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Spatial impact and triggering conditions of the exceptional hydro-geomorphological event of December 1909 in Iberia

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Abstract. According to the DISASTER database the 20–28 December 1909 event was the hydro-geomorphologic event with the highest number of flood and landslide cases that occurred in Portugal in the period 1865–2010 (Zêzere et al., 2014). This event also caused important social impacts over the Spanish territory, especially in the Douro Basin, having triggered the highest floods in more than 100 years at the river's mouth in the city of Oporto.

This work has a dual purpose: (i) to characterize the spatial distribution and social impacts of the December 1909 hydro-geomorphologic DISASTER event over Portugal and Spain; (ii) to analyse the meteorological conditions that triggered the event and the spatial distribution of the precipitation anomalies.

Social impacts that occurred in Portugal were obtained from the Disaster database (Zêzere et al., 2014) whereas the data collection for Spain was supported by the systematic analysis of Spanish daily newspapers. In addition, the meteorological conditions that triggered the event are analysed using the 20th Century Reanalysis data set from NOAA and precipitation data from Iberian meteorological stations.

The Iberian Peninsula was spatially affected during this event along the SW–NE direction spanning from Lisbon, Santarém, Oporto, and Guarda (in Portugal), to Salamanca, Valladolid, Zamora, Orense, León, and Palencia (in Spain). In Iberia, 134 DISASTER cases were recorded (130 flood cases; 4 landslides cases) having caused 89 casualties (57 due to floods and 32 due to landslides) and a further total of 3876 affected people, including fatalities, injured, missing, evacuated, and homeless people. This event was associated with outstanding precipitation registered at Guarda (Portugal) on 22 December 1909 and unusual meteorological conditions characterized by the presence of a deep low-pressure system located over the NW Iberian Peninsula with a stationary frontal system striking the western Iberian Peninsula. The presence of an upper-level jet (250 hPa) and low-level jet (900 hPa) located SW–NE oriented towards Iberia along with upper-level divergence and lower-level convergence favoured large-scale precipitation. Finally, associated with these features it is possible to state that this extreme event was clearly associated with the presence of an elongated Atmospheric River, crossing the entire northern Atlantic Basin and providing a continuous supply of moisture that contributed to enhance precipitation.

This work contributes to a comprehensive and systematic synoptic evaluation of the second most deadly hydrogeomorphologic DISASTER event that has occurred in Portugal since 1865 and will help to better understand the meteorological system that was responsible for triggering the event.

1 Introduction

In the Iberian Peninsula, extreme precipitation events that occur during winter (December-January-February-March) have been historically associated with progressive flooding events in the major rivers (Salgueiro et al., 2013), and exceptionally with flash floods that typically occur in small watersheds or urban areas (Llasat et al., 2005; Fragoso et al., 2012) and also landslide events (Zêzere and Trigo, 2011; Zêzere et al., 2014, 2015). These events have been the source of major socio-economic impacts, with human and material damage (Barriendos and Rodrigo, 2006; García et al., 2007; Vicente-Serrano et al., 2011; Trigo et al., 2014, 2015).

Regularly, these extreme events have been evaluated based on case studies, like for instance the large flood in the Guadiana River in 1876 (Trigo et al., 2014), the November 1967 flash flood in the Lisbon region (Trigo et al., 2015), the 10 June 2000 flash flood in Catalonia (Llasat et al., 2003), and the 20 February 2010 flash floods in Madeira (Fragoso et al., 2012). Major floods that occurred in the Iberian Peninsula were extensively analysed in the Tagus Basin (Benito et al., 2003; Salgueiro et al., 2013; Machado et al., 2015) and to a lesser extent in the Douro Basin (e.g. Morán-Tejeda et al., 2012).

Recently, а long-term database of hydrogeomorphological cases was developed for Portugal covering the period 1865-2010 in the project DISASTER (Zêzere et al., 2014). This database accounts for 1621 flood cases and 281 landslide cases that were responsible for 1251 fatalities and more than 41000 homeless people (Zêzere et al., 2014). This database includes social consequences (fatalities, injured, missing, evacuated, and homeless people) caused by flood and landslide cases referred to in Portuguese newspapers during the period of 1865-2010. The DISASTER database is the only database containing detailed data on the social impacts of hydrogeomorphologic disasters in Portugal as it accounts for any flood or landslide that, independently of the number of affected people, caused casualties or injuries, or missing, evacuated, or homeless people. Therefore, each DISASTER case is a unique hydro-geomorphologic occurrence, which is related to a unique geographic location and a specific period of time (e.g. the place where the flood or landslide harmful consequences occurred on a specific date). Within this database, a DISASTER event is a set of flood and landslide cases sharing the same trigger mechanism and thus clustered in the temporal domain, which may have a widespread spatial extension and a certain magnitude (Zêzere et al., 2014).

The December 1909 flood in Portugal is considered the most extraordinary flood of the 20th century along the Douro River since systematic records of the river flow began (Rodrigues et al., 2003). This particular flood also had a strong contribution from the Spanish area of the Douro Basin, at a time when the hydrologic regime was only conditioned by natural factors due to the absence of dams to control the river flow (Rodrigues et al., 2003). It is worth stressing that after the flash flood of November 1967 (with circa 500 deaths) the 1909 flood was the second deadliest Disaster event in Portugal in the period 1865–2010, and generated 35 fatalities (29 caused by floods and six caused by landslides). In addition, according to the DISASTER database the 20–28 December 1909 was the DISASTER event with the highest number of flood and landslide cases (67 flood cases and three landslide cases) occurring in Portugal in the period 1865–2010 (Zêzere et al., 2014). As we show later in the descriptive analysis there were an even larger number of fatalities in Spain (54), raising the total number of deaths to 89.

The societal impacts associated with the 1909 event were particularly high over the NW Iberian Peninsula (Fig. 1). Thus, while most observed floods were concentrated in Portuguese and Spanish provinces of the Douro Basin, other areas were equally affected, including the Minho Basin and Tagus valley (Santarém), Lisbon region and also the northern sector of the Tagus Basin in the Spanish territory (Fig. 1).

Besides the intrinsic relevance of this extreme episode from the impact perspective, this event was associated with some outstanding precipitation values and unusual meteorological conditions. However, to the best of our knowledge this extreme hydro-meteorological episode has never been evaluated in detail, namely by Portuguese and Spanish authorities or research teams. This lack of previous assessments resulted in our view from several distinct factors, including (1) lack of sufficient long-term time series of precipitation covering most of the affected territory, (2) lack of surface and upper-air meteorological fields capable of describing the most important atmospheric circulation features of the episode, and (3) the inexistence of a list of places with the affected people and socio-economic impacts for the complete event over Iberia. These three shortcomings result from the time that has passed since the event in the early 20th century. However, these caveats have been minimized to a large extent in recent years. Thus, a number of initiatives in Portugal and Spain have favoured the digitalization of time series from the late 19th century and beginning of 20th century, allowing the assessment of long-term changes in daily precipitation (Gallego et al., 2011; Cortesi et al., 2014), or related to major historical floods and extreme precipitation events (Trigo et al., 2014; Domínguez-Castro et al., 2015). Likewise, a new type of long-term reanalysis has been developed (Compo et al., 2011), and despite being based on surface data, it is capable of reproducing upper-air fields (e.g. Trigo et al., 2014). Finally, as stated previously, a comprehensive list of people affected (fatalities, injured, displaced, evacuated, and disappeared) has been obtained through the DISASTER database (Zêzere et al., 2014) although covering only the Portuguese territory.

In this context, this work has a dual purpose: (i) to characterize the spatial distribution and social impacts of the December 1909 hydro-geomorphologic DISASTER event over Portugal and Spain; (ii) to analyse the meteorological conditions that triggered the event and the spatial distribution of the precipitation anomalies. We aim at characterizing the meteorological synoptic forcing of a major hydrogeomorphologic event that occurred in the last 150 years and registered in the DISASTER database. However, besides this meteorological/climatological assessment we are equally interested in obtaining a full characterization of the associated human cost and the spatial extent of associated social im-



Figure 1. Study area and flood and landslide DISASTER cases from 20–28 December 1909 event over Iberia. Source: Disaster database for Portugal and Spanish newspapers (see Table 1). Numbers in the map correspond to the affected hydrographic basins: (1) Coruña basins; (2) Minho Basin; (3) Lima and Cávado basins; (4) Douro Basin; (5) Águeda, Mondego, and West basins; (6) Tagus Basin; (7) Sado Basin.

pacts, as such evaluation has never been done before. Overall, we expect to gain a better understanding of the climatic driving forces of hydro-geomorphologic disaster events and corresponding societal impacts over Iberia.

2 Study area and precipitation regime

The Iberian Peninsula is located southwest of Europe and includes the countries of Andorra, Portugal, and Spain. It is surrounded by the Mediterranean Sea (on the southeast and east), the Atlantic Ocean (on the west), and the Cantabrian Sea (on the north). The central part of the Iberian Peninsula is occupied by the Meseta, a vast plateau (Douro and Tagus Cainozoic basins) that is characterized by hilly and smooth terrain. The central Meseta is surrounded by mountains (e.g. Cantabrian Mountains, Pyrenean Mountains, the Betic Range, and Central Mountains) and is sculptured by the fluvial network (Santisteban and Schulte, 2007) (Fig. 1).

There are four large international Iberian rivers (Minho, Douro, Tagus, and Guadiana), all flowing predominantly along the NE–SW direction, with their lower courses in Portugal (Fig. 1). The seasonal precipitation regime in the Iberian Peninsula controls the major river flow regime (Trigo et al., 2004) like in other Mediterranean regions, where most of the river flow is registered in winter and spring time (Daveau, 1998).

In the Iberian Peninsula, there is a high spatial and temporal variability of precipitation (Serrano et al., 1999; Trigo et al., 2004; Cortesi et al., 2013) due to its location in the southern rim of the Atlantic zonal circulation (Trigo and Palutikof, 2001). The successive passage of frontal systems from west to east over the North Atlantic Basin (most frequent in the winter) produces continuous and persistent daily precipitation, mostly in the western Iberian Peninsula, which can cause major river floods (Salgueiro et al., 2013). Precipitation distribution over the Iberian Peninsula depends also on the wind trajectory (SW, W, or NW) and on the local orographic effects (Goodess and Jones, 2002; Salgueiro et al., 2013; Cortesi et al., 2014) that influence precipitation intensity and type (snow vs. hail or rainfall) at the regional scale.

Numerous studies have shown that the NAO has a strong impact on the occurrence of extreme flooding in the major rivers, namely along the Tagus River (Benito et al., 2003; Silva et al., 2012; Salgueiro et al., 2013; Machado et al., 2015) and the Guadiana River (Ortega and Garzón, 2004).

In addition, Ramos et al. (2015) show that the presence of the so-called Atmospheric River (AR) structure over the Iberian Peninsula can enhance the precipitation amount, as was the case of another outstanding historical flood event in southern Iberia in 1876 (Trigo et al., 2014). Therefore this will be also taken into account when analysing the 1909 hydro-geomorphologic DISASTER event. ARs are identified as narrow plumes of enhanced moisture transport that are usually present in the core of the warm conveyor belt and occur on the oceans along the warm sector of extratropical cyclones (Ramos et al., 2015). They are usually W-E (or SW–NE) oriented and steered by pre-frontal low-level jets along the trailing cold front and subsequently feed the precipitation in the extra-tropical cyclones. The large quantity of water vapour within the AR can generate heavy precipitation and floods when they make landfall (Dettinger, 2011; Lavers et al., 2011). Regarding the Iberian Peninsula, the

Newspaper	Reference period	Category	Distribution	Spatial incidence			
			Portugal				
Diário de Notícias	Since 1865	Daily	National	Mainly the metropolitan area of Lisbon and the Tagus valley region			
Jornal de Notícias	Since 1888	Daily	Regional	North region (mainly the metropolitan area of Oporto)			
Correio de Mirandela	1907–1937	Weekly	Local	North region (Trás-os-Montes)			
Spain							
La Época	1849–1936	Daily	National	Madrid			
La Correspondencia de España	1859–1925	Daily	National	Madrid			
El Imparcial	1867–1933	Daily	National	Madrid			
El Siglo Futuro	1875–1936	Daily	National	Madrid			
El Liberal	1879–1939	Daily	Regional	Madrid, Seville, Barcelona, Bilbao, Murcia			

 Table 1. Newspapers explored for the December 1909 event data collection.

impact of ARs in extreme precipitation has been assessed systematically (Ramos et al., 2015) but also in relation to several well-known events of floods and extreme precipitation with major socio-economic consequences (Couto et al., 2012; Liberato et al., 2013). There are several methodologies to identify ARs (Gimeno et al., 2014), here we adopted the method requiring the existence of a long (> 2000 km length) and narrow (< 500 km width) band of water vapour at low levels from the subtropics (> 5 g kg⁻¹ of specific humidity at 900 hPa), while wind speed should be considered above 12.5 ms⁻¹ (Lavers et al., 2011).

3 Materials and methods

3.1 Historical data sources

The data source used to analyse the event of December 1909 in Portugal was the DISASTER Database (Zêzere et al., 2014). The data collection process used to construct the DIS-ASTER database was based on painstaking analysis of numerous daily and weekly newspapers, and is explained in full detail in Zêzere et al. (2014).

In 1909 there were available only four newspapers for data collection in Portugal (Zêzere et al., 2014). The national daily newspaper "Diário de Notícias" published since 1865 in Lisbon provided the longest time period to data collection. As well as this daily newspaper the "Jornal de Notícias", published in Oporto (North Portugal) since 1888, was also analysed. The remaining two newspapers are weekly local and regional newspapers, one located in the North ("Correio de Mirandela") and another one in the South of the country ("O Algarve: o semanário independente").

In Spain there is no equivalent to the DISASTER database built with similar criteria in order to allow data comparison between the two countries. As a consequence, we have adopted the same methodology and criteria used in the DIS- ASTER database to ensure data integrity and to enable data comparison. The data collection process was supported by the systematic analysis of Spanish daily newspapers of the entire month of December 1909. Four daily national newspapers ("La Época", "La Correspondencia de España", "El Imparcial", "El Siglo Futuro") published in Madrid were used for data collection and a regional newspaper ("El Liberal") with spatial incidence in Madrid, Sevilla, Barcelona, Bilbao and Murcia was also explored (Table 1). In summary, we can state that, although not included in the original DISASTER database, all locations selected in Spain relative to this event fulfil the DISASTER requirements as described in Zêzere et al. (2014). In this regard they can be considered DISASTERlike cases and the use of this denomination will be applied indiscriminately throughout this work, to both cases and locations in Portugal and Spain.

Each DISASTER case includes details on process characteristics and associated damages. The first includes data on type (flood or landslide), subtype, date, location, triggering factor, and data source. In addition, the following damage data were recorded: number of casualties, injuries, missing, evacuated or homeless people, type of damages in buildings, number of affected buildings, type of damage in road and railroad, extent of interruptions in road and railroad circulation (Zêzere et al., 2014). Information about the gender of the fatalities was also obtained from the content analysis of the newspapers as well as the circumstances surrounding the fatalities (Pereira et al., 2015). The latter were classified into four classes according to the location where the fatalities occurred - namely, inside a building in a rural area; inside a building in an urban area; outdoors in a rural area; outdoors in an urban area.

Location of DISASTER cases in both countries was obtained using a point shapefile. The precision of the location depends on the detail of the news and was classified in the following classes: (i) exact coordinates (accuracy associated with scale 1 : 1000); (ii) local toponymy (accuracy associated with scale $1:10\,000$; (iii) local geomorphology (accuracy associated with scale $1:25\,000$ scale); (iv) centroid of the parish; and (v) centroid of the municipality (Zêzere et al., 2014).

Additionally, in order to evaluate the relative level of damages along the affected regions during this historical event, three damage indices (fatalities index, homeless index, and total affected index) were computed, taking into account the resident population per region obtained from the Census data in Portugal (1911) and Spain (1910):

Fatalities index =
$$\frac{\text{No. of fatalities}}{\text{Resident population}} \cdot 100\,000$$
 (1)

Homeless index =
$$\frac{\text{No. of homeless people}}{\text{Resident population}} \cdot 100\,000$$
 (2)

Total affected index =
$$\frac{\text{No. of total affected people}}{\text{Resident population}} \cdot 100\,000.$$
 (3)

The number of total affected people includes the total number of fatalities, injuries, missing people, evacuated people, and homeless people. Damage indices were mapped and classified into five classes (low, moderate, high, very high, extreme) according to natural breaks of index distribution. Natural breaks classes are based on natural groupings inherent in the data that best group similar values and maximize the differences between classes. With this methodology each affected region ("Distrito" in Portugal and "Provincia" in Spain) was classified according to the fatalities index, homeless index, and total affected index. These indices are important to compare the societal impacts of floods and landslides observed in different regions; in addition these indices can be applied to other historical events in other study areas.

3.2 Precipitation data

Precipitation data for 27 meteorological stations of Iberia (Fig. 1) used here is summarised in Table 2 including the name, elevation (in m) mean annual precipitation (MAP), mean monthly precipitation (MMP) in December, and total precipitation in 1909 (December, 22 December, and 20– 28 December).

Seven meteorological stations are located in Portugal (Beja, Campo Maior, Coimbra, Évora, Guarda, Lisbon, and Porto) and were digitized by the Geophysical Institute Infante Dom Luiz (IDL) within the ERA-CLIIM project (Stickler et al., 2014). In addition, 20 meteorological stations are located in Spain (Albacete, Alicante, Badajoz, Barcelona, Burgos, Cádiz, Ciudad Real, Granada, Huelva, Huesca, La Coruña, Madrid, Málaga, Murcia, Salamanca, Sevilla, Soria, Valencia, Valladolid, and Zaragoza) and were digitized in the framework of the project EMULATE which are the most complete and reliable daily precipitation records extending back to the mid-19th century (Brunet et al., 2007). A large fraction of these stations were used in a long-term assessment of tendencies in the frequency of days characterised by low, medium, and intense precipitation (Gallego et al., 2011).

Long-term daily precipitation data of Oporto (1906–2007), Guarda (1906–1994), Lisbon (1864–2013), and Salamanca (1894–2003) stations were further explored in detail considering the period from September to December. For the complete set of stations, climatological reference periods correspond to the 1971–2000 normal in Portugal (IPMA, Portuguese Meteorological Office) and to the 1981–2010 normal in Spain (AEMet, Spanish Meteorological Office).

3.3 Synoptic data

The atmospheric circulation during the December 1909 event and prior months was evaluated at different timescales. This was achieved with the recently developed 20th Century Reanalysis version 2 (Compo et al., 2011), hereafter 20 CR.

The 20 CR data set has been shown to be suitable for long-term assessment over Iberia, such as major floods event in 1876 (Trigo et al., 2014) or the decadal change of wind resources (Kirchner-Bossi et al., 2014). Several fields related to both surface and tropospheric levels were extracted. From the surface the 6 h sea level pressure field (hPa) and 6 h accumulated precipitation were used). In addition the following variables were also extracted and analysed: (i) for the lower levels (900 hPa), the 6 h fields of wind vector (m s⁻¹) and specific humidity (g kg⁻¹), and (ii) for the upper levels (500 and 250 hPa) the 6 h fields of wind vector (m s⁻¹), geopotential height, and temperature. The ensemble mean fields for all variables were assessed using a global grid of $2^{\circ} \times 2^{\circ}$.

3.4 NAO data

NAO data of monthly mean normalized difference of sea level pressures between Lisbon (Portugal) and Stykkisholmur/Reykjavik (Iceland) (Hurrell, 1995) were obtained from the National Centre of Atmospheric Research (NCAR) covering the period from 1865 onwards.

4 Results

4.1 Societal impacts of the December 1909 event

In Iberia, 134 DISASTER cases were recorded – 130 flood cases and four landslides cases (Fig. 1). Table 3 summarizes the social consequences of the December 1909 event in Iberia. Considering the data sources from both affected countries, floods that occurred during the December 1909 event caused 57 fatalities, 4 injured, 15 missing people, 1104 evacuated people, and 2564 homeless people (Table 3). Landslides (debris flows) caused 32 fatalities and 100 homeless people (Table 3). In total this event affected 3876 people in the Iberian Peninsula.

In this section the spatial distribution of the societal impacts will be explored per hydrographic basin whereas the results of the damage indices are presented per region. The temporal distribution of the DISASTER cases per day (20–

Rain	Elevation	MAP	MMP in	TP in	TP (mm)	TP (mm)			
gauge	(m)	(mm)	December	December	in 22	20-28			
			(mm)	1909 (mm)	December	December			
					1909	1909			
			Portugal						
Beia	246	571.8	100.6	132.9	0.0	58.9			
Campo Maior	280	*	*	125.1	2.0	57.3			
Coimbra	141	905.1	126.8	225.3	31.3	83.9			
Évora	200	609.4	102.7	172.2	1.5	85.8			
Guarda	1019	882.0	141.8	451.1	185.5	343.9			
Lisbon	77	725.8	121.8	139.1	60.9	84.0			
Oporto	93	1253.5 194.7 287.0 6.		63.0	158.6				
Spain									
Albacete	699	353.0	31.0	14.5	0.0	0.0			
Alicante	82	311.0	25.0	12.6	0.0	0.0			
Badajoz	185	447.0	69.0	129.8	0.0	51.3			
Barcelona	420	588.0	40.0	13.9	0.0	0.0			
Burgos	881	546.0	63.0	119.2	23.2	53.0			
Cádiz	30	523.0	92.0	81.7	9.1	27.5			
Ciudad Real	627	402.0	59.0	70.0	6.0	22.6			
Granada	685	365.0	56.0	45.8	10.0	17.0			
Huelva	19	525.0	99.0	48.8	5.7	13.1			
Huesca	541	480.0	44.0	52.6	0.0	16.9			
La Coruña	67	1014.0	131.0	166.1	13.0	53.2			
Madrid	679	421.0	51.0	88.8	0.5	32.0			
Málaga	7	534.0	100.0	49.1	12.8	15.5			
Murcia	57	297.0	29.0	11.6	0.0	0.0			
Salamanca	790	372.0	42.0	80.7	20.6	36.4			
Sevilla	31	539.0	99.0	69.3	19.2	36.3			
Soria	1083	512.0	50.0	80.4	7.5	25.9			
Valencia	11	475.0	48.0	33.3	0.0	0.0			
Valladolid	691	433.0	53.0	137.0	18.0	51.5			
Zaragoza	245	322.0	21.0	8.2	0.0	0.0			

Table 2. Available precipitation stations for December 1909 event in Iberia.

Note: MAP – mean annual precipitation; MMP – mean monthly precipitation; TP – total precipitation; * Campo Maior station was only in operation between December 1872 and December 1909.

28 December) will also be analysed as well as the surrounding circumstances of fatalities.

4.1.1 Spatial distribution of the impacts

The Iberian Peninsula was spatially affected during this event along the SW–NE direction spanning from Lisbon to the south slope of the Cantabrian Mountains (Fig. 1). The Douro Basin and the Tagus Basin registered 62 and 15% of total DISASTER cases, respectively (Fig. 2a), but the highest DISASTER cases density (1.4/1000 km²) was registered in the Lima and Cávado basins located in the NW sector of the Iberian Peninsula.

The locations of the DISASTER cases were also analysed taking into account their relative position in three main sectors of the corresponding hydrographic basin (headwater sector, middle sector, and lower sector). These sectors were identified within each hydrographic basin by dividing the corresponding DEM (pixel = 1 km) into three elevation classes of equal area. The lower sector of the affected hydrographic basins contains more than half of DISASTER cases considering the total hydrographic basins affected by the December 1909 event (Fig. 2b). In the western smallest hydrographic basins of Lima and Cávado basins and Sado Basin (3 and 7, Table 3) the DISASTER cases only occurred in the lower sector. Disaster cases located in the headwater sector were only registered in the major river basins of Douro and Tagus (10.6 and 4.8 %, of total cases, respectively).

The spatial distribution of fatalities generated by floods and landslides in Iberia during the 1909 event is presented in Fig. 3a. Floods were particularly disastrous along the Douro Basin, where 60% of total fatalities associated with floods were registered (Table 3). However, the DISASTER case

Hydrographic basin	Area (km ²)	No. of disaster cases	% disaster cases	Disaster cases density (#/1000 km ²)	No. of fatalities	No. of injuries	No. of evacuated people	No. of homeless people	No. of missing people	Total affected people
				Floods						
1 – Coruña basins	12 663	2	1.5	0.16	0	0	3	84	0	87
2 – Minho Basin	16993	8	6.2	0.47	9	0	18	196	0	223
3 – Lima and Cávado basins	6277	9	6.9	1.43	2	0	14	0	8	24
4 – Douro Basin	97 502	83	63.8	0.85	34	2	595	2104	6	2741
5 – Águeda, Mondego, and West basins	14119	7	5.4	0.50	4	0	410	3	0	417
6 – Tagus Basin	81 1 15	20	15.4	0.25	8	2	61	177	1	249
7 – Sado Basin	10074	1	0.8	0.10	0	0	3	0	0	3
Floods Total		130	100		57	4	1104	2564	15	3744
Landslides										
2 – Minho Basin	16993	1	25	0.06	26	0	0	100	0	126
4 – Douro Basin	97 502	2	50	0.02	3	0	0	0	0	3
6 – Tagus Basin	81 115	1	25	0.01	3	0	0	0	0	3
Landslides total		4	100		32	0	0	100	0	132
Total		134			89	4	1104	2664	15	3876

Table 3. Social consequences of the December 1909 event per hydrographic basin.

Note: hydrographic basins are shown in Fig. 1.

that generated more fatalities (26 fatalities) was produced by a single landslide (debris flow) that occurred in Orense province in the Minho Basin.

Floods generated the vast majority (96.2%) of total homeless people, and among these, 82% were located in the Douro Basin, mostly on the Spanish side of the basin (Fig. 3b, Table 3). In Spain, the provinces of Salamanca, Palencia, Zamora Orense, and León registered the highest number of homeless people due to floods along the Douro tributaries (Tormes, Carrión, Esla, Tera, and Luna rivers) (Fig. 4). In Portugal, the highest number of homeless people was registered at Santarém, in the lower sector of the Tagus hydrographic basin.

During the DISASTER event of December 1909 it was also necessary to evacuate people (Fig. 3c) rescued from the waters or flooded buildings. There were two major spots of evacuated people in Portugal, the first in the Lisbon district and the second at the mouth of the Douro River in Oporto City (Oporto district), where 400 and 105 people, respectively, were evacuated. In Spain, there was an important spot of evacuated people (101 people) near the confluence of the Carrión River with the Douro River in the province of Palencia (Fig. 4). The Douro Basin includes 53.9 % of total evacuated people registered during the December 1909 event.

Table 3 highlights the Douro Basin as the major zone impacted by floods. This basin registered 64% of total Disaster flood cases and 73% of total affected people by floods. In the case of landslides, the Minho Basin had only a single DISAS-TER case, however, this single event was responsible for the highest number of fatalities (26) and homeless people (100) generated by landslides.

Figure 5 illustrates the spatial variability for different Iberian regions of the three damage indices considered (fatalities index, homeless index, and total affected index). The Fatalities Index was classified as Extreme in the provinces of Zamora and Orense (Minho Basin) and Very High in Braganca and Salamanca in the middle sector of the Douro Basin and Santarém in the lower sector of the Tagus Basin. In the provinces of Palencia, Zamora, and Salamanca, located in the Douro Basin headwater and middle sectors, the Homeless Index was Extreme (Fig. 5b). The same provinces registered an Extreme Total Affected Index (Fig. 5c), which was mainly justified by the very high number of homeless people located in the headwater and middle sectors of the Douro Basin. The total affected index was very high in the lower section of the Tagus Basin (Lisbon and Santarém districts) as well as in the headwater and middle sectors of the Douro Basin (León province), and the Minho Basin (Orense Province).

Despite the described social impacts, the December 1909 event also caused huge economic losses in Iberia. In Spain, several settlements were completely submerged by water and a great number of buildings were destroyed, especially in the areas drained by the main tributaries of the Douro River (e.g. in the provinces of Zamora, Salamanca, León, Palencia; Fig. 4). Also, in Portugal Peso da Régua and the riverside areas of Oporto and Vila Nova de Gaia were submerged by the Douro River.

Major cities of Spain (e.g. Salamanca) and Portugal (e.g. Oporto) were kept isolated, without communications (mail, telegraph) with the outside world due to widespread cuts in road and rail networks (submerged or destroyed roads



Figure 2. Percentage of Disaster cases and Disaster cases density per hydrographic basin (**a**) and percentage of Disaster cases per hydrographic basin sector (**b**). Hydrographic basin code: 1 – Coruña basins; 2 – Minho Basin; 3 – Lima and Cávado basins; 4 - Douro Basin; 5 – Águeda, Mondego and West basins; 6 – Tagus Basin; 7 – Sado Basin.

and railroads, bridges collapsed). Rail connection between Portugal and Spain was also interrupted for several days.

Finally, it was also reported that some cattle herds were killed by the floods in the Spanish provinces affected by the event, as well as widespread destruction of cropped areas (e.g. vegetables, grain) that caused food scarcity after the floods. ("La Correspondencia de España", 22–28 December 1909).

4.1.2 Temporal distribution of the impacts

Daily evolution of fatalities and affected people generated during the December 1909 event is shown in Fig. 6. Between 20 and 22 December 48.8 % of the Disaster cases were registered (Fig. 6a). In addition, the highest daily number of DISASTER cases (58) and fatalities (52) were registered in 22 December. In terms of total affected people, it is possible to recognize two peaks, the major one on 22 December (1842 affected people) and a second on 25 December (1014 affected people).

In the Douro Basin the daily evolution of the social impacts (Fig. 6b) is quite similar to the complete set of the event. In contrast, the Tagus Basin only registered human impacts on 22, 24, and 27 December (Fig. 6c), with a peak of Disaster cases and affected people on 22 December.

The spatial distribution of the DISASTER cases per day is shown in Fig. 7. In general terms, DISASTER cases that occurred between 20 and 23 December arose predominantly in the western part of Portugal with some isolated spots in Spain (Salamanca, Zamora, León, and Orense). Disaster cases registered between 24 and 28 December were widespread in Spain in the Douro and Tagus headwater and middle sectors. Some isolated spots occurred in Portugal on 27 and 28 December in the lower sectors of the Minho, Lima and Cávado, and Tagus basins. Therefore, overall the temporal distribution of DISASTER cases is consistent with an atmospheric circulation in the Iberian Peninsula coming from SW to NE. However, we must keep in mind that information flow at the time was considerably slower than today, particularly during disruptive events such as major natural hazards. In this regard, it is possible that some of these later dates can correspond to cases that were reported a few days after their real occurrence without correction of the reported timing.

4.1.3 Circumstances surrounding fatalities

Circumstances surrounding the fatalities by gender during the December 1909 event are presented in Fig. 8. Fatalities were dominant in rural areas, both inside buildings and outdoors (87.6% of total fatalities). For the hydrogeomorphological cases where the gender of the victim is known, it is clear that male fatalities were dominant (81% of total fatalities). Also male fatalities were prevalent in rural areas. The predominance of male fatalities generated by floods and landslides is consistent with previous observations on mortality associated with hydro-geomorphologic disasters in Portugal (Pereira et al., 2015). The highest percentage of total fatalities with gender not reported that occurred inside buildings in rural areas (Fig. 8) correspond mostly to fatalities generated by landslides (81.3%).

4.2 Precipitation event and hydrologic context

Daily long-term cumulative precipitation from 1 September to 31 December is shown in Fig. 9 (grey thin lines) for Oporto, Lisbon, Guarda, and Salamanca meteorological stations. In addition, the long-term mean (magenta line) and the 95th percentile (black line) for 1909 (red line) are also highlighted. It is clear that the December 1909 event was preceded by an intense precipitation in late autumn, especially after 15 November where it surpasses the mean in all locations and also a period of enhanced precipitation after the 17 December. These features are particularly notable at the Guarda station, where the cumulative precipitation between 17 and 31 December 1909 reached a total of 406.7 mm. Mean monthly precipitation for December at Guarda station is 141.8 mm, while the total precipitation registered in De-



Figure 3. Social consequences of the DISASTER cases: fatalities (a), homeless people (b), and evacuated people (c).



Figure 4. Douro Basin and DISASTER cases of the December 1909 event.

cember 1909 was 451.1 mm (Table 2). In December 1909 a total of 80.7 mm was registered at Salamanca (although much lower than the value observed at Guarda, it corresponds to twice the amount of the mean monthly precipitation for December) (Table 2). It is important to mention that in November and December 1909 strong negative values of NAO (-1.9 and -2.5, respectively) were registered. The NAO plays an important role in shaping European precipitation variability (Trigo et al., 2008) especially in the Iberian Peninsula, where negative values of NAO are associated with above-normal precipitation in the region. This significant impact results from the displacement of the storm track further south of its normal position during months with negative NAO index, leading to a higher frequency of Atlantic lowpressure systems (mid-latitude cyclones) that travel into the region (Trigo et al., 2004, 2008). The impacts of the negative values of the NAO help to put into context the above-average precipitation that occurred in November and December.

After 22 December 1909 the cumulative precipitation (red line) was higher than the 95th percentile (black line) at Guarda and Salamanca stations, both located in the Douro Basin. The extreme rainy period is very clear at the Guarda station where a daily maximum precipitation of 185.2 mm was recorded on 22 December 1909, corresponding to a return period of 141 years. Regarding Oporto and Lisbon, both stations recorded values of accumulated precipitation up to 22 December below the 95th percentile but still above the long-term mean. In addition, for Lisbon, two particular years with major floods are also highlighted for comparison: 1876 (blue line) and 1967 (green line).

The relevance of above-average precipitation before a major event depends on the temporal and spatial scales of the phenomenon being analysed. The December 1909 event along with the 1876 event (Trigo et al., 2014) highlights the key role played by previous accumulated precipitation in the occurrence of floods during events lasting several days in the major river basins of the western Iberian Peninsula, particularly for the period prior to the construction of major dams (Trigo et al., 2014). In contrast, floods occurring after intense bursts of precipitation over relatively confined areas of western Iberia can induce flash floods, independently of the precipitation that occurred in previous weeks or months. That was the case of the sub-daily 1967 event which occurred during the night between 25 and 26 November 1967, which was confined to the Lisbon metropolitan area, i.e. without affecting large river basins. The 1967 event was the deadliest storm affecting Portugal since at least the early 19th century with several flash floods in heavily constructed suburban areas around Lisbon causing more than 500 fatalities. In this case, the accumulated precipitation prior to 25 and 26 November 1967 was not relevant (being below average) and the flash floods occurred mainly as a result of very intense hourly precipitations, ranging in duration from 4 to 9 h, compatible with return periods of 100 years or more (Trigo et al., 2015).

As already mentioned, the precipitation record corresponding to the December 1909 event was particularly notable at the Guarda station. Therefore, in order to further evaluate the time duration and magnitude of the event, the accumulated precipitation values (from 1, 2, 5, 10, 15, 30, and 45 days) for the Guarda station (88 years) were computed and all values were ranked, with the top 10 events being represented in Fig. 10. A similar methodology was applied in a previous analysis of an outstanding event, the record floods of 7 December 1876 in southern Iberia (Trigo et al., 2014). Figure 10 shows on the *x* axis the 10 highest



Figure 5. Fatalities index (a), homeless index (b), and total affected index (c) of the December 1909 event for affected regions over Iberia.

values of maximum accumulated precipitation for different temporal durations (y axis). The blue bars correspond to accumulated precipitation (z axis) for different durations of the December 1909 event. For the complete precipitation series of Guarda station, the December 1909 event occupies the first position in the cumulated precipitation for 1 day (185.2 mm) and the second position for 2 days (279.8 mm) (Fig. 10).

Interestingly, the extreme precipitation values observed in Guarda and to a less extent in Salamanca were not observed in most locations with pluviometers. Weather stations located near the Atlantic coast (Oporto and Lisbon) did not record extreme accumulated precipitation in the December 1909 event (Figs. 9 and 11) and the stations located in the eastern part of Spain registered null or low precipitation (below 20 mm). Estimated precipitation of 22 and 20–28 December 1909 was interpolated spatially using a co-kriging geostatistical technique considering the precipitation measured in the stations (points) and the Iberia SRTM 90 m DEM's



Figure 6. Daily number of DISASTER cases, fatalities, and affected people generated by the December 1909 event (a), in the Douro Basin (b), and in the Tagus Basin (c). Notice the logarithmic scale used.

with 90 m resolution (Jarvis et al., 2008) (Fig. 11a and b). Estimated precipitation shows that the event entered in the Atlantic coast (south of Lisbon region and crossed the peninsula in a SW–NE direction. In addition, the extreme precipitation that occurred on 22 December in the Douro Basin, namely in the Guarda and Salamanca regions (Fig. 11a), is impressive.

According to Rodrigues et al. (2003) the flood of 1909 was referenced as the second historical flood measured at Peso da Régua, in the Portuguese sector of the Douro River (Fig. 4) in the last 300 years. In the historical flood of 1909 a flow rate of 16 700 m³ s⁻¹ was reached, which was only surpassed by the flood of 1739 (18 000 m³ s⁻¹).

A graph with the evolution of sub-daily flow rate at Peso da Régua for the period of 18–30 December 1909 is shown in Fig. 12, using data from Rodrigues et al. (2003). At the beginning of this period a constant rise in the flow rate was recorded, from $2000 \text{ m}^3 \text{ s}^{-1}$ on 18 December to $6000 \text{ m}^3 \text{ s}^{-1}$ in the early hours of 21 December. Then, the flow rate sud-

denly increased to $16000 \text{ m}^3 \text{ s}^{-1}$ during the day of 23 December, as a consequence of the extreme precipitation that was registered along the Douro Basin during the previous days, especially on 22 December. The maximum flow rate of $16700 \text{ m}^3 \text{ s}^{-1}$ corresponds to a return period of 80 years (Rodrigues et al., 2003). At Peso da Régua the river level rose 25 m during this flood ("Jornal de Notícias", 24 December 1909). It should be noticed that at Peso da Régua the flow rate of $6000 \text{ m}^3 \text{ s}^{-1}$ has a 16-year return period and implies a 13 m river level rise, which is considered an extraordinary flood in the region (Rodrigues et al., 2003).

The flow rate starts to decrease at the end of 23 December but a second peak was registered at the end of 24 December with similar magnitude ($16\,200\,\text{m}^3\,\text{s}^{-1}$), following a secondary precipitation peak that is not evident in precipitation daily records. Only afterwards, i.e. after 25 December, did the flow rate decrease progressively until the end of the month ($2800\,\text{m}^3\,\text{s}^{-1}$ on 30 December). It is important to note that the flow rate was higher than 6000 m³ s⁻¹ between 21 and 28 December, coinciding with the Disaster event time interval.

4.3 Assessing atmospheric circulation

Here we intend to characterize the atmospheric circulation affecting western Iberia between 15 November and 28 December. However, the main focus of this section will be on 22 December, in order to understand the occurrence of highest precipitation values within this day.

The importance of the accumulated precipitation observed between mid-November and mid-December culminating with the intense precipitation on 22 December has already been shown in Sect. 4.2. The 6-hourly sequence of SLP and corresponding accumulated precipitation (both obtained from the 20 CR) between 15 November at 00:00 UTC and 28 December at 18:00 UTC can be seen in the Supplement video V1. The week between the 16 and 21 November was characterized by a stationary low-pressure system that affected the Iberian Peninsula with precipitation occurring during these 6 days. The most intense precipitation over these days was recorded on 16 November at 18:00 UTC with high precipitation rates. According to the specific humidity and wind vectors at 900 hPa (shown in video V2 in the Supplement also between 15 November to 28 December) high moisture availability was present on 16 November at 18:00 UTC due to high moisture advection from the western African Atlantic coast as a consequence of intense southwest winds.

Between 22 November and 12 December the atmospheric circulation in the Iberian Peninsula was mainly of anticyclonic nature with sporadic precipitation in the north Iberian Peninsula due to the passage of some low-pressure systems north of the domain (video V1 in the Supplement). On 13 December a low-pressure system had its cyclogenesis west of the Azores and become stationary near the Azores until 21 December. From 17 December (video V1 in the Supplement)



Figure 7. DISASTER cases of the December 1909 event per day.



Figure 8. Circumstances surrounding the fatalities by gender during the December 1909 event in Iberia.

onwards it affected the western Iberian Peninsula with almost continuous precipitation every day until 22 December, when the most intense precipitation occurred (as seen in Fig. 13).

A thorough analysis is now provided covering the days of most intense precipitation, as seen by the sequential panels every 6 h of SLP and the corresponding accumulated precipitation (both from the 20 CR) represented from 22 December at 00:00 UTC to 23 December at 00:00 UTC (Fig. 13). This figure reveals the presence of a large deep low-pressure system (966 hPa) located over the Atlantic Ocean, northwest of the Iberian Peninsula and extending from Iceland to southern Portugal and to Central France. This low-pressure system remained quasi-stationary from 22 December 00:00 until 12:00 UTC when it started to move east towards the British Isles and its central core began to increase pressure (Fig. 13d and e). In addition, the frontal system associated with this low pressure struck the western Iberian Peninsula from a SW–NE direction. Although the 20th Century Reanalysis does not provide explicit front analysis, the precipitation pattern around the cyclone centre corresponds clearly to a frontal system when compared with a conceptual model of an extratropical cyclone (Shapiro and Keyser, 1990). In fact we must emphasize that the vast majority of the Atlantic midlatitude cyclones that struck Iberia are characterized by the presence of frontal systems (sometimes already in the occlusion phase), that extend the precipitation impacts considerably to the south of the low-pressure system centre (Trigo et al., 2004).

With the aim of further analysing the upper- and lowerlevel dynamics of this event we computed the divergence fields simultaneously with the wind speeds for both the 900 and 250 hPa (Fig. 14). The wind speed and associated divergence at the 250 hPa level is shown in the upper panels of Fig. 14, for 00:00 and 12:00 UTC respectively. The most interesting feature for this event is associated with the presence of an upper-level jet branch, oriented along a SW-NE direction towards the Iberian Peninsula but with its origin in the eastern coast of USA. This feature presents an area with strongest upper-level divergence over Portugal at 00:00 and 12:00 UTC, with values above $3 \times 10^{-5} \text{ s}^{-1}$ favouring deep convection. In addition, at lower levels (900 hPa), a very intense low-level jet (winds up to $30 \,\mathrm{m \, s^{-1}}$ close to Portugal) associated with the frontal system embedded within the southern flank of the low-pressure system is also present with a similar SW-NE orientation. This intense advection of mass



Figure 9. Cumulative precipitation from 1 September to 31 December using daily precipitation data from (**a**) Oporto, (**b**) Lisbon, (**c**) Guarda, and (**d**) Salamanca. Each year of cumulative precipitation is represented in grey. The mean and the 95th percentile of precipitation are represented in magenta and black respectively and accumulated precipitation in the year of 1909 is represented by the red line. In addition, the historical storm years of 1876 (blue) and 1967 (green) in Lisbon are also highlighted.

implies strong low-level convergence at the western coast of the Iberian Peninsula $(-3 \times 10^{-5} \text{ s}^{-1})$. In summary, we can state that the large-scale conditions in the area of the western Iberian Peninsula were optimal for intense precipitation, i.e. frontal uplift associated with upper-level divergence and low-level convergence. The displacement of the jet stream so far south was also responsible for the stationary position of the low-pressure system for so long.

However, one should bear in mind that the presence of the low-pressure system and its associated frontal system and the favourable dynamical mechanisms for intense precipitation explained before cannot alone explain the intense precipitation that occurred on 22 December.

The analysis of Fig. 15 (upper panel, geopotential height and temperature at 500 hPa) shows that the geopotential height is characterized by a strong trough located south of the Azores thus favouring the southerly displaced position of the jet in the middle of the Atlantic (Fig. 14). Along with this trough there is a negative temperature gradient between 60 and 30° N with dry and cold air aloft near the Iberian Peninsula which also could enhance the formation of convection within the frontal system.



Figure 10. The 10 highest values (rank) of maximum accumulated precipitation from 1, 2, 5, 10, 15, 30, and 45 days in Guarda for the period 1906–1994. Blue bars correspond to the December 1909 event.



Figure 11. Co-kriging interpolation of total precipitation of 22 December (**a**) and event cumulated precipitation in 20–28 December 1909 (**b**) in Iberia.



Figure 12. Flow rate $(m^3 s^{-1})$ at Peso da Régua from 18–30 December 1909. Source: SNIHR (Rodrigues et al., 2003). The dashed line corresponds to an extraordinary flood at Peso da Régua (Portugal).

To analyse the moisture availability, which can enhance the precipitation intensity, the specific humidity and wind vectors at 900 hPa are also shown for the 00:00 and 12:00 UTC (Fig. 15, lower panels). In both time steps, a narrow and prolonged band of moisture being transported from sub-tropical latitudes can be seen, with origin close to the Caribbean islands. This narrow band presents an intense core with values above 9 g kg^{-1} , over the Atlantic Ocean, and moves towards extratropical latitudes, by a southwestern low-level jet with wind speeds clearly above 24 m s^{-1} (Fig. 14). Therefore the two main criteria for the identification of AR are fulfilled (see Sect. 2) on 22 December, enhancing the precipitation intensity during this particular day.

On 23 December, the low-pressure system moves NW towards the British Isles and the Iberian Peninsula becomes less under the influence of this low pressure. Despite this,



Figure 13. The sea level pressure field (contour, hPa) and 6 h accumulated precipitation (shaded, mm) for 22 December 1909 at (a) 00:00 UTC, (b) 06:00 UTC, (c) 12:00 UTC, (d) 18:00 UTC, and (e) 23 December at 00:00 UTC.

some light precipitation continues to occur between 23 and 25 December as shown in video V1 in the Supplement.

Summarizing, the synoptic analysis shows that 22 December was characterized by the presence of a deep low-pressure system located over the NW Iberian Peninsula with a stationary frontal system striking the western Iberian Peninsula. The presence of an upper-level jet (250 hPa) and low-level jet (900 hPa) located on SW–NE oriented towards the Iberian Peninsula along with upper-level divergence and lower-level convergence favoured large-scale precipitation. Along with these two features, the presence of an AR (with high content of moisture) contributed to a large extent to enhanced precipitation.

5 Discussion and conclusions

This study provides the first comprehensive assessment of a major hydro-geomorphologic event that occurred in Iberia in the last 150 years, including: (i) the spatial and temporal characteristics of the societal impacts based on newspapers analysis; (ii) the corresponding spatial distribution of precipitation anomalies using recently digitized daily precipitation data; and (iii) the meteorological conditions that triggered the event using the 20 CR.

Data collection of human and societal disaster impacts on newspapers was very useful to obtain the spatial distribution of the hydro-geomorphologic DISASTER cases in both countries and provided an important historical complement



Figure 14. Divergence (contours; 10^{-5} s⁻¹) at 22 December 1909 at 00:00 and 12:00 UTC for (**a**, **b**) the 250 Pa level, and (**c**, **d**) at the 900 Pa level. Along with the divergence field, the respective wind speed (shaded, m s⁻¹) is also shown.



Figure 15. The geopotential height (contour, m) and air temperature (shaded; °C) at 500 hPa level for 22 December: (a) 00:00 UTC and (b) 12:00 UTC. The wind vector (m s⁻¹) and specific humidity (shaded, $g k g^{-1}$) at 900 hPa level for the 22 December 1909 at (c) 00:00 UTC and (d) 12:00 UTC are also shown. Only wind vectors with wind speeds above 12.5 m s⁻¹ are shown.

useful to offset the lack of precipitation data in some of the affected areas. In this work new information concerning the event impacts in Spain was gathered with respect to the already available DISASTER database for Portugal. According to newspapers, the December 1909 event generated 89 fatalities, 4 injured, 1104 evacuated, and 2664 homeless people. However, the number of injured, evacuated, and home-less people is certainly underreported because the newspapers clearly highlighted flood and landslide cases that generated fatalities.

Hydro-geomorphological event damage assessment is particularly difficult to quantify due to the diversity of damage (human, economic, structural, or functional) and the scarce information reported in contemporary newspapers. This problem is more accentuated in the study of historical events like the December 1909 event because it is virtually impossible to reconstruct the level of loss verified in the different affected regions. To help in tackling this caveat we have introduced three damage indices that allowed us to compare societal damages in the affected regions, because these indices are normalized by the resident population at the time of the event. In general terms damage indices were higher in Spanish regions than in Portuguese. Fatalities associated with this event were more abundant in rural areas. In addition, this study demonstrates that males were more exposed to flood risk, which is explained by the more active role of males in agriculture activities and pasturing at that time.

The length of this event extending from 20 to 28 December 1909 represents an extreme episode affecting most of the western and northwestern sector of the Iberian Peninsula from the impact perspective, particularly along the Douro Basin. In addition, the temporal distribution of DISASTER cases is consistent with a frontal system that crossed the Iberian Peninsula travelling on a SW to NE orientation and associated with a low pressure system located near the British Isles.

The December 1909 event was characterized by a period of enhanced precipitation after 17 December and was associated with some outstanding precipitation records particularly notable at the Guarda station, where the cumulative precipitation between 17 and 31 December 1909 reached approximately three times more than the mean monthly precipitation for December. Since 22 December 1909 the cumulative precipitation was higher than the 95th percentile at Guarda and Salamanca stations, both located in the Douro Basin.

The extreme precipitation was not observed in most locations of the Iberian Peninsula where pluviometers were available. Weather stations of Oporto and Lisbon located near the Atlantic coast in the mouth of the Douro and Tagus rivers, respectively, did not record extreme accumulated precipitation in the December 1909 event. Nevertheless, these areas were also affected by floods as a result of the intense precipitation that occurred in the upper and middle sectors of the Douro and Tagus hydrographic basins. The limited available daily precipitation data do not show the precipitation amount that may have occurred in the affected area of the event especially on higher mountain ranges such as Gerês or Cantabrian Mountains. Since the ARs are rich in water vapour and are associated with strong winds that force the water vapour up mountain sides, the atmospheric conditions at that time did not inhibit upward motions which were responsible for higher precipitation values in mountain areas than over flat coastal areas. This can only be seen at Guarda rain gauge (altitude of 1019 m) that registered a daily precipitation record, but we believe that equally extremely high precipitation values must have occurred also in other mountain ranges especially in the Cantabrian Mountains due to the nature of the floods that occurred in northern Spain.

The assessment of the large-scale atmospheric conditions suggests that the extreme precipitation event verified in December 1909 resulted from the merging of unusual meteorological conditions. The presence of atmospheric instability due to the existence of a low-pressure system and its associated frontal and post-frontal sectors and also the presence of an AR structure (with particularly high content of humidity) responsible for the advection of moisture towards the Iberian Peninsula contributed to enhance precipitation. It was shown by Ramos et al. (2015) that 60 % of the most anomalous precipitation days are associated with the presence of an AR in the Iberian Peninsula. Since the average frequency of ARs during the extended winter (ONDJFM) over the Peninsula is around 10% (although at different latitudes) we believe that more historical Douro floods (Rodrigues et al., 2003; Fragoso et al., 2015) could be associated also with the presence of ARs and deserve a specific assessment.

The DISASTER event of December 1909 took place at a time when the relationship between the precipitation and the river flow was more straightforward in the absence of dams. Nowadays, Douro and Tagus rivers have several dams that control flow rates and reduce flood peaks, thus reducing also the risk of flooding even when higher than normal precipitation occurs (Trigo et al., 2004). Nevertheless, there are some lessons to be learnt from the reconstruction of this major historical hydro-geomorphologic event.

First, data collection of societal disaster impacts based on local and national newspapers can be considered a good source of information to achieve temporal and spatial distribution of impacts, characteristics of circumstances surrounding fatalities and gender of affected people. However, such a document-based disaster inventory is typically incomplete and, as a rule, the reliability of information is higher for the temporal distribution than for the spatial location of disaster cases.

Second, the characterization of the meteorological synoptic forcing of the DISASTER event is crucial to understand the rainfall distribution in space and time, which is of paramount importance to explain the observed social impacts. In addition, in the case of floods the complete fluvial system has to be considered, as the flooding on the lower part of the basins (e.g. the Douro and the Tagus river basins) is related to rainfall occurring in the middle and upper sectors of the basins.

Third, the detailed reconstruction of past hydrogeomorphological disastrous events such as the one presented here is very useful to improve our knowledge on extreme events which may occur again and may have a high impact on the territory. Therefore, this work is also relevant for assessing the exceptionality of past hydro-geomorphologic disasters.

Lastly, the correct understanding of the meteorological drivers of major flood and landslide events like the December 1909 event may improve the early warning systems for floods and landslides in Iberia as the critical meteorological conditions can be nowadays anticipated some days prior to the event.

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