

1 The catastrophic landfill flowslide at Hongao dumpsite on December 20, 2015 in Shenzhen, China

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6 **Abstract** A catastrophic flowslide occurred at Hongao dumpsite on Dec 20, 2015 in the Guangming New
7 District of Shenzhen, China. The flowslide caused 73 casualties with 4 missing and damaged 33 buildings.
8 In the absence of extreme weather condition and seismic activity, the causes of the failure were analyzed
9 on the basis of multi-temporal remote sensing images, site investigation, *in-situ* tests, laboratory tests, and
10 numerical analyses. Site investigations showed that the volume of the flowslide deposit was $2.32 \times 10^6 \text{ m}^3$
11 and the volume of dumpsite filling was $6.27 \times 10^6 \text{ m}^3$ at the time of the event, which is 3 times greater than
12 the design capacity. The flowslide was characterized by high travel velocity and long runout distance. The
13 displaced material was primarily silty soil in mixture of construction and demolition waste at high water
14 content (17.3%-42.4%). The primary causes of the failure were concluded as follow: (1) Groundwater flow
15 was stagnated in the dumpsite due to malfunctioned drainage system and the underlying impermeable
16 granite stratum; (2) The accumulation rate and total volume of the waste filling was in exceedance of the
17 design capacity. The failure may be ascribed to the development of excess pore water pressure as evidences
18 of liquefaction were observed at several locations, and it is postulated that such phenomena was related to
19 the surcharge loads imposed by the unregulated disposal activities.

20 **Keywords** Flowslide, Landslide, Municipal Solid Waste (MSW), Dumpsite, Landfill

21 1 Introduction

22 Owing to the population growth and industrialization, rapid increase in the accumulation rate of municipal
23 solid waste (MSW) poses challenges in MSW management and urban planning (Huang and Cheng, 2016).
24 As landfilling is the most common method of MSW disposal (Brunner and Fellner, 2007), transforming an
25 abandoned quarry to a dumpsite conserves resources by reclaiming the quarry space and provides MSW
26 storage solution (Zou, 2016). The landfill slope stability is critical to the MSW management, and thus the
27 selection and design of landfills require engineering assessment on both slope stability and environment
28 impact.

29 The porosity and water content of MSW is typically high in an unregulated landfill because of inadequate
30 drainage system, and therefore some failures of MSWs exhibited fluid flow behave, i.e. flowslide, with
31 extremely high mobility (e.g. Dai et al., 2016; Huang and Cheng, 2016). Landfill slope failure in MSW
32 dumpsite can be found in previous studies, wherein 6 reported cases between 1993 and 2005 had resulted
33 in approximately 500 deaths and significant economic loss (Blight, 2008; Blight and Fourie, 2005; Eid et
34 al., 2000; Kjeldsen and Fischer, 1995; Kocasoy and Curi, 1995; Merry et al., 2005; Mitchell et al., 1990).
35 The deadliest event in history killed 278 people in Manila, Philippines in 2000, and the second deadliest
36 event buried 71 houses and killed 143 people on Feb 21, 2005 at Leuwigajah dumpsite near Bandung,
37 Indonesia (Lavigne et al., 2014).

38 Shear strength of MSW is a function of various parameters, including the type and composition of the waste,
39 disposal rate, water content, surcharge, and compaction (Eid et al., 2000; Huvaj-Sarihan and Stark, 2008).
40 The movement of waste failure are complex and still poorly known due to the lack of field monitoring data.
41 Previous studies have focused on empirical methods (e.g. Blight and Fourie 2005; Srour 2011), laboratory
42 experiment, and numerical analyses. An energy balance approach was presented to estimate the maximum
43 flow velocity of MSW slope failure (Blight and Fourie, 2005). The empirical-statistical method still has
44 widespread use in practical applications, but the accuracy is commonly model dependent (Dai et al., 2016;
45 Huang and Cheng, 2016). The mixture of peat, kaolinite, and quartz sand were used in laboratory tests to
46 study the interplay between water content of filling and failure of MSW (Dai et al., 2016). With an increased
47 water content in the MSW, the mobility increases with the maximum and final thickness of deposit decrease.

48 Numerical simulation has been widely used in landfill slope stability analysis (Chang, 2002, 2005; Chugh
49 et al., 2007; Huang and Cheng, 2016), and some of the more advanced methods were recently employed,
50 such as the smoothed Particle Hydrodynamics (SPH) method for runout distance analysis of two typical
51 landfill flow slides occurred in Sarajevo and Bandung (Huang et al., 2013).

52 A MSW landfill consists of mainly construction and demolition waste failed at the Hongao dumpsite in the
53 Guangming New District of Shenzhen, China at approximately 11:40 AM on Dec 20, 2015. The flowslide
54 destroyed 33 buildings and accommodations in the industrial park, and resulted in 74 casualties with 4
55 missing. No extreme weather conditions or seismic activity were reported at the time of the event. This
56 article investigates the characteristics and causes of failure. Field investigation, *in-situ* tests, Unmanned
57 Aerial Vehicle (UAV) stereo-measurements, and remote sensing images were used to analyze the material
58 properties and characteristics of the failure. Laboratory tests and numerical analysis were performed to
59 better understand the failure mechanism.

60 **2 Data and Methodology**

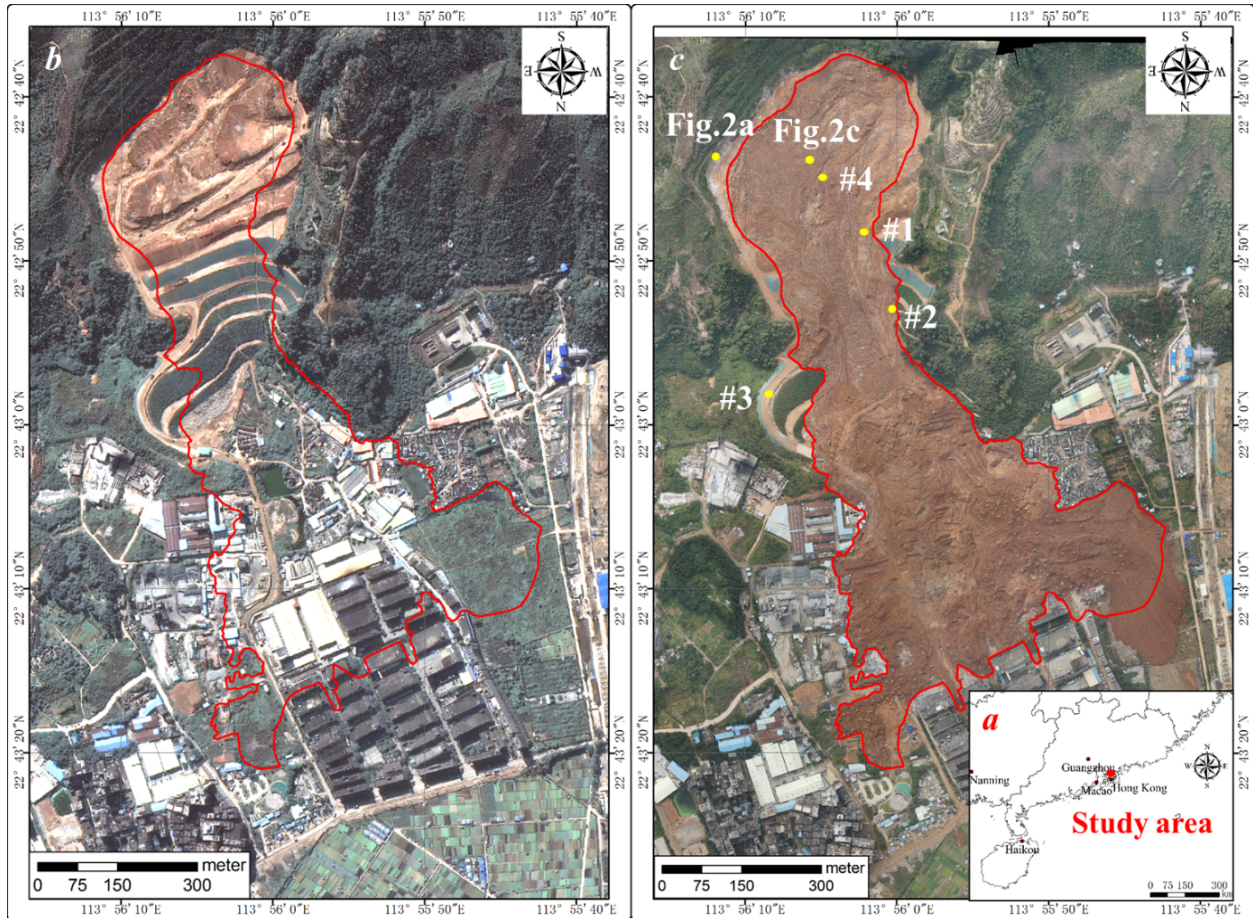
61 A series of satellite images of the site was obtained from Google Earth between Nov 2002 and Feb 2016.
62 Two high resolution pre-/post-failure images were acquired by satellite image and aerial photo. Topography
63 map of the pre-disposal site was provided by the local government, of which the design was submitted by
64 a third-party consulting firm in Dec 2013. Based on the official report on the incident, the design capacity
65 was $4 \times 10^6 \text{ m}^3$ with 9 slope benches at an average ratio of 1:2.5. Field and UAV measurement indicate that
66 the total volume of dumpsite filling was $6.27 \times 10^6 \text{ m}^3$ before failure ($5.83 \times 10^6 \text{ m}^3$ by Yin et al., (2016)). The
67 Digital Orthoimage Map (DOM) and Digital Surface Model (DSM) were reconstructed from aerial photos,
68 and a topographic map for the post-sliding dumpsite was generated from the DSM at the scale of 1:1000.
69 Structured terrains in the study site were excavated for vehicles which consisted 9 benches and 3 platforms.
70 Topographic map for the pre-sliding dumpsite was derived by combining the original, pre-sliding, and post-
71 sliding slope images, which also provided an estimation on the volume and depth of the pre-/post-sliding
72 landfill.

73 The preliminary field investigation was conducted on Dec 23, 2015, with a follow-up investigation on Jan
74 21-25, 2016. Field investigations evaluated the topographical, geological, and groundwater conditions.
75 Density, water content, and permeability of the displaced material was measured *in-situ*.

76 Undisturbed and reconstituted soil samples were obtained and used for laboratory tests, including weathered
77 silty soil and waste fillings. Grain size distribution was conducted using wet sieve analysis. Maximum dry
78 density and optimum water content of the displaced material were determined by standard compaction test.
79 The shear parameters of the displaced material (c and ϕ) were obtained by triaxial compression test. In order
80 to back-analyze the failure of dumpsite, input parameters for the material properties and model geometry
81 were measured on site or in the laboratory. The mobilized shear strength was back-calculated using the
82 conventional limit equilibrium method (LEM) to demonstrate the complexity of such progressive failure
83 involving liquefaction and post-failure landslide propagation.

84 **3 Geological and Climatic Setting**

85 The location of the site is shown in Fig. 1a, of which the pre-failure image was obtained by Pleiades image
86 on Dec 18, 2015 (2 days before failure) and the post-failure aerial image was taken by UAV (3 days after
87 failure) as shown in Fig. 1b and 1c. The dumpsite was located 23 km north of Shenzhen and 5 km south of
88 the Guangming New District (E113°56'5", N22°42'44"). The dumpsite was surrounded by three ridges with
89 a free face excavated in a mild slope gradient. The bedrock of the dumpsite is mainly Cretaceous granite
90 rock (Fig. 2a and 2b). The landfill consisted of construction waste with silty soil, clay, rock, and gravel (Fig.
91 2 and 3).



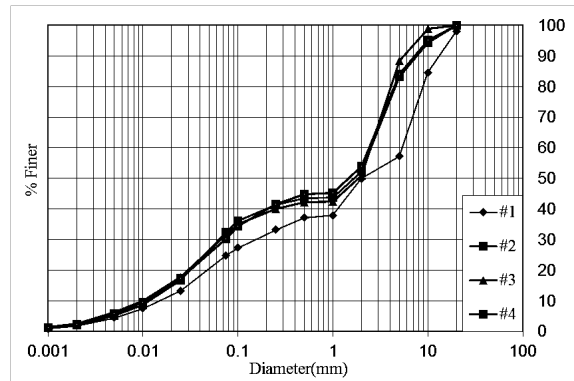
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 93 Fig. 1 Location of the study site and images of the pre-/post-failure dumpsite. The boundary of the flowslide
 94 is indicated in red. **a** Location of the study area; **b** Pleiades satellite image (Dec 18, 2015); **c** Aerial photo
 95 of the flowslide with locations of samples for laboratory test (Dec 23, 2015).

96 The study area belongs to the zone of subtropical monsoon climate with an average temperature of 22° and
 97 an average annual precipitation of 1500 mm concentrating between April and September (> 85 % of the
 98 annual rainfall) (Zhang et al., 2006). Based on the rainfall data from the nearby Tangjia Rainfall Station
 99 (Jan 1, 2013 - Dec 20, 2015) as shown in Fig. 4, a heavy rainfall occurred on Dec 9, 2015 with accumulated
 100 precipitation of 67.8 mm. The area of catchment (Fig. 5a) was estimated as $4.7 \times 10^5 \text{ m}^2$, which was nearly
 101 3 times greater than the area of the dumpsite ($1.6 \times 10^5 \text{ m}^2$). Drainage system was implemented to collect
 102 surface run-off by the drainage pipes installed at a higher elevation above the dumpsite and diverted into
 103 the peripheral drainage channels. Field evidences suggested that the drainage system was abandoned,
 104 possibly due to the lack of maintenance (Fig. 5b and 5c). With the inadequate drainage, ingress of rainwater
 105 in the dumpsite was permitted (Fig. 6c), in addition to the concentration of surface run-off and groundwater
 106 into the dumpsite. The infiltration rate of the waste filling was estimated as $5.26 \times 10^{-6} \text{ cm/s}$ by field double
 107 ring infiltrometer. Low permeability and inadequate drainage resulted in high water content and high pore
 108 water pressure in the dumpsite. The groundwater in the study area were mainly bedrock fissure water and
 109 Quaternary pore water (Fig. 6).



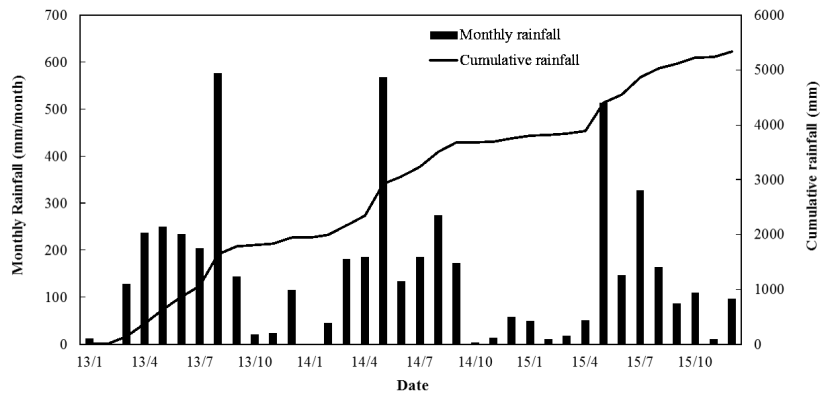
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111 Fig. 2 Photos of the rock and soil specimen collected from the displaced material. **a** Exposed granite rock
 112 near the flowslide scarp; **b** Close view of the granite rocks; **c** Displaced material in the zone of depletion;
 113 **d** Materials in the flowslide source area.



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115 Fig. 3 Particle size distribution of the waste filling (with sieve sizes of 20, 10, 5, 2.0, 1, 0.5, 0.25, 0.10,
 116 0.075, 0.025, 0.01, 0.005, 0.002, and 0.001 mm)



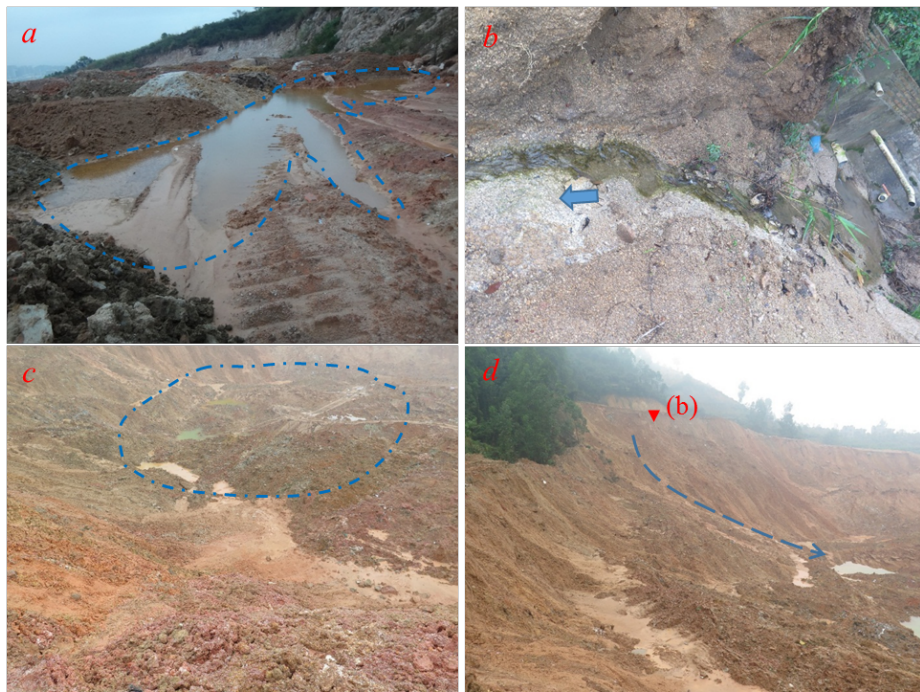
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118 Fig. 4 Monthly rainfall and cumulative precipitation between 2013 and Dec 2015.



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Fig. 5 **a** An overview of the catchment area (Google Earth image); **b** The surrounding peripheral drainage channel of the dumpsite was abandoned; **c** Damaged drainage pipes were not repaired and failed to divert surface run-off into the peripheral drainage channel. Surface run-off concentrated in the waste filling.



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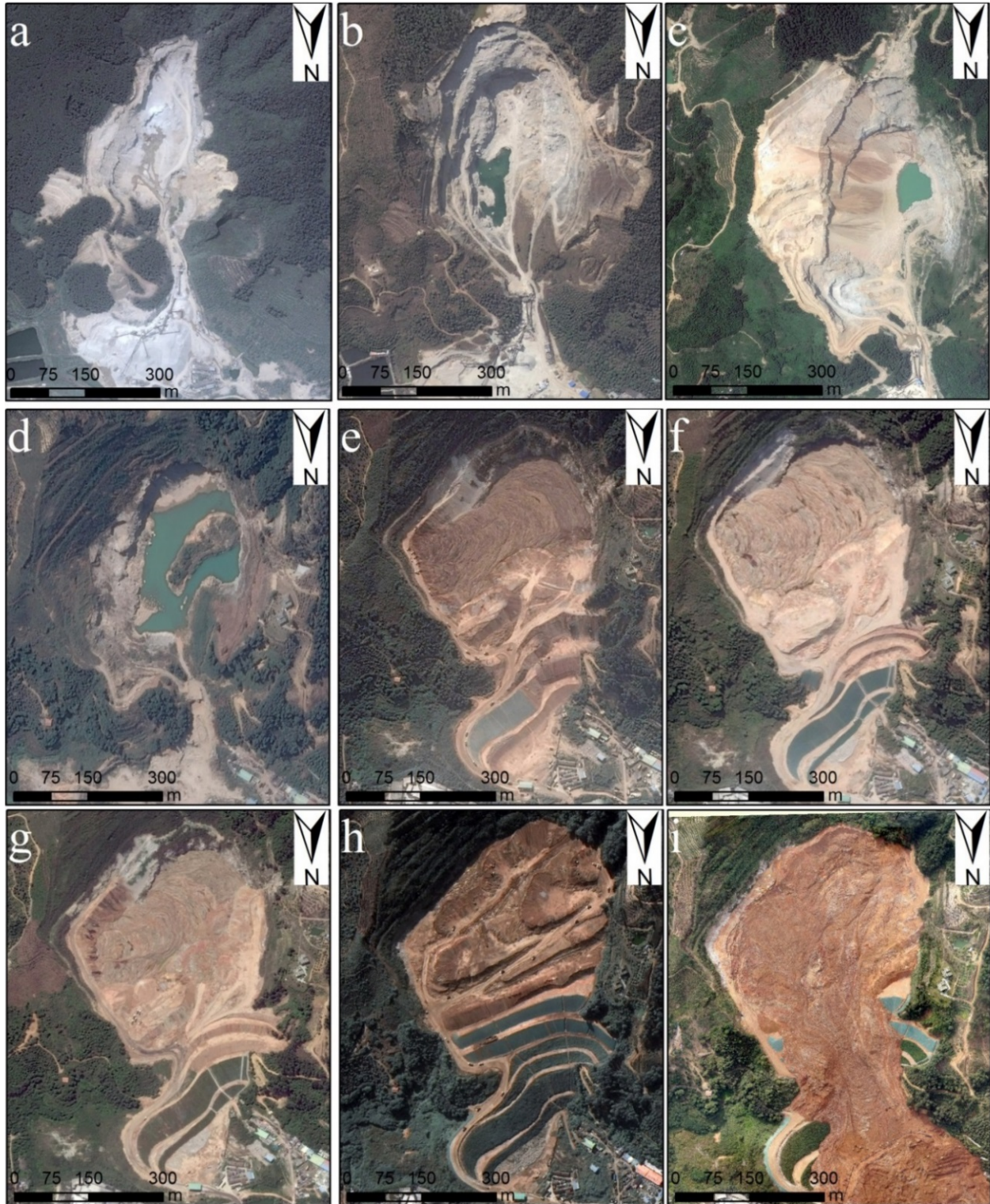
Fig. 6 **a** Groundwater overflow in the rock fissures near the flowslide crown; **b** Water accumulation in the flowslide crown; **c** Water accumulation in the source area; **d** Surface run-off and infiltration contributed to groundwater seepage.

127 **4 Multi-Temporal Remote Sensing Images**

128 The development of dumpsite and landform changes are presented by a series remote sensing images taken
129 between 2002 and 2016 (Fig. 7). It is postulated that the quarry was active between 2002 and 2008. However,
130 formation of ponds of various sizes at different spatial and temporal locations in the quarry indicated low
131 permeability and the lack of drainage network (Fig. 7b, 7c, and 7d). The quarry was abandoned in or before
132 2008 as shown in Fig. 7c (Feb 20, 2008) with a small amount of waste filling in the pit. A large-scale pond
133 was later formed due to groundwater and rainfall accumulation (Fig. 7d). The depth of the pit was over 100
134 m before the routine disposal activities. A service road was excavated between two rock hills exiting the
135 quarry and formed a small-scale gully cross the pit longitudinally (Fig. 7a, 7b, 7c, and 7d). Additionally, a
136 small-scale platform (the office area of the abandoned quarry) was excavated at the immediate downstream
137 of the quarry exit. The gully was covered by the construction of a multi-benched retaining slope between
138 the rock hills as waste disposal continued throughout the process (Fig. 7e). Disposal activity started possibly
139 in 2014 (Fig. 7e and 7f), and the accumulation rate of the waste filling was considerably fast (Zou, 2016).

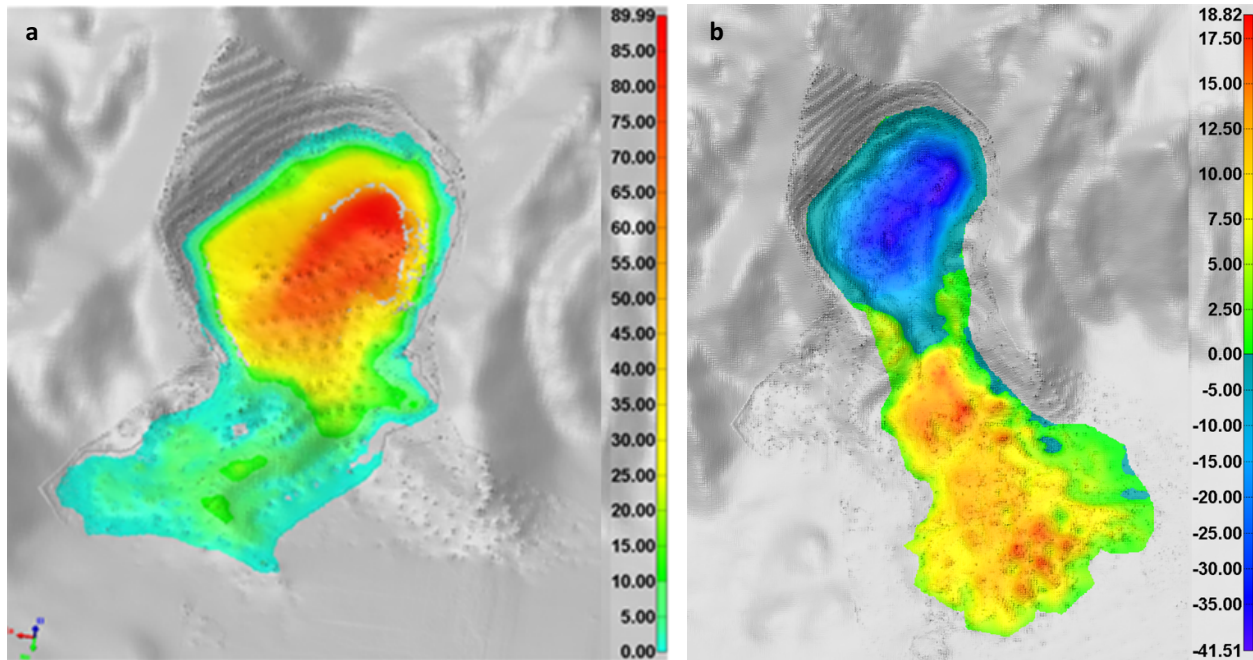
140 An unpaved road was excavated on the east of the quarry connecting the crest and toe of the hillslope as
141 shown in Fig. 7e. A large number of trucks can be seen transporting waste filling to the dumpsite. A large
142 amount of construction waste was dumped in the pit with 4 slope benches built at the exit of the pit (Fig.
143 7e). The volume of the waste increased significantly with the 1st and 2nd benches completed in late 2014,
144 and the 3rd to 4th benches were still under construction until Jan 2015 (Fig. 7f).

145 The rapid accumulation and the total volume of the waste filling received some attentions before failure.
146 An environmental assessment provided by a third-party consulting firm warned the erosion at the site and
147 its influence on the slope stability in Jan 2015 (Zou, 2016). The disposal activity was ceased for a while as
148 no trucks were seen in the image (Fig. 7f), and it was verified by the interview with the local inhabitants.
149 The disposal activity was resumed in or before April 2015, and the 4th bench was completed by then (Fig.
150 7g). By comparing Fig. 7g and Fig. 7h, significant modifications on the landform occurred with a major
151 increase in the volume of the waste filling as the landfill was close to the crest of the pit. The waste filling
152 appeared flattened and disposal activity was intense as more than 20 trucks were found in Fig. 7h. A total
153 of 8 slope benches were completed with surface drainage channels installed on the hillslope, and the 9th
154 bench was still under construction before failure. The thickness of the waste filling was around 90 m with
155 a volume of $6.3 \times 10^6 \text{ m}^3$ (Fig. 8a) by extracting the difference between the pre- and post-filling DEMs.



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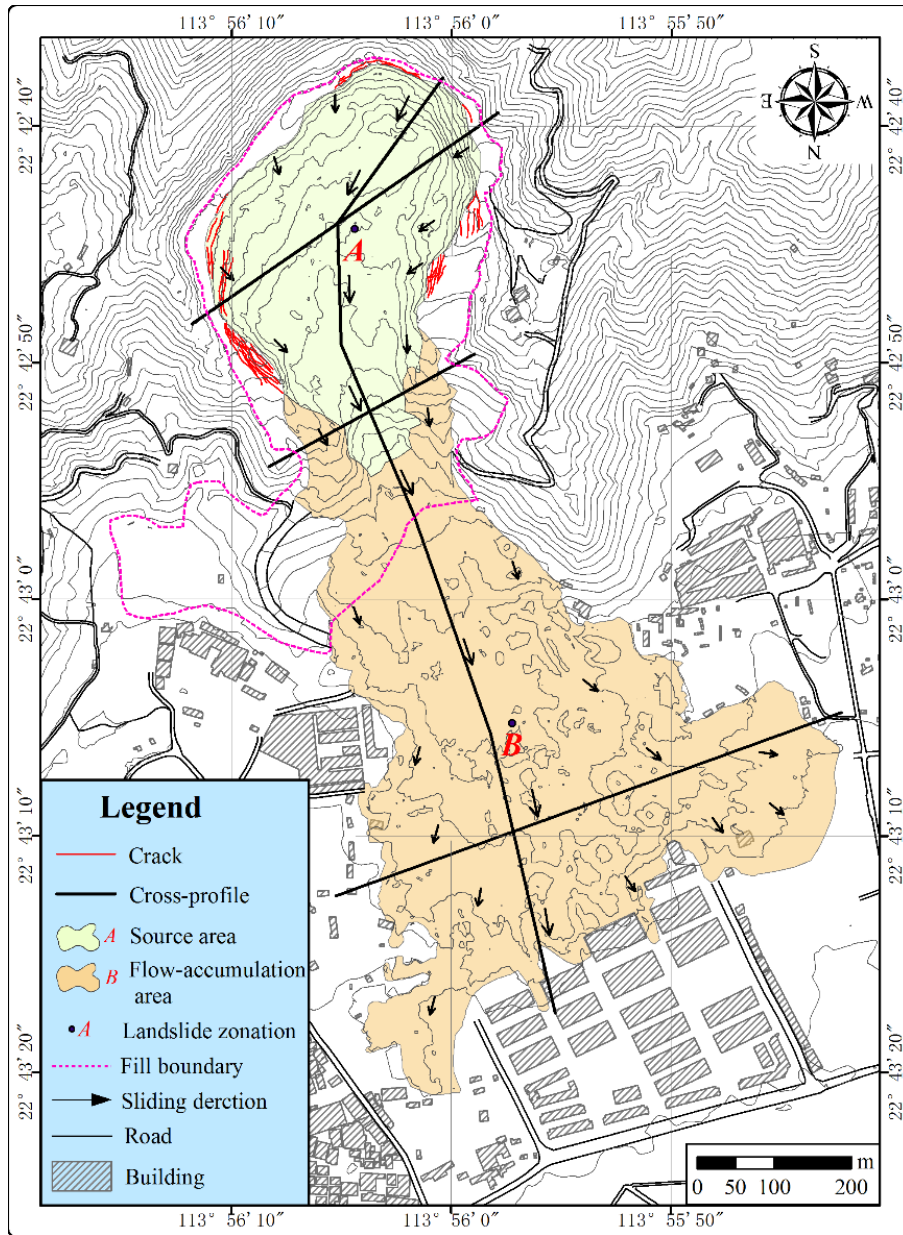
157 Fig. 7 Multi-temporal remote sensing images of the study area. **a** 08/31/2002; **b** 02/20/2008; **c** 08/30/2010;
 158 **d** 11/25/2013; **e** 11/17/2014; **f** 01/23/2015; **g** 04/14/2015; **h** Pleiades (12/18/2015); **i** Aerial (12/23/2015).
 159 Image **a** to **g** were obtained from Google Earth. A service road was excavated at the exit of the quarry with
 160 a width of nearly 70 m (see image **a**, **b**, **c**, and **d**).



161
 162 Fig. 8 **a** The thickness distribution of the dumpsite based on the pre-/post-disposal DEMs; **b** The elevation
 163 variations of the dumpsite estimated between the pre-/post-sliding DEMs.

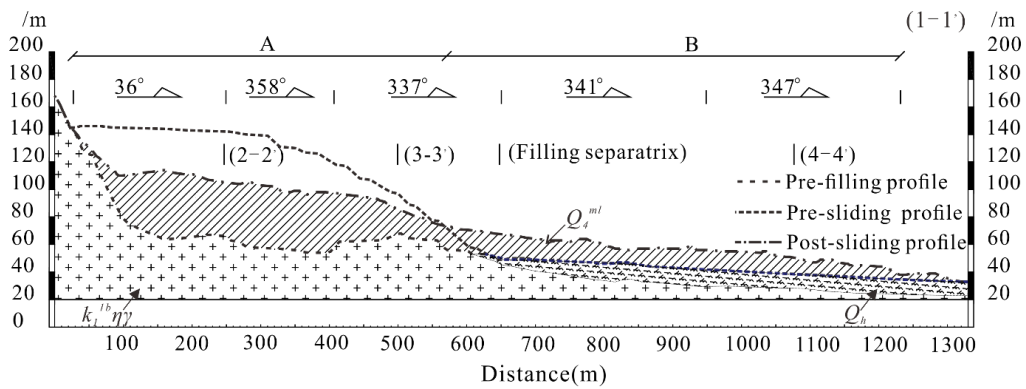
164 **5 Flowslide Characteristics**

165 The flowslide area can be divided into the source area and the flow-accumulation area as shown in Fig. 9.
 166 The source area was the abandoned quarry pit. The mass slid in the direction of 340° with the maximum
 167 traveling distance of 1203 m. The maximum thickness of the remaining material in the source area was
 168 41.51 m with an average of 20.5 m (Fig. 9 and Fig. 10). The maximum deposit thickness in the flow-
 169 accumulation area was 18.2 m with an average of 8.21 m (Fig. 8b). The geometry of a flowslide can be
 170 expressed in length (L), height (H), width (W), and area (S) (Legros, 2002; Scheidegger, 1973). The
 171 geometric parameters are indicated in the simplified flowslide geometry in Fig. 11 with values tabulated in
 172 Table 1, Table 2 and Table 3.



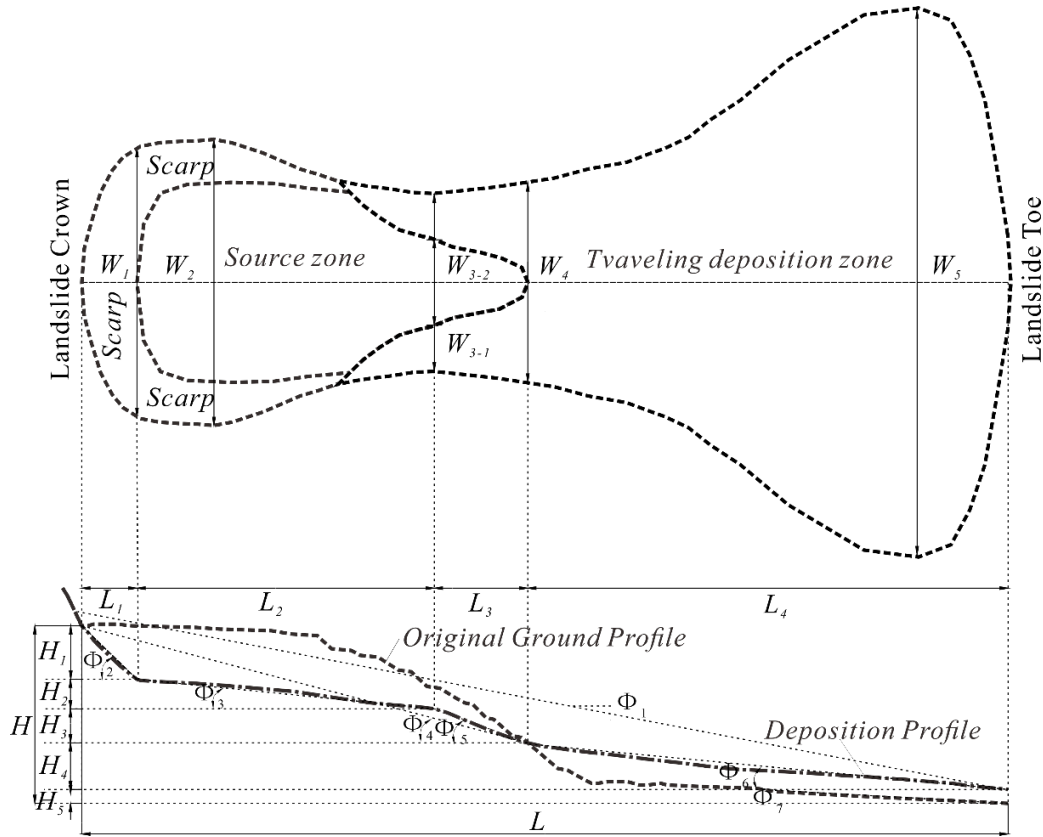
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Fig. 9 The topographic map of the flowslide.



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Fig. 10 The geological and topographical cross section (profile line 1-1')



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 178 Fig. 11 Simplified illustration of the flowslide geometry. L: runout distance; H: elevation; W₁: scarp width;
 179 W₂: max width of the source area; W₃: frontal width of the source area; W₄: width of the shear crack; W₅:
 180 max width of the flow-accumulation area; L₁: horizontal length of the scarp; L₂: horizontal length of the
 181 mild slope; L₃: horizontal length of the steep slope; L₄: horizontal length of the flow-accumulation area; H₁:
 182 height of the scarp; H₂: height of the mild slope; H₃: height of the steep slope; H₄: height of the flow-
 183 accumulation area; H₅: thickness of the toe of the flowslide; Φ₁: extension angle; Φ₂: slope gradient of the
 184 scarp; Φ₃: mild slope gradient; Φ₄: slope gradient of the source area; Φ₅: steep slope gradient; Φ₆: slope
 185 gradient of the flow-accumulation area; Φ₇: slope gradient of the foundation area

186 Table 1 Geometric parameters of the flowslide (Length and Width)

Parameter	L	H	W ₁	W ₂	W ₃₋₁	W ₃₋₂	W ₄₋₁	W ₄₋₂	W ₅
	(m)	(m)	(m)	(m)	(m)	(m)	(m)	(m)	(m)
Value	1204.67	111.31	212.5	399.32	218.31	149.38	217.71	64.63	592.25

187 Table 2 Geometric parameters of the flowslide (Angle)

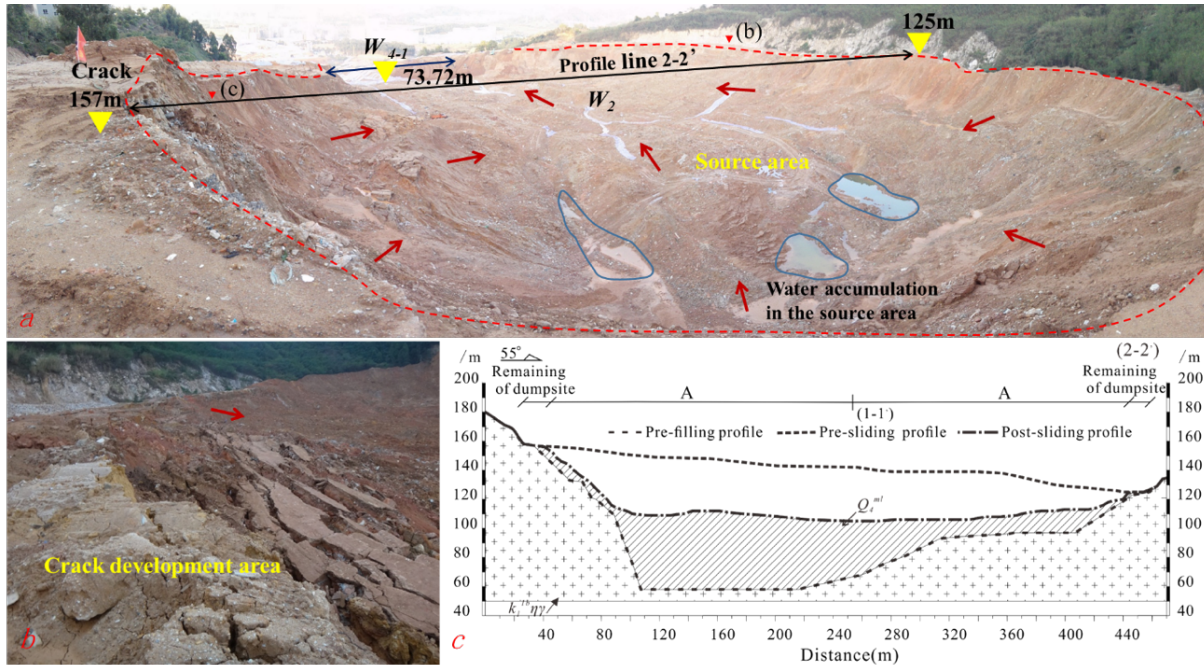
Parameter	Φ ₁	Φ ₂	Φ ₃	Φ ₄	Φ ₅	Φ ₆	Φ ₇	K	S
	(°)	(°)	(°)	(°)	(°)	(°)	(°)		(m ²)
Value	5.28	28.06	1.2	7.62	11.36	3.08	1.36	0.32	398619.6

188 Table 3 Geometric parameters of the flowslide (Area and Thickness)

Parameter	Area A	Area B
S (m ²)	11.35×10 ⁴	28.51×10 ⁴
V (m ³)	3.95×10 ⁶ (remaining)	2.34×10 ⁶
T _{max} (m)	41.51	18.82
T _{ave} (m)	20.49	8.21

189 **5.1 Source Area**

190 The source area of the flowslide is the quarry pit with a length (L_{1+2+3}) of 540.30 m and a width (W_2) of
 191 399.32 m. The area of the source area (S_A) is $11.35 \times 10^4 \text{ m}^2$ and the height (H_{1+2+3}) was 68.5 m with an
 192 apparent dip (Φ_4) of 7.62° (Fig. 12 and Table 1). The volume of the mobilized material from the source
 193 area was $2.32 \times 10^6 \text{ m}^3$ and the remaining volume was $3.95 \times 10^6 \text{ m}^3$. The maximum thickness of the source
 194 area ($T_{A-\text{max}}$) was 41.5 m with an average thickness ($T_{A-\text{ave}}$) of 20.5 m. The west of the main scarp has steep
 195 slope gradient with a mild gradient on the east. The height of the steep scarp was 25-47 m, while the height
 196 of the mild scarp was 10-20 m. The geological cross section (profile line 2-2') is shown in Fig. 12c. The
 197 maximum thickness of the source area was 40.65 m with an average of 34.83 m. The overview of the source
 198 area and surface cracks are shown in Fig. 12a and 12b, respectively.

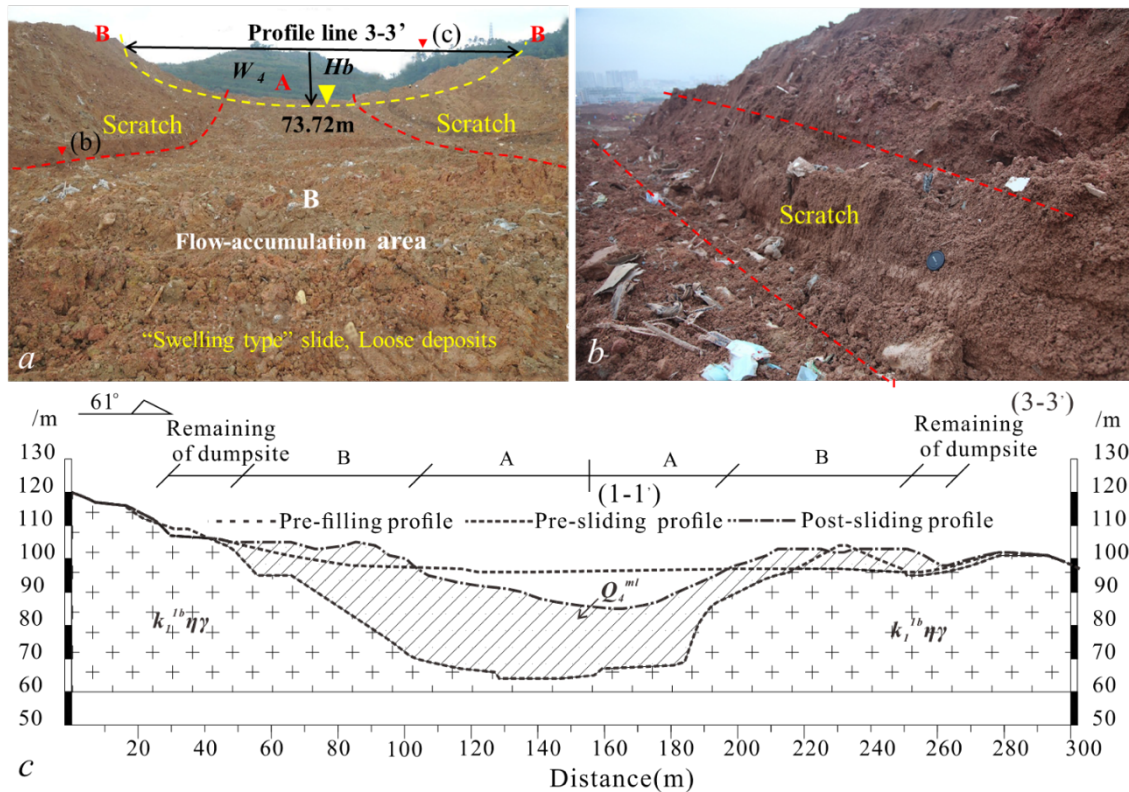


199 Fig. 12 **a** An overview of the source area (image taken at the flowslide crown facing north). A large opening
 200 (W_{4-1}) formed due to the failure of the retaining slope (the exit of the original quarry pit), which provided
 201 passage for the flowslide and permitted rapid release of kinematic energy; **b** Surface cracks were developed
 202 due to unloading near the crest and the flanks of the flowslide; **c** The geological cross section (profile line
 203 2-2') of the source area.
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205 Pore water pressure increases as rainwater and groundwater ingress in the dumpsite. The lack of drainage
 206 resulted in waste filling saturation in the basal zone of the dumpsite which appeared as the final sliding bed
 207 of the flowslide (Fig. 12a). A large amount of silty soil was observed in the source area. The water content
 208 of the remaining materials in the source area is 17.3%-42.4% (6 sampling locations). The materials on both
 209 sides of the source area were mobilized due to the debutting effect as the waste filling in the lower
 210 portion of the dumpsite slid into the downstream industrial park, and in consequence, caused collapse of
 211 the dumpsite. Step-like steep scarp was formed on the west of the main scarp with tensile cracks developed
 212 on the rear edge as well as both sides of the scarp.

213 The presence of the aforementioned rock ridges formed unfavorable topography for the stability of the
 214 retaining slope as it promoted a narrow gully for groundwater flow (Fig. 7c and Fig. 9). As a result of the
 215 retaining slope failure, a large opening was formed at the elevation of 73.7 m (between the front edge of
 216 the steep scarp and the rear edge of the flow-accumulation area) as shown in Fig. 9, Fig. 10, and Fig. 12a.
 217 The displaced material was stratified in the middle part of the flow-accumulation area during the high-speed
 218 sliding. The shape of the failed retaining slope was half-elliptic-like with a width (W_{4-1}) of 217.7 m and a

219 height (H_b) of 27.5 m (Fig. 13a and 13c).

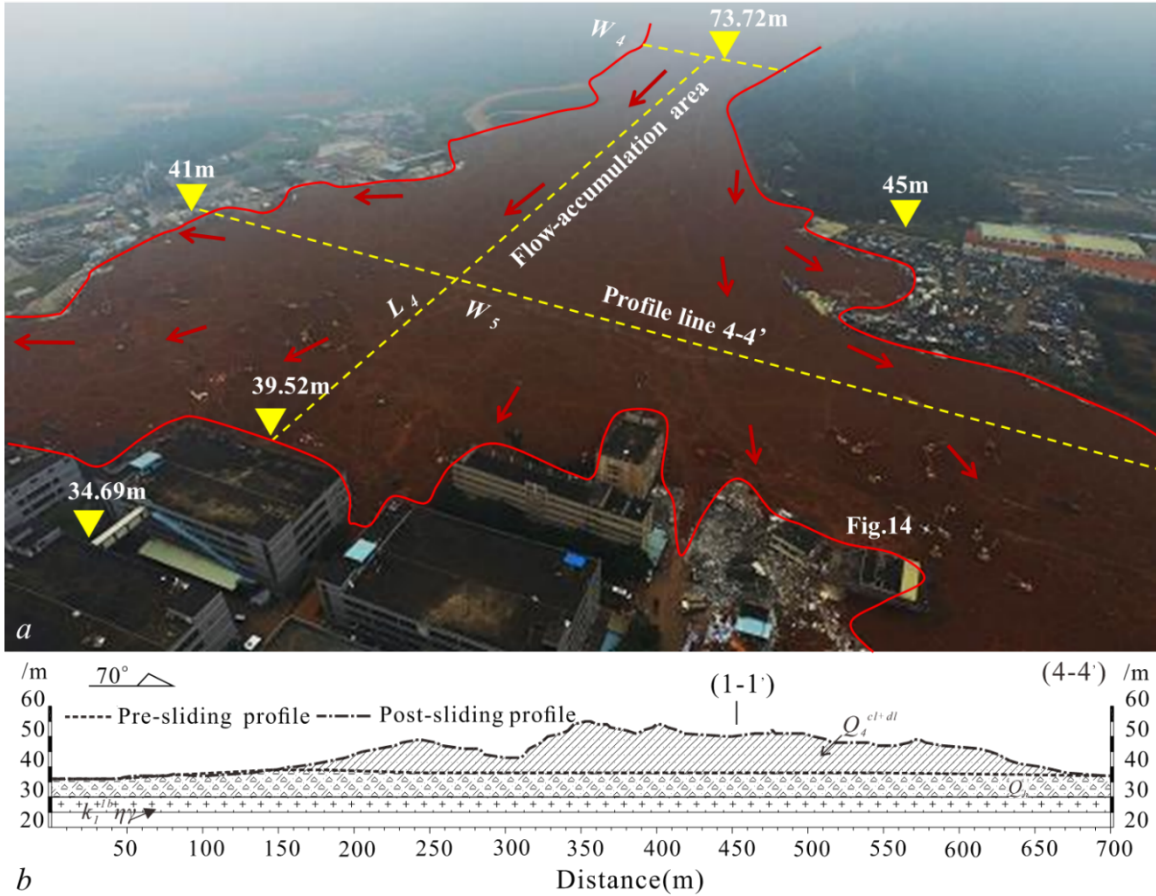


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221 Fig. 13 **a** The half-elliptic-like opening of the failed retaining slope; **b** Details on the scratch caused by the
 222 mass movement; **c** Geological cross section (profile line 3-3') of the failed retaining slope located at the
 223 original quarry pit between the two rock hills.

224 **5.2 Flow-Accumulation Area**

225 The fan-shaped flow-accumulation area mantled the retaining slope of the dumpsite as well as a large part
 226 of the industrial park as shown in Fig. 12a. The water content of the displaced material was 32.1%-37.2%
 227 (3 sampling locations). The original slope gradient of the industrial park (Φ) was 1.36° along the profile
 228 line 1-1' as shown in Fig. 12 and Table 1. The area located immediately downstream of the dumpsite was
 229 relatively flat with no major construction except a pond (area approx. 3600 m^2) and a channel (width: 7 m
 230 and length: 130 m). Most of the industrial structures were located on the east, west, and north side of the
 231 pond as shown in Fig. 9, and thus created an ideal flow channel for the flowslide. The failure of the rock
 232 retaining slope resulted in an opening for the waste fill movement which subsequently destroyed
 233 downstream buildings (Fig. 14 and 15).



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Fig. 14 **a** The overview of the flow-accumulation area with elevations (China News Agency); **b** Geological cross section (profile line 4-4') of the flow-accumulation area. The length (L_4) was 664.4 m and the width (W_3) was 218.3 m with the width of the front edge (W_5) of 592.3 m. The area of the flow-accumulation area is $28.51 \times 10^4 \text{ m}^2$, and the elevation difference (H_4) was 35.7 m with an apparent dip (Φ_6) of 3.08° . The average thickness in the flow-accumulation area was 8.21 m with the maximum thickness of 18.82 m (Fig. 12c, Table 1, and Table 3).



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Fig. 15 The damage of buildings in the industrial park (China News Agency)

243 **6 Back-Analyses of the Flowslide**

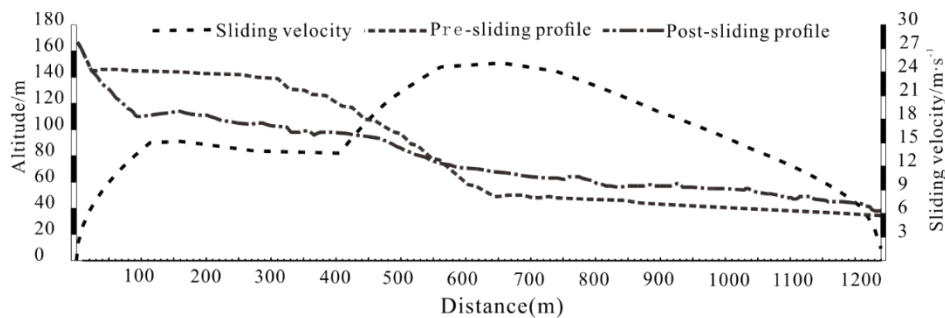
244 **6.1 Flowslide Movement**

245 The high sliding velocity and long-runout distance of the flowslide may be related to a more diffuse failure
246 due to liquefaction near the base of the slope. Evidences of localized liquefaction can be found at several
247 locations (Fig. 16). Empirical correlation was used to back-calculate the flowslide travel velocity on the
248 basis of its geometrical characteristics. Laboratory tests and numerical analyses were performed to better
249 understand the failure process.



250
251 Fig. 16 Evidences of liquefaction near the opening of the failed slope.

252 The source area of the flowslide was in the elevation of 142 m and the horizontal runout distance (L) was
253 approximately 1203 m with an elevation difference (H) of 111 m (Fig. 11). The velocity of a high-speed
254 landslide can be estimated by $v = \sqrt{2g \times (H - f \times L)}$ (Scheidegger, 1973), where v is the sliding velocity
255 (m/s), g is the gravitational acceleration (m/s^2), H and L are the elevation difference and horizontal distance
256 (m) between the crown and toe of the flowslide, respectively. f is the equivalent friction coefficient referred
257 as the ratio of height and runout distance of the flowslide ($f = H/L$). The equivalent friction coefficient of
258 the flowslide was 0.092. The sliding velocity was back-calculated and presented in Fig. 17.



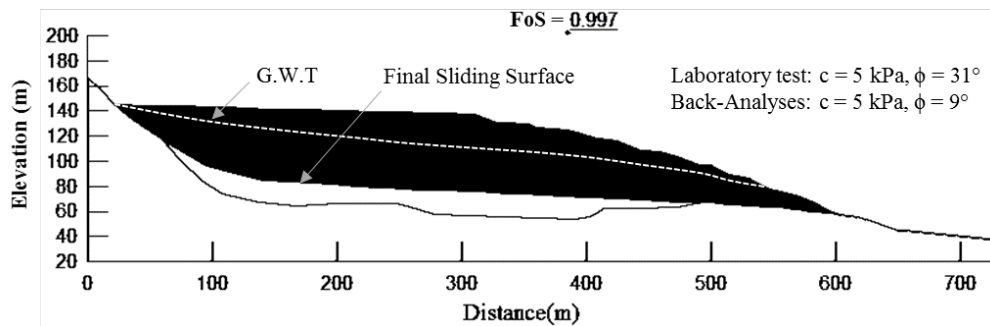
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260 Fig. 17 The relationship between sliding velocity along profile line 1-1' and pre-/post-sliding landform.

261 Based on the back-calculated velocity profile, two sharp increases were identified, including the initiation
262 of the waste filling near the steep scarp and the acceleration of the flowslide when exiting the dumpsite.
263 The sliding velocity was increased to 15.17 m/s as waste filling reached the bottom of the steep scarp. The
264 velocity displaced material gradually decreased to approximately 13 m/s before accelerate to the maximum
265 velocity of 25.15 m/s as it reached the opening of the quarry. The second sharp acceleration was followed
266 by the rapid dissipation of kinematic energy and reduction in velocity after reaching the flow-accumulation
267 area (elev. 50 m). It was estimated that the sliding velocity was reduced to 15.68 m/s when it made contact
268 with the downstream buildings, and then ceased moving due to the obstruction. The geometry and velocity
269 exhibited clear characteristics associated with high-speed long-runout flowslides.

270 6.2 Numerical Analyses

271 Basic material properties were obtained from *in-situ* and laboratory tests. Dry density of the waste filling
272 was 1.25-1.48 g/cm³ with a void ratio of 0.83-1.31. Standard compaction tests suggested that the optimal
273 water content of 15.3% with highest dry density of 1.79 g/cm³. The surface of the filling was in loose state
274 with the degree of compaction of 69.8%-82.7%. Based on the undrained shear test, the *c* and ϕ of the filling
275 were 4.7 kPa and 31.9°, respectively, and thus the friction angle was considerably higher than the gradient
276 of the slope. No strain-softening was observed in the saturated specimen under triaxial tests.

277 Numerical simulation for diffuse failure involves liquefaction and post-failure propagation is challenging
278 (Take and Beddoe, 2014), and the conventional LEB is typically not applicable for analyzing propagation
279 of landslide originating from diffuse failure induced by liquefaction (Cascini et al., 2009, 2013). Back-
280 analyses was performed using SLOPE/W with Morgenstern Prince Limit Equilibrium Analysis under the
281 assumption of a fully saturated basal zone in the dumpsite before failure. The back-analyses using LEB
282 method typically set the FoS to unity for back-calculating the mobilized strength. The final sliding surface
283 and a hypothetical groundwater level were added to determine the shear strength in an iterative approach with
284 estimated dry density of filling of 1.65 g/cm³ (Fig. 18), however the back-calculated friction angle was
285 significantly less than the experimental result.



286

287 Fig. 18 Back-calculated mobilized strength by using 2D numerical model with LEM.

288 The implausible mobilized strength shows that back-analyses using LEB is not applicable to the flowslide,
289 as the failure may involve liquefaction with subsequent progressive failure and post-failure propagation.
290 Notwithstanding the complex progressive mechanism and over-simplified force equilibrium method, the
291 misleading results of the back-analyses can be directly ascribe to the misuse of static pre-shearing pore-
292 water pressure with measured final sliding surface. Such erroneous back-analyses strategy was discussed
293 in detail by Take et al., (2004). It was postulated that the failure may be initiated at a relatively shallow
294 depth near the base of the slope as a result of static liquefaction, and followed by progressive backward
295 mobilization of the fillings in the dumpsite.

296 The exact reason of excess pore water pressure remains unclear but it was a key factor in causing the failure
297 (Ouyang et al., 2016). The excess pore water pressure can be induced by: (1) rapid surcharge on the
298 dumpsite while pore water pressure cannot dissipate sufficiently fast, and/or (2) the waste filling was
299 loosely packed with large pores in the meso-structure, and the collapse of the structure lead to the shrinkage
300 of pores, which resulted in the excess pore water pressure. Since the permeability of the waste filling was
301 considerably low, it may result in saturation in the basal zone of the waste filling. As no clear drainage
302 passage were found in the filling, the displaced material may remain undrained 'at failure'.

303 7 Conclusion

304 The Hongao dumpsite failure is of direct interest to the scientific community due to its complex progressive
305 failure mechanism and significant societal impact. The flowslide was investigated here to better understand
306 its characteristics and mechanism. The flowslide is divided into the source area and the flow-accumulation
307 area. The volume of the source area was $2.32 \times 10^4 \text{ m}^3$ with the average thickness of 20.5 m (max. 41.5 m).

308 The volume of the fan-shaped flow-accumulating area was $2.34 \times 10^6 \text{ m}^3$ with the average thickness of 8.2
309 m (max. 18.8 m). The volume expansion coefficient of the flowslide was 1.007. The maximum sliding
310 velocity of the flowslide was 25.15 m/s at the opening of the dumpsite, and reduced to 15.68 m/s as it
311 reached the industrial park and ceased moving due to the obstruction of buildings.

312 The flowslide was characterized by high-speed and long runout distance, which may be related to the (1)
313 suitable topography with a height of 124 m between the crown and toe with large potential energy needed
314 for high-speed and long runout flowslide, and (2) the low permeability and lack of drainage in the waste
315 filling in the dumpsite with impermeable bedrock resulted in groundwater stagnation, and thus high pore-
316 water pressure. Additionally, the volume of the waste filling was estimated as $6.27 \times 10^6 \text{ m}^3$, which was three
317 times larger than the design capacity of the dumpsite. The failure of the retaining slope constructed between
318 two bedrock ridges formed a narrow opening for the flowslide, which facilitated the sudden release of
319 kinematic energy generating high sliding velocity and long travel distance.

320 The failure mechanism remains unclear yet it was clear that a more diffuse failure occurred with liquefaction
321 and post-failure propagation of the flowslide. The inapplicability of the LEM demonstrated the complexity
322 of the mechanism by yield erroneous mobilized strength, which also indicates the predicament of simulating
323 liquefaction-induced slope failures with conventional numerical approaches. The cause(s) of excess pore-
324 water pressure is not clear, but the unregularly disposal activities in addition to the ingress of rainwater and
325 high pore water pressure played important roles in deformation of the dumpsite. Although the flowslide
326 destruction process was fast with excessive deposits, it is postulated that signs of deformation may have
327 already appeared in the study site but not discovered due to the absence of field monitoring. Further analysis
328 is undergoing at SKLGP to assess the cause(s) involved for generating the excess pore water pressure.

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