, as the melting of internal ice particles
532 lags behind the rate of glacial retreat resulting from ~~the~~a prolonged increase in air
533 temperature rise.

批注 [A36]: Answer to reviewer 2#:
Confused sentence

534 Glacial till is unlimited in ~~the~~ deglaciated areas, ~~while~~and its activity relies on
535 glacial retreat and internal ice particle melting. ~~Changing of~~ Glacial till changes
536 induced by increasing~~increased~~ air temperature ~~is~~are the first steps of~~in~~ forming
537 periglacial debris flows ~~comparing with~~compared to the storm-induced debris flows in
538 non-glacier areas. Glacial tills need~~require~~ a four-phase ~~experience~~process prior to
539 debris flow occurrence. In this process,~~during which the~~ ~~varied~~the variable air
540 temperature condition with~~due to~~ different factors drives the changing glacial till
541 changes, and temperature ~~series~~increases can cause glacier recession~~remove glaciers~~,
542 produce barred glacial till and enhance the glacial till activity step by step. Debris
543 flows ~~could~~can occur when a enough sufficient amount of active glacial till ~~is~~
544 exists ~~existed~~ and rainfall-induced seepage or water-runoff is more likely to generate
545 debris flows.

批注 [A37]: Answers to reviewer #1:
remove glaciers ? do you mean "cause
glacier recession or even
disappearance" ?

546 It is difficult to observe ~~the changes of~~ glacial till changes in source areas of
547 debris flows, and the analysis of the phase conversion of glacial till in this ~~research~~
548 study is based on the triggering conditions and other literature findings. Indeed, the
549 meteorological conditions, such as the antecedent air temperature and meteorological

批注 [A38]: Answer to reviewer 2#:
Confused sentence

550 triggers that drive the phase conversion, are partly ~~overlapped-coupled~~ and difficult to
551 distinguish. In ~~the first future study studies~~, we hope to ~~distinguish-determine~~ the effect
552 of each meteorological condition, and more ~~detail-detailed study studies~~ should be
553 ~~done performed in further research~~.

554 **Acknowledgements:** This research was supported by the National Natural Science Foundation
555 of China (grant ~~n~~Nos. 41190084, 41402283 and 41371038) and the “135” project of IMHE, CAS.
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557 Office and the anonymous reviewers for ~~their~~ constructive comments, which helped us ~~in~~
558 ~~improving improve~~ the contents and presentation of the manuscript.

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Table 1 ~~Changes in~~ Changes in glacier, snow, ~~bared~~ land, gully deposition and vegetation in Tianmo Valley

| Year | Glacier (km ²) | Glacier(eastern branch) (km ²) | Snow (km ²) | Bared land (km ²) | Gully deposition (km ²) | Vegetation (km ²) |
|------|-------------------------------|---|----------------------------|----------------------------------|--|----------------------------------|
| 2000 | 1.77 | 0.16 | 2.13 | 2.80 | 0.44 | 10.46 |
| 2003 | 1.71 | 0.15 | 2.44 | 2.54 | 0.44 | 10.48 |
| 2006 | 1.53 | 0.12 | 2.68 | 2.44 | 0.44 | 10.55 |
| 2009 | 1.45 | 0.096 | 2.81 | 3.03 | 0.47 | 9.90 |
| 2013 | 1.42 | 0.088 | 1.74 | 3.83 | 0.51 | 10.17 |

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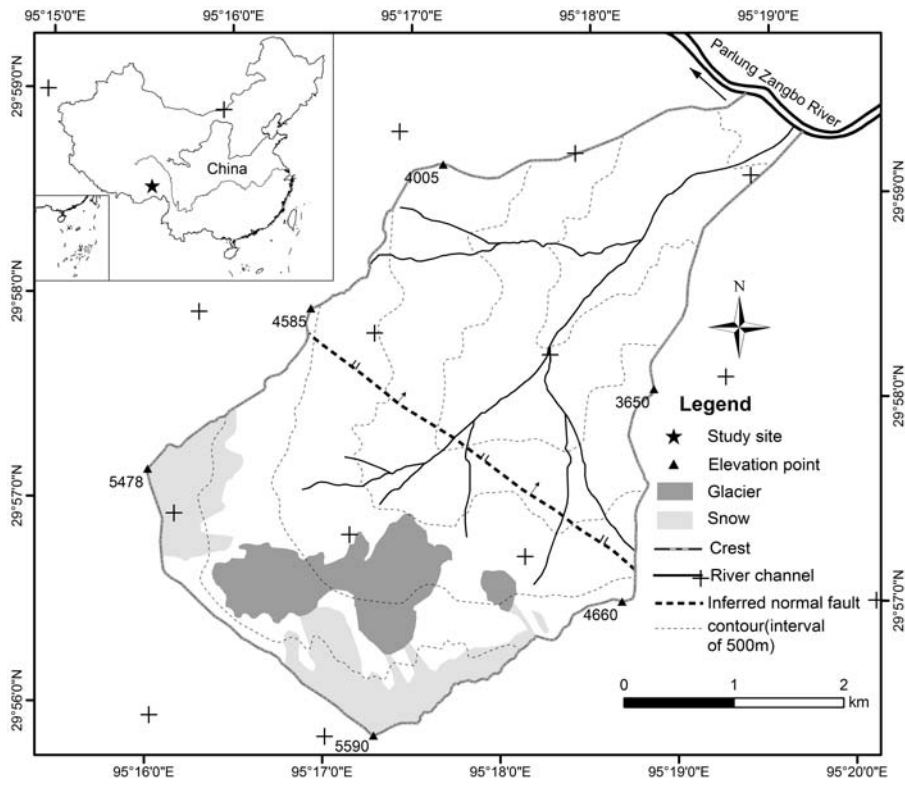
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Table 2 Basic information ~~of-regarding~~ the debris flows in Tianmo Valley and ~~the~~-nearby valleys

| No. | Name | Coordinates | Basin area (km ²) | Glacier area (in 2006) (km ²) | Date | Size class |
|-----|---------------------------|--------------------|----------------------------------|--|----------------------------|------------|
| 1 | Tianmo <u>V</u> alley | 29°59'N 95°19'E | 17.74 | 1.53 | 4 th Sep. 2007 | 6 |
| | | | | | 25 th Jul. 2010 | 5 |
| | | | | | 6 th Sep. 2010 | 5 |
| 2 | Kangbu <u>V</u> alley | 30°16'N 94°48'E | 48.7 | 1.06 | 4 th Sep. 2007 | 3 |
| | | | | | | |
| 3 | Xuewa <u>V</u> alley | 29°57'N 95°23'E | 33.22 | 0.95 | 4 th Sep. 2007 | 2 |
| | | | | | | |
| 4 | Baka <u>V</u> alley | 29°53'N 95°33'E | 22.15 | 2.46 | 7 th Sep. 2007 | 3 |
| | | | | | | |
| 5 | Jiaqing <u>V</u> alley | 30°16'N 94°49'E | 15.51 | 1.12 | 9 th Sep. 2007 | 3 |
| | | | | | | |

带格式表格

带格式的: 上标



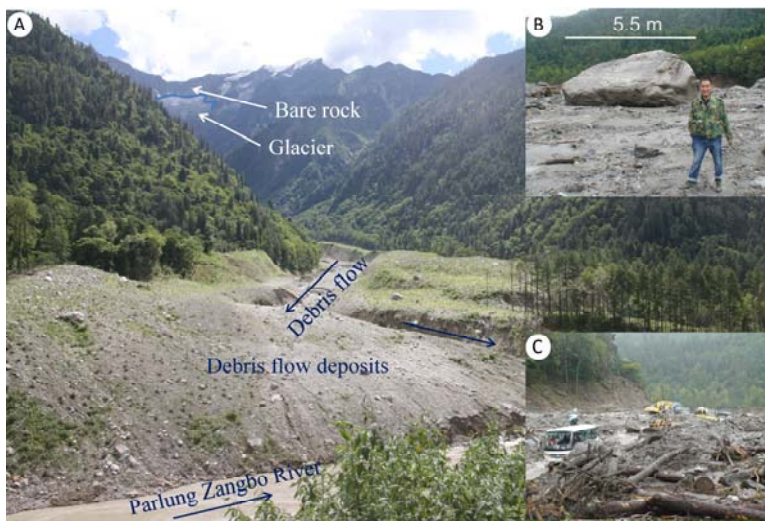
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Figure 1 Location and basic information of Tianmo Valley and related information



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Figure 2 Overview of the valley from the channel(in 2014)

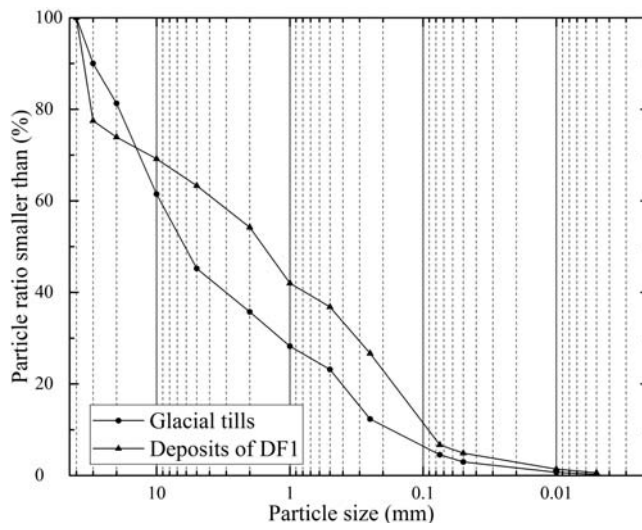


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732 | Figure 3 DF1 in 2007(A. Overview of the Tianmo debris flows from the downstream area; B& C.

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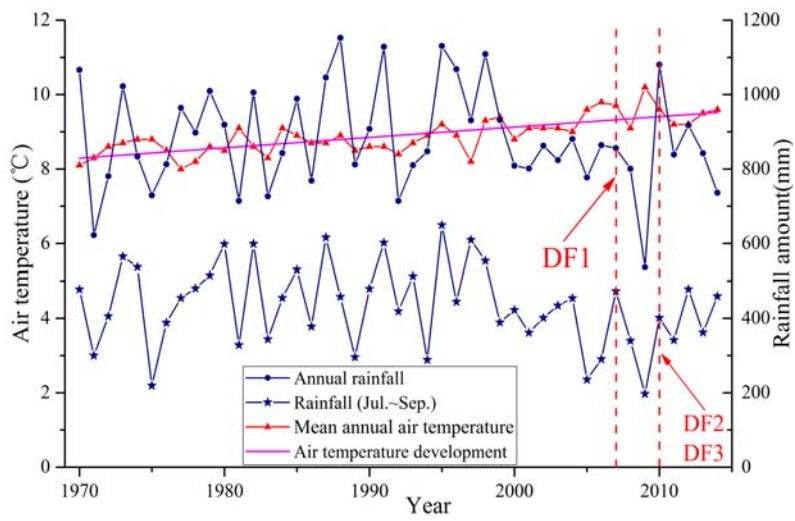
Boulder and debris flow deposits on the north side of the Parlung Zangbo River)



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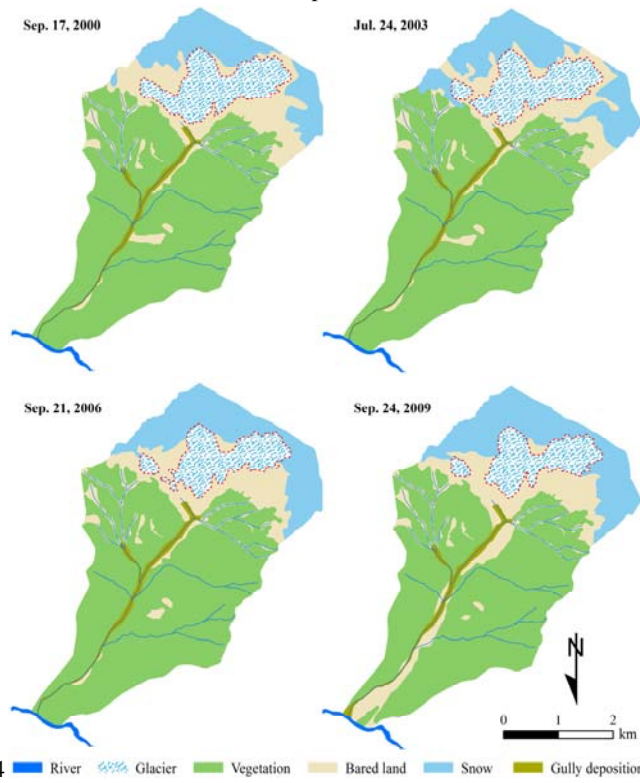
Figure 4 Particle size distributions of the glacial tills and debris flow deposits



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Figure 5 Variation of the mean annual air temperature and rainfall at Bomi from 1970 to



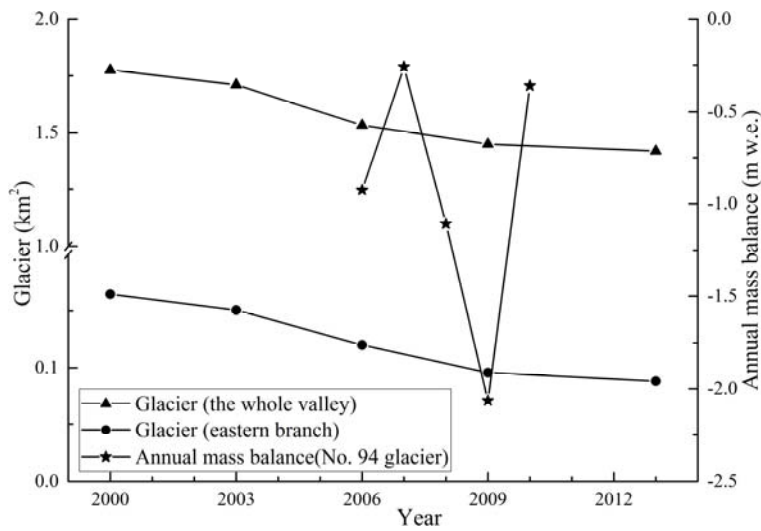
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Figure 6 Distribution and changing of changes in glacier, snow, bared land, gully deposition and vegetation in Tianmo Valley

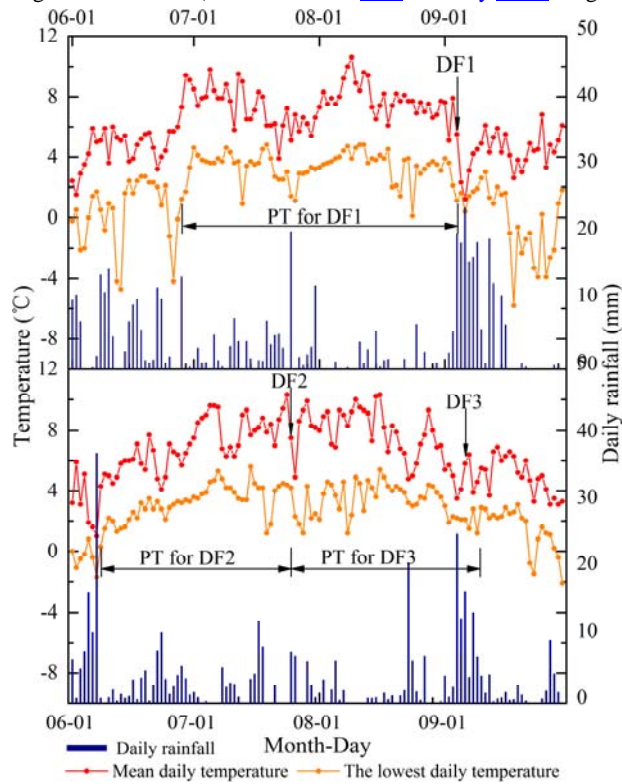


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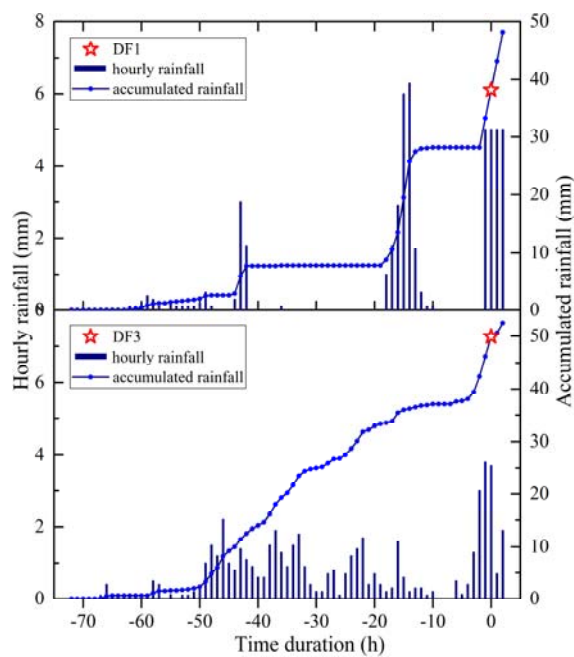
Figure 7 Changing of Changes in glacier via-over time and the measured annual mass balance for of the Parlung No. 94 Glacier (mass balance is-was edited by-from Yang et al.(2015))



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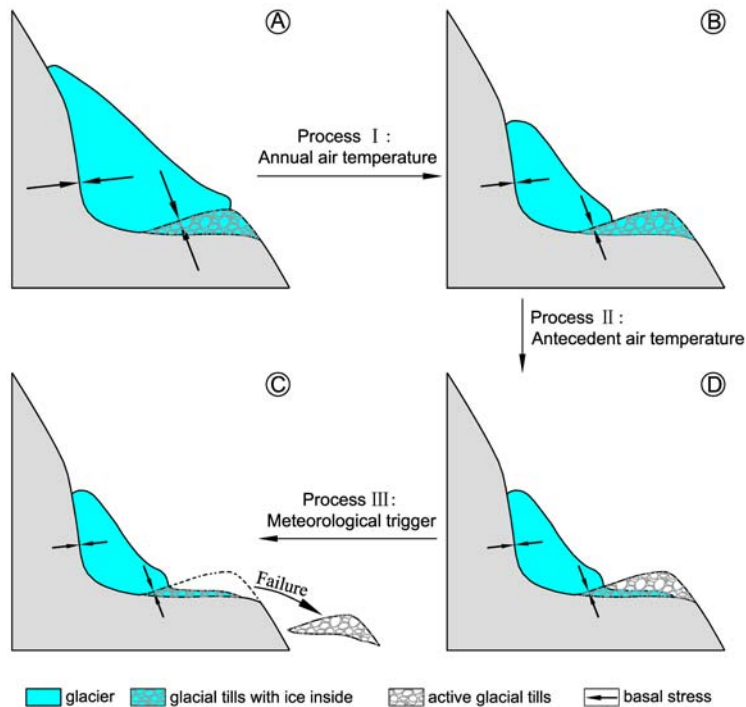
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Figure 8 Air temperature and rainfall before and after DF1, DF2 and DF3



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748 | Figure 9 Variations of the in rainfall accumulation prior to DF1 and DF3 (no rainfall before DF2)



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750 | Figure 10 Changes in a glacier and frozen glacial till before periglacial debris flow initiation (A:
751 | glacial-covered glacial tills; B: uncovered and frozen glacial tills; C: active glacial tills; D: failure
752 | of glacial tills)