



Quantifying the effect of forests on occurrence frequency and intensity of rockfalls

Christine Moos¹, Luuk Dorren¹, Markus Stoffel^{2,3,4}

5 ¹Berne University of Applied Sciences, School of Agricultural, Forest and Food Science HAFL, Länggasse 85, CH-3052 Zollikofen, Switzerland

² Climatic Change and Climate Impacts, Institute for Environmental Science, University of Geneva, 66 Bvd Carl Vogt, CH-1205 Geneva, Switzerland

³ Department of Earth Sciences, University of Geneva, rue des Maraîchers 13, CH-1205 Geneva, Switzerland

10 ⁴Dendrolab.ch, Institute of Geological Sciences, University of Bern, Baltzerstrasse 1+3, CH-3012, Bern, Switzerland

Correspondence to: Christine Moos (christine.moos@bfh.ch)

Abstract. Forests serve as a natural means of protection against rockfall. Due to their barrier effect, they reduce the intensity and the propagation probability of falling rocks and thus the occurrence frequency of an event for a given element at risk. However, despite established knowledge on the protective effect of forests, they are mostly neglected in quantitative rockfall risk analysis. Their inclusion in quantitative rockfall risk assessment would, however, be necessary to express their efficiency in monetary terms and to allow comparison of forests with other protection measures. The goal of this study is to quantify the effect of forests on the occurrence frequency and intensity of rockfalls. We therefore defined a rockfall onset frequency based on a power-law magnitude-frequency distribution, which then served as input for the simulation of rockfall events on a virtual slope and over a period of 1000 yrs. Simulations were run for different forest and non-forest scenarios under varying forest stand and terrain conditions. We determined rockfall frequencies and intensities at five different horizontal distances from the release area. Based on two multivariate statistical prediction models, we investigated which of the terrain and forest characteristics are predominantly driving the role of forest in reducing rockfall occurrence frequency and intensity and whether they are able to predict the effect of forest on rockfall occurrence and hence rockfall risk. The rockfall occurrence frequency below forested slopes is reduced between approximately 10 and 90 % as compared to non-forested slope conditions; whereas rockfall intensity is reduced by 10 to 70 %. This reduction increases with increasing slope length and decreases with decreasing tree density, tree diameter and increasing rock volume as well as in case of clustered or gappy forest structures. The statistical prediction models reveal that the cumulative basal area, rock volume and horizontal forest structure represent key variables for the prediction of the reducing effect of forests. In order to validate these results, models have to be tested on real slopes with a wide variation of terrain and forest conditions.



1. Introduction

Rockfall is a widespread and frequent natural hazard occurring below steep rocky cliffs. The occurrence of rockfall often threatens infrastructures, transportation corridors, and human life. We here define it as a fragment of rock detaching from a release area and proceeding downslope by bouncing, falling, or rolling (Whittow, 1984). Different protective measures are typically implemented in order to reduce risks in rockfall prone areas. These include structural protection measures, land-use planning, early-warning systems or biological measures, nowadays referred to as nature-based or eco- solutions (Agliardi and Crosta, 2003; Corominas et al., 2005; Sättele et al., 2016; Renaud et al., 2013). With regards to rockfall, a well-known biological measure is the protection forest. Such forests can serve as a natural means of protection against rockfall due to their barrier effect. Forests influence rockfall risk by (i) reducing the intensity of falling rocks after collisions with tree stems and by (ii) reducing the propagation probability and thus the occurrence frequency of an event at a given element at risk (Wasser and Perren, 2014).

In order to appropriately account for the positive effects of protective measures on rockfall risk and associated uncertainties, their design should be based on a quantitative risk analysis (Corominas et al., 2005; Straub and Schubert, 2008; Peila and Guardini, 2008). In doing so, the protective effect of the measure can be expressed in monetary terms, thereby allowing to evaluate its efficiency in a cost-benefit analysis (Agliardi et al., 2009). In the case of protection forests, quantitative, risk-based approaches have been only rarely applied in the past. Despite the advanced knowledge on the protective effect of forests and its maintenance (Dorren et al., 2007; Bigot et al., 2009), open questions remain on how protection forests can be quantitatively integrated into rockfall risk analyses (Masuya et al., 2009; Trappmann et al., 2014). Currently, the effect of forests is mostly neglected or only qualitatively assessed in hazard and risk analyses.

The quantification of the influence of forests on rockfall occurrence frequency is particularly demanding, especially if one aims at evaluating the effect of forests at the level of the element at risk. The occurrence frequency of a rockfall event is usually described by the annual exceedance frequency of its magnitude (expressed as the rockfall volume) or intensity (expressed as the kinetic energy of the blocks), assuming that rockfall occurrence follows a Poisson distribution (Corominas et al., 2013). Depending on the data availability and site characteristics, the annual exceedance frequency can be estimated by different approaches including the analysis of historical datasets (Hantz et al., 2003; Hungr et al., 1999; Guzzetti et al., 2003), magnitude-frequency relationships based on power laws (e.g. Agliardi et al., 2009; Lari et al., 2014; Dussauge-Peisser et al., 2002), empirical models describing rockfall frequency as a function of topographic or geological parameters (e.g. Budetta, 2004; Lan et al., 2010), or expert opinion (e.g. Romana et al., 2003). Furthermore, several techniques exist based on which the depositional ages of rocks can be reconstructed in absolute terms (e.g. Lang et al., 1999; McCarroll et al., 2001).



Dendrogeomorphology (Stoffel and Corona, 2014) represents one such approach and has proven to be a reliable method to estimate past rockfall frequencies through coupling the number of rockfall impacts with tree age (Moya et al., 2010; Corona et al., 2013; Trappmann et al., 2014; Perret et al., 2006). However, in most cases, reliable data is scarce and estimation of robust occurrence frequencies remains difficult (Hantz et al., 2003; Lari et al., 2014; Straub, 2005). Therefore, practitioners usually assume scenarios of pre-defined return periods and corresponding rock volumes (e.g. Borter et al., 1999). Such scenarios are typically derived for the current (e.g. forested) situation, but are also applied to reference (e.g. non-forested) situations (Jahn, 1988). At the same time, however, the barrier effect of forests is expected to decrease the occurrence frequency of rockfall at the location of the element at risk. Consequently, scenarios derived with the practitioner's approach may not necessarily be valid for the reference situation and might thus result in biased risk estimations.

Forests do not only reduce the occurrence frequency of rockfall events, but also reduce their intensity by stopping rocks completely and/or by absorbing (part of) their energy (Lundström et al., 2009). In this sense, the intensity of an event refers to the kinetic energy which is released by the rock at impact with the element at risk (Jaboyedoff et al., 2005; Abbruzzese et al., 2009; Lari et al., 2014).

The effect of forest on the occurrence frequency and the intensity is also expected to depend on the structure of a forest stand. Furthermore, the capacity of a tree to absorb energy will vary between species and will depend on its diameter at breast height (DBH) (Dorren et al., 2006). At the stand level, high stem densities are considered to stop falling rocks more effectively because of an enhanced impact probability (Dorren and Berger, 2005; Wehrli et al., 2006). The three-dimensional, probabilistic-deterministic rockfall simulation model RockyFor3D (Dorren, 2015) accounts for these forest effects. It integrates trees spatially explicitly and calculates the energy loss due to impacts against single trees as a function of tree species, DBH, impact height and the horizontal position of the hit (Dorren et al., 2006).

The goal of this study is to quantify the effect of forests on the occurrence frequency and intensity of rockfall by using multiple series of rockfall simulations. In this paper, we define a rockfall onset probability based on a power-law magnitude-frequency distribution, which then serves as input for the simulation of rockfall events on a virtual slope and over a period of 1000 yrs. Simulations were run for different forest and non-forest scenarios under varying forest stand and terrain conditions. These simulations do not only (i) provide input data for the determination of rockfall occurrence frequencies and intensities at five different distances from the release area, but also (ii) yield information how specific forest and terrain characteristics control rockfall occurrence frequency and intensity along a slope. The second question was addressed using multivariate statistical prediction models. Based on these approaches, we then investigate (iii) how rockfall occurrence frequency and intensity differ at a given location with an element at risk on forested and non-forested slopes; (iv) what terrain and forest characteristics are predominantly driving the role of forest in reducing rockfall occurrence frequency and intensity, and (v) whether multivariate statistical models fitted with these terrain and forest characteristics can indeed predict the effect of forest on rockfall occurrence and hence rockfall risk.



2. Material and methods

As this study aimed at an assessment of rockfall in forests under controlled conditions, it was preferred to run simulations on a virtual slope. We designed a slope raster with a resolution of 2 m, a horizontal width of 478 m and a horizontal length of 574 m. The virtual slope has a concave shape and slope angles which are increasing linearly from 20 to 40° from the slope bottom to the release area of rockfall, therefore resulting in a height difference of 328 m. The slope angle of each cell was randomly varied with $\pm 10\%$. The release area of rockfall is rectangular and has a horizontal length of 100 m and a width of 300 m. The release area was placed in the central part of the uppermost slope area and defined as a series of rockfall cliffs. Initial fall height of rocks was defined as 10 m above ground. At the bottom part of the slope, we designed a horizontal road with a width of 10 m and we further added five virtual evaluation lines located at horizontal distances of 0, 150, 300, 450, and 530 m from the downslope side of the release area to the bottom of the slope. These lines allow a systematic assessment of changes in rockfall occurrence frequency and intensity with increasing distance from the release area of rockfalls (Fig. 1). The lines were defined based on equal height differences between them. Noteworthy, the evaluation line at a slope length of 450 m is identical to the level of the road that was built into the model.

15

2.1. Rockfall simulation model

We used the model RockyFor3D for the rockfall simulations, which is a probabilistic process-based rockfall trajectory model simulating trajectories of falling rocks using 3D vectors (Dorren, 2015). RockyFor3D was developed on the basis of real-size rockfall experiments in the field and uses raster maps describing topography (Digital Elevation Model, DEM), rockfall source cells, the elasticity of the surface material, slope surface roughness, the number of trees per cell, DBH of trees in each cell and tree species per cell as input data (Dorren et al., 2004; Dorren et al., 2006). For each rockfall source cell, the trajectories of a given number of rocks are simulated by considering flying and bouncing. Rolling is simulated with short distance bouncing, similar to the approach of Pfeiffer and Bowen (1989). The trajectory of rockfall is primarily determined by topography. The elasticity of the impacted material is calculated based on the normal coefficient of restitution (R_n) which is predefined by seven different soil types or undergrounds. Surface roughness is represented by a mean obstacle height (MOH) representative for 70, 20 and 10 %, respectively, of each cell (for more details see Dorren, 2015). RockyFor3D explicitly calculates the deviation and energy loss after impacts with trees dependent on tree diameter, impact position, and the kinetic energy of the rock before the impact. Provided that the exact positions of trees within the slope are not known, trees are randomly positioned within each pixel according to the number of trees (i.e. forest stand density) assigned to each pixel. The main output of RockyFor3D consists of raster cells containing the maximum kinetic energy, the 95 % confidence interval of all maximum kinetic energy values, the maximum bounce height, the number of rocks passed through each cell, the number of deposited rocks, the maximum simulated velocity, the maximum tree impact height and the number of tree impacts per cell (Dorren et al., 2006; Dorren, 2015). We simulated one block per event following a power-law magnitude frequency distribution and did not consider rock fragmentation.

25
30



2.2. Onset probability

We assume a power-law distribution for the magnitude-frequency relationship of rocks released from the release area, since power laws have proven to fit the release volume distribution of rockfalls (e.g. Dussauge-Peisser et al., 2002; Malamud et al., 2004) with the general form:

$$F(V_i) = \alpha V_i^{-\beta} \quad (1)$$

where $F(V_i)$ is the annual frequency of volume i (V_i).

We derived the exponent of the power law (β) from Carrea et al. (2015) who used Terrestrial Laser Scanning to identify fallen blocks with their respective volumes for the “La Cornalle” site in Switzerland. At this site, the cliff consists of alternate layers of marls and sandstones. The power-law distribution was fitted to volumes ranging from 0.010 m³ to 7.63 m³. For the scope of our study, we considered blocks with volumes between 0.05 m³ and 2.0 m³. These volumes can be potentially hazardous but are still within a range for which forests are assumed to have an effect on rockfall propagation and energy (Stokes, 2006). The constant α of the cumulative power-law distribution was defined as 12 in our study corresponding to a rockfall retreat rate of approximately 0.2 mm/yr for the considered volume range (0.05 m³ and 2.0 m³). This is in the typical range of rockfall retreat rates in alpine regions (Sass and Wollny, 2001; Hoffmann and Schrott, 2002; Moore et al., 2009). The simulations were run over a period of 1000 yr. We calculated the number of expected rockfalls released in 1000 yrs for 20 volume classes (Fig. 2). The source cells were randomly sampled in the release area.

2.3. Forest and terrain scenarios

The soil scenarios (Table 1) considered scree or medium compact soil with small rock fragments (soil type 3) and talus slope or compact soil with large rock fragments (soil type 4), as these are expected to be most frequent, often continuous and with a large spatial distribution. The release area was in all cases defined as soil type 5 (bedrock with thin weathered material or soil cover) and the road was set to soil type 7 (asphalt road). As shown in Table 1, soil roughness was set to 0 m (100 %) in the scenario “zero roughness” and to 0.15 m (10 % of the surface), 0.05 m (20 %) and 0.01 m (70 %) in the scenario “rough”, respectively. Definition of the four forest types (Table 2) was based on natural rockfall protection forests as defined from the Swiss National Forest Inventory (Messmer, 2014). The forest types differ with respect to the diameter at breast height (DBH; ranging from 21-40 cm), dominant tree species (deciduous, conifers) and the number of tree stems (with DBH > 12 cm) per hectare (Nha; 200-500 trees ha⁻¹). The forest stands of each forest type were designed for four different horizontal forest structures (Fig. 2) as follows: random tree distribution, clustered tree distribution, random distribution with gaps of 20 x 20 m and random distribution with 3 aisles of 20 m in width.

The combination of the different forest types and structures and terrain scenarios yielded 49 different simulation scenarios. For each scenario, the expected number of blocks of the 20 volume classes was then simulated 100 times to obtain robust results.



2.4. Statistical analysis

Simulation results were analysed statistically as follows:

- 5 (i) zonal statistics of rockfall frequencies and energies at the level of the evaluation lines
- (ii) statistical comparison of rockfall frequency and intensity between different scenarios and by fitting power-law based intensity-frequency curves
- (iii) design of multivariate statistical models relating the frequency and the intensity reduction of forests to terrain and forest characteristics
- 10 (iv) assessment of the performance of the statistical models and sensitivity to changes in slope angle

For each volume class and simulation scenario, the number of rocks passing a cell (i.e. number of passages) was summed up for the five evaluation lines (EL) and averaged over the 100 simulations. We calculated the rockfall occurrence frequency (Freq) by dividing the sum of the number of passages (Nrp) by 1000 yrs:

$$15 \quad Freq = \frac{\sum_{EL} Nrp}{1000 \text{ yrs}} \quad (2)$$

Consequently the return period (RP) of rockfall is its reciprocal value and represented as follows:

$$RP = \frac{1}{Freq} = \frac{1000 \text{ yrs}}{\sum_{EL} Nrp} \quad (3)$$

20

We calculated an indicator for the reduction in the number of passages by the forest stand (Nrp_{red}) in order to evaluate changes in the frequency between forested and non-forested conditions. The indicator Nrp_{red} is defined as the difference between the number of passages without (Nrp_{nF}) and with forest (Nrp_F), divided by the number of passages without forest (Eq. 4):

25

$$Nrp_{red} = \frac{Nrp_{nF} - Nrp_F}{Nrp_{nF}} \quad (4)$$

We then used the 95th percentile of the maximum energy (E95 in kJ) as an indicator for rockfall intensity. For each EL, we calculated the mean of E95 and averaged it over the 100 simulations. Similarly to occurrence frequency, we calculated the intensity reduction offered by forests ($E95_{red}$). This indicator is defined as the difference between E95 without ($E95_{nF}$) and with forest stand ($E95_F$) divided by E95 without forest (Eq. 5):

30

$$E95_{red} = \frac{E95_{nF} - E95_F}{E95_{nF}} \quad (5)$$



We further determined intensity-frequency distributions of E95 (intensity) and Nrp (frequency) under different forest and non-forest scenarios and at a slope length of 300 m, to which power-law distributions (Eq. 1) were fitted based on least squares (Draper and Smith, 1998).

5 To detect possible effects of forest and terrain characteristics on the forest effect, we first assessed whether Nrp_{red} and $E95_{red}$ significantly differ between different forest and terrain scenarios based on the *Wilcoxon rank-sum test*, with a significance threshold of $p \leq 0.05$. Subsequently, we applied regression tree (RT) models (Breiman et al., 1984) and generalized linear models (GLM) (McCullagh and Nelder, 1989) relating Nrp_{red} and $E95_{red}$ to possible explanatory variables.

10 RTs are a non-parametric regression approach which recursively partitions the data based on explanatory variables. At each node, the data is split into two groups using a single predictor (Breiman et al., 1984). The splitting variable is selected aiming at impurity reduction. This means that daughter nodes have to be as homogeneous (“pure”) as possible. RTs consider parameter interactions and account for non-linearities (Vorpahl et al., 2012). RT models were fitted using the *rpart* function of the party package in the statistical software R (Ripley et al., 2015).

15 We used rock volume, soil type (categorical), soil roughness (categorical), the horizontal forest structure (categorical) and the cumulative basal area (cbA; Eq. 7) of the forest as potential explanatory variables. The latter is defined as the product of the basal area (bA; m^2/ha) for a slope width of 100 m and the forested slope length (fsL; m) from the top of the release area to the respective EL. The basal area (bA) is defined as the area per hectare which is occupied by the cross-section of tree stems (Bitterlich, 1948).

$$20 \quad cbA [m^2 ha^{-1}] = \frac{bA}{100 m} \times fsL = \frac{\sum_{EL} bA / \sum_{EL} area}{100 m} \times fsL \quad (7)$$

25 We calculated the Spearman correlation coefficients to check whether the explanatory variables are substantially correlated. The final GLM was determined using a stepwise backward variable selection with the aim to minimize the Akaike Information Criterion (AIC). The quality of the models was examined with goodness-of-fit tests and customary residual diagnostic plots (Stahel, 2013) indicating that the cumulative basal Area (cbA) should be transformed to the natural logarithm.

30 The GLM and RT were fitted with the simulation data of the concave slope. They were subsequently calibrated with a training data set representing 75% of the data. We further applied three times repeated 10-fold cross validation and calculated the average performance across the hold-out predictions with the aim to avoid over-fitting (Kohavi, 1995). The predictive performance was assessed based on the Root Mean Squared Error normalized with the range of the simulated data (nRMSE).



Furthermore, we tested the statistical prediction models for $N_{rp_{red}}$ with field data of a study site in the French Alps at which real-size rockfall experiments were conducted on forested and non-forested sites (Dorren et al., 2006). We evaluated $N_{rp_{red}}$ at a distance of 223 and 324 m from the release point (as measured along the slope).

To assess whether the forest effect on rockfall frequency and intensity depends on the slope angle, we conducted additional simulations for four linearly shaped slopes with varying slope angles (32°, 35°, 38°, 40°) for forest type 1 with random tree distribution, soil type 3 and rough conditions. On these slopes, we tested the multivariate statistical prediction models designed for the concave slope (GLM, RT) and calculated their performance. On the linearly-shaped slopes, evaluation lines were defined with the same distances along the slopes.

10 3. Results

3.1. Effect of forest on rockfall occurrence frequency

Forest stands considerably reduce rockfall frequency, with differences in the frequency between the forested and non-forested slope scenarios increasing strongly with increasing slope length. In the case of forest type 1 (*Fagus sylvatica* forest with 460 stems ha^{-1}) with randomly distributed trees, the frequency at a distance of 450 m from the release area has been shown to decrease to zero (RP > 1000 yrs) whereas on the non-forested slope, the RP remains at values ranging from 1 to 100 yrs, depending on rock volume (Fig. 4). We also show that with decreasing cbA, the effect of the forest is decreasing ($p < 0.05$; Fig. 6), and the reduction of rockfall is becoming less effective. In a pole-stand *F. sylvatica* forest (forest type 4), by contrast, the return period decreases to values between 30 and 1000 yrs at a slope length of 450 m. In the conifer forest composed of *Pinus sylvestris* and *Larix decidua* (forest type 2), the return period is slightly smaller as compared to deciduous forests. Furthermore, we also illustrate that differences between forested and non-forested slopes will chiefly depend on forest structure. In this sense, $N_{rp_{red}}$ is significantly smaller for a clustered tree distribution, gaps or aisles than for a random tree distribution ($p < 0.05$).

25 The reducing effect of the forest is decreasing with increasing rock volume (Fig. 7; $p < 0.05$). This is especially pronounced for forests with small tree diameters (e.g., forest type 4). Also, $N_{rp_{red}}$ is significantly reduced in case of zero roughness ($p < 0.05$). A significant difference in $N_{rp_{red}}$ also exist between soil types 3 and 4 (see Table 1).

30 According to the final generalized linear model (GLM_{freq}), $N_{rp_{red}}$ is significantly influenced by the cumulative basal area (cbA), rock volume, horizontal forest structure, soil type, soil roughness, and the percentage of conifers present in the forest stand (Table 3). GLM_{freq} has a R^2 of 0.80 and a normalized Root Mean Squared Error (nRMSE) of 0.14 with cross-validation for the training data set and 0.16 for the test data set. We also realize that the nRMSE changes only slightly if GLM_{freq} is applied to linear slopes (Table 4).



The variables reported above were also decisive in the regression tree model (RT_{freq}). The dataset was first partitioned based on a threshold of $60 \text{ m}^2 \text{ ha}^{-1}$ for cbA . In the case where cbA is larger than this value, Nrp_{red} is between 0.3 and 1. At the same time, however, Nrp_{red} clearly decreases in the case that rock volumes become $> \sim 1 \text{ m}^3$. On the other hand, and if cbA is smaller than $60.0 \text{ m}^2 \text{ ha}^{-1}$, the mean Nrp_{red} drops to 0 ($cbA < 28.0 \text{ m}^2 \text{ ha}^{-1}$) and 0.5 ($cbA > 28.0 \text{ m}^2 \text{ ha}^{-1}$ and a rock volume $< 0.75 \text{ m}^3$). The normalized Root Mean Squared Error (nRMSE) of RT_{freq} is 0.14 with cross-validation for the training dataset and 0.16 for the test dataset. As can be seen from Table 4, the nRMSE is in the same range of values for the linear slopes.

In the case of the field site in Vaujany (Table 5), for which real data exist from experiments, the GLM_{freq} and the RT_{freq} models predict Nrp_{red} values of 0.62 and 0.70, respectively, at a distance of 223 m (0.64 measured value) and 0.75 and 0.91, respectively, at a distance of 324 m (1.0 measured value).

3.2. Effect of forest on rockfall intensity

On the concave slope, the rocks reach energies of up to 3000 kJ at a slope length of 300 and 450 m, respectively. In the last 100 m of the slope (530 m), energies are reduced to zero. Similarly to the rockfall occurrence frequency, rockfall intensity is distinctly reduced on the forested slopes compared to the non-forested slope (Fig. 5). Again, the reducing effect is decreased with decreasing cbA , increasing rock volume and for the clustered and gappy forest structures (Fig. 6-7). Furthermore, $E95_{red}$ is significantly smaller on slopes with zero roughness ($p < 0.05$), but does not depend on soil type.

In the final GLM (GLM_{Int}), the horizontal forest structure, soil roughness, percentage of conifer trees, cbA , and rock volume have a significant effect on $E95_{red}$. GLM_{Int} has a R^2 of 0.84 and a nRMSE of 0.10 with cross-validation for the training data set and 0.16 for the test data set. If GLM_{Int} is applied to linear slopes, we observe that the nRMSE values increase slightly (Table 4).

In the regression tree model (RT_{Int}), cbA was selected as a splitting variable. Figure 7 illustrates that in the case of high cbA ($> 85.7 \text{ m}^2 \text{ ha}^{-1}$), $E95_{red}$ is distinctively smaller with a clustered or gappy forest structure (i.e. ~ 0.75 with a random tree distribution, ~ 0.55 with aisles, gaps, or clustered tree distribution). The nRMSE of RT_{Int} is 0.10 with cross-validation for the training data set and 0.09 for the test data set. Similar to GLM_{Int} , we observe that the nRMSE of RT_{Int} values hardly changes on linear slopes (Table 4).

3.3. Intensity-frequency curves

Analysis of intensity-frequency distributions of rockfalls depends strongly on the forest cover. In the case of non-forested slopes, the intensity-frequency curve is substantially shifted upward compared to forested slopes at a distance of 300 m downslope from the start area, thereby indicating a higher frequency (intensity) for a given intensity (frequency) (Figure 10).



In other words, the β and the α coefficients (Eq. 1) of the power law fitted to the intensity-frequency distributions are considerably lower when forest cover is present as compared to non-forested conditions (Table 6).

4. Discussion and conclusion

5 In this study we investigated the role of forests – in terms of stand density and species composition – on rockfall occurrence at increasing distances from the release area of rockfalls by using a hypothetical slope typical of mountain environments. Based on a large number of simulation runs using different scenarios, we show that rockfall occurrence frequency below forested slopes is reduced between approximately 10 and 90 percent as compared to non-forested slope conditions. Rockfall intensity is also reduced – although to a slightly smaller extent – by 10 and 70 percent. These findings are in agreement with
10 the study of Lopez-Saez et al. (2016) who found a distinct increase in rockfall return periods (e.g., from 143 yrs under non-forested conditions in 1850 to >1000 yrs under recently grown forest in 2013 and for a block volume of 1.2 m³). In this particular case in the Chartreuse massif (France), the disappearance of viticultural landscapes has led to intense (natural) afforestation and can thus be seen as a natural example for the validation of our theoretical results. Similar to our study, Lopez-Saez et al. (2016) also observe that the kinetic energy of rocks clearly decreases at the bottom of the slope and with
15 increasing forested surface, which is again in concert with the findings of our study. Stoffel et al. (2005) investigated spatial and temporal variations of rockfall activity in a protection forest in the Swiss Alps based on dendrogeomorphic data. They reconstructed a decrease in rockfall rates after the recolonization of part of the slope where most of the forest was destroyed after a high magnitude event in 1720. Masuya et al. (2009), on the other hand, did not find a decrease in the number of rockfalls reaching the damage potential at a distance of 350 m from the rockfall source based on three-dimensional
20 simulations taking vegetation probabilistically into account, but an increase in the spread of the rockfalls and lower rock energies. It has to be mentioned that the considered vegetation cover featured relatively small trees and low tree density. The multivariate statistical models used in this study allowed quantification of the reduction of rockfall frequency and intensity and its prediction under varying forest and slope conditions. Both models (GLM and RT model) revealed that the effect a forest stand has on rockfall will depend clearly on the cumulative basal area (cbA) of trees, the horizontal forest
25 structure, and on the rock volume. We realize that rockfall frequency and intensity are significantly reduced with decreasing cbA and increasing rock volume as well as in clustered or gappy forests, and are now able to quantify these effects. Moreover, the results also demonstrate how the reducing effect of forests is enhanced with increasing soil roughness and soil elasticity. The influence of the two slope parameters was, however, only significant in the GLM, but not in the RT model. According to the RT models, the forest effect of rockfall frequency appears to depend mainly on cbA and rock volume, whereas cbA and forest structure appear as the most decisive factors for the reduction in rockfall intensity. Rock volume, by
30 contrast, only has a marginal influence on the reducing effect of forest on rockfall intensity (Fig. 8). The maximum reduction of the rockfall energy by forests is reached for volumes between 0.3 and 0.6 m³. This appears to be the optimal combination



between a sufficiently high tree impact probability and rockfall impact energy. For larger rocks, however, impact probability increases further, but the rockfall energy cannot be dissipated during a single tree impact.

The cbA appears to be a good measure of the protection efficacy of forests, as it combines the basal area (which is determined by tree density and tree diameter) with the forested slope length – two parameters which have been promoted as
5 key variables for forest management in previous work (Perret et al., 2004; Berger and Dorren, 2007; Rammer et al., 2015).
Based on our results, we recommend a minimum cbA of about 80-90 m² ha⁻¹ for rock volumes larger than 1 m³ and a
minimum cbA of about 30-40 m² ha⁻¹ for volumes smaller than 1 m³. Compared to the minimum threshold of 20 m² ha⁻¹ for
the basal area of a rockfall protection forest as suggested by Dorren et al. (2015), this corresponds to a forested slope length
of 450 m (rock volume >1 m³) and 200 m (rock volume <1 m³), respectively.

10 According to the RT models, the horizontal forest structure is particularly important when it comes to the reduction of
rockfall intensity. We demonstrate that the kinetic energies of rocks are significantly higher in the case of forest stands with
a clustered tree structure or in forests with gaps or aisles compared to random tree distribution. The horizontal forest
structure, by contrast, is only of secondary importance for the reduction of rockfall frequency and the number of trees which
are impacted by the rock in motion will be decisive.

15 The performance of the implemented statistical prediction models is satisfactory. They yielded relatively low normalized
Root Mean Squared Error (nRMSE), also when applying cross-validation. This indicates that the generalization capacity of
the models is relatively high and over-fitting unlikely. The application of the models to four different linear slopes with
varying slope angles (32°, 35°, 38°, 40°) did not substantially change the nRMSE (Table 4) suggesting that the models are
20 relatively robust with respect to slope angles.

Various factors influence the robustness of the developed models with respect to the applicability to real slopes. The
simulated rock volume was limited to 2.0 m³ and therefore they do not necessarily apply to larger volumes. In the GLM, the
Nrp_{red} is linearly extrapolated for larger rock volumes, whereas in the RT model a threshold of 2.0 m³ is fixed and the
25 reductive effect of the forest for larger volumes might be overestimated. Furthermore, since we used the rockfall model
Rockyfor3D as an important basis for this study, we assume that this model simulates the rockfall process and impacts
against trees sufficiently realistic. It has to be considered, however, that the model takes into account two “species” only,
being coniferous and broadleaved, for calculating the energy dissipative capacity of trees. In reality, the range of this
capacity is much larger and shows huge variations due to, for example, tree vitality, tree anchoring and other site conditions
30 determining tree growth. Additionally, Rockyfor3D uses a simplified stochastic approach to account for different rock
shapes. When considering a single rockfall event with a rock shape that does not correspond to standard shapes such as
rectangular or spherical, differences between model and reality can be expected.



We could show that the intensity-frequency distributions of rockfall events can be significantly altered below forests compared to non-forested situations. This supports the importance of a coupled consideration of intensity and frequency in order to fully account for the forest effect as it was already reported for other natural hazard processes (Alila et al., 2009). Otherwise, risk analyses are expected to be biased and the risk below forests may be overestimated resulting in over-
5 dimensional structural protection measures associated with high costs.

Overall, this study substantiates the importance of forests in reducing rockfall risk. The statistical prediction models based on the simulation results for different forest and terrain scenarios allow to quantify this effect and to predict it for other slopes, given the constraints mentioned above. In order to validate these results, the models have to be tested on real slopes. Dendrogeomorphic data on tree impacts (Trappmann and Stoffel, 2013, 2015; Morel et al., 2015) might help evaluation of
10 changes in frequency reduction along the slope depending on the forest structure (Corominas and Moya, 2010). However, serious validation of the difference between forested and non-forested slopes remains difficult since data is missing.

The shown influence of the forest type and structure on rockfall frequency and intensity underlines the importance of forest management aiming at maintenance of its protection function. Disturbances, such as fire, wind, or insects, can temporarily eliminate or at least substantially reduce the protective effect of forests (Maringer et al., 2016; Cordonnier et al., 2008). Also
15 the rockfall process itself, and such as extreme rockfall events, can destroy considerable parts of the forest and, thus, encompass higher rockfall frequency and intensity in the following years (Stoffel et al., 2005).



References

- Abbruzzese, J. M., Sauthier, C., and Labiouse, V.: Considerations on Swiss methodologies for rock fall hazard mapping based on trajectory modelling, *Nat. Hazards Earth Syst. Sci.*, 9, 1095–1109, 2009.
- Agliardi, F. and Crosta, G. B.: High resolution three-dimensional numerical modelling of rockfalls, *International Journal of Rock Mechanics and Mining Sciences*, 40, 455–471, doi:10.1016/S1365-1609(03)00021-2, 2003.
- Agliardi, F., Crosta, G. B., and Frattini, P.: Integrating rockfall risk assessment and countermeasure design by 3D modelling techniques, *Nat. Hazards Earth Syst. Sci.*, 9, 1059–1073, 2009.
- Alila, Y., Kuraś, P. K., Schnorbus, M., and Hudson, R.: Forests and floods: A new paradigm sheds light on age-old controversies, *Water Resour. Res.*, 45, doi:10.1029/2008WR007207, 2009.
- 10 Berger, F. and Dorren, L. K.: Principles of the tool Rockfor.net for quantifying the rockfall hazard below a protection forest, *Schweizerische Zeitschrift für Forstwesen*, 158, 157–165, doi:10.3188/szf.2007.0157, 2007.
- Bigot, C., Dorren, L., and Berger, F.: Quantifying the protective function of a forest against rockfall for past, present and future scenarios using two modelling approaches, *Nat Hazards*, 49, 99–111, doi:10.1007/s11069-008-9280-0, 2009.
- Bitterlich, W.: Die Winkelzählprobe, *Allgemeine Forst- und Holzwirtschaftszeitung*, 59, 4–5, 1948.
- 15 Borter, P., Heinimann, H. R., Bart, R., Egli, T., and Gächter, M.: Risikoanalyse bei gravitativen Naturgefahren: Methode, *Umwelt-Materialien*, 107/I, Bundesamt für Umwelt, BAFU, Bern, 117 pp., 1999.
- Breiman, L., Friedman, J., Stone, C. J., and Olshen, R. A.: *Classification and Regression Trees*, Taylor & Francis, 1984.
- Budetta, P.: Assessment of rockfall risk along roads, *Natural Hazards and Earth System Sciences*, 4, 71–81, 2004.
- Carrea, D., Abellan, A., Derron, M.-H., and Jaboyedoff, M.: Automatic Rockfalls Volume Estimation Based on Terrestrial Laser Scanning Data, in: *Engineering Geology for Society and Territory - Volume 2*, Lollino, G., Giordan, D., Crosta, G. B., Corominas, J., Azzam, R., Wasowski, J., and Sciarra, N. (Eds.), Springer International Publishing, Cham, 425–428, 2015.
- 20 Cordonnier, T., Courbaud, B., Berger, F., and Franc, A.: Permanence of resilience and protection efficiency in mountain Norway spruce forest stands: A simulation study, *Forest Ecology and Management*, 256, 347–354, doi:10.1016/j.foreco.2008.04.028, 2008.
- 25 Corominas, J., Copons, R., Moya, J., Vilaplana, J. M., Altimir, J., and Amigó, J.: Quantitative assessment of the residual risk in a rockfall protected area, *Landslides*, 2, 343–357, doi:10.1007/s10346-005-0022-z, 2005.
- Corominas, J. and Moya, J.: Contribution of dendrochronology to the determination of magnitude–frequency relationships for landslides, *Geomorphology*, 124, 137–149, doi:10.1016/j.geomorph.2010.09.001, 2010.
- Corominas, J., van Westen, C., Frattini, P., Cascini, L., Malet, J.-P., Fotopoulou, S., Catani, F., Van Den Eeckhaut, M., Mavrouli, O., Agliardi, F., Pitilakis, K., Winter, M. G., Pastor, M., Ferlisi, S., Tofani, V., Hervás, J., and Smith, J. T.: Recommendations for the quantitative analysis of landslide risk, *Bull Eng Geol Environ*, doi:10.1007/s10064-013-0538-8, 2013.



- Corona, C., Trappmann, D., and Stoffel, M.: Parameterization of rockfall source areas and magnitudes with ecological recorders: When disturbances in trees serve the calibration and validation of simulation runs, *Geomorphology*, 202, 33–42, doi:10.1016/j.geomorph.2013.02.001, 2013.
- Dorren, L.: Rockyfor3D (v5.2) revealed – Transparent description of the complete 3D rockfall model, *ecorisQ paper*, 32, 5 2015.
- Dorren, L., Berger, F., Frehner, M., Huber, M., Kühne, K., Métral, R., Sandri, A., Schwitter, R., Thormann, J.-J., and Wasser, B.: Das neue NaiS-Anforderungsprofil Steinschlag, *Schweizerische Zeitschrift für Forstwesen*, 166, 16–23, doi:10.3188/szf.2015.0016, 2015.
- Dorren, L., Berger, F., Jonsson, M., Krautblatter, M., Mölk, M., Stoffel, M., and Wehrli, A.: State of the art in rockfall – forest interactions, *Schweizerische Zeitschrift für Forstwesen*, 158, 128–141, doi:10.3188/szf.2007.0128, 2007.
- Dorren, L., Berger, F., and Putters, U. S.: Real-size experiments and 3-D simulation of rockfall on forested and non-forested slopes, *Natural Hazards and Earth System Sciences*, 6, 145–153, 2006.
- Dorren, L. and Berger, F.: Stem breakage of trees and energy dissipation during rockfall impacts, *Tree physiology*, 26, 63–71, 2005.
- 15 Dorren, L., Maier, B., Putters, U. S., and Seijmonsbergen, A. C.: Combining field and modelling techniques to assess rockfall dynamics on a protection forest hillslope in the European Alps, *Geomorphology*, 57, 151–167, doi:10.1016/S0169-555X(03)00100-4, 2004.
- Draper, N. R. and Smith, H.: *Applied regression analysis*, 3.th ed., Wiley series in probability and statistics, John Wiley & Sons, Inc., 1998.
- 20 Dussauge-Peisser, C., Helmstetter, A., Grasso, J.-R., Hantz, D., Desvarreux, P., Jeannin, M., and Giraud, A.: Probabilistic approach to rock fall hazard assessment: potential of historical data analysis, *Natural Hazards and Earth System Sciences*, 2, 15–26, 2002.
- Guzzetti, F., Reichenbach, P., and Wieczorek, G.F.: Rockfall hazard and risk assessment in the Yosemite Valley, California, USA, *Natural Hazards and Earth System Sciences*, 3, 491–503, 2003.
- 25 Hantz, D., Dussauge-Peisser, C., Jeannin, M., and Vengeon, J. M.: Rock fall hazard assessment: From qualitative to quantitative failure probability, *Fast Slope Movements*, Naples, 263–267, 2003.
- Hoffmann, T. and Schrott, L.: Modelling sediment thickness and rockwall retreat in an Alpine valley using 2D-seismic refraction (Reintal, Bavarian Alps), *Zeitschrift für Geomorphologie, Supplementband*, 127, 153–173, 2002.
- Hungr, O., Evans, S. G., and Hazzard, J.: Magnitude and frequency of rock falls and rock slides along the main transportation corridors of southwestern British Columbia, *Can. Geotech. J.*, 36, 224–238, doi:10.1139/t98-106, 1999.
- 30 Jaboyedoff, M., Dudt, J. P., and Labiouse, V.: An attempt to refine rockfall hazard zoning based on the kinetic energy, frequency and fragmentation degree, *Nat. Hazards Earth Syst. Sci.*, 5, 621–632, doi:10.5194/nhess-5-621-2005, 2005.
- Jahn, J.: *Entwaldung und Steinschlag*, in: *International Congress Interpraevent, Conference Proceedings, Interpraevent*, Graz, 185–198, 1988.



- Kohavi, R.: A study of cross-validation and bootstrap for accuracy estimation and model selection., International Joint Conference on Artificial Intelligence (IJCAI), 1995.
- Lan, H., Martin, C. D., Zhou, C., and Lim, C. H.: Rockfall hazard analysis using LiDAR and spatial modeling, *Geomorphology*, 118, 213–223, doi:10.1016/j.geomorph.2010.01.002, 2010.
- 5 Lang, A., Moya, J., Corominas, J., Schrott, L., and Dikau, R.: Classic and new dating methods for assessing the temporal occurrence of mass movements, *Geomorphology*, 30, 33–52, doi:10.1016/S0169-555X(99)00043-4, 1999.
- Lari, S., Frattini, P., and Crosta, G. B.: A probabilistic approach for landslide hazard analysis, *Engineering Geology*, 182, 3–14, doi:10.1016/j.enggeo.2014.07.015, 2014.
- Lopez-Saez, J., Corona, C., Eckert, N., Stoffel, M., Bourrier, F., and Berger, F.: Impacts of land-use and land-cover changes on rockfall propagation: Insights from the Grenoble conurbation, *Science of the Total Environment*, 547, 345–355, doi:10.1016/j.scitotenv.2015.12.148, 2016.
- 10 Lundström, T., Jonsson, M. J., Volkwein, A., and Stoffel, M.: Reactions and energy absorption of trees subject to rockfall: a detailed assessment using a new experimental method, *Tree physiology*, 29, 345–359, doi:10.1093/treephys/tpn030, 2009.
- Malamud, B. D., Turcotte, D. L., Guzzetti, F., and Reichenbach, P.: Landslide inventories and their statistical properties, *Earth Surf. Process. Landforms*, 29, 687–711, doi:10.1002/esp.1064, 2004.
- 15 Maringer, J., Ascoli, D., Dorren, L., Bebi, P., and Conedera, M.: Temporal trends in the protective capacity of burnt beech forests (*Fagus sylvatica* L.) against rockfall, *Eur J Forest Res*, doi:10.1007/s10342-016-0962-y, 2016.
- Masuya, H., Amanuma, K., Nishikawa, Y., and Tsuji, T.: Basic rockfall simulation with consideration of vegetation and application to protection measure, *Nat. Hazards Earth Syst. Sci.*, 9, 1835–1843, doi:10.5194/nhess-9-1835-2009, 2009.
- 20 McCarroll, D., Shakesby, R. A., and Matthews, J. A.: Enhanced rockfall activity during the Little Ice Age: further lichenometric evidence from a Norwegian talus, *Permafrost and Periglacial Processes*, 12, doi:10.1002/ppp.359, 2001.
- McCullagh, P. and Nelder, J. A.: *Generalized Linear Models*, 2.th ed., Chapman & Hall/CRC, London, 1989.
- Messmer, C.: *Grundflächen im Steinschlagschutzwald: Ein Vergleich mit LFI 3 Daten*, Master thesis, Institute for Terrestrial Ecosystems, Swiss Federal Institute of Technology, Zürich, 129 pp., 2014.
- 25 Moore, J. R., Sanders, J. W., Dietrich, W. E., and Glaser, S. D.: Influence of rock mass strength on the erosion rate of alpine cliffs, *Earth Surf. Process. Landforms*, 34, 1339–1352, doi:10.1002/esp.1821, 2009.
- Morel, P., Trappmann, D., Corona, C., and Stoffel, M.: Defining sample size and sampling strategy for dendrogeomorphic rockfall reconstructions, *Geomorphology*, 236, 79–89, doi:10.1016/j.geomorph.2015.02.017, 2015.
- Moya, J., Corominas, J., Pérez Arcas, J., and Baeza, C.: Tree-ring based assessment of rockfall frequency on talus slopes at Solà d’Andorra, Eastern Pyrenees, *Geomorphology*, 118, 393–408, doi:10.1016/j.geomorph.2010.02.007, 2010.
- 30 Peila, D. and Guardini, C.: Use of the event tree to assess the risk reduction obtained from rockfall protection devices, *Nat. Hazards Earth Syst. Sci.*, 8, 1441–1450, 2008.
- Perret, S., Dolf, F., and Kienholz, H.: Rockfalls into forests: Analysis and simulation of rockfall trajectories ? considerations with respect to mountainous forests in Switzerland, *Landslides*, 1, doi:10.1007/s10346-004-0014-4, 2004.



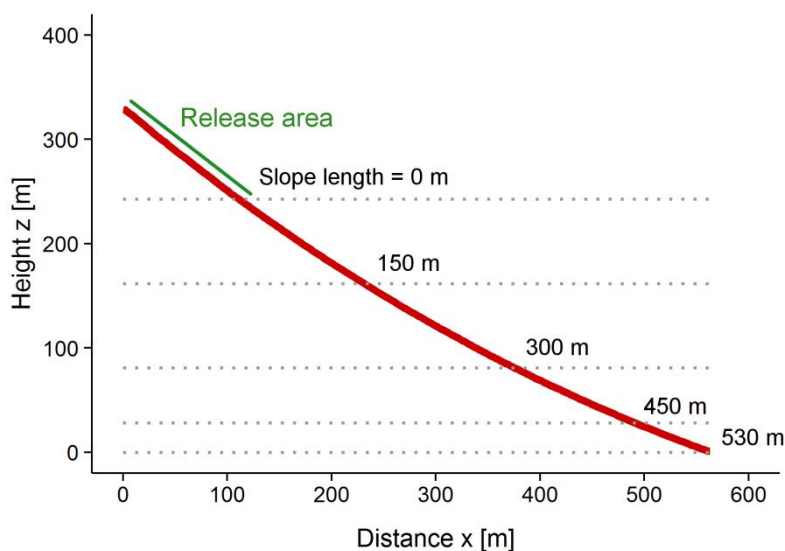
- Perret, S., Stoffel, M., and Kienholz, H.: Spatial and temporal rockfall activity in a forest stand in the Swiss Prealps - A dendrogeomorphological case study, *Geomorphology*, 74, 219–231, doi:10.1016/j.geomorph.2005.08.009, 2006.
- Pfeiffer, T. and Bowen, T.: Computer simulation of rockfalls, *Bulletin of the Association of Engineering Geologists*, 26, 135–146, 1989.
- 5 Rammer, W., Brauner, M., Ruprecht, H., and Lexer, M. J.: Evaluating the effects of forest management on rockfall protection and timber production at slope scale, *Scandinavian Journal of Forest Research*, 30, 719–731, doi:10.1080/02827581.2015.1046911, 2015.
- Renaud, F. G., Sudmeier-Rieux, K., and Estrella, M. (Eds.): *The role of ecosystems in disaster risk reduction*, United Nations University, 2013.
- 10 Ripley, B., Therneau, T., and Atkinson, B.: Package “rpart”, Ripley, Brian, 2015.
- Romana, M., Seron, J. B., and Montalar, E. (Eds.): *SMR Geomechanics classification: Application, experience and validation*, South African Institute of Mining and Metallurgy, 4 pp., 2003.
- Sass, O. and Wollny, K.: Investigations regarding Alpine talus slopes using ground-penetrating radar (GPR) in the Bavarian Alps, Germany, *Earth Surf. Process. Landforms*, 26, 1071–1086, doi:10.1002/esp.254, 2001.
- 15 Sättele, M., Bründl, M., and Straub, D.: Quantifying the effectiveness of early warning systems for natural hazards, *Nat. Hazards Earth Syst. Sci.*, 16, 149–166, doi:10.5194/nhess-16-149-2016, 2016.
- Stahel, W.: *Lineare Regression: Seminar für Statistik*, Zürich, 2013.
- Stoffel, M. and Corona, C.: Dendroecological dating of geomorphic disturbance in trees, *Tree-Ring Research*, 70, 3–20, doi:10.3959/1536-1098-70.1.3, 2014.
- 20 Stoffel, M., Schneuwly, D., Bollschweiler, M., Lièvre, I., Delaloye, R., Myint, M., and Monbaron, M.: Analyzing rockfall activity (1600–2002) in a protection forest—a case study using dendrogeomorphology, *Geomorphology*, 68, 224–241, doi:10.1016/j.geomorph.2004.11.017, 2005.
- Stokes, A.: Selecting tree species for use in rockfall-protection forests, *Forest, Snow and Landscape Research*, 80, 77–86, 2006.
- 25 Straub, D.: Natural hazards risk assessment using Bayesian networks, in: *Proceedings ICOSSAR’05*, Augusti, G., Schueller, G. I., and Ciampoli, M. (Eds.), 9th International Conference on Structural Safety and Reliability, Rome, June 19–23, Millpress, Rotterdam, 2509–2516, 2005.
- Straub, D. and Schubert, M.: Modeling and managing uncertainties in rock-fall hazards, *Georisk: Assessment and Management of Risk for Engineered Systems and Geohazards*, 2, 1–15, doi:10.1080/17499510701835696, 2008.
- 30 Trappmann, D. and Stoffel, M.: Counting scars on tree stems to assess rockfall hazards: A low effort approach, but how reliable?, *Geomorphology*, 180–181, 180–186, doi:10.1016/j.geomorph.2012.10.009, 2013.
- Trappmann, D. and Stoffel, M.: Visual dating of rockfall scars in *Larix decidua* trees, *Geomorphology*, 245, 62–72, doi:10.1016/j.geomorph.2015.04.030, 2015.



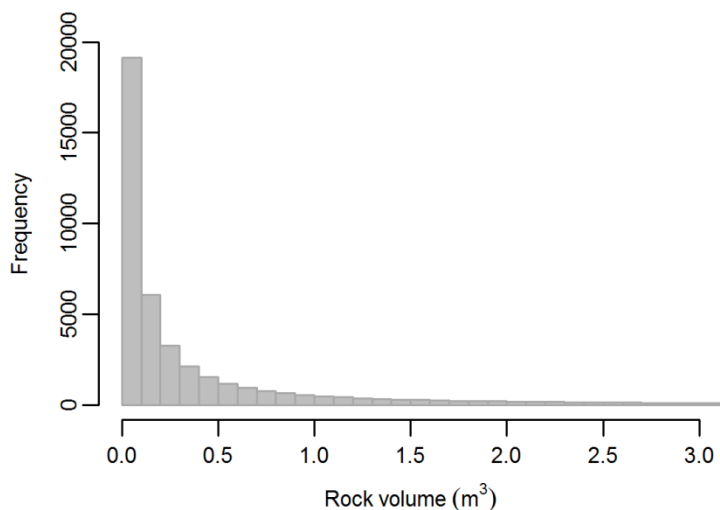
- Trappmann, D., Stoffel, M., and Corona, C.: Achieving a more realistic assessment of rockfall hazards by coupling three-dimensional process models and field-based tree-ring data, *Earth Surf. Process. Landforms*, 39, 1866–1875, doi:10.1002/esp.3580, 2014.
- Vorpahl, P., Elsenbeer, H., Märker, M., and Schröder, B.: How can statistical models help to determine driving factors of
5 landslides?, *Ecological Modelling*, 239, 27–39, doi:10.1016/j.ecolmodel.2011.12.007, 2012.
- Wasser, B. and Perren, B.: Wirkung von Schutzwald gegen gravitative Naturgefahren – Protect-Bio, *Schweizerische Zeitschrift für Forstwesen*, 165, 275–283, doi:10.3188/szf.2014.0275, 2014.
- Wehrli, A., Weisberg, P. J., Schönenberger, W., Brang, P., and Bugmann, H.: Improving the establishment submodel of a
10 forest patch model to assess the long-term protective effect of mountain forests, *Eur J Forest Res*, 126, 131–145, doi:10.1007/s10342-006-0142-6, 2006.
- Whittow, J.: *Dictionary of Physical Geography*, Penguin, London, 1984.



Figures



5 **Fig. 1:** Profile of the virtually constructed digital elevation model (in red) used for the rockfall simulations. Dotted lines indicate the levels at which rockfall occurrence frequency and intensity were evaluated. The rockfall release area is marked in green. The initial fall height of rocks was set to 10 meters above ground.



10 **Fig. 2:** Expected number of rocks released in 1000 yrs on the virtual slope. Calculations are based on a power-law volume-frequency relationship, where β is the cumulative volume frequency distribution and calculated at 0.463 (Carrea et al., 2014), and where α was set to 12.

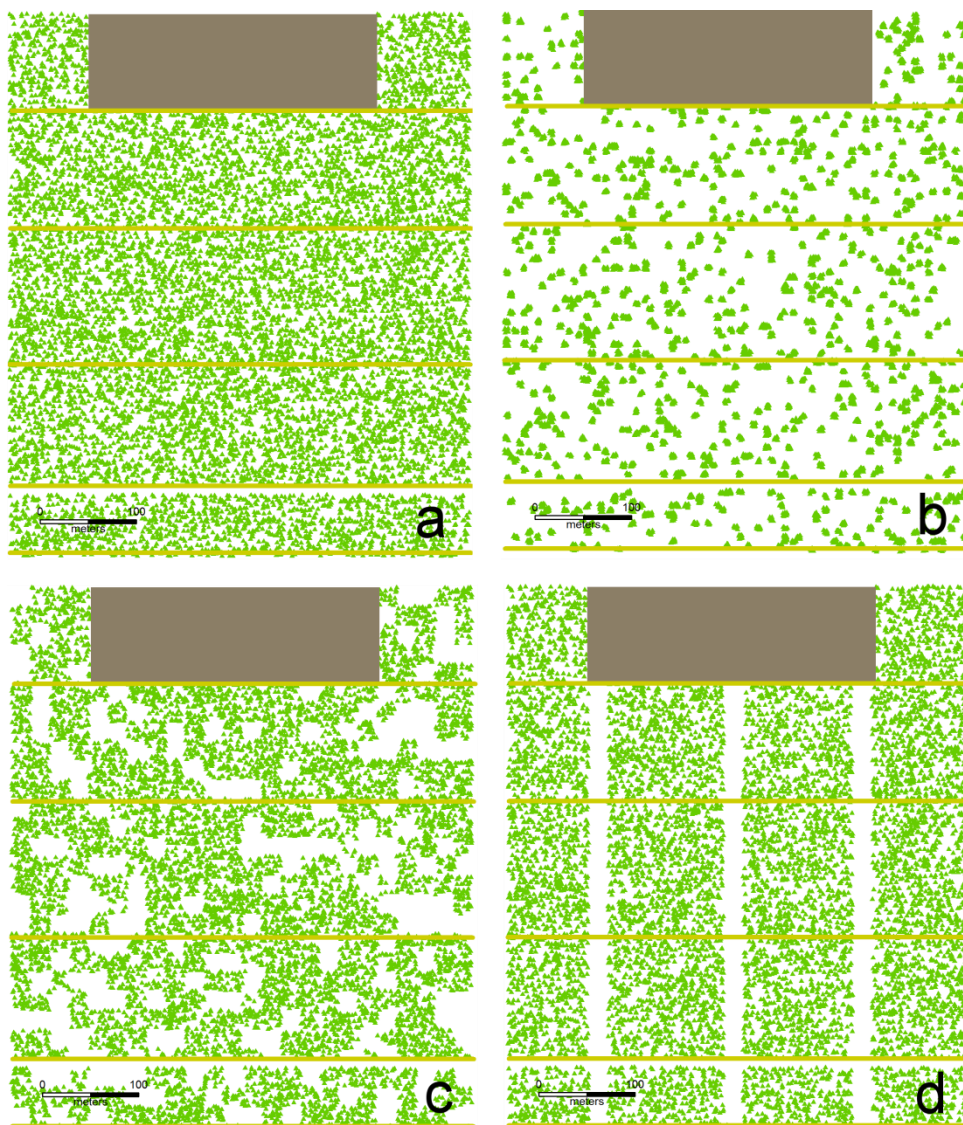


Fig. 3: Design of forest structures and release area of rockfalls for simulation runs. For each forest type, we considered four different scenarios regarding the horizontal forest structure. Forest type 1 is illustrated in (a) with a *random* tree distribution and (b) with random distribution of trees in *clusters* of 10 trees; (c) with a distribution of trees with *random gaps* (minimum 20 x 20 m); and in (d) with 3 *aisles* of 20 m in width starting below the release area of rockfall.

5

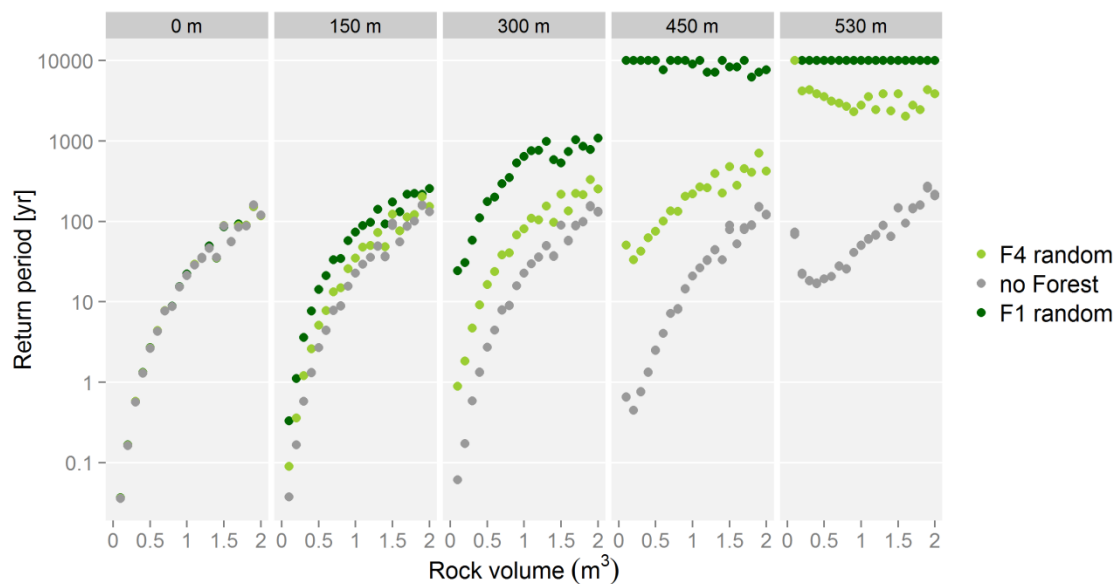


Figure 4: Return periods of rockfall (averaged over 100 simulations) at different evaluation zones located at 0-530 m downslope of the release area and for rock volumes ranging from 0.01 to 2.0 m³ under forested (forest type 1: dark green; forest type 4 (F4): light green) and non-forested conditions (grey) with a random tree distribution, soil type 3 and rough slope conditions. Note that the Y-axis is log-transformed.

5

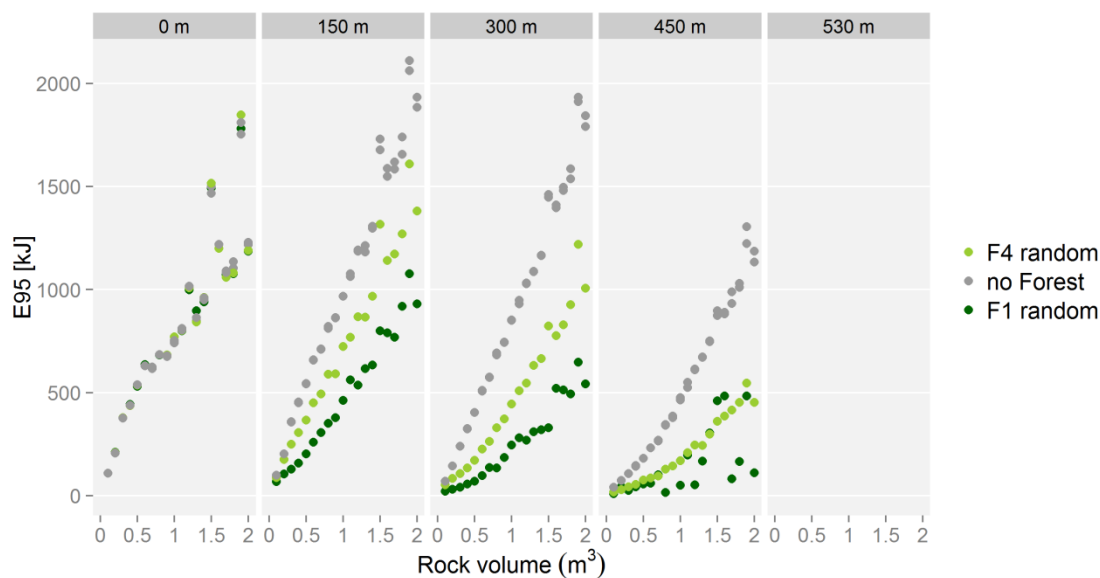
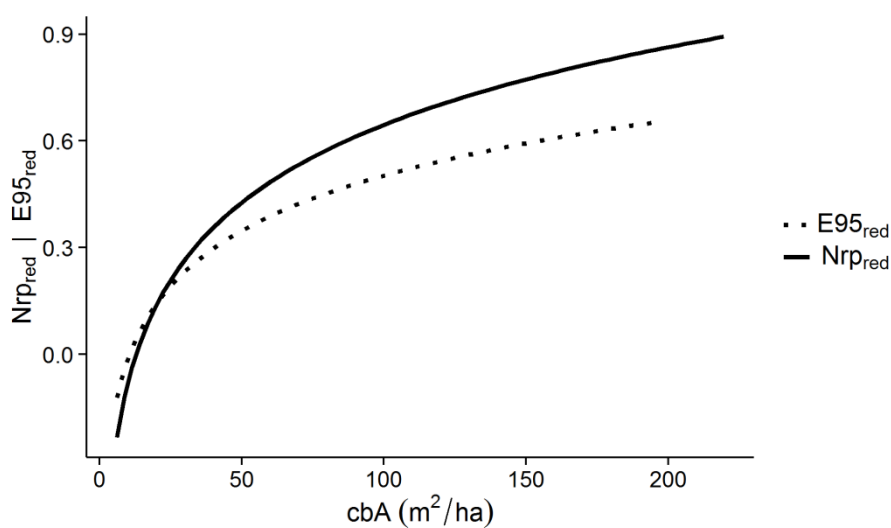


Figure 5: Illustration of the 95th percentile of maximum kinetic energies (E95) averaged over 100 simulations and at different evaluation zones located at 0-530 m downslope of the release area. As before, results include a range of rock volumes from 0.01 to 2.0 m³ under



forested (forest type 1 (F1): dark green; forest type 4 (F4): light green) and non-forested conditions (grey) and with a random tree distribution, soil type 3, and rough slope conditions.



5 **Fig. 6:** Nrp_{red} (solid) and $E95_{red}$ (dotted) based on the simulation of all forest and terrain scenarios on the concave slope and depending on cbA using a logarithmic smoothing function.

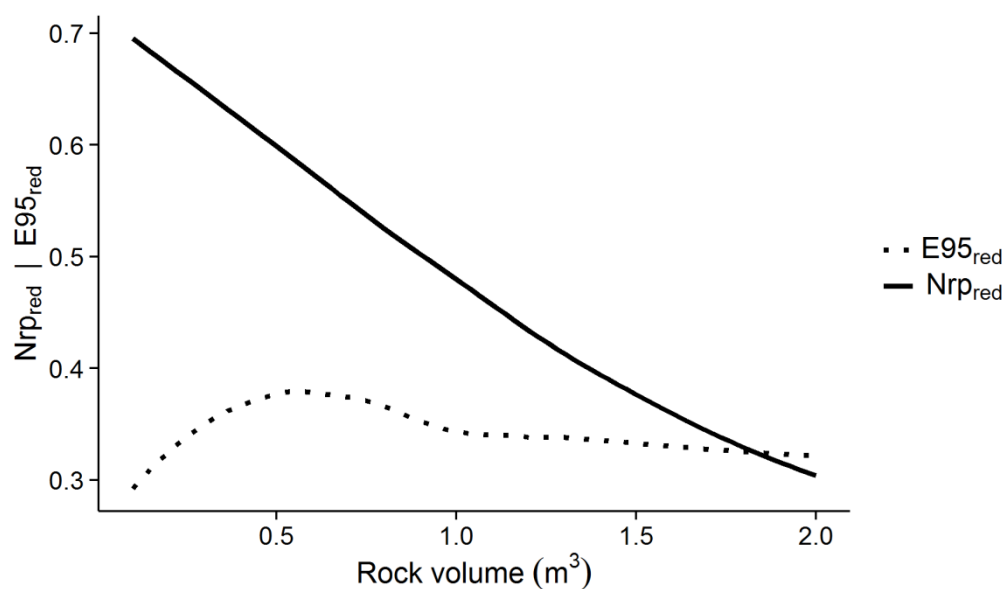




Fig. 7: Nr_{red} (solid) and $E95_{red}$ (dotted) based on the simulation of all forest and terrain scenarios on a concave slope and depending on rock volume using a “loess” smoothing function.

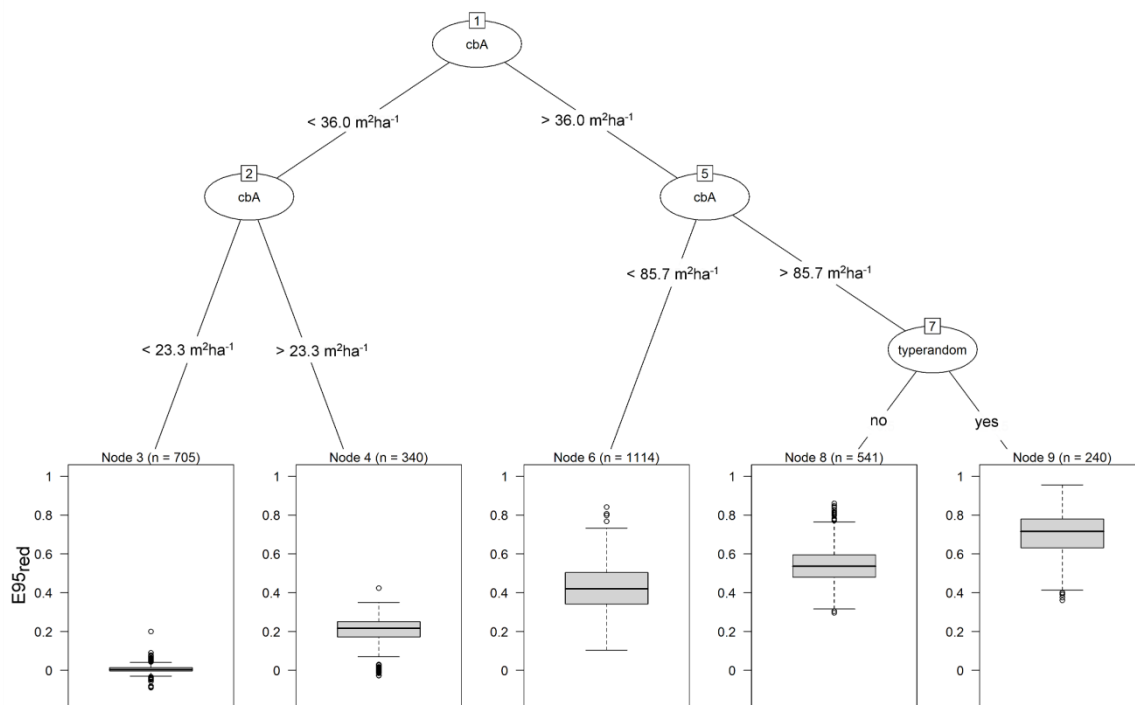
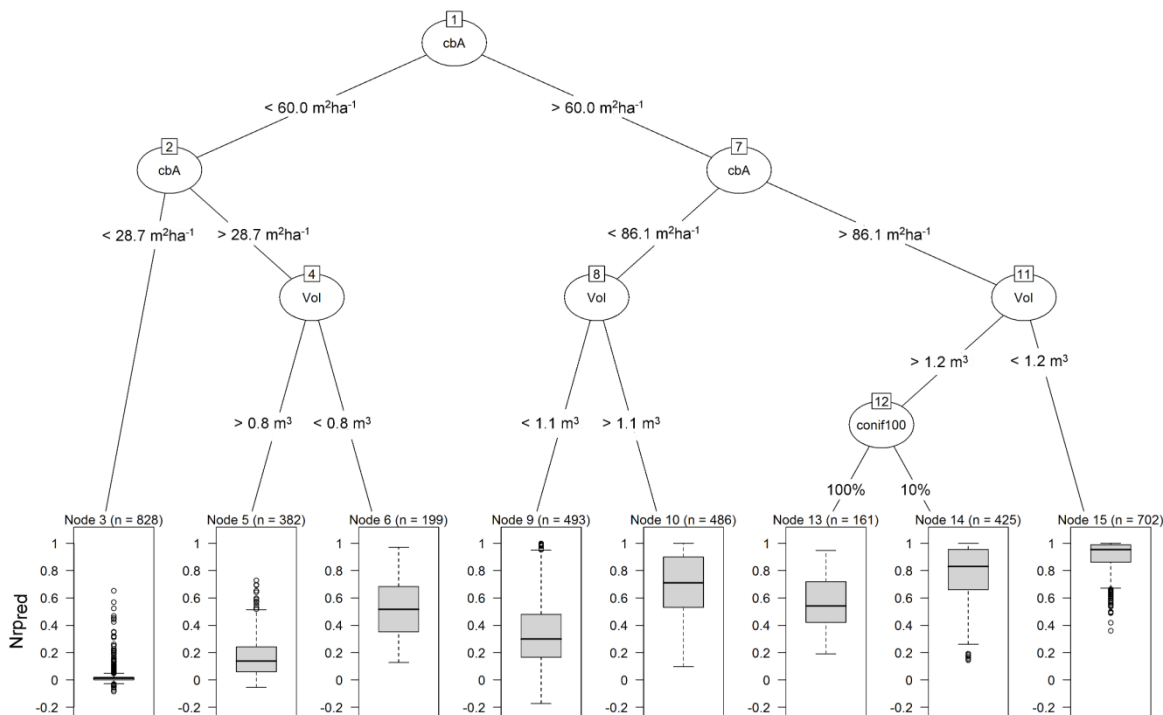




Figure 8: Regression tree models were used to predict the reduction in rockfall occurrence frequency (RT_{Freq} ; above) and the reduction in rockfall intensity (RT_{Int} ; below) by forests. The models were fitted with a training set representing 75 % of the entire dataset ($n=2940$) and by applying 3 times 10-fold cross-validation. The nodes represent the splitting variables followed by the applied threshold value. cbA = cumulative basal area [m^2ha^{-1}]; Vol = volume [m^3]; conif100 = coniferous percent [10, 100 %]; typerandom = random tree distribution [yes, no].

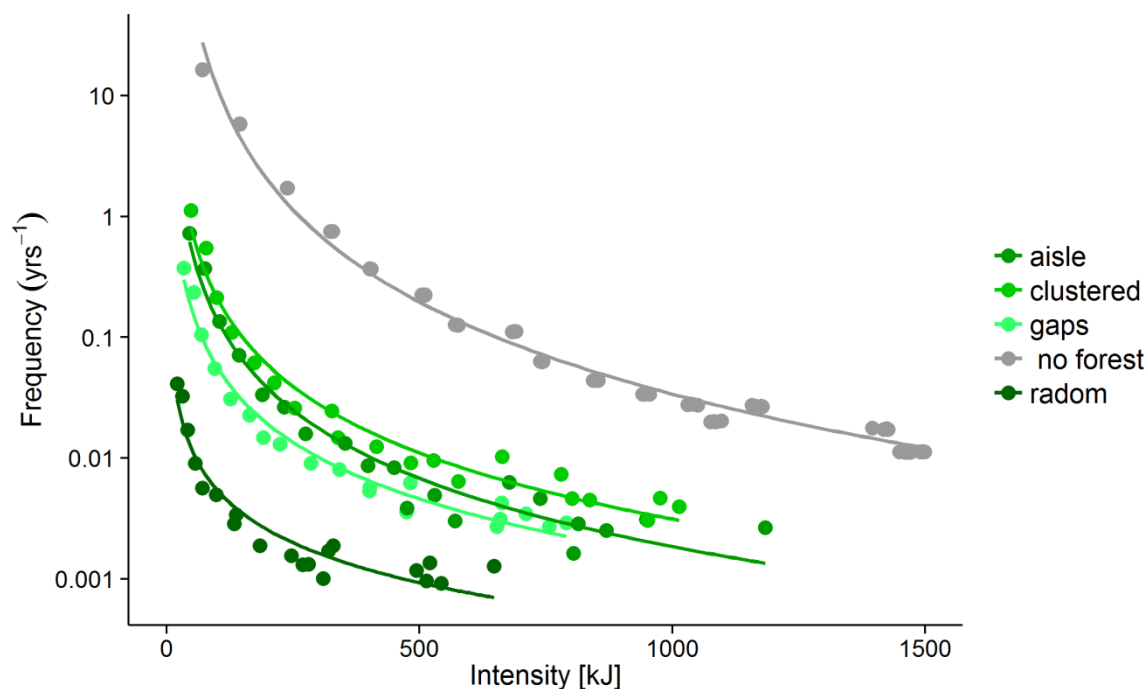


Figure 9: Frequency-intensity distributions with fitted power laws at a distance of 300 m from the release area for forest type 1 with different horizontal forest structures and without forest. The intensity is expressed as the 95th percentile of the maximum kinetic energy of rocks (E_{95}).



Tables

Table 1: Soil types and roughness used for the different simulation scenarios according to the classification of Dorren (2015). The release area and the forest road were set to no roughness and soil types 5 and 7, respectively, in all scenarios.

	Slope scenarios		Release area	Road
Soil types	soil type 3: scree $\phi < \sim 10$ cm or medium compact soil with small rock fragments	soil type 4: talus slope $\phi > \sim 10$ cm or compact soil with large rock fragments	soil type 5: bedrock with thin weathered material or soil cover	soil type 7: asphalt road
Roughness	Rough: 0.15 (10 %), 0.05 (20 %), 0.01 (70 %)	No: 0 m (100 %)	No: 0 m (100 %)	No: 0 m (100 %)

5

Table 2: Characteristics of the different forest types used for the rockfall simulations. Values have been taken from the Swiss National Forest Inventory (NFI) datasets published in Messmer (2014).

Forest type	Definition	Mean number of trees ha^{-1} (with DBH > 12 cm)	Mean DBH [cm] (DBH > 12 cm)	STD DBH	Percentage of conifers [%]
1	Fagus sylvatica 1	460	33	8.36	10
2	Pinus sylvestris-Larix decidua	304	40	10.85	100
3	Fagus sylvatica 2	200	33	8.36	10
4	Pole-stand F. Sylvatica	500	21	5.00	10

10



Table 3: Estimated regression coefficients, standard errors, Z-values (i.e. ratio of estimate and standard error), and p-values of the parametric explanatory variables of the general linear model (GLM) for the reduction in rockfall occurrence frequency by forests (GLM_{Freq}) and the GLM for the reduction in rockfall intensity (GLM_{Int}) by forests. The models were fitted with a training set representing 75 % of the entire dataset (n=2940) applying 3 times a 10-fold cross-validation. Note that R² GLM_{Freq} = 0.80 and R² GLM_{Int} = 0.84.

	Estimate		Std. Error		Z-value		p (> z)	
	GLMFreq	GLMInt	GLMFreq	GLMInt	GLMFreq	GLMInt	GLMFreq	GLMInt
(Intercept)	-0.58	-0.57	0.014	0.009	-39.23	-62.30	<2*10-16	<2*10-16
Vol	-0.21	-0.02	0.005	0.003	-44.18	-5.23	<2*10-16	1.8*10-7
log(cbA)	0.33	0.23	0.003	0.002	109.31	119.61	<2*10-16	<2*10-16
type clustered	-0.1	-0.03	0.007	0.005	-12.03	-5.99	<2*10-16	2.4*10-9
type gaps	0.06	0.08	0.008	0.005	6.95	15.58	4.3*10-12	<2*10-16
type random	0.06	0.09	0.008	0.005	7.66	17.72	2.8*10-14	<2*10-16
soil type 4	-0.05		0.006		-7.72		1.5*10-14	
Roughness 2	-0.09	-0.02	0.006	0.004	-13.63	-6.18	<2*10-16	7.2*10-10
Conifer percent 100	-0.10	-0.09	0.008	0.005	-14.26	-17.54	<2*10-16	<2*10-16

5

Table 4: Normalized Root Mean Squared Error (nRMSE) of the generalized linear models (GLM) and the regression tree models (RT) predicting Nrp_{red} (GLM_{Freq}, RT_{Freq}) and E95_{red} (GLM_{Int}, RT_{Int}) with 3 times 10-fold cross-validation (cv) and for predictions of the test dataset (25 % of the data) and linear slopes with varying slope angle (slope 2-5).

Model	nRMSE cv	nRMSE test	nRMSE slope 2 (32°)	nRMSE slope 3 (35°)	nRMSE slope 4 (38°)	nRMSE slope 5 (40°)
-------	----------	---------------	------------------------	------------------------	------------------------	------------------------

10



GLM_{Freq}	14 %	16 %	14 %	15 %	14 %	15 %
RT_{Freq}	14 %	17 %	15 %	16 %	10 %	8 %
GLM_{E95}	10 %	9 %	14 %	13 %	16 %	18 %
RT_{E95}	10 %	9 %	10 %	7 %	9 %	11 %

Table 5: Model input parameters and predicted values of $N_{p_{red}}$ with the GLM and the RT model as well as the measured value for $N_{p_{red}}$ for the study site in Vaujany where (Dorren et al., 2006) performed real-size rockfall experiments.

Position	Vol [m ³]	cbA [m ² ha ⁻¹]	Forest type	Soil type	Roughness	$N_{p_{red}}$ (true)	$N_{p_{red}}$ (pred, GLM)	$N_{p_{red}}$ (pred, RT)
Middle slope	0.5	70.5	Random	4	Rough	0.64	0.62	0.70
Bottom slope	0.5	102.4	random	4	rough	1.0	0.75	0.91

5

Table 6: α and β coefficient and adjusted R^2 of the fitted power-laws of the frequency-intensity distributions at a distance of 300 m from the release area for forest type 1 with different horizontal forest structures and without forest.

Forest structure	α	β	R^2
No forest	14.11	2.53	0.99
Random	-0.17	1.09	0.93
Clustered	6.82	1.82	0.97
Aisle	6.64	1.87	0.98
Gaps	4.29	1.56	0.96

10