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Rainfall thresholds for shallow landslides occurrence in Calabria, southern Italy

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Abstract

In many areas, rainfall is the primary trigger of landslides. Determining the rainfall conditions responsible for landslide occurrence is important, and may contribute to save lives and properties. In a long-term national project for the definition of rainfall thresholds for possible landslide occurrence in Italy, and for the implementation of a national landslide warning system, we compiled a catalogue of 186 rainfall events that have resulted in 251 shallow landslides in Calabria, southern Italy, from January 1996 to September 2011. Landslides were located geographically using Google Earth®, and were given a mapping and a temporal accuracy. We used the landslide information, and sub-hourly rainfall measurements obtained from two complementary networks of rain gauges, to determine cumulated event vs. rainfall duration (*ED*) thresholds for Calabria. For the purpose, we adopted an existing method used to prepare rainfall thresholds and to estimate their associated uncertainties in central Italy. The regional thresholds for Calabria were found nearly identical to previous *ED* thresholds for Calabria obtained using a reduced set of landslide information, and slightly higher than the *ED* thresholds obtained for central Italy. We segmented the regional catalogue of rainfall events with landslides on lithology, soil regions, rainfall zones, and seasonal periods. The number of events in each subdivision was insufficient to determine reliable thresholds, but allowed for preliminary conclusions on the role of the environmental factors on the rainfall conditions responsible for shallow landslides in Calabria. We further segmented the regional catalogue based on administrative subdivisions used for hydro-meteorological monitoring and operational flood forecasting, and we determined separate *ED* thresholds for the Tyrrhenian and the Ionian coasts of Calabria. We expect the *ED* rainfall thresholds for Calabria to be used in regional and national landslide warning systems. The thresholds can also be used for landslide hazard and risk assessments, and for erosion and landscape evolution studies, in the study area and in similar physiographic regions in the Mediterranean area.

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1 Introduction

In Italy, rainfall-induced landslides are common geomorphological events that cause economic and environmental damage, and casualties (Salvati et al., 2010). Due to the morphological, geologic and climatic settings, rainfall-induced landslides are frequent and abundant in Calabria, a region in southern Italy (Sorriso-Valvo, 1993; Trigila et al., 2010; Gullà et al., 2008; Iovine et al., 2009). In the framework of a national project for the definition of rainfall thresholds for possible landslide occurrence, and for the implementation of a nationwide landslide warning system in Italy (Brunetti et al., 2010; Rossi et al., 2012), we have defined empirical rainfall thresholds for shallow landslide occurrence in Calabria.

In this work, we introduce the landslide information and the rainfall data used to define the thresholds in Calabria, we present the regional thresholds and their uncertainties, and we discuss limits to the definition of rainfall thresholds for possible landslide occurrence in the region. The paper is organized as follows. In Sect. 2, we provide background information on the definition and use of empirical rainfall thresholds for landslide occurrence, focusing on cumulated event rainfall – rainfall duration (*ED*) thresholds (Guzzetti et al., 2007, 2008). This is followed by a description of the study area (Sect. 3), and of the data used and the method adopted to define the thresholds (Sect. 4). In Sect. 5, we present the *ED* thresholds, including regional thresholds, and thresholds prepared for different subdivisions based on lithology, soil types, rainfall and seasonal characteristics, and administrative subdivisions for hydro-meteorological monitoring and forecasting. Next, we discuss the results obtained and we compare the *ED* thresholds defined for Calabria to similar thresholds available for other areas in Italy. We conclude (Sect. 7), summarizing the lessons learned.

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2 Background

In the literature, rainfall-induced shallow landslides, including soil slips, soil slides, debris flows, and rock falls (Cruden and Varnes, 1996) are predicted using empirical rainfall thresholds or spatially distributed, physically based numerical models. In this work, we concentrate on the definition of rainfall thresholds for the possible occurrence of shallow landslides. Most commonly, a threshold defines the lower bound of known rainfall conditions that have resulted in landslides (e.g., Reichenbach et al., 1998; Corominas, 2000; Crosta and Frattini, 2001; Aleotti, 2004; Wieczorek and Glade, 2005). To establish a threshold, different types of rainfall measurements can be used (Guzzetti et al., 2007). Inspection of the literature reveals that most of the thresholds are determined considering the duration (D) of the rainfall events, the event cumulated rainfall (E) or the mean rainfall intensity (I) of the events that have resulted in landslides. Rainfall thresholds can be defined for different geographical areas, from a local to global thresholds (Guzzetti et al., 2007, 2008).

Endo (1969), working in the Hokkaido island, Japan, was probably first to introduce the concept of a minimum amount of rainfall to trigger landslides. Onodera et al. (1974), also working in Japan, were first to propose quantitative rainfall thresholds for the possible initiation of landslides. Campbell (1975) studied the rainfall conditions that triggered soil slips in southern California in 1969, and suggested that the failures were the result of the combination of a cumulated rainfall and of the event rainfall intensity. Caine (1980) collected information on 73 rainfall intensity I vs. rainfall duration D conditions responsible for shallow landslides globally, and was first to propose a global (minimum) ID threshold expressed by the power-law equation $I = 14.82 \cdot D^{-0.39}$. Based on this relationship, when $I < 14.82 \cdot D^{-0.39}$ landslide activity is unlikely. Innes (1983) used information on 35 rainfall events that had resulted in debris flows worldwide to establish a global minimum cumulated rainfall event E – rainfall duration D (ED) threshold for possible debris flow occurrence, given by the power-law equation $E = 4.9355 \cdot D^{0.5041}$. Again, based on this relationship, when $E < 4.9355 \cdot D^{0.5041}$ debris flow activity is un-

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likely. More recently, Kanji et al. (2003) collected information on the rainfall conditions that have resulted in debris flows and landslides in several geographical areas, and proposed the global minimum *ED* threshold $E = 22.4 \cdot D^{0.41}$.

Guzzetti et al. (2007, 2008) reviewed the literature on rainfall thresholds for possible landslide occurrence and recognized that many of the thresholds were defined using ill-formalized, poorly documented, or non-reproducible methods. The lack of standards has negative consequences on the possible use of the thresholds. In an attempt to measure the uncertainty inherent to the thresholds, Brunetti et al. (2010) and Peruccacci et al. (2012) proposed an objective method for the definition of empirical rainfall thresholds based on a statistical *frequentist* approach. In particular, Brunetti et al. (2010) collected information on 753 rainfall events that triggered landslides in Italy from 1841 to 2009, and defined *ID* thresholds for Italy and for the Abruzzo region, central Italy. The regional *ID* threshold for Abruzzo was found to be lower than the threshold defined for Italy, and lower than similar regional *ID* thresholds defined for areas in northern Italy by Aleotti (2004) and Ceriani et al. (1994) and in southern Italy by Calcaterra et al. (2000). More recently, Peruccacci et al. (2012) compiled a new catalogue of 442 rainfall events that triggered landslides in the Abruzzo, Marche, and Umbria regions, central Italy, between 2002 and 2010, modified the method proposed by Brunetti et al. (2010) to consider the uncertainties associated with the definition of the parameters defining the thresholds, and proposed *ED* thresholds for different exceedance probabilities for the entire study area and for three administrative subdivisions, for the main lithological domains, and for different seasonal periods. The cumulated rainfall necessary to trigger landslides in the study area was found slightly larger for flysch deposits than for soft post-orogenic sediments (clay, silt, sand, gravel) and for sedimentary carbonate rocks. The thresholds for the dry and the wet seasonal periods were found to be indistinguishable for rainfall durations in the range $12 \leq D \leq 100$ h.

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3 Study area

Our study area comprises the Calabria region that extends for 15 080 km² in southern Italy with elevation in the range from sea level to 2267 m at Monte Serra Dolcedorme, in the Pollino range (Fig. 1a). Most of the area consists of crystalline-metamorphic nappes, Palaeozoic in age, that form the Calabrian Arc (Fig. 1b), and that were thrust on Mesozoic–Cenozoic units comprising the Apennines mountain chain, and were then covered by sedimentary rocks, Miocene to Quaternary in age (Amodio Morelli et al., 1976). The Calabrian–Sicilian rift zone (Monaco and Tortorici, 2000), a system of active normal faults along the western side of the Calabrian Arc, produced intense regional uplift. The uplift started in the Pleistocene and it is responsible for the intense historical seismicity of the area (Monachesi and Stucchi, 1997; Boschi et al., 2000). Due to the complex geodynamic history (Monaco and Tortorici, 2000; Tortorici et al., 1995; Westaway, 1993), in Calabria rocks are highly deformed, fragmented, and deeply weathered. Landscape is hilly or mountainous, with lowlands confined to local tectonic depressions and along the coasts. The mountains are made of crystalline-metamorphic rocks mantled by thick soils, except in the northernmost portion of the region and in tectonic windows where carbonate rocks crop out. In the hills crop out clastic sediments, and clay, silt, and sand are most abundant in the coastal plains. Orography controls the precipitation regime (Bellecci et al., 2002), which is driven by regional meteorological fronts approaching transversally the region, and by convective cells forming along the steep coastal slopes. Cumulated annual precipitation ranges from less than 500 mm along the coasts to more than 2000 mm in the mountains. About 70 % to 80 % of the annual precipitation falls from October to March, with only 10 % of the precipitation in the summer. Intense summer storms are more frequent along the Ionian (eastern) coast than along the Tyrrhenian (western) coast (Terranova and Iaquina, 2011). Due to the local setting, landslides of different types and sizes are abundant and frequent in Calabria (Sorriso-Valvo, 1993; Iovine et al., 2009). Recent regional studies have shown

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that approximately 23 % of the territory is highly susceptible to shallow landslides (Gullà et al., 2008).

4 Method

To determine rainfall thresholds for possible shallow landslide occurrence in Calabria, we first prepared a catalogue of rainfall events that have resulted in shallow landslides in the region in the 16 yr period from January 1996 to September 2011. The catalogue was compiled searching national and local newspapers and journals, local historical archives, Internet blogs, and the scientific and technical literature. Additional valuable information was obtained searching the archive of the local Fire Brigades of the Cosenza Province. For each landslide we collected information on: (i) the geographical location of the landslide, including latitude, longitude, and elevation, (ii) the time or period of occurrence of the slope failure, and (iii) the landslide type, classified according to Cruden and Varnes (1996). The geographical coordinates and the elevation of the individual landslides were obtained using Google Earth®. Individual landslides were mapped as points, and were given a level of geographic accuracy, in three classes: high ($P_1 < 1 \text{ km}^2$), medium ($1 \leq P_{10} < 10 \text{ km}^2$), or low ($10 \leq P_{100} < 100 \text{ km}^2$) (Peruccacci et al., 2012). Since accurate information on the time (or the period) of occurrence of a landslide is difficult to obtain and the temporal information is often affected by considerable uncertainty (Guzzetti et al., 2007), a level of temporal accuracy was assigned to each landslide, in three classes: high when the hour of occurrence of the failure was known, intermediate when the period of the day (e.g., morning, mid day, afternoon, evening, night) was known, and low when only the day of occurrence was known.

When the geographical location and the time (or period) of occurrence of a landslide were established, the rainfall duration D and the cumulated event rainfall E that (presumably) caused the failure were determined using two synergic sources of rainfall information (Fig. 1a). The first source of information was the national database of

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rainfall measurements assembled and managed by the National Department for Civil Protection. For Calabria, and at the time of the study, this database contained sub-hourly rainfall measurements obtained by a network of 113 rain gauges. A second source of rainfall information was the regional database of sub-hourly rainfall measurements assembled and managed by the Calabria Regional Agency for Environmental Protection (ARPACAL). At the time of the study, this database contained rainfall measurements obtained by a network of 154 rain gauges. Figure 1a shows the geographical location and the spatial density of the rain gauges, and reveals that 75 rain gauges belong to both networks.

To determine the rainfall conditions considered responsible for a landslide, a “representative” rain gauge had to be selected. To select the rain gauge, the following selection criteria were decided: (i) the geographical distance between the landslide and the rain gauge shall not exceed 12 km, (ii) the elevation of the rain gauge and of the landslide shall be comparable, and (iii) the general physiographical settings for the rain gauge and the landslide should also be comparable e.g., they should be located in slopes facing in the same direction, or in the same or in similar valleys.

When the representative rain gauge was selected, the rainfall conditions considered responsible for the landslide were computed. To determine the rainfall measurements the most critical step was the selection of the “start time” and the “end time” of the rainfall event. For the purpose, the following criteria were adopted. The “end time” of the rainfall event was taken to coincide with the known or inferred time of the landslide occurrence (i.e., the hour, the end of the period of the day, or the end of the day for landslides characterized by a high, medium or low temporal accuracy, respectively). The “start time” of the rainfall event was decided based on a heuristic procedure that considered the (known or presumed) meteorological conditions resulting in the rainfall event. For events produced by regional frontal systems, a dry (no rain) period of two days between April and October (A–O), and of four days between November and March (N–M) was used to separate successive events. The difference between the two periods is justified by the Mediterranean climate that dominates the study area. For

rainfall events produced by local convective conditions, the “start time” for the event was decided searching visually for an abrupt (distinct) increase in the slope of the cumulated rainfall curve. This is similar to what was proposed by Aleotti (2004). When the “start time” and the “end time” of the rainfall event were decided, the rainfall duration D (in hours) and the corresponding cumulated event rainfall E (in mm), were computed.

To determine the ED rainfall thresholds, we adopted the method proposed by Brunetti et al. (2010), and improved by Peruccacci et al. (2012). The method allows for the definition of empirical rainfall thresholds for different exceedance probability levels, and to estimate the uncertainty associated to the thresholds. This is a significant advantage over pre-existing, heuristic or subjective methods for the definition of rainfall thresholds (Guzzetti et al., 2007, 2008). For the method, the threshold curves are described by power laws of the general form $E = (\alpha \pm \Delta\alpha) \cdot D^{(\gamma \pm \Delta\gamma)}$, where α is the scaling parameter (the intercept), γ is the shape parameter (the exponent) that controls the slope of the power law curve, and $\Delta\alpha$ and $\Delta\gamma$ are the uncertainties associated to α and γ , respectively. The method adopts a “bootstrap” non-parametric statistical technique (Efron, 1979; Efron and Tibshirani, 1994) to estimate the mean values of α and γ , and the associated uncertainties $\Delta\alpha$ and $\Delta\gamma$ (Peruccacci et al., 2012).

5 Results

Our final catalogue lists 186 rainfall events that have resulted in 251 shallow landslides in Calabria in the period from January 1996 to September 2011. Table 1 and Fig. 2 summarize the abundance and the temporal distribution of the 186 triggering events in the five provinces in Calabria. The number of events is largest for the Cosenza Province that covers 44.1 % of the region and comprises 60.2 % of the listed events. The temporal distribution of the events is similar for the five provinces. The events listed in the catalogue have rainfall durations in the range $1 \leq D \leq 451$ h and cumulated event rainfall in the range $12.8 \leq E \leq 542.0$ mm. Figure 3 shows the cumulated number and

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the frequency of the rainfall duration D , and of the cumulated rainfall E , for the 186 rainfall events in the catalogue.

5.1 Cumulated event rainfall – rainfall duration (ED) regional thresholds

The 186 blue dots in Fig. 4 portray the rainfall duration vs. cumulated event rainfall (D , E) conditions that have resulted in shallow landslides in Calabria in the investigated period. Visual inspection of the plot reveals that the cloud of points is reasonably well distributed in the range of durations and cumulated event rainfall to allow for the definition of reliable thresholds (Brunetti et al., 2010; Peruccacci et al., 2012). In the plot, the blue line and the associated shaded area show the 5 % ED threshold ($T_{5,C}$), $E = (8.6 \pm 1.1) \cdot D^{(0.41 \pm 0.03)}$. Table 2 lists the equations and the uncertainties associated to the 1 % ($T_{1,C}$) and the 5 % ($T_{5,C}$) ED thresholds defined for Calabria.

We have analysed the uncertainty associated to the ED thresholds (Fig. 5). The analysis was performed adopting the bootstrap non-parametric statistical technique proposed by Peruccacci et al. (2012). As expected, the variation in the mean values of the two parameters α and γ that controls the power law threshold curves, and the associated uncertainties ($\Delta\alpha$ and $\Delta\gamma$), decreases when the number of events increases. The mean value of the scaling parameter converges to $\alpha = 5.8$ for $T_{1,C}$ and to $\alpha = 8.6$ for $T_{5,C}$ (Fig. 5a and c). Similarly, the mean value of the exponent of the power law model converges rapidly to $\gamma = 0.41$, for both thresholds $T_{1,C}$ and $T_{5,C}$ (Fig. 5e). The uncertainties $\Delta\alpha$ and $\Delta\gamma$, represented by the error bars in Fig. 5a, c, e and by the square symbols in Fig. 5b, d, and f, decrease as the number of events increases. In particular, for $n < 75$, the values of α have not yet converged to their limits. For $n \geq 75$, the average values of the α and γ parameters do not vary significantly (e.g. α varies < 0.2), but the uncertainties remain significant (for $n = 75$ the relative uncertainties are $\Delta\alpha/\alpha = 17\%$ and $\Delta\gamma/\gamma = 11\%$).

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5.2 Role of environmental factors

To investigate the role of different environmental factors on the rainfall conditions responsible for shallow landslides in Calabria, we segmented the regional catalogue in different sub-sets. The segmentation was performed for (i) the main lithological domains, (ii) the soil regions, (iii) the rainfall regions, and on (iv) two seasonal periods. We further segment the catalogue on administrative alert zones decided by the National Civil Protection Department for hydro-meteorological forecasting and warning.

To attribute a landslide to a specific environmental class (e.g., a lithological type, soil region, rainfall region), the individual landslides originally mapped as points were transformed to circles, with the size (area) of the circle dependent on the geographical accuracy of each landslide. Specifically, the area of the circles was $A = 0.5 \text{ km}^2$ for landslides attributed a high level of accuracy (P_1), $A = 5 \text{ km}^2$ for landslides attributed a medium level of accuracy (P_{10}), and $A = 50 \text{ km}^2$ for landslides attributed a low level of accuracy (P_{100}). Next, the GIS layer containing the landslide circles was intersected with the individual layers representing the single environmental factors, or the administrative subdivision, and the proportion of each class of terrain in the circle was computed. Only landslides for which at least 75 % on a single environmental class was present in the representative circle were selected, and used for the subsequent analyses. The selection was adopted to reduce the uncertainty associated to the presence of multiple environmental factors conditioning a single landslide (Peruccacci et al., 2012).

5.2.1 Lithological domains

Using the lithological map of Italy compiled by Compagnoni et al. (1976–1983) at 1 : 500 000 scale, we partitioned the Calabria region in four broad lithological domains encompassing (Fig. 6a): (i) carbonate rocks (CC, 13.6 % of the region), including limestone, dolomitic limestone, dolostone, hard sandstone, and conglomerate, (ii) flysch deposits (FD, 30.1 %), comprising marl, shale, evaporite, and very-low to low grade metamorphic rocks, (iii) metamorphic and crystalline rocks (MC, 30.6 %), comprising

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medium to high-grade metamorphic rocks (gneiss), intrusive rocks (granite, granodiorite), and ophiolites, and (iv) post-orogenic sediments (PO, 25.7 %), chiefly continental and marine clay, silt, sand, and gravel.

Intersection in the GIS of the landslide layer with the lithological layer (Fig. 6a) identified 142 landslides (56.5 %) for which at least 75 % of the representative circle was attributed to a single (prevalent) lithological class, including: (i) 56 landslides (39.4 %) in the PO domain (14.4 landslides per 1000 km²), (ii) 49 landslides (34.5 %) in the MC domain (10.6 landslides per 1000 km²), (iii) 24 landslides (16.9 %) in the FD domain (5.3 landslides per 1000 km²), and (iv) 13 landslides (9.2 %) in the CC domain (6.3 landslides per 1000 km²) (Table 3).

Figure 7a portrays the 119 rainfall (D , E) conditions that have resulted in the 142 shallow landslides in the four lithological domains. The reduced number of empirical data points in each lithological domain (11 for CC, 23 for FD, 39 for MC, and 46 for PO) was insufficient to determine reliable thresholds. However, significant differences are not evident in the distribution and the general trend of the (D , E) data points, for the different domains. Inspection of Fig. 7a shows that the rainfall conditions responsible for landslides in the MC and the PO lithological domains cover the entire range of rainfall durations ranging from a few hours to several days ($1 \leq D \leq 357$ h), whereas the conditions that have caused landslides in the FD and in the CC domains are more numerous for durations $D > 10$ h. The apparent lack of events for short durations is probably due to the reduced number of events collected in the two FD and CC domains. The trend for each lithological domain has to be confirmed by collecting a larger number of events.

5.2.2 Soil regions

The Calabria Regional Agency for the Development and Services in Agriculture published a Soil Map of Calabria, at 1 : 250 000 scale (Fig. 6b) (ARSSA, 2003). The map shows four soil regions characterized by distinct morphological, lithological, climatic, and soil properties, including: (i) the Pollino range and the neighbouring hills (SR1,

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7.8 % of the region), comprising mountains and hills where carbonate rock crop out and climate is Mediterranean, (ii) the northern Ionian coastal zone (SR2, 4.1 %), comprising hills, coastal plains, and piedmont zones underlined by flysch rocks and where climate is Mediterranean mountainous, (iii) the coastal hills and plains (SR3, 50.2 %), comprising alluvial and coastal plains and hills where the bedrock consists of sediments Tertiary to Quaternary in age, and climate is sub-continental to sub-tropical Mediterranean, and (iv) the Sila, Serre, and Aspromonte ranges (SR4, 37.9 %), comprising mountains, hills, and minor internal plains where the bedrock consists of crystalline and metamorphic rocks, and the climate is Mediterranean mountainous (ARSSA, 2003).

Intersection in the GIS of the landslide layer with the soil map (Fig. 6b) singled out 199 landslides (79.3 %) for which at least 75 % on the representative circle was attributed to a single (prevalent) soil class, including: (i) 130 landslides (65.3 %) in the SR3 soil region (17.2 landslides per 1000 km²), (ii) 57 landslides (28.6 %) in the SR4 region (10 landslides per 1000 km²), and (iii) 12 landslides (6 %) in the SR1 region (10.2 landslides per 1000 km²) (Table 3). No landslides were identified in the SR2 soil region, limited to the NE part of the region (Fig. 6b).

Figure 7b portrays the 157 rainfall (*D*, *E*) conditions that have produced 199 shallow landslides in the three soil regions. As for the lithological domains, the number of empirical data points in each soil region was insufficient for a reliable determination of specific rainfall thresholds. For the SR1 and the SR3 soil regions, Fig. 7b shows minor differences in the distribution of the empirical data points. Rainfall conditions in the SR1 soil align along a straight line (power law) with a very low dispersion. However, this may be an artefact due to the very reduced number of data points in this soil region (9). We conclude that rainfall conditions for landslide occurrence in the three soil regions for Calabria cannot be discriminated currently.

5.2.3 Rainfall regions

For operational flood forecasting and hazard assessment, Versace et al. (1989) analysed the annual maxima of high-intensity, short-duration rainfall events in Calabria, and

identified three homogeneous rainfall regions (Fig. 6c), including: (i) a Tyrrhenian region (T, 25.8 % of the Region) along the western Tyrrhenian coast, (ii) a Central region (C, 44.3 %), comprising the mountain ranges along the main divide, and (iii) a Ionian region, along the eastern Ionian coast (I, 29.9 %).

Intersection in the GIS of the landslide layer with the map showing the three rainfall regions (Fig. 6b) recognized 212 landslides (84.5 %) for which at least 75 % on the representative circle was attributed to a prevalent rainfall region, including: (i) 117 landslides (55.2 %) in the Tyrrhenian region (17.5 landslides per 1000 km²), (ii) 63 landslides (29.7 %) in the Central region (16.2 landslides per 1000 km²), and (iii) 32 landslides (15.1 %) in the Ionian region (7.1 landslides per 1000 km²) (Table 3).

Figure 7c shows the 158 rainfall (*D*, *E*) conditions responsible for the 212 shallow landslides in the three rainfall geographical subdivisions. Again, the number of empirical points in each rainfall region was insufficient to establish reliable thresholds for the individual regions. Inspection of Fig. 7c shows that the rainfall conditions responsible for landslides in the Ionian (I) and in the Central (C) regions are more abundant than the rainfall events in the Tyrrhenian region (T). This is a consequence of the spatial distribution of rainfall events in the catalogue. Inspection of Table 1 reveals that about 60 % of the events (151 over 251) is located in the northern part of the region (in the Cosenza province), where the Ionian area is smaller than the Tyrrhenian and the Central areas. Figure 7c shows that, on average, the rainfall (*D*, *E*) conditions that triggered landslides in the Tyrrhenian region are less severe than in the Ionian and Central regions.

5.2.4 Dry and wet seasonal periods

In Calabria, climate is Mediterranean with distinct dry and wet seasonal periods. To investigate the dependence of the rainfall events with the seasonal rainfall conditions, we have segmented the catalogue in two seasonal periods: (i) a wet period from November to March (N–M), and (ii) a dry period from April to October (A–O). Of the 186 rainfall (*D*, *E*) conditions in the catalogue, 67.7 % are in the wet (N–M) period, and 32.3 % in the dry (A–O) period (Fig. 8, Table 3). Rainfall events in the N–M period have durations

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ranging from a few hours to several days ($3 \leq D \leq 451$ h, $\bar{D} = 98$ h), and low rainfall mean intensities ($\bar{I} = 3 \text{ mm h}^{-1}$). Conversely, events in the A–O period exhibit shorter durations ($1 \leq D \leq 290$ h, $\bar{D} = 38$ h) and higher mean intensities ($\bar{I} = 7.3 \text{ mm h}^{-1}$). In detail, 65 % of the events occurred in the dry (A–O) period are characterised by rainfall durations not exceeding 24 h, and 75 % of the events in the wet (N–M) period have durations longer than 24 h. The average number of rainfall events per month is 25.2 for the N–M period, and 8.6 for the A–O period. Rainfall events with landslides are most frequent in January (43) and are rare in July (4).

Figure 8 shows the rainfall (D , E) conditions for 126 events in the wet (N–M) period (grey dots), and for 60 events in the dry (A–O) period (blue dots). For both periods, the number of rainfall events was insufficient to determine reliable thresholds for the two seasonal subdivisions. Nonetheless, differences exist in the distribution and the general trend of the (D , E) empirical data points, for the considered seasonal periods. In Calabria, most of the rainfall events that have resulted in landslides and are listed in the catalogue occurred in the winter (53.8 %) and in the autumn (20.0 %), with 23.5 % of the events in January and 19.1 % in December. The remaining events occurred in the spring (15.1 %) and in the summer (11.1 %), with 10.4 % of the events in September (10.4 %). The result is in agreement with the outcomes of an historical analysis of 2982 historical rainfall induced landslides in Calabria between 1921 and 2006 that revealed that 67 % of the landslides occurred in the 4 month period between November and February (Polemio and Petrucci, 2010). The distribution of the event duration in the two periods is different. The majority of rainfall events with landslides in the A–O period are characterised by durations $D < 10$ h, while in the N–M period in most of the triggering conditions $D > 10$ h. This result is in agreement with the results of Peruccacci et al. (2012) for the Abruzzo, Marche and Umbria regions.

5.3 Sub-regional thresholds for shallow landslide forecasting

For hydro-meteorological monitoring and operational flood forecasting, the Italian national Department for Civil Protection has divided the national territory into 134 alert

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zones, based on administrative boundaries and regional hydro-meteorological conditions. The subdivision partitions Calabria in six alert zones, three along the Tyrrhenian coast and three along the Ionian coast (Fig. 6d). The number of rainfall events with landslides in each division was insufficient to establish reliable rainfall thresholds for the single alert zones. Thus, considering the general hydro-meteorological characteristics in the alert zones (Terranova, 2003, 2004; Terranova et al., 2013) we grouped them in two broad alert regions (Fig. 6d) including: (i) a Tyrrhenian alert region (59.4 % of the total area), comprising the three alert zones along the Tyrrhenian coast, and (ii) a Ionian alert region (40.6 % of the total area), comprising the three alert zones along the Ionian coast. Intersection in the GIS of the landslide layer with the map showing the two alert regions (Fig. 6d) assigned 245 landslides (97.6 %) to an alert region, including: (i) 157 landslides (64.1 %) to the Tyrrhenian alert region (25.6 landslides per 1000 km²), and (ii) 88 landslides (35.9 %) to the Ionian alert region (9.8 landslides per 1000 km²) (Table 3).

Figure 7d shows the 186 rainfall (D , E) conditions responsible for 245 shallow landslides in the two alert regions. The ranges of rainfall duration in the two alert regions are similar, $1 \leq D \leq 357$ h for the Ionian alert region, and $1 \leq D \leq 451$ h for the Tyrrhenian alert region. The landslide triggering rainfall events in the Ionian alert region have, on average, a lower duration and a larger cumulated event rainfall than the events in the Tyrrhenian alert region (Table 3, Fig. 9a). For cumulated event rainfall exceeding about $E > 30$ mm, a larger amount of rainfall is necessary to trigger landslides in the Ionian region than in the Tyrrhenian alert region. The result was manifest in the subdivision shown in Fig. 7c. No significant difference exists for the duration of the events in the two alert regions (Fig. 9b). In the Ionian region rainfall events responsible for landslides are more frequent in September and from December to January, whereas in the Tyrrhenian alert region the events that have resulted in landslides were more abundant between December and March (Fig. 9c).

Figure 10 portrays the 186 rainfall (D , E) conditions that have produced 245 shallow landslides in the two alert regions. Despite the limited number of empirical data point,

we determined the rainfall thresholds for the possible occurrence of shallow landslides for the Tyrrhenian alert region ($T_{1,TAR}$ and $T_{5,TAR}$) and for the Ionian alert region ($T_{1,IAR}$ and $T_{5,IAR}$) (Table 2). Inspection of Fig. 10 reveals that the $T_{5,IAR}$ threshold defined for the Ionian region is steeper than the similar $T_{5,TAR}$ threshold established for the Tyrrhenian region. For rainfall durations exceeding one day, $D > 24$ h, the two thresholds are statistically different (i.e., their uncertainties do not overlap completely). We conclude that for rainfall events exceeding one day lower amounts of cumulated event rain are required to trigger landslides in the Tyrrhenian alert region than in the Ionian alert region.

6 Discussion

Our catalogue lists 186 rainfall events that have resulted in 251 shallow landslides in Calabria in the 16 yr period from January 1996 to September 2011 (Fig. 2). This is an average frequency of 11.8 landslide triggering rainfall events per year. The figure compares to 52.0 landslide-triggering rainfall events per year found by Peruccacci et al. (2012) for the Abruzzo, Marche and Umbria regions, central Italy. Peruccacci and her co-workers used the same strategy and similar sources of landslide and rainfall information used in this study. Considering that Calabria covers about half the area covered by the Abruzzo, Marche and Umbria regions ($15\,080\text{ km}^2$ vs. $28\,950\text{ km}^2$), the density of events (# of events per year per square kilometre) for Calabria is less than half of the density of events for central Italy. The difference is conditioned by the availability of information used to compile the catalogues, and by the criteria adopted to determine the rainfall conditions that have resulted in landslides. Thus, the difference cannot be used to measure climatic, meteorological, geomorphological or other physiographical differences, or to infer differences in landslide vulnerability or risk between the study areas.

The empirical (D , E) data collected for Calabria were sufficient to determine regional ED rainfall thresholds for possible landslide occurrence in the region, for different ex-

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ceedance probability levels (Brunetti et al., 2013; Peruccacci et al., 2012) (Fig. 4). However, the evaluation of the relative uncertainty associated with the thresholds revealed that the uncertainties were significant. Peruccacci et al. (2012), working in central Italy, considered acceptable values of the relative uncertainty $< 10\%$, which they reached for a number of empirical data points $n = 175$. Based on this evidence, they concluded that 175 well-distributed (D , E) empirical points were sufficient to estimate reliable rainfall thresholds. In Calabria, the relative uncertainty obtained using the entire set of 186 events is larger than 10% , at least for the scaling parameter α of the power law model i.e., $\Delta\alpha/\alpha = 13.8\%$. The relative difference of the shape parameter γ is $\Delta\gamma/\gamma = 7.3\%$. This suggests that the minimum number of empirical data points necessary to determine reliable rainfall thresholds varies regionally. Variations depend on the shape of the distribution (e.g., data scattering around the mean tendency line, slope of the line).

For a 1% (5%) threshold, 1% (5%) of the empirical (D, E) data points should be below the threshold (Brunetti et al., 2013). In the regional dataset for Calabria, only one empirical point (0.5%) is below the $T_{1,C}$ threshold, and five points (2.7%) are below the $T_{5,C}$ threshold. We conclude that the established ED thresholds ($T_{1,C}$ and $T_{5,C}$) are slightly lower than expected. The result is directly related to the statistical technique used to determine the thresholds, and to the asymmetric distribution of the empirical (D, E) points around their mean tendency line (Brunetti et al., 2010). The asymmetry may be an artefact due to the (relatively) limited number of empirical data, or it may reflect a specific characteristic of the landslide-triggering rainfall events in Calabria. More empirical data are needed to resolve this ambiguity.

Most of the landslides triggered by the 186 rainfall events listed in the catalogue were shallow failures, and they occurred in post-orogenic sediments (PO lithological domain) and where metamorphic and crystalline rocks (MS lithological domain) crop out. We attribute the large density of landslides in the PO lithological domain (39.4% , 14.4 failures per 1000 km^2) to the poor geotechnical properties of the continental and marine sediments in this domain, and subordinately to the local terrain conditions (e.g., steep slopes). We explain the large proportion of events in the metamorphic and crystalline

rocks (34.5 %, 10.6 failures per 1000 km²) with the abundance of weathered materials mantling the bedrock in this domain. For short duration rainfall events ($D \leq 24$ h), a smaller amount of rainfall is necessary to trigger landslides in the metamorphic and the crystalline rocks, compared to the post-orogenic sediments. For longer duration events ($D > 24$ h) the trend is reversed, and landslides in metamorphic and crystalline rocks require more rain than in the post-orogenic sediments to occur. We attribute the large abundance of landslides in the coastal hills and plains soil region (SR3, 50.2 %) to the peculiar setting of the region, where soft sediments, Tertiary to Quaternary in age, crop out. Landslides were more abundant in the Tyrrhenian rainfall region than in the Central and the Ionian regions, and the amount of rainfall needed to trigger the failures in the Ionian region was higher than in the Central and in the Tyrrhenian regions. Most of the landslides listed in the catalogue occurred in the wet period between November and March (in particular, in January and December) and where triggered by short duration rainfall events with $D > 10$ h. Further analysis revealed that, in Calabria, shallow landslides are more frequent in clay, silt, sand, and gravel deposits (PO lithological domain). In this domain, the failures occur primarily in the SR3 soil region (coastal hills and plains) of the Tyrrhenian rainfall region, between November and March. Conversely, shallow landslides are less frequent where marl, shale, evaporites, and very-low to low grade metamorphic rocks crop out (FD lithological domain). In this domain, the failures are more rare in the Ionian rainfall region, and in the seasonal period between March and May.

It is worth comparing the *ED* thresholds curves defined for Calabria with similar *ED* thresholds. Figure 11 shows the new regional threshold ($T_{5,C}$) and the sub-regional thresholds for the Tyrrhenian ($T_{5,TAR}$) and the Ionian ($T_{5,IAR}$) alert regions, for the 5 % exceedance probability level, together with the *ED* threshold curves established by Brunetti et al. (2013) for Calabria (T_{Cal}), and by Peruccacci et al. (2012) for the Abruzzo, Marche and Umbria ($T_{5,AMU}$) regions, central Italy. Also shown are the global *ED* thresholds proposed by Innes (1983) and by Kanji et al. (2003). The new $T_{5,C}$ threshold and the pre-existing T_{Cal} thresholds are based on different sets of (D , E) empirical data:

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(i) 113 events with rainfall duration $1 \leq D \leq 1080$ h for T_{Cal} , and (ii) 186 events with rainfall duration $1 \leq D \leq 451$ h for $T_{5,C}$. Despite the significant increase in the number of empirical data (73, a 65 % increase), the two thresholds are virtually identical. As discussed before, the sub-regional thresholds for the Tyrrhenian ($T_{5,\text{TAR}}$) and the Ionian ($T_{5,\text{IAR}}$) alert zones are different from the regional threshold ($T_{5,C}$), particularly the $T_{5,\text{IAR}}$ threshold. However, the dataset used to prepare this threshold is small (69 data points), and the difference may be an artefact caused by the scarcity of the data. The new regional threshold for Calabria ($T_{5,C}$) is higher than the threshold obtained for the Abruzzo, Marche and Umbria regions, central Italy by Peruccacci et al. (2012). Based on this result, for the same rainfall duration the cumulated amount of event rainfall necessary to trigger landslides is larger in Calabria than in central Italy. The difference increases with increasing rainfall durations (Fig. 11). Finally, for all the considered rainfall durations, the *ED* threshold for Calabria ($T_{5,C}$) is higher than the threshold proposed by Innes (1983), and significantly lower than the threshold proposed by Kanji et al. (2003).

7 Conclusions

We compiled a catalogue of 186 rainfall events that have resulted in 251 landslides in Calabria, southern Italy, from January 1996 to September 2011. Using this information and rainfall measurements obtained from two synergic networks of rain gauges, we determined new regional cumulated event rainfall – rainfall duration (*ED*) thresholds for possible shallow landslide occurrence in Calabria. The thresholds were determined for the 1 % and the 5 % exceedance probability levels adopting the procedure proposed by Brunetti et al. (2010) and changed by Peruccacci et al. (2012) to evaluate the uncertainty associated to the thresholds. The new regional *ED* thresholds are slightly lower than expected. The result is due to an asymmetric distribution of the empirical (*D*, *E*) points around their mean tendency line (Brunetti et al., 2010). The asymmetry may

be a statistical artefact due to the limited number of data, or it may reflect a specific characteristic of the landslide-triggering rainfall events in Calabria.

Segmentation of the catalogue of rainfall events that have triggered landslides on lithological domains, soil regions, rainfall zones, and seasonal periods resulted in a number of empirical data points in each environmental class that was insufficient to determine reliable thresholds. However, it allowed for two conclusions on the role of the environmental factors on the rainfall conditions responsible for shallow landslides in Calabria. First, conditions for landslide occurrence in different soil regions in Calabria cannot be discriminated currently. Second, for short duration ($D \leq 24$ h) rainfall events, lower amounts of cumulated event rainfall are required to trigger landslides in the Tyrrhenian (western) alert region than in the Ionian (eastern) alert region.

We expect the new regional *ED* rainfall thresholds defined for Calabria, and the sub-regional thresholds defined for the Tyrrhenian and the Ionian alert region to be used in regional and national landslide warning systems (Brunetti et al., 2009; Rossi et al., 2012). The thresholds can also be used for landslide hazard and risk assessments, and for erosion and landscape studies in southern Italy, and in similar physiographic regions in the Mediterranean area.

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Table 1. Distribution of landslides in the five provinces of Calabria. The last three columns list the number of the landslides mapped for each geographical mapping accuracy class (see text).

Province		Area (km ²)	(%)	Landslides			P_1 (#)	P_{10} (#)	P_{100} (#)
				(#)	(# per 1000 km ²)	(%)			
Catanzaro	CZ	2392	15.9	29	12.1	11.6	6	15	8
Cosenza	CS	6651	44.1	151	22.7	60.2	78	62	11
Crotone	KR	1718	11.4	10	5.8	4.0	6	3	1
Reggio di Calabria	RC	3180	21.1	38	11.9	15.1	13	20	5
Vibo Valentia	VV	1139	7.6	23	20.2	9.2	13	2	8
Total		15 080	100.0	251	16.6	100.0	116	102	33

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Table 2. Rainfall ED thresholds for shallow landslides in Calabria, and in other areas in Italy. Number shows the number of rainfall events used to define the threshold. Source: 1 to 6, this work; 7, Peruccacci et al. (2012); 8, Brunetti et al. (2013); 9, Innes (1983); 10, Kanji et al. (2003).

Source	Label	Area	Number	Equation	Validity range (h)	Uncertainty
1	$T_{1,C}$	Calabria	186	$E = 5.8 \cdot D^{0.41}$	$1 \leq D \leq 451$	$\Delta\alpha = 0.8, \Delta\gamma = 0.03$
2	$T_{5,C}$	Calabria	186	$E = 8.6 \cdot D^{0.41}$	$1 \leq D \leq 451$	$\Delta\alpha = 1.1, \Delta\gamma = 0.03$
3	$T_{1,TAR}$	Tyrrhenian Alert Region	117	$E = 6.7 \cdot D^{0.36}$	$1 \leq D \leq 451$	$\Delta\alpha = 1.1, \Delta\gamma = 0.04$
4	$T_{5,TAR}$	Tyrrhenian Alert Region	117	$E = 9.7 \cdot D^{0.36}$	$1 \leq D \leq 451$	$\Delta\alpha = 1.5, \Delta\gamma = 0.04$
5	$T_{1,IAR}$	Ionian Alert Region	69	$E = 6.3 \cdot D^{0.50}$	$1 \leq D \leq 357$	$\Delta\alpha = 1.4, \Delta\gamma = 0.05$
6	$T_{5,IAR}$	Ionian Alert Region	69	$E = 9.0 \cdot D^{0.50}$	$1 \leq D \leq 357$	$\Delta\alpha = 1.8, \Delta\gamma = 0.05$
7	$T_{5,AMU}$	Abruzzo, Marche, Umbria	442	$E = 7.4 \cdot D^{0.38}$	$1 \leq D \leq 1212$	$\Delta\alpha = 0.5, \Delta\gamma = 0.02$
8	$T_{5,Cal}$	Calabria	113	$E = 7.85 \cdot D^{0.43}$	$1 \leq D \leq 1080$	n.a.
9	Innes	Global threshold	35	$E = 4.9355 \cdot D^{0.5041}$	$0.2 \leq D \leq 100$	n.a.
10	Kanji	Global threshold	n.a.	$E = 22.4 \cdot D^{0.41}$	n.a.	n.a.

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Table 3. Descriptive statistics for landslides and for the corresponding rainfall events in Calabria in the period January 1996–September 2011. Rows: RE, number of rainfall events; LE, number of landslide events; *D*, rainfall duration (h); *E*, cumulated event rainfall (mm); *I*, rainfall mean intensity (mm h^{-1}). Columns: CC, carbonate rocks; FD, flysch deposits; MC, metamorphic and crystalline rocks; PO, post-orogenic sediments; SR1, Pollino range and the neighbouring hills; SR2, northern Ionian coastal zone; SR3, coastal hills and plains; SR4, Sila, Serre and Aspromonte ranges; T, Tyrrhenian rainfall region; C, Central rainfall region; I, Ionian rainfall region; TAR, Tyrrhenian alert region; IAR, Ionian alert region; A–O, April–October seasonal period; N–M, November–March seasonal period.

	Extent (km^2)	RE	LE	<i>D</i> (h)			<i>E</i> (mm)			<i>I</i> (mm h^{-1})		
				Min	Max	Mean	Min	Max	Mean	Min	Max	Mean
Calabria	15 080	186	251	1	1128	98.9	12.8	542.0	135.9	0.30	66.40	4.30
Lithological domains												
CC	2054	11	13	7	451	75.4	42.0	273.6	136.2	0.61	9.58	4.32
FD	4514	23	24	2	310	101.0	13.0	542.0	187.0	0.75	7.41	2.90
MC	4618	39	49	1	357	75.1	12.8	477.6	135.2	0.63	29.04	4.68
PO	3894	46	56	1	451	68.2	16.8	312.4	98.4	0.34	66.40	7.05
Total	15 080	119	142									
Soil regions												
SR1	1175	9	12	2	176	55.7	24.2	206.3	82.7	1.11	12.10	4.45
SR2	618	0	0									
SR3	7575	102	130	1	451	82.3	13.0	542.0	131.9	0.34	23.20	4.30
SR4	5712	46	57	1	357	81.5	12.8	477.6	157.4	0.63	20.44	4.03
Total	15 080	157	199									
Rainfall regions												
T	6681	81	117	1	451	90.3	12.8	309.4	98.5	0.30	451.0	90.3
C	3886	49	63	1	290	54.4	13.6	477.6	148.5	0.51	29.04	6.10
I	4513	28	32	2	282	78.4	13.0	542.0	196.4	0.50	18.53	4.77
Total	15 080	158	212									
Dry/wet seasonal periods												
A–O		60	80	1	290	38.1	13.0	542.0	106.3	0.50	66.40	7.30
N–M		126	171	3	451	98.4	12.8	433.0	141.2	0.34	29.04	3.06
Total		186	251									
Alert regions												
TAR	6122	117	157	1	451	82.9	12.8	357.4	109.0	0.30	66.40	4.10
IAR	8958	69	88	1	357	72.5	13.0	542.0	163.8	0.50	20.40	5.06
Total	15 080	186	245									

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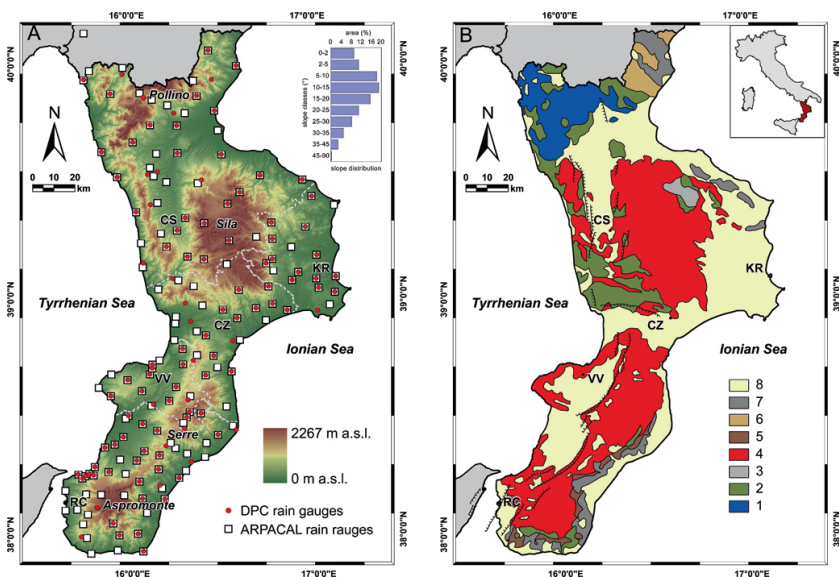


Fig. 1. (A) Terrain elevation map for Calabria, southern Italy. Shades of colour show terrain elevations from sea level (dark green) to 2267 m (brown). Histogram shows distribution of terrain slope. Red dots show location of rain gauges of the National Department of Civil Protection. White squares show location of rain gauges of the Regional Agency for Environmental Protection in Calabria. White dashed lines are province boundaries. Legend: CS, Cosenza; CZ, Catanzaro; KR, Crotona; RC, Reggio di Calabria; VV, Vibo Valentia. **(B)** Main lithological domains in Calabria (modified after Sorriso-Valvo, 1993). Legend: (1) carbonate units; (2) very low to low-grade metamorphic units, with ophiolites; (3) sedimentary, flysch-type sequences; (4) intermediate to high-grade metamorphic and intrusive rocks; (5) turbiditic, mainly coarse-grained deposits; (6) flysch-type nappes mainly marly-clayey; (7) flysch-type nappes with chaotic structure; (8) sedimentary autochthonous units. Hachured black lines show main tectonic structures pertaining to the active Calabrian-Sicilian Rift Zone (modified after Monaco and Tortorici, 2000).

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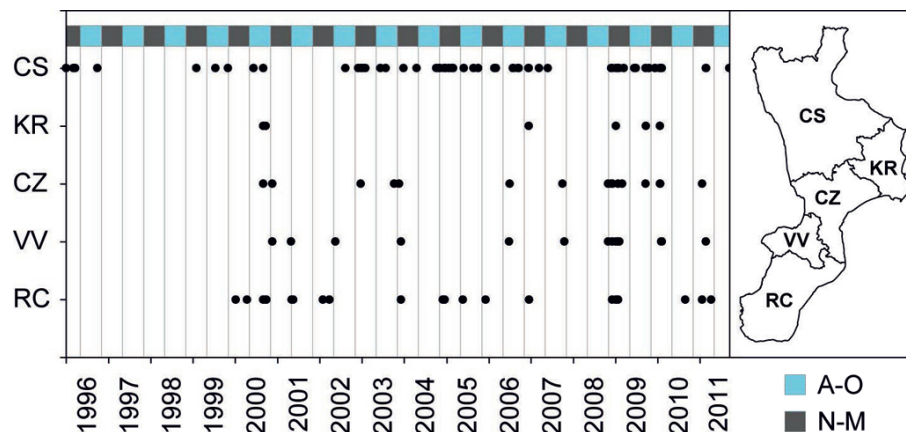


Fig. 2. Time distribution of 186 rainfall events that have resulted in 251 landslides in the 16 yr period from January 1996 to September 2011 in the five provinces of Calabria. Legend: A–O = April to October seasonal period, N–M = November to March seasonal period. CS, Cosenza; CZ, Catanzaro; KR, Crotone; RC, Reggio di Calabria; VV, Vibo Valentia.

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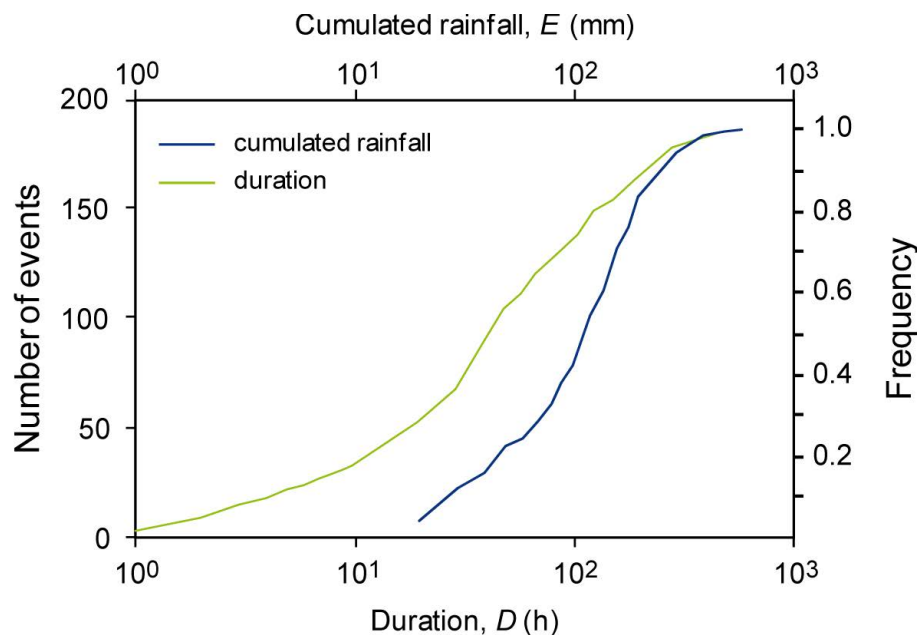


Fig. 3. Frequency curves of the rainfall duration D (in h) and the cumulated event rainfall E (in mm) for the 186 rainfall (D , E) conditions that have resulted in shallow landslides in Calabria in the 16 yr period from January 1996 to September 2011.

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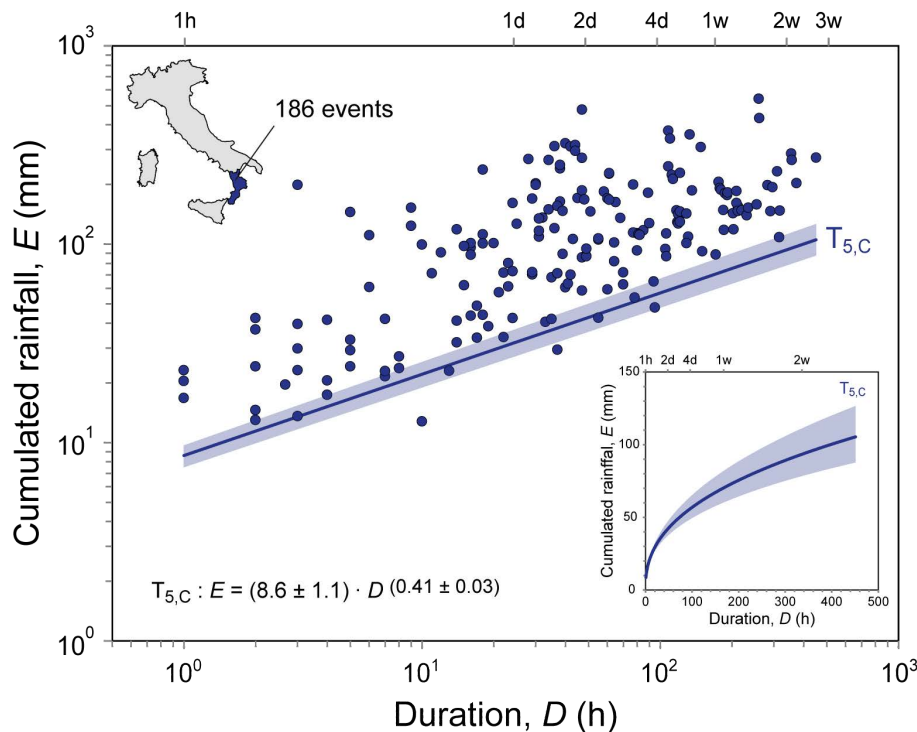


Fig. 4. Rainfall duration vs. cumulated event rainfall (D , E) conditions that have resulted in shallow landslides in Calabria (blue dots), and related ED threshold at the 5% exceedance probability level. Inset shows the threshold in linear coordinates. Shaded areas show uncertainties associated to the threshold.

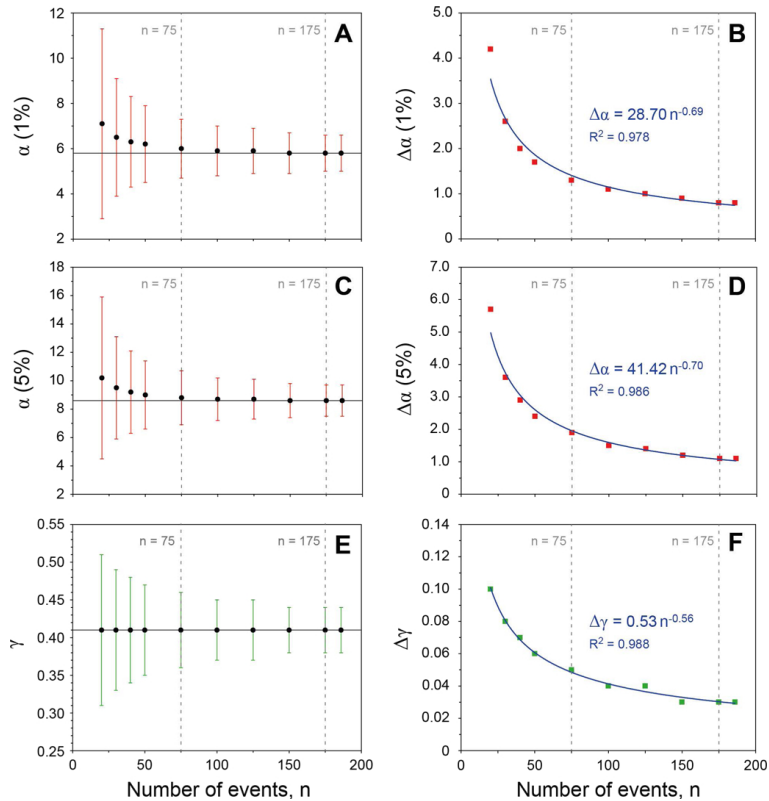


Fig. 5. Variation of the threshold model parameters (α and γ) and of their uncertainties ($\Delta\alpha$ and $\Delta\gamma$) as a function of the number of events n , for the 1% and the 5% exceedance probability levels. In the left plots, black dots show the average values of α (A, C) and γ (E), and the coloured bars show the relative variations. In the right plots, coloured squares represent the values of $\Delta\alpha$ (B, D) and $\Delta\gamma$ (F), and the blue lines are best fits (linear regression method) to the empirical data.

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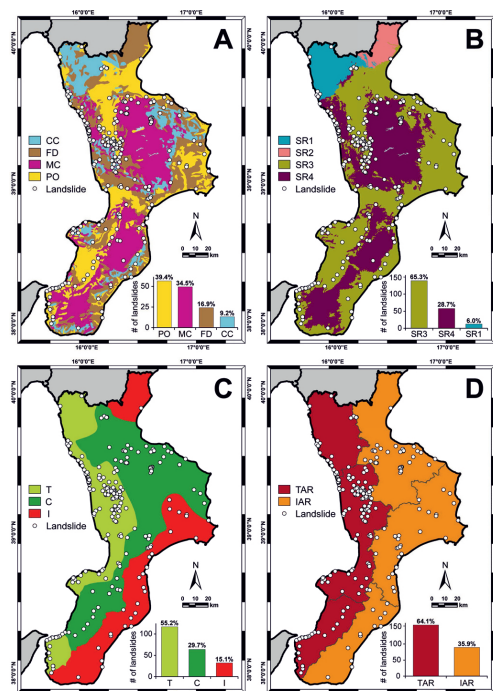


Fig. 6. Geographical distribution of shallow rainfall-induced landslides (white dots) in Calabria, in the 16 yr period from January 1996 to September 2011, superimposed on maps of different environmental factors. Histograms show the number and the proportion of shallow landslides for each factor class. **(A)** Lithological domains: CC, carbonate rocks; FD, flysch deposits; MC, metamorphic and crystalline rocks; PO, post-orogenic sediments. **(B)** Soil regions: SR1, Pollino range and neighbouring hills; SR2, northern Ionian coastal zone; SR3, coastal hills and plains; SR4, Sila, Serre and Aspromonte ranges. **(C)** Rainfall regions: T, Tyrrhenian region; C, Central region; I, Ionian region. **(D)** Civil Protection alert regions: TAR, Tyrrhenian alert region; IAR, Ionian alert region.

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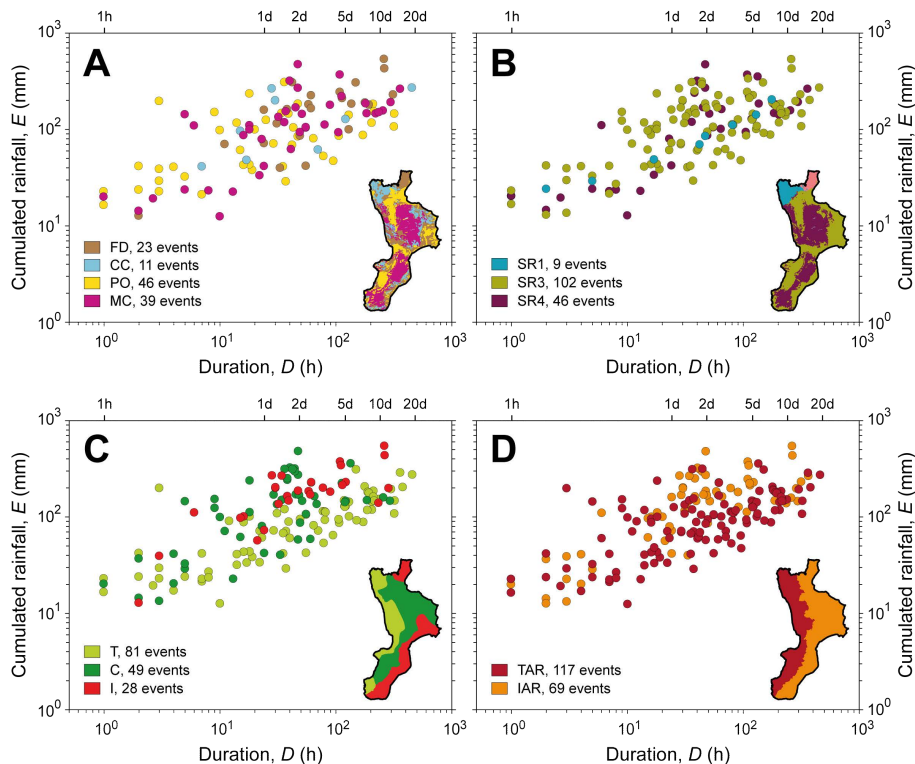


Fig. 7. Rainfall duration vs. cumulated event rainfall (D , E) conditions that have resulted in shallow landslides in Calabria for different environmental factors. **(A)** Four lithological domains. **(B)** Four soil regions. **(C)** Three rainfall regions. **(D)** Two Civil Protection alert regions. See Fig. 6 for location of the environmental factors and for colour legends.

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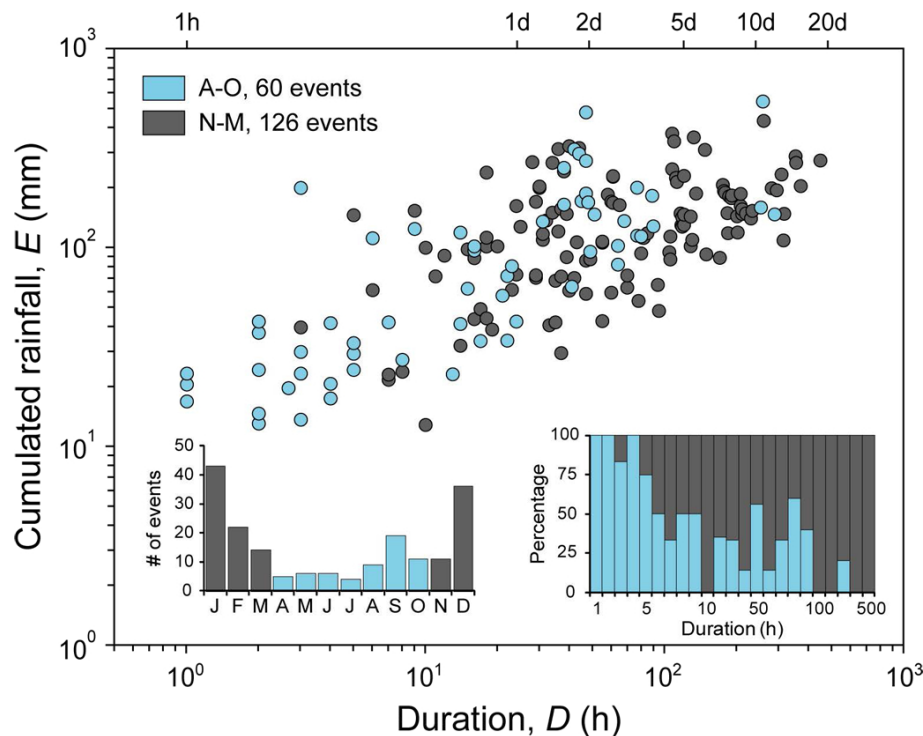


Fig. 8. Rainfall duration vs. cumulated event rainfall (D , E) conditions that have resulted in shallow landslides in Calabria between January 1996 and September 2011 for the April–October seasonal period (A–O, light blue dots) and for the November–March seasonal period (N–M, dark grey dots). Bar charts show the number of events per month (left), and the percentage of events for different durations (right), for the two periods.

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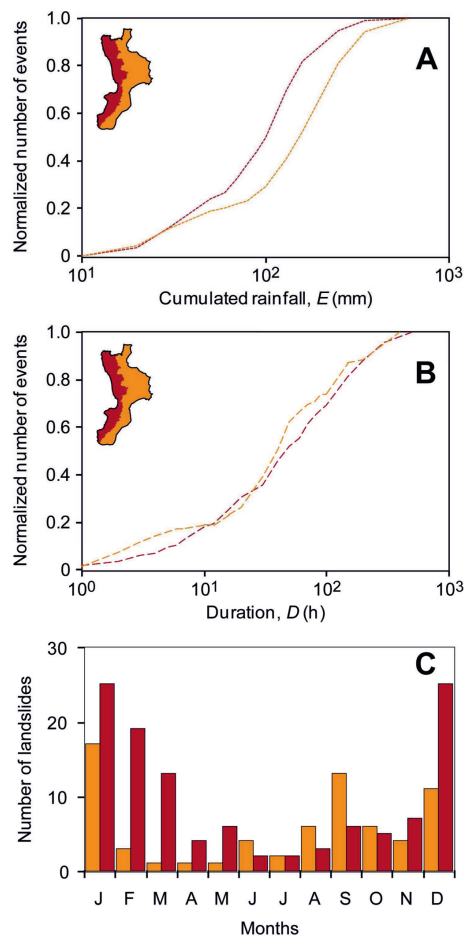


Fig. 9. Frequency of **(A)** the cumulated event rainfall and **(B)** the duration of the events, for the Tyrrhenian and the Ionian alert regions. **(C)** Monthly distribution of landslides in the two alert regions.

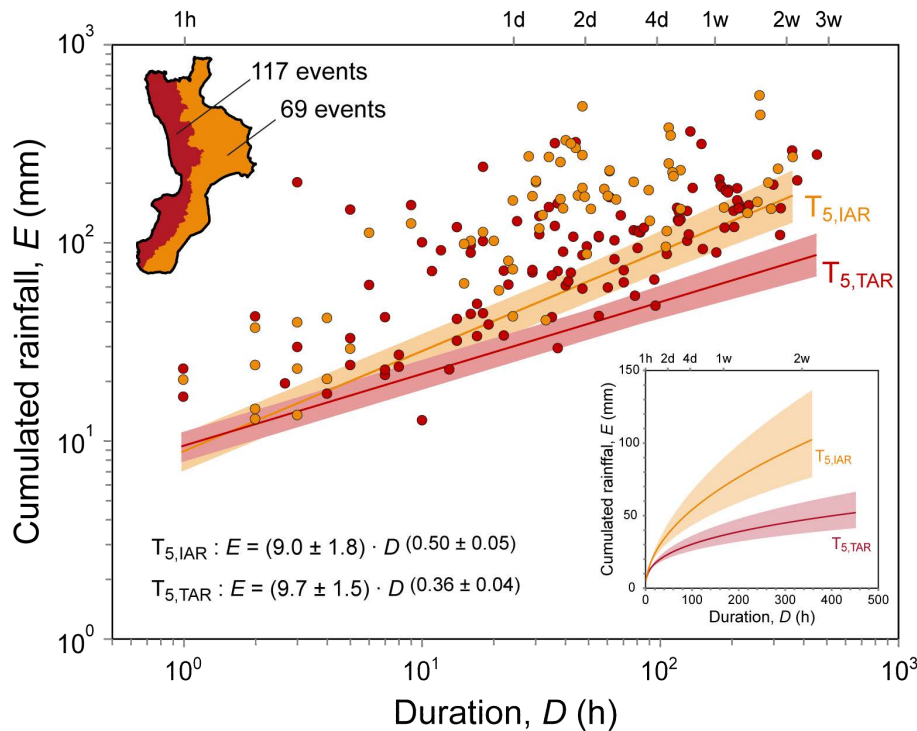


Fig. 10. Rainfall duration vs. cumulated event rainfall (D , E) conditions that have resulted in shallow landslides in Calabria in the Ionian alert region (orange dots) and in the Tyrrhenian alert region (red dots), and related ED thresholds at the 5% exceedance probability levels ($T_{5,IAR}$ and $T_{5,TAR}$). Inset shows the thresholds in linear coordinates. Shaded areas show uncertainties associated to the thresholds.

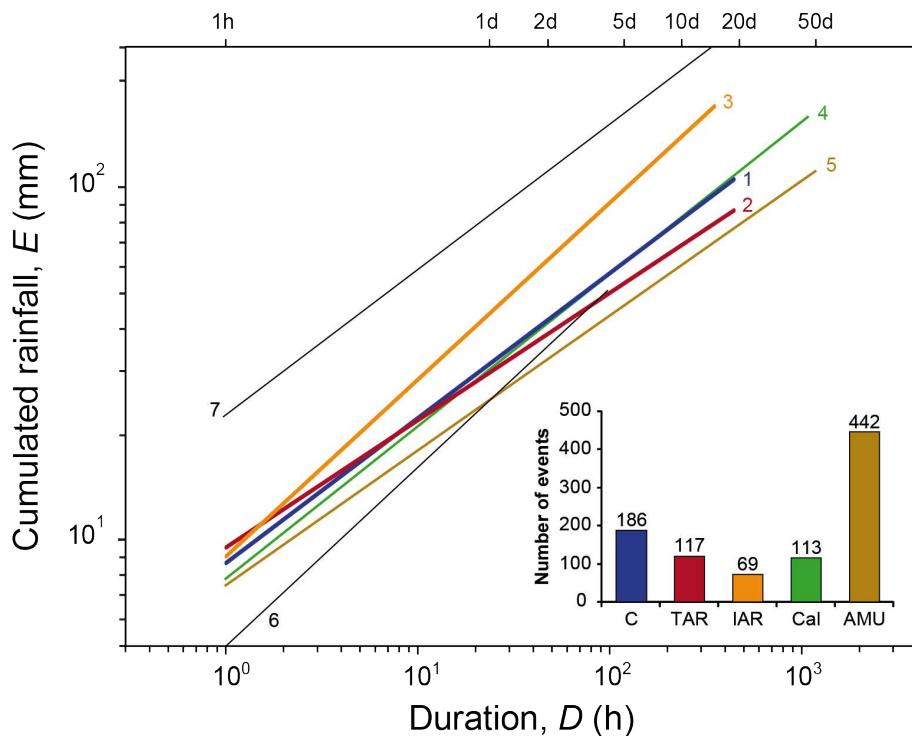


Fig. 11. Cumulated event rainfall – rainfall duration (ED) thresholds for shallow landslide in Calabria: $T_{5,C}$ (line 1), $T_{5,TAR}$ (line 2), $T_{5,IAR}$ (line 3), $T_{5,Cal}$ (line 4). $T_{5,Cal}$ is the 5% frequentist threshold proposed for Calabria by Brunetti et al. (2013). $T_{5,AMU}$ (line 5) is the 5% frequentist threshold proposed for Abruzzo, Marche and Umbria (central Italy) by Peruccacci et al. (2012). Black line 6 is the global threshold for debris flows of Innes (1983). Black line 7 is the global thresholds of Kanji et al. (2003). Bar chart shows number of rainfall events that have resulted in landslides in different study areas in Italy. Equations for the thresholds are given in Table 2.